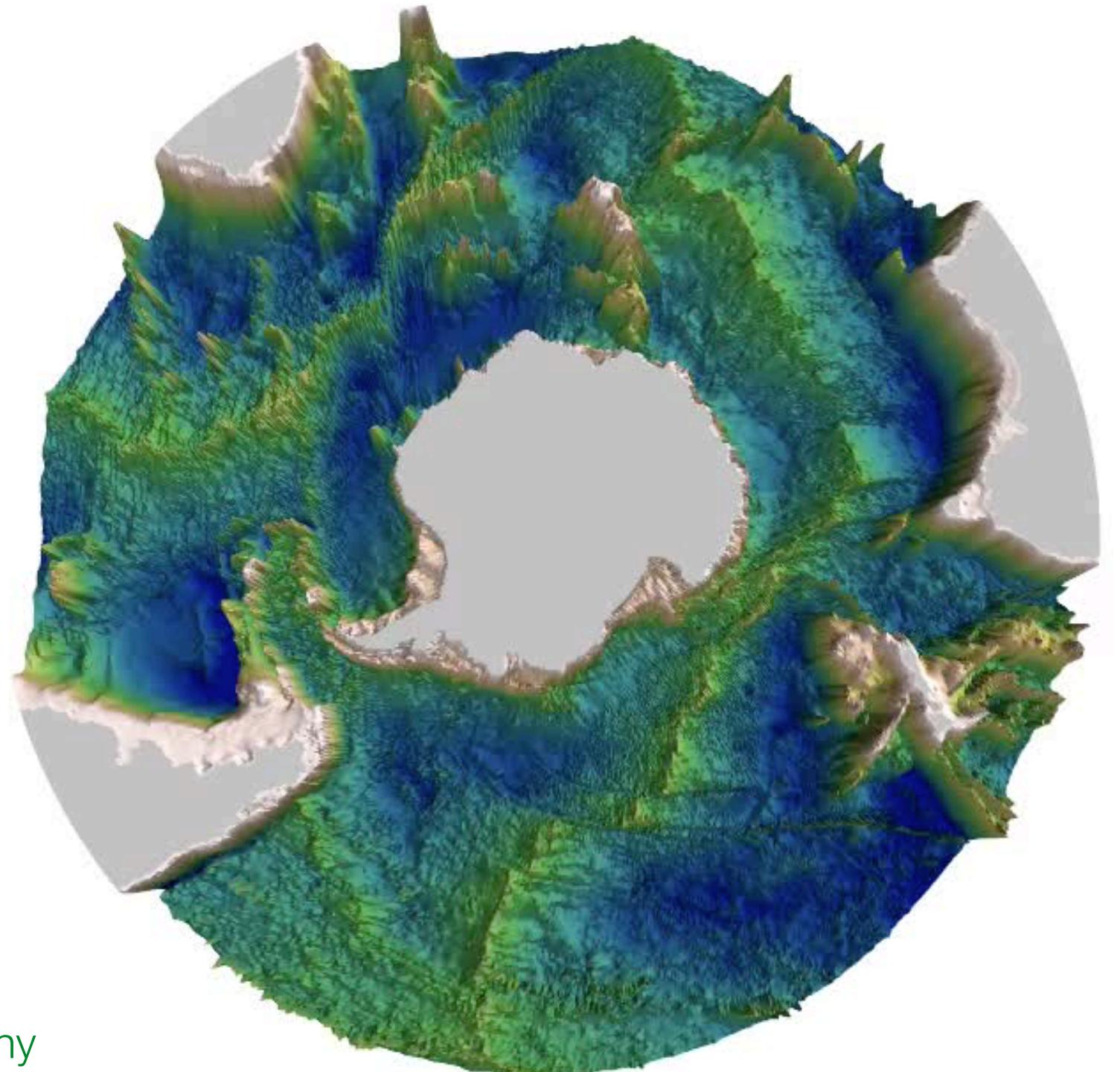


Eddy saturation in a barotropic model



Navid C. Constantinou & William R. Young
Scripps Institution of Oceanography, UC San Diego



Animated view of the
Southern Ocean topography
by V. Tamsitt.

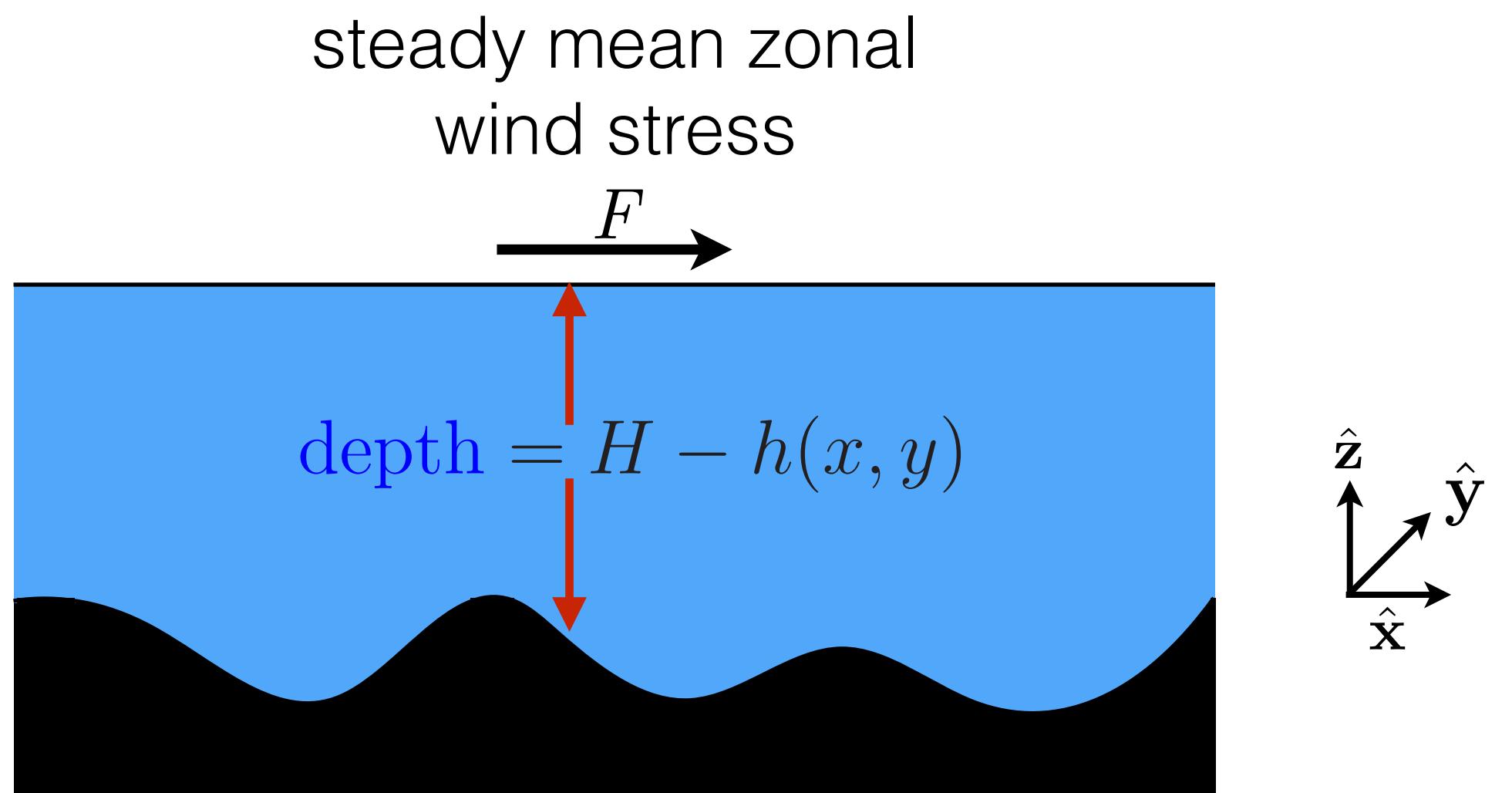
Can a **barotropic** QG model
exhibit “eddy saturation”?

What is the minimum requirements
for “eddy saturation”?

the plan

Revisit an old **barotropic** QG model on a β -plane.

(Hart 1979, Davey 1980, Bretherton & Haidvogel 1976,
Holloway 1987, Carnevale & Fredericksen 1987)



A distinctive feature of this model is a
“large-scale **barotropic** zonal flow” $U(t)$.

