



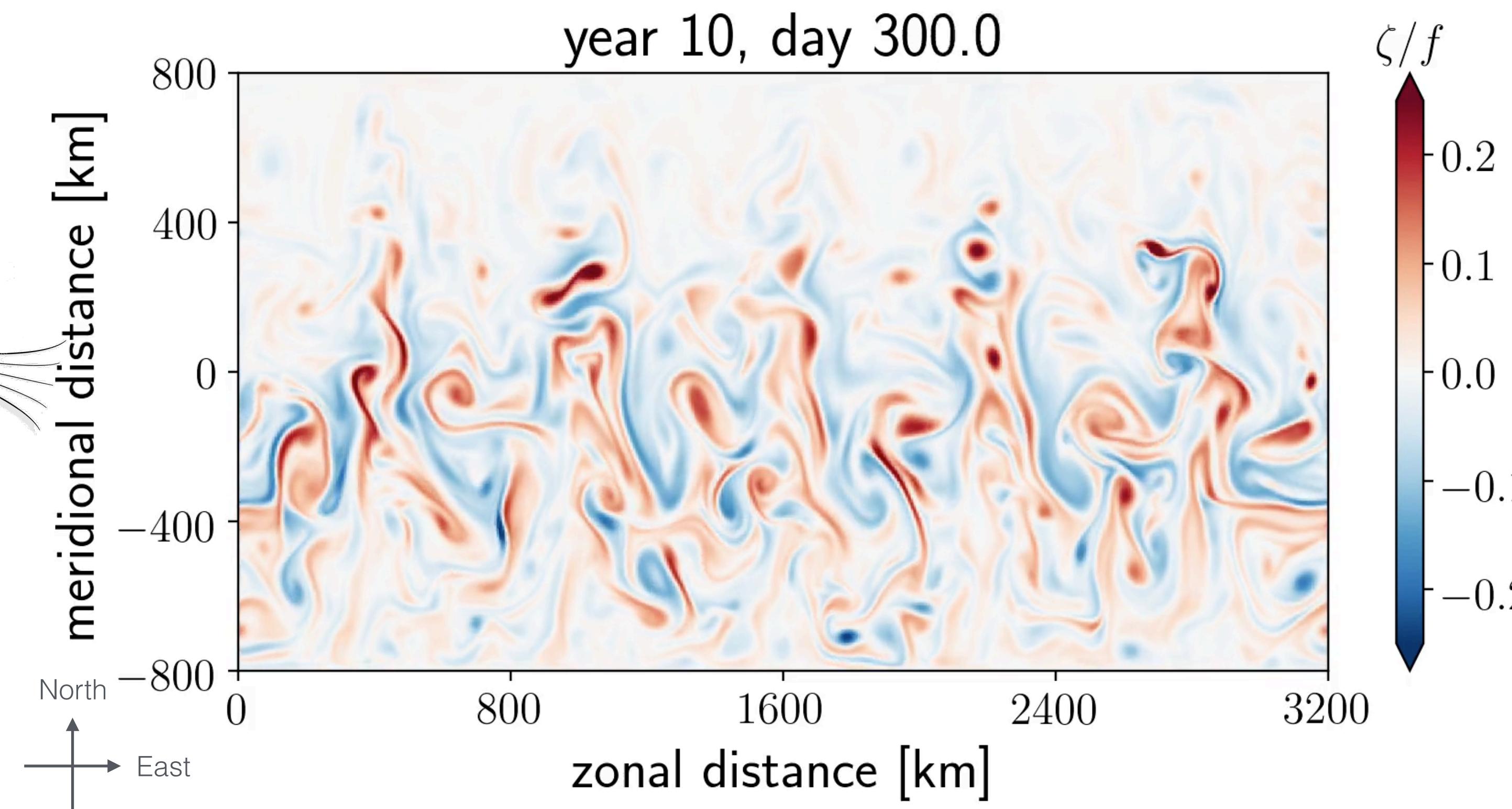
Australian National  
University

# Barotropic versus Baroclinic eddy saturation: implications to Southern Ocean dynamics

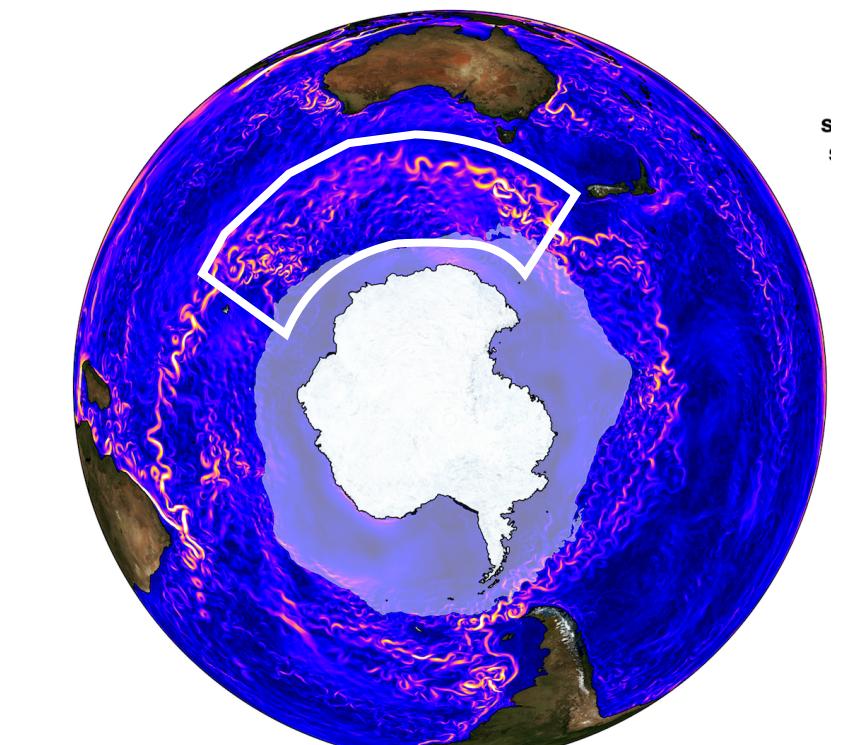


ARC Centre of Excellence  
for Climate Extremes

Navid Constantinou & Andy Hogg

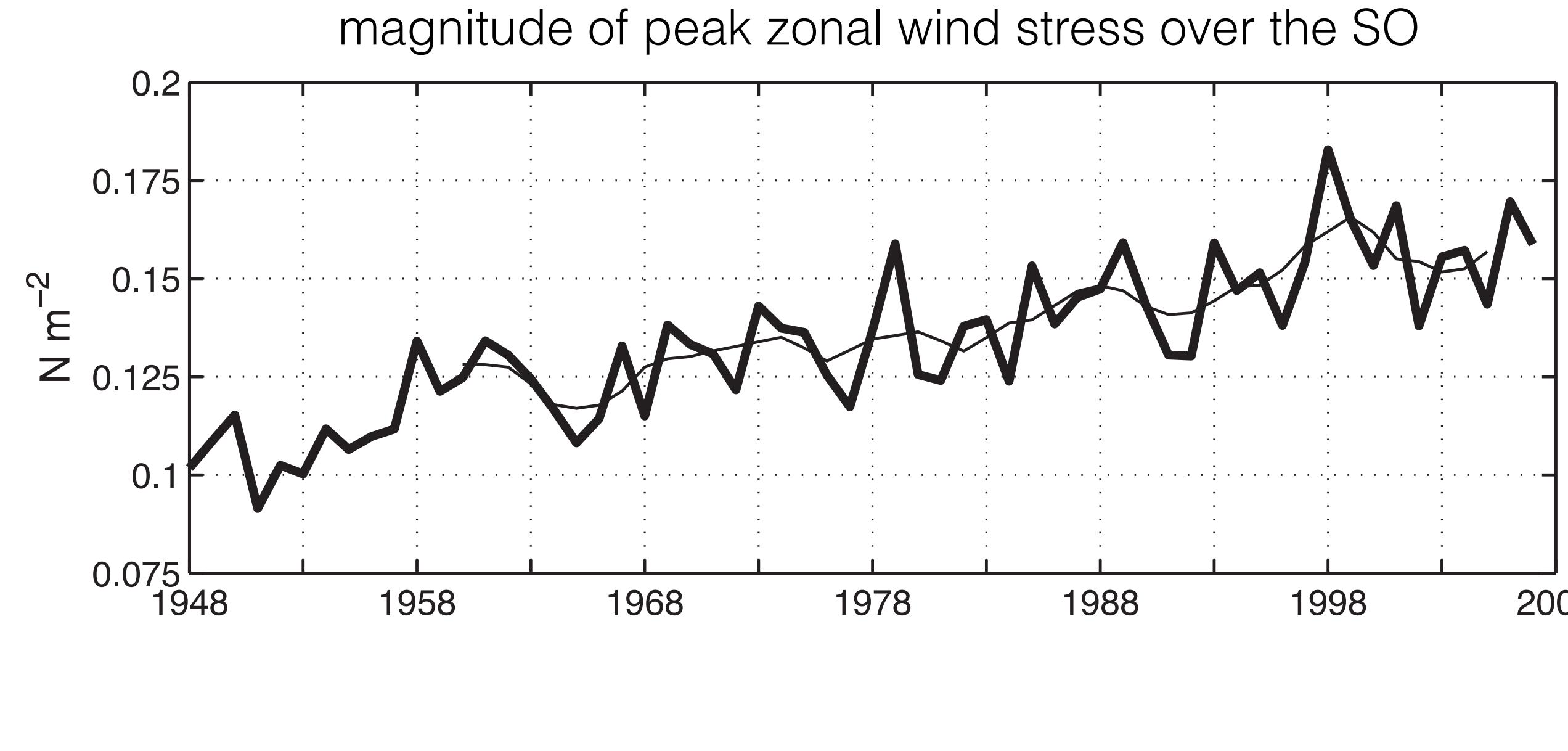


top-layer  
relative vorticity  
 $\zeta = \partial_x v - \partial_y u$



22<sup>nd</sup> AOFD Conference  
Portland ME, June 25th, 2019

# winds over Southern Ocean are getting stronger

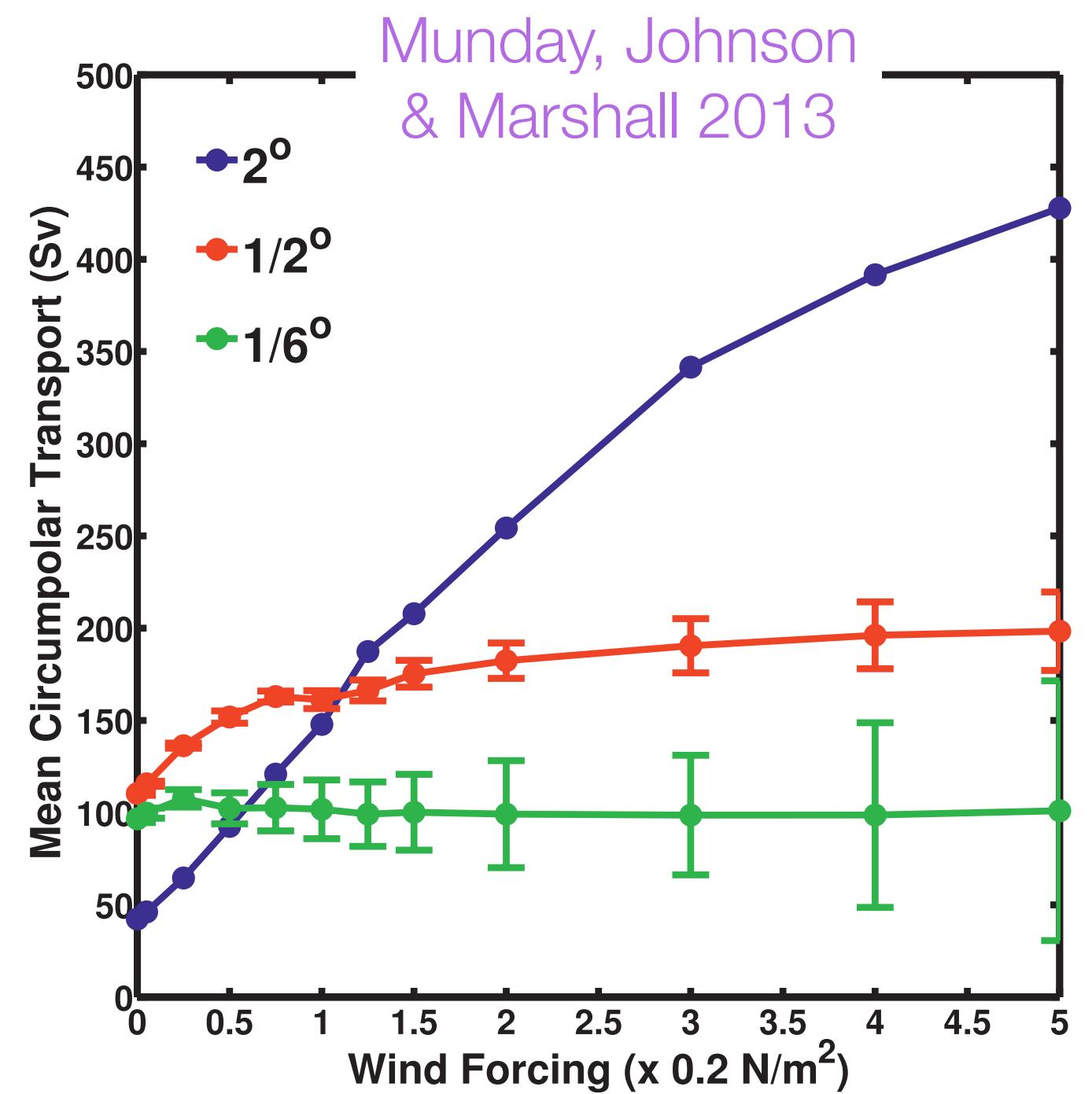


how will the Antarctic Circumpolar Current (ACC) respond?

does doubling the winds imply double ACC the transport?  
not always — “eddy saturation”

