# On Recovering IP Information from EM Data

Doug Oldenburg\*, Dave Marchant, Dikun Yang, Seogi Kang, and Eldad Haber





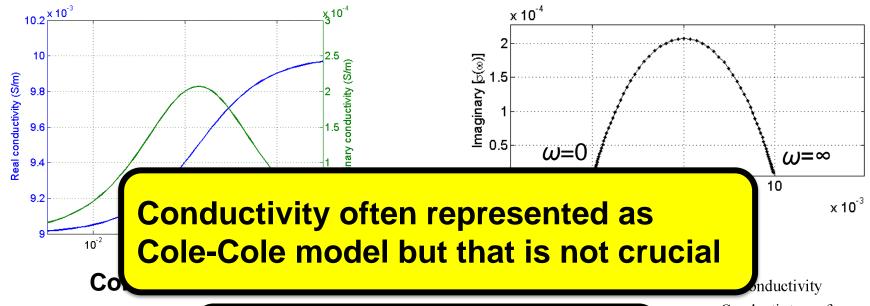
#### Outline

- Background for IP problem
- Forward problem in the time domain
- Inverse problem
- Field example Mt. Milligan
- Conclusions and the path ahead



#### | Electrical conductivity is complex

$$\vec{J} = \sigma(\omega)\vec{E}$$



$$\sigma(\omega) = \sigma_{\infty} \left[ (1 - \frac{\eta}{1 + (1 - \eta)(i\omega\tau)^{c}}) \right]$$

 $\sigma_a$ : Conducti at zero frequency

 $\eta$ : Chargeability

 $\omega$ : Angular frequency

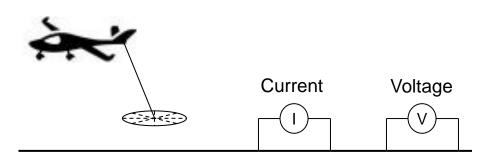
 $\tau$ : Time constant

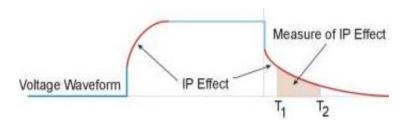
c: Frequency dependence

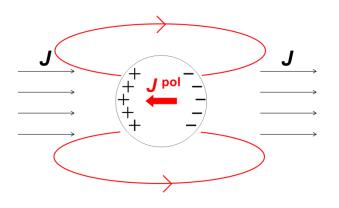


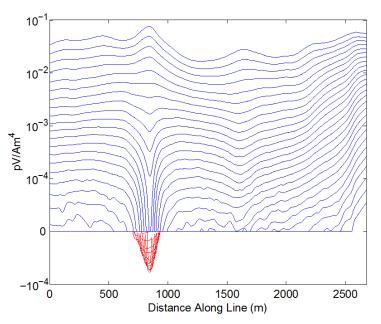


# Different experiments











#### Inverse problem: two paths

1. Directly in time domain

2. Linearized inversion



# **Complex** conductivity

$$\sigma(\omega) = \sigma_{\infty} \left[ \left( 1 - \frac{\eta}{1 + (1 - \eta)(i\omega\tau)^{c}} \right) \right] = \sigma_{\infty} + \Delta\sigma(\omega)$$

 $\sigma(\omega)$ : Conductivity

 $\sigma_o$ : Conducti at zero frequency

 $\eta$ : Chargeability

 $\omega$ : Angular frequency

 $\tau$ : Time constant

c: Frequency dependence

- σ<sub>∞</sub> is background conductivity
- Obtained from early times that are not contaminated with IP effects.
- The IP effect is a perturbation. This is a good approximation when η is small and has been traditionally used



#### IP data and EM coupling removal

Conductivity

$$\sigma(\omega) = \sigma_{\infty} + \Delta \sigma(\omega) \qquad \tilde{\eta}(\omega) = \frac{\eta}{1 + (1 - \eta)(i\omega\tau)^{c}}$$

**Background conductivity** 

Removing EM coupling

Removing EWI coupling 
$$+ \frac{dF}{d\sigma} \Delta \sigma(t)$$
 
$$F[\sigma_{\infty} + \Delta \sigma(t)] - F[\sigma_{\infty}] = d^{IP}(t) = -\frac{dF(t)}{d\sigma} \sigma_{\infty} \eta(t)$$
 
$$d^{IP}(t) = G(t)\eta(t)$$
 
$$G(t) = -\frac{dF(t)}{d\log(\sigma)}$$

**Routh and Oldenburg (2001)** 





#### Need background conductivity ( $\sigma_{\infty}$ )

 Carry out 3D inversion of the AEM data (leave out time channels that have negative transients)

 Methodology: Yang and Oldenburg [AGU, 2012, ASEG 2012]

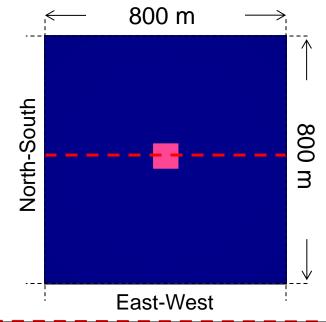
Local mesh for forward modelling and sensitivities

