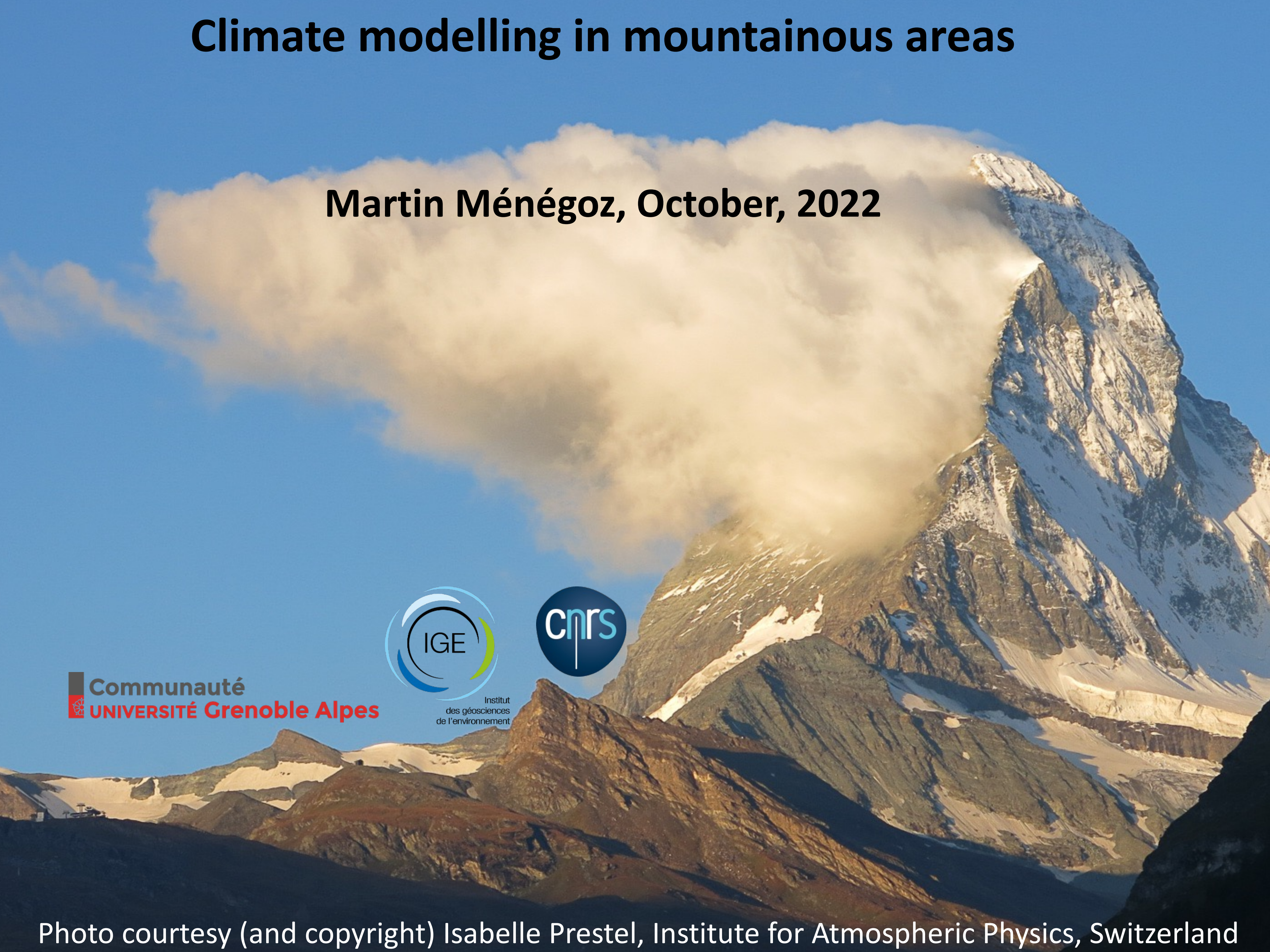
Climate modelling in mountainous areas

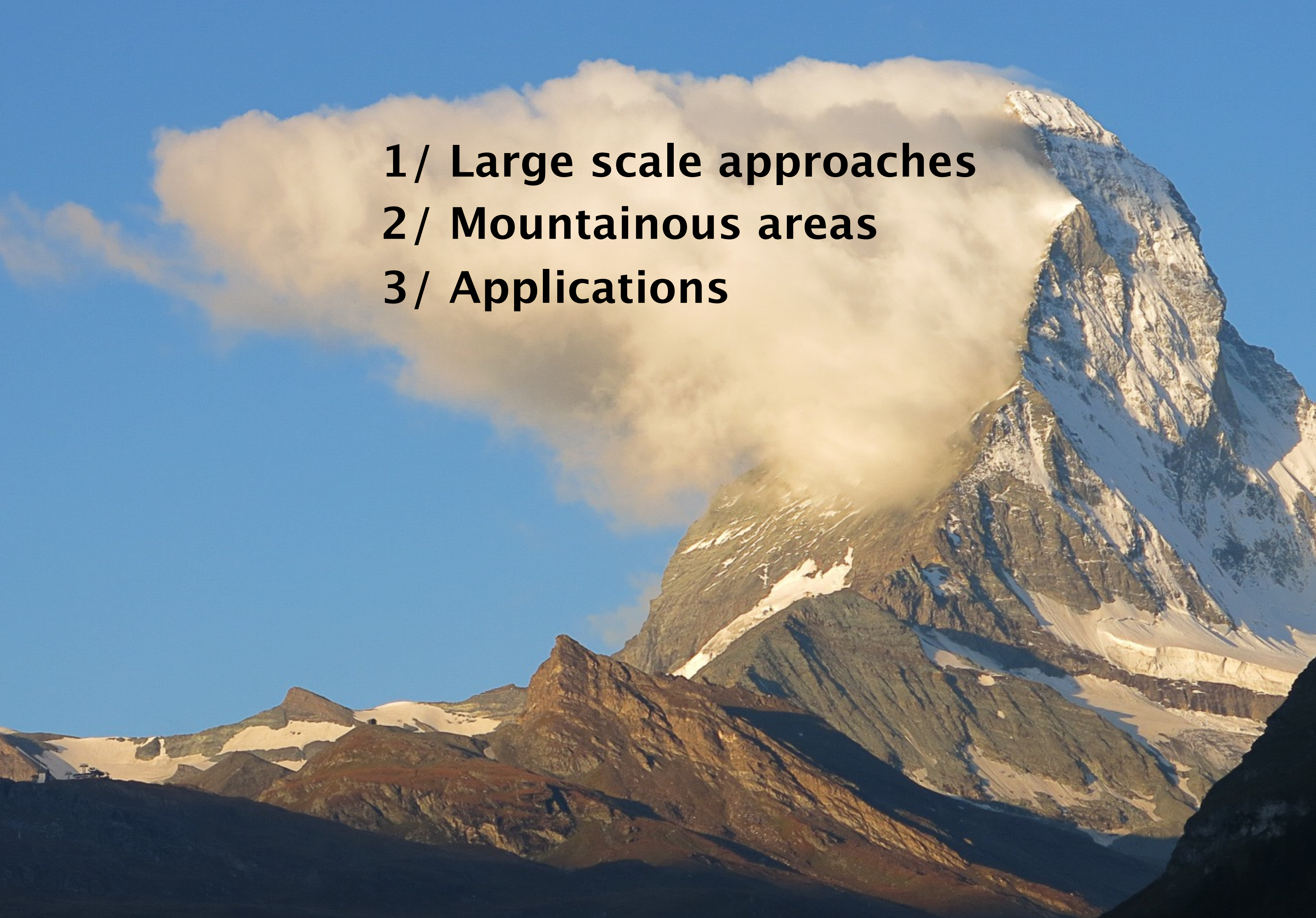
Martin Ménégoz, October, 2022



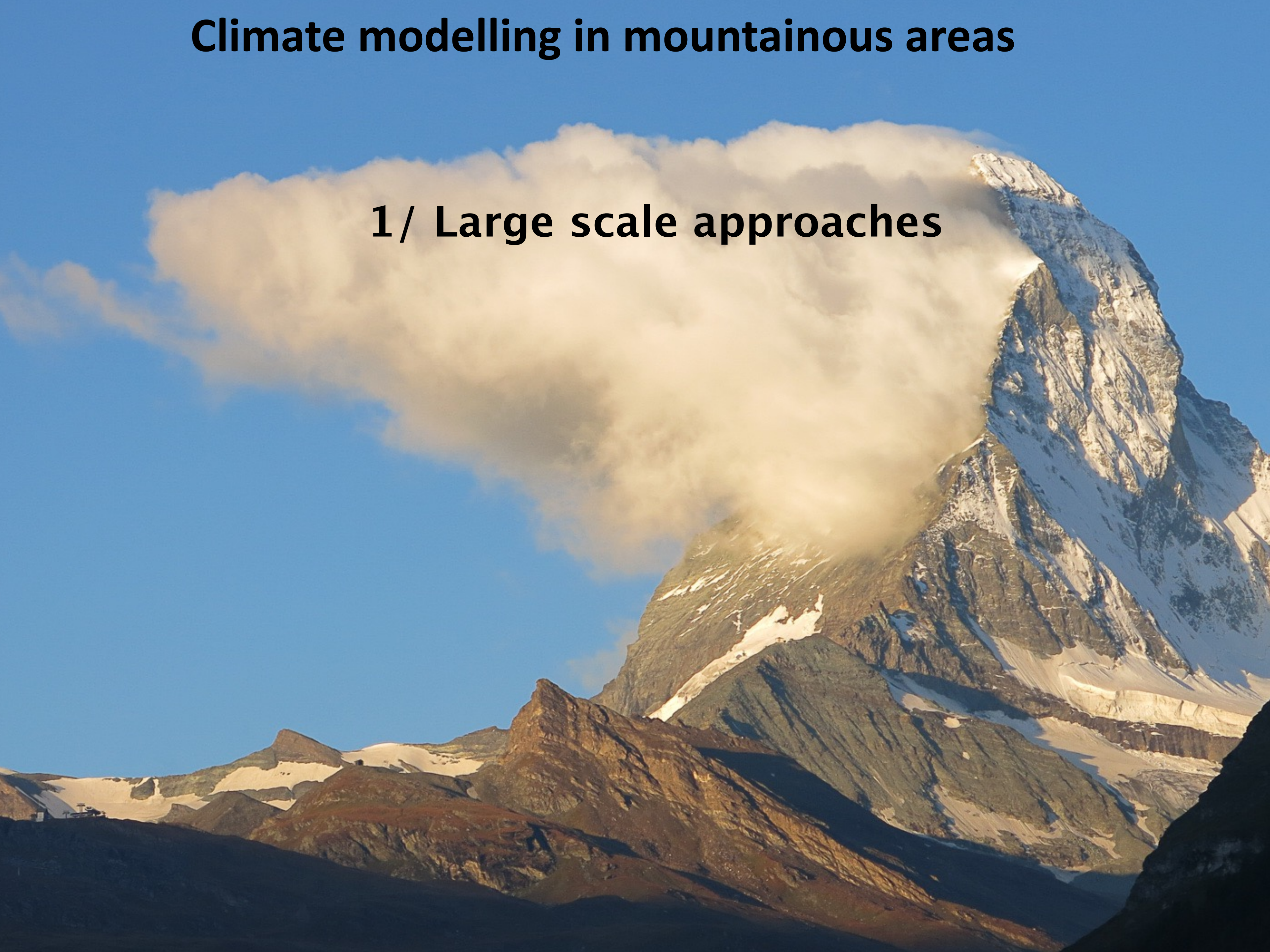
Communauté UNIVERSITÉ Grenoble Alpes
IGE Institut des géosciences de l'environnement
cnrs

Climate modelling in mountainous areas

1/ Large scale approaches
2/ Mountainous areas
3/ Applications



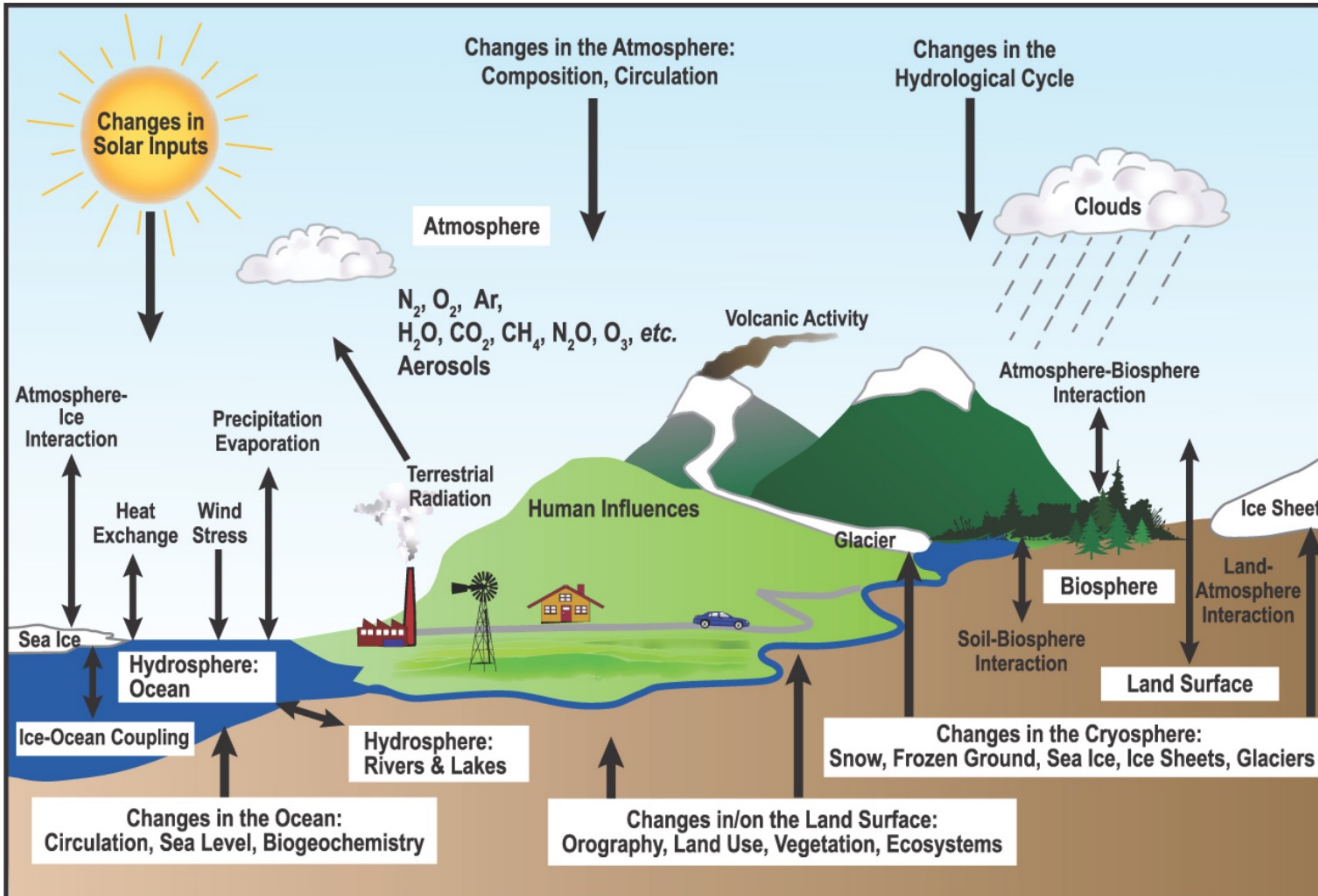
Climate modelling in mountainous areas



1/ Large scale approaches

The climate system

Ocean <-> Atmosphere <-> continental surfaces <-> other components



Source : IPCC AR4 (2007) FAQ 1.2 Fig. 1.

