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1 Phenomenology of decaying turbulence beneath 2 surface waves

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8 This paper explores decaying turbulence beneath surface waves that is initially isotropic and
9 shear-free. We start by presenting phenomenology revealed by wave-averaged numerical
10 simulations: an accumulation of angular momentum in coherent vortices perpendicular
11 to the direction of wave propagation, suppression of kinetic energy dissipation, and the
12 development of depth-alternating jets. We interpret these features through an analogy with
13 rotating turbulence (Holm 1996), wherein the curl of the Stokes drift, $\nabla \times \mathbf{u}^S$, takes on the
14 role of the background vorticity (for example, $(f_0 + \beta y)\hat{z}$ on the beta plane). We pursue
15 this thread further by showing that a two-equation model proposed by Bardina *et al.* (1985)
16 for rotating turbulence reproduces the simulated evolution of volume-integrated kinetic
17 energy. This success of the two-equation model — which explicitly parametrizes wave-driven
18 suppression of kinetic energy dissipation — carries implications for modeling turbulent
19 mixing in the ocean surface boundary layer. We conclude with a discussion about a wave-
20 averaged analogue of the Rossby number appearing in the two-equation model, which we
21 term the “pseudovorticity number” after the pseudovorticity $\nabla \times \mathbf{u}^S$. The pseudovorticity
22 number is related to the Langmuir number in an integral sense.

23 1. Introduction

24 Surface waves enhance near-surface ocean turbulent mixing, especially in summertime and
25 in the tropics where boundary layers are often shallow, sunny, windy, and wavy (Sullivan &
26 McWilliams 2010). Turbulence driven by surface wind stress and affected by surface waves
27 is usually called “Langmuir turbulence” (McWilliams *et al.* 1997), implying a connection
28 between wave-catalyzed turbulent coherent structures and the structure of a laminar, wave-
29 catalyzed shear instability that Craik & Leibovich (1976) called “Langmuir circulation”.

30 In this paper we investigate decaying turbulence beneath surface waves using numerical
31 simulations of the wave-averaged Navier–Stokes equations (Craik & Leibovich 1976).
32 Decaying turbulence is fundamental (Batchelor 1953) but has seen little attention beneath
33 surface waves. Most work on wave-modified turbulence involves strong ambient shear and
34 also invokes surface forcing by winds and buoyancy fluxes (McWilliams *et al.* 1997; Polton
35 *et al.* 2005; Harcourt & D’Asaro 2008; Van Roekel *et al.* 2012; Large *et al.* 2019; Fan *et al.*

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36 We show that some of the essential phenomenology of turbulence beneath surface
 37 waves is revealed by focusing on the evolution of unforced, initially-shear-free turbulence.

38 We interpret the results of these simulations in section 2, leveraging an analogy between
 39 wave-averaged and rotating dynamics first proposed by Holm (1996). Like rotation, surface
 40 waves *catalyze* a transfer of energy from smaller to larger scales (hereby referred to as
 41 ‘inverse cascade’) within the turbulent flow beneath the waves — without exchanging energy
 42 with the flow — and cause angular momentum to accumulate in coherent vortices, while
 43 suppressing kinetic energy dissipation. This contrasts our case with the interaction between
 44 balanced flows and internal waves, wherein non-catalytic interactions can occur even for
 45 steady wave fields (Thomas 2023). Moreover, unlike in shear-driven Langmuir turbulence
 46 where coherent structures tend to align with the direction of the Lagrangian-mean shear
 47 (Sullivan & McWilliams 2010; Van Roekel *et al.* 2012), the coherent vortices that we
 48 observe forming in the absence of shear are instead oriented *perpendicular to the direction*
 49 *of surface wave propagation*. We also observe the development of alternating jets, as in
 50 beta-plane turbulence.

51 In section 3 we further the analogy by adapting a phenomenological model proposed by
 52 Bardina *et al.* (1985) for decaying rotating turbulence to the wave-averaged case, showing that
 53 the model also describes the decay of volume-averaged kinetic energy beneath surface waves.
 54 This model suggests a new way to incorporate wave effects into two-equation models (for
 55 example, Harcourt (2015)) by explicitly representing the surface-wave-driven suppression of
 56 kinetic energy dissipation.

57 We conclude in section 4 by proposing a new non-dimensional number to characterize the
 58 importance of surface waves for the evolution of turbulent flows analogous to the Rossby
 59 number. We show how this new number may be related to the Langmuir number for wind-
 60 forced cases, while also generalizing to purely convective or decaying situations.

