

MTMW14: Numerical Modelling of Atmosphere and Oceans

Project 1: Recharge Oscillator Model for ENSO

Introduction

A Recharge Oscillator Model (ROM) is given by Jin (1997b) for modelling Seas Surface Temperature (SST) and thermocline depth anomalies caused by the El Nino Southern Oscillation (ENSO) in the Pacific ocean along the equator, as-

$$\frac{dh_w}{dt} = -rh_w - \alpha bT_E - \alpha\xi_1 \quad (1)$$

$$\frac{dT_E}{dt} = RT_E + \gamma h_w - e_n(h_1 + bT_E)^3 + \gamma\xi_1 + \xi_2, \quad (2)$$

Here T_w is the SST anomaly and h_w is the thermocline depth anomaly. I have used the 4th order Runge-Kutta time scheme to implement the above ROM. All temperature, depth and time values are non-dimensionalised before inputting into the above model. They are dimensionalised again before making the plots.

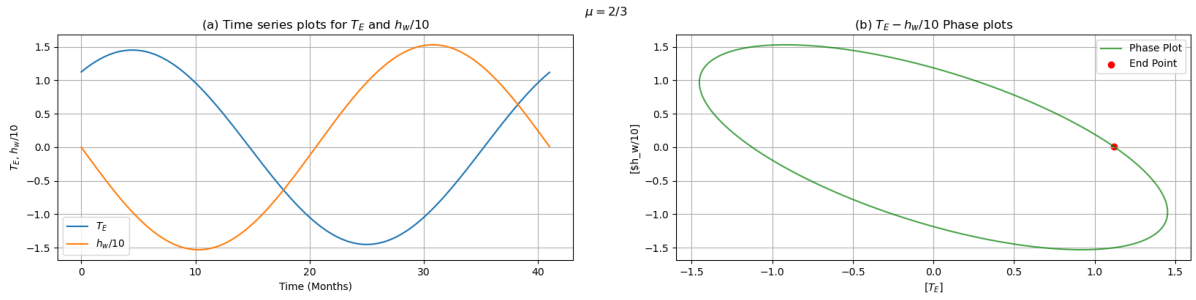
Runge-Kutta Time Scheme

Runge-Kutta is a 4th order iterative scheme. It is a highly accurate scheme and allows the use of longer time steps, but is computationally expensive. Since in this project we will use a range of time-steps, but is not complicated enough to demand high computational cost, I have chosen to use this scheme. As seen in Task A, it is also highly stable. See Appendix for detailed formulation and implementation.

```
In [1]: import numpy as np
import matplotlib.pyplot as plt
from Constants import *
from funcs import *
```

Task A: The neutral linear (deterministic) ROM

```
In [2]: endTime=(2*np.pi/np.sqrt(3/32)) #dimensionalised
nt=1000
dt=endTime/nt
mu=mu_c
E1,E2=0,0
en=0
time=np.linspace(0,endTime,nt) *t_nd # time array, dimensionalised
#plots
Te,hw=RKplots(time,'$\mu=2/3$',nt,dt,T_nd,h_nd,t_nd,E1,E2,mu_c,0,en)
```

