

# COMMON ELEMENTS

In this chapter some of the elements that are common for all the functions such as code format, instructions for use on various operating systems, elements of the input/output, and the GetData function are described.

## IRF Formatting:

The code has been written to fit the Initialize, Run, Finalize format, as shown in the Figure. The Main function for each program includes all the data structure preparation (variable declaration, data structure definitions, etc.) and the three functions called Initialize, Run, and Finalize. What each consists of is explained below:

- I The Initialize function contains the sub-functions that a) call the data from the user's inputted text file, b) compute the constant parameters that will be used by the Run function to carry out the calculations, and c) prepare the data for use later on.
- R The Run function contains sub-functions that actually do the calculations. In codes with a time loop, the Run function calls all the sub-functions needed in the time loop. In codes without any time loop, it is simply the function that carries out the calculation.
- F The Finalize function sends the output to the user. It organizes the data computed by the Run function, and prints it out to a tab separated text file.

```
#include <stdio.h>
#include <stdlib.h>
#include <math.h>
#define G 9.81

//Prototypes

int Initialize(double [], double *, double *, double *);
int Run(double *, double *, double *, double, double, double);
int Finalize(double, double, double);

int main (int argc, const char * argv[]) {
    double viscosity=0.0, spec_grav=0.0, grain_size=0.0, velocity=0.0;
    double Rep0, Rf0, date[3];

    Initialize(date, &viscosity, &spec_grav, &grain_size);
    Run(&velocity, &Rep0, &Rf0, viscosity, spec_grav, grain_size);
    Finalize(Rep0, Rf0, velocity);

    return 0;
}
```

In the main function of some models there is another function called SaveDataToMatrix, which is called after the Run function to store the data in a matrix.

## Input/Output:

The models are organized in folders. Each folder also contains the documentation that describes the input/output tab separated text files. The documentation consists of three or four text files that are called *Key for Inputs.txt*, *test.txt*, *Output.txt* and *Output1.txt*.

The *Key for Inputs* file explains how to prepare the input text file. In particular it contains information on a) the format of the text file, b) the units of the input parameters, and c) the physical meaning of the input parameters.

```

Key for Inputs.txt

f time step in flood days
N number of hydrograph entries [between 2 and 16 values]
c number of cycles per year
e bed elevation at downstream end [m]
r roughness height [m]
S initial bed slope
M number of intervals
I number of iterations before a print is made
P number of prints
D factor by which Deltap is multiplied for roughness height
n factor by which Deltap is multiplied for active layer thickness
r coefficient in Manning-Strickler relation
R submerged specific gravity of gravel
b bed porosity, grain size
U upwinding coefficient for local spatial derivatives in Exner equation (>0.5 suggested)
o coefficient for material transferred to substrate as bed upgrades

qv [m^2/s]          dtbf [m^2/s] # of time steps
qp [m^2/s]          dtbf [m^2/s] # of time steps
qn [m^2/s]          dtbf [m^2/s] # of time steps

.
.

[number should match the n given above]

Diameter [mm]    GSD Feed   GSD Surface   GSD Substrate
Diameter [mm]    GSD Feed   GSD Surface   GSD Substrate
Diameter [mm]    GSD Feed   GSD Surface   GSD Substrate
.
.

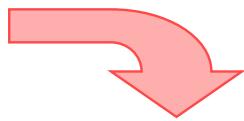
[between 2 and 9 values should be entered]

```

```
N 18  
c 1  
e 0  
S 0.0025  
L 10000  
l 1  
M 10  
p 6  
V 300  
k 2  
n 1  
r 1  
R 1.65  
l 9.4  
q 0.75  
g 0.95  
  
3 0.00005 4  
16 0.00005 1  
11 0.00005 1  
11 0.00005 2  
9 0.00005 3  
7 0.00005 3  
6 0.00005 3  
3.25 0.00005 4  
3 0.00005 5  
  
256 100 100 100  
128 80 80 80  
64 39 39 39  
32 19 19 19  
16 10 10 10  
8 5 5 5  
4 2 2 2  
2 0 0 0
```

The test file is an example of an input file that has been run during the test of the model. In the *Output* file the results of the *Test* run are reported. When the user opens the output file, the data may look somewhat convoluted at first, but once the text file is resized, they should line up in nice columns (see figure below). Some of the folders contain a fourth text file called *Output1*, which is the text file that contains the information stored by the *GetData* function.

	time (yrs):	0.000000	200.000000	400.000000	600.000000	800.000000	1000.000000	1200.000000	eta
x (m)	eta (m)	eta (m)	eta (m)	eta (m)	eta (m)	eta (m)	eta (m)	eta (m)	
0	26.200000	15.427539	15.427539	15.427539	15.427539	15.427539	15.427539	15.427539	
15.427536	15.029380	15.827894	15.827894	15.827894	15.827894	15.827894	15.827894	15.827894	
15.827893	23.530000	13.899999	13.899999	13.899999	13.899999	13.899999	13.899999	13.899999	
13.899991	13.599881	13.807618	13.807618	13.807618	13.807618	13.807618	13.807618	13.807618	
2000	20.960000	12.366722	12.366722	12.366722	12.366722	12.366722	12.366722	12.366722	
12.002114	12.000000	12.000000	12.000000	12.000000	12.000000	12.000000	12.000000	12.000000	
3000	18.349000	18.844514	18.844514	18.844514	18.844514	18.844514	18.844514	18.844514	
10.513985	10.506945	10.505648	10.505648	10.505648	10.505648	10.505648	10.505648	10.505648	
4000	15.729000	9.308234	9.308234	9.308234	9.308234	9.308234	9.308234	9.308234	
9000	9.001980	9.000036	9.000036	9.000036	9.000036	9.000036	9.000036	9.000036	
13.000000	13.109000	7.765731	7.543303	7.543303	7.543303	7.543303	7.543303	7.543303	
7.511482	7.505962	7.584936	7.594742	7.594742	7.594742	7.594742	7.594742	7.594742	
0	10.489000	6.211605	6.631803	6.631803	6.631803	6.631803	6.631803	6.631803	
5.695251	6.008947	6.008021	5.999387	5.999387	5.999387	5.999387	5.999387	5.999387	
2000	7.360000	4.668792	4.528399	4.528399	4.528399	4.528399	4.528399	4.528399	
4.598175	4.594648	4.583974	4.583974	4.583974	4.583974	4.583974	4.583974	4.583974	
5.209238	5.240000	5.188553	5.015385	5.015385	5.015385	5.015385	5.015385	5.015385	
2.999751	2.999333	2.999333	2.999333	2.999333	2.999333	2.999333	2.999333	2.999333	
9000	2.620000	1.559951	1.511258	1.511258	1.511258	1.511258	1.511258	1.511258	
1.594141	1.592862	1.592623	1.592576	1.592576	1.592576	1.592576	1.592576	1.592576	
3000	0.000000	0.000000	0.000000	0.000000	0.000000	0.000000	0.000000	0.000000	
0.000000	0.000000	0.000000	0.000000	0.000000	0.000000	0.000000	0.000000	0.000000	



	time (yrs):	0.000000	200.000000	400.000000	600.000000	800.000000	1000.000000	1200.000000	eta
x (m)	eta (m)	eta (m)	eta (m)	eta (m)	eta (m)	eta (m)	eta (m)	eta (m)	
0	25.300000	25.407539	25.407539	25.407539	25.407539	25.407539	25.407539	25.407539	
1000	23.500000	15.899999	13.564139	13.564139	13.564139	13.564139	13.564139	13.564139	
2000	23.500000	15.899999	13.564139	13.564139	13.564139	13.564139	13.564139	13.564139	
3000	10.349000	10.844514	10.255384	10.513965	10.506945	10.505648	10.505648	10.505648	
4000	12.500000	9.240000	9.000000	9.000000	9.000000	9.000000	9.000000	9.000000	
5000	13.100000	7.765731	7.543303	7.511462	7.506969	7.504742	7.504742	7.504742	
7000	10.489000	6.211605	6.015385	6.005211	6.000000	6.000000	6.000000	6.000000	
8000	5.240000	3.188553	3.015385	3.002039	3.000000	3.000000	3.000000	3.000000	
9000	2.999751	1.559951	1.511258	1.504411	1.502022	1.501576	1.501576	1.501576	
12000	0.000000	0.000000	0.000000	0.000000	0.000000	0.000000	0.000000	0.000000	

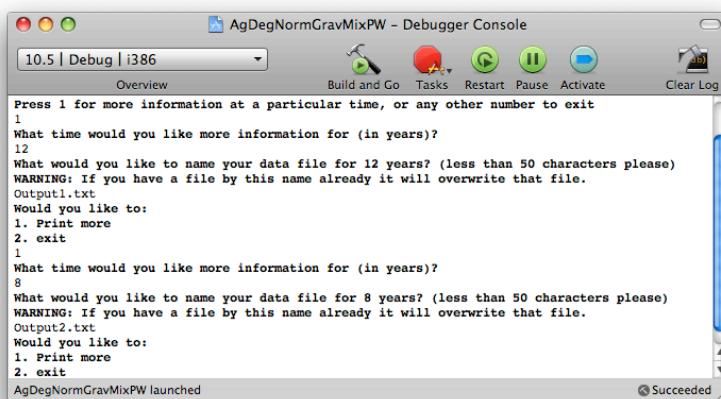
A synthetic description of the input and output parameters of the codes is reported in the continuation of this User's Guide. For a more detailed description of the problem and of the models the user should refer to Gary Parker's e-book 1D Sediment Transport Morphodynamics with applications to rivers and turbidity currents (simply called e-book in the continuation of the guide) downloadable at

[http://vtchliuc.edu/people/parkerg/morphodynamics\\_e-book.htm](http://vtchliuc.edu/people/parkerg/morphodynamics_e-book.htm).

## GetData Function:

The GetData function is included in all the codes with a time loop to store additional information at each time step that might be useful for the user.

For example, if the program outputs the bed elevation, the GetData function allows the user to get additional information (i.e. water depth, sediment transport rate, shear stress, etc.) in the computational nodes. Values of the additional parameters are stored when the code sends the main output parameters. If Output.txt contains data for year 1, 2, 3, 4, and 5, then the program has not stored the data for year 2.5. In order to have information at  $t = 2.5$  years, the user has to go back to the input file, and edit their initial parameters in order for the program to store the data at that time. The GetData function is called in the Finalize function.



## Operating Systems:

The codes of the 1D Sediment Transport Morphodynamics C Function Library can run on various operating systems (i.e. Macintosh, Linux and Windows). In each folder there is a file with the extension .xcodeproj which is for use on a Macintosh computer. The active directory folder is the same folder in which all the text documents are stored. For documentation on XCode, see the following site

<http://developer.apple.com/documentation/DeveloperTools/Xcode-date.html> .

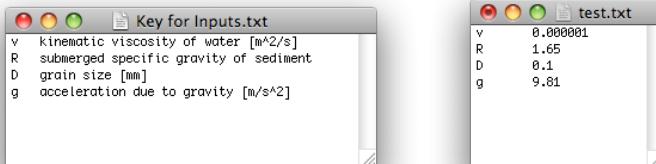
Otherwise, for use on a Linux machine, set the active directory to the folder in which all the .c documents are stored, and compile with

gcc -Wall <insert the names of all the functions here> -lm

The -lm is important, because the programs use the <math.h> library, and therefore they will not compile properly without it.

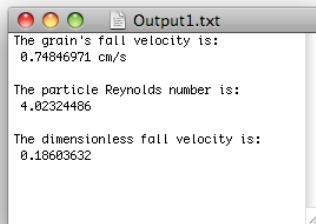
# 1) Fall Velocity

The program computes the settling velocity,  $v_s$ , of a particle with the formulation of Dietrich (1982). Input parameters are, as shown in the windows below, the sediment particle size, D, the kinematic viscosity of the liquid,  $\nu$ , the acceleration of gravity, g, and the submerged specific gravity of the sediment, R, defined as  $(\rho_s - \rho)/\rho$ , where  $\rho_s$  and  $\rho$  respectively denote the densities of the water and of the sediment.



Outputs of the model are, the particle (or grain) settling velocity in cm/s, the particle Reynolds number  $Re_p$  and the dimensionless fall velocity  $R_f$ , as shown in the window below. These last two non-dimensional parameters are respectively defined as:

$$Re_p = \frac{\sqrt{RgD}}{\nu} \quad \text{and} \quad R_f = \frac{v_s}{\sqrt{RgD}} \quad (1.i)$$



For a more detailed description of the theoretical formulation of the problem and of the code, the user should refer to Chapter 2 "Characterization of Sediment and Grain Size Distributions" of the e-book.

## Notes:

- This formulation is only valid for Reynold's numbers less than or equal to  $2.5 \cdot 10^6$ . If  $Re_p$  is greater than this upper limit, the function will alert the user, and exit the program.

List of variables

INPUT

v kinematic viscosity of water in m<sup>2</sup>/s;  
R submerged specific gravity of sediment, nondimensional  
D grain size in mm;  
g acceleration of gravity in m/s<sup>2</sup>;

OUTPUT

$v_s$  particle settling velocity in cm/s;  
 $Re_p$  particle Reynolds number;  
 $R_f$  dimensionless fall velocity.