Definitions

Climate = Mean state + climate variability

Climate variability = internal variability + external forcings

External forcings = anthropogenic forcings + natural forcings

Natural variability = natural forcing + internal variability

Definitions

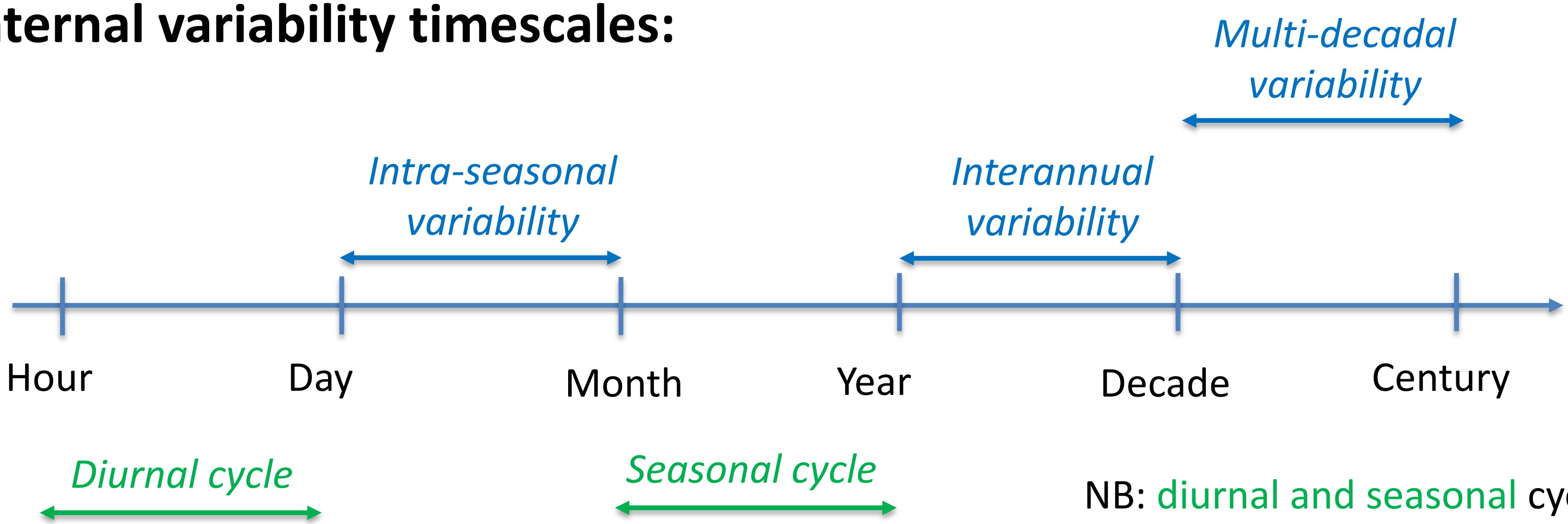
Climate = Mean state + climate variability

Climate variability = internal variability + external forcings

External forcings = anthropogenic forcings + natural forcings

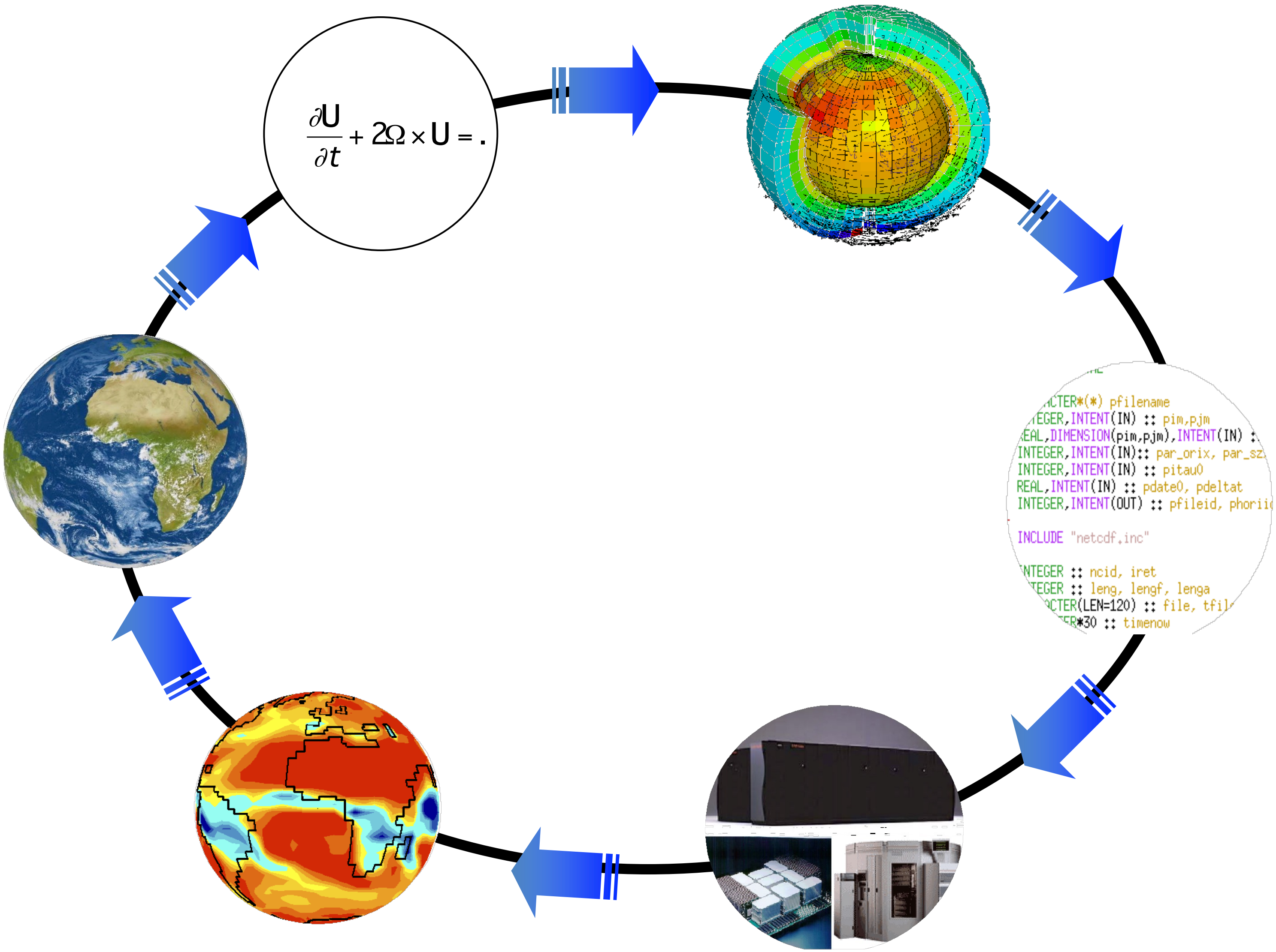
Natural variability = natural forcing + internal variability

Internal variability timescales:



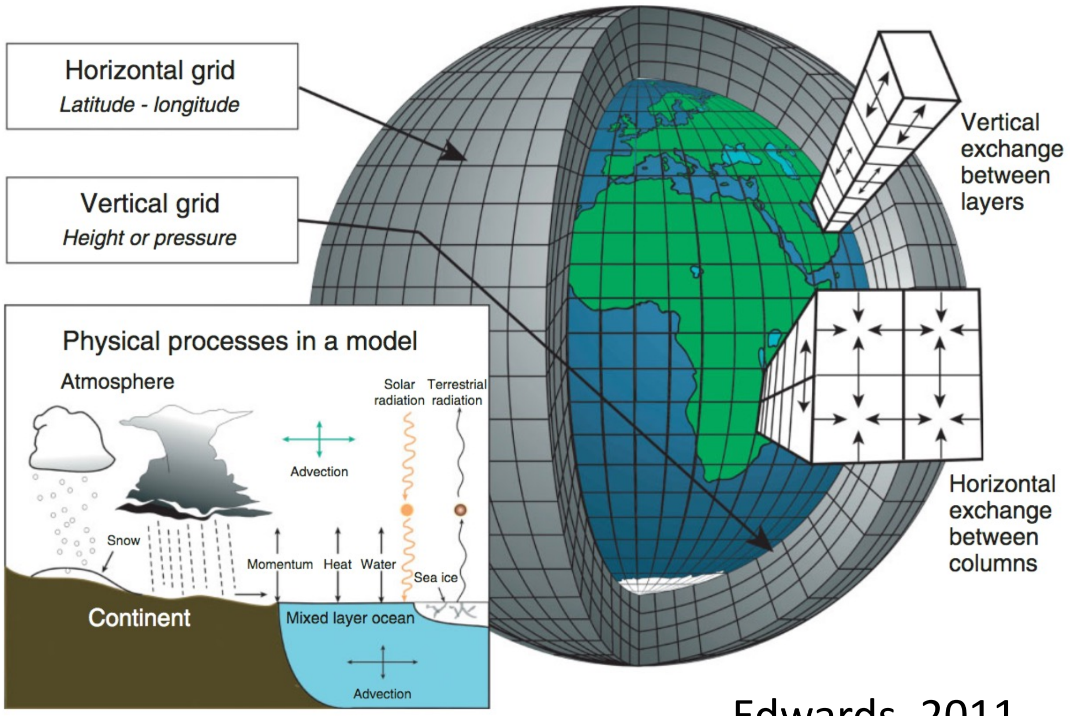
NB: diurnal and seasonal cycles are not internal variability

Model developpments



Model developments

- Space (x,y,z) and time (t) discretisation
- 3D resolution of dynamical processes (u,v, and tracers)
- 1D physical scheme : surface energy balance, radiative transfert, thermal conduction in the ground…
- Sub grid-scale parametrisations: cloud fraction, snow cover fraction, atmospheric convection, etc…



Edwards, 2011

Model developments

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- 1D physical scheme : surface energy balance, radiative transfert, thermal conduction in the ground…
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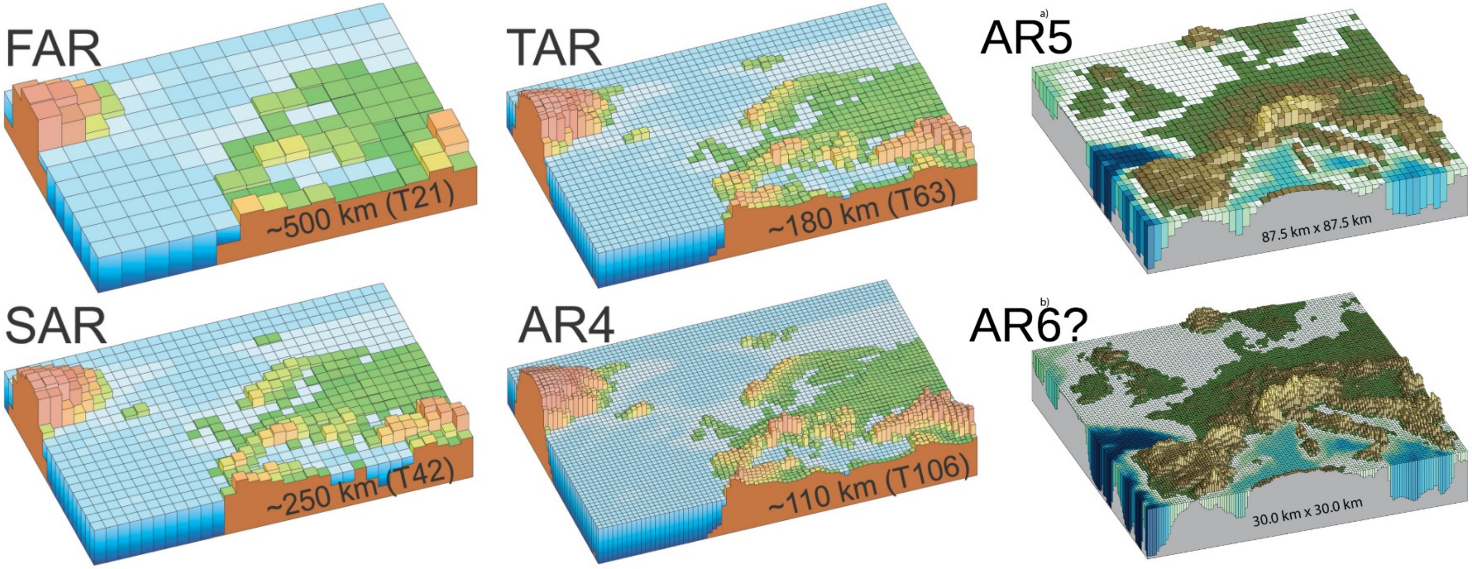
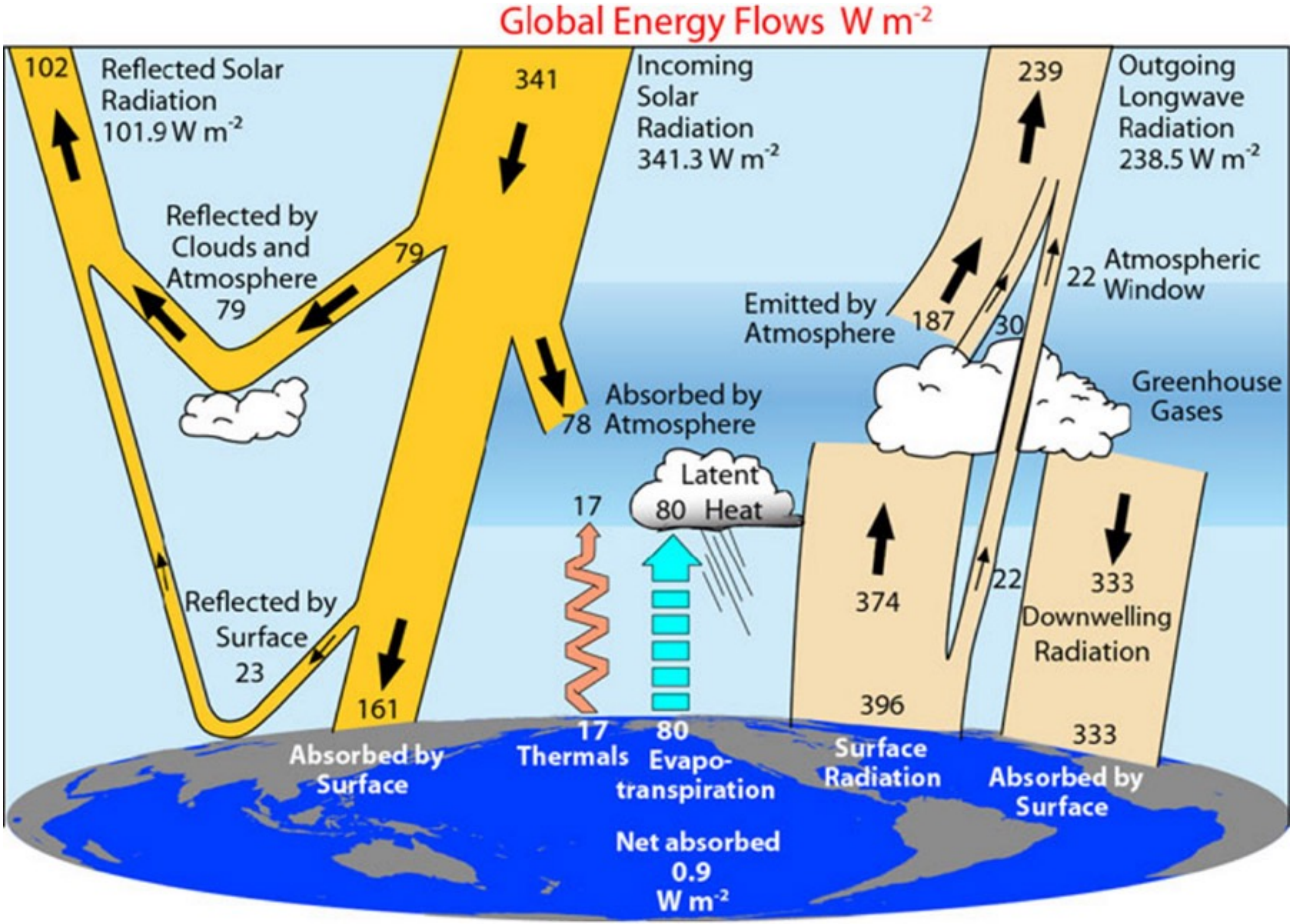
Grille typique des modèles lors des rapports successifs du GIEC
(1991, 1995, 2001, 2007, 2013, 202?)