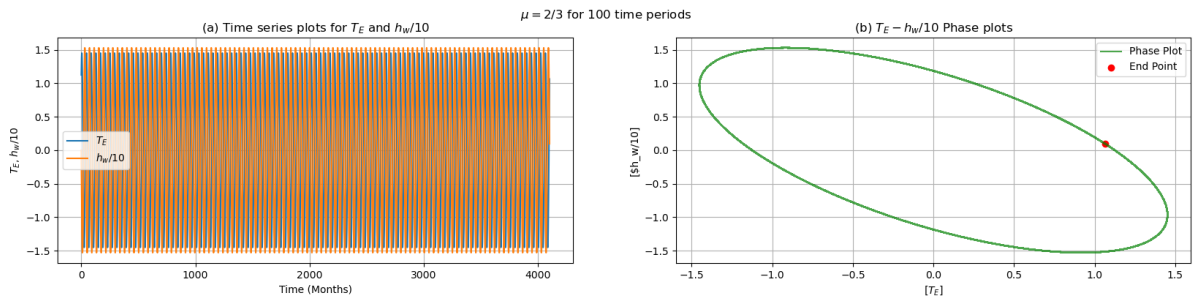


We can see from the time series plot that T_E and h_w are periodic when the coupling coefficient, $\mu = \mu_c = 2/3$ and we have no external forcings ($\xi_1, \xi_2 = 0$) or non-linearity ($e_n = 0$). This is also evident from the $T_E - h_w$ phase plot. The initial conditions are $T = 1.125K$ and $h = 0m$. These initial conditions are used for all parts in this project. When oscillations are created by giving initial perturbation only to T_E in this way, h_w oscillations lag behind. As stated by Jin (1997b), there is no growth in either anomalies and they oscillate uniformly. The period of oscillation plotted above is $\approx 41months$, which is consistent with Jin (1997a).

Stability analysis

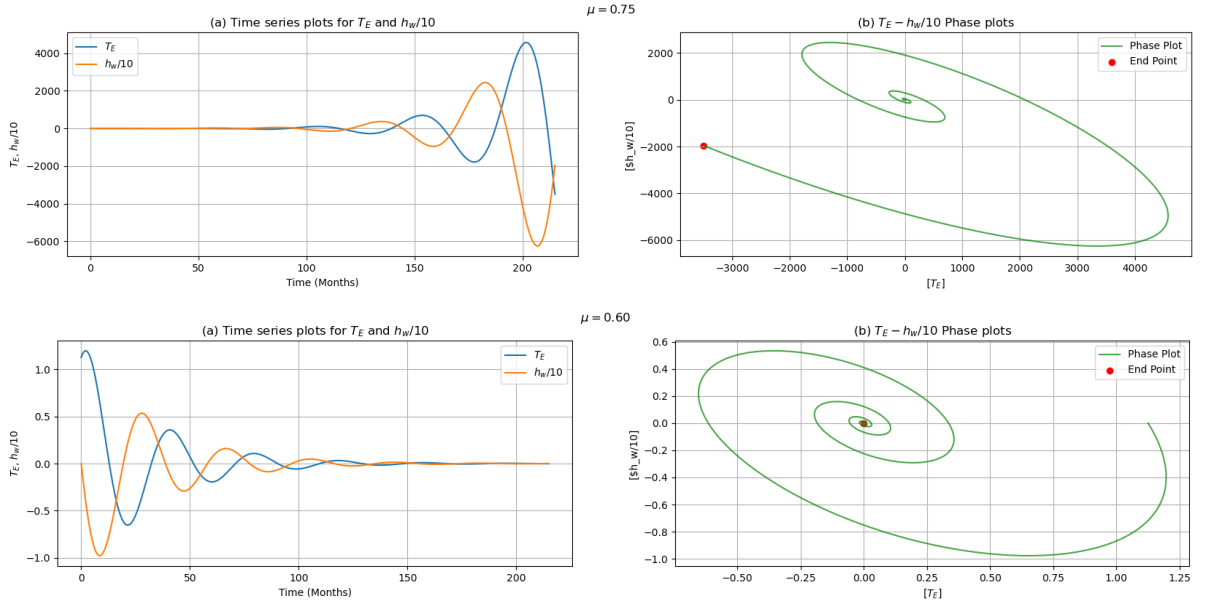
I have checked the stability numerically by running the above model for 100 time periods, and we can see below that RK4 remains stable for both variables. Thus the choice of time scheme is justified.

```
In [3]: endTime=(2*np.pi/np.sqrt(3/32))*100 #dimentionalised
nt=1000*10
dt=endTime/nt
time=np.linspace(0,endTime,nt) *t_nd # time array, dimentionalised
Te,hw=RKplots(time,'$\mu=2/3$ for 100 time periods',nt,dt,T_nd,h_nd,t_nd,0,0,mu_c,0)
```



Task B Testing ROM behaviour around sub-critical and super-critical settings of the coupling parameter

