

¹ **neonSoilFlux: An R Package for Continuous
2 Sensor-Based Estimation of Soil CO₂ Fluxes**

³ John Zobitz¹ Edward Ayres² Zoey Werbin³ Ridwan Abdi¹
⁴ Natalie Ashburner-Wright⁵ Lillian Brown⁵
⁵ Ryan Frink-Sobierajski⁵ Lajntxiag Lee¹ Dijonë Mehmeti¹
⁶ Christina Tran⁵ Ly Xiong¹ Naupaka Zimmerman^{4,5}

⁷ ¹ Augsburg University, 2211 Riverside Avenue, Minneapolis, MN 55454

⁸ ² National Ecological Observatory Network, Battelle, 1685 38th Street, Suite 100, Boulder, CO
⁹ 80301

¹⁰ ³ Boston University, 5 Cummington Street, Boston, MA 02215

¹¹ ⁴ University of Kansas, 1450 Jayhawk Boulevard, Lawrence, KS 66045

¹² ⁵ University of San Francisco, 2130 Fulton Street, San Francisco, CA 94117

¹³ **Acknowledgments**

¹⁴ John Zobitz acknowledges Kathleen O'Rourke for code development. Naupaka Zimmerman
¹⁵ thanks technical staff at USF for support with field gear assembly and shipping. We thank the
¹⁶ NEON field staff and assignable assets teams for facilitating each of the six NEON site visits.

¹⁷ We are grateful to LI-COR technical staff for helpful discussions about optimal soil chamber
¹⁸ sampling methods. This work was supported by NSF DEB grant #2017829 awarded to John
¹⁹ Zobitz, and NSF DEB grant #2017860 awarded to Naupaka Zimmerman. This material is
²⁰ based in part upon work supported by the National Ecological Observatory Network (NEON),
²¹ a program sponsored by the U.S. National Science Foundation (NSF) and operated under
²² cooperative agreement by Battelle. We also thank the reviewers and subject editor for their
²³ constructive feedback.

²⁴ **Conflict of Interest Statements**

²⁵ None of the authors have a financial, personal, or professional conflict of interest related to this
²⁶ work.

²⁷ **Author Contributions**

²⁸ Conceptualization: John Zobitz, Naupaka Zimmerman; Methodology: Edward Ayres, John
²⁹ Zobitz, Naupaka Zimmerman; Software: John Zobitz, Naupaka Zimmerman, Zoey Werbin,
³⁰ Edward Ayres, Dijonë Mehmeti, Ridwan Abdi, Ly Xiong, Lajntxiag Lee; Validation: John
³¹ Zobitz, Naupaka Zimmerman; Formal Analysis: John Zobitz, Naupaka Zimmerman, Dijonë
³² Mehmeti, Ridwan Abdi, Ly Xiong, Lajntxiag Lee; Investigation: John Zobitz, Naupaka Zimmerman,
³³ Ryan Frink-Sobierajski, Christina Tran, Natalie Ashburner-Wright, Lillian Brown;
³⁴ Resources: John Zobitz, Naupaka Zimmerman; Data curation: John Zobitz, Naupaka Zimmerman,
³⁵ Dijonë Mehmeti, Ly Xiong; Writing – original draft: John Zobitz, Naupaka Zimmerman;
³⁶ Writing – review and editing: John Zobitz, Naupaka Zimmerman, Zoey Werbin, Edward Ayres,
³⁷ Christina Tran, Dijonë Mehmeti, Ly Xiong; Visualization: John Zobitz, Naupaka Zimmerman,

³⁸ Dijonë Mehmeti, Ridwan Abdi, Ly Xiong; Supervision: John Zobitz, Naupaka Zimmerman;
³⁹ Project Administration: John Zobitz, Naupaka Zimmerman; Funding Acquisition: John Zobitz,
⁴⁰ Naupaka Zimmerman.

⁴¹ **Data Availability**

⁴² Data available via <https://doi.org/10.5281/zenodo.17516319> (Zobitz & Zimmerman, 2025).
⁴³ Field-collected data, `neonSoilFlux` calculated outputs, and manuscript-generating code are
⁴⁴ provided within this repository.

45 **1 Abstract**

- 46 1. Accurate quantification of soil carbon fluxes is essential to reduce uncertainty in estimates
47 of the terrestrial carbon sink. However, these fluxes vary over time and across ecosystem
48 types and so it can be difficult to estimate them accurately across large scales. The flux
49 gradient method estimates soil carbon fluxes using co-located measurements of soil CO₂
50 concentration, soil temperature, soil moisture, and other soil properties. The National
51 Ecological Observatory Network (NEON) provides such data across 20 ecoclimatic domains
52 spanning the continental U.S., Puerto Rico, Alaska, and Hawai‘i.
- 53 2. We present an R software package (`neonSoilFlux`) that acquires soil environmental data
54 to compute half-hourly soil carbon fluxes for each soil replicate plot at a given terrestrial
55 NEON site. To assess the computed fluxes, we visited six focal NEON sites and measured
56 soil carbon fluxes using a closed-dynamic chamber approach.
- 57 3. Outputs from the `neonSoilFlux` showed agreement with measured fluxes (R^2 between
58 measured and `neonSoilFlux` outputs ranging from 0.04 to 0.81 depending on calculation
59 method used); measured outputs generally fell within the range of calculated uncertainties
60 from the gradient method. Calculated fluxes from `neonSoilFlux` aggregated to the daily
61 scale exhibited expected site-specific seasonal patterns.
- 62 4. While the flux gradient method is broadly effective, its accuracy is highly sensitive
63 to site-specific inputs, including the extent to which gap-filling techniques are used to
64 interpolate missing sensor data and to estimates of soil diffusivity and moisture content.
65 Future refinement and validation of `neonSoilFlux` outputs can contribute to existing
66 databases of soil carbon flux measurements, providing near real-time estimates of a critical
67 component of the terrestrial carbon cycle.

68 **1.1 Keywords**

69 Soil carbon, carbon dioxide, flux gradient, carbon cycle, field validation, soil respiration,
70 ecosystem variability, diffusion

71 **2 Introduction**

72 Soils contain the planet's largest reservoir of terrestrial carbon (Jobbág & Jackson, 2000). A
73 critical component of this reservoir is soil organic matter, the accumulation of which is influenced
74 by biotic factors such as above-ground plant inputs (Jackson et al., 2017). These inputs in
75 turn are influenced by environmental factors such as growing season length, temperature, and
76 moisture (Desai et al., 2022), which also affect the breakdown of soil organic matter and its
77 return to the atmosphere. Across heterogeneous terrestrial landscapes, the interplay between
78 these biotic and abiotic factors influence the size of the soil contribution to the terrestrial
79 carbon sink (Friedlingstein et al., 2025). However, the heterogeneity of these processes across
80 diverse ecosystems in the context of rapid environmental change leads to large uncertainty
81 about the magnitude of this sink in the future, and thus there remains a pressing need to
82 quantify changes in soil carbon pools and fluxes across scales.

83 Ecological observation networks such as the United States' National Ecological Observatory
84 Network (NEON) and others (e.g. the globally-distributed FLUXNET or the European Inte-
85 grated Carbon Observation System) present a significant advancement in the nearly continuous
86 observation of biogeochemical processes at the continental scale. Notably, at 47 terrestrial
87 sites across the continental United States that span 20 ecoclimatic domains, NEON provides
88 half-hourly measurements of soil CO₂ concentration, temperature, and moisture at different
89 vertical depths. Each of these NEON sites also encompasses measurements of the cumulative
90 sum of all ecosystem carbon fluxes in an airshed using the eddy covariance technique (Baldocchi,

91 Soil observations provided by NEON are on the same timescale and standardized with
92 eddy covariance measurements from FLUXNET. These types of nearly continuous observational
93 data (NEON and FLUXNET) can be used to reconcile differences between model-derived
94 or data-estimated components of ecosystem carbon flux (Jian et al., 2022; Luo et al., 2011;
95 Phillips et al., 2017; J. Shao et al., 2015; P. Shao et al., 2013; Sihi et al., 2016).

96 Estimated or observed soil carbon fluxes are a key metric for understanding change in soil
97 carbon pools over time (Bond-Lamberty et al., 2024). A soil carbon flux to the atmosphere (F_S ,
98 units $\mu\text{mol m}^{-2} \text{s}^{-1}$), represents the aggregate process of transfer of soil CO_2 to the atmosphere
99 from physical and biological processes (e.g. diffusion and respiration). Soil carbon fluxes can
100 be assumed to encompass soil carbon respiration from autotrophic or heterotrophic sources
101 (Davidson et al., 2006) and modeled with a exponential Q_{10} paradigm (Bond-Lamberty et al.,
102 2004; Chen & Tian, 2005; Hamdi et al., 2013).