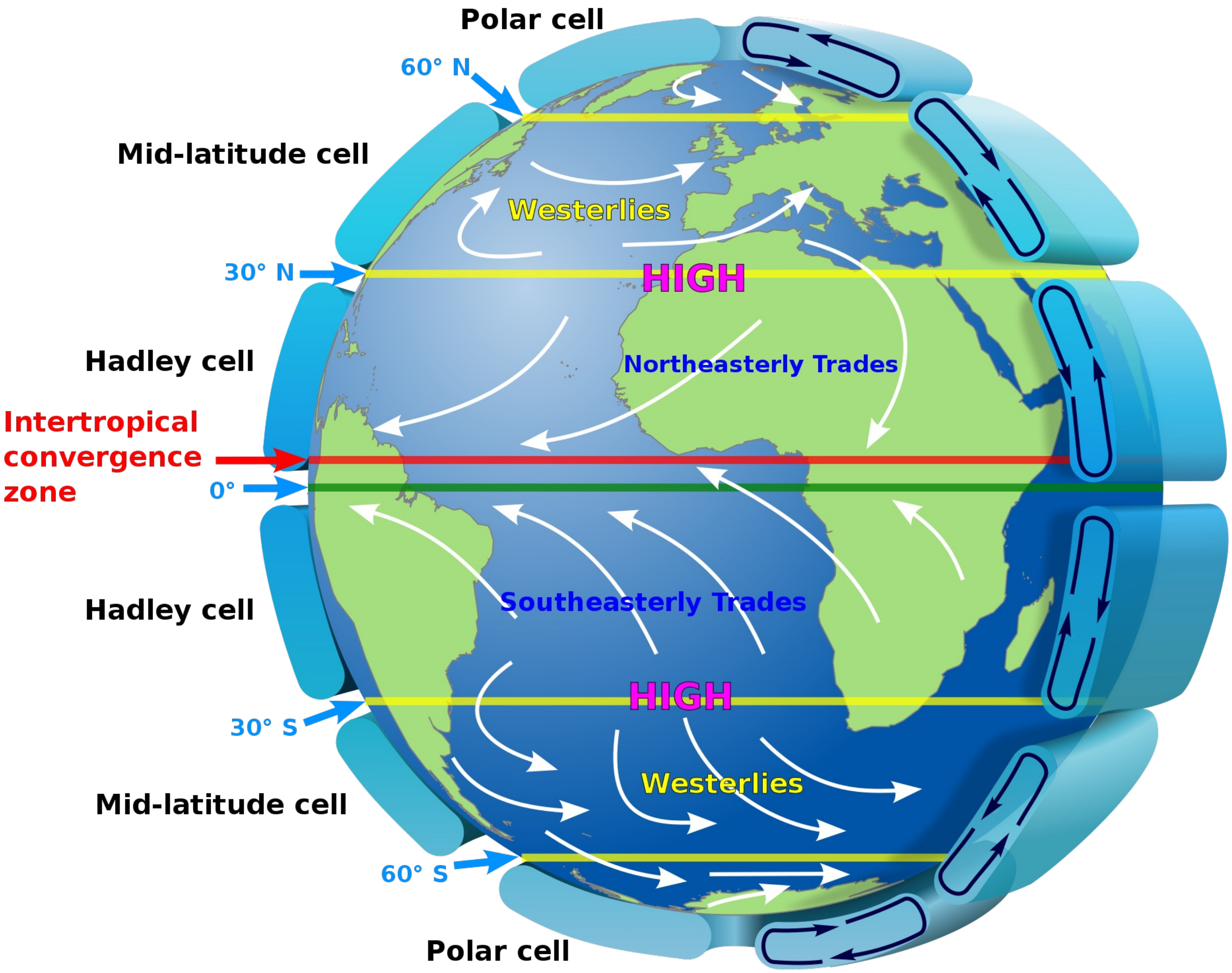


Source : IPCC AR4 (2007) Fig. 1.4 et AR5 (2013) Fig 1.14.



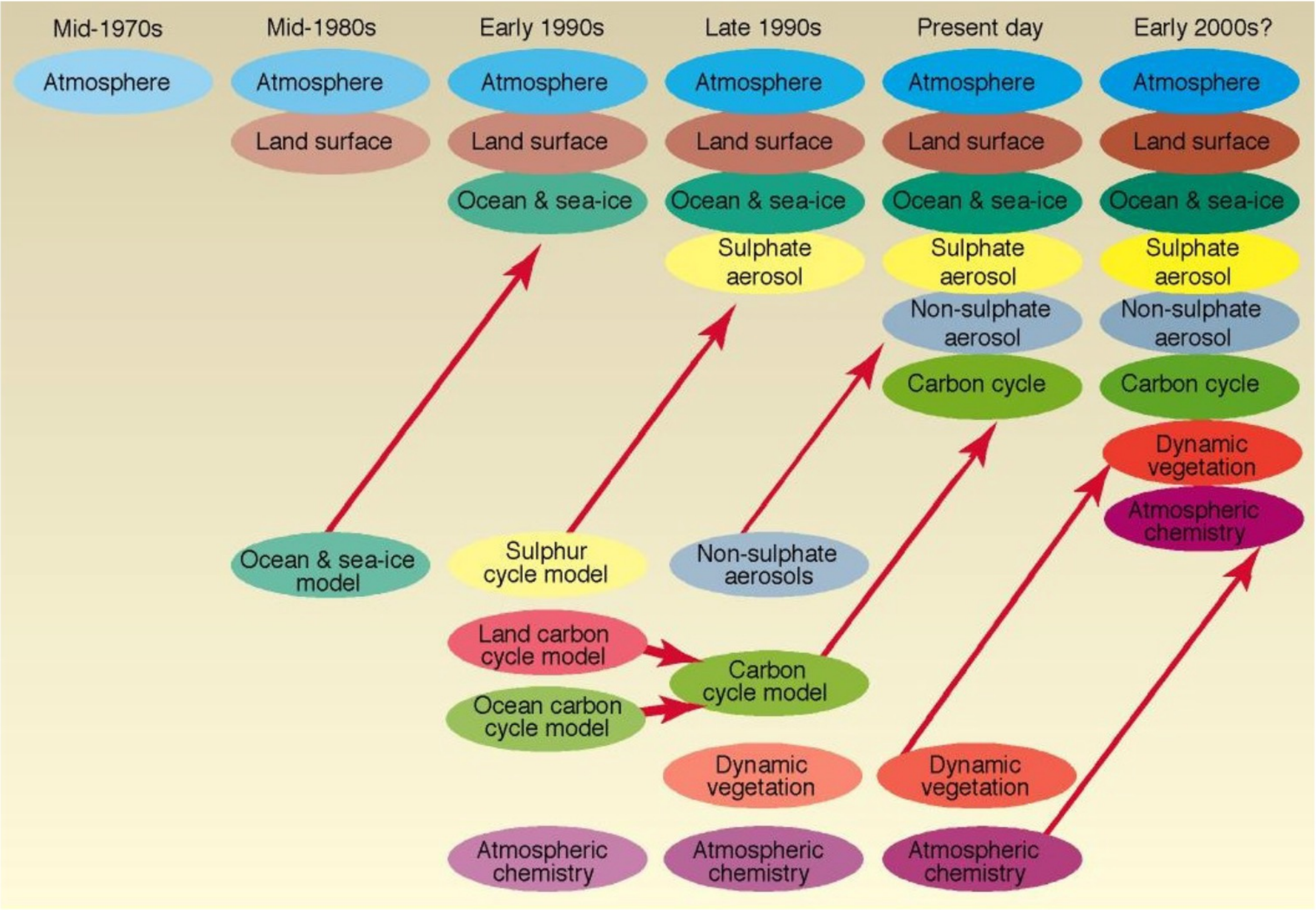
Global annual estimation of the atmospheric energy balance,
Source : Kiehl et Trenberth (1997)

Atmospheric circulation



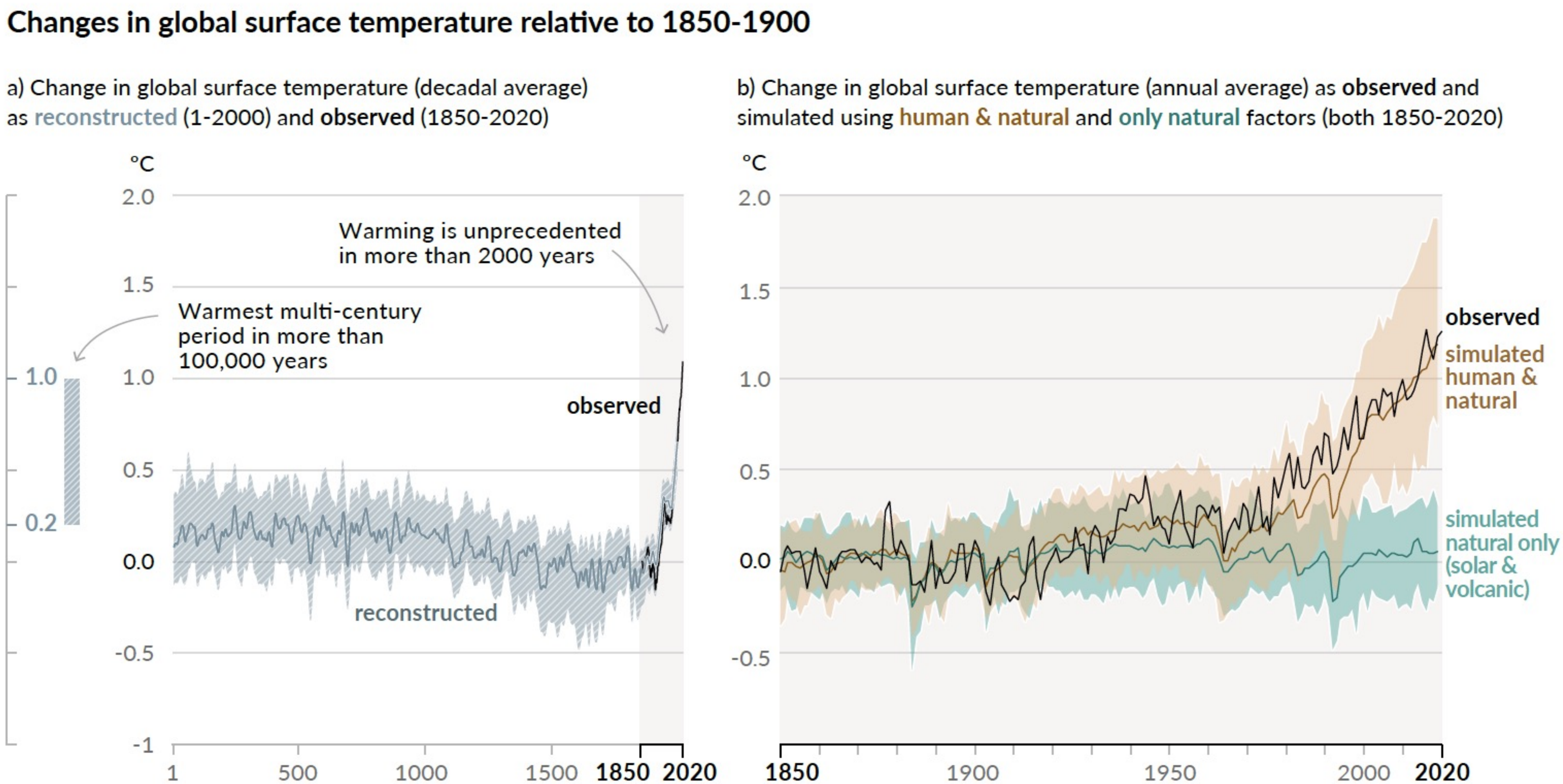
https://en.wikipedia.org/wiki/Atmospheric_circulation

From AOGCMs to ESMs



Source : IPCC TAR (2001), Fig. TS Box 3.