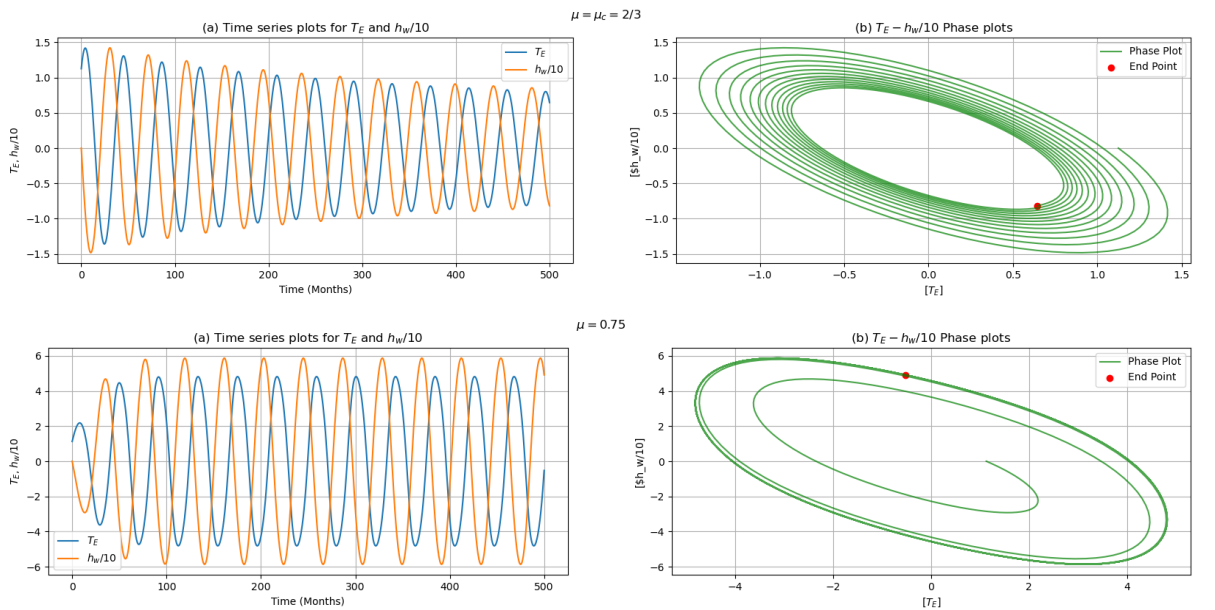
```
In [4]: endTime=215/t_nd
dt=endTime/nt
time=np.linspace(0,endTime,nt) *t_nd # time array, dimentionalised
mu=0.75
Te1,hw1=RKplots(time,'$\mu=0.75$',nt,dt,T_nd,h_nd,t_nd,E1,E2,mu,0,en)
mu=0.60
Te2,hw2=RKplots(time,'$\mu=0.60$',nt,dt,T_nd,h_nd,t_nd,E1,E2,mu,0,en)
```



We run the model with same parameters as part A, except we test it for a super-critical and a sub-critical value of μ , i.e., $\mu = 0.75$ and $\mu = 0.60$ respectively. For $\mu > \mu_c$, the amplitudes blow-up, i.e., becomes self-exciting (explored in more detail in Task D) and for $\mu < \mu_c$, the amplitudes decay. But in both cases, the anomalies remain in a harmonic motion. The period of oscillation has increased in both cases.

Task C: Extending ROM to include the impact of non-linearity

```
In [5]: endTime=500/t_nd
nt=1000
dt=endTime/nt
time=np.linspace(0,endTime,nt) * t_nd # time array, dimentionalised
en=0.1
mu=2/3
Te1,hw1=RKplots(time, '$\mu=\mu_c=2/3$', nt, dt, T_nd, h_nd, t_nd, E1, E2, mu, 0, en)
mu=0.75
Te2,hw2=RKplots(time, '$\mu=0.75$', nt, dt, T_nd, h_nd, t_nd, E1, E2, mu, 0, en)
```