61 2. Simulations of decaying turbulence beneath surface waves

62 The incompressible, inviscid momentum equation in the presence of a background vortic-
 63 ity Ω ,

$$64 \quad \partial_t \mathbf{u} + (\mathbf{u} \cdot \nabla) \mathbf{u} + \Omega \times \mathbf{u} + \nabla p = 0, \quad \text{with} \quad \nabla \cdot \mathbf{u} = 0, \quad (2.1)$$

65 describes *both* rotating flows without surface waves and *also* wave-averaged flows beneath
 66 steady, horizontally-uniform surface-wave fields (see for example Craik & Leibovich 1976;
 67 Suzuki & Fox-Kemper 2016; Holm 1996; Wagner *et al.* 2021). In the wave-averaged
 68 case, p and \mathbf{u} in (2.1) represent the Eulerian-mean kinematic pressure and Lagrangian-mean
 69 velocity, respectively. The Lagrangian-mean velocity is the “total” velocity responsible for
 70 the advection of mass, momentum, tracers, and particles, and may be decomposed into
 71 an Eulerian-mean component plus “Stokes drift” correction (van den Bremer & Breivik
 72 2018; Vanneste & Young 2022, see for example). The wave-averaged momentum equation
 73 may be derived by asymptotic expansion leveraging a time-scale separation between rapidly
 74 oscillating surface waves and the slower evolution of \mathbf{u} (Craik & Leibovich 1976; Leibovich
 75 1980; Vanneste & Young 2022). The appearance of surface waves in (2.1) as an effective
 76 Coriolis force is discussed by Holm (1996).

77 For rotating flows, the background vorticity Ω in (2.1) takes the form

$$78 \quad \Omega_{\text{rotating}} = f \hat{z}, \quad (2.2)$$

79 where f is the Coriolis parameter and \hat{z} is the axis of rotation. For wave-averaged flows
 80 beneath a uniform surface wave field propagating in the x direction, the background

81 vorticity Ω is alternatively

82
$$\Omega_{\text{waves}} = -\nabla \times \mathbf{u}^S = -\partial_z u^S \hat{\mathbf{y}}, \quad (2.3)$$

83 where $\mathbf{u}^S = u^S(z) \hat{\mathbf{x}}$ is the surface wave Stokes drift (Craik & Leibovich 1976; Suzuki &
84 Fox-Kemper 2016; van den Bremer & Breivik 2018; Vanneste & Young 2022). We refer to
85 $\nabla \times \mathbf{u}^S$ — the curl of the surface-wave pseudomomentum \mathbf{u}^S (Andrews & McIntyre 1978)
86 — as the “pseudovorticity”. In the context of equation (2.1), the main difference between
87 rotating turbulence and turbulence beneath steady surface waves is the spatial structure of
88 $\partial_z u^S$: in shallow water $\partial_z u^S \sim z$ to leading-order, while $\partial_z u^S$ decays exponentially for deep
89 water waves.

90 *2.1. Coherent structures in rotating and wave-averaged turbulence*

91 To illustrate the similarity of rotating turbulence and turbulence beneath surface waves we
92 conduct large eddy simulations of equation (2.1) with implicit grid-scale dissipation in a unit
93 cube with three choices for Ω ,

94
$$\Omega_{\text{rotating}} = \frac{1}{4} \hat{\mathbf{z}}, \quad \Omega_{\text{waves}} = -\frac{1}{2} z \hat{\mathbf{y}}, \quad \text{and} \quad \Omega_{\text{isotropic}} = 0. \quad (2.4)$$

95 Above, $\frac{1}{2} z$ is the pseudovorticity associated with a shallow water wave in $z \in [0, 1]$ with
96 Stokes drift $\mathbf{u}^S \approx \frac{1}{2}(1 + \frac{1}{2}z^2) \hat{\mathbf{x}}$. The domain is horizontally-periodic in x, y , and we use
97 free-slip boundary conditions at the top and bottom boundaries in z . Implicit dissipation of
98 grid-scale kinetic energy is provided by an upwind-biased, nominally 9th-order Weighted,
99 Essentially Non-Oscillatory advection scheme (Pressel *et al.* 2017; Shu 2020; Silvestri
100 *et al.* 2024). Our implicit dissipation method allows us to simulate decaying turbulence
101 that undergoes a transition from isotropic to nearly two-dimensional dynamics.