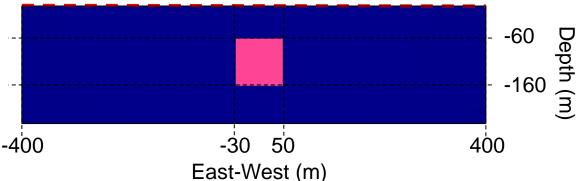
Stochastic selection of transmitters



#### Synthetic example

True model





- ✓ Discretization
  - 20 m core cells
  - 69 x 69 x 50 cells
- ✓ Background

$$\sigma_{half} = 5e-4 \text{ S/m}$$

✓ Target

$$-\sigma_{\infty} = 6.25e-2 \text{ S/m}$$

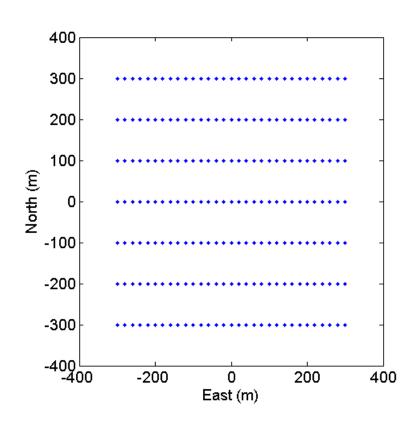
$$-\eta = 0.2$$

$$-\tau = 0.01$$

- 80m x 80m x 80m
- 60m below surface

#### Synthetic example

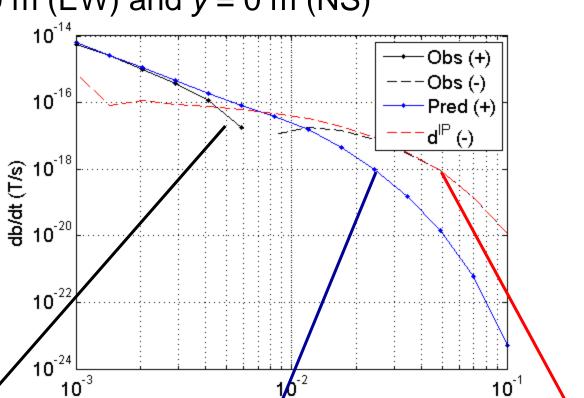
#### Geometry



- ✓ VMD source30m above surfaceCollocated Rx
- $\checkmark$  7×31 = 217 Txs
- ✓ Receivers
  - Collocated with Tx
  - Measuring db/dt
  - 14 time channels
  - $-10^{-3}-10^{-1}$  seconds



• At x = 0 m (EW) and y = 0 m (NS)



 $F[\sigma_{\infty} + \Delta \sigma(t)]$ 

 $F[\sigma_{\infty}]$ 

Time (s)