103 One common method by which F_S is measured in the field is through the use of soil chambers
104 in a closed, well-mixed system (Norman et al., 1997) with headspace trace gas concentrations
105 measured with an infrared gas analyzer (IRGA). F_S can also be estimated from soil CO_2
106 measurements at different depths in the soil using the flux-gradient method (Maier & Schack-
107 Kirchner, 2014). Closed-chamber IRGA measurements, while being the most common method,
108 require either frequent in-person site visits or expensive and fragile automated systems. The
109 potential of the gradient method is that fluxes can be estimated from continuous data recorded
110 by robust solid-state sensors. The flux-gradient method is an approach that uses conservation of
111 mass to calculate flux at a vertical soil depth z at steady state by applying Fick's law of diffusion.
112 A simplifying assumption for the flux-gradient method is that there is no mass transfer in the
113 other spatial dimensions x and y (Maier & Schack-Kirchner, 2014). The diffusivity profile, a
114 key component of this calculation, varies across the soil depth as a function of soil temperature,
115 soil volumetric water content, atmospheric air pressure, and soil bulk density (Millington &

116 Shearer, 1971; Moldrup et al., 1999; Sallam et al., 1984).

117 Databases such as the Soil Respiration Database (SRDB) or the Continuous Soil Respiration
118 Database (COSORE) add to the growing network of resources for making collected observations
119 of soil fluxes available to other researchers (Bond-Lamberty, 2018; Bond-Lamberty et al., 2020;
120 Bond-Lamberty & Thomson, 2010; Jian et al., 2021; Jiang et al., 2024). However, these
121 databases currently encompass primarily direct soil measurements of fluxes (i.e. those using
122 methods like the closed-chamber method described above). Currently, NEON provides all
123 measurements to calculate F_S from Fick's law, but soil flux as a derived data product was
124 descoped from the initial network launch due to budget constraints (Berenbaum et al., 2015).
125 Deriving estimates of F_S using continuous sensor data across NEON sites using NEON data
126 thus remains a high priority.

127 This study describes an R software package, `neonSoilFlux`, that computes a standardized
128 estimate of F_S at all terrestrial NEON sites using the flux-gradient method. Using direct
129 chamber-based field observations of soil carbon dioxide flux from a subset of terrestrial NEON
130 sites spanning six states, we provide a direct validation of F_S from `neonSoilFlux`. While
131 open source R software tools currently exist for processing chamber-based flux measurements
132 (Jurasinski et al., 2022; Pedersen, 2024; Rheault et al., 2024; Wilson et al., 2024; Zhao, 2019),
133 to our knowledge this is the first package that incorporates NEON data directly.

134 Key objectives of this study are to:

135 1. Apply the flux-gradient method to estimate soil CO₂ flux from continuous sensor mea-
136 surements across six NEON sites.
137 2. Benchmark estimated soil carbon fluxes against field measurements (e.g. direct chamber
138 measurements of soil flux).

- 139 3. Identify sources of error in the flux-gradient approach across diverse sites in order to
140 guide future work.

141 **3 Materials and Methods**

142 **3.1 Field methods**

143 **3.1.1 Focal NEON Sites**

144 In order to acquire field data to validate model predictions of flux, we selected six terrestrial
145 NEON sites for analysis. We conducted roughly week-long field measurement campaigns at
146 these sites, which span a range of environmental gradients and terrestrial domains (Table 1).
147 SJER, SRER, and WREF were visited during May and June of 2022, and WOOD, KONZ, and
148 UNDE during May and June of 2024. Permits or waivers were sought and approved prior to
149 field work at all six sites. In 2022, research activities were conducted whole or in part on the
150 Wind River Experimental Forest within the Gifford Pinchot National Forest. No permit was
151 required for this work. Approval for research at San Joaquin Experimental Range was granted
152 by Dr. Angela White in May 2022 and for research at Santa Rita Experimental Range by
153 Dr. Mitch McClaran in May 2022. In 2024, permits were received for work at WOOD (Chase
154 Lake WMD; permit number 62515-24-020), KONZ (Konza Prairie Biological Station; permit
155 number 766), and UNDE (University of Notre Dame Environmental Research Center; permit
156 number UNDERC-2024-5).

157 **3.1.2 Soil collar placement**

158 Either one (2022 sampling campaign) or two (2024 sampling campaign) PVC soil collars (20.1
159 cm inside diameter) were installed in close proximity to the permanent NEON soil sensors at
160 each site (Figure 1). As instruments in the NEON soil sensor arrays can occasionally break
161 down or stop working, the specific soil plot where we made measurements was chosen at each
162 site in consultation with NEON staff to maximize likelihood of quality soil sensor measurements
163 during the duration of the IRGA measurements. The plot selected at each site (out of the 5 in
164 each replicate array at each site) are presented in the last column of Table 1. After installation,
165 collar(s) were left to equilibrate for approximately 24 hours prior to any measurements being
166 taken.

167 **3.1.3 Infrared gas analyzer measurements of soil CO₂ flux**

168 In 2022, we then made measurements of flux on an hourly interval for 8 hours each day.
169 Measurements were taken from roughly 8 am to 4 pm, with the time interval selected to
170 capture the majority of the diurnal gradient of soil temperature each day. These measurements
171 were made using a LI-6800 infrared gas analyzer instrument (LI-COR Environmental, Lincoln,
172 NE) fitted with a soil chamber attachment (attachment 6800-09). In 2024, we again used the
173 same LI-6800 instrument, but made half-hourly measurements over an approximately 8 hour
174 period. In addition, in 2024 we also installed a second collar and used a second instrument, an
175 LI-870 CO₂ IRGA, connected to an automated robotic chamber (LI-COR chamber 8200-104)
176 controlled by an LI-8250 multiplexer to make automated measurements. The multiplexer was
177 configured to take half-hourly measurements 24 hours a day for the duration of our sampling
178 bout at each site. Each instrument was paired with a soil temperature and moisture probe
179 (Stevens HydraProbe, Stevens Water, Portland, OR) that was used to make soil temperature
180 and moisture measurements concurrent with the CO₂ flux measurements. Chamber volumes

were set by measuring collar offsets at each site. System checks were conducted daily for the LI-6800 and weekly for the LI-8250. Instruments were factory calibrated before each field season.



Figure 1: Spatial layout of field sampling using a closed-dynamic chamber setup at a representative NEON site (KONZ). Left image shows collars (white arrows) and permanent soil sensor installation (black arrow) and right image shows the LI-6800 (foreground) and LI-8200-104 (background) instruments placed on the collars.

Table 1: Listing of NEON sites studied for field work and analysis. Site refers to NEON site codes: Santa Rita Experimental Range (SRER), San Joaquin Experimental Range (SJER), Wind River Experimental Forest (WREF), Chase Lake National Wildlife Refuge (WOOD), Konza Prairie Biological Station (KONZ), and the University of Notre Dame Environmental Research Center (UNDE). Location is reported in decimal degrees of latitude and longitude. Other abbreviations include Mean Annual Temperature (MAT); \bar{T}_S : average soil temperature during field measurements; \bar{SWC} : average soil water content during field measurements. Dates refer to field measurement dates for each site. Plot refers to the particular location in the soil sensor array (denoted as HOR by NEON) where field measurements were made.

Site	Location	Ecosystem	MAT	\bar{T}_S	MAP	\bar{SWC}	Dates	Plot
SRER	31.91068, -110.83549	Shrubland	19.3 °C	47.6 °C	346 mm	4.0%	May 29– June 1 2022	004
SJER	37.10878, -119.73228	Oak woodland	16.4 °C	41.7 °C	540 mm	1.2%	June 1–4 2022	005
WREF	45.82049, -121.95191	Evergreen forest	9.2 °C	15.3 °C	2225 mm	27.2%	June 7–9 2022	001

Table 1: Listing of NEON sites studied for field work and analysis. Site refers to NEON site codes: Santa Rita Experimental Range (SRER), San Joaquin Experimental Range (SJER), Wind River Experimental Forest (WREF), Chase Lake National Wildlife Refuge (WOOD), Konza Prairie Biological Station (KONZ), and the University of Notre Dame Environmental Research Center (UNDE). Location is reported in decimal degrees of latitude and longitude. Other abbreviations include Mean Annual Temperature (MAT); \bar{T}_S : average soil temperature during field measurements; \bar{SWC} : average soil water content during field measurements. Dates refer to field measurement dates for each site. Plot refers to the particular location in the soil sensor array (denoted as HOR by NEON) where field measurements were made.

Site	Location	Ecosystem	MAT	\bar{T}_S	MAP	\bar{SWC}	Dates	Plot
WOOD	47.1282, -99.241334	Restored prairie	4.9 °C	14.9 °C	495 mm	14.9%	June 3–9 2024	001
KONZ	39.100774, -96.563075	Tallgrass prairie	12.4 °C	23.4 °C	870 mm	23.4%	May 29– June 1 2024	001
UNDE	46.23391, -89.537254	Deciduous forest	4.3 °C	13.0 °C	802 mm	13.0%	May 22–25 2024	004

3.1.4 Post-collection processing of field data

We used LI-COR SoilFluxPro software (v 5.3.1) to assess the data after collection and to inform sampling parameters. We checked appropriateness of dead band and measurement durations using built-in evaluation tools. Based on this, the deadband period was set for 30-40 seconds, depending on the site, and the measurement duration was 180 seconds with a 30 second pre-purge and a 30 second post-purge at most sites, and a 90 second pre- and post-purge at sites with higher humidity due to recent precipitation events. We also assessed the R^2 of linear and exponential model fits to measured CO₂ to verify measurement quality.

3.2 neonSoilFlux R package

We developed an R package called `neonSoilFlux` (Zobitz et al., 2024) to compute half-hourly soil carbon fluxes and uncertainties from NEON data. The objective of the `neonSoilFlux`

195 package is a unified workflow (Figure 2) for soil data acquisition and analysis that supplements
196 the existing `neonUtilities` data acquisition R package (Lunch et al., 2025).