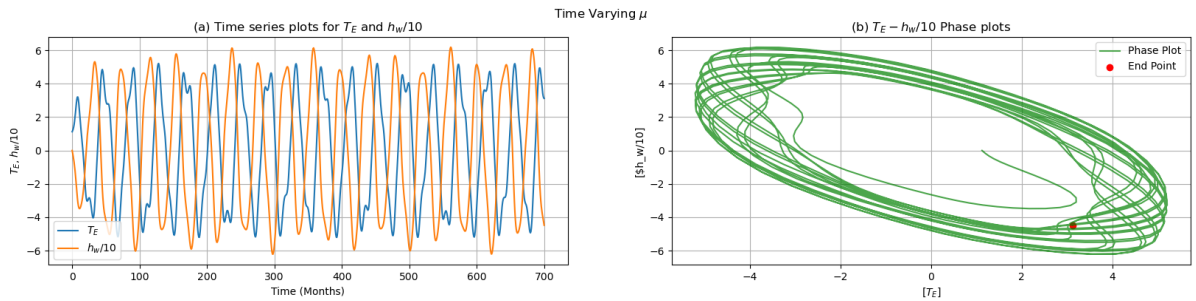


We run the model with non-linearity introduced ($e_n = 0.1$). For $\mu = \mu_c$, comparing with Task A, we see that the non-linearity make the anomalies decay slowly. When μ is increased to a

super-critical value of 0.75, the anomalies show an initial growth, and then oscillate with uniform amplitudes. But amplitude increase for h_w is larger than that for T_E . As seen from Equation (2) above, making $e_n = 0.1$ makes the growth rate a cubic function of T_E and h_w , causing a decay or growth. As stated by Jin (1997a), this non-linearity caps the linear growth, previously seen in Task B. This is because of thermocline feedback, i.e., a huge variation in h_w will not cause a similar variation in T_E and vice-versa due to the now non-linear relationship between them (Jin, 1997a).

Task D: Test the self-excitation hypotheses

```
In [6]: endTime=700/t_nd
nt=1000
dt=endTime/nt
time=np.linspace(0,endTime,nt) * t_nd # time array, dimensionalised
mu_0=0.75
Te,hw=RKplots(time, 'Time Varying  $\mu$ ', nt, dt, T_nd, h_nd, t_nd, E1, E2, mu, mu_0, en, varyn
```



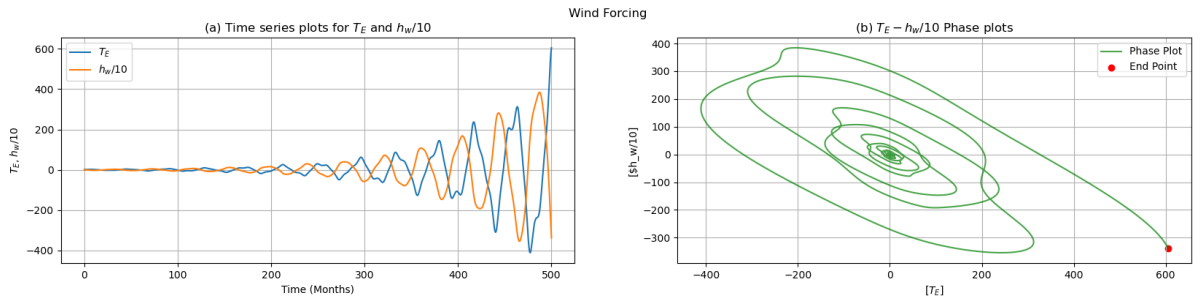
We test the self-excitation hypothesis by introducing an annual forcing by varying μ in an annual cycle as-

$$\mu = \mu_0 \left(1 + \mu_{ann} \cos \left(\frac{2\pi t}{\tau} - \frac{5\pi}{6} \right) \right) \quad (3)$$

, where the period of forcing is $\tau = 12 \text{ months}$, $e_n = 0.1$, $\mu_0 = 0.75$ and $\mu_{ann} = 0.2$. In the SST anomaly, T_E , time series plot we can see wiggles at regular intervals. These are also visible in the phase plot. They are a result of the annual forcing. The time series of both anomalies are still periodic, but with a time period greater than that of the annual forcing. After an initial rise, growth for T_E and h_w is capped, similar to that seen in Task C. The h_w amplitude also shows an oscillation every 3 time periods.