**References:**

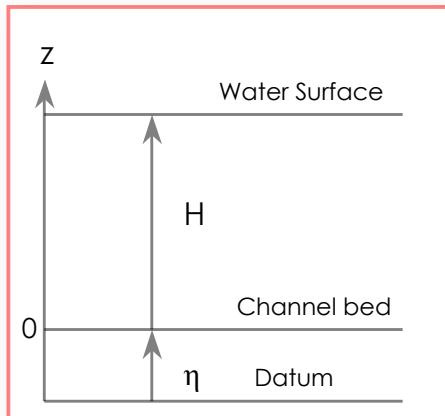
Dietrich, E. W., 1982, Settling velocity of natural particles, Water Resources Research, 18 (6), 1626-1982.

## 2) Rouse-Vanoni Equilibrium

The program computes the Rouse-Vanoni profile of suspended sediment concentration at equilibrium. For a more detailed description of the formulation the user should refer to Chapter 10 of the e-book "Relations for the Entrainment and 1D Transport of Suspended Sediment".

The parameters of the model are:

- a) H, water depth (see Figure below);
- b)  $\xi$ , vertical coordinate in the cross section, i.e.  $\xi = 0$  on the bed surface and  $\xi = H$  on the water surface (see Figure below);
- c)  $\xi_b$ , position near the bed surface where the volumetric concentration of suspended sediment is equal to  $c_b$  (see point g below);
- d)  $u^*$ , shear velocity;
- e)  $v_s$ , particle settling velocity computed with the formulation presented by Dietrich (1982) described in the previous chapter;
- f)  $c$ , concentration of suspended sediment in the water column at elevation  $\xi$  averaged over turbulence;
- g)  $c_b$ , concentration of suspended sediment in the water column at  $\xi = \xi_b$  averaged over turbulence.  $c_b$  can be computed, for example, with one of the entrainment relations presented in Chapter 10 of the e-book.

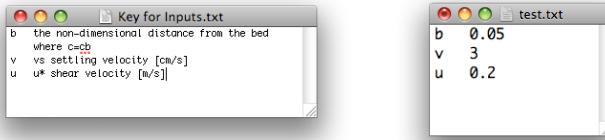


The non-dimensional expression of the Rouse-Vanoni profile is

$$\frac{c}{c_b} = \left[ \frac{(1-z)/z}{(1-b)/b} \right]^{\frac{v_s}{ku^*}} \quad (2.i)$$

where  $k$  denotes the constant of Von Karman, set equal to 0.4,  $z$  is a non-dimensional vertical coordinate defined as  $\xi/H$  and  $b$  is the non-dimensional near-bed distance equal to  $\xi_b/H$ .

Input parameters for the model are the non-dimensional distance from the bed surface where  $c_b$  is computed,  $b$ , the settling velocity of the particles,  $v_s$ , in cm/s and the shear velocity,  $u^*$ , in m/s. In the windows below the files "Key for Inputs.txt" and "test.txt" are represented.



The output of the model is the non-dimensional equilibrium profile of suspended sediment concentration, i.e. the height in the water column is expressed in terms of  $z = \xi/H$  and the volumetric concentration averaged over the turbulence as  $c/c_b$ , in the next window the output text file is represented.

The following are the data for the Rouse-Vanoni Profile:	
$z/H$	$c/c_b$
0.050000	1.000000
0.100000	0.755629
0.150000	0.635288
0.200000	0.557494
0.250000	0.500481
0.300000	0.455469
0.350000	0.418104
0.400000	0.385924
0.450000	0.357935
0.500000	0.334468
0.550000	0.305159
0.600000	0.284739
0.650000	0.262815
0.700000	0.241255
0.750000	0.219557
0.800000	0.197104
0.850000	0.172969
0.900000	0.145421
0.950000	0.109984
0.990000	0.077027
0.995000	0.045548
1.000000	0.000000

## Notes:

- To compute the equilibrium profile the user can choose between
  - the grid of the excel workbook *RTe-bookRouseSpreadsheetFun.xls*. This grid has 22 points in the vertical direction. The lowest 19 are equally spaced between  $z = b$  and  $z = b + 18(1-b)/19$ . The upper three points are located at  $z = 0.98$ ,  $z = 0.995$  and  $z = 1$ , where the concentration of suspended sediment goes to zero;
  - an equally spaced grid with a user specified number of points in the vertical.

#### List of variables

- H water depth in m;  
 $\xi$  dimensional vertical coordinate,  $\xi = 0$  on the channel bed and  $\xi = H$  at the water surface;  
z non-dimensional height in the water column equal to  $\xi/H$ ;  
c volumetric concentration of sediment averaged over turbulence;  
 $c_b$  volumetric near-bed concentration of suspended sediment averaged over turbulence, i.e. value of c at elevation  $\xi_b$ ;

#### INPUT

- b non-dimensional distance from the bed where  $c = c_b$ ,  $b = \xi_b/H$ ;  
v settling velocity of the particles in cm/s;  
 $u^*$  shear velocity in m/s;

#### OUTPUT

Non-dimensional profile of suspended sediment concentration in a two-column tab separated text file. In the first column the user finds the non-dimensional height of the point in the water column and in the second column the corresponding ratio  $c/c_b$  is reported. For  $\xi = b$ ,  $c/c_b = 1$  and for  $\xi = 1$ ,  $c/c_b = 0$ .

### 3) GSD Calculator

Given a grain size distribution, the program computes the geometric mean diameter,  $D_g$ , the geometric standard deviation,  $\sigma_g$ . Characteristic diameters based on percent finer,  $D_x$  (i.e. size such that  $x$  percent of the sample is finer than  $D_x$ ) can also be computed if requested by the user.

If the inputted size distribution does not have a lower bound,  $D_{xL}$ , such that  $xL = 0$  and an upper bound,  $D_{xU}$ , such that  $xU = 100$ , the program computes these bounds with a linear interpolation of the data. To linearly interpolate the data and to perform the calculations the code uses the  $\psi$  scale defined as

$$\psi = \log_2 D \quad (3.i)$$

where  $D$  is expressed in millimeters.

For a more detailed description of the theoretical formulation of the problem and of the code, the user should refer to Chapter 2 "Characterization of Sediment and Grain Size Distributions" of the e-book.

The input grain size distribution is specified, as indicated in the text files "Key for Inputs" and "test" represented in the windows below, in terms of  $M+1$  (with  $M < 100$ ) bound diameters in millimeters and the corresponding value of percent finer (or passing) with a scale that can be either 0 - 1 or 0 - 100.

Grain Size (mm)	Percent Passing
Grain Size (mm)	Percent Passing
:	
(no more than 100 values)	

test.txt	
4	100
2	99
1	97
0.5	83.4
0.25	42
0.125	10
0.062	3.2
0.031	2

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**Comment [1]:** How can the user specify the characteristic grain sizes he wants to compute?

**Answer:** the program prompts the user for whatever grain sizes they want to put in. The user will enter them when prompted. Do we want to put those into the inputs?

The output file, shown in the next window, is organized as follows. First the grain size distribution used in the calculations is reported, and then the values of the geometric mean diameter and standard deviation are printed. Finally the computed characteristic diameters based on percent finer,  $D_x$ , are reported, if requested by the user

```

Output.txt
The following are your organized data points:

diameter      percent passing
0.009764      0.000000
0.031000      0.020000
0.062000      0.032000
0.125000      0.100000
0.250000      0.420000
0.500000      0.834000
1.000000      0.976000
2.000000      0.999000
4.000000      1.000000

The geometric mean is: 0.272923 mm
The geometric standard deviation is: 2.168175

The size such that 20.0% is passing is: 0.155232 mm
The size such that 50.0% is passing is: 0.285831 mm
The size such that 90.0% is passing is: 0.699936 mm

```

### Notes:

- Input data may be entered either from the finer to the coarser size or from the coarser to the finer. The program will automatically reorder the data;
- Data may be on either a 0.00-1.00 scale or a 0-100 scale. The program will always use a 0-1 scale to do the calculations;
- If there are no lower and upper bounds in the inputted distribution such that their percent finer are respectively equal to 0 and 100 (or 1), the program will compute these bounds with a linear interpolation on the  $\psi$  scale and will add these diameters to the input size distribution;
- The program will prompt users if they want to calculate characteristic diameters based on percent finer,  $D_x$ , and what diameters they would like to know;
- The program can calculate up to 10 characteristic diameters based on percent finer;
- The geometric mean diameter,  $D_g$ , the geometric standard deviation,  $\sigma_g$ , and the user-defined characteristic diameters based on percent finer will be appended to a file with the reorganized, scaled, and bounded grain size distribution

## 4) Backwater Calculator

The program solves the backwater equation for subcritical flow with a predictor – corrector scheme. To compute the water depth, H, everywhere in the channel for a given water discharge per unit channel width,  $q_w$ , and downstream boundary condition, i.e. a user specified water depth, the equation is formulated as

$$\frac{dH}{dx} = \frac{S - S_f}{1 - Fr^2} \quad (4.i)$$

where  $x$  is a streamwise coordinate,  $S$  denotes the bed slope,  $S_f$  the friction slope, and  $Fr$  is the Froude number.

The bed slope,  $S$ , is assumed constant in the streamwise direction, the friction slope,  $S_f$ , and the Froude number,  $Fr$ , are defined below

$$S_f = C_f \cdot Fr^2 \quad Fr = \frac{q_w}{\sqrt{gH^3}} \quad (4.ii \text{ a, b})$$

where  $C_f$  represents a non-dimensional friction coefficient that can be evaluated with both a Chézy and a Manning-Strickler formulation, and  $g$  is the acceleration of gravity.

In the Chezy formulation (that is implemented in the excel workbook *RTe-bookBackwate.xlsr*) the friction coefficient is a user specified constant in space and time:

$$C_f = \frac{1}{Cz^2} \quad Cz = \frac{K_{Cz}}{\sqrt{g}} \quad (4.iii \text{ a, b})$$

where  $Cz$  denotes the non-dimensional Chezy friction coefficient and  $K_{Cz}$  is the dimensional Chezy friction coefficient.

In the Manning-Strickler formulation the friction coefficient is a function of the water depth. In the present model it is assumed that the cross section is wide, therefore the friction coefficient is computed as

$$C_f^{-1/2} = \alpha_r \left( \frac{H}{k_s} \right)^{1/6} \quad (4.iv)$$

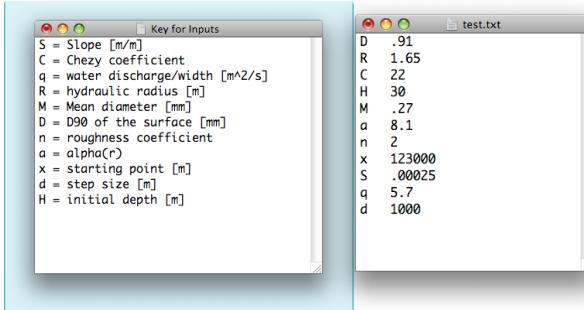
where  $\alpha_r$  is a user specified parameter and  $k_s$  is a roughness height due to skin friction defined as

$$k_s = n \cdot D_{90} \quad (4.v)$$

For a more detailed description of the theoretical formulation of and of the code, the user should refer to Chapter 5 "Review of 1D open channel hydraulics" of the e-book.

Input parameters of the model are the bed slope,  $S$ , the water discharge per unit width,  $q_w$ , a starting water depth at the downstream end of the channel,  $H_1$ , a spatial step length,  $\Delta x$ , and a starting point  $x$ . If the user wants to perform the calculation with a Chezy formulation, he should also input the non-dimensional Chezy friction coefficient,  $Cz$ , defined in equation (4.iii a, b). If a Manning-Strickler formulation is chosen, the user is asked to input the parameter  $\alpha_r$  in equation (iv) and the parameters to evaluate the roughness height,  $k_s$ , i.e. the coefficient  $n$  and the  $D_{90}$  of the bed surface defined in equation (4.v). The text file with

the definition of the input parameters and an example of input file are shown in the two windows below.



Output parameters of the model are the water depth, H, the mean flow velocity, U, the shear stress,  $\tau_b$ , the bed elevation,  $\eta$ , and the water surface elevation,  $\xi$ , equal to the sum of the bed elevation and the water depth. In the output file water depth, mean flow velocity, Froude number and shear stress are also reported for normal flow conditions with the critical values of water depth and mean flow velocity. An example of the output file is reported in the window below.

