# Study Nuclear Winter with a radiative-convective climate model

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July 31, 2023

#### Abstract

- 1 Introduction
- 2 Methods

# 2.1 Hypotheses and conventions

To simplify some calculations for radiative fluxes, the definition of wavenumber  $\nu = \frac{1}{\lambda}$  is used in this work, where  $\lambda$  is the wavelenght.

Some assumptions are made on the dynamical state and the composition of Earth's atmosphere. Air parcels are supposed to be in hydrostatic equilibrium, with each component obeying the ideal gas law. Gravitational acceleration is constant (cfr. Table ?? for values).

#### 2.2 Longwave radiation

At wavelengths  $\lambda \geq 4\,\mu\mathrm{m}$ , solar radiation has lower intensities than radiation emitted by Earth's surface and atmosphere at the same wavelengths. Moreover, it presents negligible scattering in atmosphere with respect to absorption. For these reasons longwave radiation is considered to be emitted only by Earth's surface and atmosphere.[1, p. 468]

#### 2.3 Shortwave radiation

At wavelengths  $\lambda < 4\,\mu\text{m}$ , solar radiation has much greater intensity than radiation emitted from Earth's surface and atmosphere. Both scattering and absorption by gases, aerosols and clouds of atmosphere dissipates solar radiation.[1, p. 469]

Specific intensity of solar radiation can be expressed by a differential equation whose resolution

is complex even applying approximations and numerical methods.[1, p. 469]

A lower complexity parametrisation is adopted instead, where the atmosphere is divided in a given number of layers and radiation is absorbed, scattered and reflected between each layer. Multiple reflections can occur from each layer but only one is considered in this model because successive reflections from amospheric layers have negligible intensities compared to the first one. [1, p. 470]

## 2.4 Numerical approach

Euler's method is used to solve the Ordinary Differential Equation (ODE) for the time-dependant temperature function.[1, p. 472] A solution for each atmospheric layer is evaluated, hence the resulting values are triplets of temperature, altitude (i.e. proxy for the atmospheric layer) and time (i.e. simulation time). Further information on storage and plotting of data are presented in Section B.1.

- 3 Results
- 3.1 Stability analysis
- 4 Discussion

### A Source code

In this section the C++ code used to obtain the results presented in this work is shown and commented.

#### A.1 Classes

# B Supplementary information

#### B.1 Plotting

Software Gnuplot is used to generate plots shown in this work. Output values from the simulation are stored in a DAT file with the following structure, line spacing between data blocks is important:

Value N\_t is not fixed a priori since it is the number of temporal steps needed to reach convergence, instead N\_P is the number of atmospheric layers.

#### C Mathematical derivations

In this appendix mathematical derivations of some ancillary results and formulae used in the main text are explicitly shown.

# C.1 Relation between pressure and altitude

A general result regarding planetary atmosphere is that atmospheric pressure decreases with increasing altitude. Theoretical relations which approximate this behaviour can be obtained. Pressure P depends linearly from altitude z if density is assumed constant:

$$dP = -\rho g dz \iff P(z) = P_0 - \rho g(z - z_0)$$
, (1)

and point  $(z_0, P_0)$  is a reference point in the atmosphere.

# References

[1] V. Ramanathan and J. A. Coakley Jr., "Climate modeling through radiative-convective models," Reviews of Geophysics, vol. 16, no. 4, pp. 465-489, 1978. DOI: https://doi.org/10.1029/RG016i004p00465. eprint: https://agupubs.onlinelibrary.wiley.com/doi/pdf/10.1029/RG016i004p00465. [Online]. Available: https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/RG016i004p00465.