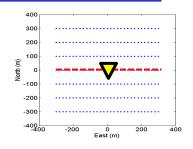
 $d^{L}$ 

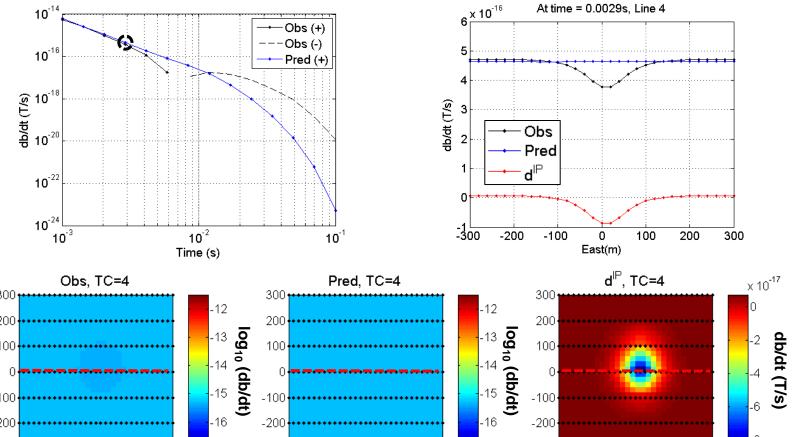
-100





• db/dt data at time = 0.003s





200

-200

-300

-200

200

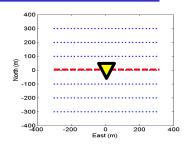
-200

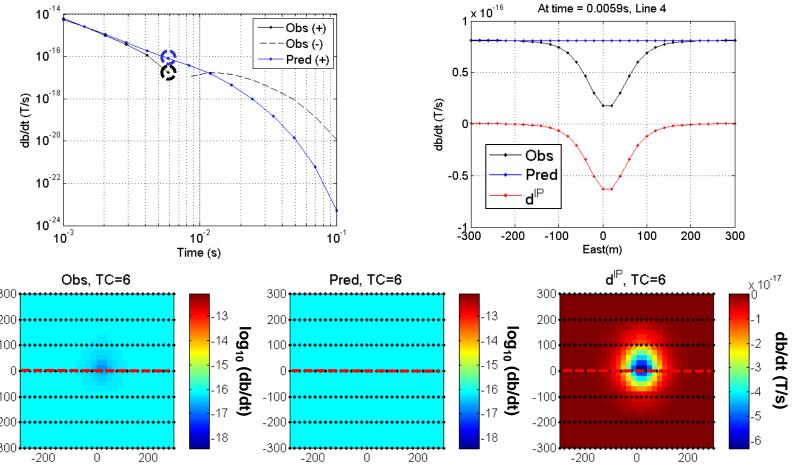


200

0

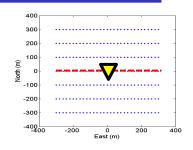
db/dt data at time = 0.006s

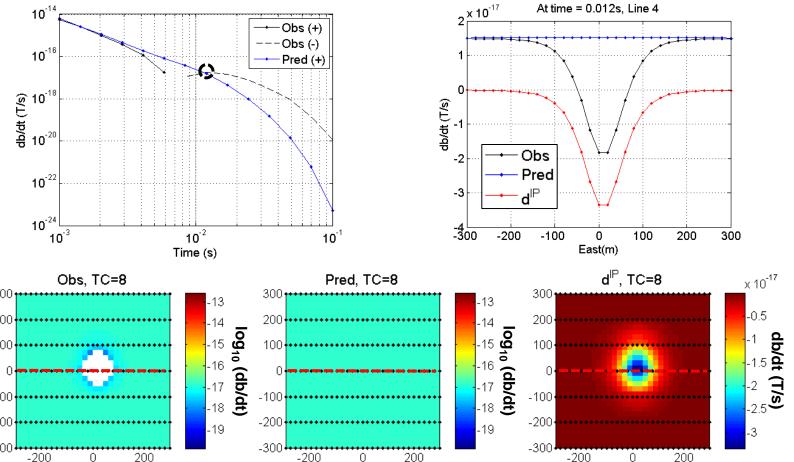






• db/dt data at time = 0.012s

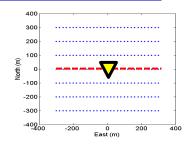


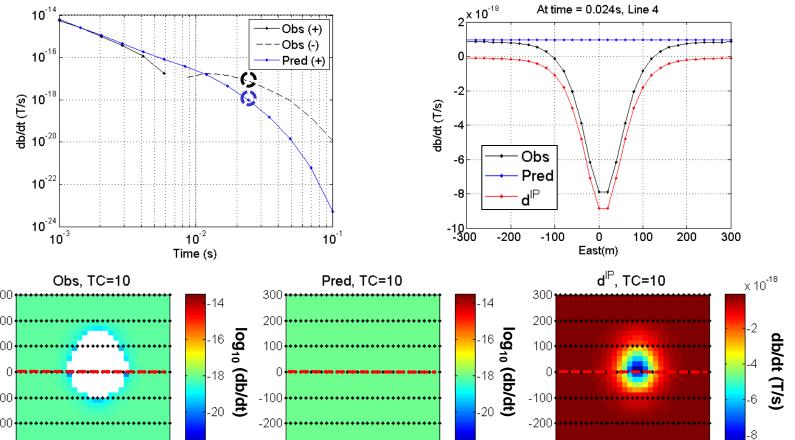






• db/dt data at time = 0.024s





200

0

-200

-300

-200

0

200

n

200

-200



#### |Linear Inversion

#### At each time

$$\min \quad \phi = \phi_d(\mathbf{m}) + \beta \phi_m(\mathbf{m})$$
  
s.t.  $0 \le \mathbf{m}$ 

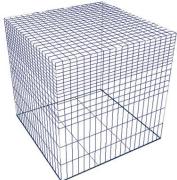
$$\phi_d = \sum_{i=1}^{N} \left( \frac{\vec{d}_i^{pred} - \vec{d}_i^{obs}}{\epsilon_i} \right)^2$$