Task E: Test the stochastic initiation hypotheses by adding noisy wind forcing to the linear model

```
In [7]: endTime=500/t_nd
dt=(1/30)/t_nd
nt=int(endTime/dt)
time=np.linspace(0,endTime,nt) * t_nd # time array, dimensionalised
en=0
mu_0=0.69
Te,hw=RKplots(time, 'Wind Forcing', nt, dt, T_nd, h_nd, t_nd, E1, E2, mu, mu_0, en, varymu=True
```



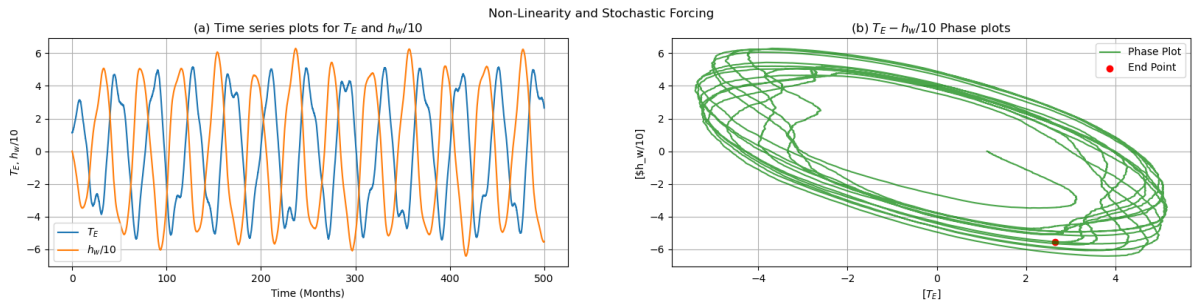
We introduce an annual stochastic forcing in the form of noisy wind stress forcing-

$$\xi_1 = f_{ann} \cos\left(\frac{2\pi t}{\tau}\right) + f_{ran} W \frac{\tau_{cor}}{\Delta t} \quad (4)$$

, where W is a random number between -1 and 1. Nonlinearity is removed ($e_n = 0$) and μ_0 is chosen as 0.69. Since we are in the super-critical region, there is a linear growth in amplitude, similar to Task B, and since $e_n = 0$, there is no cap on the growth of the anomalies, unlike Tasks C,D. The annual forcing creates wiggles in T_E similar to Task D, and the random wind forcing creates

Task F: Testing the non-linearity and the stochastic forcing together

```
In [8]: endTime=500/t_nd
dt=(1/30)/t_nd
nt=int(endTime/dt)
time=np.linspace(0,endTime,nt) *t_nd # time array, dimentionalised
en=0.1
mu_0=0.75
T,h=RKplots(time,'Non-Linearity and Stochastic Forcing',nt,dt,T_nd,h_nd,t_nd,E1,E2,
```



We run the model with similar parameters to Task E, except we turn ON the nonlinearity ($e_n = 0.1$) and set $\mu_0 = 0.75$. The initial linear growth is now capped, similar to Task C and D. The annual forcing also creates trends in T_E and h_w similar to those in Task D. But here we can also see the effect of wind stress as random vibrations in the phase trajectory.

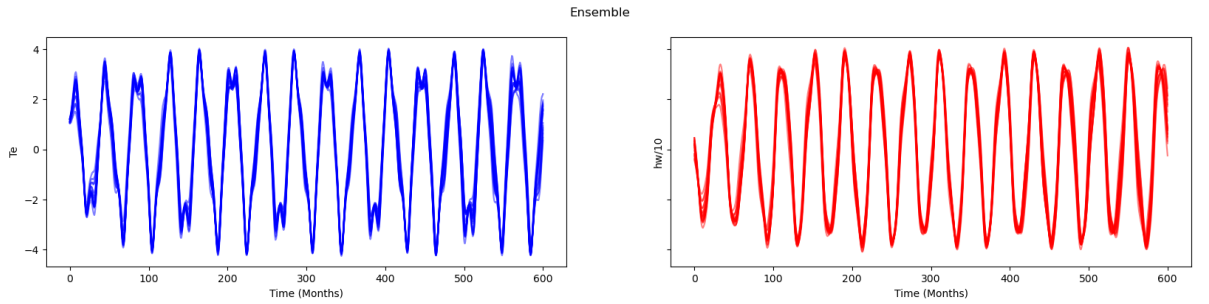
Task G: Testing whether chaotic behaviour can be triggered through addition of initial condition uncertainty

```
In [9]: endTime=600/t_nd
nt=5000
dt=endTime/nt
time=np.linspace(0,endTime,nt) *t_nd # time array, dimentionalised
en=0.1
mu_0=0.70
```

```

no_en=20 # no of ensemble members
Trange=0.1/T_nd
hrange=5/h_nd
T_purtinit=np.zeros(no_en)
h_purtinit=np.zeros(no_en)
fig, (ax1, ax2) = plt.subplots(1,2,figsize=(20,4), sharey=True)
fig.suptitle('Ensemble')
for i in range(no_en):
    T_purtinit[i]=Tinit+np.random.uniform(-Trange,Trange)
    h_purtinit[i]=hinit+np.random.uniform(-hrange,hrange)
    Te,hw= RKplots(time,'Time Varying $\mu$',nt,dt,T_nd,h_nd,t_nd,E1,E2,mu,mu_0,en,
    ax1.plot(time,Te,label='Te',alpha=0.5,color='blue')
    ax2.plot(time,hw/10,label='hw/10',alpha=0.5,color='red')
    ax1.set_xlabel("Time (Months)")
    ax2.set_xlabel("Time (Months)")
    ax1.set_ylabel("Te")
    ax2.set_ylabel("hw/10")