102 Following previous isotropic homogeneous turbulence studies (e.g., Orszag & Patterson
103 (1972)), we produce an initial condition for the three simulations by conducting a preliminary
104 simulation initialized with kinetic energy spectrum

105
$$\frac{1}{2} |\hat{\mathbf{u}}|^2 \sim |\mathbf{K}|^2 \exp \left\{ -2 \left(|\mathbf{K}| / K_i \right)^2 \right\}, \quad (2.5)$$

106 where $\hat{\mathbf{u}}$ is the Fourier transform of \mathbf{u} , \mathbf{K} is the Fourier wavenumber vector, and $K_i = 32 \times 2\pi$
107 is the 32nd wavenumber in the domain. The amplitude of the preliminary initial condition is
108 scaled to produce an initial root-mean-square vorticity

109
$$\sqrt{\int |\omega|^2 dV} \stackrel{\text{def}}{=} \omega_{\text{rms}}(t = -t_0) = 1000, \quad \text{where} \quad \omega \stackrel{\text{def}}{=} \nabla \times \mathbf{u} = \xi \hat{\mathbf{x}} + \eta \hat{\mathbf{y}} + \zeta \hat{\mathbf{z}}. \quad (2.6)$$

110 The initial simulation is run for a duration t_0 until $t = 0$, defined as the time when the
111 mean-square vorticity has decayed to $\omega_{\text{rms}}(t = 0) = 10$. The velocity field is then saved
112 to disk to be used as an initial condition in subsequent runs starting from $t = 0$. The
113 simulations are conducted with Oceananigans (Ramadhan *et al.* 2020; Wagner *et al.* 2025),
114 which discretizes (2.1) with a finite volume method. Scripts that reproduce simulations in
115 this paper are stored on GitHub; see the Data availability statement.

116 We focus first on the evolution of the relative vorticity ω defined in equation (2.6).
117 Figure 1 shows vorticity components for the three cases after $t = 1000$ time units: figure 1(a)
118 shows vertical vorticity ζ , while figures 1(b) and (c) show the horizontal vorticity η .
119 Figure 1(a) and (b) for rotating and wave-affected turbulence, respectively, both exhibit the
120 formation of coherent structures and relatively greater vorticity levels than the unorganized,
121 small amplitude isotropic vorticity in figure 1(c). Figure 2 is similar, except that figure 2(a)
122 shows ζ in rotating turbulence in the xy -plane, while figures 2(b) and 2(c) show η in the xz
123 plane for wave-averaged and isotropic turbulence, respectively.

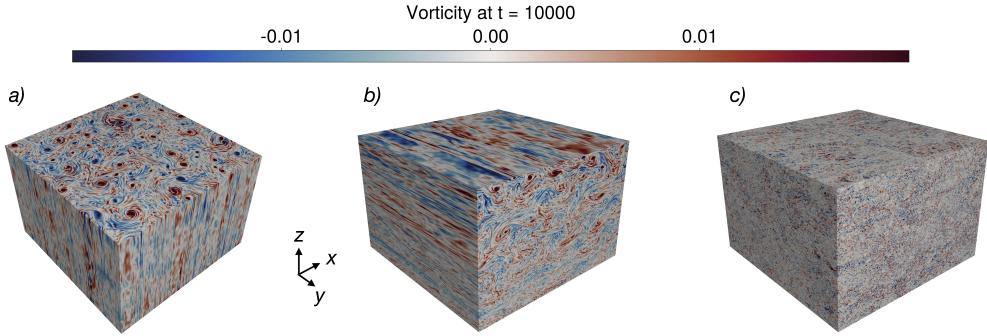


Figure 1: Vorticity in simulations decaying turbulence with 512^3 finite volume cells in a unit cube after 1000 time units. (a) Vertical relative vorticity $\zeta = \hat{z} \cdot (\nabla \times \mathbf{u})$ in rotating turbulence with Coriolis parameter $f = 1/4$, (b) Horizontal relative vorticity $\eta = \hat{y} \cdot (\nabla \times \mathbf{u})$ in turbulence beneath surface waves with Stokes shear $\partial_z u^S = -z/2$, and (c) η in isotropic turbulence.