AOGCMs are used
-> to estimate the climate response to anthropohenic forcings

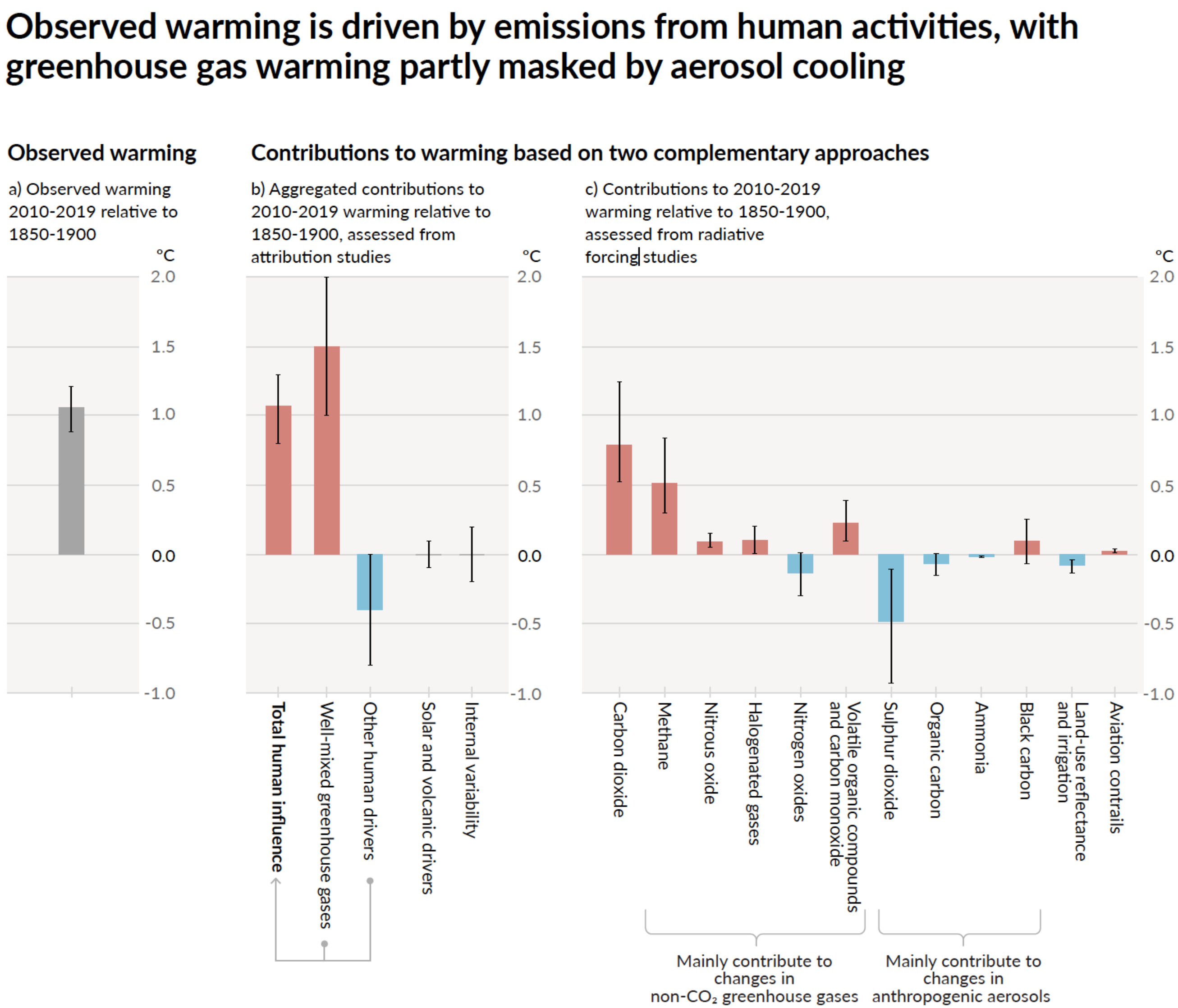


IPCC, AR6

AOGCMs are used

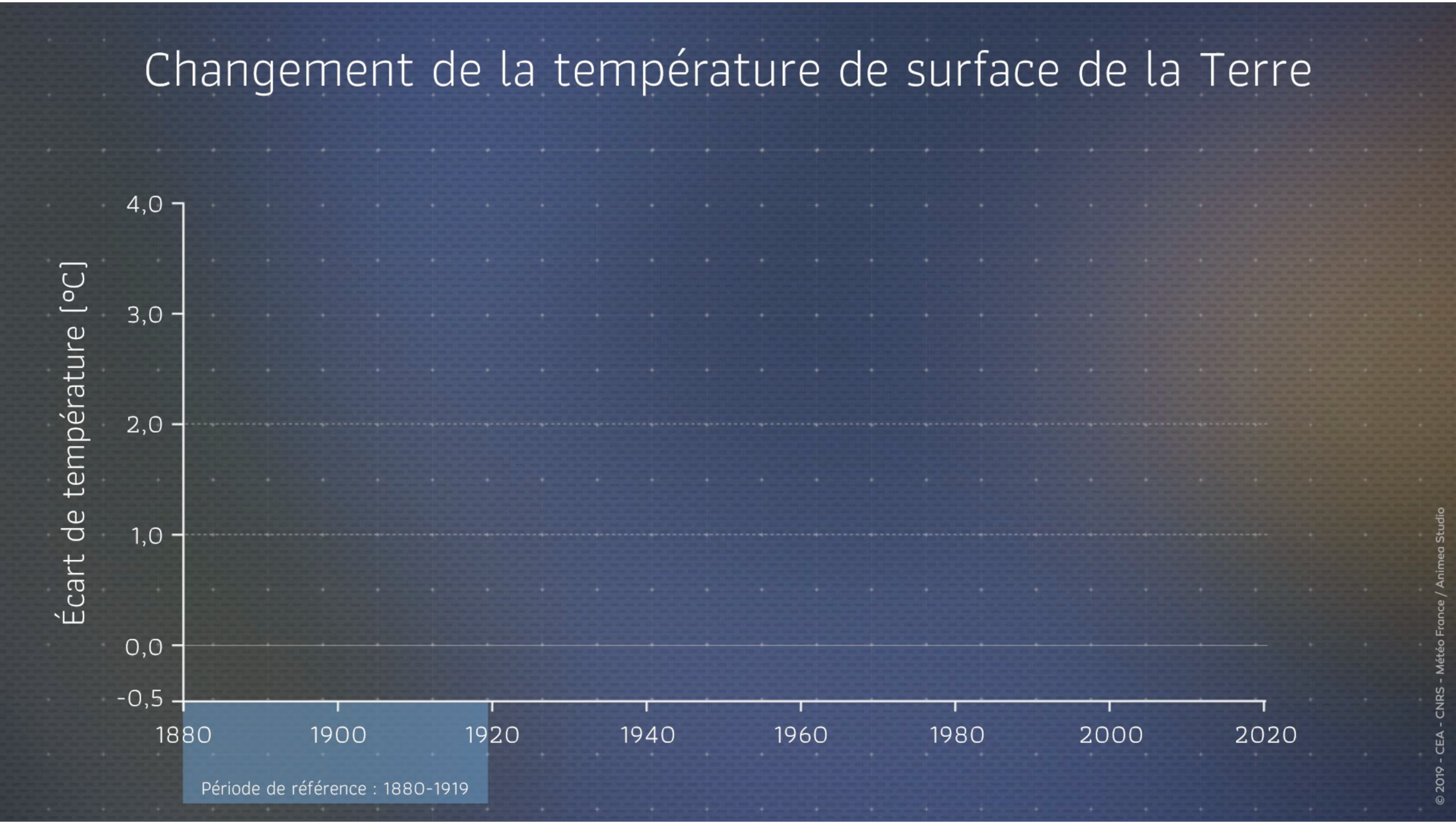
-> to estimate the climate response to anthropohenic forcings

-> to disentangle the imprints of the different forcings

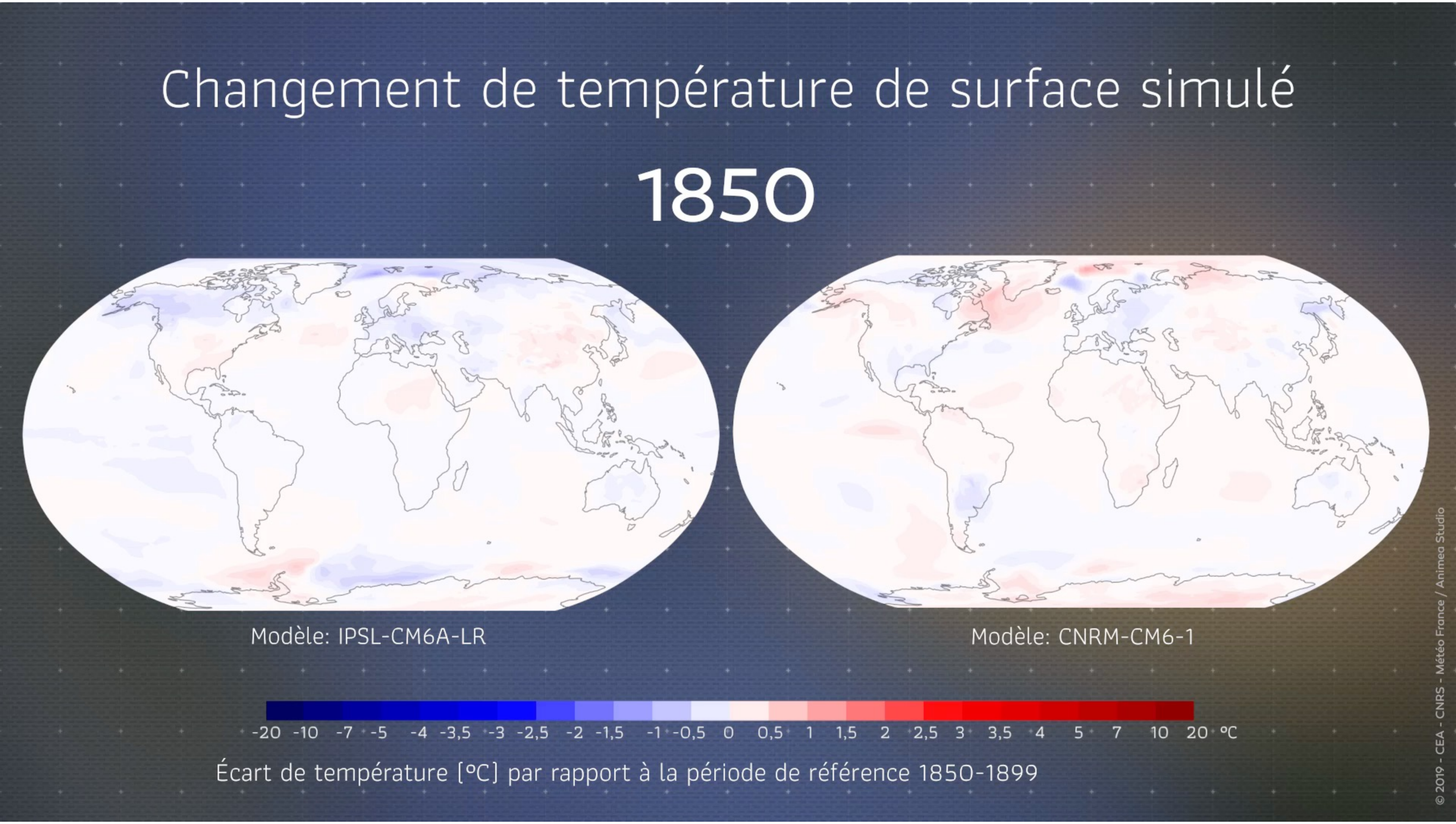


IPCC, AR6

Simulating global annual temperature changes

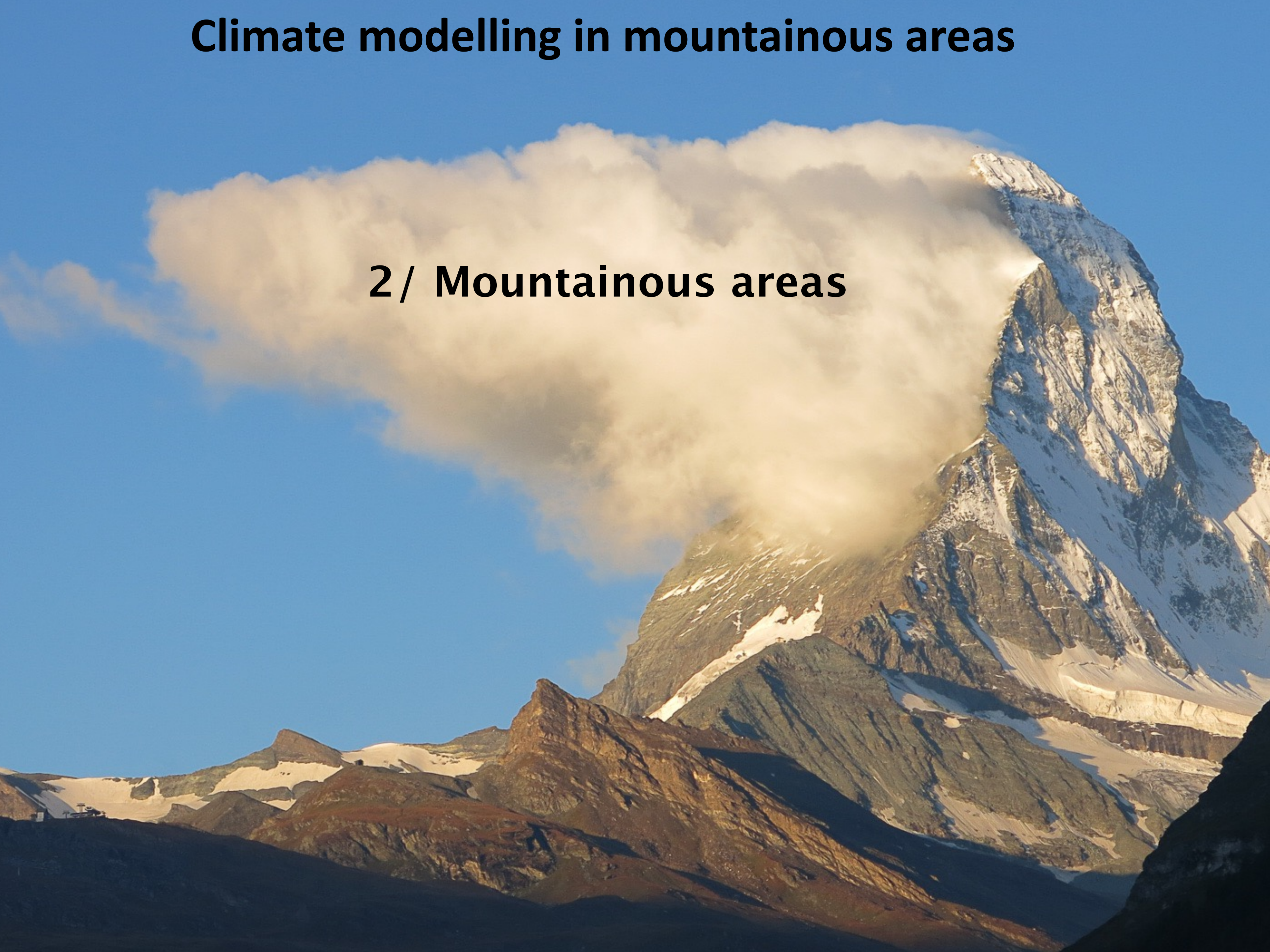


Simulating global annual temperature changes



Climate modelling in mountainous areas

2/ Mountainous areas



Climate modelling over mountain -> tricky!!!