$$\phi_m = ||\mathbf{W_m}(m - m_{ref})||_2^2$$

$$d^{IP}(t) = G(t)\tilde{\eta}(t)$$

$$G(t) = -\frac{dF(t)}{d\log(\sigma)}$$

#### 3D volume of $\tilde{\eta}(t)$

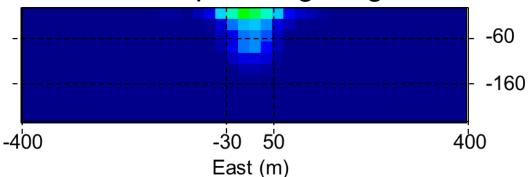




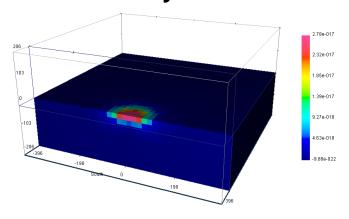
#### Recovered pseudo chargeability

#### <Cross section>

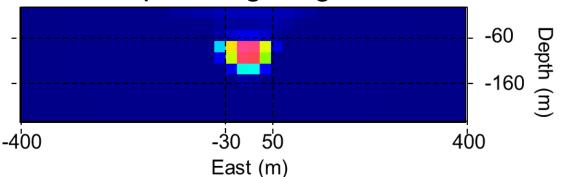
#### Without depth weighting



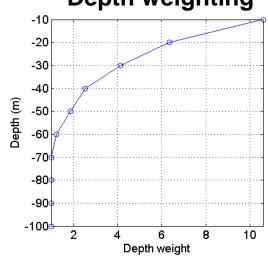
#### Sensitivity volume



#### With depth weighting



#### **Depth weighting**

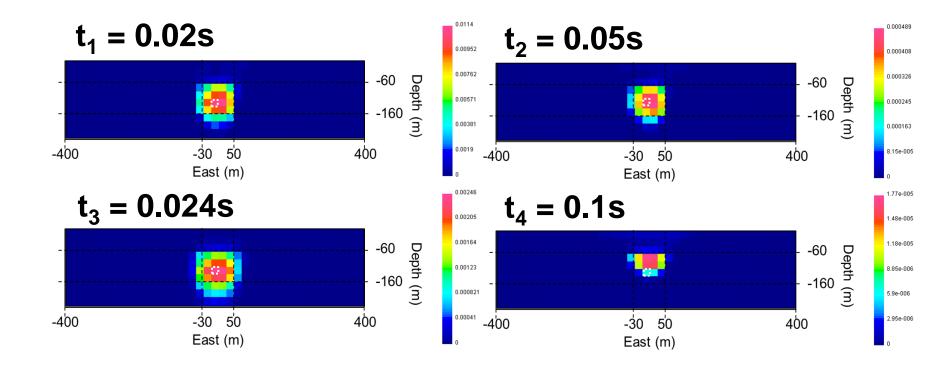






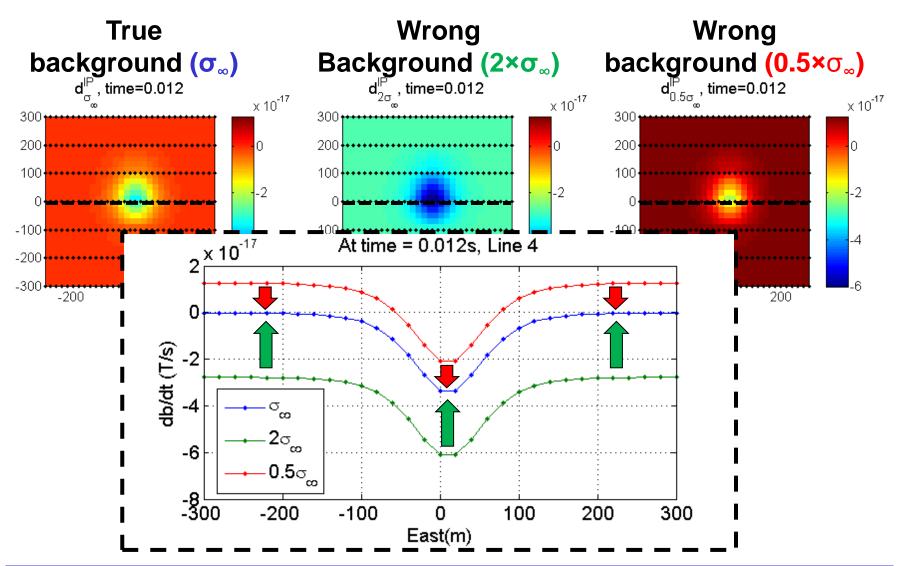
#### Recovered pseudo chargeability

At four different time channels





#### Effects of background conductivity







#### Procedure

Invert TEM to recover a background 3D σ model

Compute d<sup>IP</sup> data at multiple time channels

Estimate a regional field

Subtract regional to obtain IP data

Invert individual time channels

- ✓ Attempt to avoid soundings/times that have IP coupling effects
- ✓ Easy if there are negative transients

Assume IP data are composed of a smoothly varying regional plus true IP data

✓ Not easy; but this issue is confronted before

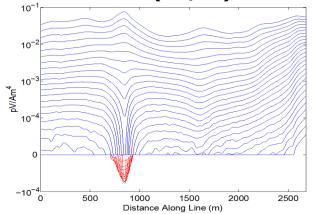


# Field example: Mt. Milligan VTEM

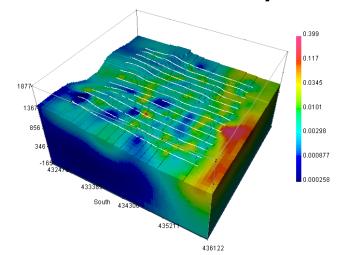


Geotech (2007)