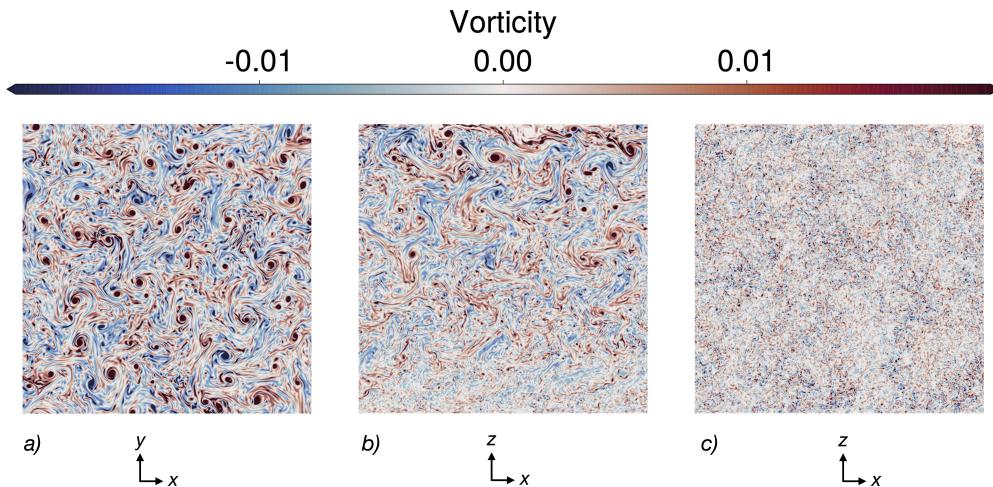


Figure 2: As figure 1 but showing (a) ζ for rotating turbulence in the xy -plane; (b) η for turbulence beneath surface waves in the xz -plane, (c) η for isotropic turbulence in the xz -plane.

124 2.2. *Zonation and the analogy of wave-averaged turbulence with beta-plane turbulence*
 125 Note that turbulence beneath shallow water waves is homeomorphic to β -plane turbulence
 126 Rhines (1975) — with $\Omega_{\beta\text{-plane}} = \beta y \hat{z}$ — modulo on the xz plane instead of on the xy plane
 127 and with the Stokes drift curvature $\partial_z^2 u^S$ playing the role of β . One of the most striking results
 128 of this similarity is the propensity for turbulence beneath surface waves to develop “zonal jets”
 129 — coherent, alternating jets in the direction of surface wave propagation (and perpendicular
 130 to the pseudovorticity direction). This similarity allows us, therefore, to borrow intuition on
 131 structure formation in β -plane turbulence (for example, Huang & Robinson (1998); Farrell
 132 & Ioannou (2007); Srinivasan & Young (2012); Constantinou *et al.* (2014); for an overview
 133 see Constantinou (2015); Farrell & Ioannou (2019); Marston & Tobias (2023).)