↑
this is
the “ACC”

Study how momentum is balanced by
topographic form stress and investigate the
requirements for eddy saturation.

topographic PV

total streamfunction

QGPV

$$\eta = \frac{f_0 h}{H}$$

$$-U(t)y + \psi(x, y, t)$$

$$\nabla^2 \psi + \eta + \beta y$$

a barotropic QG model for mid-ocean region

total streamfunction $-U(t)y + \psi(x, y, t)$

QGPV $\nabla^2\psi + \eta + \beta y$

Material conservation of QGPV

$$\begin{aligned}\nabla^2\psi_t + U(\nabla^2\psi + \eta)_x + J(\psi, \nabla^2\psi + \eta) \\ + \beta\psi_x = -\mu \nabla^2\psi + \text{hyper visc.}\end{aligned}$$

Large-scale zonal momentum

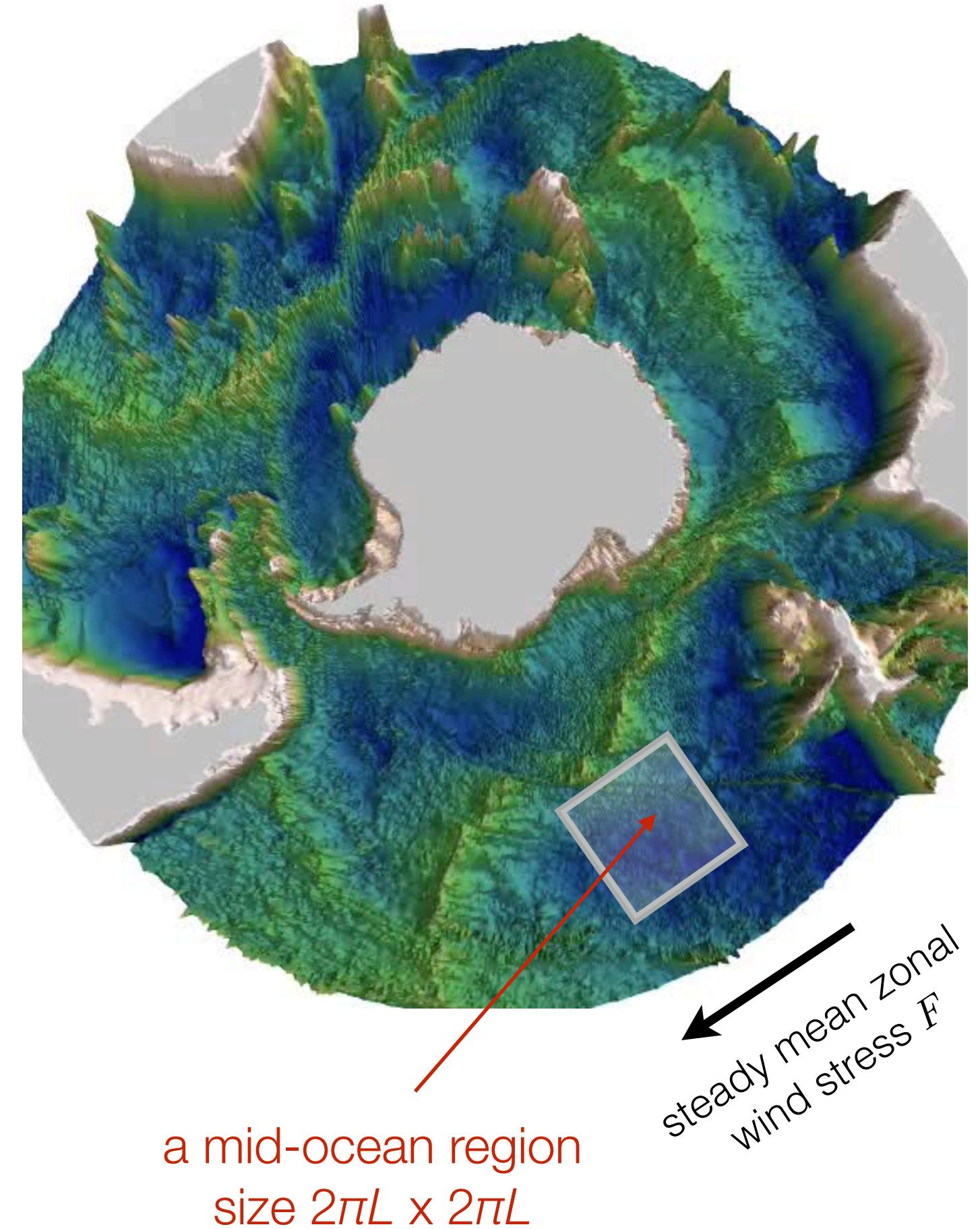
$$U_t = F - \mu U - \langle \psi \eta_x \rangle$$

↑
topographic
form stress

vertically integrated and
horizontally averaged
zonal angular
momentum equation

$\langle \rangle$ is domain average

$$F = \frac{\tau_s}{\rho_0 H} : \text{wind stress forcing}$$



periodic boundary conditions

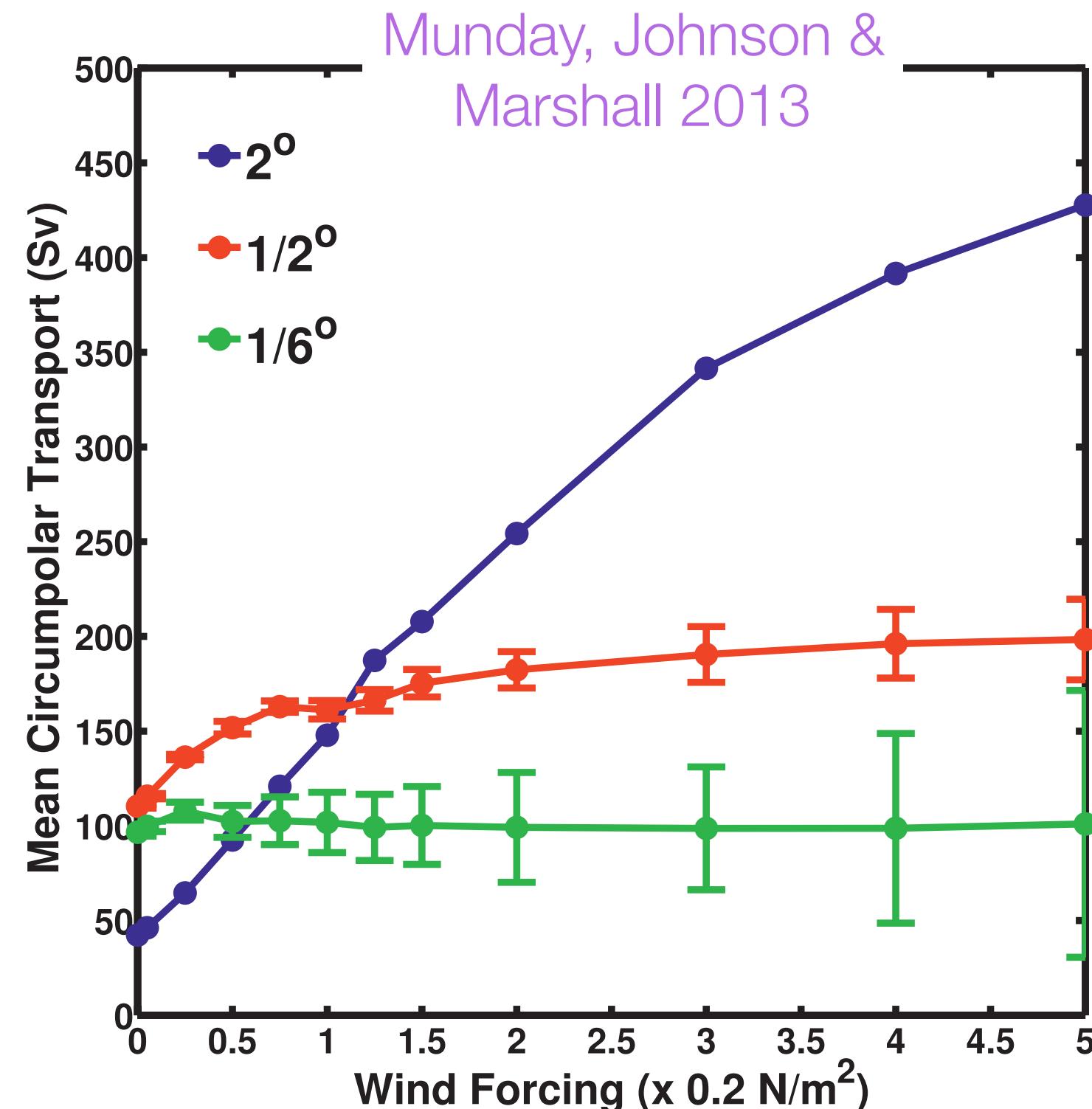
Question:

Does this **barotropic** QG model
show “eddy saturation”?

