

# Nonlinear Test Models

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Test models for developing data assimilation schemes:

- Simplified, efficient to compute
- Capture the key dynamical process of interest, for which there are challenges in performing data assimilation.

Lorenz 1963 model: (L63)

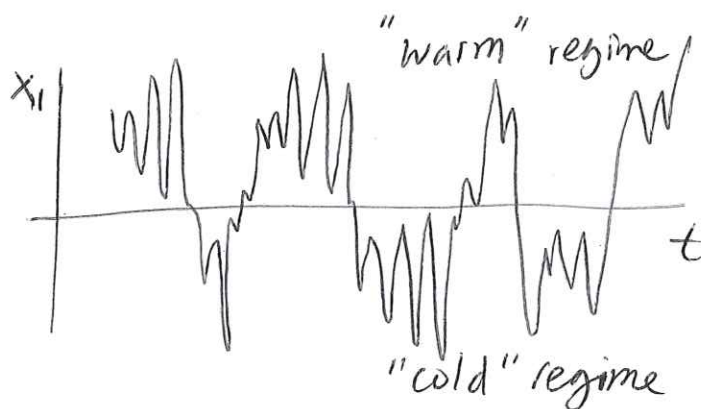
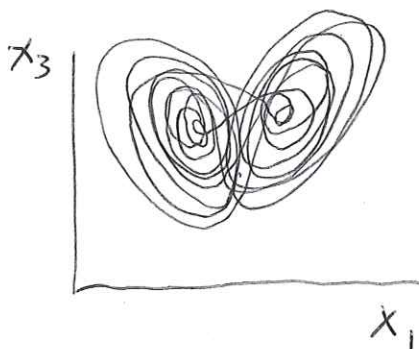
$$\vec{x} = \begin{pmatrix} x_1 \\ x_2 \\ x_3 \end{pmatrix}$$

$$\begin{cases} \frac{dx_1}{dt} = \sigma(x_2 - x_1), \\ \frac{dx_2}{dt} = x_1(\rho - x_3) - x_2, \\ \frac{dx_3}{dt} = x_1 x_2 - \beta x_3 \end{cases}$$

→ an approximation of Navier-Stokes equations, which describes convective movement in the atmosphere.

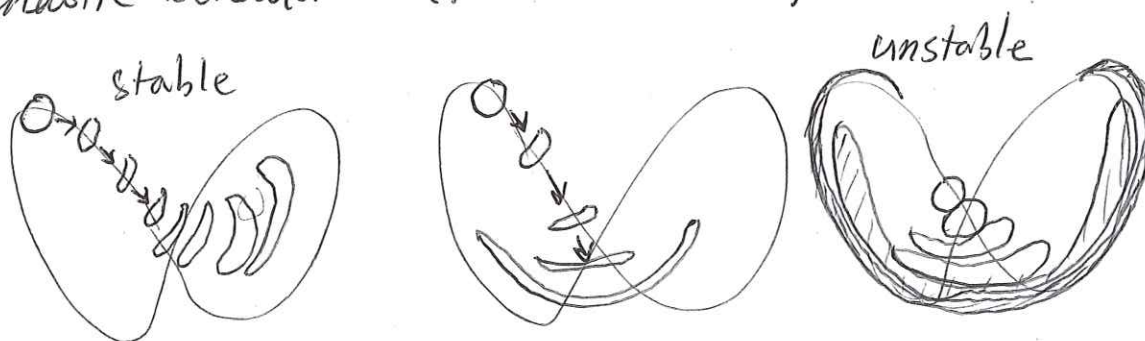
→ a typical configuration:  $\sigma=10$ ,  $\beta=8/3$ ,  $\rho=28$   
displays "chaotic" behavior: sensitive to initial condition

Solution phase space:



chaotic behavior: (Palmer et al. 2007)

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- Consider an initial volume of uncertain conditions the trajectories of points inside the volume end up in a new volume after some integration time, and the volume changes shape.
- Somewhere in the phase space (leftmost one). features relatively stable trajectories, with volumes slowly stretching. "errors grow slowly"; while somewhere (rightmost one) features very rapid error growth and bifurcation of the volume into two regimes.
- sensitivity to initial condition  $\rightarrow$  predictability
- linear error growth: stretching into a hyper ellipsoid  
nonlinear " " phase: folding is needed for the solution to stay within the bounds.  
 $\rightarrow$  asymptote to a "strange attractor"

## Lorenz 1996 model (L96)

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Defined on a 1D cyclic domain with  $n$  grid points:

$$i = 1, 2, \dots, n$$

$$\frac{dx_i}{dt} = \underbrace{(x_{i+1} - x_{i-2}) x_{i-1}}_{\substack{\text{similar to} \\ \text{nonlinear} \\ \text{advection}}} - \underbrace{x_i^2}_{\substack{\text{dissipation} \\ \sim \nabla^2 u \\ \sim -u}} + \underbrace{F}_{\text{constant "solar" forcing}}$$

- typical choice  $n=40$ ,  $F=8$
- the cyclic domain mimics the mid-latitude belt of 32000 km with grid spacing 800 km, which is about the Rossby deformation radius.
- Solution features "Rossby-like" waves with unstable westward phase velocity, collectively move eastward.
- 0.2 time unit  $\sim$  1 day in real atmosphere

- conserves energy norm:  $E = \sum_{i=1}^n x_i^2 / 2$

$$\sum_{i=1}^n x_i \frac{dx_i}{dt} = \sum_{i=1}^n \left[ x_i x_{i-1} (x_{i+1} - x_{i-2}) - x_i^2 + F x_i \right]$$

$$\frac{dE}{dt} = -E + F \sum_{i=1}^n x_i / 2 = 0$$

- solar forcing scale  $\sim$  1 day, dissipation  $\sim$  5 days.  
Decorrelation scale  $\sim$  3 days.

How to numerically integrate a model?

- See "Numerical Method" courses for more details.

Discretization in space and time: (L63, L96 already discretized in space).

Euler forward in time:

for a model  $\frac{d\vec{x}}{dt} = f(\vec{x})$

$$\vec{x}^{t+1} = \vec{x}^t + \Delta t \cdot f(\vec{x}^t) \quad (\text{explicit scheme}).$$

Runge-Kutta 4th-order scheme: (RK4).

$$\vec{x}_{(1)} = \vec{x}^t + \frac{\Delta t}{2} f(\vec{x}^t)$$

$$\vec{x}_{(2)} = \vec{x}^t + \frac{\Delta t}{2} f(\vec{x}_{(1)})$$

$$\vec{x}_{(3)} = \vec{x}^t + \Delta t f(\vec{x}_{(2)})$$

$$\vec{x}^{t+1} = \vec{x}^t + \Delta t \left[ \frac{f(\vec{x}^t)}{6} + \frac{f(\vec{x}_{(1)})}{3} + \frac{f(\vec{x}_{(2)})}{3} + \frac{f(\vec{x}_{(3)})}{6} \right]$$

• Euler forward requires shorter  $\Delta t$  to achieve numerical stability than RK4.

Implicit scheme, e.g. Crank-Nicholson

$$\vec{x}^{t+1} = \vec{x}^t + \Delta t f\left(\frac{\vec{x}^t + \vec{x}^{t+1}}{2}\right)$$

→ since  $\vec{x}^{t+1}$  is on both left and right hand side, need to solve matrix inversion!

→ numerically stable!