# what's eddy saturation?

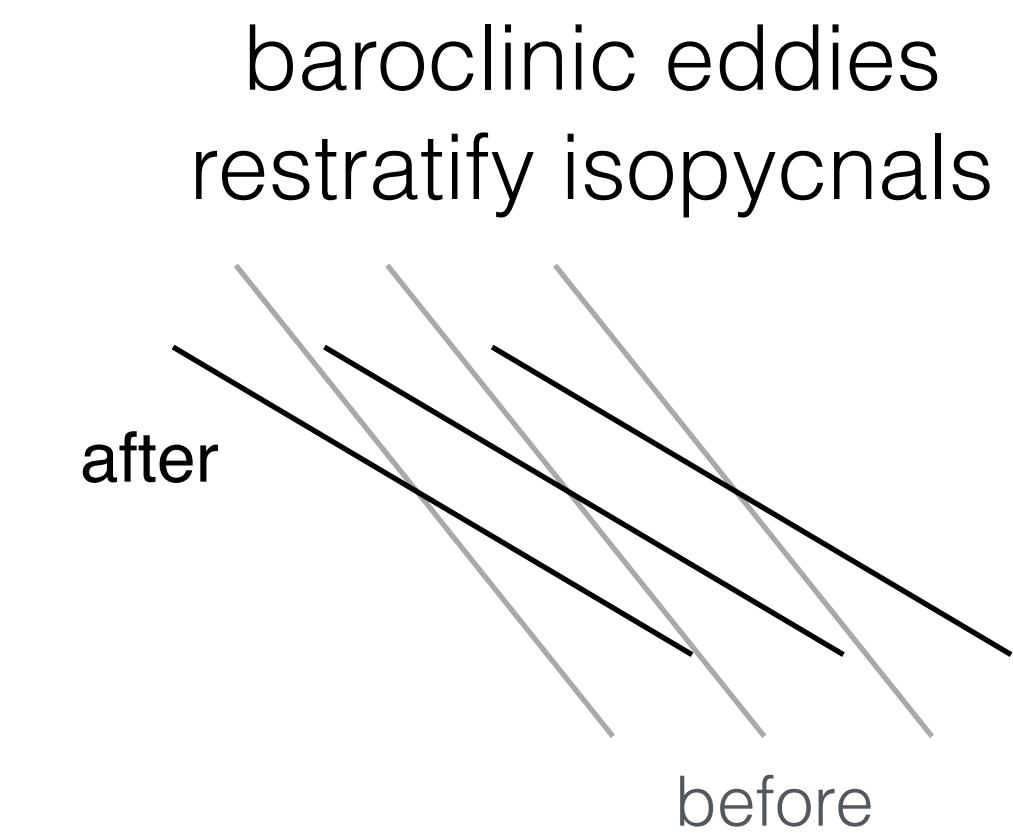
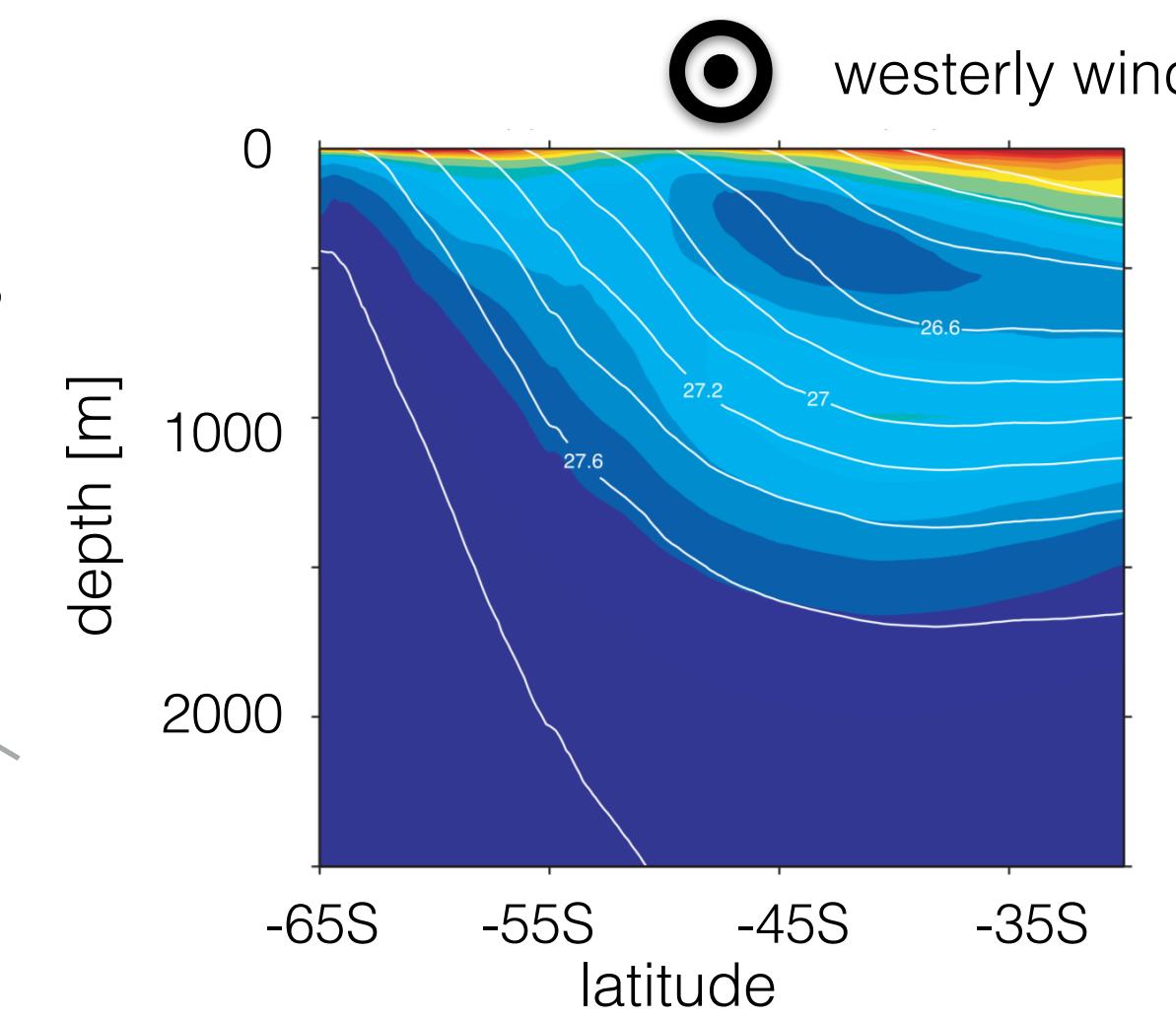
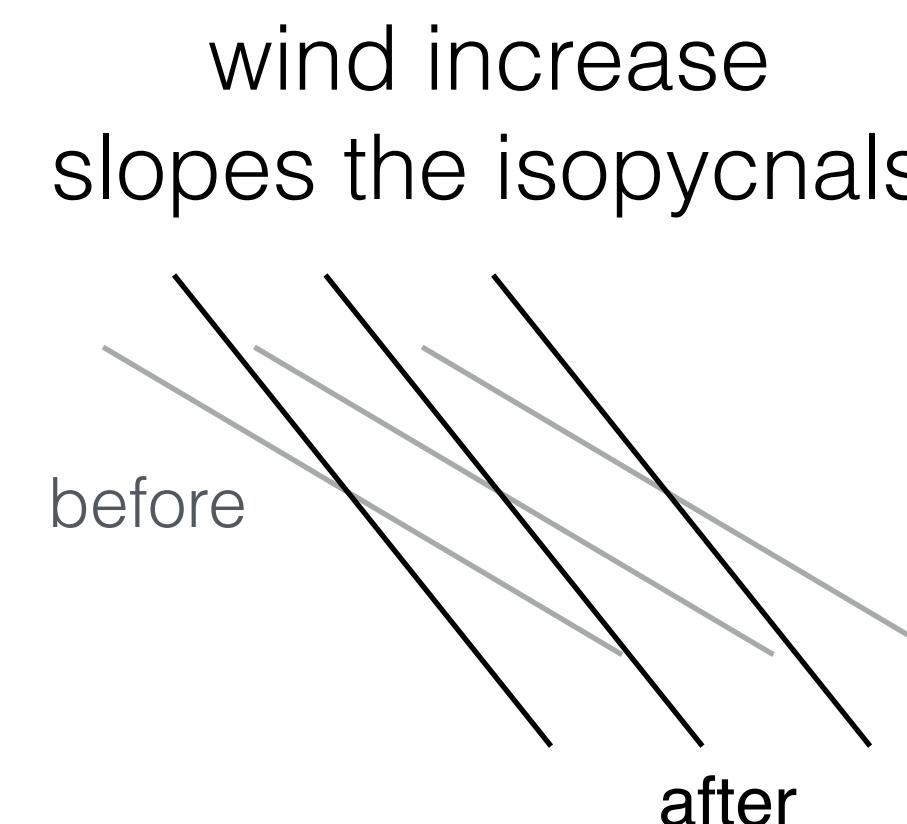
The *insensitivity* of the time-mean ACC volume transport to wind stress increase.



higher resolution → eddy saturation “occurs”

Eddy saturation is seen in eddy-resolving ocean models.  
(some hints also in obs.)

Eddy saturation was theoretically predicted by Straub (1993)  
but with an *entirely baroclinic* argument.



[Other examples: Hallberg & Gnanadesikan 2001, Tansley & Marshall 2001, Hallberg & Gnanadesikan 2006, Hogg et al. 2008, Nadeau & Straub 2009, 2012, Farneti et al. 2010, Meredith et al. 2012, Morisson & Hogg 2013, Abernathey & Cessi 2014, Farneti et al. 2015, Nadeau & Ferrari 2015, Marshall et al. 2017.]

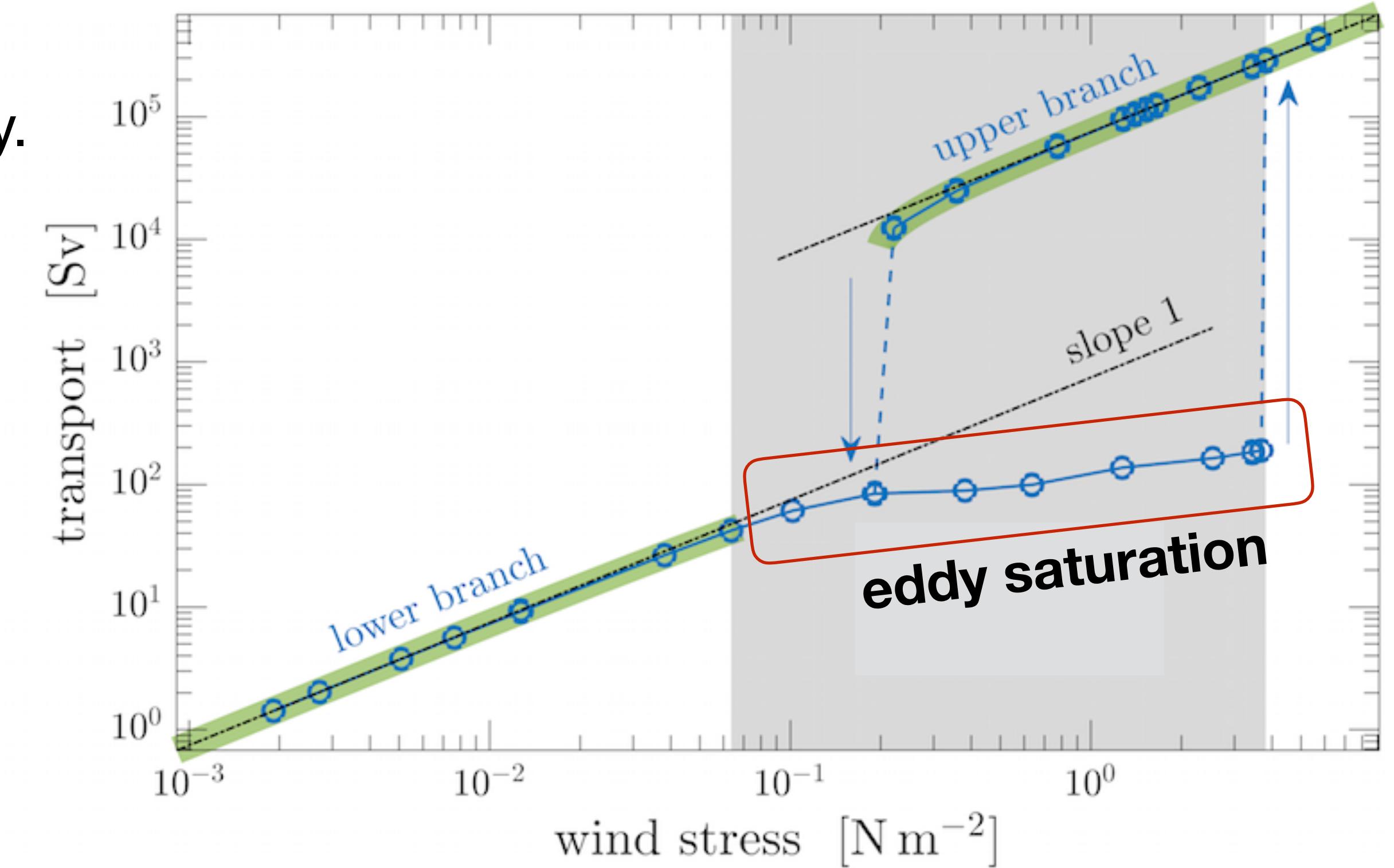
# in the previous episode

[back in 2017, in a Portland far far away]

Eddy saturation can occur *without* baroclinicity  
in a homogeneous QG barotropic model with bathymetry.