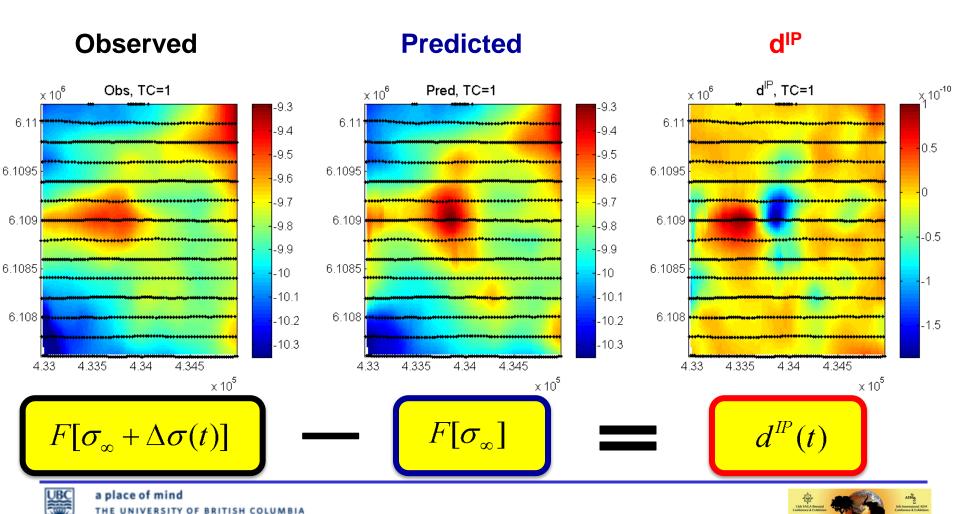
#### √ Observed data (db/dt)