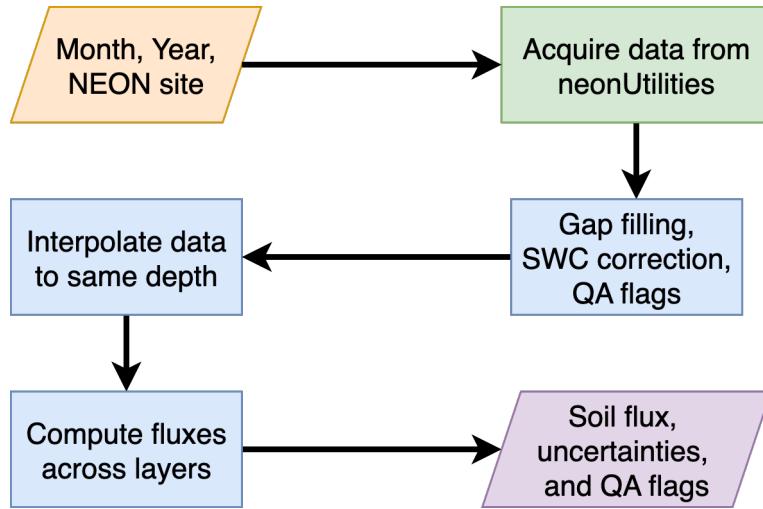


Figure 2: Diagram of `neonSoilFlux` R package. For a given month, year, and NEON site (orange parallelogram), the package acquires all relevant data to compute F_S using the `neonUtilities` R package (green rectangle). Data are gap-filled according to reported QA flags and adjusted for changes in soil water content (SWC) calibration coefficients, then interpolated to the same measurement depth before computing the soil flux, uncertainties, and final QA flags (blue rectangles). The package reports the associated soil flux, uncertainties, and quality assurance (QA) flags for the user (purple parallelogram).

197 At a given NEON site there are five replicate soil plots, each with measurements of soil
198 CO_2 concentration, soil temperature, and soil moisture at different depths (Figure 3). The
199 `neonSoilFlux` package acquires measured soil CO_2 concentration (NEON, 2024b), soil temper-
200 ature (NEON, 2024d), soil water content (NEON, 2024e), barometric pressure from the nearby
201 tower (NEON, 2024a), and soil properties (e.g. bulk density) (NEON, 2024c) from a range of
202 different NEON data products. The static soil properties were collected by NEON staff from a
203 nearby soil pit during initial site characterization and are assumed to be constant at each site.
204 A soil flux calculation is computed at each replicate soil plot.

205 The workflow to compute a value of F_S with `neonSoilFlux` consists of three primary steps,

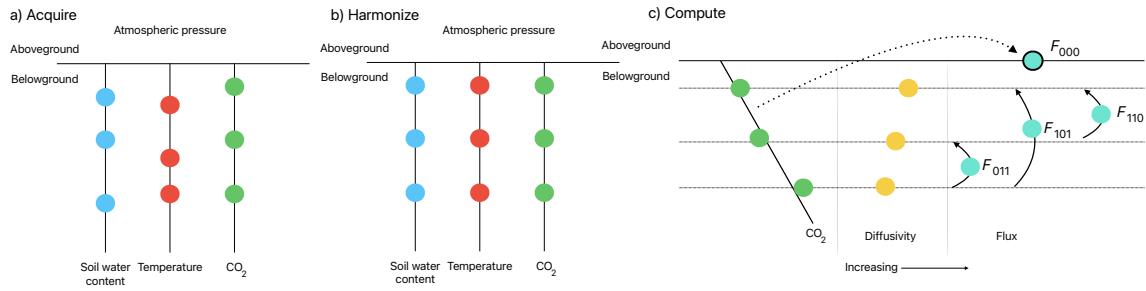


Figure 3: Model diagram of the data workflow for the `neonSoilFlux` R package. a) Acquire: Data are obtained for a given NEON location and horizontal sensor location, which includes soil water content, soil temperature, CO_2 concentration, and atmospheric pressure. All data are screened for quality assurance; if gap-filling of missing data occurs, it is flagged for the user. b) Harmonize: Any belowground data are then harmonized to the same depth as CO_2 concentrations using linear regression. c) Compute: The flux across a given depth is computed via Fick's law, denoted with F_{ijk} , where i , j , or k are either 0 or 1 denoting the layers the flux is computed across ($i =$ closest to surface, $k =$ deepest). F_{000} represents a flux estimate where the gradient dC/dz is the slope of a linear regression of CO_2 with depth.

206 illustrated in Figure 3. First, NEON data are acquired for a given site and month via the
 207 `neonUtilities` R package (yellow parallelogram and green rectangle in Figure 2 and Panel a
 208 in Figure 3). Acquired environmental data can be exported to a comma separated value file for
 209 additional analysis. Quality assurance (QA) flags are reported as an indicator variable. Since
 210 the calibration coefficients on the soil water content sensors have changed over time (NEON,
 211 2024e), raw sensor measurements were back-calculated and soil-specific calibrations were applied
 212 following Ayres et al. (2024) to generate a consistent time series at each measurement location.

213 The second step is harmonizing the data to compute soil fluxes across soil layers. This
 214 step consists of three different actions (blue rectangles in Figure 2 and Panel b in Figure 3).
 215 If a given observation by NEON is reported as not passing a quality assurance check, we
 216 applied a gap filling method to replace that measurement with its monthly mean at that same
 217 depth (Section 3.2.1). Belowground measurements of soil water and soil temperature are then
 218 interpolated to the same depth as soil CO_2 measurements. The diffusivity (Section 3.2.2) and

219 soil flux across different soil layers (Section 3.2.3) are then computed.

220 The third and final step is computing a surface soil flux through extrapolation to the surface
221 (purple parallelogram in Figure 2 and Panel c in Figure 3). Uncertainty on a soil flux
222 measurement is computed through quadrature. An aggregate quality assurance (QA) flag for
223 each environmental measurement is also reported, representing if any gap-filled measurements
224 were used in the computation of a soil flux. Within the soil flux-gradient method, several
225 different approaches can be used to derive a surface flux (Maier & Schack-Kirchner, 2014); the
226 `neonSoilFlux` package reports four different possible values for soil surface flux (Section 3.2.3)
227 for each of two different methods of diffusivity estimation, for a total of eight estimates of
228 flux.

229 3.2.1 Gap-filling routine

230 NEON reports QA flags as binary values for each measurement and half-hourly interval. For a
231 given half-hour, if any input variable (soil CO₂ concentration, soil temperature, or soil moisture)
232 at depth z is flagged, computation of F_S is not possible. To address this, flagged measurements
233 and their uncertainties were replaced with a bootstrapped monthly mean (\bar{m}) and monthly
234 standard deviation (\bar{s}) (Efron & Tibshirani, 1994).

235 For each month, depth z , and variable, we computed bootstrapped estimates of \bar{m} and \bar{s}
236 from the vectors of unflagged measurements (**m**), reported standard errors (σ), and the 95%
237 confidence interval (ϵ , or expanded uncertainty; Farrance & Frenkel (2012)). We also defined a
238 bias vector $\mathbf{b} = \sqrt{\epsilon^2 - \sigma^2}$, which quantifies the spread of uncertainty in a given period and is
239 incorporated into \bar{m} .

240 From these, 5000 bootstrap samples were generated for **m**, σ , and **b**. For each sample (m_k, b_k, σ_k),
241 we generated a vector **n** (length $N = 5000$) by drawing from a normal distribution with mean

242 $m_k + b_k$ and standard deviation σ_k . The sample mean and standard deviation were then
243 computed from \mathbf{n} . The resulting distributions of sample means and sample standard deviations
244 provided the bootstrapped monthly mean (\bar{m}) and standard error (\bar{s}) respectively.

245 This gap-filling procedure provides a consistent treatment across all data streams. However,
246 alternative approaches may be better suited for longer gaps (e.g., correlations with other NEON
247 measurement levels or soil plots) or for variable-specific conditions. We discuss the effect of
248 gap-filling on our results in Section 5.1.

249 3.2.2 Soil diffusivity

250 Soil diffusivity D_a at a given measurement depth is the product of the diffusivity in free air
251 $D_{a,0}$ ($\text{m}^2 \text{ s}^{-1}$) and the tortuosity ξ (no units) (Millington & Shearer, 1971).

252 We compute $D_{a,0}$ with Equation 1:

$$D_{a,0} = 0.0000147 \cdot \left(\frac{T_i + 273.15}{293.15} \right)^{1.75} \cdot \left(\frac{P}{101.3} \right) \quad (1)$$

253 where T_i is soil temperature ($^\circ\text{C}$) at depth i (NEON, 2024d) and P surface barometric pressure
254 (kPa) (NEON, 2024a).

255 Previous studies by Sallam et al. (1984) and Tang et al. (2003) demonstrated the sensitivity
256 of modeled F_S depending on the tortuosity model (ξ) used to compute diffusivity. At low
257 soil water content, the choice of tortuosity model can lead to order-of-magnitude differences
258 in D_a , which in turn affect modeled F_S . The `neonSoilFlux` package currently includes two
259 approaches to calculate ξ , representing the range of tortuosity behavior reported in Sallam et
260 al. (1984).

261 The first approach is the Millington-Quirk model (Millington & Shearer, 1971), in which
262 tortuosity depends on both porosity and soil water content:

$$\xi = \frac{(\phi - SWC_i)^{10/3}}{\phi^2} \quad (2)$$

263 In Equation 2, SWC is the soil water content at depth i (NEON, 2024e) and ϕ is the porosity,
264 which in turn is a function of soil physical properties (NEON, 2024c):

$$\phi = \left(1 - \frac{\rho_s}{\rho_m}\right) (1 - f_V) \quad (3)$$

265 In Equation 3, ρ_m is the particle density of mineral soil (2.65 g cm^{-3}), ρ_s the soil bulk density
266 (g cm^{-3}) excluding coarse fragments greater than 2 mm (NEON, 2024c), and f_V is a site-specific
267 value that accounts for the proportion of soil fragments between 2-20 mm. Soil fragments
268 greater than 20 mm were not estimated due to limitations in the amount of soil that can be
269 analyzed (NEON, 2024c). We assume that rock fragments contain no internal pores.