Output.txt					
The chezy formulation H is: 3.013699 The chezy formulation U is: 1.8941379 The chezy formulation Fr is: 0.347851 The chezy formulation tau(b) is: 7.391072					
The critical value for H is: 1.498597 The critical value for U is: 3.823971					
For the Chezy formulation:					
x	U	tau(b)	et.b	k.s	
123000	30.000000	0.190000	0.074587	0.200000	38.000000
122000	29.760225	0.191955	0.075944	0.250000	38.002025
121000	29.500456	0.193217	0.077134	0.500000	38.009456
120000	29.230687	0.194527	0.078367	0.750000	38.016892
119000	29.000937	0.195455	0.079584	1.000000	38.00937
118000	28.751196	0.196253	0.081287	1.250000	38.001186
117000	28.501442	0.199999	0.082636	1.500000	38.001142
116000	28.251705	0.202659	0.084194	1.750000	38.001795
115000	28.002000	0.203577	0.085748	2.000000	38.001776
114000	27.752254	0.203599	0.087158	2.250000	38.002054
113000	27.502541	0.207254	0.088748	2.500000	38.002541
112000	27.252834	0.209153	0.090382	2.750000	38.002534
111000	27.003127	0.210452	0.092051	3.000000	38.002536
110000	26.753445	0.213957	0.093787	3.250000	38.003445
109000	26.503763	0.215057	0.095563	3.500000	38.003763
108000	26.254092	0.217189	0.097339	3.750000	38.004091
107000	26.004429	0.219320	0.099106	4.000000	38.004429
106000	25.754767	0.221418	0.100872	4.250000	38.004776
105000	25.505133	0.223494	0.102193	4.500000	38.005133
104000	25.255593	0.225693	0.103243	4.750000	38.005581
103000	25.005950	0.227796	0.104254	5.000000	38.005950
102000	24.756312	0.229845	0.105250	5.250000	38.006312
101000	24.506674	0.232590	0.111773	5.500000	38.006674
100000	24.257098	0.234983	0.114885	5.750000	38.007098
99000	24.007517	0.237426	0.116469	6.000000	38.007517
98000	23.757936	0.239700	0.118092	6.250000	38.007936
97000	23.508415	0.242466	0.121467	6.500000	38.008415
96000	23.258813	0.245668	0.124087	6.750000	38.008813
95000	23.009373	0.247725	0.126793	7.000000	38.009373
94000	22.759830	0.250000	0.129474	7.250000	38.00974
93000	22.501093	0.252326	0.132476	7.500000	38.010393
92000	22.252492	0.255694	0.135462	7.750000	38.010929
91000	22.003894	0.258956	0.138559	8.000000	38.011484
90000	21.762058	0.262192	0.141744	8.250000	38.012058
89000	21.512651	0.265458	0.145050	8.500000	38.012651
88000	21.263244	0.268642	0.148354	8.750000	38.013244

## Notes:

- The program will automatically recognize the formulation (i.e. Chezy or Manning-Strickler) the user wants to implement from the input file. If a value of the Chezy friction coefficient is specified in the input file but the parameters to compute the friction coefficient with equation (4.v) are not given, the code will use a Chézy formulation. On the contrary, if values for  $\alpha_r$ ,  $D_{90}$  and  $n$  are specified while there is no Chezy friction

Enrica 11/8/2009 11:06 PM

**Comment [2]:** The order of the input parameters is different. Is this a problem?

**Answer:** this does not make a difference, I'm not sure how you would like to proceed here. I can rearrange them, or we can make a comment in the beginning of the user guide that the order doesn't matter (as long as the right letter is in front of it)

coefficient in the input file, the program will automatically use a Manning-Strickler formulation. If all the four parameters are specified, the program will ask the user which formulation he would prefer to use

#### List of variables

x streamwise coordinate, in m

#### INPUT

S bed slope

$q_w$  water discharge per unit width in  $m^2/s$

$C_z$  non-dimensional Chézy friction coefficient

$D_{90}$  diameter of the bed surface such that 90% of the distribution is finer in mm

n parameter to compute roughness height with equation (v)

$\alpha_r$  dimensionless constant in equation (iv)

$x_1$  starting position in m

$\Delta x$  step size in m

$H_1$  initial depth in m

#### OUTPUT

H water depth in m

U mean flow velocity in m/s

$\tau(b)$  shear stress on the bed surface in  $N/m^2$

$\eta$  bed surface elevation in m

$c_s$  water surface elevation in m

$H_n$  water surface at normal flow in m

$H_c$  critical water depth in m

$F_r$  Froude number at normal flow

$U_n$  mean flow velocity at normal flow in m/s

$U_c$  critical flow velocity

$\tau(b)_n$  bed shear stress at normal flow in  $N/m^2$

## 5) AgDegNormal

The program calculates a) an ambient mobile-equilibrium sediment transport rate and b) the morphodynamic evolution of a reach due to a change in sediment input rate. For a detailed description of the theory and of the model, the user should refer to Chapter 14 "1D aggradation and degradation of rivers: normal flow assumption" of the e-book, and to the word file *RTE-bookAgDegNormalFormul.doc*, downloadable at  
[http://vtchl.uiuc.edu/people/parkerg/word\\_files.htm](http://vtchl.uiuc.edu/people/parkerg/word_files.htm).

The assumptions and the simplifications at the basis of the formulation are:

- 1) the quasi steady approximation (deVries, 1965), i.e. the bed changes so slowly compared to the characteristic response time of the flow that the flow can be approximated as responding immediately. This assumption greatly simplifies the problem, but it limits the applicability of the model. In particular, this model cannot be applied in cases when a) it is desired to characterize the sediment transport over an entire rapidly varying hydrograph; b) one wishes to capture the effect of a flood wave (with a high water surface slope on the upstream side of the wave and a low water surface slope on the downstream side) on sediment transport; and c) the flow makes transitions between subcritical and supercritical flow, in which case a shock-capturing method capable of automatically locating hydraulic jumps is required. For more information on the quasi-steady approximation, the user can refer to Chapter 13 "The quasi-steady approximation" of the e-book;
- 2) the channel width,  $B_c$ , is constant in the streamwise direction and in time;
- 3) the floodplain is absent;
- 4) the flow is assumed normal;
- 5) the sediment is uniform;
- 6) the Exner equation of channel bed sediment conservation is based on the computation of total bed material load;
- 7) the relevant morphodynamic changes of the river occur during floods. The full hydrograph or flow duration curve is replaced by flood intermittency factor,  $I_f$ , (Paola et al., 1992) and a constant flood discharge,  $Q_w$ , (or discharge per unit channel width  $q_w = Q_w/B_c$ ).

The water depth can be computed either with a Manning-Strickler (implemented in the excel file *RTE-bookAgDegNormal.xls*) or with a Chezy formulation. In the Chezy formulation, the non-dimensional friction coefficient  $C_z$ , eq. (4.iii a, b), is a user specified parameter. In the Manning-Strickler formulation the friction coefficient,  $C_f$ , is still computed as

$$C_f = \alpha_r \left( \frac{H}{k_c} \right)^{\frac{1}{n}} \quad (5.i)$$

This expression has the same form of eq. (4.iv).  $\alpha_r$  is a user specified parameter and the roughness height,  $k_s$ , due to skin friction is substituted with a composite roughness,  $k_c$ , which may include the effect of bedforms, if present. For further information on the normal flow calculation, the user can refer to Chapter 5 "Review of 1D open channel hydraulics" of the e-book". Some techniques to estimate  $k_c$  when bedforms are expected are explained in

Chapters 9 "Relations for Hydraulic Resistance in Rivers", 10 "Relations for the Entrainment and 1D Transport of Suspended Sediment" and 14 "1D Aggradation and Degradation of Rivers: Normal Flow Assumption" of the e-book.

The total bed material load per unit width,  $q_t$ , is computed with the generic equation

$$\frac{q_t}{\sqrt{RgDD}} = \alpha_r \left( \frac{\varphi_s \tau_b}{\rho R g D} - \tau_c^* \right)^{n_t} \quad (5.ii)$$

where

- $q_t$  is computed in  $m^2/s$ ;
- $R$  denotes the submerged specific gravity of the sediment and is defined as  $(\rho_s - \rho)/\rho$ , where  $\rho$  and  $\rho_s$  are the densities of the water and of the sediment;
- $g$  is the acceleration of gravity;
- $D$  is the diameter of the sediment;
- $\varphi_s \leq 1$  is a constant to convert total boundary shear stress to that due to skin friction (if necessary);
- $\tau_b$  is the total boundary shear stress defined as  $\rho g H S$  (with  $H$  to denote the water depth and  $S$  the bed slope) for the normal flow assumption (refer to Chapter 5 of the e-book);
- $\alpha_r, \tau_c^*$  and  $n_t$  are user specified parameters that depend on the load relation. For the version of the Meyer-Peter and Muller bedload relation due to Wong and Parker (2006)  $\alpha_r = 3.97$ ,  $\tau_c^* = 0.0495$  and  $n_t = 1.5$ . For the relation of Engelund and Hansen (1967) to compute the total load of sand,  $\alpha_r = 0.05/C_f$ ,  $\tau_c^* = 0$  and  $n_t = 2.5$ . For more details on the relations to compute bedload and total sediment transport rates, the user can refer to Chapters 7 "Relations for 1D Bedload Transport" and 12 "Bulk Relations for Transport of Total Bed Material Load" of the e-book.

The Exner equation, eq. (5.iii) below, is solved with a finite difference scheme. For the derivation of the Exner equation the user can refer to Chapter 4 "Relations for the Conservation of Bed Sediment" of the e-book.

$$(1 - \lambda_p) \frac{\partial \eta}{\partial t} = -I_f \frac{\partial q_t}{\partial x} \quad (5.iii)$$

where  $\lambda_p$  denotes the bed porosity,  $\eta$  is the bed elevation above the datum,  $x$  is the streamwise coordinate and  $t$  is the temporal coordinate. The length of the fluvial reach is denoted as  $L$ . The domain is divided in  $M$  sub-reaches bounded by  $(M+1)$  computational nodes. The spatial step length  $\Delta x$  is thus equal to  $L/M$ . The annual sediment transport rate is fed in a ghost node one step length upstream of the first node, i.e.  $q_{t,ghost} = q_{t,feed}$ .

The Exner equation is discretized as

$$\eta|_{i,t+\Delta t} = \eta|_{i,t} - \frac{I_f}{1 - \lambda_p} \frac{\Delta q_{t,i}}{\Delta x} \Delta t \quad (5.iv)$$

where the index  $i$  denotes the computational node, the index  $t$  denotes the time and  $\Delta t$  is the time step. The spatial derivative of the total bed material load per unit width,  $\Delta q_{t,i}/\Delta x$  is computed as

$$\frac{\Delta q_{t,i}}{\Delta x} = a_u \frac{q_{t,i} - q_{t,i-1}}{\Delta x} + (1 - a_u) \frac{q_{t,i+1} - q_{t,i}}{\Delta x} \quad (5.v)$$

where  $a_u$  is an upwinding coefficient. In a pure upwinding scheme  $a_u = 1$ , in a central difference scheme  $a_u = 0.5$ . In computing eq. (5.vi) at the first node it is assumed that  $q_{t,-1} = q_{t,ghost} = q_{t,feed}$ .

The bed slope is computed in each node with the following relation

$$S = \begin{cases} \frac{\eta_1 - \eta_2}{\Delta x} & i = 1 \\ \frac{\eta_{i-1} - \eta_{i+1}}{2\Delta x} & i = 2 \dots M \\ \frac{\eta_M - \eta_{M+1}}{\Delta x} & i = M + 1 \end{cases} \quad (5.vi)$$

The upstream boundary conditions are specified in terms of a water discharge,  $Q_w$ , flood intermittency,  $I_f$ , and annual sediment transport rate in metric tons per year,  $G_{tf}$ . The downstream boundary condition is expressed as a fix bed elevation at the downstream end of the fluvial reach. Thus the Exner equation is not implemented at node  $M+1$ . The bed elevation at the downstream end of the modeled reach is set equal to zero.

The initial condition is that of a given bed profile. In the model the initial profile,  $\eta(x, t)_{t=0}$  is given in terms of a specified initial downstream bed elevation  $\eta_{ld}$  and constant initial slope  $S_i$

$$\eta(x, t)_{t=0} = \eta_{ld} + S_i(L - x) \quad (5.vii)$$

Input parameters of the model are

- the characteristic flood discharge,  $Q_w$ , and the flood intermittency,  $I_f$ ;
- the channel width,  $B_c$ ;
- the characteristic diameter of the sediment,  $D$ ;
- the bed porosity,  $\lambda_p$ ;
- the roughness height,  $k_c$ ;
- the ambient bed slope,  $S$ ;
- the imposed annual sediment transport rate,  $G_{tf}$ ;
- the length of the fluvial reach,  $L$ ;

The user also has to specify the following parameters to perform the calculations and to control the output:

- the number of sub-reaches,  $M$ ;
- the temporal step length,  $\Delta t$ ;
- the upwinding coefficient  $a_u$ ;
- the number of time steps to printout,  $N_{toprint}$ ;
- the numbers of printout in addition to the initial equilibrium state,  $N_{print}$ ;
- the coefficient  $\alpha_r$  in eq. (5.i);
- the parameters to define the load relation,  $\alpha_r, \tau_c^*$  and  $n_t$ , defined in eq. (5.ii);

- the constant to convert the total boundary shear stress to that due to skin friction,  $\varphi_s$ ;
  - the submerged specific gravity of the sediment,  $R$ .