Credit CHRISTOPH HORMANN / SCIENCE PHOTO LIBRARY

Clouds: one of the complex point in models

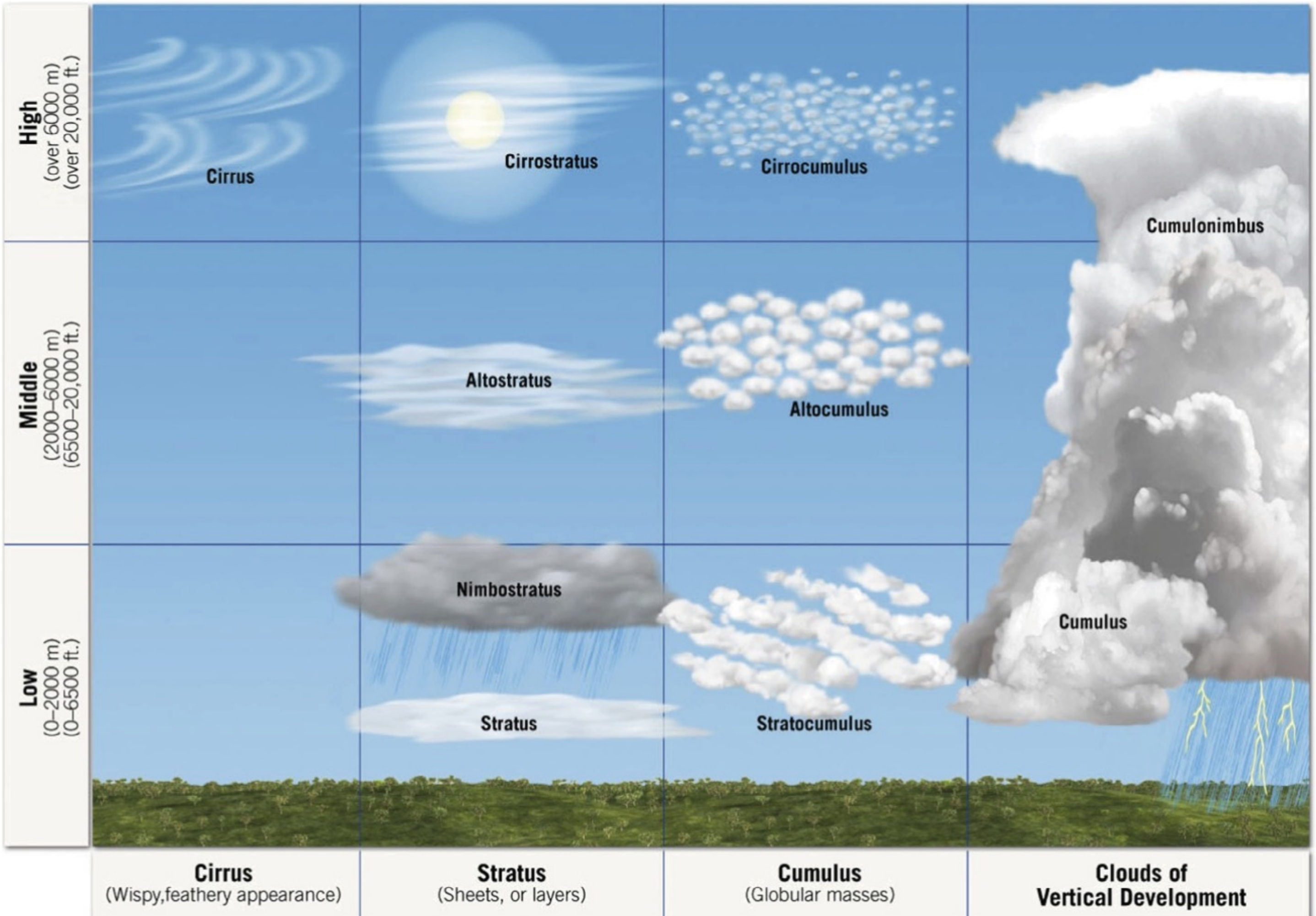
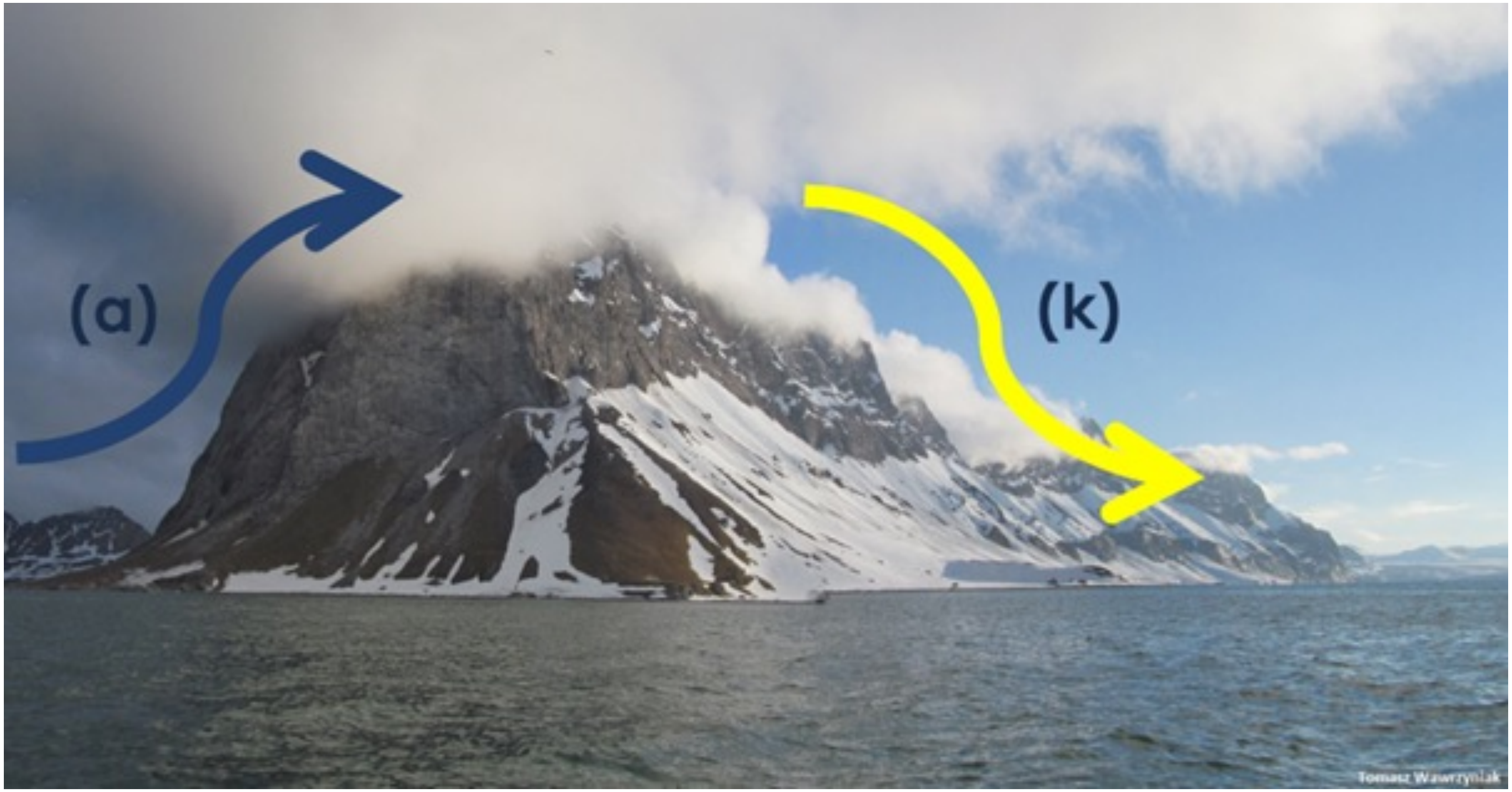


Figure 1 Ten major cloud types. Reproduced from Lutgens et al., 2016. http://www.pearsonhighered.com/bookseller/product/Atmosphere-An-Introduction-to-Meteorology-The-Plus-MasteringMeteorology-with-eText-Access-Card-Package/9780321984425.page

In mountains, it’s even more complex!



Orographic clouds and foehn processes
(source: https://polarpedia.eu)

In mountains, it's even more complex!

(source: wikipedia)

In mountains, it's even more complex!

(source: wikipedia)

In mountains, it's even more complex!

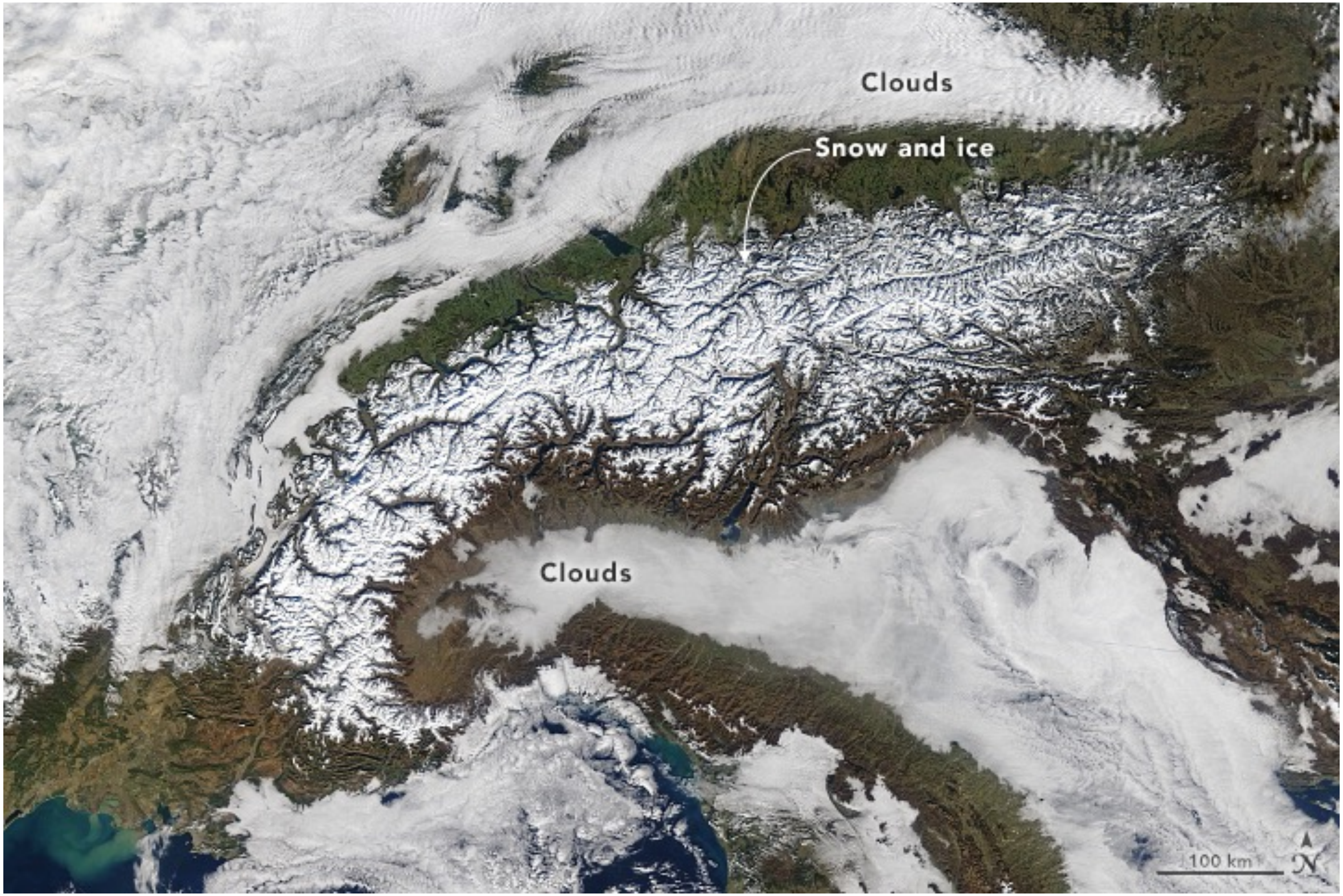
Photo courtesy (and copyright) Isabelle Prestel, Institute for Atmospheric Physics, Switzerland

In mountains, it's even more complex!

Summer cumulonimbus in French Alps
Crédit : Laurence Martin via twitter
@meteocentreisere

Sea-cloud in Chamonix
Credit : Daniel Simpson

How to simulate the snow cover in mountainous areas?



Credits: Nasa

How to simulate aire pollution trapped in the valleys?

Winter atmospheric pollution in the Alps

Elevation-dependant warming?

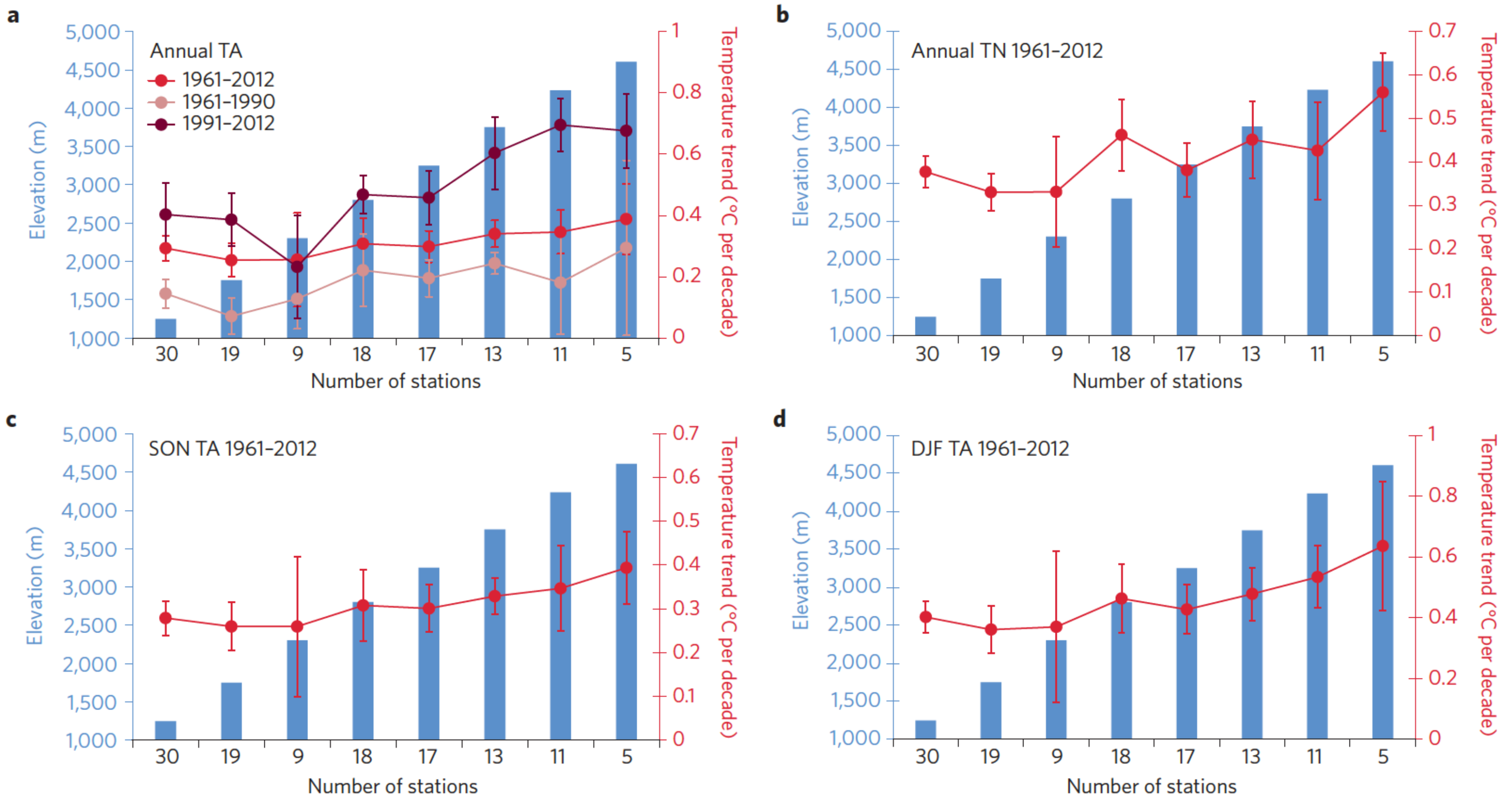
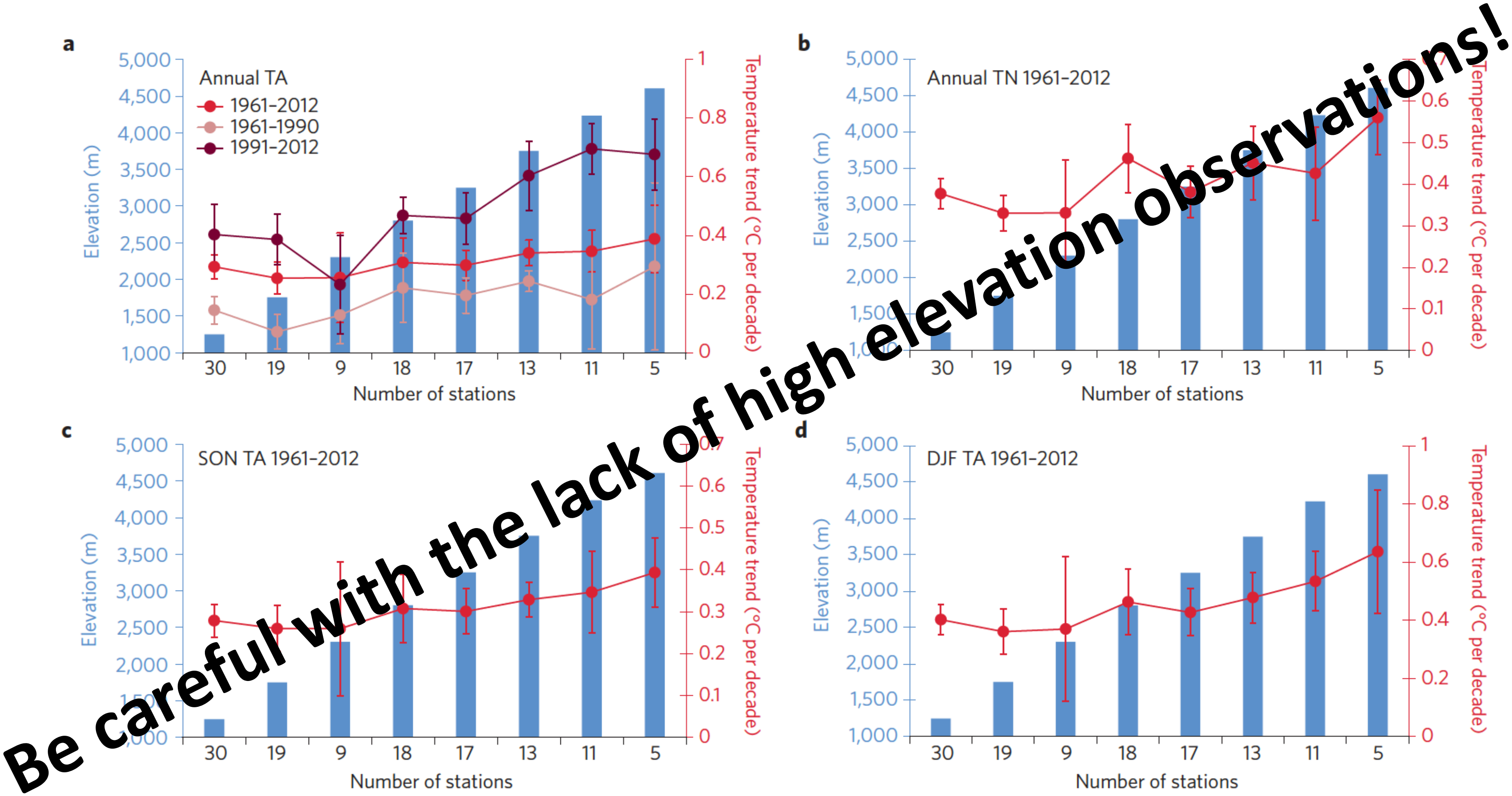


Figure 1 | Elevation-dependent warming over and around the Tibetan Plateau. a, Annual mean surface air temperature (TA) over 3 time periods. **b**, Annual mean minimum temperature (TN) from 1961–2012. **c**, Mean autumn (September-November, SON) surface temperature from 1961–2012. **d**, Mean winter (December–February, DJF) surface temperature from 1961–2012 . Bars represent elevation and trend magnitude is plotted on the *y* axis according to the 8 elevation ranks of 122 stations. The presentation format is similar to ref. 76 for ease of comparison. Error bars are based on 95% confidence intervals around the mean. The vertical scale for winter warming rate (**d**) and annual warming rate (**a**) have been adjusted to reflect the more rapid warming.