270 The Millington-Quirk model assumes ξ is modulated by the amount of fluid saturation in soil
271 pores (Millington & Shearer, 1971). In contrast, the Marshall model (Marshall, 1959) expresses
272 tortuosity as only a function of porosity ($\xi = \phi^{1.5}$), with ϕ defined from Equation 3. The
273 Marshall model is independent of soil water content and assumes tortuosity is only governed
274 by soil structure. The `neonSoilFlux` package allows users to choose the tortuosity model most
275 appropriate for site-specific conditions and research goals.

276 **3.2.3 Soil flux computation**

277 We applied Fick's law (Equation 4) to compute the soil flux F_{ij} ($\mu\text{mol m}^{-2} \text{s}^{-1}$) across two soil
278 depths i and j :

$$F_{ij} = -D_a \frac{dC}{dz} \quad (4)$$

279 where D_a is the diffusivity ($\text{m}^2 \text{s}^{-1}$) and $\frac{dC}{dz}$ is the gradient of CO_2 molar concentration (μmol
280 m^{-3} , so the gradient has units of $\mu\text{mol m}^{-3} \text{m}^{-1}$). The soil surface flux is theoretically defined
281 by applying Equation 4 to measurements collected at the soil surface and directly below the
282 surface. Measurements of soil temperature, soil water content, and soil CO_2 molar concentration
283 across the soil profile allow for application of Equation 4 across different soil depths. Each
284 site had three measurement layers, so we denote the flux as a three-digit subscript F_{ijk} with
285 indicator variables i , j , and k indicate if a given layer was used (written in order of increasing
286 depth), according to the following:

- 287 • F_{000} is a surface flux estimate using the intercept of the linear regression of D_a with
288 depth and the slope from the linear regression of CO_2 with depth (which represents $\frac{dC}{dz}$
289 in Fick's Law). Tang et al. (2003) used this approach to compute fluxes in an oak-grass
290 savannah.
- 291 • F_{110} is a flux estimate across the two shallowest measurement layers.
- 292 • F_{011} is a flux estimate across the two deepest measurement layers.
- 293 • F_{101} is a flux estimate across the shallowest and deepest measurement layers.

294 For F_{110} , F_{011} , and F_{101} , the diffusivity used in Fick's Law is always at the deeper measurement
295 layer. When used as a surface flux estimate we assume CO_2 remains constant above this flux
296 depth. Uncertainty in all F_{ijk} values was quantified using quadrature (Taylor, 2022). These

297 computed fluxes could provide the basis for additional soil flux estimates. For example, Tang et
298 al. (2005) estimated surface flux by linearly extrapolating F_{110} and F_{011} to the soil surface.

299 **3.3 Post processing evaluation**

300 Following collection of field measurements and calculation of the soil fluxes from `neonSoilFlux`
301 package, we compared measured F_S based on closed-dynamic chamber measurements with the
302 LI-COR instruments to a given soil flux calculation from `neonSoilFlux` for each site and flux
303 computation method and quantified the relationship statistically (R^2). Finally, for a half-hourly
304 interval we also computed a *post hoc* diffusivity (D_a) using the LI-COR flux along with the
305 CO_2 surface gradient reported by NEON using the measurement levels closest to the surface.

306 **4 Results**

307 **4.1 Concordance between modelled and measured soil CO_2 flux**

308 The sites we visited ranged substantially in both their annual average temperature and
309 precipitation as well as their biome type (Table 2). These differences also influenced the wide
310 range of observed flux rates across sites.

311 The timeseries of the measured fluxes from the LI-COR 6800 and 870/8250 were compared
312 to modeled soil fluxes from the `neonSoilFlux` R package (Figure 4). We also assessed year-
313 long estimated flux time series and compared those to field measurements made at each site
314 (Figure 5). Results are reported in local time. Where applicable, sites are displayed from left
315 to right by increasing soil temperature (Table 1). Positive values of the flux indicate that there
316 is a flux moving towards the surface. Overall, with the exception of SRER (discussed later) the
317 computed fluxes determined using a variety of plausible methods spanned the field-measured

Table 2: Summary of measured soil characteristics and flux results from field measurements across six NEON sites using a LI-COR 6800 (LI-870/8250 measurements omitted to enable direct comparability) via the closed-dynamic chamber method. Numeric values for soil CO₂ flux, soil temperature, and volumetric soil water content (VSWC) are the mean and standard deviation of field measurements at each site.

Site	Flux μmol m ⁻² s ⁻¹	Soil temp °C	VSWC cm ³ cm ⁻³	n
UNDE	2.55 ± 0.26	14.33 ± 0.77	0.33 ± 0.02	61
WOOD	3.02 ± 0.4	16.01 ± 1.54	0.28 ± 0.01	53
WREF	3.62 ± 0.3	15.34 ± 1.76	0.27 ± 0.06	21
KONZ	6.35 ± 0.97	27.28 ± 4.14	0.37 ± 0.01	44
SJER	0.94 ± 0.02	41.68 ± 11.22	0.01 ± 0.01	32
SRER	0.72 ± 0.09	47.64 ± 7.46	0.04 ± 0.01	32

³¹⁸ fluxes, but the specific flux-gradient method that best approximated field measurements varied
³¹⁹ by site.

³²⁰ We calculated a statistical relationship between the various estimates of soil flux computed by
³²¹ `neonSoilFlux` and the field-measured fluxes within daily interval periods. Statistics for these
³²² comparisons are reported in Figure 6, which also shows how these fall relative to a 1:1 line.

³²³ 4.2 Effects of method choice on diffusivity estimates

³²⁴ In four of six field sites, the *post hoc* D_a estimate fell roughly between the two diffusion
³²⁵ estimation methods; however this was less the case in the two driest sites, SJER and SRER
³²⁶ (Table 1), where the field estimate of diffusivity was either lower or higher than both of the
³²⁷ other methods (Figure 7).

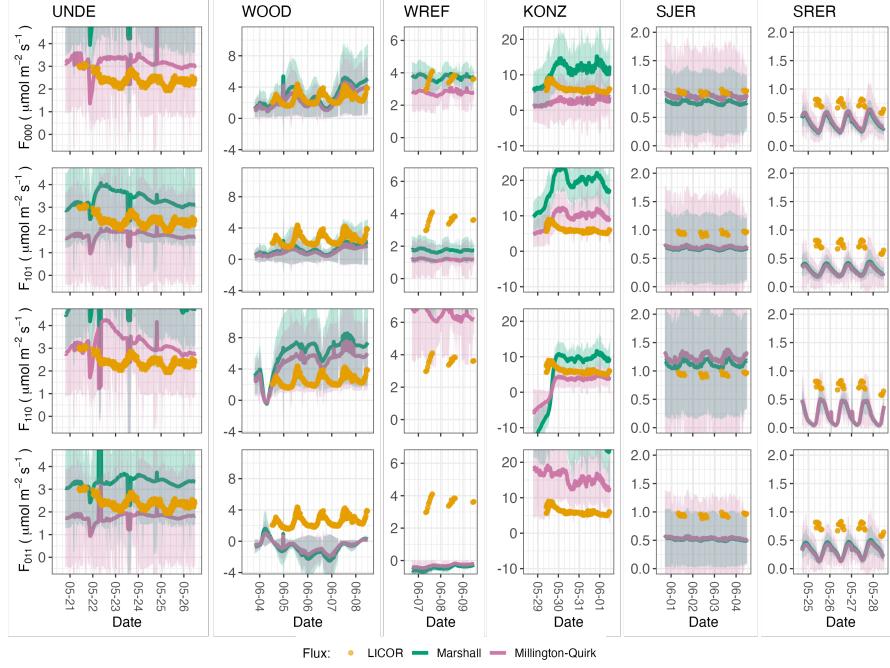


Figure 4: Timeseries of soil surface flux (F_S) from field-measured (yellow lines) and modeled soil fluxes (green or purple lines) by the `neonSoilFlux` R package. Fluxes from the `neonSoilFlux` R package are separated by the diffusivity model used (Millington-Quirk or Marshall, Section 3.2.2). Individual vertical axis labels in the first column represent the measurement levels where the flux-gradient approach is applied (Section 3.2.3). Ribbons for modeled soil fluxes represent ± 1 standard deviation. Results are reported in local time. WREF, SJER, and SRER were sampled in 2022, and UNDE, WOOD, and KONZ were sampled in 2024. Sites (columns) are arranged from left to right in terms of increasing mean annual temperature.