The text file with the definition of the input parameters and an example of input file are shown in the two windows below.



The outputs of the model are  $N_{\text{print}}$  longitudinal profiles at different times  $t_{\text{plot},j}$

$$t_{\text{plot},j} = j \cdot N_{\text{toprint}} \Delta t, \quad j = 1 \dots N_{\text{print}} \quad (5.\text{viii})$$

An example of the output file *Output.txt* is reported in the window below with an example of the additional parameters that the user may want to plot in the text file *Output1.txt*, i.e. bed slope,  $S$ , water depth,  $H$ , Shield's number,  $\tau_b$ , and the total bed material load,  $q_t$ . The initial bed profile is the bed profile at mobile-bed equilibrium, given the water discharge, the flow intermittency and the initial bed slope.

## Notes:

- The maximum number of computational nodes, M, is 99 (this is the case for all of the AgDeg functions)
  - Calculates the water depth with a Chézy formulation, if only the Chézy coefficient is specified in the input text file. The code uses a Manning-Strickler formulation, when only the roughness height,  $k_c$ , and the coefficient  $\alpha_r$  are given in the input text file. If all these parameters are in the text file, the program will ask the user which formulation he would like to use

- Prompts user whether they would like to append certain parameters (i.e. flow depth at flood, Shield's number at flood, Einstein number at flood, etc.) to the data file, or write them in a separate file

#### List of variables

$x$	streamwise coordinate in m
$\Delta x$	spatial step length in m
$t$	temporal coordinate in seconds
$C_f$	non-dimensional friction coefficient
INPUT	
$Q_w$	flood discharge in $m^3/s$
$I_f$	flood intermittency
$B_c$	channel width in m
$D$	characteristic grain size in mm
$\lambda_p$	bed porosity
$k_c$	composite roughness height in mm
$S_l$	ambient bed slope
$G_{tf}$	imposed annual sediment transport rate in tons/annum
$L$	length of reach in m
$\Delta t$	time step in yr
$N_{ttoprint}$	number of timesteps to printout
$N_{print}$	number of printouts
$M$	number of spatial intervals
$\alpha_u$	upwinding coefficient (1=full upwind, 0.5=central difference)
$\alpha_r$	coefficient in Manning-Strickler
$\alpha_s$	coefficient in sediment transport relation
$n_t$	exponent in sediment transport relation
$\tau_c^*$	reference Shields number in sediment transport relation
$\varphi_s$	fraction of bed shear stress due to skin friction
$R$	submerged specific gravity
$C_z$	non-dimensional Chézy friction coefficient
OUTPUT	
$\eta$	bed surface elevation in m
$S$	bed slope
$H$	water depth in m
$\tau_b$	total boundary shear stress on bed surface N/m <sup>2</sup>
$q_t$	total bed material load in m <sup>2</sup> /s

## References:

- de Vries, M. 1965. Considerations about non-steady bed-load transport in open channels. *Proceedings, 11th Congress, International Association for Hydraulic Research*, Leningrad: 381-388.
- Engelund, F. and E. Hansen, 1967, *A Monograph on Sediment Transport in Alluvial Streams*, Technisk Vorlag, Copenhagen, Denmark.
- Paola, C., Heller, P. L. & Angevine, C. L., 1992, The large-scale dynamics of grain-size variation in alluvial basins. I: Theory, *Basin Research*, 4, 73-90.
- Wong M. and G. Parker, 2006, Reanalysis and correction of bed-load relation of Meyer-Peter and Müller using their own database, *Journal of Hydraulic Engineering*, 132(11), 1159-1168;

## 6) AgDegBW

The model calculates a) an ambient mobile-bed equilibrium, and b) the response of a river reach to either 1) changed sediment input rate at the upstream end of the reach starting from  $t = 0$  or 2) changed downstream water surface elevation at the downstream end of the reach starting from  $t = 0$ , where  $t$  is the temporal coordinate. The code is very similar to AgDegNorm described in Chapter 5 of this user's guide. The main difference between the two codes is in the procedure to compute the water depth. In AgDegNorm the flow is assumed normal (i.e. steady and uniform), while in AgDegBW the flow is assumed steady and it is computed solving the backwater equation. The case of Froude-subcritical flow, for which  $Fr < 1$  (refer to Chapter 5 "Review of 1D open channel hydraulics" of the e-book), is considered herein. This implies that integration of the backwater equation must proceed upstream from  $x = L$ , with  $x$  streamwise coordinate and  $L$  length of the modeled reach. Both a Chezy and a Manning-Strickler formulation can be used to compute the flow.

For a detailed description of the theory and of the model, the user should refer to Chapter 20 "Aggradation and Degradation of Rivers: Backwater Formulation" of the e-book, and to the word file *RTe-bookAgDegBWFormul.doc*, downloadable at

[http://vtchli.uiuc.edu/people/parkerg/word\\_files.htm](http://vtchli.uiuc.edu/people/parkerg/word_files.htm).

The upstream boundary conditions are specified in terms of a sediment feed rate in metric tons per year,  $G_{tf}$ , a characteristic flood discharge,  $Q_w$ , and a flow intermittency,  $I_f$ . The downstream boundary condition is given in terms of a known water surface elevation,  $\xi_d$ , with

$$\xi_d = [\eta(x, t) + H(x, t)]_{x=L} \quad (6.i)$$

where  $\eta$  denotes the bed elevation and  $H$  the water depth. As opposed to the normal flow calculation, downstream bed elevation  $\eta(L, t)$  is no longer specified, and is free to vary during the run, i.e. the Exner equation is implemented at node M+1.

The initial bed profile is the same as the one used for the calculations using the normal flow approximation, eq. (5.vii) i.e. at  $t = 0$  the initial bed elevation at the downstream end of the reach is set equal to zero.

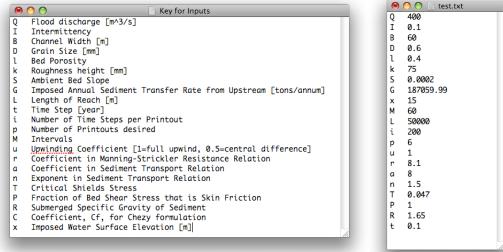
Input parameters of the model are

- the characteristic flood discharge,  $Q_w$ , and the flood intermittency,  $I_f$ ;
- the channel width,  $B_c$ ;
- the characteristic diameter of the sediment,  $D$ ;
- the bed porosity,  $\lambda_p$ ;
- the roughness height,  $k_c$ ;
- the ambient bed slope,  $S$ ;
- the imposed annual sediment transport rate,  $G_{tf}$ ;
- the imposed water surface elevation at the downstream end of the reach,  $\xi_d$ ;
- the length of the fluvial reach,  $L$ ;

The user also has to specify the following parameters to perform the calculations and to control the output:

- the number of sub-reaches,  $M$ ;
  - the temporal step length,  $\Delta t$ ;
  - the upwinding coefficient  $a_u$ ;
  - the number of time steps to printout,  $N_{\text{tprint}}$ ;
  - the numbers of printout in addition to the initial equilibrium state,  $N_{\text{print}}$ ;
  - the coefficient  $\alpha_r$  in eq. (5.i);
  - the parameters to define the load relation,  $\alpha_s, \tau_c^*$  and  $n_t$ , defined in eq. (5.ii);
  - the constant to convert the total boundary shear stress to that due to skin friction,  $\varphi_s$ ;
  - the submerged specific gravity of the sediment,  $R$ .

The text file with the definition of the input parameters and an example of input file are shown in the two windows below.



The output parameters of the model printed in the text file *Output.txt* are the bed elevation,  $\eta$ , the channel slope,  $S$ , the water depth,  $H$ , and the water surface elevation,  $\xi$ , at the different times  $t_{plot,j}$  defined in eq. (5.viii). An example of output file is reported in the window below. The values of bed elevation, water surface and water elevation at  $t = 0$  are representative of the condition of mobile-bed equilibrium antecedent to the change in annual sediment input rate or downstream water elevation.

The additional parameters that the user may print at times  $t_{plot,i}$  in the text file *Output1.txt* are the channel slope,  $S$ , the Shields number,  $\tau_b$ , and the total sediment transport rate,  $q_t$ . An example of the file *Output1.txt* is reported in the windows below. The parameters at  $t = 0$  are

representative of the condition of mobile-bed equilibrium antecedent to the change in annual sediment input rate or downstream water elevation.

x(n)	t(n)	H(n)	k1(n)	t0	z0
0.9	120yr	120yr	4.469461	20.46292	0.894398
0.933..3	15.973553	0.000211	4.481061	20.454649	0.888535
1666..7	15.798637	0.000209	4.493828	20.29465	0.882648
2590..8	15.625008	0.000207	4.506793	20.131877	0.876752
3524..3	15.452379	0.000206	4.519753	19.97703	0.870856
4166..7	15.282080	0.000204	4.533206	19.815366	0.864917
5000..0	15.112632	0.000203	4.546817	19.659448	0.858922
5833..3	14.944553	0.000201	4.560538	19.505991	0.852886
6666..7	14.777042	0.000199	4.574454	19.352295	0.846759
7500..0	14.610531	0.000198	4.588376	19.200455	0.840639
8333..3	14.4440518	0.000196	4.602395	19.051395	0.834492
9166..7	14.285992	0.000194	4.617399	18.903291	0.828494
10000..0	14.124647	0.000193	4.632187	18.756754	0.822388
10833..3	13.964758	0.000191	4.647831	18.611781	0.816268
11666..7	13.805869	0.000190	4.663475	18.467803	0.810153
12500..0	13.646982	0.000189	4.677511	18.322533	0.803942
13333..3	13.493185	0.000188	4.693871	18.186256	0.797633
14166..7	13.336861	0.000185	4.708848	18.047543	0.791486
15000..0	13.185549	0.000183	4.724844	17.918392	0.785166
15833..3	13.035239	0.000182	4.741063	17.791914	0.778918
16666..7	12.885928	0.000181	4.757277	17.664771	0.772658
17500..0	12.734117	0.000178	4.774198	17.538297	0.766322
18333..3	12.586293	0.000177	4.791085	17.377379	0.760084
19166..7	12.439787	0.000175	4.808225	17.249813	0.753743
20000..0	12.293433	0.000173	4.825364	17.122019	0.747423
20833..3	12.147084	0.000172	4.842525	16.994269	0.741095
21666..7	12.000814	0.000170	4.861891	16.860294	0.734741
22500..0	11.866814	0.000169	4.879297	16.746821	0.728387
23333..3	11.726799	0.000167	4.897576	16.624375	0.722024
24166..7	11.586569	0.000166	4.915203	16.501262	0.715653
25000..0	11.446344	0.000164	4.932830	16.378149	0.709281
25833..3	11.314373	0.000163	4.950448	16.266822	0.703893
26666..7	11.179418	0.000161	4.973676	16.153884	0.696498
27500..0	11.045687	0.000160	4.993383	16.039867	0.690087
28333..3	10.912199	0.000158	5.015336	15.925651	0.683679
29166..7	10.778715	0.000157	5.038444	15.812444	0.677267
30000..0	10.645185	0.000155	5.061414	15.700869	0.670845
30833..3	10.512988	0.000154	5.078285	15.598973	0.664419
31666..7	10.380506	0.000153	5.096264	15.491570	0.657998
32500..0	10.248292	0.000151	5.117250	15.386556	0.651551

## Notes:

- The downstream water surface elevation must exceed the  $H_c$ , critical depth, which is equal to  $(Q_w^2/(B_c^2g))^{1/3}$ , otherwise the user is alerted, and the program exits;
- In the original version of the code embedded in the excel file RTe-bookAgDegBW.xls the Manning-Strickler formulation is implemented;
- The water depth is calculated using a Chézy formulation, when only the Chézy coefficient is specified in the input text file. The Manning-Strickler formulation is implemented, when only the roughness height,  $k_c$ , and the coefficient  $\alpha$ , in eq. (5.i) are given in the input text file. When all the three parameters are present, the program will ask the user which formulation they would like to use
- Program prompts user whether they would like to append certain parameters to their output file (i.e. Output.txt), or whether they would like to create a separate file for these parameters. Included in these parameters are values such as flow depth at flood, Einstein number at flood, initial normal Froude number, ultimate normal Froude number, etc.