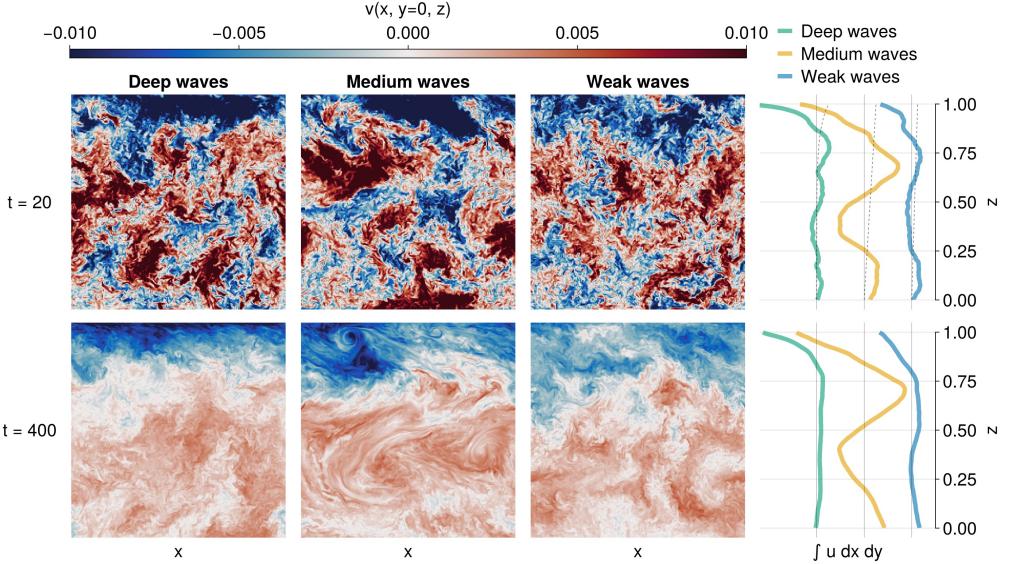


Figure 3: The evolution of cross-wave momentum v in the xz -plane (6 panels on the left) and horizontally-averaged along-wave-momentum u (2 panels on the right) at $t = 40, 400$ and for three wave fields: “deep” (left panel, red lines), “medium” (middle panel, orange lines), and “weak” (right panel, blue lines). Dashed lines in the top right plot show the Stokes drift profile (normalized) for each case. The “medium” waves case uses the background vorticity Ω_{waves} in (2.4). Both the u profiles and the light gray “zero lines” are spaced apart by the scale $\delta u = 10^{-2}$. The u -profiles at right show the development of depth-alternating jets, including a counter-wave surface jet for all cases. The v -slices are similar between the three cases at early times, but at later times exhibit strong, localized wave-impacts in their respective regions of significant Stokes shear. Note that v plays the role that vertical velocity plays in rotating turbulence. Note that the Rhines scale may be estimated as $\ell_R = 2\pi\sqrt{U/|\partial_z^2 u^S|}$. With $|\partial_z^2 u^S| = \frac{1}{2}$ and $U \approx 10^{-2}$ for the medium waves case, for example, we estimate $\ell_R \approx 0.8$, consistent with the jet spacing apparent on the right panels.

134 To illustrate this, we consider two additional cases with background vorticity

135
$$\Omega_{\text{deep}} = -\frac{1}{4}e^{8(z-1)}\hat{y}. \quad \text{and} \quad \Omega_{\text{weak}} = -\frac{1}{8}z\hat{y}. \quad (2.7)$$

136 These and subsequent simulations use 384^3 finite volume cells to reduce their computational
137 expense for the purpose of performing longer simulations to $t = 10^4$.

138 Figure 3 shows time-series of the y -momentum v , which plays the role that vertical
139 velocity plays in rotating turbulence. The right side of figure 3 shows vertical profiles of the
140 horizontally-averaged x -momentum,

141
$$U(z, t) \stackrel{\text{def}}{=} \int u \, dx \, dy. \quad (2.8)$$

142 These U -profiles exhibit the development of depth-alternating jets, which are most promi-
143 nently exhibited by the zig-zagging structure in the medium waves case on the right side
144 of figure 3. To our knowledge, depth-alternating jets have not been observed in shear-free
145 wave-modified turbulence. Comparing the slices of v for medium and weak waves reveals
146 how the medium-strong waves induce a strong inverse cascade, more coherent vortices, and
147 fewer small-scale motions than the weak waves.

148 **3. Phenomenological model for the evolution of kinetic energy**

149 We turn to the evolution of the domain-averaged kinetic energy,

150
$$k(t) \stackrel{\text{def}}{=} \int \frac{1}{2} (u^2 + v^2 + w^2) dV. \quad (3.1)$$

151 Figure 4 plots time-series of the normalized kinetic energy $k(t)/k(t = 0)$ for the three
 152 cases presented in figure 1, as well as three additional cases that use $\Omega = Sz\hat{y}$ with $S =$
 153 $(1/2, 1/8, 1/16)$. Figure 4 illustrates another common feature to rotating turbulence and
 154 turbulence beneath surface waves: a suppression of the kinetic energy dissipation rate,
 155 such that at long times the kinetic energy levels off. We theorize that the suppression of
 156 dissipation is linked to the formation of coherent structures: rather than cascading to the grid
 157 scale, where dissipation can occur, kinetic energy accumulates in much larger-scale coherent
 158 structures. Similar to the dynamics of two-dimensional turbulence McWilliams (1984), the
 159 concentration of kinetic energy in coherent structures suppresses the forward cascade of
 160 kinetic energy.

161 Inspired by Bardina *et al.* (1985), we model the kinetic energy $k(t)$ with the phenomeno-
 162 logical two-equation system

163
$$\frac{d}{dt} k = -\epsilon, \quad (3.2)$$

164
$$\frac{d}{dt} \epsilon = -a \frac{\epsilon^2}{k} - b\Omega\epsilon, \quad (3.3)$$

165 where $\epsilon(t)$ describes the dissipation rate of k , Ω is a characteristic scale for Ω , and a and b
 166 are $O(1)$ free parameters. Equation (3.2) follows directly from $\int \mathbf{u} \cdot (2.1) dV$ after accounting
 167 for the effect of the implicit numerical dissipation in our simulations. (Alternative models
 168 that include surface-wave-associated source terms for turbulent kinetic energy, such as the
 169 one proposed by Axell (2002), are inconsistent with (2.1).) The catalytic nature of rotation
 170 or surface waves may also be deduced from the fact that Ω does not appear in equation (3.2).
 171 The first term in (3.3) models the destruction of ϵ on the turbulent time-scale $\tau = k/\epsilon$, since
 172 $d\epsilon/dt \sim -\epsilon/\tau$ when $\Omega\epsilon \ll \epsilon/\tau$.