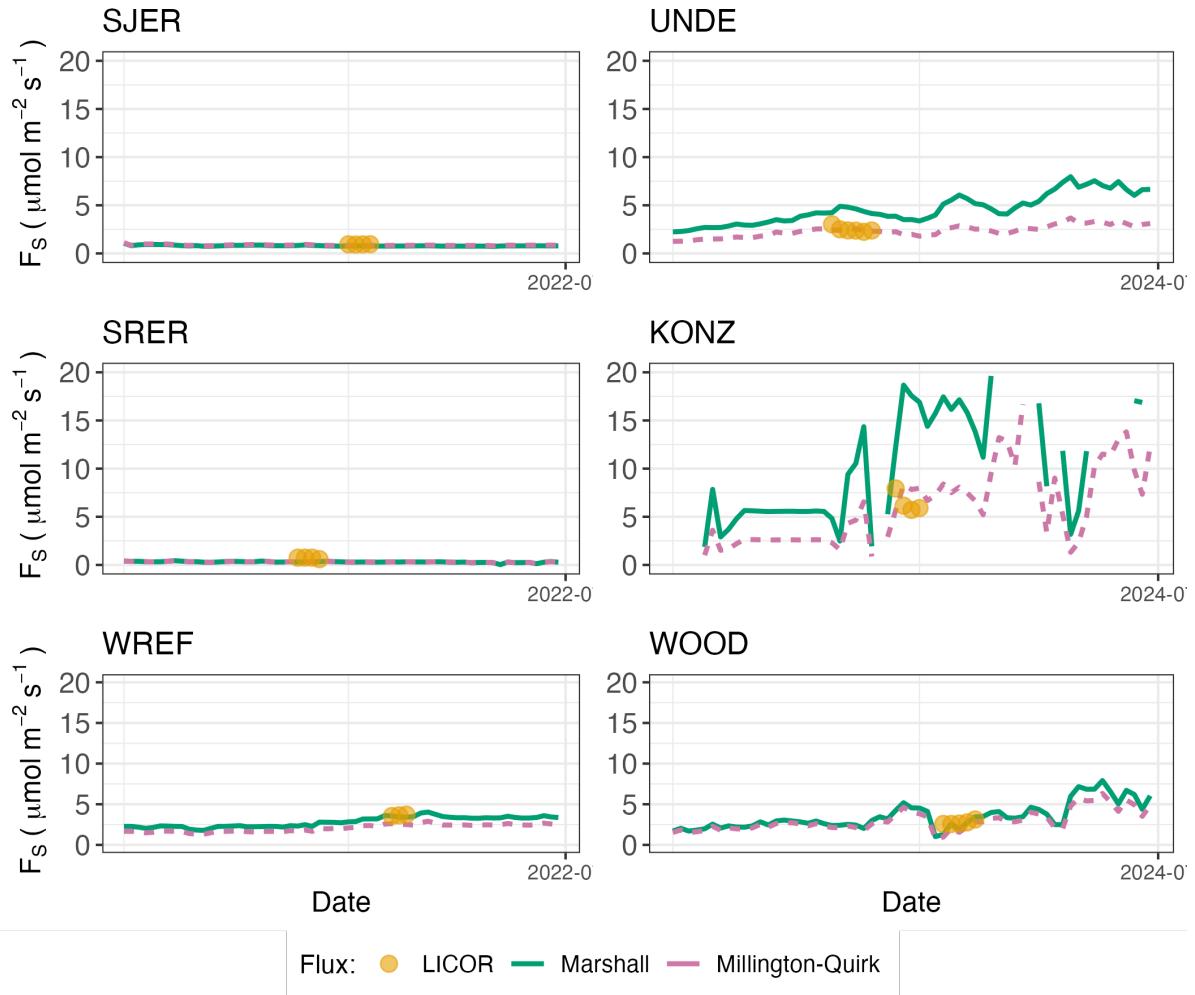


Figure 5: Timeseries of both daily-averaged field F_S (yellow circles) and daily ensemble averaged soil fluxes (average of F_{000} , F_{101} , F_{011} , F_{110} , Section 3.2.3) by the `neonSoilFlux` R package, separated by the diffusivity model used (green or purple lines, Millington-Quirk or Marshall, Section 3.2.2).

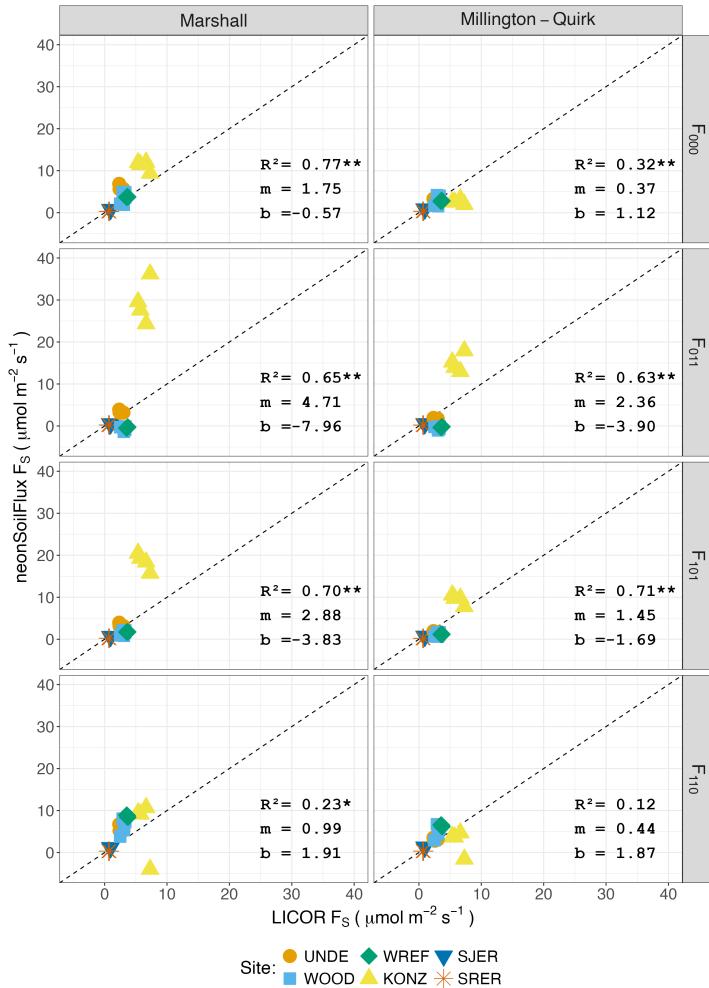


Figure 6: Statistical comparison between measured fluxes at each NEON site with fluxes reported by `neonSoilFlux` with the different flux calculation approaches and diffusivity calculations applied. Points are daily averages and LICOR F_S values are from the 6800 instrument only, for consistency. The dotted line represents a 1:1 relationship, and the reported R^2 quantifies the relationship between field-measured and `neonSoilFlux` estimated fluxes. * = significance at the 5% level, ** = significance at the 1% level. The slope (m) and intercept (b) of the linear regression between measured and modeled fluxes are also reported. The low-value outlier from KONZ in the F_{110} Marshall plot is an example of the effect of inverted CO₂ gradients causing an estimated flux to be negative, bringing down the daily mean, which later resolved as the soils dried back out. Intervals where calculated fluxes were < 0 ($n = 12$) were excluded from the plot; 3 each from WREF and WOOD when using (F_{011} approach).

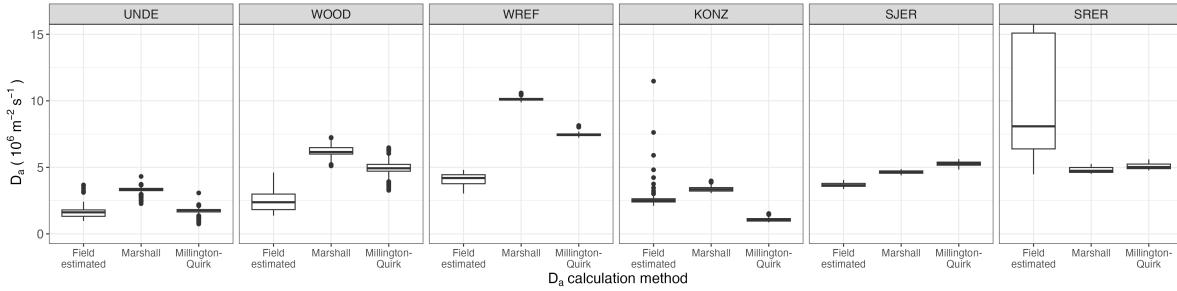


Figure 7: Distribution of diffusivity (D_a) at each study site. Values of D_a were provided by the `neonSoilFlux` package, computed from the Millington-Quirk or Marshall models (Section 3.2.2). A post-hoc estimate of diffusivity (labeled as “Field estimated”) was computed through the field measured flux (Figure 4), divided by the CO₂ gradient from the measurement levels closest to the soil surface, as reported by NEON. We only used F_S measured by the LICOR 6800 at all sites to standardize comparisons. Some outliers (n = 2 from the field estimated values at KONZ and n = 9 from field estimated values at SRER) are excluded from the plot to allow better comparative visualization across sites.

328 5 Discussion

329 This study presents a unified data science workflow to efficiently process automated measurements
 330 of belowground soil CO₂ concentrations, soil water content, and soil temperature to
 331 infer estimates of soil surface CO₂ effluxes through application of Fick’s Law (Equation 4).
 332 Our core goals in this study were: (1) to generate estimates of soil flux from continuous soil
 333 sensor data at terrestrial NEON sites using the flux-gradient method and then (2) to compare
 334 those estimates to field-measured fluxes based on the closed chamber approach at six NEON
 335 focal sites. We discuss our progress toward these core goals through (1) an overall evaluation
 336 of the flux-gradient approach (and uncertainty calculation) and (2) site-specific evaluation of
 337 differences in estimated vs measured fluxes.

338 **5.1 General evaluation of flux-gradient approach**

339 Key assumptions of the flux-gradient approach are that CO₂ concentrations increase throughout
340 the soil profile such that the highest concentrations are observed in the deepest layers. Addition-
341 ally, field flux measurements should correlate with F_{000} because they represent surface fluxes.
342 Periods where this gradient condition are not met generally are connected to processes that occur
343 during soil wetting events, where more shallow soil layers produce higher concentrations of CO₂
344 due to microbial respiration pulses following rewetting. This effect is likely to be largest at sites
345 with rich organic soils (e.g. KONZ). Based on this reasoning, in these types of situations we would
346 *a priori* expect F_{011} (deepest layers) $\leq F_{101} \leq F_{110}$ (shallow layers) $\leq F_{000}$ (all layers) be-
347 cause the previous flux estimates rely primarily on CO₂ concentrations at deeper depths, and
348 could miss high concentrations of CO₂ produced in shallower layers.