Do we need **baroclinicity**?
Do we even need channel walls?

what is “eddy saturation”?

The *insensitivity* of the total ACC volume transport to wind stress increase.



Eddy saturation is seen in eddy-resolving ocean models.

Higher resolution → eddy saturation “occurs”

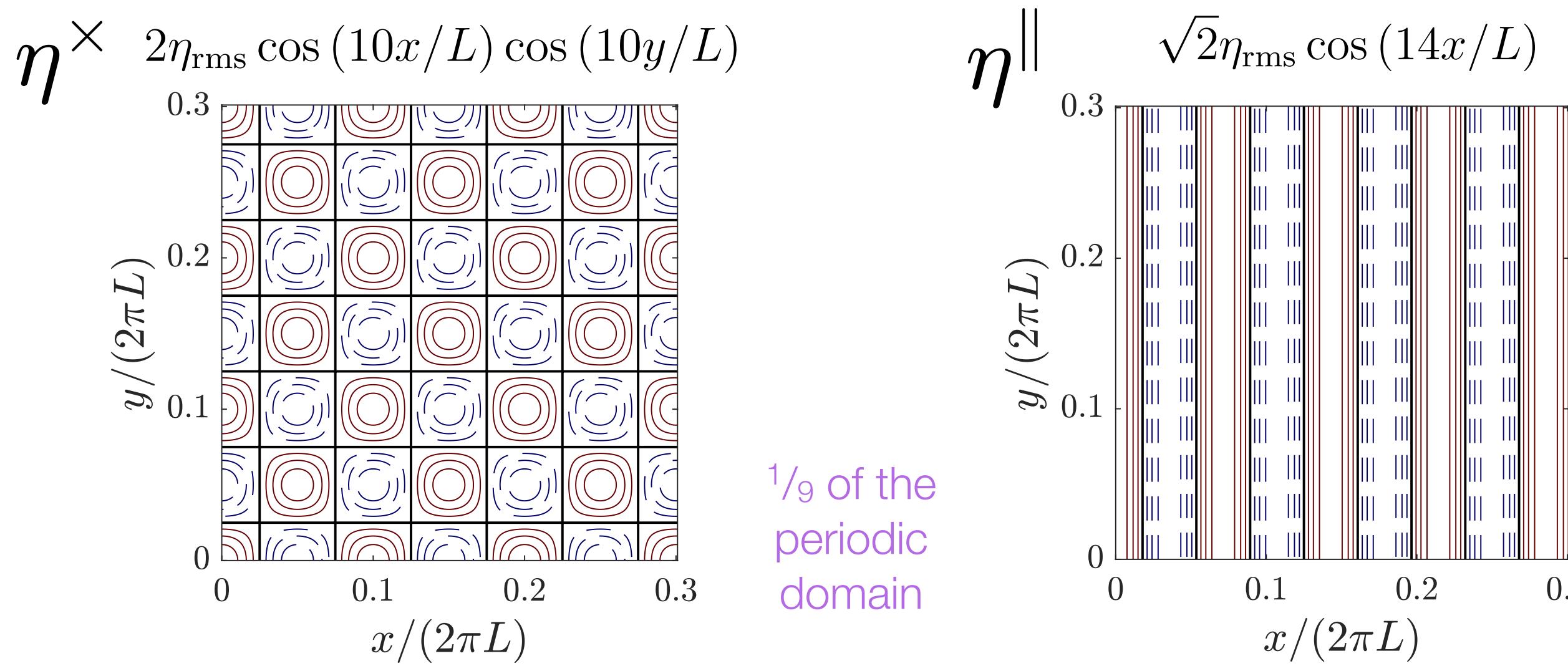
Eddy saturation was theoretically predicted by Straub (1993)
but with an *entirely baroclinic* argument.
(vertical momentum transfer interfacial eddy form stress)

[There are many other examples: Hallberg & Gnanadesikan 2001, Tansley & Marshall 2001, Hallberg & Gnanadesikan 2006, Hogg et al. 2008, Nadeau & Straub 2009, Farneti et al. 2010, Nadeau & Straub 2012, Meredith et al. 2012, Morisson & Hogg 2013, Abernathey & Cessi 2014, Farneti et al. 2015, Nadeau & Ferrari 2015, Marshall et al. 2016.]

Question:

So, does this **barotropic** QG model
show eddy saturation or not?

let's use these two topographies



(both topographies imply the same length-scale: $\ell_\eta = \sqrt{\frac{\langle \eta^2 \rangle}{\langle |\nabla \eta|^2 \rangle}} = 0.07L$)

put some “quasi-realistic” numbers

$$L = 775 \text{ km} \quad H = 4 \text{ km} \quad \rho_0 = 1035 \text{ kg m}^{-3}$$