173 The second term in (3.3) models the suppression of kinetic energy dissipation — or
 174 alternatively, the growth of the correlation length due to the coalescence of coherent structures
 175 — by the background vorticity Ω . The relative importance of “intrinsic” destruction of ϵ and
 176 background-vorticity-induced destruction of ϵ is measured by the non-dimensional number
 177 $\Omega\tau = \Omega k/\epsilon$, whose significance is revisited in section 4. We note that (3.2) is exact —
 178 and unchanged whether or not the system is rotating or modulated by surface waves. The
 179 modulation of turbulence by surface waves is therefore fundamentally “catalytic”, and can
 180 only affect k indirectly by changing (3.3).

181 The free parameter a may be constrained by considering isotropic turbulence with $\Omega = 0$.
 182 In this case we expect the turbulent kinetic energy to decay according to $k \sim t^{-6/5}$ (Saffman
 183 1967), which implies $\frac{d}{dt} k = -\frac{6}{5} \frac{k}{t}$, and thus in turn, via (3.2), leads to

184
$$\epsilon = \frac{6}{5} \frac{k}{t}, \quad \text{so that} \quad \frac{d}{dt} \epsilon = -\frac{66}{25} \frac{k}{t^2}. \quad (3.4)$$

185 Inserting (3.4) into (3.3) yields $a = 11/6$.

186 From (3.2)–(3.3) we deduce that the rate of change of the combination $\log \epsilon - a \log k$ is
 187 proportional to b , that is,

188
$$\frac{d}{dt} (\log \epsilon - a \log k) = -b\Omega. \quad (3.5)$$

189 Integrating (3.5) produces

$$190 \quad \frac{\epsilon}{\epsilon_0} = \left(\frac{k}{k_0} \right)^a e^{-b\Omega t}, \quad (3.6)$$

191 where ϵ_0 and k_0 are the dissipation and kinetic energy at $t = 0$ respectively. If we insert the
192 expression for the dissipation $\epsilon(t)$ from (3.6) into (3.2) and integrate in time, we obtain

$$193 \quad \frac{k}{k_0} = \left[1 + \frac{1}{nb\Omega} \frac{\epsilon_0}{k_0} \left(1 - e^{-b\Omega t} \right) \right]^{-n}, \quad (3.7)$$

194 where $n = 1/(a - 1) = 6/5$. (See also equation (10) in *Bardina et al.* (1985).) Taking the
195 limit $\Omega \rightarrow 0$ (or $t \rightarrow 0$ with finite Ω), we obtain the corresponding solution for isotropic
196 turbulence,

$$197 \quad k_{\text{isotropic}}(t) = k_0 \left(1 + \frac{1}{n} \frac{\epsilon_0}{k_0} t \right)^{-n}, \quad (3.8)$$

198 which yields the expected power law $k \sim t^{-n}$ when $\epsilon_0 t / nk_0 \gg 1$.

199 At long times the isotropic and vortical solutions diverge: $k_{\text{isotropic}} \rightarrow 0$ as $t \rightarrow \infty$, while
200 in the vortical case k/k_∞ limits to the constant

$$201 \quad \frac{k_\infty}{k_0} = \left(1 + \frac{\epsilon_0}{nb\Omega k_0} \right)^{-n}, \quad (3.9)$$

202 where $k_\infty = \lim_{t \rightarrow \infty} k(t)$. Equation (3.9) yields a formula for b in terms of k_∞ ,

$$203 \quad b\Omega = \frac{\epsilon_0}{nk_0} \frac{1}{\left(\frac{k_\infty}{k_0} \right)^{1/n} - 1}, \quad (3.10)$$

204 which we use to diagnose b from our numerical simulations.