349 When modeling soil respiration, typically a non-linear response function that also considers soil
350 type is used (Bouma & Bryla, 2000; Yan et al., 2016, 2018). For the `neonSoilFlux` package,
351 soil type is connected to the measurement of bulk density, which was characterized at each
352 NEON site. This bulk density estimate is based on replicate samples collected from the site
353 megapit at a subset of soil horizons, with an estimated uncertainty of $\pm 5\%$ (NEON, 2024c).
354 Coarse fragment estimates also have very large uncertainties, but because the volume fraction
355 tends to be low in surface soils it is unlikely to contribute much additional flux uncertainty.

356 Our results suggest that the most important way to improve reliability of the flux estimate is
357 to reduce the usage of gap-filled data. The current approach to gap filling in `neonSoilFlux`
358 uses monthly mean data to gap fill—this approach decreases the ability of the estimate to be
359 responsive to short-term pulses that occur with rapid weather shifts. Four sites (KONZ, SRER,
360 WREF, and UNDE) had more than 75% of half-hourly periods with no-gap filled measurements
361 (Figure S1, Supplementary Information). Two sites (SJER and WOOD) had more than 75% of
362 half-hourly intervals with just one gap-filled measurement. The large uncertainty evident in

363 Figure 4 for estimates from WOOD and SJER are thus due in part to the gap-filling used in
364 these sites (Figure S1). While we did not need to use gap-filled measurements to compute the
365 flux at WREF, field data collection occurred following a severe rainstorm, with soils at the
366 beginning of the sampling week near their water holding capacity. In general, we recommend
367 that whenever possible, knowledge of local field conditions should influence analysis decisions
368 in addition to any QA filtering protocols in the `neonSoilFlux` package.

369 We recognize that this gap-filling approach may lead to gap-filled values that are quite different
370 from the actual values, such as an underestimate of soil moisture following rain events. Further
371 extensions of the gap filling method could use more sophisticated gap-filling routines, similar to
372 what is used for net ecosystem carbon exchange (Falge et al., 2001; Liu et al., 2023; Mariethoz
373 et al., 2015; Moffat et al., 2007; Zhang et al., 2023). Additionally, since the deepest temperature
374 and soil moisture sensors are located below the deepest CO₂ sensors at NEON sites, it is
375 possible that excluding these deeper layers from consideration prior to analysis would lead to a
376 reduced need for gap filling. Future iterations of the `neonSoilFlux` package may incorporate
377 this as an option. The current gap-filling routine provides a consistent approach that can be
378 applied to each data stream, but further work may explore alternative gap-filling approaches.

379 **5.2 Evaluation of flux-gradient approach at each site**

380 Derived results from the `neonSoilFlux` package have patterns that are broadly consistent with
381 those directly measured in the field (Figure 4 and Figure 5), even though statistical comparisons
382 between the field-measured and `neonSoilFlux` values were quite variable (e.g. R^2 ranging
383 from 0.04 to 0.81; Figure 6). One advantage of the `neonSoilFlux` package is its ability to
384 calculate fluxes across different soil depths (Figure 3), which allows for additional site-specific
385 customization. We believe the package can provide a useful baseline estimate of soil fluxes that
386 can always be complemented through additional field measurements.

387 The six locations studied provide a range of case studies that suggest different considerations
388 may apply to different sites when applying the flux-gradient method. For example, the Santa
389 Rita Experimental Range (SRER) is a desert site characterized by sandy soil, which also was
390 the location of the highest field soil temperatures that we observed (Table 2). At SRER the
391 flux across the top two layers (F_{110}) produced a pattern of soil flux most consistent with the
392 observed field data. The remaining methods F_{101} , F_{011} , or F_{000} are derived from information
393 taken from the deepest layer, which seems to have been decoupled from the surface layers both
394 in terms of temperature and CO₂ concentration. This may be a general circumstance where
395 there are large diurnal temperature extremes that rapidly change during the course of a day
396 and overnight, leading to lags in the timing of when temperature increases propagate down to
397 deeper soil layers.

398 Immediately prior to our visit to Konza Prairie (KONZ), that site that experienced a significant
399 rain event that led to wet soils that gradually dried out over the course of our time there.
400 This pulse of precipitation increased the soil CO₂ concentration at the top layer above the
401 concentrations in lower layers, leading to negative estimated flux values at the start of the field
402 sampling period. In this case it was only when the soil began to return to a baseline level that
403 the assumptions of the flux-gradient method were again met.

404 Both of the previous cases also provide context for the variable statistical comparisons between
405 field-measured soil fluxes and `neonSoilFlux` outputs (Figure 6). When considering systematic
406 deployment of this method across a measurement network, there are a number of independent
407 challenges that require careful consideration. There are clear tradeoffs between (1) accuracy of
408 modeled fluxes (defined here as closeness to field-measured F_S and the uncertainty reduction
409 factor ϵ), (2) precision (which could be defined by the signal to noise ratio), and (3) the
410 choice of the diffusivity model (Section 3.2.2) or flux computation method (Section 3.2.3). A
411 sensitivity analysis (Figure S2, Supplemental Information) found that flux output uncertainty

412 was dominated by measurement uncertainty (T_S , P , SWC , or CO_2) rather than by the diffusivity
413 method used to compute soil flux. Notably, the F_{110} method was least sensitive to measurement
414 uncertainty likely because it best aligns with the surface chamber measurement assumptions.

415 Finally, comparing the effects of different diffusivity estimation methods on the match between
416 modeled and measured fluxes (Figure 5) highlights the sensitivity of F_{ijk} to diffusivity. The
417 comparison between diffusivity estimates compared to field estimated diffusivity (Figure 7)
418 demonstrates that site parameters can dictate which measure of diffusivity is most likely to be
419 accurate in a given environmental context. Site-specific differences are largely a reflection of
420 differences in soil moisture across the sites (Table 1), as not all diffusivity estimation methods
421 incorporate soil moisture equivalently. While we here have compares two approaches to calculate
422 diffusivity (the Millington-Quirk and Marshall models), it may be valuable to evaluate other
423 diffusivity models (e.g. the Moldrup model; Moldrup et al. (1999)) as well. Ultimately the
424 choice of a particular diffusivity model could be determined based on knowledge of site-specific
425 evaluations or a set of these models could be used to generate a model ensemble average as a
426 means to trade precision for a more general approach.

427 **5.3 Recommendations for future method development**

428 The `neonSoilFlux` package provides several approaches to estimate soil flux using the gradient
429 method. We believe these approaches enable the software to be used across a range of site-
430 specific assumptions (Maier & Schack-Kirchner, 2014). We note, however, that this choice
431 can have a determinative approach on the calculated values. Ensemble averaging approaches
432 (Elshall et al., 2018; Raftery et al., 2005) may be one way to address this problem if the goal is
433 to calculate fluxes using the same method at a diverse range of different sites. Two other ideas
434 would be to apply machine learning algorithms (e.g. random forest) to generate a single flux
435 estimate across diverse sites, or using co-located estimates of net ecosystem carbon exchange

⁴³⁶ from eddy-flux towers to further constrain results or to assess soil flux results for plausibility
⁴³⁷ (Phillips et al., 2017).

⁴³⁸ These challenges notwithstanding, the method used here and made available in the
⁴³⁹ `neonSoilFlux` R package has the potential to produce nearly continuous estimates of flux
⁴⁴⁰ across all terrestrial NEON sites. These estimates are a significant improvement on available
⁴⁴¹ approaches to constrain the portion of ecosystem respiration attributable to the soil. This, in
⁴⁴² turn, also aids in our ability to understand the soil contribution to the net ecosystem flux
⁴⁴³ measured at these sites using the co-located eddy flux towers.

⁴⁴⁴ **6 Conclusions**

⁴⁴⁵ We used the R package `neonSoilFlux` to estimate soil CO₂ fluxes with the flux-gradient method
⁴⁴⁶ using data from buried soil sensors at NEON terrestrial sites. We compared the predicted
⁴⁴⁷ fluxes to those measured directly using a field-based closed chamber approach. Soil fluxes
⁴⁴⁸ from `neonSoilFlux` were broadly effective at producing estimates of flux comparable to those
⁴⁴⁹ measured in the field using a chamber-based technique. However `neonSoilFlux` outputs are
⁴⁵⁰ quite sensitive to a number of issues, including: missing data (and thus gap-filling of input
⁴⁵¹ measurement datasets), the selection of soil depths used to best calculate the gradient (which
⁴⁵² may vary between sites), and finally the choice of method used for estimating soil diffusivity.
⁴⁵³ The flexibility of the `neonSoilFlux` package allows the user to evaluate each of these issues
⁴⁵⁴ with site-specific knowledge and contexts. Future refinements and subsequent validation of
⁴⁵⁵ `neonSoilFlux` outputs will feed forward into evaluating soil carbon fluxes broader spatial scales
⁴⁵⁶ to enhance understanding of the ways in which soils across diverse ecosystems are responding
⁴⁵⁷ to a changing climate.