Pépin et al., 2015

Elevation-dependant warming?



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Pépin et al., 2015

Elevation-dependant warming? Which mecanisms?

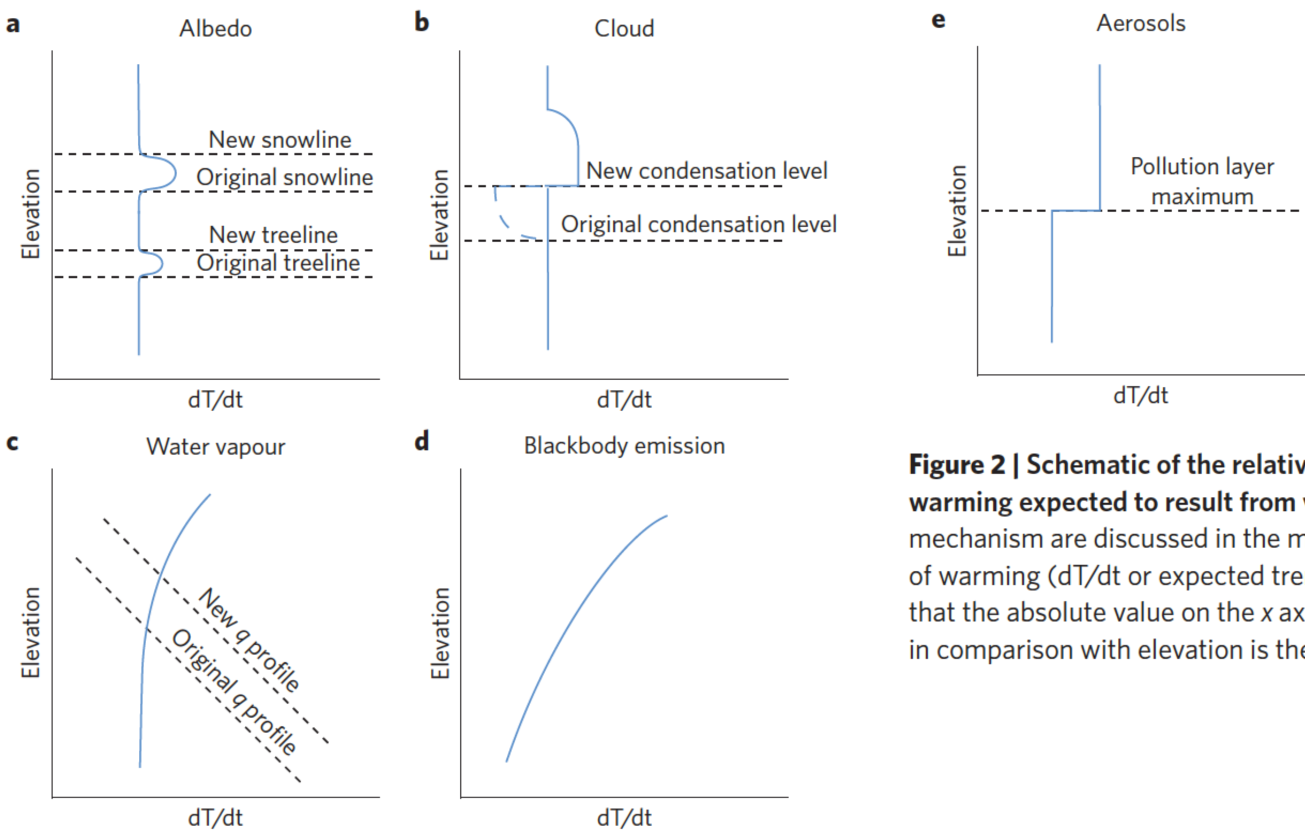


Figure 2 | Schematic of the relative vertical profile in atmospheric warming expected to result from various mechanisms. The details of each mechanism are discussed in the main text. The *x* axis represents the rate of warming (dT/dt or expected trend magnitude). The curves are relative in that the absolute value on the *x* axis is unimportant; the shape of the signal in comparison with elevation is the principal focus.

Pépin et al., 2015

GCMs biases in HMA in current climate models

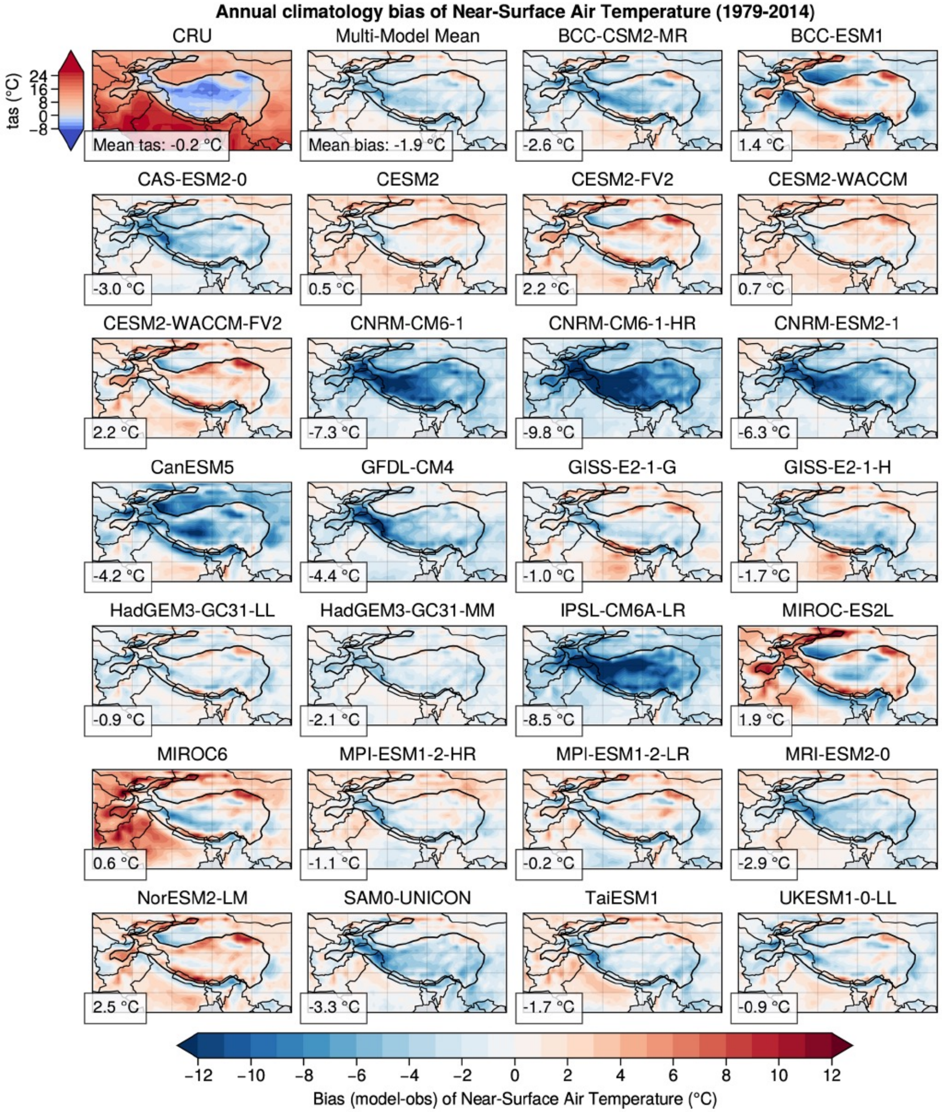


Figure 3. Annual bias (model minus observation) computed over 1979–2014 for temperature, except the top left panel that shows the climatology estimated from the CRU observation, used as the reference for the bias computation. The panel located at the right side of the CRU observation shows the bias of the multimodel mean based on the 26 models shown in the figure. The black contour shows the political frontiers and the bold black line the HMA domain located above 2500 m a.s.l., for which the spatial average of the bias is given in the bottom left of each panel.

Lalande et al., 2021

GCM biases in HMA: seasonal contrasts

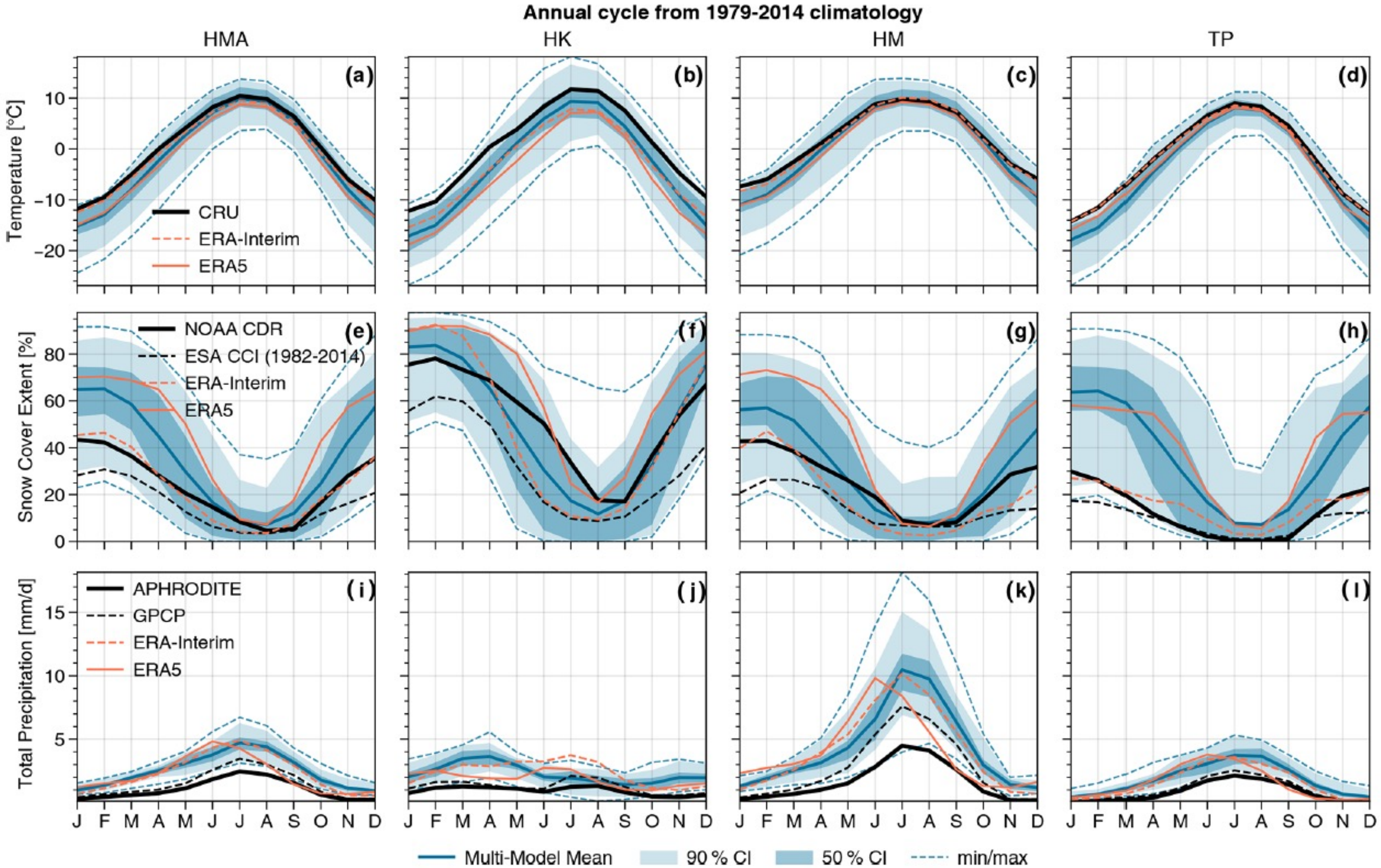


Figure 2. 1979–2014 climatology of the annual cycle of temperature **(a–d)**, snow cover **(e–h)** and precipitation **(i–l)** averaged over HMA **(a, e, i)** HK **(b, f, j)** HM **(c, g, k)** and TP **(d, h, l)**, excluding the surface area located below 2500 m a.s.l. (red contours in Fig. 1). The multimodel mean (dark blue line) is shown with the 50 % confidence interval (CI, dark blue shading), the 90 % CI (light blue shading) and the minimum and maximum (dashed blue lines) of the ensemble. The black curves correspond to the observational datasets: CRU, NOAA CDR and APHRODITE, respectively, for temperature, snow cover and precipitation. The ERA-Interim and ERA5 reanalyses are shown, respectively, with the dashed and solid orange curves. GPCP and ESA CCI datasets are also shown for snow cover and precipitation respectively (dashed black line). The ESA CCI covers only the 1982–2014 period.