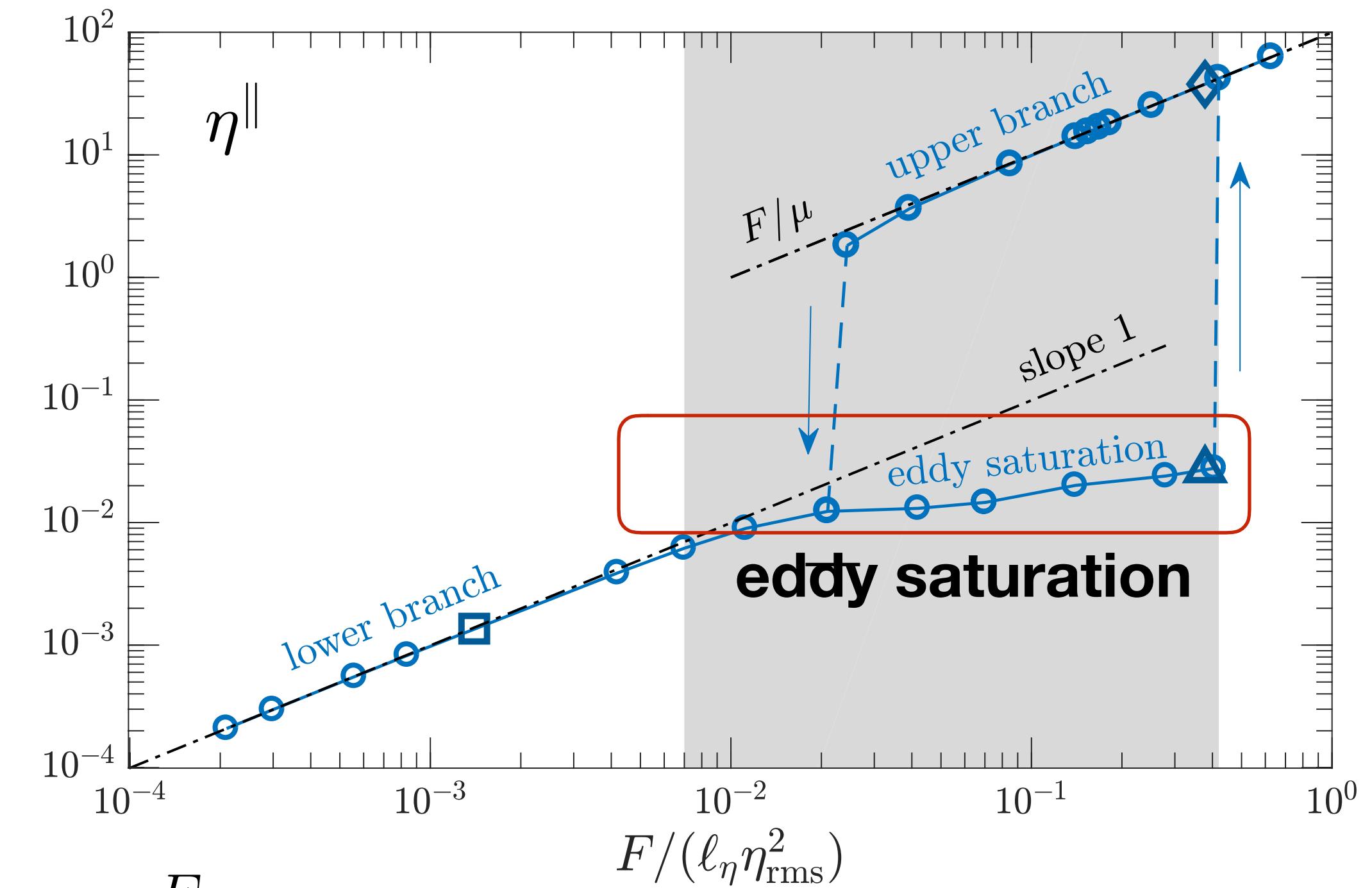
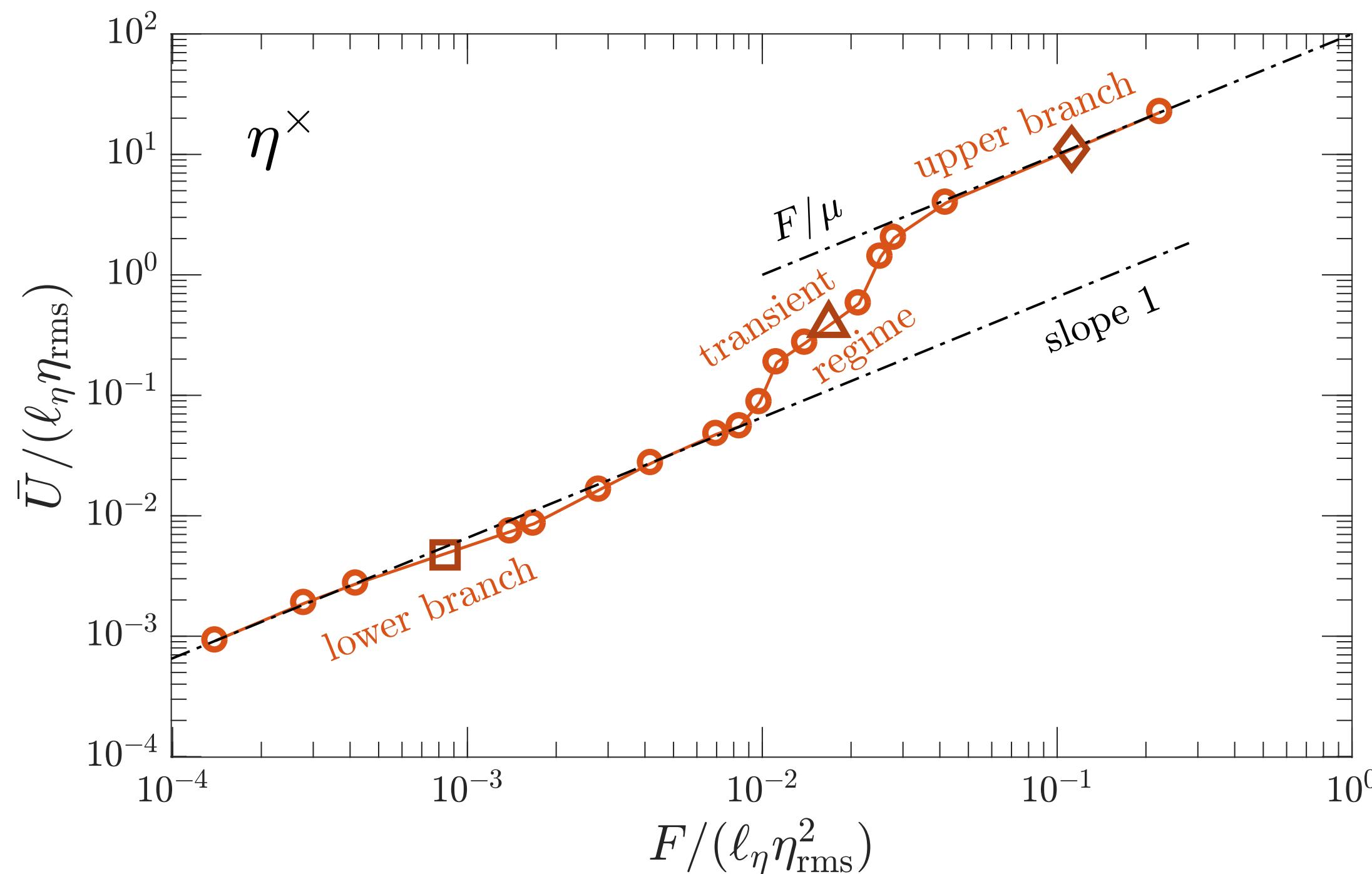
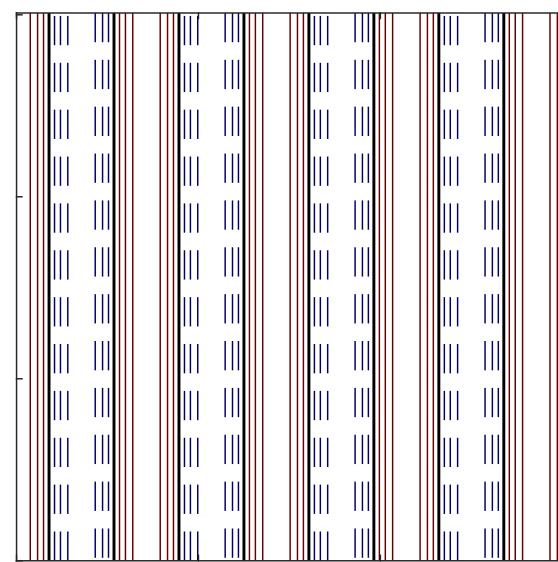
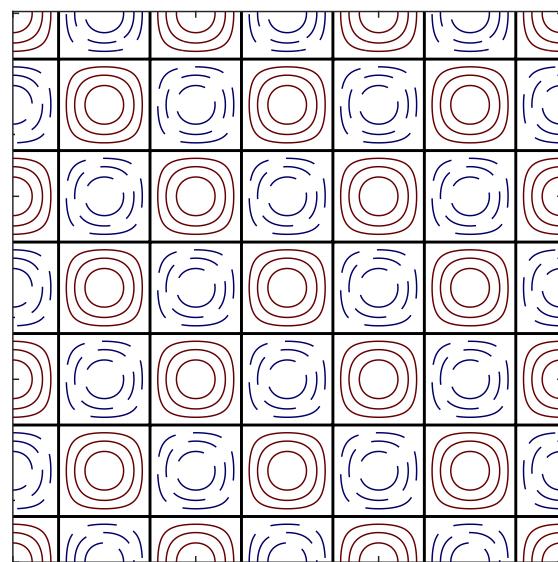
$$f_0 \quad \& \quad \beta \quad \text{for } 60^\circ\text{S}$$

$$\mu = (180 \text{ days})^{-1}$$

$$h_{\text{rms}} = 200 \text{ m} \Rightarrow \eta_{\text{rms}} = (1.8 \text{ days})^{-1}$$

and vary the wind stress...

how does the transport \bar{U} vary with wind stress?