### List of variables

$x$	streamwise coordinate in m
$\Delta x$	spatial step length in m
$t$	temporal coordinate in seconds
$C_f$	non-dimensional friction coefficient

#### INPUT

$Q_w$	flood discharge in $m^3/s$
$I_f$	flood intermittency
$B_c$	channel width in m
$D$	characteristic grain size in mm
$\lambda_p$	bed porosity
$k_c$	composite roughness height in mm
$S_l$	ambient bed slope
$\xi_d$	imposed downstream water elevation in m
$G_f$	imposed annual sediment transport rate in tons/annum
$L$	length of reach in m
$\Delta t$	time step in yr
$N_{toprint}$	number of time steps to printout
$N_{print}$	number of printouts
$M$	number of spatial intervals
$a_u$	upwinding coefficient (1=full upwind, 0.5=central difference)
$\alpha_r$	coefficient in Manning-Strickler
$\alpha_s$	coefficient in sediment transport relation
$n_t$	exponent in sediment transport relation
$\tau_c^*$	reference Shields number in sediment transport relation
$\varphi_s$	fraction of bed shear stress due to skin friction
$R$	submerged specific gravity
$C_z$	non-dimensional Chézy friction coefficient

#### OUTPUT

$\eta$	bed surface elevation in m
$S$	bed slope
$H$	water depth in m
$\xi$	water surface elevation in m
$\tau_b$	total boundary shear stress on bed surface N/m <sup>2</sup>
$q_t$	total bed material load in $m^2/s$

## 7) AgDegNormalSub

The program computes the approach to mobile-bed equilibrium in a river carrying uniform material and flowing into a subsiding basin. It is a descendant of AgDegNormal. Three relatively minor changes have been implemented as follows:

- a) The input parameters have been modified to include the following parameters: subsidence rate  $\sigma$ , ratio of depositional width to channel width  $r_B$ , ratio of wash load deposited per unit bed material load  $\Lambda$  and channel sinuosity  $\Omega$ ;
- b) The code has been modified so as to include subsidence in the calculation of mass balance;
- c) The output has been modified to show the time evolution of not only the profile of bed elevation  $\eta$ , but also the profiles of bed slope  $S$  and the ratio  $q_t/q_{tf}$ , where  $q_{tf}$  denotes the volume feed rate of bed material load per unit width.

The new input parameters, i.e. ratio of depositional to channel width,  $r_B$ , units of wash load deposited per unit bed material load,  $\Lambda$ , and channel sinuosity,  $\omega$ , have been introduced to generalize the Exner equation for uniform sediment from a flume-like setting, eq. (6.i) below, to a river, eq. (6.ii).

$$(1 - \lambda_p) \frac{\partial \eta}{\partial t} = - \frac{\partial q_t}{\partial x} \quad (7.i)$$

$$(1 - \lambda_p) \left( \frac{\partial \eta}{\partial t} + \sigma \right) = - \frac{l_f}{r_B} (1 + \Lambda) \Omega \frac{\partial q_t}{\partial x} \quad (7.ii)$$

where  $\lambda_p$  denotes the bed porosity,  $\eta$  is the bed elevation,  $q_t$  is the total sediment transport rate per unit channel width,  $l_f$  is the flow intermittency,  $t$  is the temporal coordinate and  $x$  is the streamwise (i.e. downchannel) coordinate.

The ratio of depositional to channel width,  $r_B$ , has been introduced to model the fact that in an aggrading river sediment deposits not only in the channel itself, but also in a much wider belt (e.g. the floodplain or basin width, due to overbank deposition, channel migration and avulsion). Here channel width is denoted as  $B_c$  (which can be taken to be synonymous with bankfull width) and effective depositional width is denoted as  $B_d$ . Both of these are taken as constant here for simplicity.  $r_B$  is thus defined as

$$r_B = \frac{B_d}{B_c} \quad (7.iii)$$

The parameter  $\Lambda$  that represents the units of wash load deposited per unit of bed material is introduced to consider that in the 1D formulation implemented in this model it is assumed that deposition occurs not only in the channel but on a much wider area (e.g. the floodplain). Sediment deposited in the channel is mostly made of bed material but sediment deposited around the channel contains a significant amount of wash load. A precise mass balance for wash load is beyond the scope of this model. For simplicity it is assumed that for every unit of sand deposited in the system,  $\Lambda$  units of wash load are deposited. It is also assumed that the supply of wash load from upstream is always sufficient for deposition at

such a rate. This is not likely to be strictly true, but should serve as a useful starting assumption.

The parameter  $\Omega$  has been introduced to consider that channels may be sinuous. Here it is assumed that the channel has a sinuosity,  $\Omega$ , but that the depositional surface across which it wanders is rectangular. In the present formulation the sinuosity is defined as the ratio of downchannel distance per unit of downvalley distance.

For a detailed description of the theory and of the model, the user should refer to Chapters 4 "Relations for the Conservation of Bed Sediment", 25 "Long Profiles of Rivers, with an Application on the Effect of Base Level Rise on Long Profiles" and 26 "Rivers Flowing into Subsiding Basins: Upward Concavity of Long Profile and Downstream Fining" of the e-book.

Boundary and initial conditions are equal to that implemented for the ancestor model AgDegNormal.

The input parameters of the model are specified in a slightly different form than those of AgDegNormal. The input parameters of AgDegNormalSub are listed below:

- the characteristic flood discharge per unit channel width,  $q_w$ , and the flood intermittency,  $I_f$ ;
- the characteristic diameter of the sediment,  $D$ ;
- the bed porosity,  $\lambda_p$ ;
- the roughness height,  $k_c$ ;
- the ambient bed slope,  $S$ ;
- the volume sediment feed rate per unit channel width,  $q_{tf}$ ;
- the subsidence rate,  $\sigma$ ;
- the ratio of depositional width to channel width,  $r_B$ ;
- the channel sinuosity,  $\Omega$ ;
- the units of wash load deposited in the system per unit of bed material load,  $\Lambda$ ;
- the length of the fluvial reach,  $L$ ;

The user also has to specify the following parameters to perform the calculations and to control the output:

- the number of sub-reaches,  $M$ ;
- the temporal step length,  $\Delta t$ ;
- the upwinding coefficient  $a_u$ ;
- the number of time steps to printout,  $N_{toprint}$ ;
- the numbers of printout in addition to the initial equilibrium state,  $N_{print}$ ;
- the coefficient  $\alpha_r$  in eq. (5.i);
- the parameters to define the load relation,  $\alpha_s, \tau_c^*$  and  $n_t$ , defined in eq. (5.ii);
- the constant to convert the total boundary shear stress to that due to skin friction,  $\varphi_s$ ;
- the submerged specific gravity of the sediment,  $R$ .

The text file with the definition of the input parameters and an example of input file are shown in the two windows below.

Key for Inputs.txt

```

q Water Discharge per unit Width During Flood [m^2/s]
I Intensity of Rainfall [mm/h]
D Grain Size [mm]
B Bed Porosity
k Roughness height [mm]
S Ambient Channel Slope
W Wash Sediment Feed Rate Per Unit Width [m^2/s]
L Length of Reach [m]
T Time Step [year]
N Number of Steps per Printout
P Number of Printsouts desired
M Intervals
U Upwinding Coefficients [1=Mull upwind, 0=central difference]
C Coefficient in Manning-Strickler Resistance Relation
G Coefficient in Sediment Transport Relation
E Exponent in Sediment Transport Relation
Tc Critical Sediment Transport Capacity
F Fraction of Bed Shear Stress that is a Skin Friction
R Submerged Specific Gravity of Sediment
C Coefficient Cf, for Chezy formulation
S Subcritical reach length
B Ratio of Depositional Width to Channel Width
O Channel Sinuosity
Y Unit Wash Load Deposited in Channel-Floodplain per Unit Bed Material Load

```

test.txt

```

Q 3
I 0.025
D 20
L 0.4
K 60
S 0.005
Q 0.004
L 52000
T 1
N 1000
P 5
M 50
U 0.5
C 8.1
G 4
E 1.5
Tc 0.0495
F 1.65
S 5
B 60
O 1.5
Y 1

```

The output parameters of the model printed in the text file *Output.txt* are the bed elevation,  $\eta$ , the channel slope, S, and the ratio between the transport and the feed rate of bed material load.

Output.txt										
t(s)	S1	qb/qf1	eta(t)	eta(t)	S1	qb/qf1	eta(t)	eta(t)	S1	qb/qf1
0.00	20.00000	0.00000	0.00000	0.00000	20.00000	0.00000	0.00000	0.00000	20.00000	0.00000
0.001	20.40000	0.00000	0.393775	0.393775	20.40000	0.00000	0.395663	0.395663	20.40000	0.00000
0.002	20.80000	0.00000	0.787550	0.787550	20.80000	0.00000	0.794566	0.794566	20.80000	0.00000
0.003	21.20000	0.00000	0.990000	0.990000	21.20000	0.00000	0.996323	0.996323	21.20000	0.00000
0.004	21.60000	0.00000	0.993775	0.993775	21.60000	0.00000	0.998338	0.998338	21.60000	0.00000
0.005	22.00000	0.00000	0.996323	0.996323	22.00000	0.00000	0.999471	0.999471	22.00000	0.00000
0.006	22.40000	0.00000	0.998338	0.998338	22.40000	0.00000	0.999686	0.999686	22.40000	0.00000
0.007	22.80000	0.00000	0.999471	0.999471	22.80000	0.00000	0.999800	0.999800	22.80000	0.00000
0.008	23.20000	0.00000	0.999686	0.999686	23.20000	0.00000	0.999904	0.999904	23.20000	0.00000
0.009	23.60000	0.00000	0.999800	0.999800	23.60000	0.00000	0.999954	0.999954	23.60000	0.00000
0.010	24.00000	0.00000	0.999904	0.999904	24.00000	0.00000	0.999975	0.999975	24.00000	0.00000
0.011	24.40000	0.00000	0.999954	0.999954	24.40000	0.00000	0.999986	0.999986	24.40000	0.00000
0.012	24.80000	0.00000	0.999975	0.999975	24.80000	0.00000	0.999994	0.999994	24.80000	0.00000
0.013	25.20000	0.00000	0.999986	0.999986	25.20000	0.00000	0.999997	0.999997	25.20000	0.00000
0.014	25.60000	0.00000	0.999994	0.999994	25.60000	0.00000	0.999998	0.999998	25.60000	0.00000
0.015	26.00000	0.00000	0.999997	0.999997	26.00000	0.00000	0.999999	0.999999	26.00000	0.00000
0.016	26.40000	0.00000	0.999998	0.999998	26.40000	0.00000	0.999999	0.999999	26.40000	0.00000
0.017	26.80000	0.00000	0.999999	0.999999	26.80000	0.00000	0.999999	0.999999	26.80000	0.00000
0.018	27.20000	0.00000	0.999999	0.999999	27.20000	0.00000	0.999999	0.999999	27.20000	0.00000
0.019	27.60000	0.00000	0.999999	0.999999	27.60000	0.00000	0.999999	0.999999	27.60000	0.00000
0.020	28.00000	0.00000	0.999999	0.999999	28.00000	0.00000	0.999999	0.999999	28.00000	0.00000
0.021	28.40000	0.00000	0.999999	0.999999	28.40000	0.00000	0.999999	0.999999	28.40000	0.00000
0.022	28.80000	0.00000	0.999999	0.999999	28.80000	0.00000	0.999999	0.999999	28.80000	0.00000
0.023	29.20000	0.00000	0.999999	0.999999	29.20000	0.00000	0.999999	0.999999	29.20000	0.00000
0.024	29.60000	0.00000	0.999999	0.999999	29.60000	0.00000	0.999999	0.999999	29.60000	0.00000
0.025	30.00000	0.00000	0.999999	0.999999	30.00000	0.00000	0.999999	0.999999	30.00000	0.00000
0.026	30.40000	0.00000	0.999999	0.999999	30.40000	0.00000	0.999999	0.999999	30.40000	0.00000
0.027	30.80000	0.00000	0.999999	0.999999	30.80000	0.00000	0.999999	0.999999	30.80000	0.00000
0.028	31.20000	0.00000	0.999999	0.999999	31.20000	0.00000	0.999999	0.999999	31.20000	0.00000
0.029	31.60000	0.00000	0.999999	0.999999	31.60000	0.00000	0.999999	0.999999	31.60000	0.00000
0.030	32.00000	0.00000	0.999999	0.999999	32.00000	0.00000	0.999999	0.999999	32.00000	0.00000
0.031	32.40000	0.00000	0.999999	0.999999	32.40000	0.00000	0.999999	0.999999	32.40000	0.00000
0.032	32.80000	0.00000	0.999999	0.999999	32.80000	0.00000	0.999999	0.999999	32.80000	0.00000
0.033	33.20000	0.00000	0.999999	0.999999	33.20000	0.00000	0.999999	0.999999	33.20000	0.00000
0.034	33.60000	0.00000	0.999999	0.999999	33.60000	0.00000	0.999999	0.999999	33.60000	0.00000
0.035	34.00000	0.00000	0.999999	0.999999	34.00000	0.00000	0.999999	0.999999	34.00000	0.00000
0.036	34.40000	0.00000	0.999999	0.999999	34.40000	0.00000	0.999999	0.999999	34.40000	0.00000
0.037	34.80000	0.00000	0.999999	0.999999	34.80000	0.00000	0.999999	0.999999	34.80000	0.00000
0.038	35.20000	0.00000	0.999999	0.999999	35.20000	0.00000	0.999999	0.999999	35.20000	0.00000
0.039	35.60000	0.00000	0.999999	0.999999	35.60000	0.00000	0.999999	0.999999	35.60000	0.00000
0.040	36.00000	0.00000	0.999999	0.999999	36.00000	0.00000	0.999999	0.999999	36.00000	0.00000
0.041	36.40000	0.00000	0.999999	0.999999	36.40000	0.00000	0.999999	0.999999	36.40000	0.00000
0.042	36.80000	0.00000	0.999999	0.999999	36.80000	0.00000	0.999999	0.999999	36.80000	0.00000
0.043	37.20000	0.00000	0.999999	0.999999	37.20000	0.00000	0.999999	0.999999	37.20000	0.00000
0.044	37.60000	0.00000	0.999999	0.999999	37.60000	0.00000	0.999999	0.999999	37.60000	0.00000
0.045	38.00000	0.00000	0.999999	0.999999	38.00000	0.00000	0.999999	0.999999	38.00000	0.00000
0.046	38.40000	0.00000	0.999999	0.999999	38.40000	0.00000	0.999999	0.999999	38.40000	0.00000
0.047	38.80000	0.00000	0.999999	0.999999	38.80000	0.00000	0.999999	0.999999	38.80000	0.00000
0.048	39.20000	0.00000	0.999999	0.999999	39.20000	0.00000	0.999999	0.999999	39.20000	0.00000
0.049	39.60000	0.00000	0.999999	0.999999	39.60000	0.00000	0.999999	0.999999	39.60000	0.00000
0.050	40.00000	0.00000	0.999999	0.999999	40.00000	0.00000	0.999999	0.999999	40.00000	0.00000
0.051	40.40000	0.00000	0.999999	0.999999	40.40000	0.00000	0.999999	0.999999	40.40000	0.00000
0.052	40.80000	0.00000	0.999999	0.999999	40.80000	0.00000	0.999999	0.999999	40.80000	0.00000
0.053	41.20000	0.00000	0.999999	0.999999	41.20000	0.00000	0.999999	0.999999	41.20000	0.00000
0.054	41.60000	0.00000	0.999999	0.999999	41.60000	0.00000	0.999999	0.999999	41.60000	0.00000
0.055	42.00000	0.00000	0.999999	0.999999	42.00000	0.00000	0.999999	0.999999	42.00000	0.00000
0.056	42.40000	0.00000	0.999999	0.999999	42.40000	0.00000	0.999999	0.999999	42.40000	0.00000
0.057	42.80000	0.00000	0.999999	0.999999	42.80000	0.00000	0.999999	0.999999	42.80000	0.00000
0.058	43.20000	0.00000	0.999999	0.999999	43.20000	0.00000	0.999999	0.999999	43.20000	0.00000
0.059	43.60000	0.00000	0.999999	0.999999	43.60000	0.00000	0.999999	0.999999	43.60000	0.00000
0.060	44.00000	0.00000	0.999999	0.999999	44.00000	0.00000	0.999999	0.999999	44.00000	0.00000
0.061	44.40000	0.00000	0.999999	0.999999	44.40000	0.00000	0.999999	0.999999	44.40000	0.00000
0.062	44.80000	0.00000	0.999999	0.999999	44.80000	0.00000	0.999999	0.999999	44.80000	0.00000
0.063	45.20000	0.00000	0.999999	0.999999	45.20000	0.00000	0.999999	0.999999	45.20000	0.00000
0.064	45.60000	0.00000	0.999999	0.999999	45.60000	0.00000	0.999999	0.999999	45.60000	0.00000
0.065	46.00000	0.00000	0.999999	0.999999	46.00000	0.00000	0.999999	0.999999	46.00000	0.00000
0.066	46.40000	0.00000	0.999999	0.999999	46.40000	0.00000	0.999999	0.999999	46.40000	0.00000
0.067	46.80000	0.00000	0.999999	0.999999	46.80000	0.00000	0.999999	0.999999	46.80000	0.00000
0.068	47.20000	0.00000	0.999999	0.999999	47.20000	0.00000	0.999999	0.999999	47.20000	0.00000
0.069	47.60000	0.00000	0.999999	0.999999	47.60000	0.00000	0.999999	0.999999	47.60000	0.00000
0.070	48.00000	0.00000	0.999999	0.999999	48.00000	0.00000	0.999999	0.999999	48.00000	0.00000
0.071	48.40000	0.00000	0.999999	0.999999	48.40000	0.00000	0.999999	0.999999	48.40000	0.00000
0.072	48.80000	0.00000	0.999999	0.999999	48.80000	0.00000	0.999999	0.999999	48.80000	0.00000
0.073	49.20000	0.00000	0.999999	0.999999	49.20000	0.00000	0.99			