Lalande et al., 2021

HMA trends: model and observations

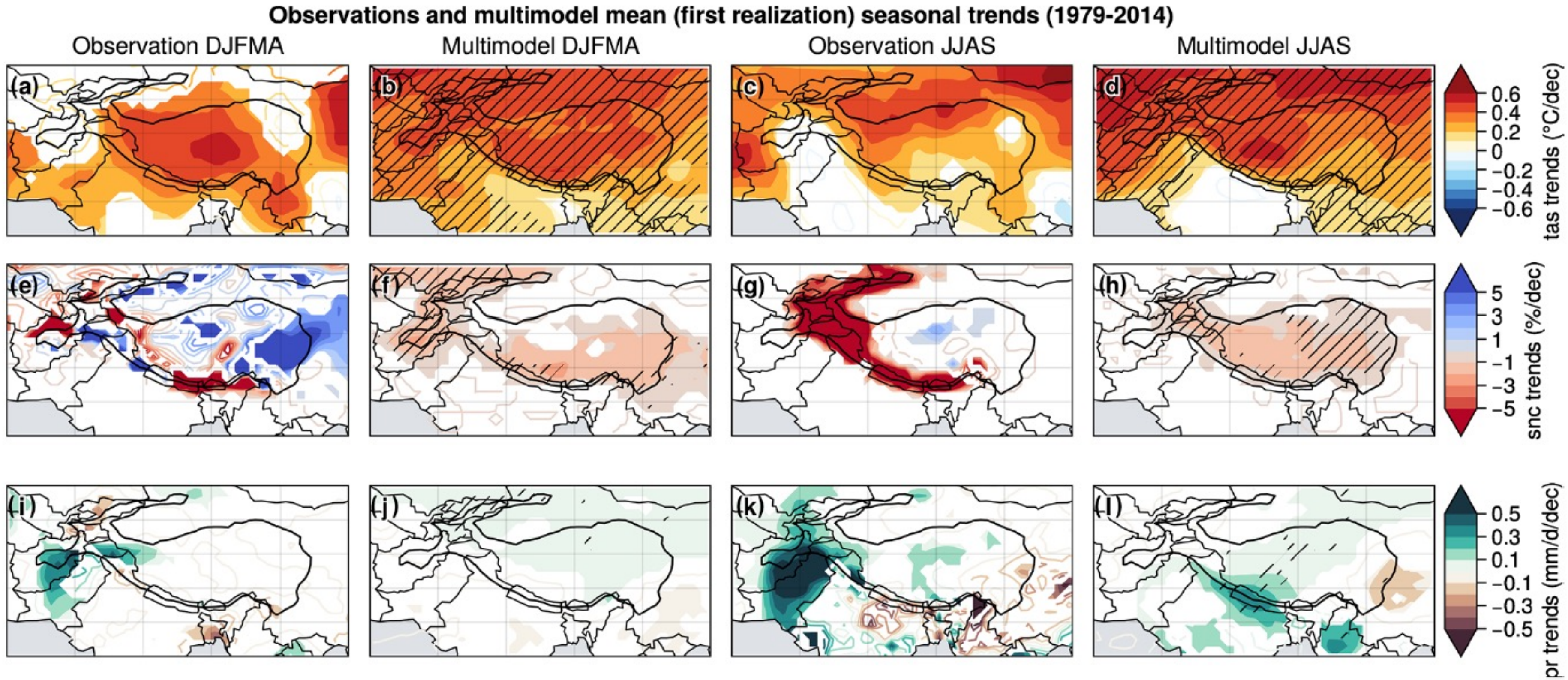


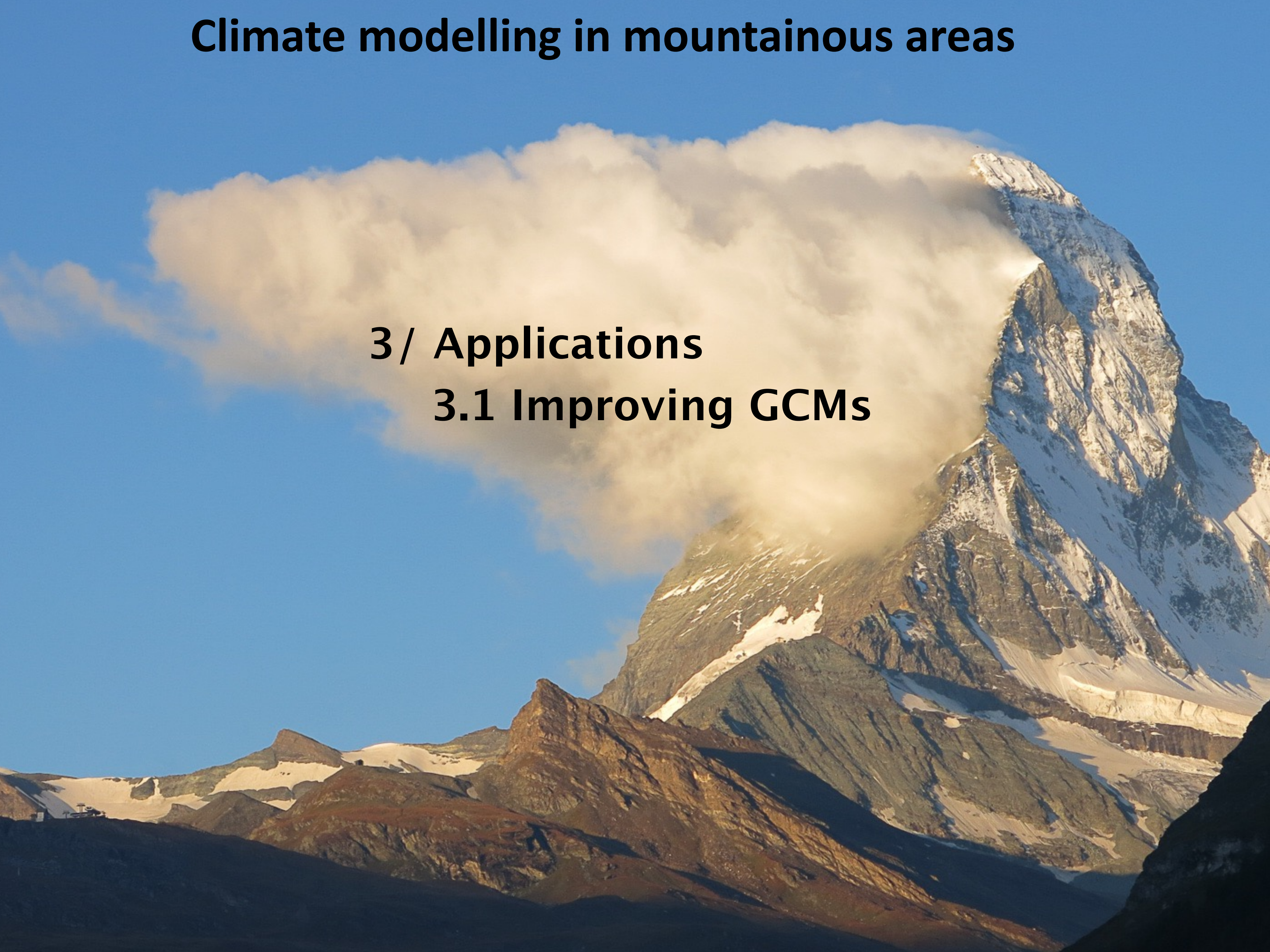
Figure 7. DJFMA (left) and JJAS (right) trends computed over 1979–2014 for temperature **(a–d)**, snow cover **(e–h)** and precipitation **(i–l)**. CRU temperature, NOAA CDR snow cover and APHRODITE precipitation observation trends (DJFMA: **a, e, i** and JJAS: **c, g, k**) are compared to the multimodel mean computed with the first realization for each model (DJFMA: **b, f, j** and JJAS: **d, h, l**). Contours are used for non-significant trends, shading for significant trends (p value > 0.05) and hatching for points where $> 80\,\%$ of the models agree on the sign of the trend.

Lalande et al., 2021

Climate modelling in mountainous areas

3/ Applications

Climate modelling in mountainous areas

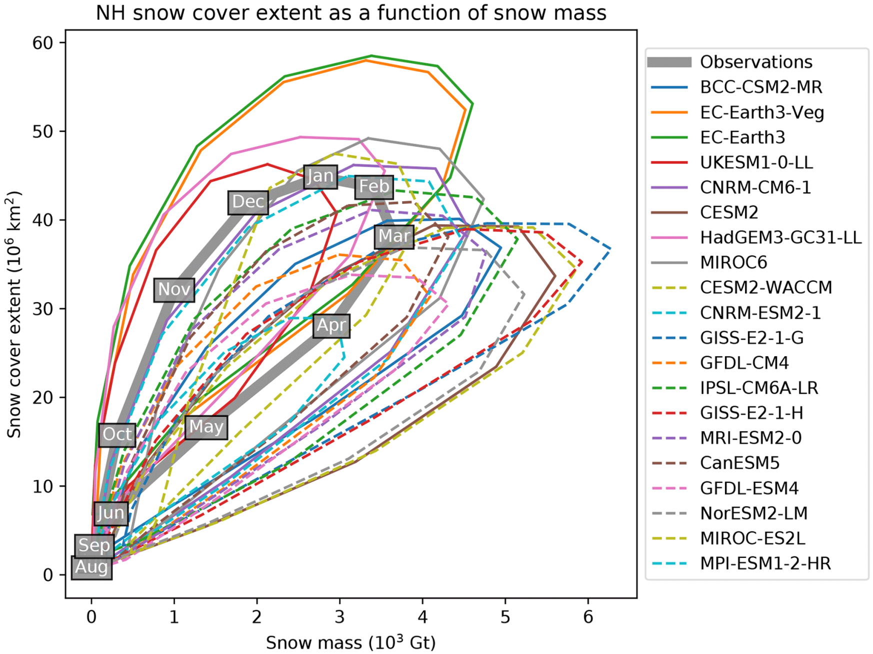


3/ Applications

3.1 Improving GCMs

Improving GCMs:

Example: subgridscale representation of the snow cover area

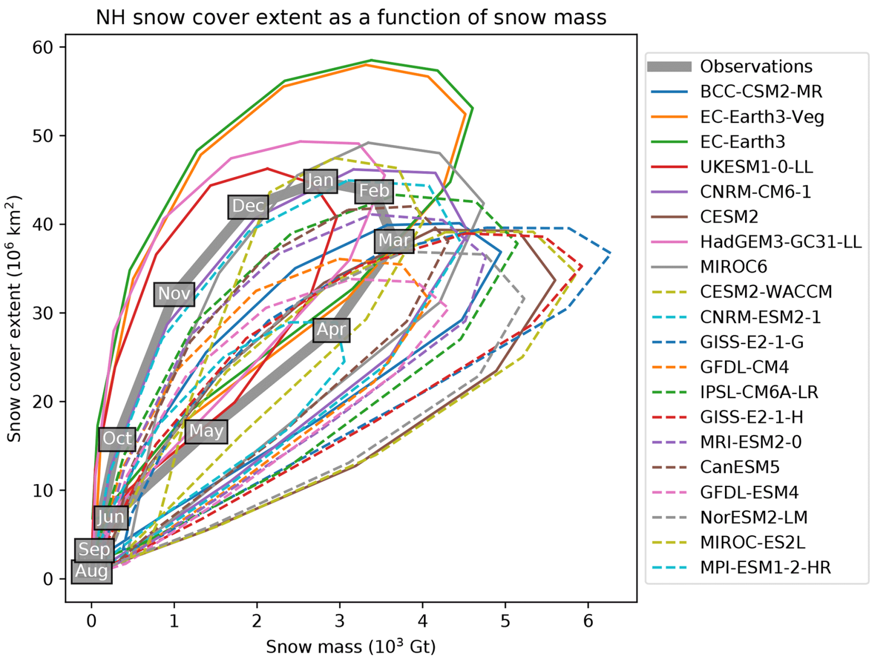


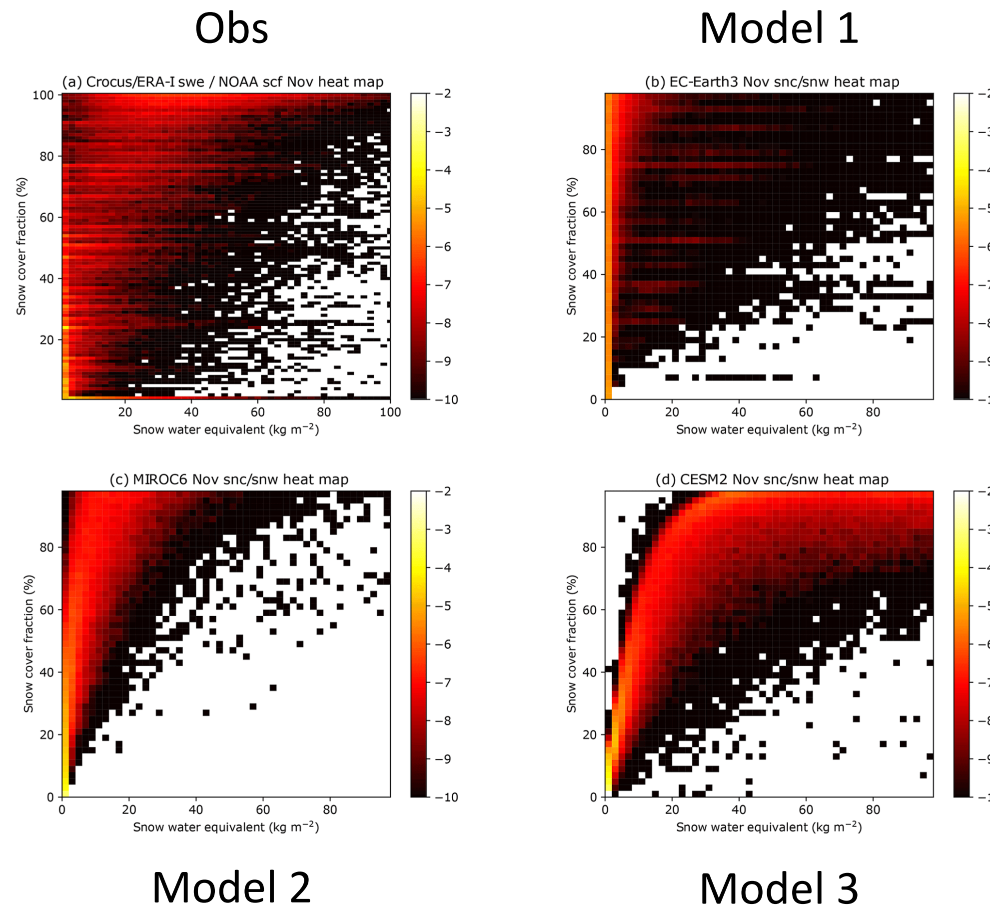
Snow cover fraction as a function of the snow depth.

Mudrick et al., 2020

Improving GCMs:

Example: subgridscale representation of the snow cover area





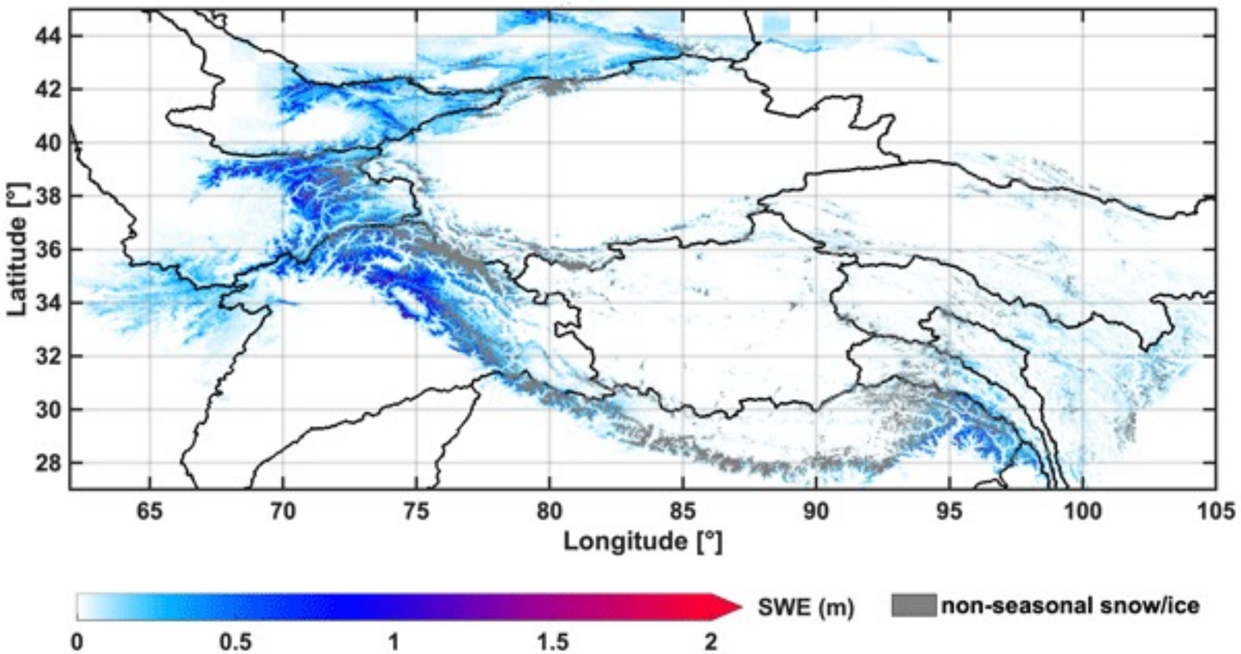
Snow cover fraction as a function of the snow depth.