Surprising!

All previous arguments *relied* on baroclinic instability  
for producing transient eddies.



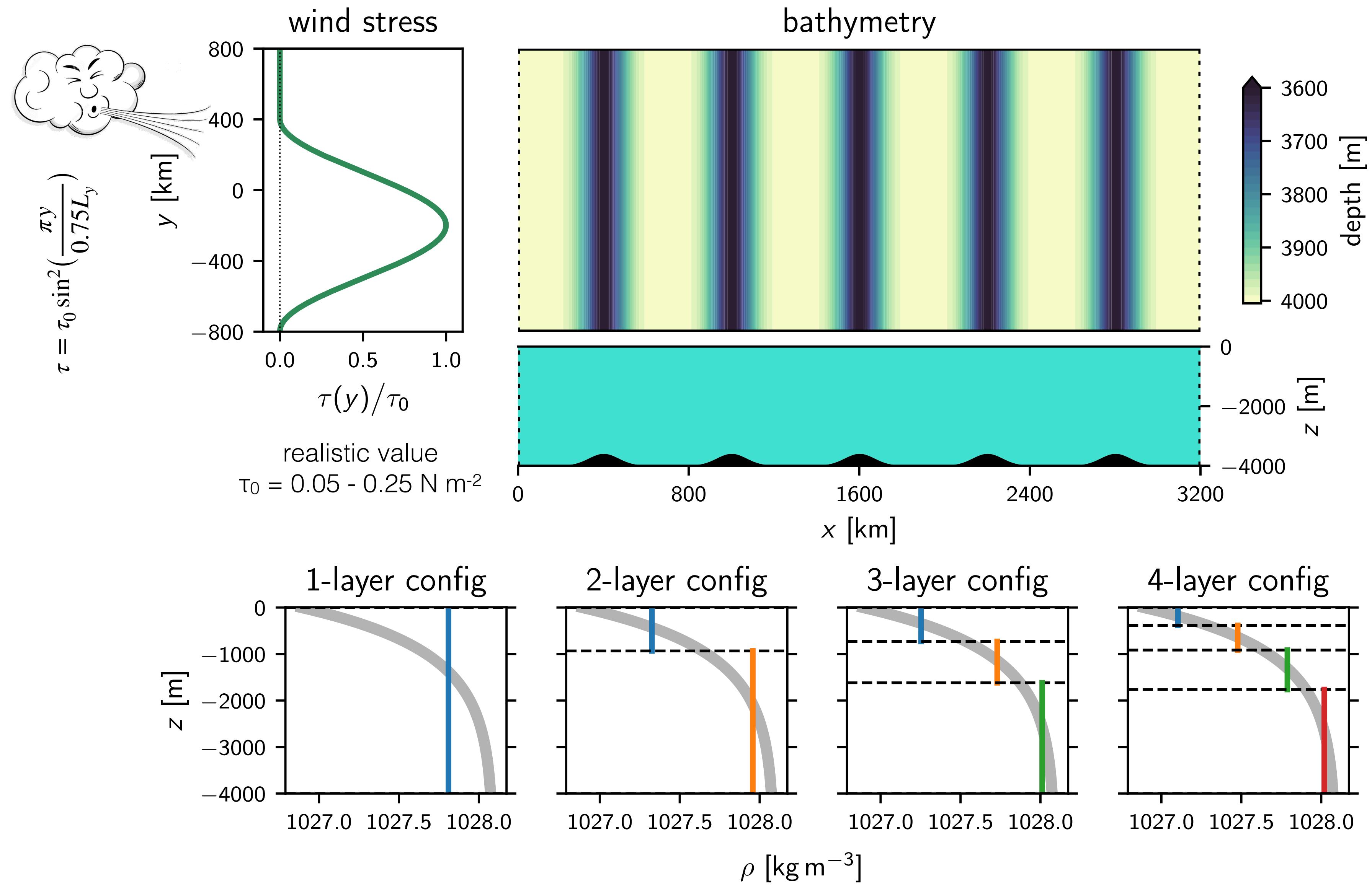
# what's the plan

Assess the relative role of  
**barotropic** versus **baroclinic** dynamics  
in establishing "eddy saturated" ocean states.

Use an isopycnal layered model  
with varying number of fluid layers.

# model setup

GFDL's MOM6  
primitive equations  
in isopycnal coordinates  
Boussinesq approximation



$\beta$ -plane  $f = f_0 + \beta y$   
1st deformation radius  $\approx 19 \text{ km}$   
zonally re-entrant  
free surface  
free-slip walls  
quadratic bottom drag  
grid spacing 4 km