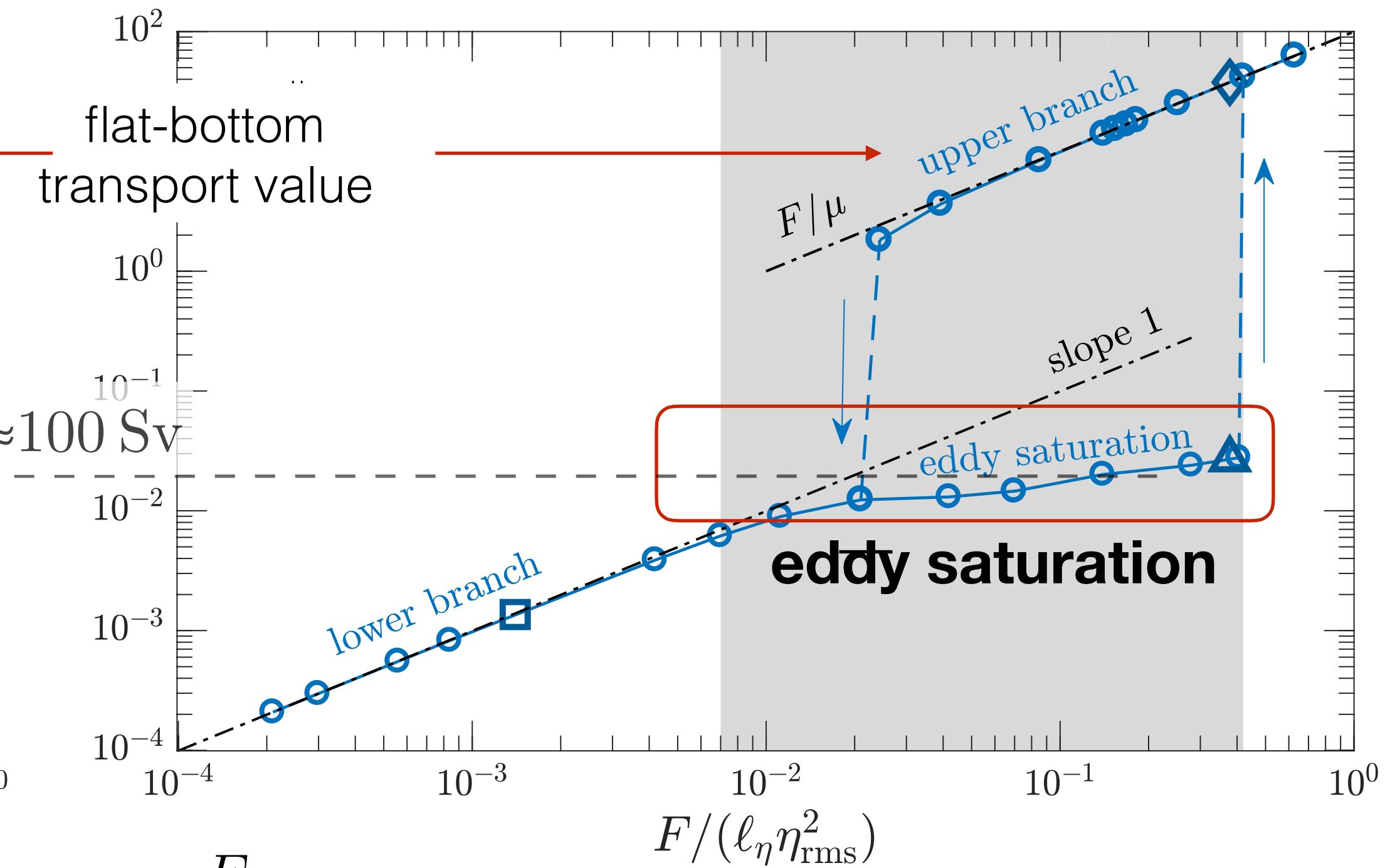
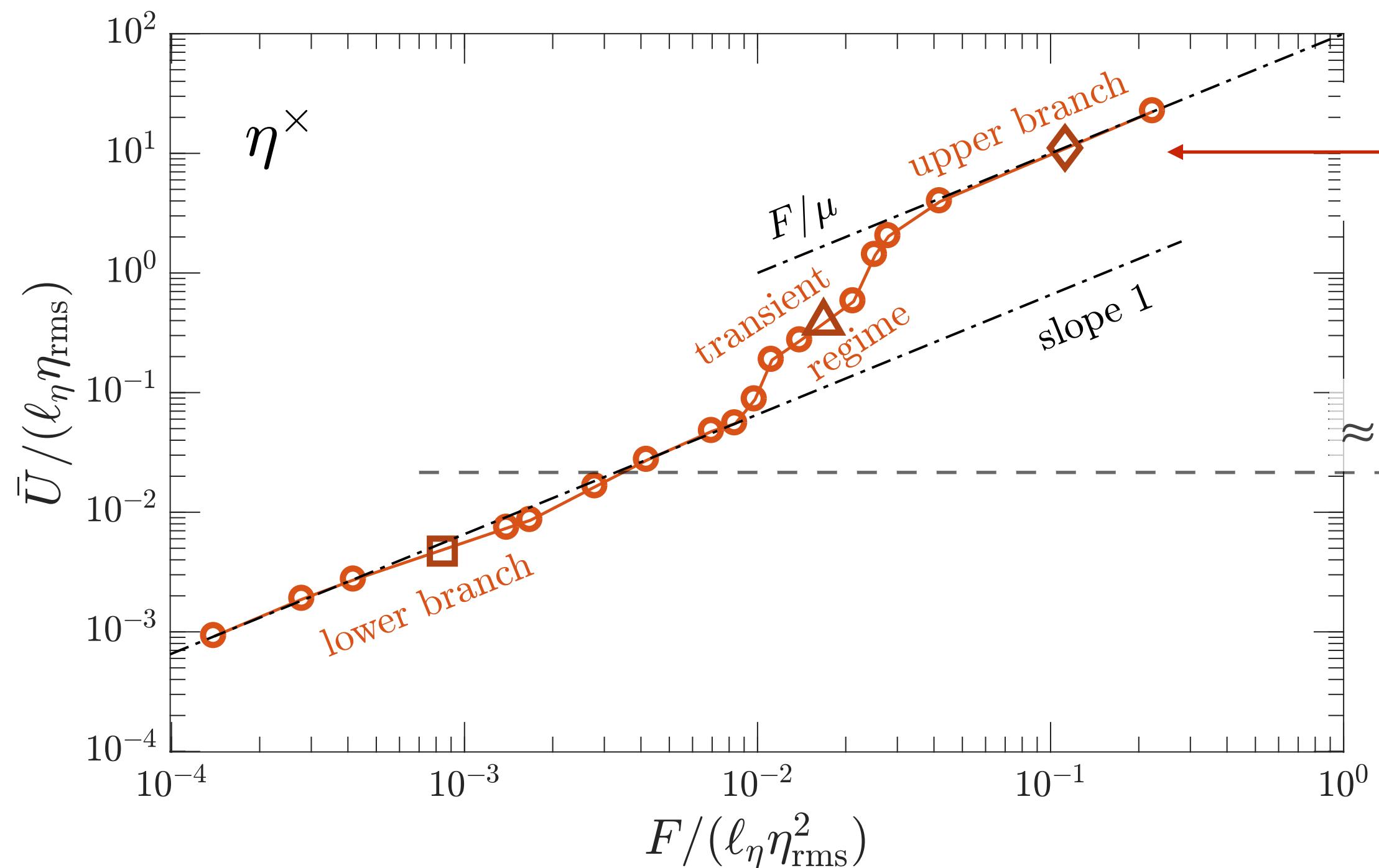
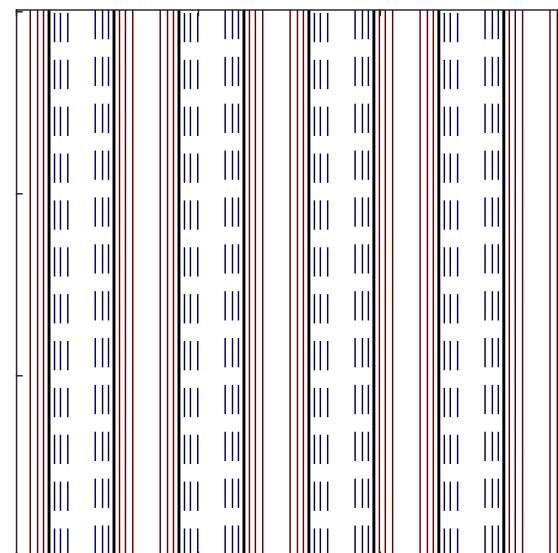
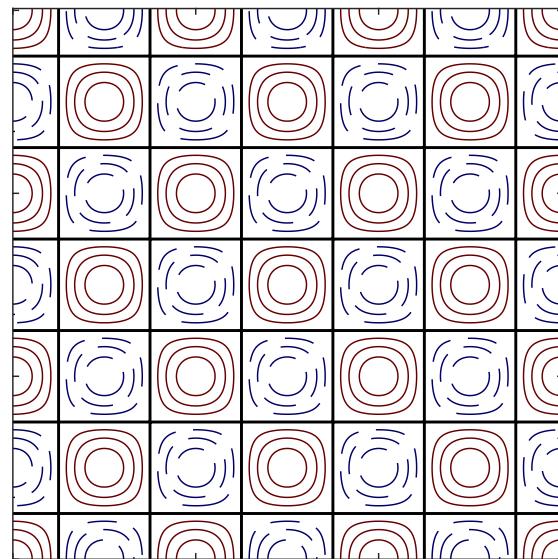


$$\tau = 0.2 \text{ N m}^{-2} \Leftrightarrow \frac{F}{\ell_\eta \eta_{\text{rms}}^2} \approx 0.02$$

do we understand why?

U increases 4 times
over a 60-fold wind stress increase

how does the transport \bar{U} vary with wind stress?



$$\tau = 0.2 \text{ N m}^{-2} \Leftrightarrow \frac{F}{\ell_\eta \eta_{\text{rms}}^2} \approx 0.02$$

do we understand why?