205 The dashed curves in figure 4 show solutions to the two-equation system in (3.2)–(3.3)
206 that correspond to the solid-line simulated results. Figure 4 shows that a single value of
207 $b = 0.036$ qualitatively describes the evolution of kinetic energy beneath surface waves for
208 a wide range of background vorticity magnitudes Ω , where we estimate Ω via the z -average
209 of $\partial_z u^S$, that is $\Omega \approx \int \partial_z u^S dz$. For the rotating case, we use $b = 0.033$ and $\Omega = 1/4$. While
210 qualitatively excellent considering that $b \approx 0.036$ describes a wide range of conditions,
211 we also find that the phenomenological model overestimates the dissipation rate — and
212 therefore underestimates the kinetic energy k — during the transition between isotropic and
213 background-vorticity-dominated regimes.

214 4. Discussion

215 In this paper we point out the similarity between rotating turbulence on the beta-plane and
216 turbulence beneath surface waves. In particular, turbulence beneath surface waves exhibits
217 the formation of coherent structures perpendicular to the direction of wave propagation, the
218 development of zonal jets, and the suppression of kinetic energy dissipation. These features
219 are consistent with known properties of turbulence beneath surface waves, but the connection
220 with rotating turbulence is obscured in studies that also involve surface wind stress, ambient
221 Lagrangian-mean shear, and surface forcing. In particular, we hypothesize that some of the
222 generic features of Langmuir turbulence — such as the formation of coherent structures,
223 a suppression of kinetic energy dissipation, and an increase in mixing — are driven by
224 similar dynamics as the phenomenon observed in our decaying situations. At the same time,
225 however, the presence of Lagrangian-mean shear produces additional and distinctly different

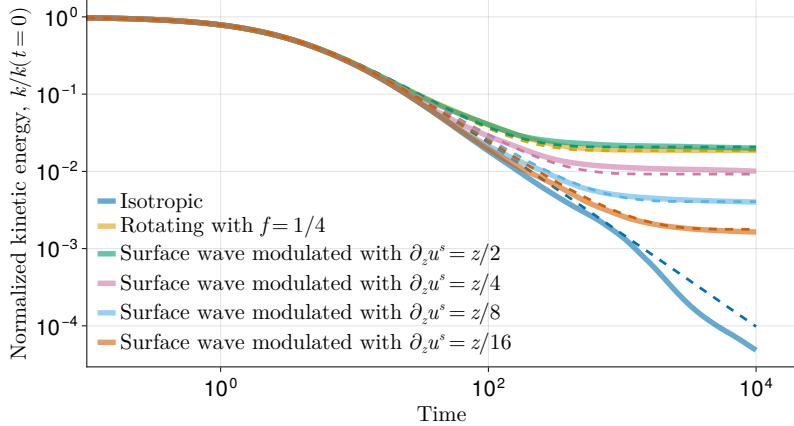


Figure 4: The decay of kinetic energy $k(t)$ in isotropic, rotating, and surface-wave-modulated turbulence. Solid lines show kinetic energy normalized by its initial value, k/k_0 , computed from large eddy simulations. Dashed lines show k/k_0 given by (3.7), which solves the phenomenological two-equation system in (3.2)–(3.3). The initial dissipation rate ϵ_0 in (3.7) is estimated from the numerical solution; specifically we evaluate (3.8) at Δt , substitute $k(t = \Delta t)$ from the numerical solution, and solve for ϵ_0 . We use $a = 11/6$ and estimate b from (3.10). For the rotating case, $\Omega = 1/4$ yields $b = 0.033$.

For the four surface-wave-modulated cases we use $b = 0.036$, where

$$\Omega \stackrel{\text{def}}{=} \int \partial_z u^S dz = (1/4, 1/8, 1/16, 1/32). \text{ Note that using } a = 1.75 \text{ along with commensurate adjustments to } b \text{ matches the simulation data even more closely.}$$

phenomenon, most clearly in the alignment of coherent vortices with the Lagrangian-mean shear, rather than perpendicular to the direction of wave propagation.

We exploit the connection to rotating turbulence by adapting a phenomenological two-equation model proposed by Bardina *et al.* (1985). In this two-equation model, the evolution of dissipation is affected by two terms: one “classical” term producing power-law decay of kinetic energy, and a second term that describes the suppression of kinetic energy dissipation by the presence of surface waves, which in turn effectively enhances kinetic energy levels and turbulent mixing relative to pure isotropic turbulence.

This phenomenological model hints at a new way to understand how surface waves enhance turbulent mixing. In the paradigm proposed by McWilliams *et al.* (1997), the effect of surface waves on turbulence is associated with the “Stokes contribution” to shear production in the turbulent kinetic energy budget. However, our results show that this interpretation must be incomplete, because surface waves also control the evolution of initially shear-free flows.