- Flow is calculated assuming normal flow approximation;
- In the original version of the code embedded in the excel file RTe-bookAgDegNormSub.xls the Manning-Strickler formulation is implemented;
- The water depth is calculated using a Chézy formulation, when only the Chézy coefficient is specified in the input text file. The Manning-Strickler formulation is implemented, when only the roughness height,  $k_c$ , and the coefficient  $\alpha$  in eq. (5.i) are given in the input text file. When all the three parameters are present, the program will ask the user which formulation they would like to use
- If the input channel length is longer than the maximum possible length of the fluvial reach, the program cannot perform the calculation. The maximum possible length of the fluvial reach,  $L_{max}$ , is defined as the maximum length of basin that the sediment supply can fill; at this length the sediment transport rate out of the basin drops precisely to zero (see Chapter 26 “Rivers Flowing into Subsiding Basins: Upward Concavity of Long Profile and Downstream Fining” of the e-book). For the Exner equation generalized for a river, eq. (7.ii),  $L_{max}$  is computed as

$$L_{max} = \frac{l_f(1 + \Lambda)\rho}{r_B} \frac{q_{tf}}{(1 - \lambda_p)^2} \quad (7.iv)$$

### List of variables

$x$	downchannel coordinate in m
$\Delta x$	spatial step length in m
$t$	temporal coordinate in seconds
$B_c$	channel width in m
$B_d$	width of the depositional area in m
$L_{max}$	maximum possible length of the modeled reach in m

#### INPUT

$q_w$	flood discharge per unit channel width $m^2/s$
$I_f$	flood intermittency
$D$	grain size in mm
$\lambda_p$	bed porosity
$k_c$	composite roughness height in mm
$SI$	initial bed slope
$q_{tf}$	sediment input rate of bed material load per unit channel width in $m^2/s$
$\sigma$	subsidence rate in mm/yr
$r_B$	ratio between the depositional and the channel width
$\Omega$	channel sinuosity
$\Lambda$	units of wash load deposited per unit of deposited bed material load
$L$	length of reach in m
$\Delta t$	time step in yr
$N_{tprint}$	number of time steps to printout
$N_{print}$	number of printouts
$M$	number of spatial intervals
$\alpha_u$	upwinding coefficient (1=full upwind, 0.5=central difference)
$\alpha_r$	coefficient in Manning-Strickler
$\alpha_s$	coefficient in sediment transport relation
$n_t$	exponent in sediment transport relation
$\tau_c^*$	reference Shields number in sediment transport relation
$\varphi_s$	fraction of bed shear stress due to skin friction
$R$	submerged specific gravity
$C_z$	non-dimensional Chézy friction coefficient

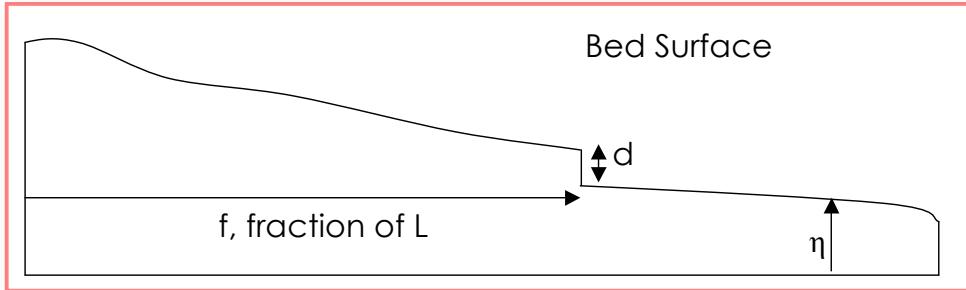
#### OUTPUT

$\eta$	bed surface elevation in m
$S$	bed slope
$H$	water depth in m
$\tau_b$	total boundary shear stress on bed surface N/m <sup>2</sup>
$q_t$	total bed material load in $m^2/s$
$q_t/q_{tf}$	ratio between the sediment transport and feed rate of bed material



## 8) AgDegNormalFault

The program AgDegNormalFault is an extension of AgDegNormal for sudden vertical faulting of the bed, as shown in the figure below.



The bed downstream of the point  $x = r_f L$  ( $0 < r_f < 1$ ), where  $x$  is a streamwise coordinate and  $L$  is the length of the modeled reach, is suddenly faulted downward by an amount  $\Delta\eta_f$  at time  $t_f$ . The eventual smearing out of the long profile is then computed.

The initial condition for the model is the bed at mobile equilibrium for the specified boundary conditions, i.e. characteristic flood discharge,  $Q_w$ , flood intermittency,  $I_f$ , and bed slope.

Input parameters of the model are

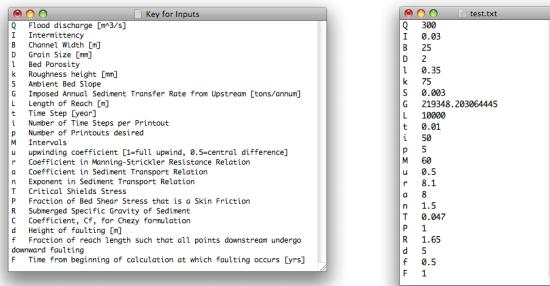
- the characteristic flood discharge,  $Q_w$ , and the flood intermittency,  $I_f$ ;
- the channel width,  $B_c$ ;
- the characteristic diameter of the sediment,  $D$ ;
- the bed porosity,  $\lambda_p$ ;
- the roughness height,  $k_c$ ;
- the ambient bed slope,  $S$ ;
- the imposed annual sediment transport rate,  $G_{tf}$ ;
- the fraction of reach length such that all points downstream of  $x = r_f L$  undergo downward faulting,  $r_f$ ;
- the height of faulting,  $\Delta\eta_f$ ;
- the time from beginning of calculation at which faulting occurs,  $t_f$ ;
- the length of the fluvial reach,  $L$ ;

The user also has to specify the following parameters to perform the calculations and to control the output:

- the number of sub-reaches,  $M$ ;
- the temporal step length,  $\Delta t$ ;
- the upwinding coefficient  $a_u$ ;
- the number of time steps to printout,  $N_{toprint}$ ;
- the numbers of printout in addition to the initial equilibrium state,  $N_{print}$ ;
- the coefficient  $\alpha_r$  in eq. (5.i);

- the parameters to define the load relation,  $\alpha_s, \tau_c^*$  and  $n_t$ , defined in eq. (5.ii);
- the constant to convert the total boundary shear stress to that due to skin friction,  $\varphi_s$ ;
- the submerged specific gravity of the sediment,  $R$ .

The text file with the definition of the input parameters and an example of input file are shown in the two windows below.



The outputs of the model are  $N_{\text{print}}$  longitudinal profiles at different times  $t_{\text{plot},i}$  defined by eq. (5.viii).

An example of the output file *Output.txt* is reported in the window below with an example of the additional parameters that the user may want to plot in the text file *Output1.txt*, i.e. bed slope, S, water depth, H, Shields number,  $\tau_b$ , and the total bed material load,  $q_t$ . The initial bed profile is the bed profile at mobile-bed equilibrium, given the water discharge, the flow intermittency and the initial bed slope.

	eta(w)	eta(s)	eta(w)	eta(w)	eta(s)	eta(s)
x(n)	0.00yr	0.59yr	1.00yr	1.58yr	2.09yr	2.50yr
0..0	30.000000	26.582256	24.809158	23.249491	21.527944	20.891666
166..7	29.500000	26.111788	24.455281	22.889826	21.182861	19.737065
333..3	29.000000	26.034626	24.266424	22.740465	20.975747	19.553769
500..0	28.500000	25.957464	24.145551	22.889803	20.887757	19.473247
666..7	28.000000	25.674499	24.119488	22.541379	20.833229	19.397696
833..3	27.500000	25.047334	23.514422	21.892524	20.241347	18.849952
1000..0	27.000000	24.935181	23.208546	21.536492	19.892597	18.527362
1166..7	26.500000	25.065362	23.517854	21.962503	20.426298	18.839165
1333..3	26.000000	25.447279	24.011413	21.644113	19.849948	18.761602
1500..0	25.500000	23.744817	23.379953	20.641488	19.888775	17.761638
1666..7	25.000000	23.576363	22.015897	20.224676	18.666281	17.379687
1833..3	24.500000	23.782189	22.399332	20.685474	19.110618	17.775636
2000..0	24.000000	23.233108	22.037682	20.261818	18.694553	17.397662
2166..7	23.500000	22.246000	21.073720	19.626242	17.742424	16.515588
2333..3	23.000000	22.000000	20.000002	19.697154	17.000057	16.200000
2500..0	22.500000	22.297128	21.189828	19.249527	17.782762	16.544535
2666..7	22.000000	21.668576	20.729977	18.750241	17.324638	16.133869
2833..3	21.500000	20.573796	19.617964	17.504133	15.222378	15.115089
3000..0	21.000000	20.329299	19.204815	16.956827	15.729491	14.676716
3166..7	20.500000	20.740207	19.740696	17.400000	16.000002	15.000000
3333..3	19.500000	19.320182	19.285938	17.841883	15.788848	14.723989
3500..0	19.500000	18.796166	18.035949	15.633991	14.549943	13.575219
3666..7	19.000000	18.559491	17.669416	15.830241	14.825828	13.119459
3833..3	18.500000	18.872538	18.097985	15.730441	14.623481	13.633566
4000..0	18.000000	18.035467	17.665102	15.137956	14.102865	13.179386
4166..7	17.500000	18.104044	17.830309	15.200055	14.112005	13.112000
4333..3	17.000000	16.711985	15.895239	12.969832	12.209521	11.446689
4500..0	16.500000	16.367932	16.421739	13.726331	12.834267	11.983784
4666..7	16.000000	16.066204	13.253797	13.080113	12.296631	11.519294
4833..3	15.500000	15.019411	13.685736	11.464694	10.833867	10.141865
5000..0	15.000000	14.526296	11.546381	10.000000	10.000000	9.400000
5166..7	14.500000	14.349516	14.549605	11.597116	10.391087	9.759857
5333..3	14.000000	14.055951	9.193649	18.926388	10.837117	8.200714
5500..0	13.500000	13.199416	7.737643	9.247061	8.298441	7.821974
5666..7	13.000000	12.905249	7.262189	8.571441	8.948358	8.374688
5833..3	12.500000	12.701962	7.827641	9.387317	8.948358	8.374688