-> Should depend on the topography!

Mudrick et al., 2020

Improving GCMs:

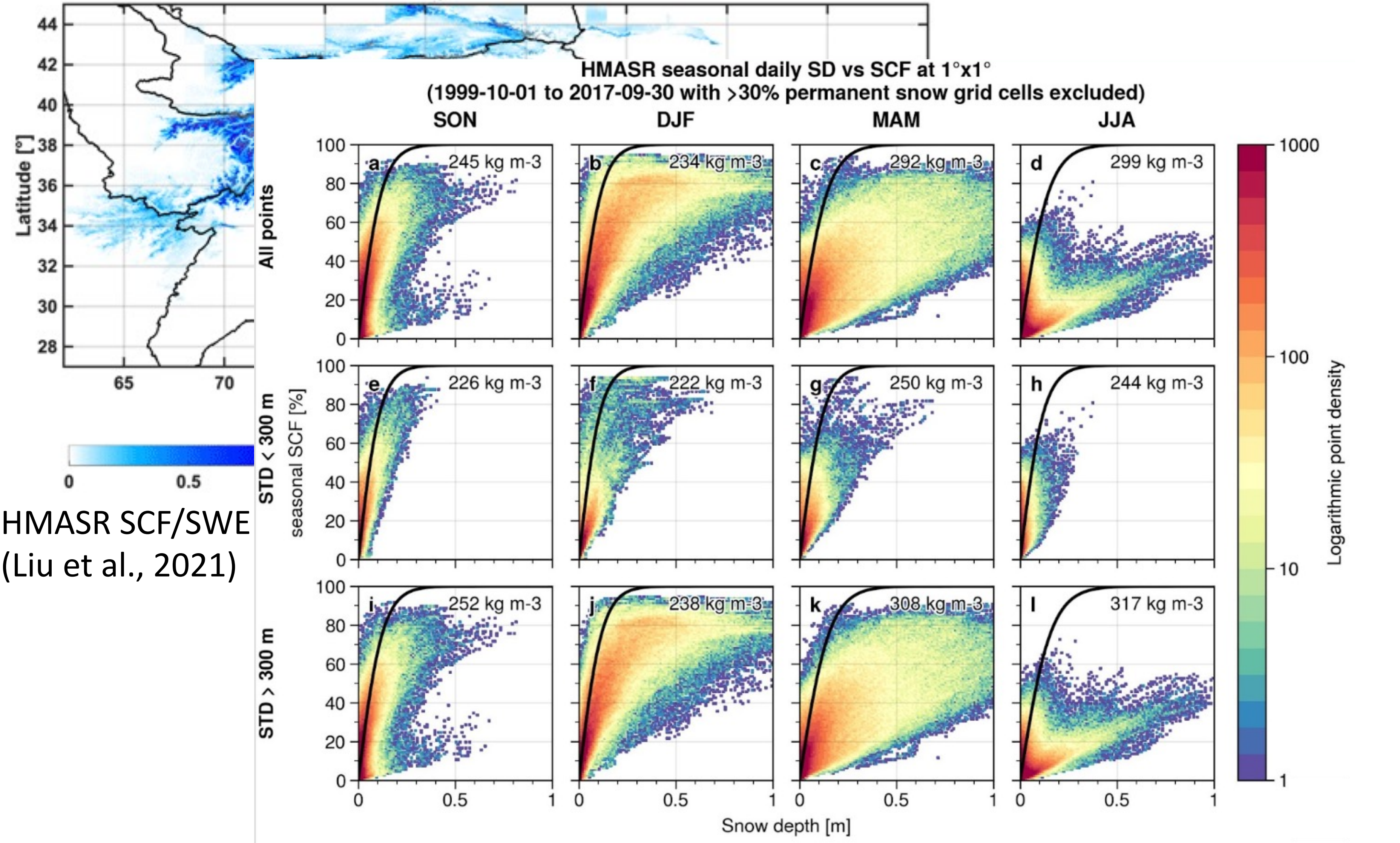
Example: subgridscale representation of the snow cover area



HMASR SCF/SWE reanalysis
(Liu et al., 2021)

Improving GCMs:

Example: subgridscale representation of the snow cover area



HMASR SCF/SWE
(Liu et al., 2021)

Lalande et al., in prep

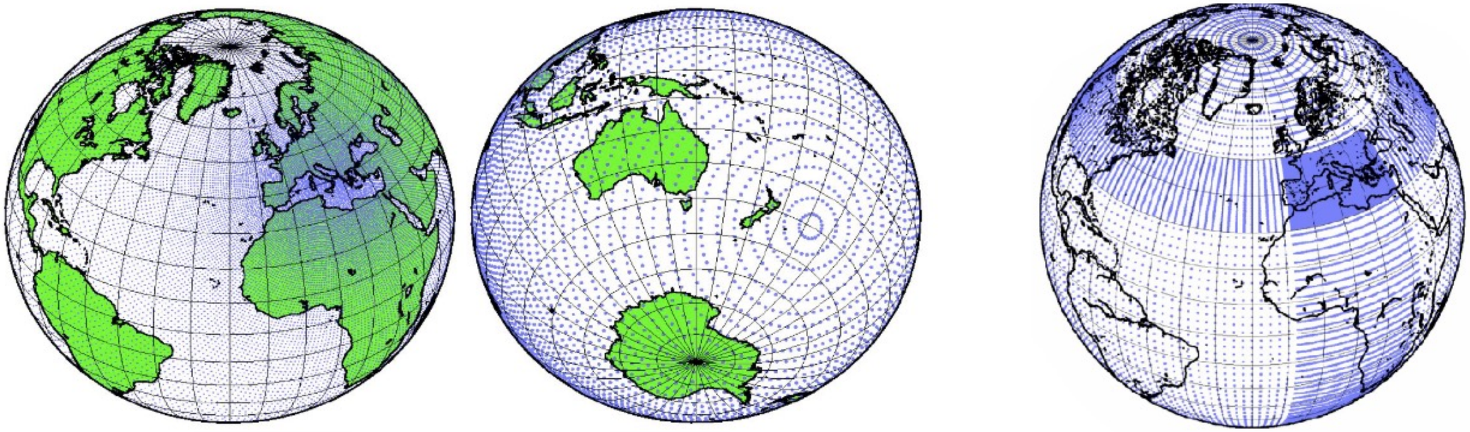
Climate modelling in mountainous areas

3/ Applications

3.2 Downscaling

Dynamical downscaling

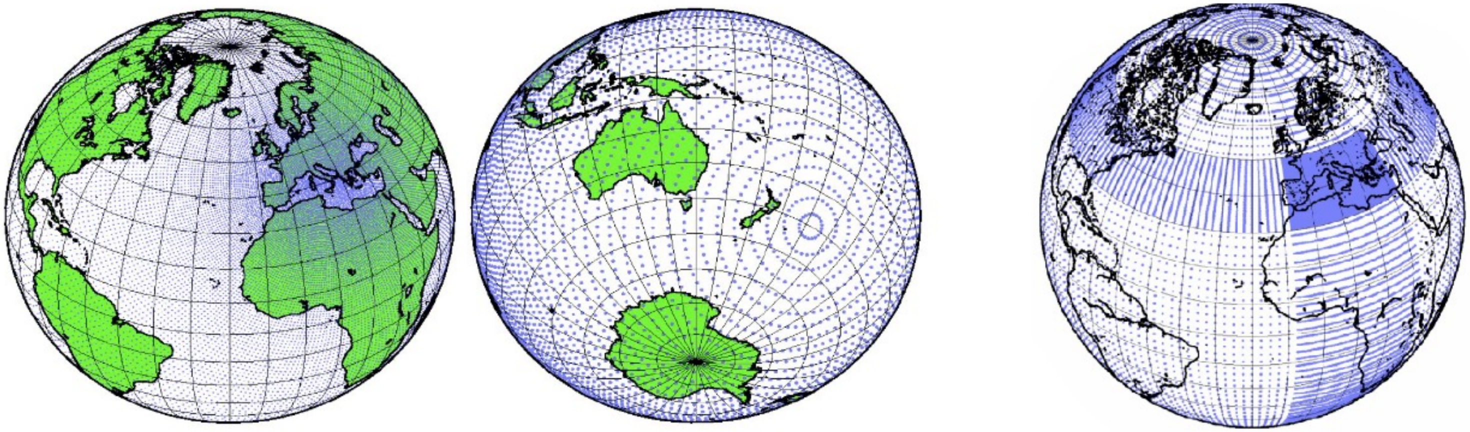
Stretched grid in GCMs
(e.g. LMDZ, ARPEGE)



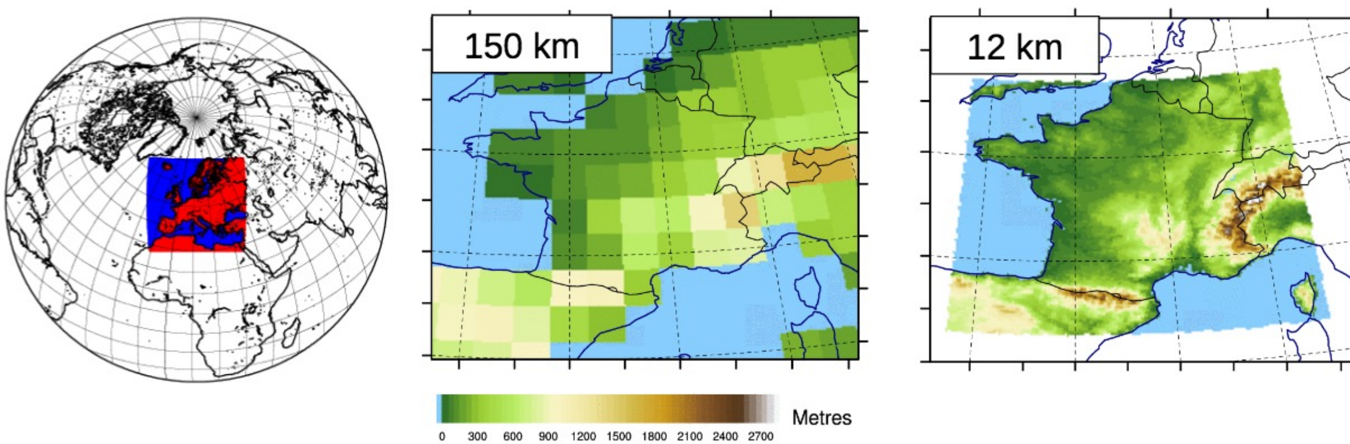
Courtesy B. Pohl, C. Cassou, J. Cattiaux

Dynamical downscaling

Stretched grid in GCMs
(e.g. LMDZ, ARPEGE)

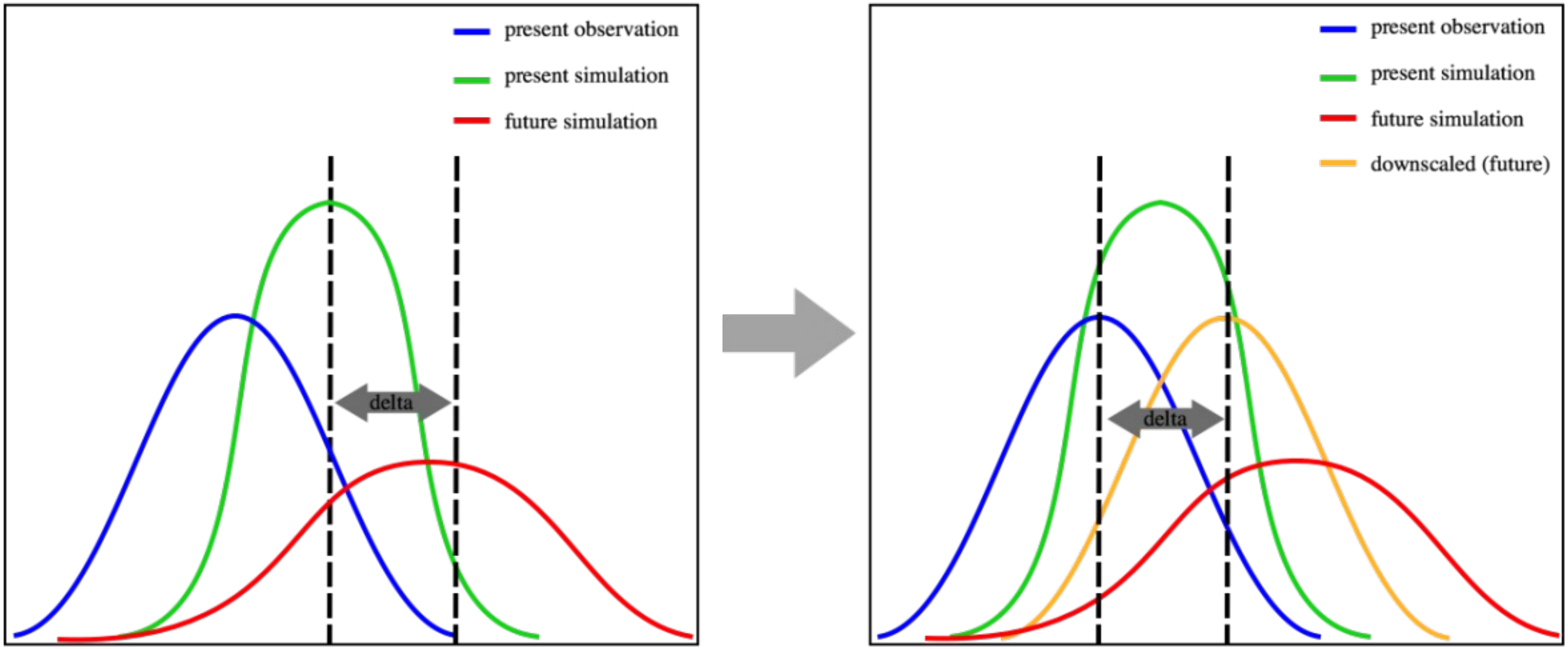


Regional Climate Models
(RCMs, limited area models)



Courtesy B. Pohl, C. Cassou, J. Cattiaux

Statistical downscaling



https://rcmes.jpl.nasa.gov/content/statistical-downscaling

Statistical versus dynamical downscaling

Dynamical	Statistical
Physical realism, catching feedbacks	Lack of physical realism
High computational cost, big data!	Low computational cost, possibility to produce large ensembles
RCMs reproduce a part of the GCMs biases trough lateral conditions	Biases and large scale - small scale teleconnections might be not stationnary
Spatial resolution still too coarse for mountainous areas	Possibility to get data at very local scale

Synthesis on the climat of HMA: https://www.encyclopedie-environnement.org/climat/glaciers-hautes-montagnes-asie-face-au-changement-climatique/

Thanks!!!