We therefore propose an alternative theory which supposes that the impact of surface waves may be instead linked to their tendency to catalyze, without exchanging energy, an increase in the correlation times and length scales of turbulent motions. A benefit to our alternative interpretation is that turbulent shear production can be interpreted in the standard way: as a transfer of kinetic energy from the horizontally-averaged, Lagrangian-mean velocity, whose energy is otherwise conserved. In other words, our descriptive analysis leads to an alternative paradigm for wave-modified turbulence wherein shear production is “unchanged” relative to wave-free turbulence (as in rotating turbulence). Instead of modifying shear production directly, then, waves impact both mixing and kinetic energy by inducing an inverse cascade, increasing the turbulent mixing length, and suppressing kinetic energy dissipation.

249 4.1. *The “pseudovorticity” number*

250 The analogy with rotating turbulence inspires the definition of a new non-dimensional
 251 number for characterizing surface wave effects on turbulence. In rotating turbulence, the
 252 Rossby number measures the relative magnitude of the relative vorticity $\nabla \times \mathbf{u}$ and the
 253 background vorticity, $f \hat{\mathbf{z}}$, such that

254
$$\text{Ro} \stackrel{\text{def}}{=} \frac{|\nabla \times \mathbf{u}|}{f} \sim \frac{U}{fL}, \quad (4.1)$$

255 where U is a characteristic horizontal velocity scale and L is a characteristic turbulent
 256 horizontal scale. We propose the analogous “pseudovorticity number” for boundary layer
 257 turbulence,

258
$$\text{Ps} \stackrel{\text{def}}{=} \frac{|\nabla \times \mathbf{u}|}{|\partial_z \mathbf{u}^S|} \sim \frac{W}{\Omega H}, \quad (4.2)$$

259 where W is a turbulence velocity scale, Ω is the magnitude of the Stokes drift shear, and H
 260 is a turbulence length scale. In boundary layer turbulence, for example, W and H are most
 261 straightforwardly measured by the magnitude of the vertical velocity and the boundary layer
 262 depth.

263 The pseudovorticity number Ps measures the average role of surface waves with Stokes
 264 shear magnitude Ω on turbulent motions with turbulent velocity scale W and turbulent length
 265 scale H . In the two-equation model in (3.2)–(3.3), $\text{Ps} \sim \epsilon/k\Omega$. In our decaying scenarios, Ps
 266 eventually vanishes at $t \rightarrow \infty$ and dissipation is completely suppressed.

267 4.2. *Connection with the Langmuir number*

268 The pseudovorticity number is connected between the Langmuir number — the usual
 269 non-dimensional quantity used to characterize the effect of waves on turbulence — when
 270 considering an estimate of Ps integrated over deep boundary layers. One definition of the
 271 Langmuir number is (McWilliams *et al.* 1997),

272
$$\text{La} \stackrel{\text{def}}{=} \sqrt{\frac{u_\star}{u^S(z=0)}}, \quad (4.3)$$

273 where u_\star is the friction velocity, which is the square root of the kinematic wind stress. At
 274 face value, La does not apply to decaying cases, but we can amend this by interpreting u_\star
 275 more generally as a turbulent velocity scale.

276 To connect La and Ps , we consider a “bulk” estimate of Ps over a wind-forced boundary
 277 layer of depth H , where the turbulent velocity scale is $W = u_\star$ and the turbulent length scale
 278 is H . If $u^S(z = -H)$ is negligible, then estimating the pseudovorticity scale Ω as the average
 279 value over the boundary layer yields

280
$$\Omega \sim \frac{1}{H} \int_{-H}^0 \partial_z u^S dz \sim \frac{u^S(z=0)}{H}, \quad (4.4)$$

281 such that

282
$$\text{bulk Ps} = \frac{W}{\Omega H} \sim \frac{u_\star}{u^S(z=0)} \sim \text{La}^2. \quad (4.5)$$

283 We thus find that La^2 may be regarded as a bulk estimate of the more locally-applicable Ps ,
 284 computed over the depth of the boundary layer, H . The difference between the bulk
 285 estimate La and the more specific estimate Ps resolves a paradox associated with La in
 286 that it depends on u^S , despite that only $\partial_z u^S$ appears in (2.1).

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294 **Data availability statement.** Scripts to reproduce the simulations and analyses in this of this study are
 295 openly available at the GitHub repository github.com/glwagner/WaveAveragedDecayingTurbulence.

296 Simulation output used to produced the figures in this paper will be uploaded in a Zenodo repository
 297 upon acceptance.

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