## Notes:

- If the channel slope is negative and the water depth is not a number, “nan”, check the time step and the spatial step length. In particular, the time step may be too large or equivalently the spatial step length may be too small. Change these values and run the model again;
- The water depth is calculated using a Chézy formulation, when only the Chézy coefficient is specified in the input text file. The Manning-Strickler formulation is implemented, when only the roughness height,  $k_c$ , and the coefficient  $\alpha_r$  in eq. (5.i) are given in the input text file. When all the three parameters are present, the program will ask the user which formulation they would like to use.

### List of variables

$x$	streamwise coordinate in m
$\Delta x$	spatial step length in m
$t$	temporal coordinate in seconds
$C_f$	non-dimensional friction coefficient
INPUT	
$Q_w$	flood discharge in $m^3/s$
$I_f$	flood intermittency
$B_c$	channel width in m
$D$	characteristic grain size in mm
$\lambda_p$	bed porosity
$k_c$	composite roughness height in mm
$S_i$	ambient bed slope
$G_{tf}$	imposed annual sediment transport rate in tons/annum
$r_f$	fraction of reach length such that all points downstream of $x = r_f L$ undergo downward faulting.
$\Delta \eta$	height of faulting in m
$t_f$	time from beginning of calculation at which faulting occurs in years
$L$	length of reach in m
$\Delta t$	time step in years
$N_{toprint}$	number of time steps to printout
$N_{print}$	number of printouts
$M$	number of spatial intervals
$\alpha_u$	upwinding coefficient (1=full upwind, 0.5=central difference)
$\alpha_r$	coefficient in Manning-Strickler
$\alpha_s$	coefficient in sediment transport relation
$n_t$	exponent in sediment transport relation
$\tau_c^*$	reference Shields number in sediment transport relation
$\varphi_s$	fraction of bed shear stress due to skin friction
$R$	submerged specific gravity
$C_z$	non-dimensional Chézy friction coefficient
OUTPUT	
$\eta$	bed surface elevation in m
$S$	bed slope
$H$	water depth in m
$\tau_b$	total boundary shear stress on bed surface N/m <sup>2</sup>
$q_t$	total bed material load in $m^2/s$



## 9) AgDegNormGravMixPW

The program AgDegNormGravMixPW is an extension of AgDegNormal for sediment mixtures in gravel bed rivers where the channel bed material is transported as bedload only. Gravel-bed rivers tend to be poorly-sorted. During floods, bed material load consists almost exclusively of bedload. (Sand is often transported in copious quantities as washload during floods.) The surface material (armor or pavement) tends to be coarser than the substrate. By definition the median size  $D_{\text{sub}50}$  or geometric mean size  $D_{\text{sub}g}$  of the substrate is in the gravel range, but the substrate may contain up to 30% sand in the interstices of an otherwise clast-supported deposit. The Exner equation of conservation of channel bed sediment, eq. (5.iii), is thus written as

$$(1 - \lambda_p) \frac{\partial \eta}{\partial t} = -I_f \frac{\partial q_{bt}}{\partial x} \quad (9.i)$$

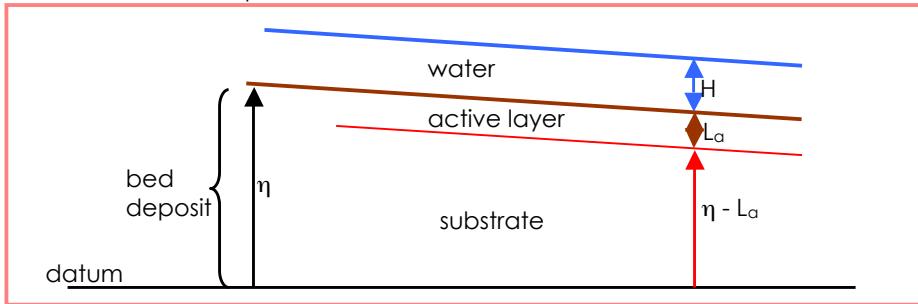
where  $\lambda_p$  is the bed porosity,  $\eta$  is the bed elevation above a datum,  $t$  is a temporal coordinate,  $I_f$  is the flow intermittency,  $x$  is a streamwise coordinate and  $q_{bt}$  denotes the total volumetric bedload transport rate per unit channel width.

The grain size distribution of the bed material is specified in terms of 12 size bounds,  $D_{bi}$  with  $i = 1 \dots 12$ , such that  $f_{fi}$  denotes the mass fraction of the sample that is finer than  $D_{bi}$ . The 12 bound diameters specify 11 grain size ranges defined by  $(D_{bi}, D_{b,i+1})$  and  $(f_{fi}, f_{f,i+1})$ . For each size range the model computes the characteristic diameter,  $D_i$ , and the fraction of sample in the  $i$ th size range,  $f_i$ , with the equations (9.ii a, b) reported below.

$$D_i = (D_{bi} \cdot D_{b,i+1})^{1/2} \quad f_i = f_{f,i+1} - f_{fi} \quad (9.\text{ii a, b})$$

The flow is assumed normal and the water depth can be computed with either a Manning-Strickler (implemented in the worksheet *Rte-bookAgDegNormGravMixPW.xls*) or a Chezy formulation can be used.

The exchange of sediment between the bedload and the bed deposit is modeled with a two-layer model for the channel bed. The bed deposit is divided in two regions, 1) the substrate and 2) the active (or surface or armor) layer, as represented in the figure below, where  $\eta$  denotes the channel bed elevation above a datum,  $L_a$  is the thickness of the active layer and  $H$  is the water depth.



The grain size distribution of the active layer is assumed to be constant in the vertical and it may vary in the streamwise direction and in time, i.e. the active layer is assumed well-mixed. The grain size distribution of the substrate, in principle, may vary in both streamwise and vertical direction, but it is constant in time. The only way the grain size distribution of the substrate may vary in time is by creating new substrate via bed aggradation. In the present model, the grain size distribution of the substrate is assumed to be constant in space and in time, therefore it works only for the cases of aggradation always and everywhere or degradation always and everywhere. In the case of aggradation followed by degradation, it is necessary to modify the code so that the vertical variation of the grain size distribution of the new substrate created by aggradation is stored in memory.

The evolution in time and space of the grain size distribution of the active layer is described by the grain-size based Exner equation, eq. (9.iii), that expresses the conservation of sediment in each grain size range.

$$(1 - \lambda_p) \left[ L_a \frac{\partial F}{\partial t} + (F_i - f_{ii}) \frac{\partial L_a}{\partial t} \right] = -l_i \frac{\partial Q_{bi} p_i}{\partial x} + l_i f_{ii} \frac{\partial q_{bi}}{\partial x} \quad (9.\text{iii})$$

$F_i$ ,  $p_i$  and  $f_{ii}$  respectively represent the fraction of sediment in the  $i$ th grain size range in the active layer, in the bedload and at the active-layer substrate interface. The grain size distribution of the sediment at the active layer substrate interface is defined in eq. (9.iv). It is equal to the weighted average between the grain size distribution of the active layer and of the bedload when the bed aggrades and the grain size distribution of the substrate when the bed degrades.

$$f_{ii} = \begin{cases} \alpha F_i + (1 - \alpha) p_i & \text{if } \frac{\partial \eta}{\partial t} > 0 \\ f_{sub,i} & \text{if } \frac{\partial \eta}{\partial t} < 0 \end{cases} \quad (9.\text{iv})$$

where  $\alpha$  is a user specified parameter and  $f_{sub,i}$  denotes the fraction of substrate material in the  $i$ th size range.

No attempt is made in this code to decompose the bed resistance into skin friction and form drag. The constant to convert total boundary shear stress to that due to skin friction,  $\varphi_s$ , is set equal to 1 and consequently the composite roughness height for the Manning-Strickler formulation,  $k_c$ , is equal to the roughness height due to skin friction,  $k_s$ , eq. (4.iv). The roughness height and the thickness of the active layer are computed with the liner functions of the diameter of the bed surface such that the 90% of the sediment is finer,  $D_{s90}$ , reported in eq. (9.v a, b)

$$k_s = n_k D_{s90} \quad L_a = n_a D_{s90} \quad (9.\text{v a, b})$$

where  $n_k$  and  $n_a$  are user specified order-one non dimensional constants.

To compute the bedload transport rate the user can choose from two surface-based bedload transport formulations; those of Parker (1990) and Wilcock and Crowe (2003). In the relation of Parker (1990) the surface grain size distributions need to be renormalized to exclude sand before specification as input to the program. This step is neither necessary nor desirable in the case of the relation of Wilcock and Crowe (2003), where the sand plays an important role in mediating the gravel bedload transport.

For a detailed description of the theory and of the model the user may refer to Chapters 4 "Relations for the Conservation of Bed Sediment", 7 "Relations for 1D Bedload Transport", 17 "Aggradation and Degradation of Rivers Transporting Gravel Mixtures" and 18 "Mobile and Static Armor in Gravel-bed Streams" of the e-book.

The upstream boundary conditions are given in terms of specified values of water discharge per unit width,  $q_w$ , and volumetric total gravel input rate per unit width,  $q_{btf}$ , with its grain size distribution, given in terms of bound diameters,  $D_{bi}$ , and percent finer,  $p_{ff}$ . The downstream boundary condition is a fixed bed elevation  $\eta_d$  at the downstream end of the modeled reach, i.e. the Exner equation (9.i) is not implemented at node M+1.

The initial conditions are a specified initial bed profile here simplified to a specified initial bed slope  $S_i$  ( $S_{tbl}$  in the excel file *RTe-bookAgDegNormGravMixPW.xls*), see eq. (5.vii). The grain size distributions of the active layer,  $F_{fi}$ , and of the substrate,  $F_{subfi}$ , have to be specified and they are assumed to be the same in each computational node. These grain size distributions are expressed in terms of bound diameters,  $D_{bi}$ , and percent finer, respectively  $F_{fi}$  for the bed surface and  $F_{subfi}$  ( $F_{fsi}$  in the excel file *RTe-bookAgDegNormGravMixPW.xls*).

Input parameters of the model are

- the characteristic flood discharge per unit channel width,  $q_w$ , and the flood intermittency,  $I_r$ ;
- the volumetric gravel input rate per unit channel width,  $q_{btf}$ ;
- the downstream bed elevation,  $\eta_d$ ;
- the bed porosity,  $\lambda_p$ ;
- the initial bed slope,  $S_i$ ;
- the length of the fluvial reach,  $L$ ;
- the grain size distributions of the substrate of the active layer and of the gravel feed rate. The bound diameters,  $D_{bi}$ , are the same for the three distributions;

The user also has to specify the following parameters to perform the calculations and to control the output:

- the number of sub-reaches,  $M$ ;
- the temporal step length,  $\Delta t$ ;
- the upwinding coefficient  $a_u$ ;
- the number of time steps to printout,  $N_{toprint}$ ;
- the numbers of printout in addition to the initial equilibrium state,  $N_{print}$ ;
- the coefficient  $\alpha_r$  in eq. (4.iv);
- the coefficient  $n_k$  in eq. (9.v a) to compute the roughness height;
- the coefficient  $n_a$  in eq. (9.v b) to compute the thickness of the active layer;
- the load relation, i.e. Parker (1990) or Wilcock and Crowe (2003);
- the submerged specific gravity of the sediment,  $R$ ;
- the parameter  $\alpha$  in eq. (9.iv) that governs the grain size distribution of the sediment at the active layer-substrate interface during bed aggradation.