bathymetry:  
Gaussian ridges  
400 m tall, half-width 165 km

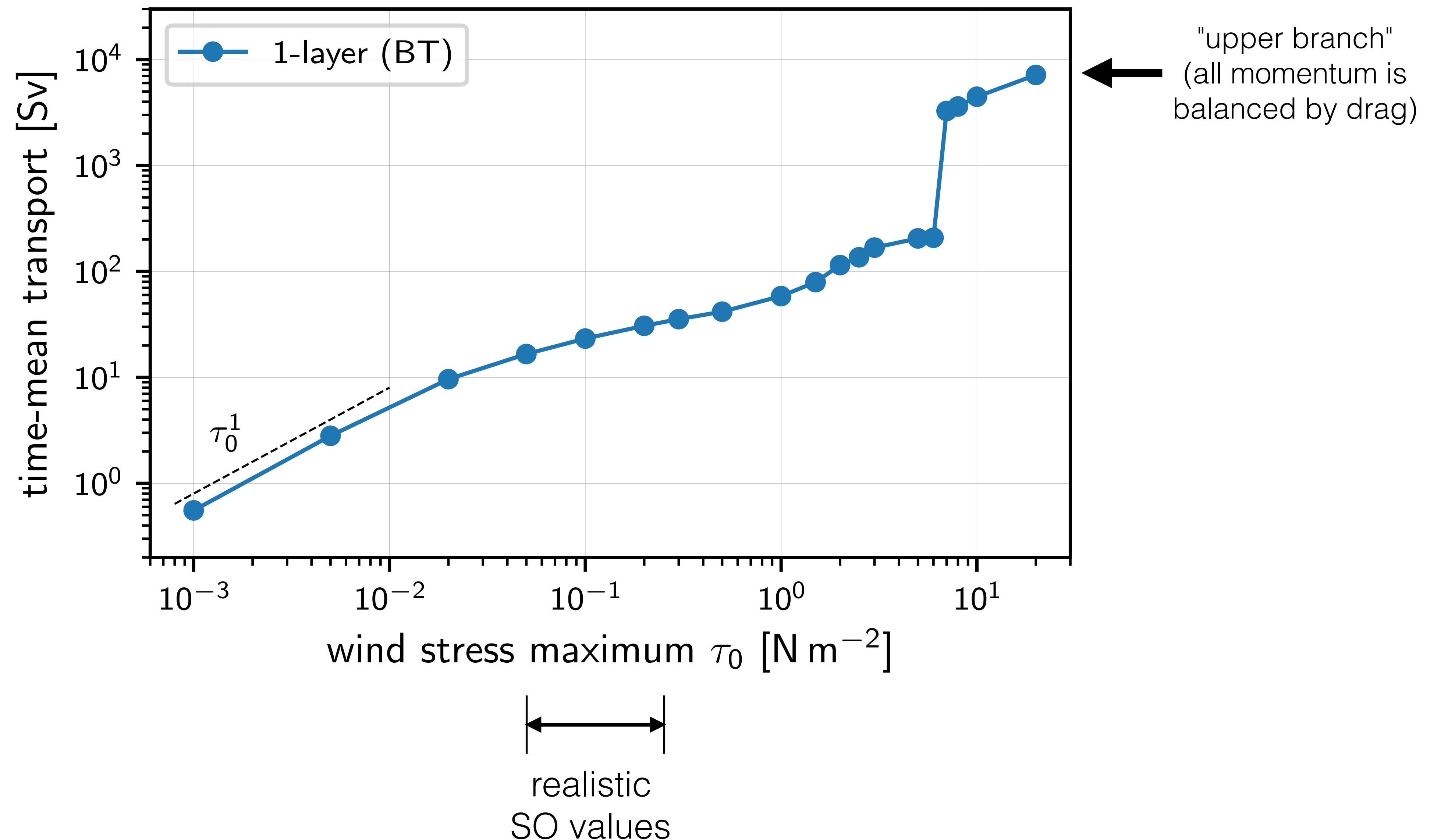
exponential density profile  

$$\rho = \rho_0 + \Delta\rho (1 - e^{z/d})$$
  
 $\Delta\rho = 1.2 \text{ kg m}^{-3}, d = 1 \text{ km}$

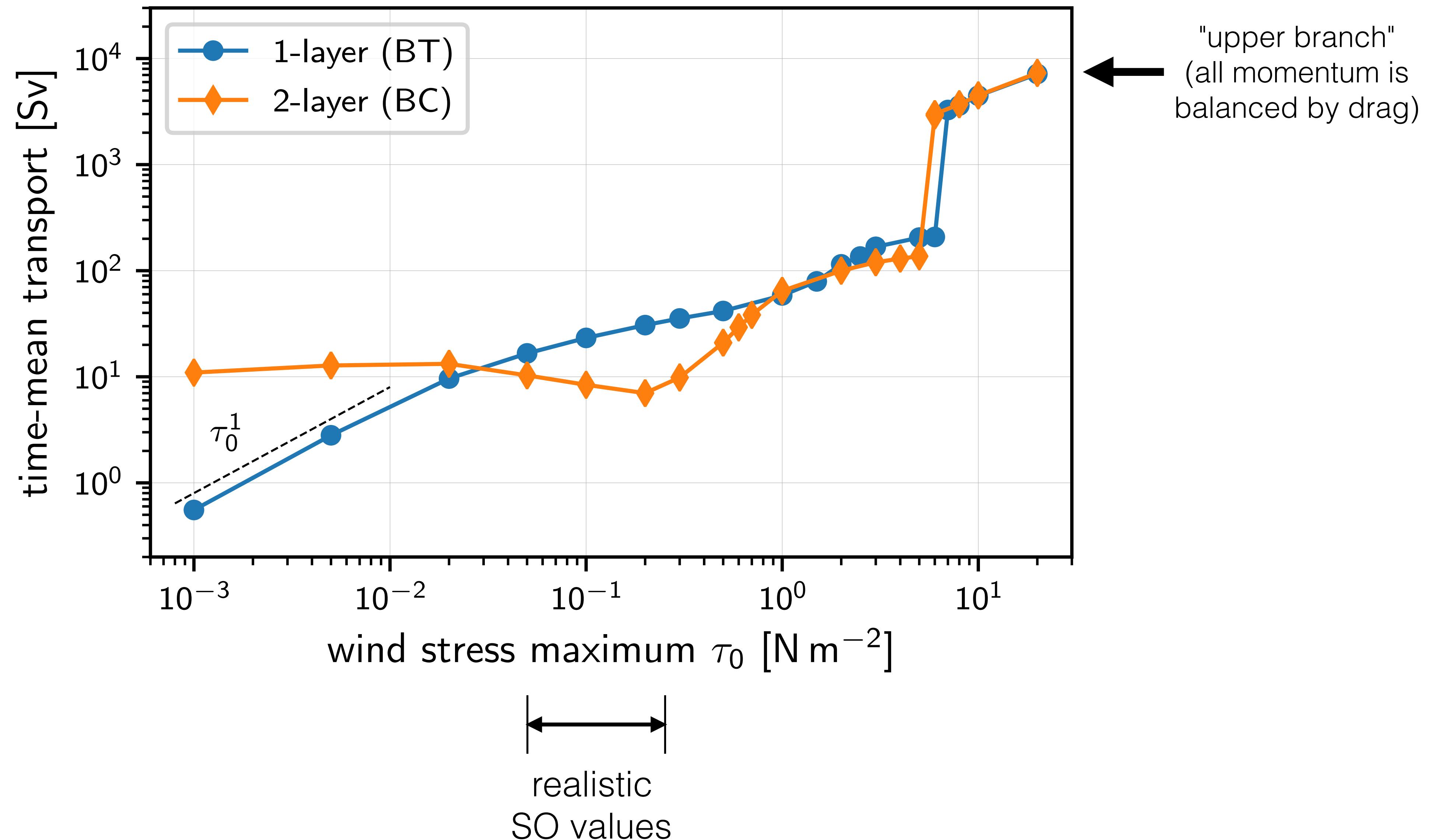
layered approximations

vary the wind stress amplitude  $\tau_0$   
and see how the time-mean zonal transport changes