458 **Sources Cited**

- 459 Ayres, E., Reichle, R. H., Colliander, A., Cosh, M. H., & Smith, L. (2024). Validation of
460 Remotely Sensed and Modeled Soil Moisture at Forested and Unforested NEON Sites.
461 *IEEE Journal of Selected Topics in Applied Earth Observations and Remote Sensing*, 17,
462 14248–14264. <https://doi.org/10.1109/JSTARS.2024.3430928>
- 463 Baldocchi, D. (2014). Measuring fluxes of trace gases and energy between ecosystems and the
464 atmosphere - the state and future of the eddy covariance method. *Global Change Biology*,
465 20(12), 3600–3609. <https://doi.org/10.1111/gcb.12649>
- 466 Berenbaum, M. R., Carpenter, S. R., Hampton, S. E., Running, S. W., & Stanzione, D. C.
467 (2015). *Report from the NSF BIO Advisory Committee Subcommittee on NEON Scope
468 Impacts*.
- 469 Bond-Lamberty, B. (2018). New Techniques and Data for Understanding the Global Soil
470 Respiration Flux. *Earth's Future*, 6(9), 1176–1180. <https://doi.org/10.1029/2018EF000866>
- 471 Bond-Lamberty, B., Ballantyne, A., Berryman, E., Fluet-Chouinard, E., Jian, J., Morris, K.
472 A., Rey, A., & Vargas, R. (2024). Twenty Years of Progress, Challenges, and Opportunities
473 in Measuring and Understanding Soil Respiration. *Journal of Geophysical Research: Biogeosciences*, 129(2), e2023JG007637. <https://doi.org/10.1029/2023JG007637>
- 475 Bond-Lamberty, B., Christianson, D. S., Malhotra, A., Pennington, S. C., Sihi, D., AghaK-
476 ouchak, A., Anjileli, H., Altaf Arain, M., Armesto, J. J., Ashraf, S., Ataka, M., Baldocchi,
477 D., Andrew Black, T., Buchmann, N., Carbone, M. S., Chang, S.-C., Crill, P., Curtis, P.
478 S., Davidson, E. A., ... Zou, J. (2020). COSORE: A community database for continuous
479 soil respiration and other soil-atmosphere greenhouse gas flux data. *Global Change Biology*,
480 26(12), 7268–7283. <https://doi.org/10.1111/gcb.15353>
- 481 Bond-Lamberty, B., & Thomson, A. (2010). A global database of soil respiration data.
482 *Biogeosciences*, 7(6), 1915–1926. <https://doi.org/10.5194/bg-7-1915-2010>

- 483 Bond-Lamberty, B., Wang, C., & Gower, S. T. (2004). A global relationship between the
484 heterotrophic and autotrophic components of soil respiration? *Global Change Biology*,
485 10(10), 1756–1766. <https://doi.org/10.1111/j.1365-2486.2004.00816.x>
- 486 Bouma, T. J., & Bryla, D. R. (2000). On the assessment of root and soil respiration for soils of
487 different textures: Interactions with soil moisture contents and soil CO₂ concentrations.
488 *Plant and Soil*, 227(1), 215–221. <https://doi.org/10.1023/A:1026502414977>
- 489 Chen, H., & Tian, H.-Q. (2005). Does a General Temperature-Dependent Q10 Model of Soil
490 Respiration Exist at Biome and Global Scale? *Journal of Integrative Plant Biology*, 47(11),
491 1288–1302. <https://doi.org/10.1111/j.1744-7909.2005.00211.x>
- 492 Davidson, E. A., Janssens, I. A., & Luo, Y. (2006). On the variability of respiration in
493 terrestrial ecosystems: Moving beyond Q10. *Global Change Biology*, 12, 154–164. <https://doi.org/10.1111/j.1365-2486.2005.01065.x>
- 494 Desai, A. R., Murphy, B. A., Wiesner, S., Thom, J., Butterworth, B. J., Koupaei-Abyazani, N.,
495 Muttaqin, A., Paleri, S., Talib, A., Turner, J., Mineau, J., Merrelli, A., Stoy, P., & Davis,
496 K. (2022). Drivers of Decadal Carbon Fluxes Across Temperate Ecosystems. *Journal of
497 Geophysical Research: Biogeosciences*, 127(12), e2022JG007014. <https://doi.org/10.1029/2022JG007014>
- 498 499
- 500 Efron, B., & Tibshirani, R. J. (1994). *An Introduction to the Bootstrap*. Chapman and
501 Hall/CRC. <https://doi.org/10.1201/9780429246593>
- 502 Elshall, A. S., Ye, M., Pei, Y., Zhang, F., Niu, G.-Y., & Barron-Gafford, G. A. (2018). Relative
503 model score: A scoring rule for evaluating ensemble simulations with application to microbial
504 soil respiration modeling. *Stochastic Environmental Research and Risk Assessment*, 32(10),
505 2809–2819. <https://doi.org/10.1007/s00477-018-1592-3>
- 506 Falge, E., Baldocchi, D., Olson, R., Anthoni, P., Aubinet, M., Bernhofer, C., Burba, G.,
507 Ceulemans, R., Clement, R., Dolman, H., Granier, A., Gross, P., Grünwald, T., Hollinger,
508 D., Jensen, N.-O., Katul, G., Keronen, P., Kowalski, A., Lai, C. T., ... Wofsy, S. (2001).

- 509 Gap filling strategies for defensible annual sums of net ecosystem exchange. *Agricultural*
510 and *Forest Meteorology*, 107(1), 43–69. [https://doi.org/10.1016/S0168-1923\(00\)00225-2](https://doi.org/10.1016/S0168-1923(00)00225-2)
- 511 Farrance, I., & Frenkel, R. (2012). *Uncertainty of Measurement: A Review of the Rules*
512 for Calculating Uncertainty Components through Functional Relationships. *The Clinical*
513 *Biochemist Reviews*, 33(2), 49–75.
- 514 Friedlingstein, P., O’Sullivan, M., Jones, M. W., Andrew, R. M., Hauck, J., Landschützer,
515 P., Le Quéré, C., Li, H., Luijckx, I. T., Olsen, A., Peters, G. P., Peters, W., Pongratz,
516 J., Schwingshackl, C., Sitch, S., Canadell, J. G., Ciais, P., Jackson, R. B., Alin, S. R., ...
517 Zeng, J. (2025). Global Carbon Budget 2024. *Earth System Science Data*, 17(3), 965–1039.
518 <https://doi.org/10.5194/essd-17-965-2025>
- 519 Hamdi, S., Moyano, F., Sall, S., Bernoux, M., & Chevallier, T. (2013). Synthesis analysis
520 of the temperature sensitivity of soil respiration from laboratory studies in relation to
521 incubation methods and soil conditions. *Soil Biology and Biochemistry*, 58, 115–126.
522 <https://doi.org/10.1016/j.soilbio.2012.11.012>
- 523 Jackson, R. B., Lajtha, K., Crow, S. E., Hugelius, G., Kramer, M. G., & Piñeiro, G. (2017).
524 The Ecology of Soil Carbon: Pools, Vulnerabilities, and Biotic and Abiotic Controls.
525 *Annual Review of Ecology, Evolution and Systematics*, 48(Volume 48, 2017), 419–445.
526 <https://doi.org/10.1146/annurev-ecolsys-112414-054234>
- 527 Jian, J., Bailey, V., Dorheim, K., Konings, A. G., Hao, D., Shiklomanov, A. N., Snyder,
528 A., Steele, M., Teramoto, M., Vargas, R., & Bond-Lamberty, B. (2022). Historically
529 inconsistent productivity and respiration fluxes in the global terrestrial carbon cycle. *Nature*
530 *Communications*, 13(1), 1733. <https://doi.org/10.1038/s41467-022-29391-5>
- 531 Jian, J., Vargas, R., Anderson-Teixeira, K., Stell, E., Herrmann, V., Horn, M., Kholod, N.,
532 Manzon, J., Marchesi, R., Paredes, D., & Bond-Lamberty, B. (2021). A restructured and
533 updated global soil respiration database (SRDB-V5). *Earth System Science Data*, 13(2),
534 255–267. <https://doi.org/10.5194/essd-13-255-2021>

- 535 Jiang, J., Feng, L., Hu, J., Liu, H., Zhu, C., Chen, B., & Chen, T. (2024). Global soil
536 respiration predictions with associated uncertainties from different spatio-temporal data
537 subsets. *Ecological Informatics*, 82, 102777. <https://doi.org/10.1016/j.ecoinf.2024.102777>
- 538 Jobbágy, E. G., & Jackson, R. B. (2000). The Vertical Distribution of Soil Organic Carbon
539 and its Relation to Climate and Vegetation. *Ecological Applications*, 10(2), 423–436.
540 [https://doi.org/10.1890/1051-0761\(2000\)010%5B0423:TVDOSO%5D2.0.CO;2](https://doi.org/10.1890/1051-0761(2000)010%5B0423:TVDOSO%5D2.0.CO;2)
- 541 Jurasinski, G., Koebisch, F., Guenther, A., & Beetz, S. (2022). *Flux: Flux Rate Calculation
542 from Dynamic Closed Chamber Measurements*.
- 543 Liu, K., Li, X., Wang, S., & Zhang, H. (2023). A robust gap-filling approach for European
544 Space Agency Climate Change Initiative (ESA CCI) soil moisture integrating satellite
545 observations, model-driven knowledge, and spatiotemporal machine learning. *Hydrology
546 and Earth System Sciences*, 27(2), 577–598. <https://doi.org/10.5194/hess-27-577-2023>
- 547 Lunch, C., Laney, C., Mietkiewicz, N., Sokol, E., Cawley, K., & Network), N. (National. E. O.
548 (2025). *neonUtilities: Utilities for Working with NEON Data*.
- 549 Luo, Y., Ogle, K., Tucker, C., Fei, S., Gao, C., LaDeau, S., Clark, J. S., & Schimel, D. S. (2011).
550 Ecological forecasting and data assimilation in a data-rich era. *Ecological Applications*,
551 21(5), 1429–1442. <https://doi.org/10.1890/09-1275.1>
- 552 Maier, M., & Schack-Kirchner, H. (2014). Using the gradient method to determine soil gas
553 flux: A review. *Agricultural and Forest Meteorology*, 192–193, 78–95. <https://doi.org/10.1016/j.agrformet.2014.03.006>
- 555 Mariethoz, G., Linde, N., Jougnot, D., & Rezaee, H. (2015). Feature-preserving interpolation
556 and filtering of environmental time series. *Environmental Modelling & Software*, 72, 71–76.
557 <https://doi.org/10.1016/j.envsoft.2015.07.001>
- 558 Marshall, T. J. (1959). The Diffusion of Gases Through Porous Media. *Journal of Soil Science*,
559 10(1), 79–82. <https://doi.org/10.1111/j.1365-2389.1959.tb00667.x>
- 560 Millington, R. J., & Shearer, R. C. (1971). Diffusion in aggregated porous media. *Soil Science*,