U increases 4 times
over a 60-fold wind stress increase

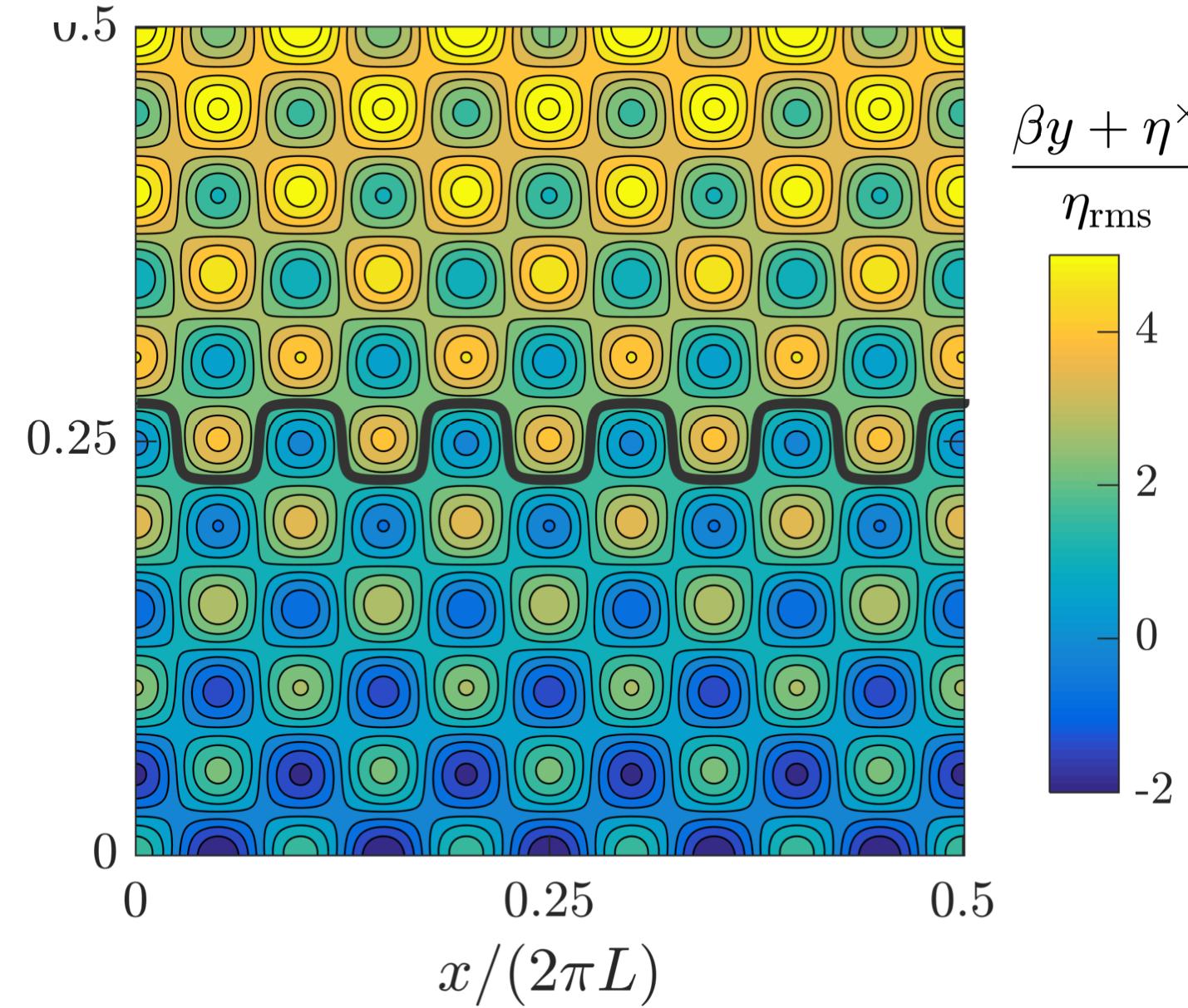
geostrophic contours

$$\beta y + \eta(x, y)$$

this is small-Rossby number expansion of $f/(H+h)$

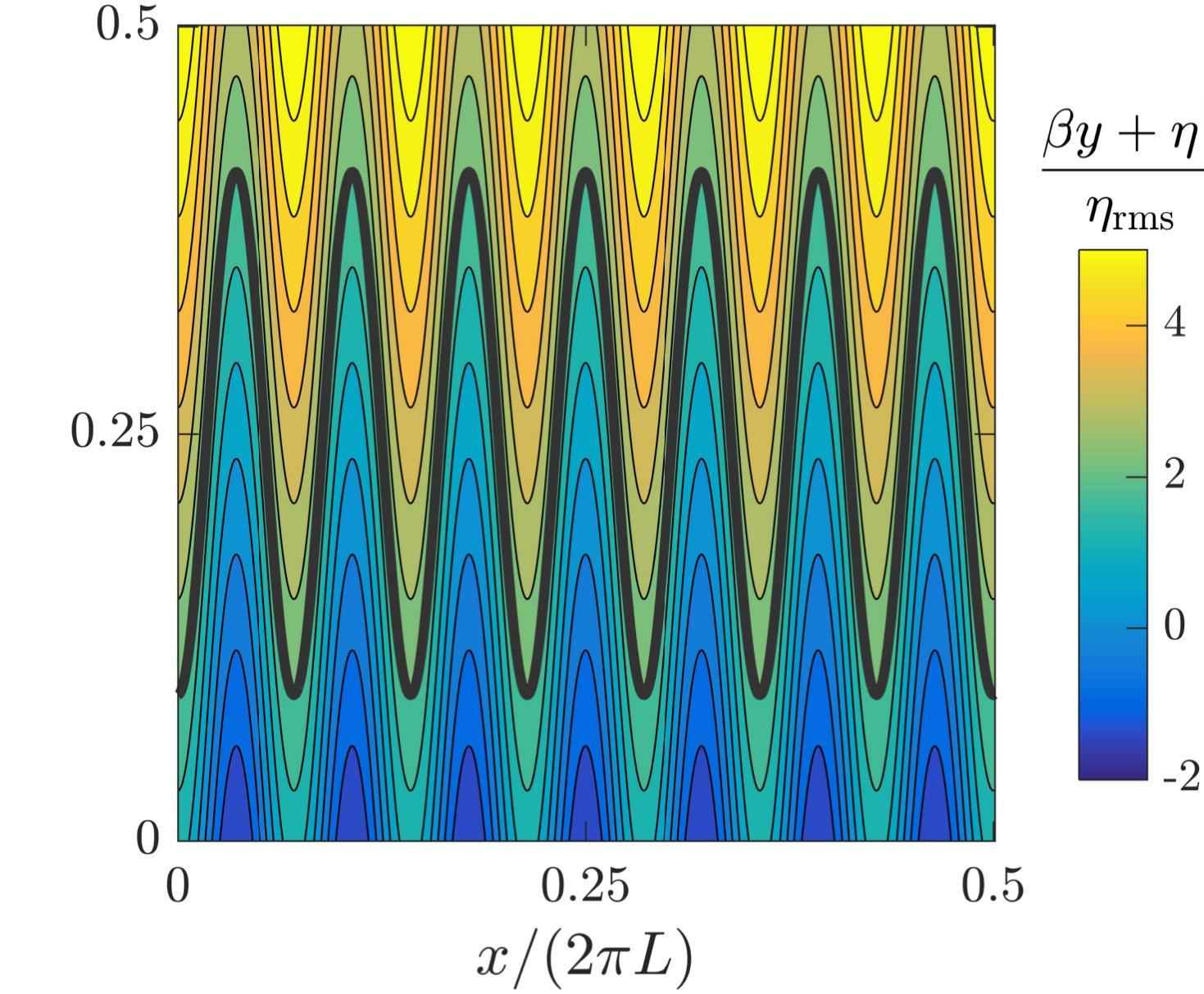
The *main control* parameter for whether eddy saturation occurs is the structure of the geostrophic contours.