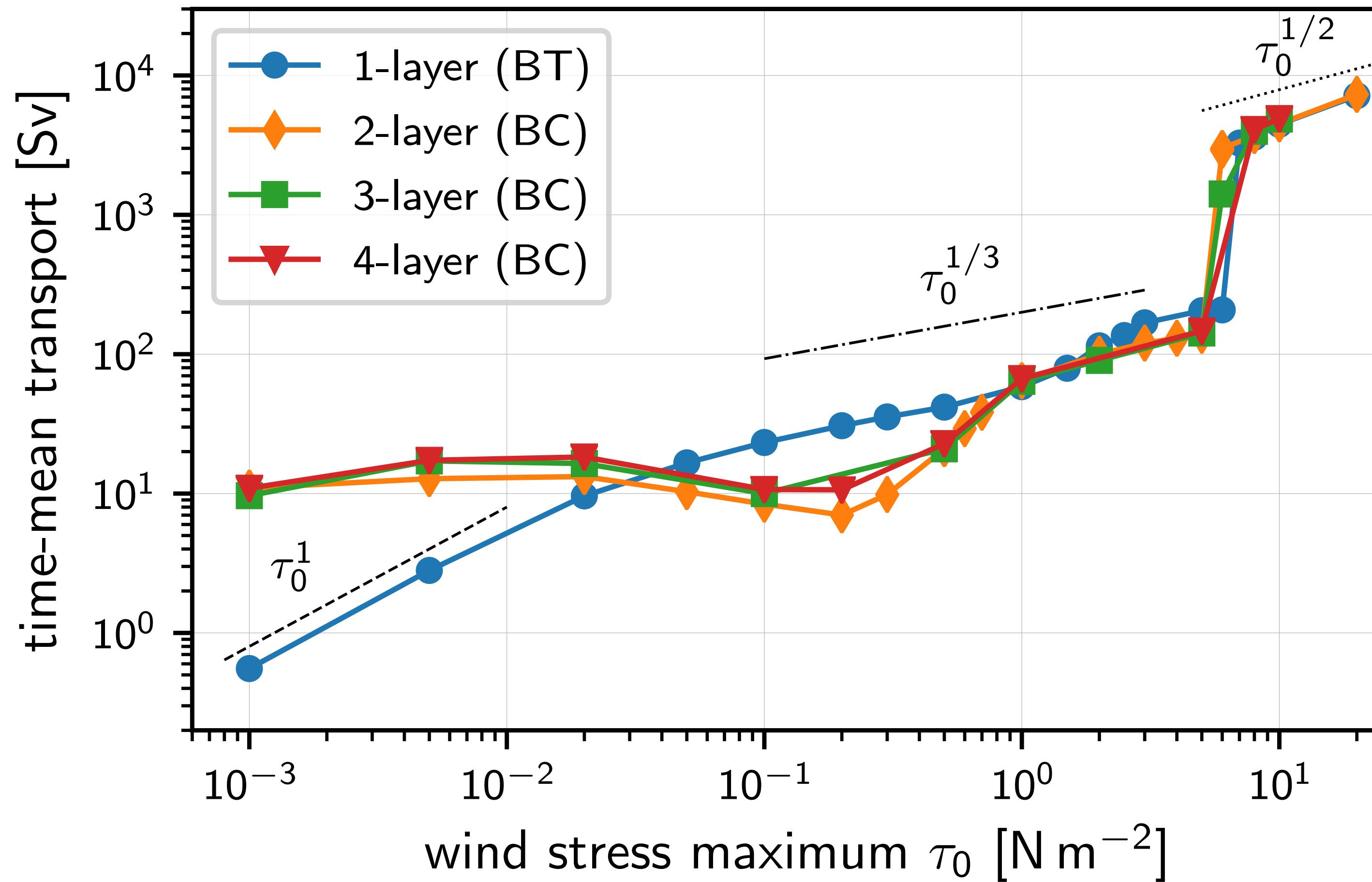
# mean zonal transport Vs wind stress



# mean zonal transport Vs wind stress

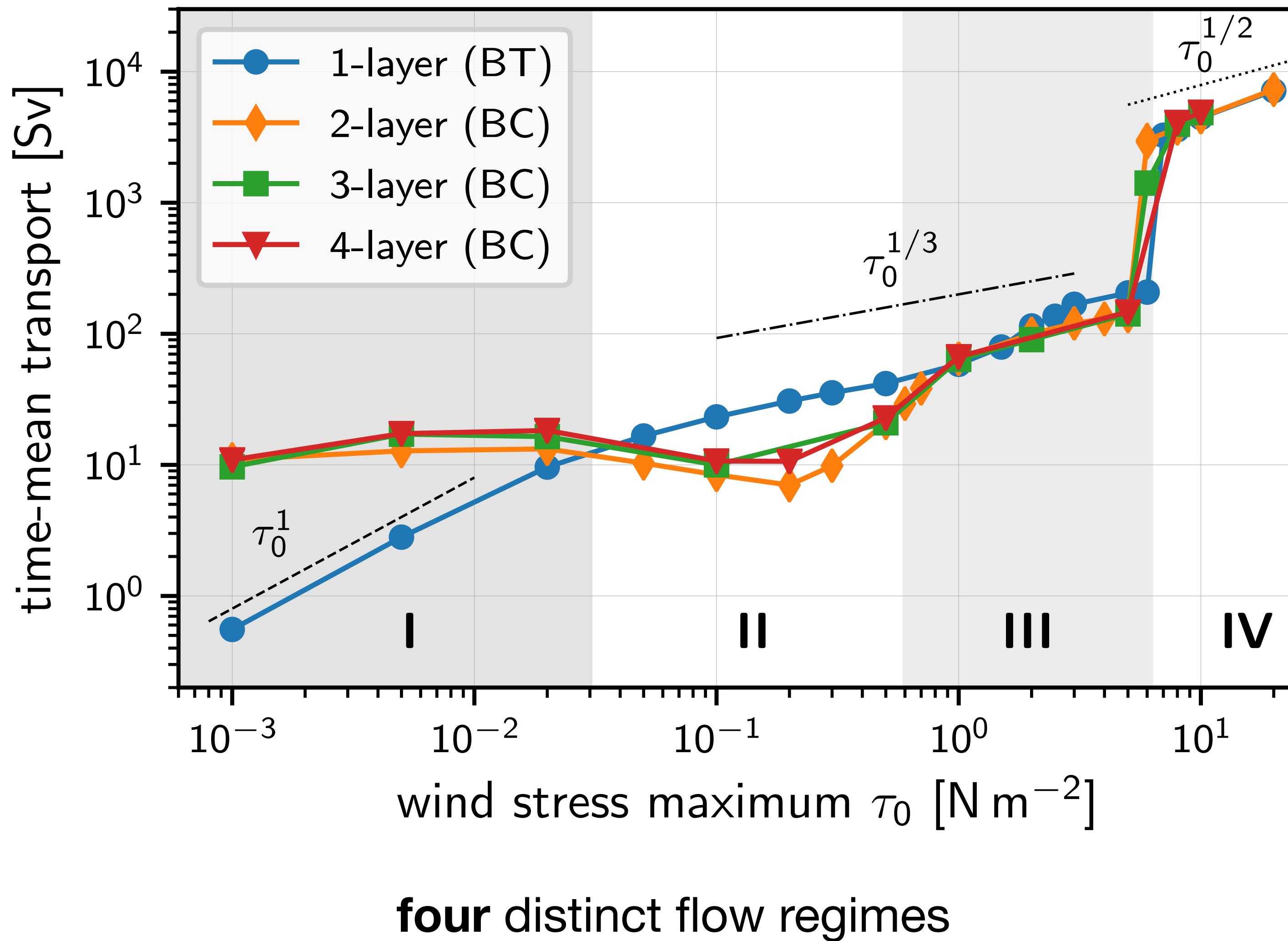


# mean zonal transport Vs wind stress

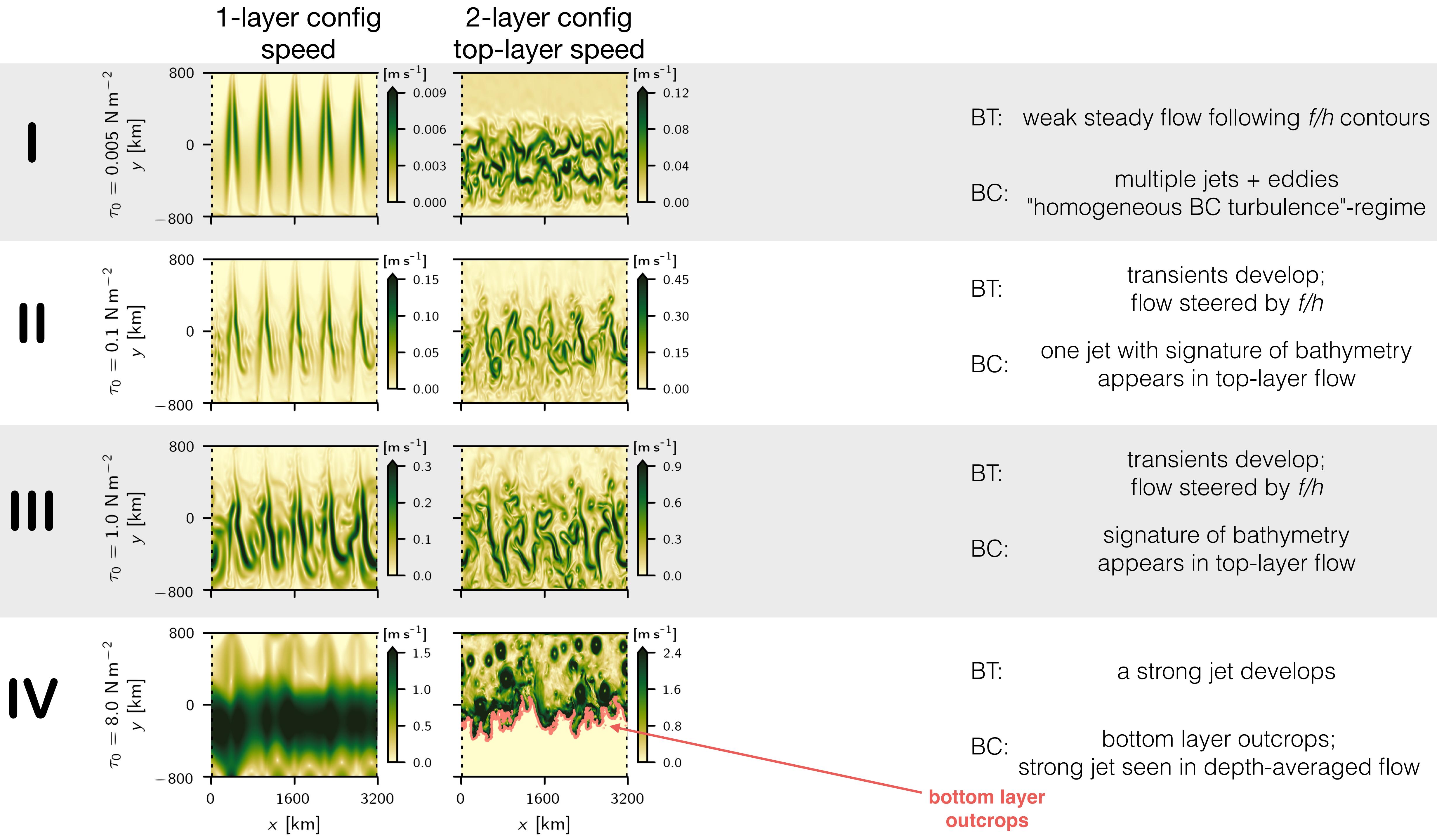


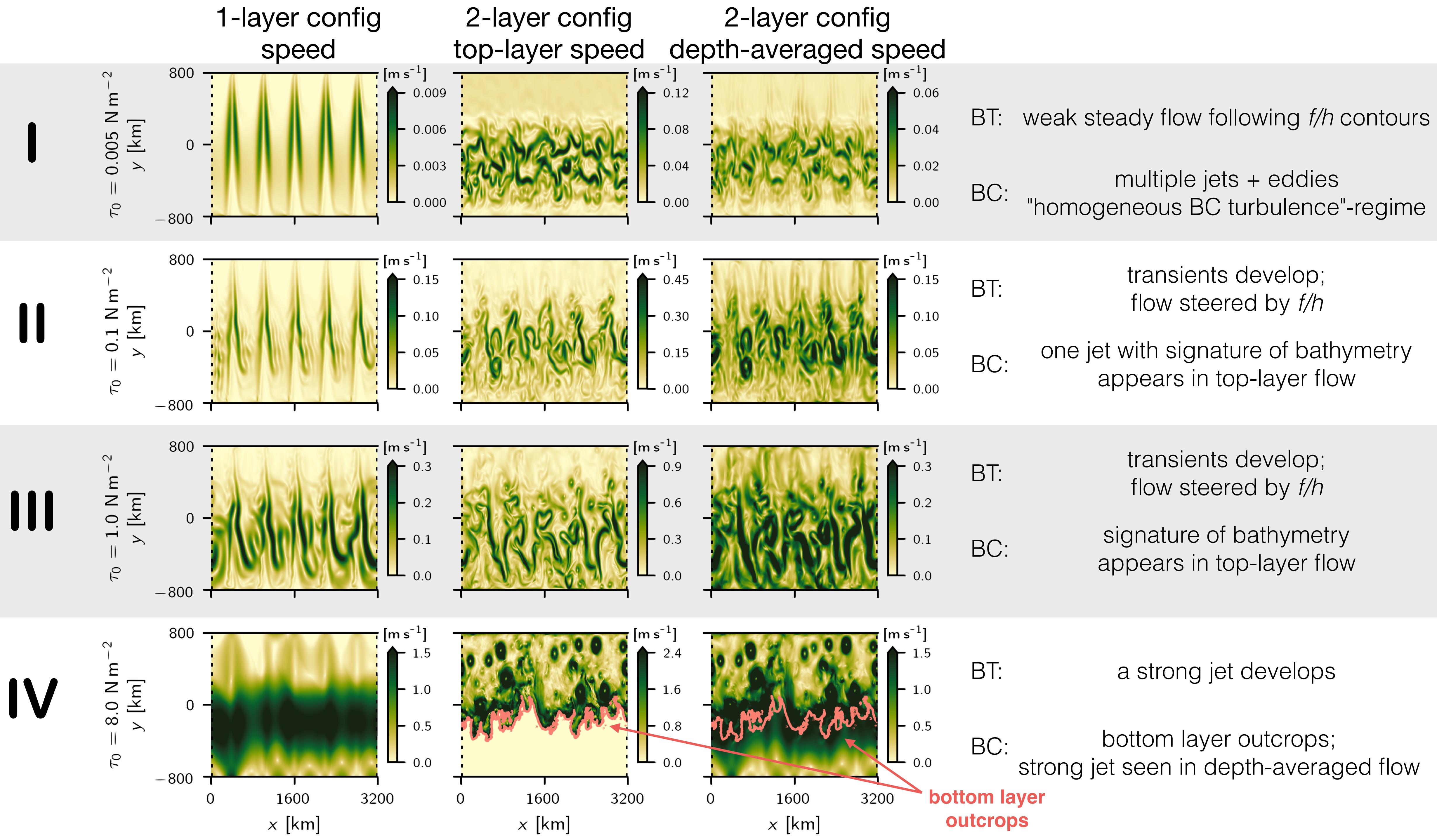
>3-layer configurations are the same as 2-layers  
(as far as the mean zonal transport is concerned)