- 561 111(6), 372–378.
- 562 Moffat, A. M., Papale, D., Reichstein, M., Hollinger, D. Y., Richardson, A. D., Barr, A. G.,
563 Beckstein, C., Braswell, B. H., Churkina, G., Desai, A. R., Falge, E., Gove, J. H., Heimann,
564 M., Hui, D., Jarvis, A. J., Kattge, J., Noormets, A., & Stauch, V. J. (2007). Comprehensive
565 comparison of gap-filling techniques for eddy covariance net carbon fluxes. *Agricultural and*
566 *Forest Meteorology*, 147(3), 209–232. <https://doi.org/10.1016/j.agrformet.2007.08.011>
- 567 Moldrup, P., Olesen, T., Yamaguchi, T., Schjønning, P., & Rolston, D. E. (1999). Modeling
568 diffusion and reaction in soils: 9. The Buckingham-Burdine-Campbell equation for gas
569 diffusivity in undisturbed soil. *Soil Science*, 164(2), 75.
- 570 NEON. (2024a). *Barometric pressure (DP1.00004.001)*. National Ecological Observatory
571 Network (NEON). <https://doi.org/10.48443/RT4V-KZ04>
- 572 NEON. (2024b). *Soil CO₂ concentration (DP1.00095.001)*. National Ecological Observatory
573 Network (NEON). <https://doi.org/10.48443/E7GR-6G94>
- 574 NEON. (2024c). *Soil physical and chemical properties, Megapit (DP1.00096.001)*. National
575 Ecological Observatory Network (NEON). <https://doi.org/10.48443/S6ND-Q840>
- 576 NEON. (2024d). *Soil temperature (DP1.00041.001)*. National Ecological Observatory Network
577 (NEON). <https://doi.org/10.48443/Q24X-PW21>
- 578 NEON. (2024e). *Soil water content and water salinity (DP1.00094.001)*. National Ecological
579 Observatory Network (NEON). <https://doi.org/10.48443/A8VY-Y813>
- 580 Norman, J. M., Kucharik, C. J., Gower, S. T., Baldocchi, D. D., Crill, P. M., Rayment,
581 M., Savage, K., & Striegl, R. G. (1997). A comparison of six methods for measuring
582 soil-surface carbon dioxide fluxes. *Journal of Geophysical Research: Atmospheres*, 102(D24),
583 28771–28777. <https://doi.org/10.1029/97JD01440>
- 584 Pedersen, A. R. (2024). *HMR: Flux Estimation with Static Chamber Data*.
- 585 Phillips, C. L., Bond-Lamberty, B., Desai, A. R., Lavoie, M., Risk, D., Tang, J., Todd-Brown,
586 K., & Vargas, R. (2017). The value of soil respiration measurements for interpreting and

- 587 modeling terrestrial carbon cycling. *Plant and Soil*, 413(1), 1–25. <https://doi.org/10.1007/s11104-016-3084-x>
- 588
- 589 Raftery, A. E., Gneiting, T., Balabdaoui, F., & Polakowski, M. (2005). *Using Bayesian Model*
590 *Averaging to Calibrate Forecast Ensembles*. <https://doi.org/10.1175/MWR2906.1>
- 591 Rheault, K., Christiansen, J. R., & Larsen, K. S. (2024). goFlux: A user-friendly way to
592 calculate GHG fluxes yourself, regardless of user experience. *Journal of Open Source*
593 *Software*, 9(96), 6393. <https://doi.org/10.21105/joss.06393>
- 594 Sallam, A., Jury, W. A., & Letey, J. (1984). Measurement of Gas Diffusion Coefficient under
595 Relatively Low Air-filled Porosity. *Soil Science Society of America Journal*, 48(1), 3–6.
596 <https://doi.org/10.2136/sssaj1984.03615995004800010001x>
- 597 Shao, J., Zhou, X., Luo, Y., Li, B., Aurela, M., Billesbach, D., Blanken, P. D., Bracho, R.,
598 Chen, J., Fischer, M., Fu, Y., Gu, L., Han, S., He, Y., Kolb, T., Li, Y., Nagy, Z., Niu, S.,
599 Oechel, W. C., ... Zhang, J. (2015). Biotic and climatic controls on interannual variability
600 in carbon fluxes across terrestrial ecosystems. *Agricultural and Forest Meteorology*, 205,
601 11–22. <https://doi.org/10.1016/j.agrformet.2015.02.007>
- 602 Shao, P., Zeng, X., Moore, D. J. P., & Zeng, X. (2013). Soil microbial respiration from
603 observations and Earth System Models. *Environmental Research Letters*, 8(3), 034034.
604 <https://doi.org/10.1088/1748-9326/8/3/034034>
- 605 Sihi, D., Gerber, S., Inglett, P. W., & Inglett, K. S. (2016). Comparing models of microbial–
606 substrate interactions and their response to warming. *Biogeosciences*, 13(6), 1733–1752.
607 <https://doi.org/10.5194/bg-13-1733-2016>
- 608 Tang, J., Baldocchi, D. D., Qi, Y., & Xu, L. (2003). Assessing soil CO₂ efflux using continuous
609 measurements of CO₂ profiles in soils with small solid-state sensors. *Agricultural and Forest*
610 *Meteorology*, 118(3), 207–220. [https://doi.org/10.1016/S0168-1923\(03\)00112-6](https://doi.org/10.1016/S0168-1923(03)00112-6)
- 611 Tang, J., Misson, L., Gershenson, A., Cheng, W., & Goldstein, A. H. (2005). Continuous
612 measurements of soil respiration with and without roots in a ponderosa pine plantation

- 613 in the Sierra Nevada Mountains. *Agricultural and Forest Meteorology*, 132(3), 212–227.
- 614 <https://doi.org/10.1016/j.agrformet.2005.07.011>
- 615 Taylor, J. R. (2022). *An Introduction to Error Analysis: The Study of Uncertainties in Physical*
616 *Measurements, Third Edition* (3rd ed.). University Science Press.
- 617 Wilson, S. J., Bond-Lamberty, B., Noyce, G., Bittencourt Peixoto, R., & Megonigal, J. P.
618 (2024). Fluxfinder: An R Package for Reproducible Calculation and Initial Processing
619 of Greenhouse Gas Fluxes From Static Chamber Measurements. *Journal of Geophysical*
620 *Research: Biogeosciences*, 129(11), e2024JG008208. <https://doi.org/10.1029/2024JG008208>
- 621 Yan, Z., Bond-Lamberty, B., Todd-Brown, K. E., Bailey, V. L., Li, S., Liu, C., & Liu, C. (2018).
622 A moisture function of soil heterotrophic respiration that incorporates microscale processes.
623 *Nature Communications*, 9(1), 2562. <https://doi.org/10.1038/s41467-018-04971-6>
- 624 Yan, Z., Liu, C., Todd-Brown, K. E., Liu, Y., Bond-Lamberty, B., & Bailey, V. L. (2016).
625 Pore-scale investigation on the response of heterotrophic respiration to moisture conditions
626 in heterogeneous soils. *Biogeochemistry*, 131(1), 121–134. <https://doi.org/10.1007/s10533-016-0270-0>
- 627 Zhang, R., Kim, S., Kim, H., Fang, B., Sharma, A., & Lakshmi, V. (2023). Temporal Gap-Filling
628 of 12-Hourly SMAP Soil Moisture Over the CONUS Using Water Balance Budgeting. *Water*
629 *Resources Research*, 59(12), e2023WR034457. <https://doi.org/10.1029/2023WR034457>
- 630 Zhao, J. (2019). FluxCalR: A R package for calculating CO₂ and CH₄ fluxes from static
631 chambers. *Journal of Open Source Software*, 4(43), 1751. <https://doi.org/10.21105/joss.01751>
- 632 Zobitz, J., Ayres, E., O'Rourke, K., Werbin, Z., Lee, L., Abdi, R., Mehmeti, D., & Xiong, L.
633 (2024). *neonSoilFlux: Compute Soil Carbon Fluxes for the National Ecological Observatory*
634 *Network Sites*.
- 635 Zobitz, J., & Zimmerman, N. (2025). *Supporting Code and Data for neonSoilFlux: An R*
636 *Package for Continuous Sensor-Based Estimation of Soil CO₂ Fluxes*. Zenodo. <https://doi.org/10.5281/zenodo.5500000>

639 //doi.org/10.5281/zenodo.17516320