η^\times closed geostrophic contours



most geostrophic contours
are **closed**

η^\parallel open geostrophic contours



all geostrophic contours
are **open**

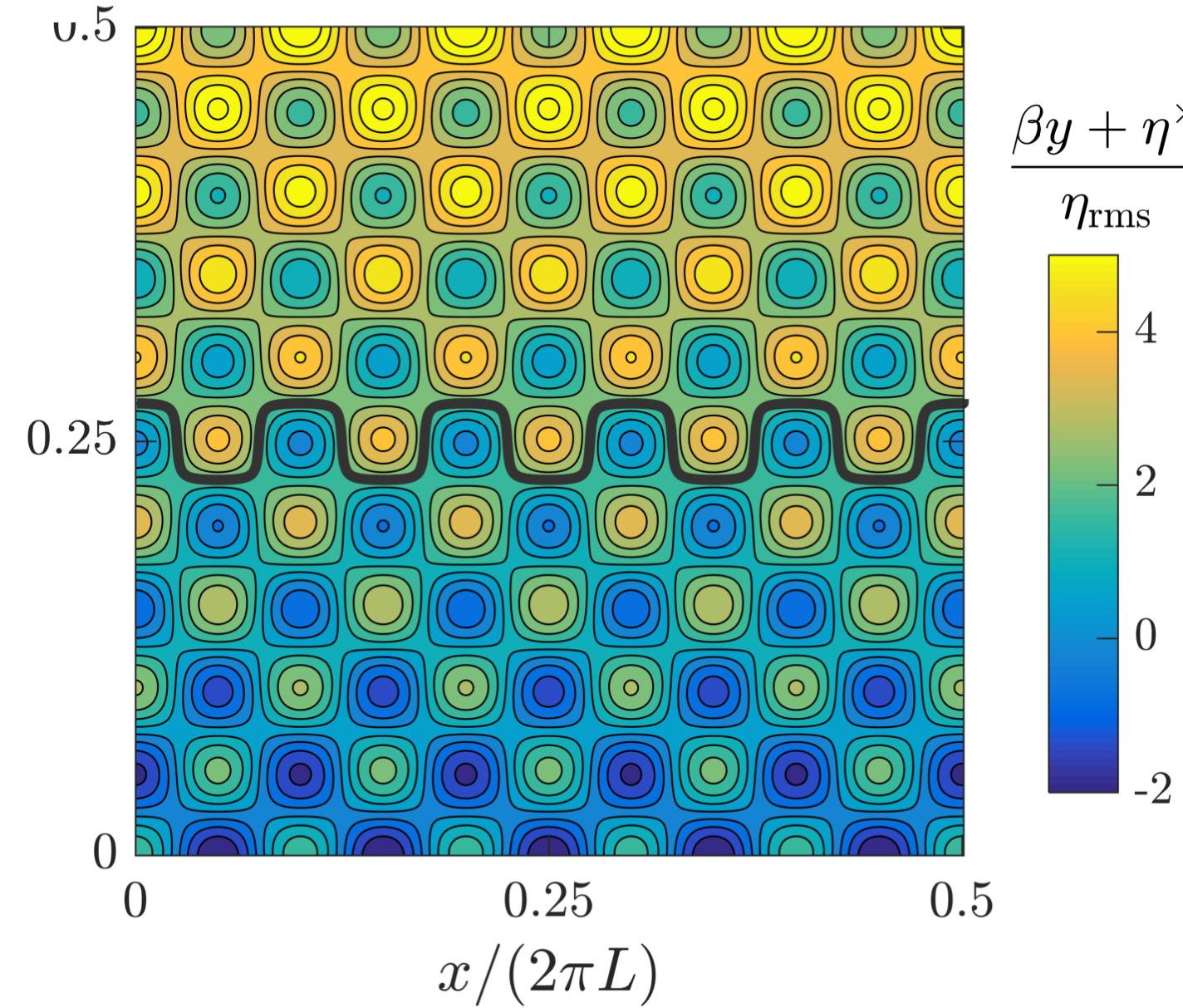
geostrophic contours

$$\beta y + \eta(x, y)$$

this is small-Rossby number expansion of $f/(H+h)$

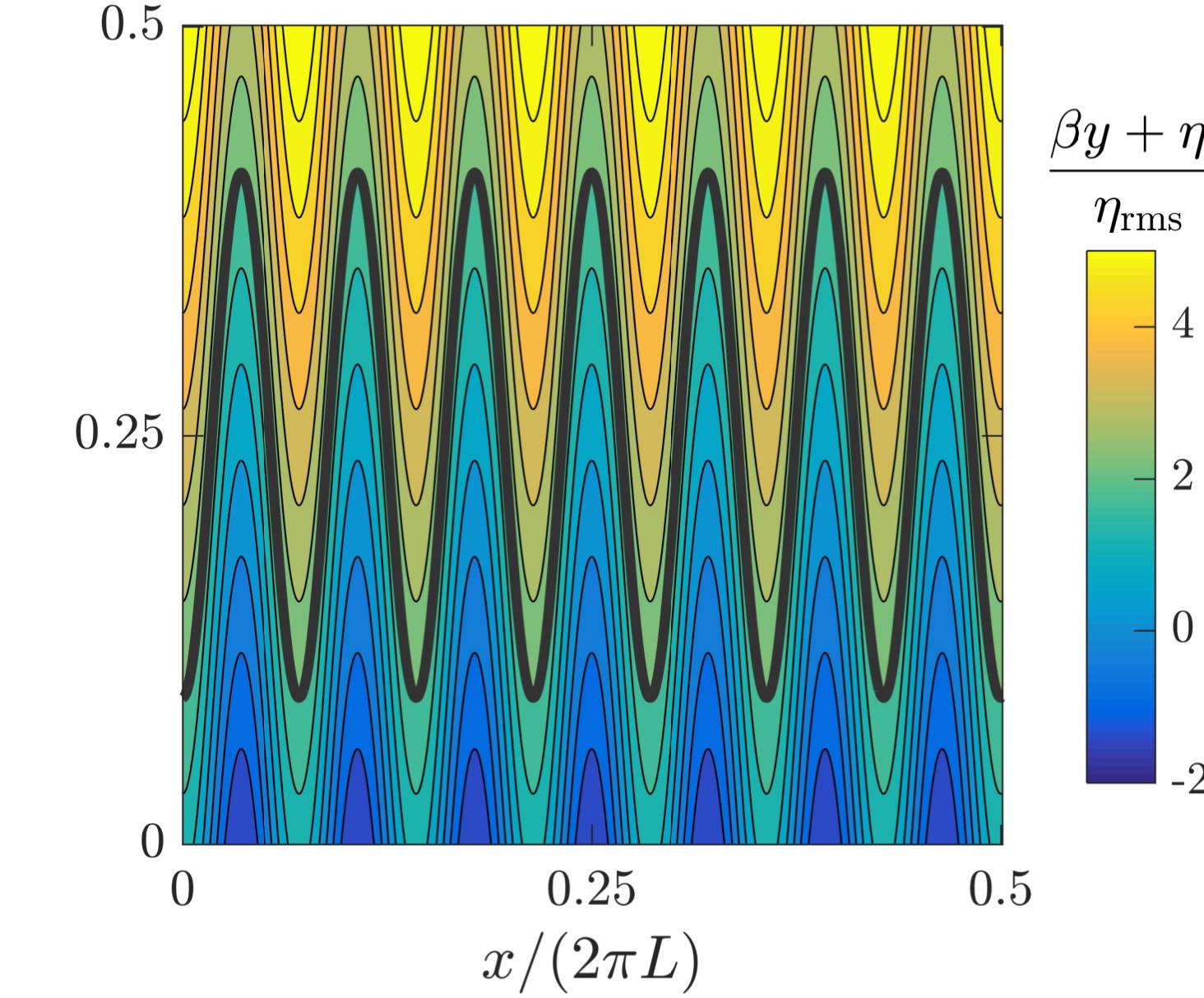
The *main control* parameter for whether eddy saturation occurs is the structure of the geostrophic contours.

η^{\times} closed geostrophic contours



most geostrophic contours
are **closed**

η^{\parallel} open geostrophic contours



all geostrophic contours
are **open**

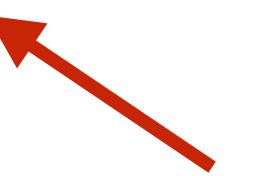
this
topography
exhibits
eddy
saturation