# mean zonal transport Vs wind stress



how does the flow look like in the four flow regimes?





# depth-integrated zonal momentum balance

$\langle \rangle$  : layer average  
 $\overline{\phantom{x}}$  : time average

$$\langle \tau \rangle = \langle \overline{p_{\text{bot}} \partial_x h_{\text{bot}}} \rangle + \langle \rho_m c_D \overline{u_{\text{bot}} | \mathbf{u}_{\text{bot}} |} \rangle$$

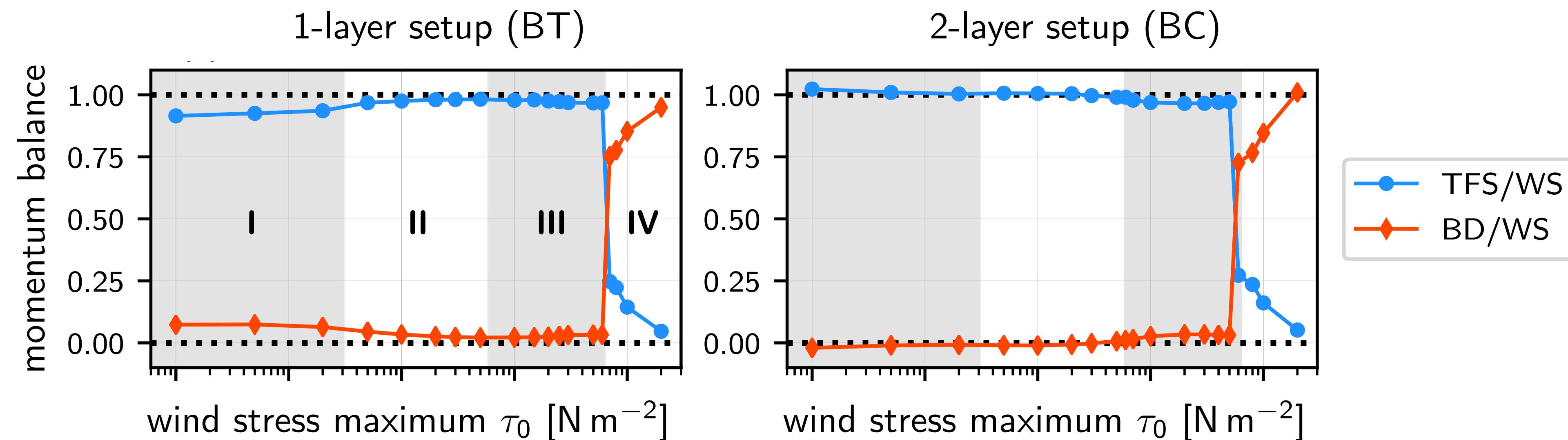
$$\langle \overline{p_{\text{bot}} \partial_x h_{\text{bot}}} \rangle = \langle \overline{p_{\text{bot}}} \partial_x h_{\text{bot}} \rangle$$

only standing flow  
contributes to TFS

wind  
stress  
(WS)

topographic  
form stress  
(TFS)

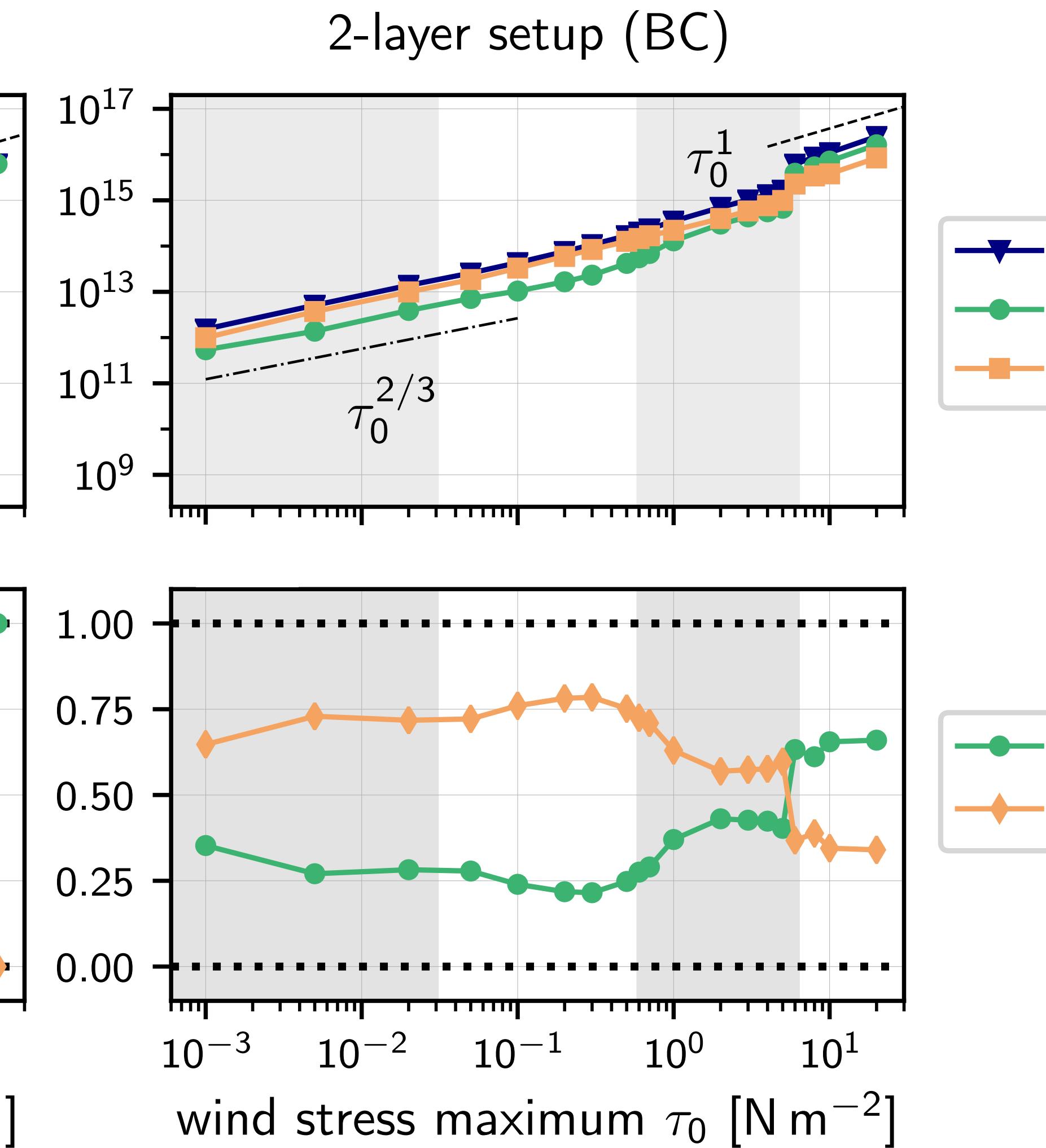
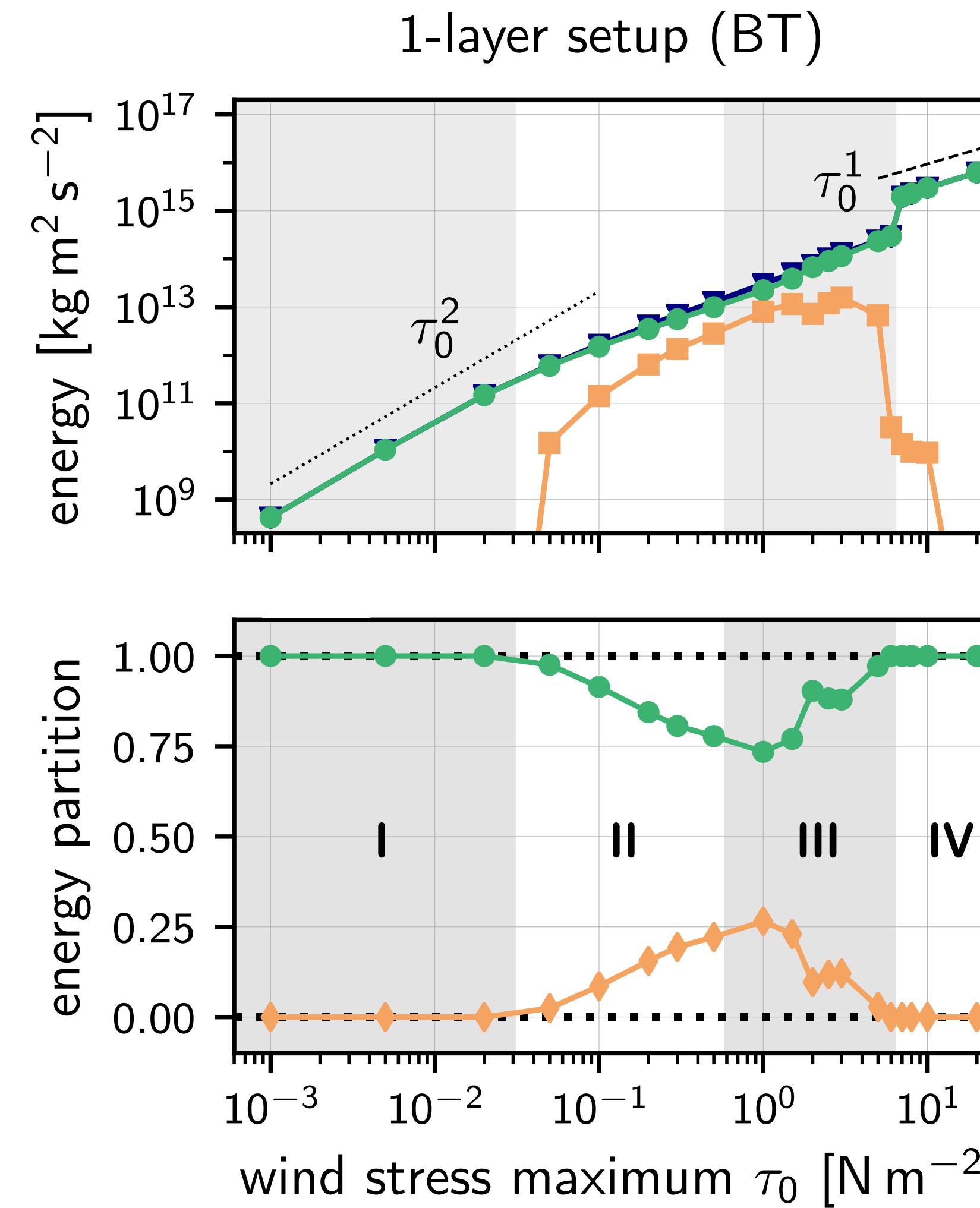
bottom drag  
(BD)