The text file with the definition of the input parameters and an example of input file are shown in the two windows below.

```

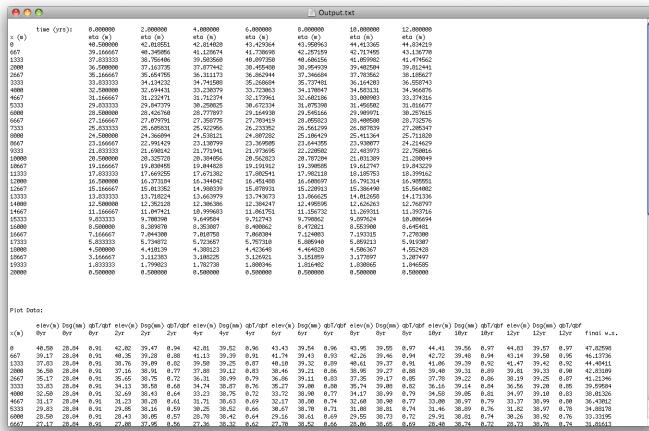
d water discharge/width [m^2/s]
g gravel input [m^2/s]
i intermittency
e base level [m]
s initial bed slope
l reach length [m]
t total time [days]
M no. of intervals (100 or less)
p no. of prints
i no. of iterations per print
K factor by which Ds90 is multiplied for roughness height
R factor by which Ds90 is multiplied for active layer thickness
r coefficient in Manning-Strickler relation
S submerged specific gravity of gravel
l bed porosity, gravel
u upwinding coefficient for load spatial derivatives in Exner equation (>0.5
suggestion
c coefficient for material transferred to substrate as bed aggrades
C coefficient, Cf, in the Chezy formulation

Diameter [mm] GSD Feed GSD Surface GSD Substrate
Diameter [mm] GSD Feed GSD Surface GSD Substrate
Diameter [mm] GSD Feed GSD Surface GSD Substrate
.
.
.
Diameter [mm] GSD Feed GSD Surface GSD Substrate

[exactly 12 different diameters should be entered]

```

The output parameters of the model printed in the text file *Output.txt* are the bed elevation,  $\eta$ , the geometric mean diameter of the bed surface,  $D_{sg}$ , and the ratio between the bedload transport rate and the feed rate.



The additional parameters that the user may print at times  $t_{plot,j}$  in the text file *Output1.txt* are the channel slope,  $S$ , the water depth,  $H$ , the total bedload transport rate,  $q_{bt}$ , the diameter of the bed surface such that 90% of the sediment is finer,  $D_{s90}$ , the Shields number,  $\tau_{sg}^*$ , i.e. the non-dimensional Shields number, defined as

$$\tau_{sg}^* = \frac{\tau_b}{(\rho_s - \rho)gD_{sg}} \quad (9.vi)$$

where  $\tau_b$  is the bed shear stress,  $\rho_s$  and  $\rho$  respectively denote the density of the sediment and of the water,  $g$  is the acceleration of gravity and  $D_{sg}$  is the geometric mean diameter of the bed surface.

An example of the file *Output1.txt* is reported in the windows below.

	eta(s)	S1	H(m)	tousg	qbT	dsg(mm)	D90s(mm)
x(n)	12yr	12yr	2.991764	8.116669	0.000974	39.569752	126.419631
0.0	44.834219	0.002546	2.991764	8.116669	0.000949	39.496145	126.188260
6666.7	43.136778	0.002528	3.000591	8.116018	0.000923	39.417175	125.942750
1333.3	41.474562	0.002493	3.009952	8.115395	0.000898	39.325903	125.692382
2000.0	39.812441	0.002467	3.018653	8.114738	0.000874	39.253107	125.451495
2666.7	38.185627	0.002440	3.027837	8.114034	0.000850	39.189994	125.227017
3333.3	36.558743	0.002414	3.037996	8.113396	0.000825	39.095668	125.006236
4000.0	34.966877	0.002386	3.046386	8.112853	0.000803	38.984543	124.793592
4666.7	33.374316	0.002362	3.055806	8.112132	0.000781	38.867957	124.612186
5333.3	31.816677	0.002337	3.065108	8.111356	0.000759	38.749807	124.436668
6000.0	30.257615	0.002313	3.074338	8.110887	0.000737	38.764448	124.241219
6666.7	28.732576	0.002289	3.083552	8.110377	0.000715	38.747638	124.054262
7333.3	27.207537	0.002265	3.092766	8.109558	0.000693	38.747538	123.869395
8000.0	25.701398	0.002243	3.101569	8.108937	0.000670	38.747627	123.769490
8666.7	24.246429	0.002221	3.110269	8.108600	0.000648	38.498895	123.692667
9333.3	22.759016	0.002200	3.118827	8.107973	0.000626	38.476462	123.469581
10000.0	21.288949	0.002188	3.127867	8.107298	0.000604	38.445981	123.334834
10666.7	19.843226	0.002161	3.134835	8.106926	0.000583	38.436492	123.182882
11333.3	18.399162	0.002143	3.142419	8.106528	0.000561	38.322188	123.049933
12000.0	16.985551	0.002126	3.149530	8.105970	0.000540	38.270399	122.926537
12666.7	15.564802	0.002110	3.156179	8.105568	0.000519	38.247872	122.797666
13333.3	14.171333	0.002094	3.162300	8.105267	0.000503	38.178994	122.675527
14000.0	12.768797	0.002083	3.167995	8.104896	0.000487	38.117732	122.451389
14666.7	11.393716	0.002071	3.173018	8.104569	0.000476	38.089084	122.334834
15333.3	10.006694	0.002061	3.177536	8.104325	0.000458	38.050833	122.210943
16000.0	8.645461	0.002052	3.181413	8.104099	0.000452	38.008768	122.363821
16666.7	7.278907	0.002044	3.185176	8.103807	0.000437	37.981310	122.277338
17333.3	5.912307	0.002036	3.187770	8.103573	0.000422	37.954447	122.188262
18000.0	4.552428	0.002029	3.189446	8.103365	0.000409	37.926462	122.139952
18666.7	3.207497	0.002029	3.191394	8.103059	0.000396	37.916469	122.079428
19333.3	1.846505	0.002028	3.199785	8.103554	0.000377	37.917211	122.079463
20000.0	0.500000	0.002028	3.195551	8.103272	0.000358	37.976328	121.947362

## Notes:

- In the case of the load relation due to Parker (1990), the grain size distributions are automatically renormalized because the relation is for the transport of gravel only
- In the case of the load relation due to Wilcock-Crowe (2003), the sand and the fine sediment are retained for the computation;
- The user will be prompted by the program as to which bedload relation he would like to use;
- The input grain size distributions may be on a 0-100% or a 0.00-1.00 scale, and the program will automatically scale;
- The input grain size distributions must have bounds at 0% and 100% (1.00) to properly perform the calculation. If the user does not input the bounds the program will automatically interpolate upper and lower bounds  $D_{bu}$  and  $D_{bl}$  such that  $f_{fu} = 100$  (1.00) and  $f_{fl} = 0$ ;
- The water depth is calculated using a Chézy formulation, when only the Chézy coefficient is specified in the input text file. The Manning-Strickler formulation is implemented, when only the coefficients  $a_r$  and  $n_r$  in equations (4.iv) and (4.v) are given in the input text file. When all the three parameters are present, the program will ask the user which formulation they would like to use.

### List of variables

$x$	downchannel coordinate in m
$\Delta x$	spatial step length in m
$t$	temporal coordinate in seconds
$k_s$	roughness height due to skin friction only in mm
$C_f$	Manning-Strickler friction coefficient
$L_a$	thickness of the active layer in m
$f_{fi}$	fraction of sediment finer than the $i$ th bound diameter
$f_i$	fraction of sediment in the $i$ th grain size range
INPUT	
$q_w$	flood discharge per unit channel width $m^2/s$
$I_f$	flood intermittency
$\lambda_p$	bed porosity
$S_i$	initial bed slope
$\eta_d$	bed elevation at the downstream end of the reach in m
$Q_{btf}$	total bedload input rate per unit channel width in $m^2/s$
$L$	length of reach in m
$\Delta t$	time step in days
$N_{toprint}$	number of time steps to printout
$N_{print}$	number of printouts
$M$	number of spatial intervals
$a_u$	upwinding coefficient (1=full upwind, 0.5=central difference)
$\alpha_r$	coefficient in Manning-Strickler formulation
$\alpha$	parameter in eq. (9.iv) that governs the grain size distribution of the sediment at the active layer-substrate interface during bed aggradation
$n_k$	coefficient to estimate the roughness height as a function of the $D_{s90}$
$n_a$	coefficient to estimate the thickness of the active layer as a function of the $D_{s90}$
$R$	submerged specific gravity of the sediment
$C_z$	non-dimensional Chézy friction coefficient
$D_{bi}$	bound diameters of the grain size distribution
$D_i$	characteristic diameters of the grain size distribution defined in eq. (9.ii a)
$F_i, f_{subi}, p_i, f_{li}$	fractions of sediment in the $i$ th grain size range of the bed surface, of the substrate, of the bedload and at the active layer-substrate interface

OUTPUT

$\eta$	bed surface elevation in m
$S$	bed slope
$H$	water depth in m
$\tau_{sg}$	Shields number computed with eq. (9.vi)
$q_{bt}$	total bedload transport rate per unit channel width in $m^2/s$
$q_{bt}/q_{btf}$	ratio between the total bedload transport and the feed rate per unit channel width
$D_{sg}$	geometric mean diameter of the bed surface in mm
$D_{s90}$	diameter of the active layer such that the 90% of the sediment is finer in mm

References:

- Parker, G., 1990, Surface-based bedload transport relation for gravel rivers, *Journal of Hydraulic Research*, 28(4): 417-436.  
Wilcock, P. R., and Crowe, J. C., 2003, Surface-based transport model for mixed-size sediment, *Journal of Hydraulic Engineering*, 129(2), 120-128.

# 10) AgDegNormGravMixHyd

This program is a close relative of AgDegNormGravMixP. It computes aggradation and degradation in gravel-bed river subject to a repeated hydrograph. The sediment is modeled as mixture of different grain sizes and the bedload formulation is that of Parker (1990) that was derived to compute the transport of gravel only. For a more extensive description of the theory and of the model the user can refer to Chapter 19 "Effect of Hydrograph on Morphology of Gravel-bed Streams" of the e-book.

To perform a numerical calculation with a flow hydrograph, the actual hydrograph must be specified in terms of W constant water discharges  $Q_w$ , where  $w = 1 \dots W$ , each extending for time duration  $\Delta t_w$ . The river is assumed to be morphologically inactive when it is not in flood. The morphodynamic evolution is computed solving the equation of sediment continuity (i.e. Exner equation). A short flood time step,  $\Delta t_f$ , is used to solve the Exner equation. It is convenient to choose each time duration  $\Delta t_w$  to be an integer number,  $n_{step,w}$ , of the flood time step, so that

$$\Delta t_w = n_{step,w} \Delta t_f \quad (10.i)$$

Once  $\Delta t_f$  is specified along with  $Q_w$  and  $n_{step,w}$  for all the W constant water discharges, it is possible to determine a discharge  $Q_p$  for each time step  $p = 1 \dots P$  in the hydrograph, where

$$P = \sum_{w=1}^W n_{step,w} \quad (10.ii)$$

The flow is assumed locally normal and it is computed with a Manning-Strickler formulation, eq. (4.iv). No attempt is made to decompose the bed resistance into skin friction and form drag. The roughness height and the active layer thickness are assumed functions of the grain size distribution of the bed surface and they are computed with eqs. (9.v a, b).

Initial conditions are the initial bed profile and the grain size distribution of the bed surface and of the substrate. The initial bed profile, as in AgDegNormGravMixPW, is specified in terms of a constant slope,  $S_l$  ( $S_{fbl}$  in the excel worksheet RTe-bookAgDegGravMixHyd.xls), and a downstream bed elevation,  $\eta_d$ . The grain size distributions of the bed surface and of the substrate are specified in terms of percent finer and  $n_{pp}$  bound diameters,  $D_{bi}$ , with  $n_{pp} = 2 \sim 9$ . All the grain size distributions must be specified with the same  $n_{pp}$  bound diameters and the fractions of sediment finer than the bound diameters can be given either on a 0 – 100 or 0 – 1 scale. If the fraction of sediment finer than the lower (upper) bound of the input distributions is grater than zero (100 or 1), the program automatically computes a new lower (upper),  $D_{bnew}$ , bound of the distribution such that the fraction of sediment finer than  $D_{bnew}$  is equal to 0 (100 or 1). The fraction of sand ( $D_{b,i} < 2 \text{ mm}$ ) in the input the grain size distributions has to be removed and the distributions need to be renormalized before starting the calculations.

The upstream boundary condition is given in terms of W values of water discharge, bedload input rate and  $n_{step,w}$ .

Enrica 11/19/2009 2:02 AM

**Comment [1]:** Andrew, did you also implement a Chezy formulation?

**Answer:** I did not implement a Chezy formulation... would you like me too? It would just be adding some inputs and adding a line of code to the depth calculation.