```



Setup

We create an ensemble by perturbing the initial conditions to check if the model is chaotic, meaning if it deviates significantly from the unperturbed system if a small external forcing is introduced. A 20 member ensemble is created with T_E , *initial* in range $1.125 \pm 0.1K$ and h_w , *initial* in the range $0 \pm 5m$. Values in the given ranges are chosen at random for every member of the ensemble. Nonlinearity is ON, and annual and stochastic wind forcings are applied.

Discussion

The perturbations to initial conditions only produce small, temporary deviations in some members of the ensemble. All members show characteristics similar to the system modelled in Task F. Deviation of ensemble members from reference state are visible following places on the plots where effects of external forcings like annual forcing and random wind stress are prominent, but deviations go away. This also shows that the system is not sensitive to periodic or random external forcings applied after initial time.

Appendix: RK4

The coupled Equations (1) and (2) can be re-written as-

$$\frac{d}{dt} \begin{bmatrix} T \\ h \end{bmatrix} = \begin{bmatrix} RT + \gamma h - \epsilon(h + bT)^3 + \gamma\xi_1 + \xi_2 \\ -rh_w - \alpha bT_E - \alpha\xi_1 \end{bmatrix} = F\left(\begin{bmatrix} T \\ h \end{bmatrix}, t, \dots\right) \quad (5)$$

The RK4 solution is then given by-

$$T^{n+1} = T^n + \Delta t(k_1 + 2k_2 + 2k_3 + k_4) \quad (6)$$

and

$$h^{n+1} = h^n + \Delta t(l_1 + 2l_2 + 2l_3 + l_4) \quad (7)$$

, where the super script n denotes the value at n^{th} time step. The coefficeints k_i and l_i are calculated at-

$$\begin{bmatrix} k_1 \\ l_1 \end{bmatrix} = F\left(\begin{bmatrix} T^n \\ h^n \end{bmatrix}, t, \Delta t, \mu, \dots\right) \quad (8)$$

$$\begin{bmatrix} k_2 \\ l_2 \end{bmatrix} = F\left(\begin{bmatrix} T^n + k_1\Delta t/2 \\ h^n + l_1\Delta t/2 \end{bmatrix}, t + \Delta t/2, \Delta t, \mu, \dots\right) \quad (9)$$

$$\begin{bmatrix} k_3 \\ l_3 \end{bmatrix} = F\left(\begin{bmatrix} T^n + k_2\Delta t/2 \\ h^n + l_2\Delta t/2 \end{bmatrix}, t + \Delta t/2, \Delta t, \mu, \dots\right) \quad (10)$$

$$\begin{bmatrix} k_4 \\ l_4 \end{bmatrix} = F\left(\begin{bmatrix} T^n + k_3\Delta t \\ h^n + l_3\Delta t \end{bmatrix}, t + \Delta t, \Delta t, \mu, \dots\right) \quad (11)$$

References

- Jin, F.-F. (1997a). An equatorial ocean recharge paradigm for ENSO: Part I: Conceptual model. *J. Atmos. Sci.*, 54, 811–829
- Jin, F.-F. (1997b). An equatorial ocean recharge paradigm for ENSO: Part II: A stripped down coupled model. *J. Atmos. Sci.*, 54, 830–847.