Almost all momentum is balanced by topographic form stress  
(except when flow transitions to "upper branch").

# standing–transient kinetic energy decomposition

BT config  
has transients  
only in **II & III**



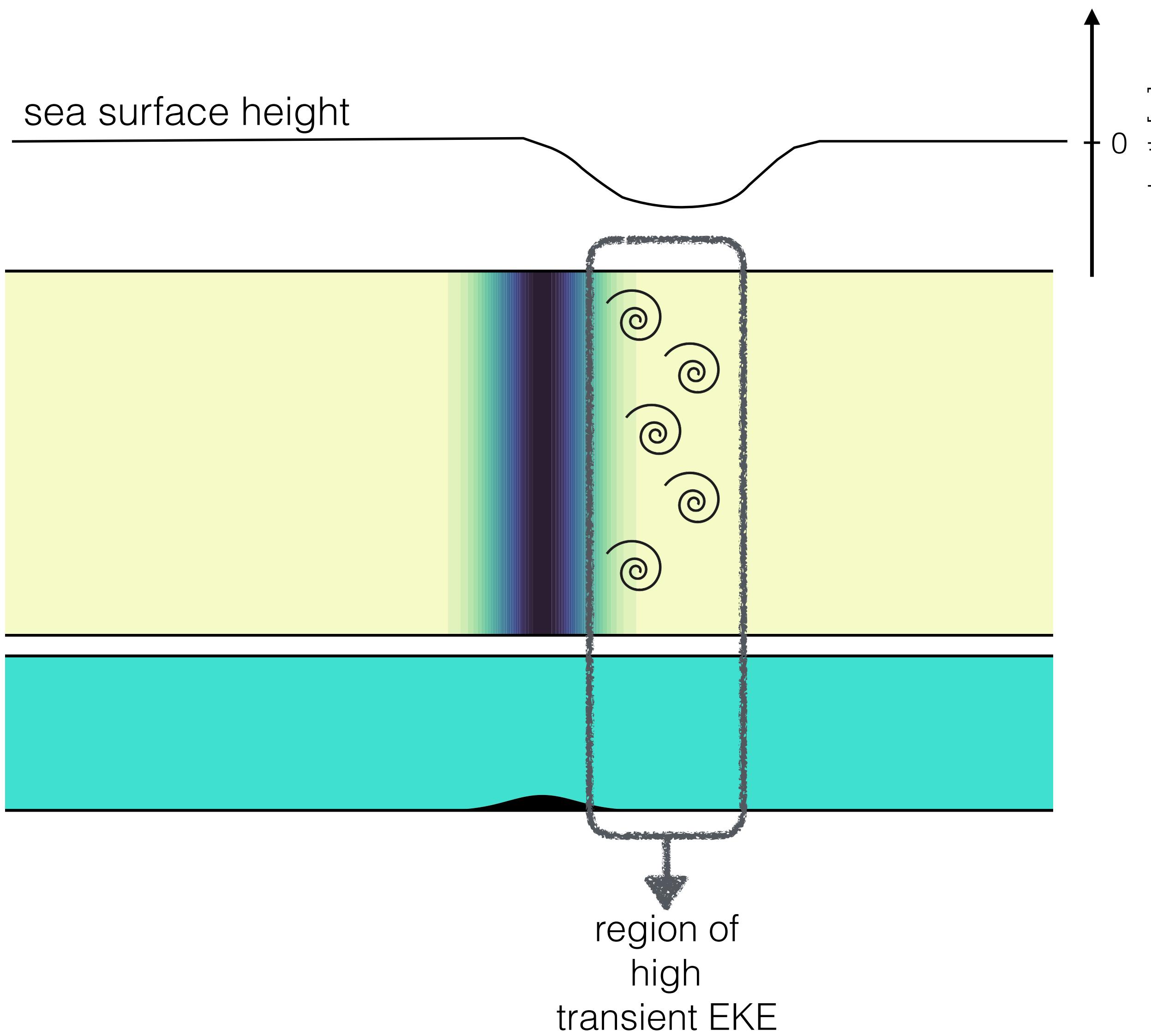
Despite the great differences in flow fields,  
both **BT** and **BC** configs show same mean zonal transport for regimes **III & IV**.

$$\langle \overline{p_{\text{bot}} \partial_x h_{\text{bot}}} \rangle = \langle \overline{p_{\text{bot}}} \partial_x h_{\text{bot}} \rangle$$

only standing flow contributes to  
topographic form stress

how transients affect  
topographic form stress?

# how transients lead to time-mean topographic form stress?



[Same process as described by Youngs et al. 2017]

# take home messages

when transient eddies exist (both in **barotropic** or **baroclinic** configs)  
the mean zonal transport becomes eddy saturated  
[transport is much less sensitive to wind stress increase]

**proposal:**

eddy saturation occurs due to  
transient eddies shaping the standing flow  
to produce topographic form stress that balances the wind stress  
(*regardless* of the process from which transient eddies originate)

our results show that the (oftentimes ignored) barotropic flow-component  
plays an important role in setting up the ACC transport

[in agreement with recent obs. evidence, e.g., Thompson & Naveira Garabato 2014,  
Peña-Molino et al. 2014, Donohue et al. 2016 (cDrake exp)]

*thank you*