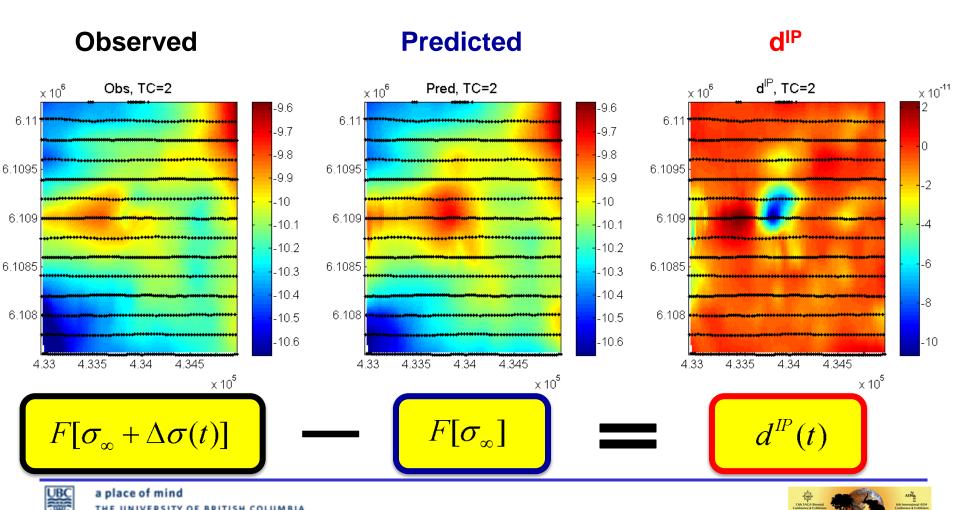
#### ✓ Recovered 3D conductivity model



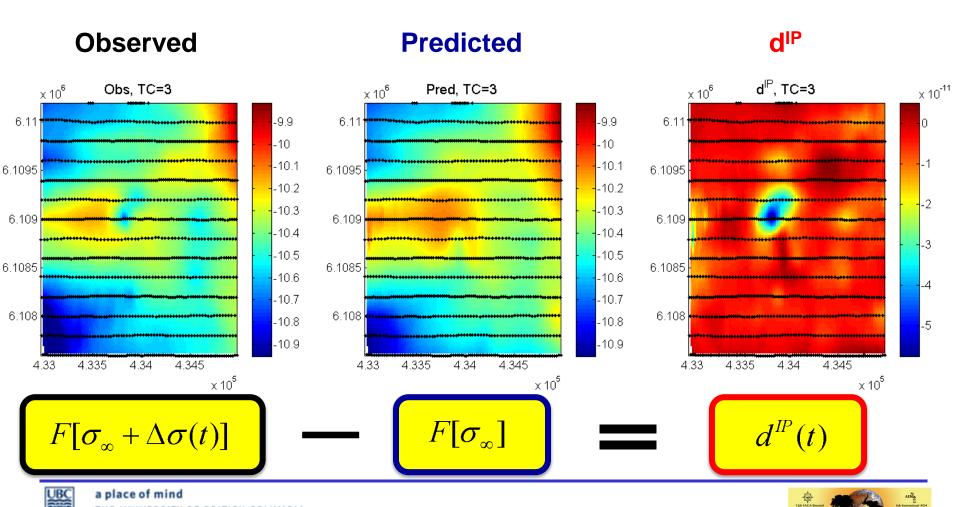
db/dt data at time = 0.00463s



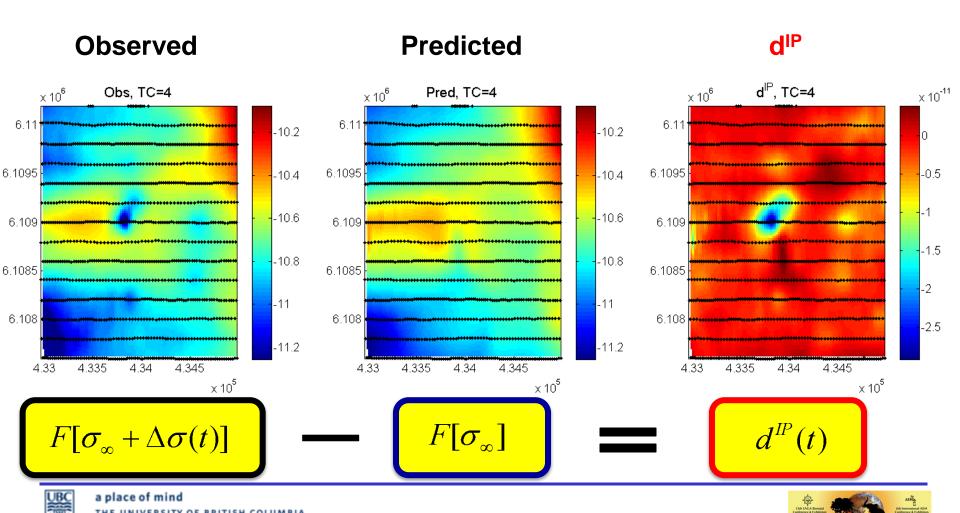
db/dt data at time = 0.00473s



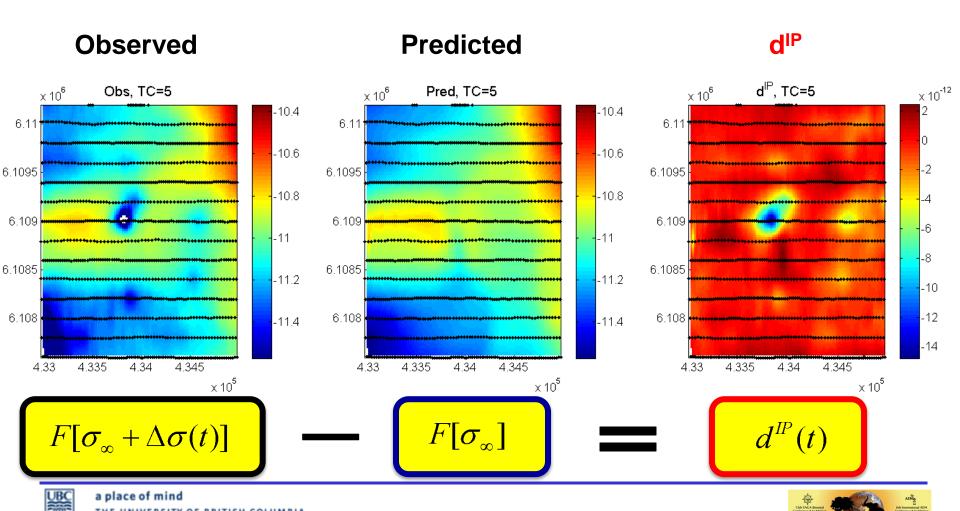
db/dt data at time = 0.00488s



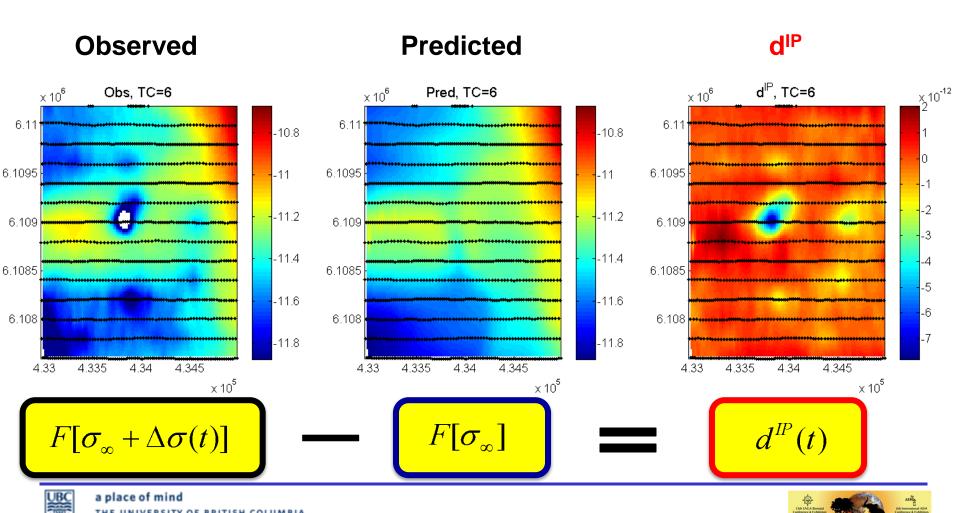
• db/dt data at time = 0.00508s



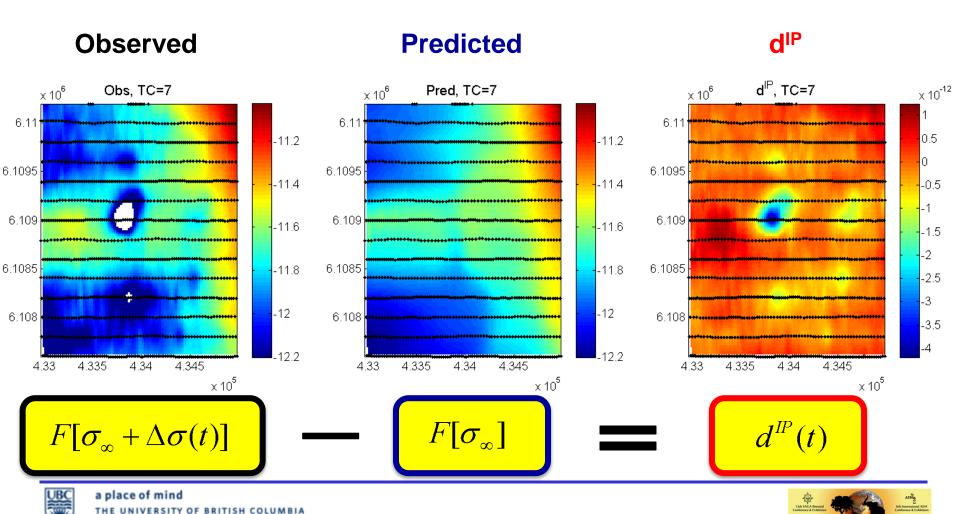
db/dt data at time = 0.00537s



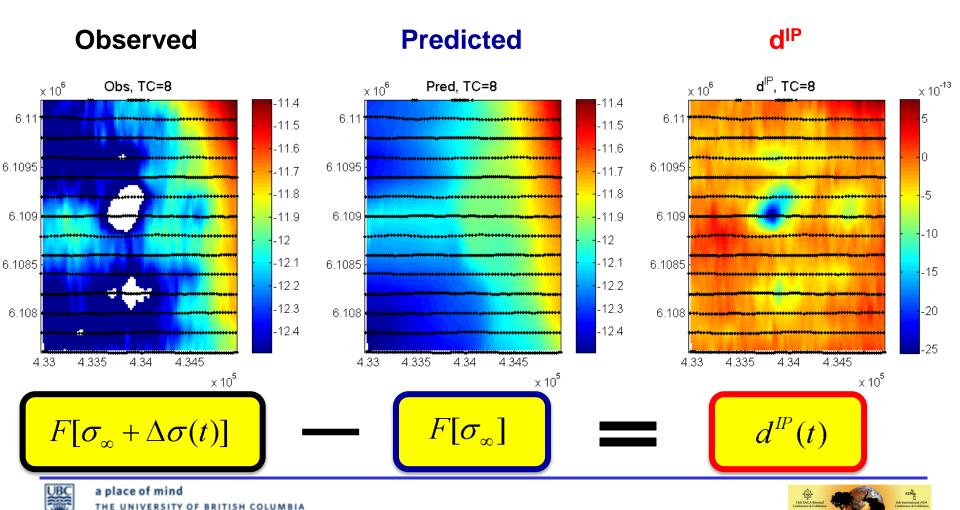
db/dt data at time = 0.00577s