geostrophic contours

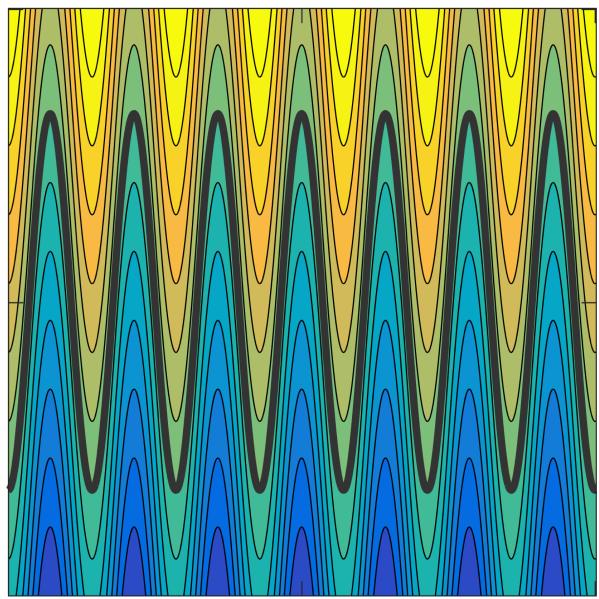
$$\beta y + \eta(x, y)$$

this is small-Rossby number
expansion of $f/(H+h)$

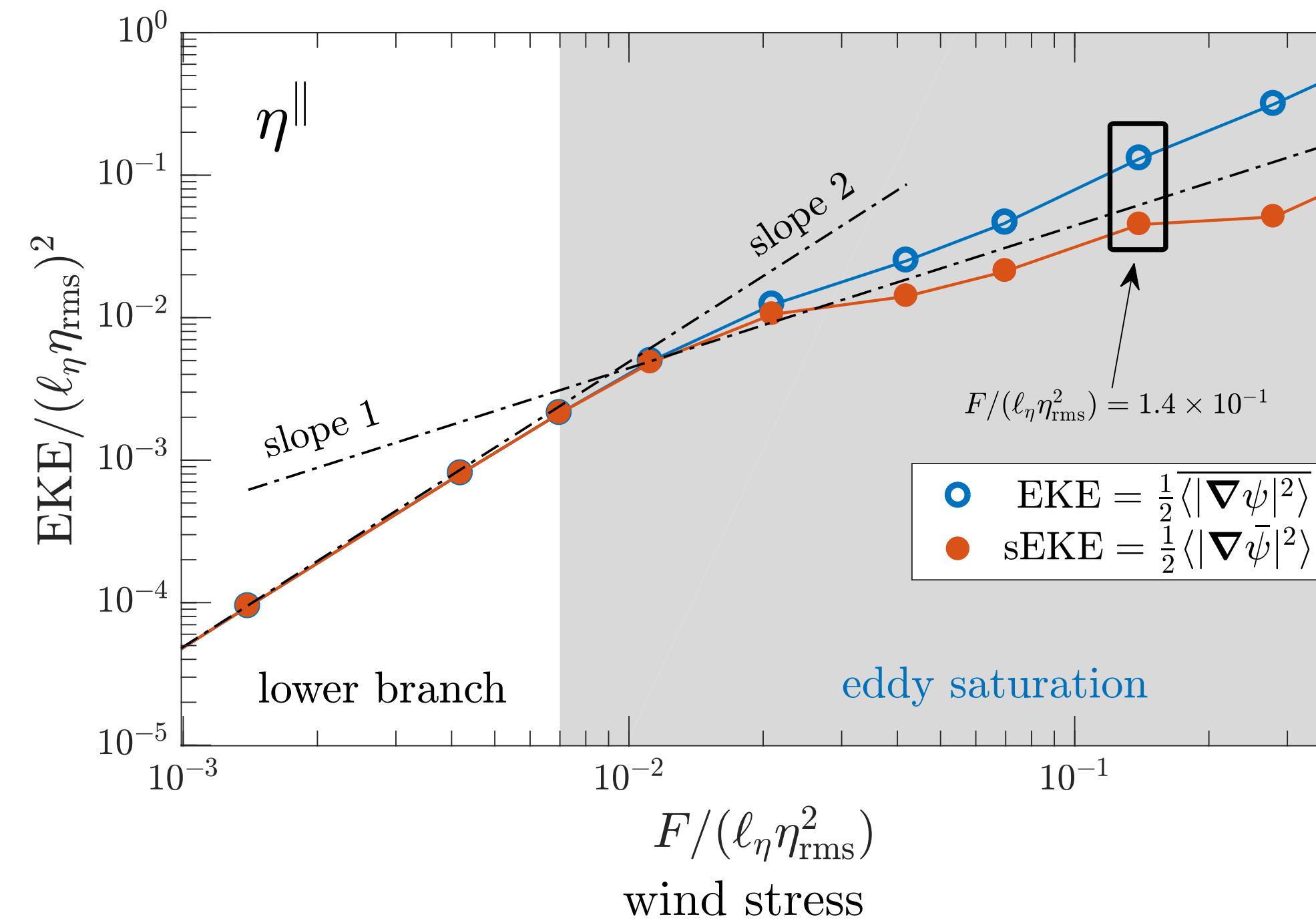
Eddy saturation occurs when
the geostrophic contours are “open”,
that is, when the geostrophic contours
span the domain in the zonal direction.



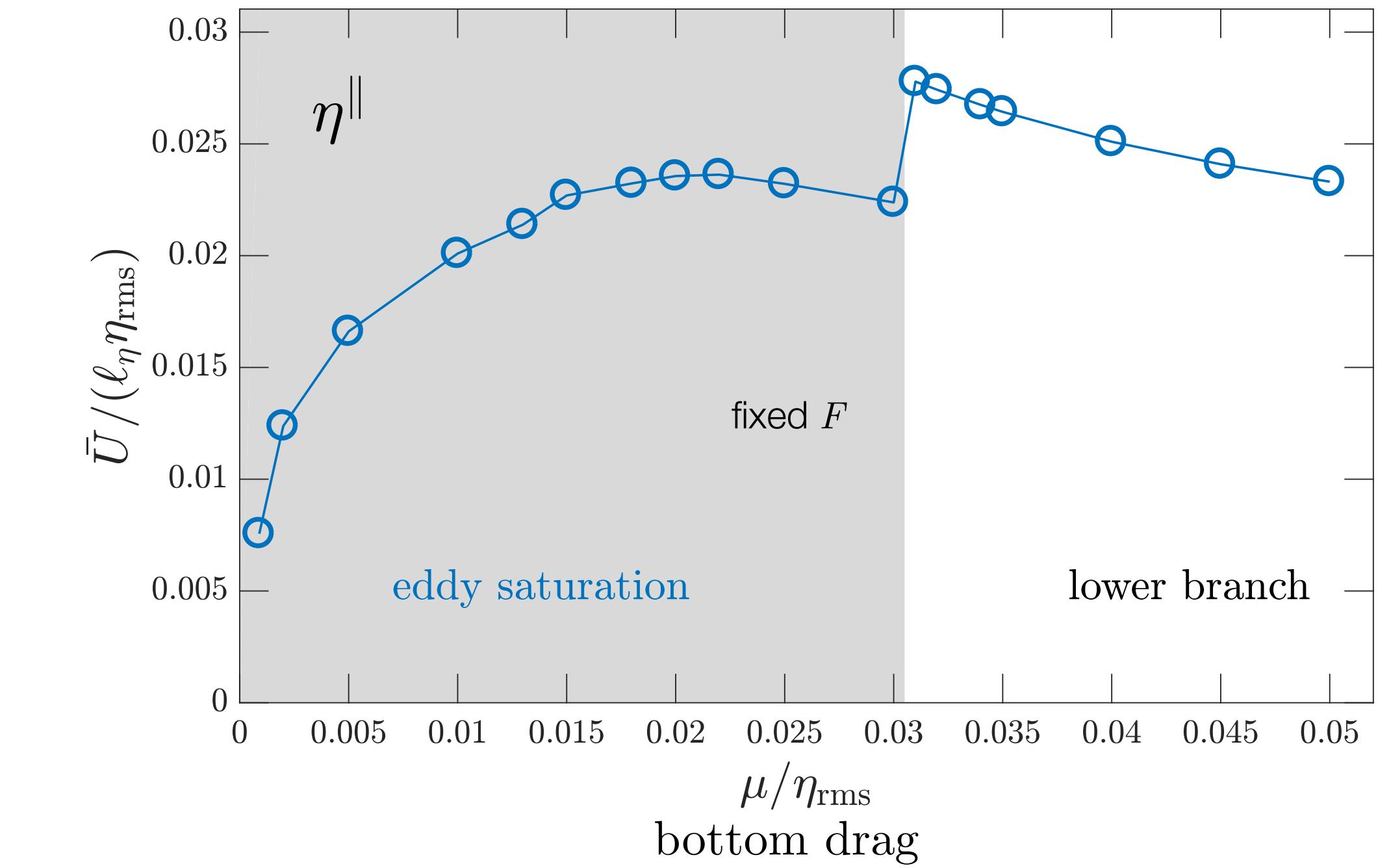
this is a general result
we've seen it in various cases
whatever the topography



further “symptoms” of eddy saturation



EKE grows roughly linearly
with wind stress



transport grows
with increasing bottom drag

take home messages

Eddy saturation *can occur without baroclinicity!*

The **barotropic** QG model shows eddy saturation
when geostrophic contours are **open**.

This is surprising! All previous arguments were based on **baroclinicity**.

This, does not preclude the role of **baroclinic** processes in the ACC equilibration.
But it does argues that **barotropic** processes do contribute.

We need new process models of **baroclinic** turbulence in which
the mean flow is wind-driven and topography exerts form stress.

(In progress)

thank you

Constantinou (2017). A barotropic model of eddy saturation. *JPO* (submitted, arXiv:1703.06594)

Constantinou and Young (2017). Beta-plane turbulence above monoscale topography. *JFM* (in revision, arXiv:1612.03374)