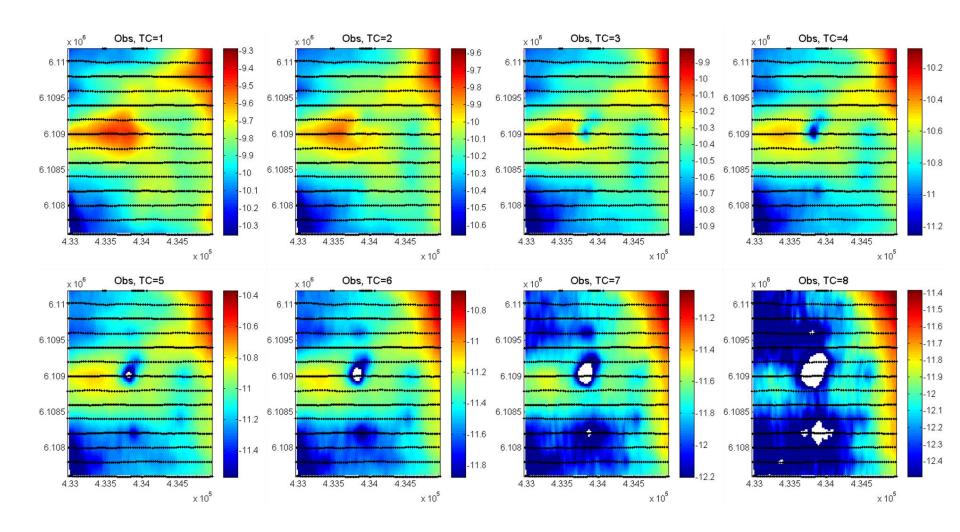
db/dt data at time = 0.00635s



db/dt data at time = 0.00715s



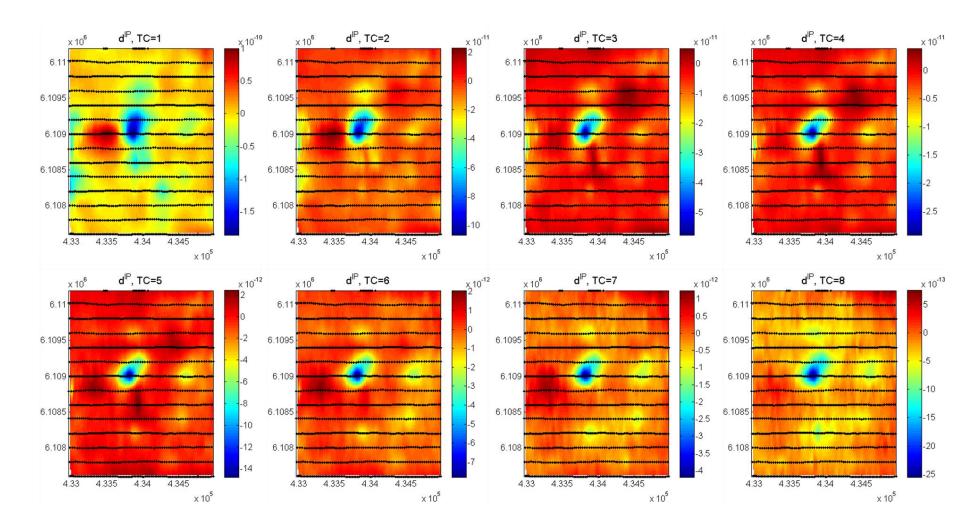
#### Observed data for all channels





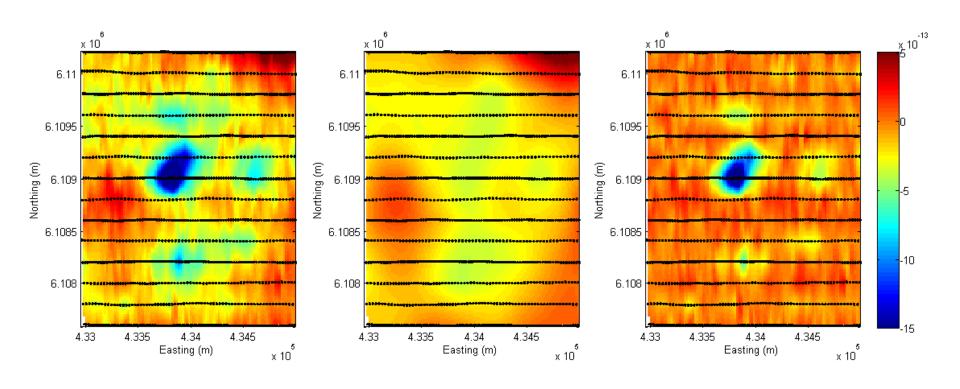


# d<sup>IP</sup> data for all channels





# Removal of a regional



 $d^{IP}$ 

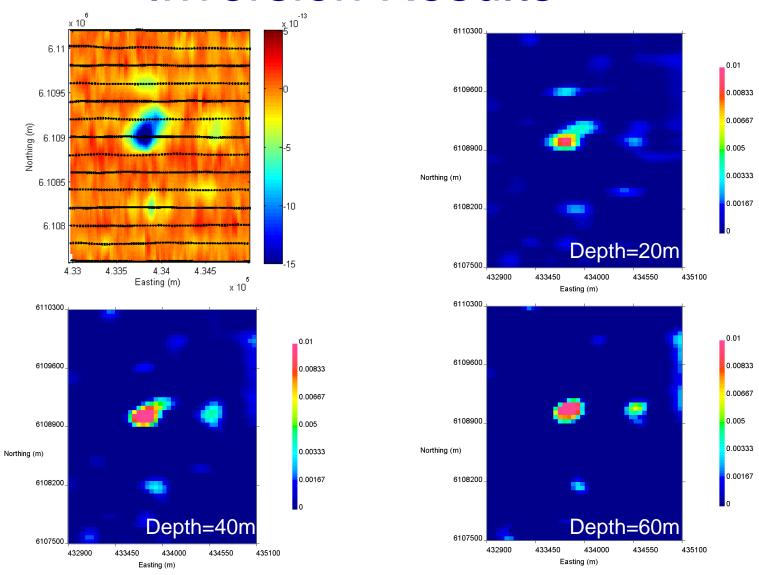
regional

data for inversion





#### **Inversion Results**







#### Summary and path forward

- Estimating σ<sub>∞</sub>, 3D background conductivity. (be mindful of IP coupling)
- $\mathbf{d}^{IP} = \mathbf{d}^{obs} \mathbf{F}[\boldsymbol{\sigma}_{\infty}]$
- For each time channel:
  - ✓ Assume: d<sup>IP</sup> = d<sup>IP</sup>(true) + smooth background signal
- Estimate background signal (challenge)
- Individual inversions and then attempt to extract spectral information



#### Summary and path forward

 More advanced inversions to work with all time channels at once.

 Current progress provides optimism that we can invert the data if it is sufficiently high quality



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