

Probabilistic Seismic Hazard Assessment of Central Nepal Himalayan Region

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ABSTRACT

1 The Central Nepal Himalayan region is one of the most seismically hazardous
2 areas on Earth due to the ongoing collision of the Indian and Eurasian tectonic
3 plates. This study presents a comprehensive Probabilistic Seismic Hazard Assess-
4 ment (PSHA) for this region, utilizing an updated earthquake catalog and mod-
5 ern analytical techniques. We compiled seismic data for the central Himalayas,
6 removed aftershocks through declustering, and estimated earthquake recurrence
7 parameters using the Gutenberg-Richter relation. The frequency-magnitude anal-
8 ysis yields a Gutenberg-Richter b-value of approximately 0.9, indicating a slightly
9 higher proportion of larger earthquakes compared to the global average. Using
10 these parameters and appropriate ground motion models, we computed the likeli-
11 hood of exceeding various levels of ground shaking across the region. The results
12 are synthesized in a seismic hazard map for peak ground acceleration (PGA) with a
13 10% probability of exceedance in 50 years (equivalent to a 475-year return period).
14 The hazard map reveals significant potential ground motions, with PGA values
15 ranging from 0.3 to 0.5 g in portions of central Nepal. These levels of seismic haz-
16 ard are comparable to or exceed previous estimates, underscoring the significant
17 earthquake risk faced by the region's population centers. Our findings highlight
18 the critical need for earthquake-resistant design and risk mitigation measures in
19 central Nepal. This study not only updates the seismic hazard profile of the re-
20 gion, considering recent data (including the 2015 Gorkha earthquake), but also
21 provides a methodological framework for ongoing hazard assessment efforts. The
22 comprehensive PSHA presented here will aid engineers, urban planners, and pol-
23 icymakers in developing effective strategies to improve resilience against future
24 earthquakes.

25 **Key words:** PSHA; Seismic Hazard in Himalaya

1 INTRODUCTION

26 Nepal lies in the active Himalayan collision zone where the Indian Plate thrusts
27 under the Eurasian Plate, resulting in frequent earthquakes and a high seismic
28 hazard. The country's historical record includes devastating large earthquakes,
29 such as the Mw 8.0 1934 Nepal–Bihar earthquake, which severely damaged Kath-
30 mandu and caused approximately 10,600 fatalities, and the more recent Mw 7.8
31 Gorkha earthquake of 2015, which struck central Nepal, resulting in nearly 9,000
32 deaths (Avouac et al., 2015; Bai et al., 2019; Dixit et al., 2015; Kurashimo et al.,
33 2019; Pandey et al., 1999b). These events starkly highlight Nepal's vulnerability
34 to seismic disasters. Indeed, despite large earthquakes being relatively infrequent
35 on the Main Himalayan Thrust, the region's dense population and infrastructure
36 exposure make even moderate events potentially catastrophic. This context under-
37 scores the importance of robust seismic hazard assessment for Nepal.

38 Probabilistic Seismic Hazard Assessment (PSHA) is a methodology to quan-
39 tify the likelihood of different levels of earthquake ground shaking occurring in
40 a region over a given time period. Unlike deterministic scenarios, PSHA consid-
41 ers the full range of possible earthquakes and their uncertainties in size, location,
42 and recurrence. The approach, first formalized by Cornell & Toro (1970), combines
43 information on earthquake recurrence with models of ground motion attenuation
44 to estimate the probability of exceeding various ground-motion levels. PSHA re-
45 sults are often expressed as hazard curves or maps for specified probability levels,
46 which are crucial for developing building codes and risk mitigation strategies.

47 In Nepal, systematic seismic hazard assessments began in the 1990s. A UNDP-
48 supported project in 1993 produced the first seismic hazard map for Nepal as part
49 of the national building code development (NBC 105:1994). That study estimated
50 peak ground accelerations (PGA) for a 500-year return period (10% probability
51 in 50 years) and delineated seismic zones used in the building code. Subsequent
52 studies have refined the hazard model. Notably, Pandey et al. (1999a) carried out
53 a country-wide PSHA using the CRISIS99 software, dividing Nepal into 12 seis-
54 mic source zones and employing an attenuation relation from Youngs et al. (1997).
55 Their results, shown in Figure 1, provided an updated seismic hazard map for
56 Nepal with PGA values ranging roughly from 0.10g to 0.45g (on rock, 10% in 50
57 years) across the country. This map informed Nepal's engineering design practices
58 in the early 2000s. More recent analyses have continued to update Nepal's hazard

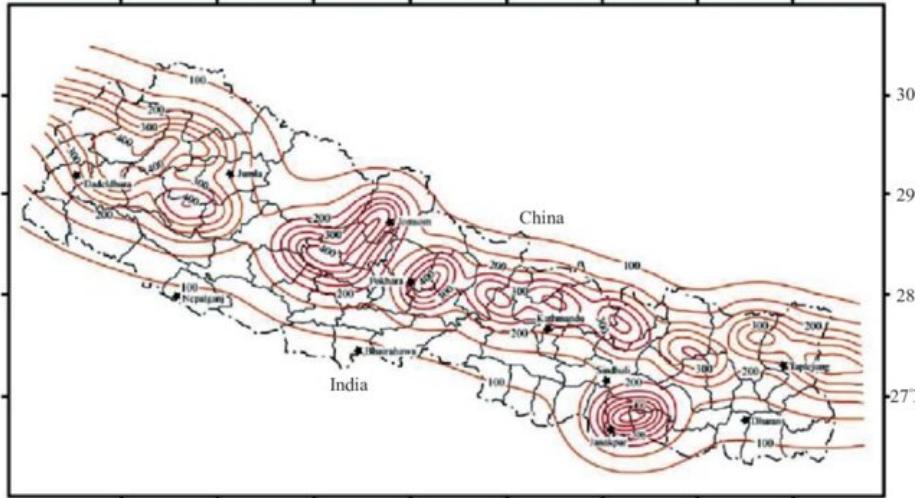


Figure 1: Seismic hazard map of Nepal from Pandey et al. (2002), showing contours of peak horizontal acceleration (in units of %g) with 10% probability of exceedance in 50 years. This prior study identified significant hazard across the Nepal Himalaya, with values ranging up to 0.45g in parts of the country. It provided the basis for Nepal's building code seismic zonation.

59 assessment: for example, Parajuli et al. (2010) applied kernel-based earthquake
 60 density estimation and multiple ground-motion prediction equations to account
 61 for epistemic uncertainties, finding especially high hazard around the Kathmandu
 62 Valley. Thapa & Wang (2013) further refined the source zonation (delineating 23
 63 seismic source zones) and produced hazard maps for various probability levels,
 64 noting significant hazard in the far-western and eastern Nepal Himalayas. These
 65 studies commonly conclude that central and eastern Nepal face very high seismic
 66 risk and emphasize the need to update hazard models as new data become avail-
 67 able periodically.

68 Given the advances in data collection and methodology since the early 2000s
 69 – including an expanded earthquake catalog and lessons from the 2015 Gorkha
 70 event – it is imperative to reassess the seismic hazard in central Nepal with up-to-
 71 date information. The central Nepal Himalayan region, encompassing the greater
 72 Kathmandu area and surrounding districts, is of particular concern due to its con-
 73 centrated population and infrastructure in a zone of high tectonic strain. This pa-
 74 per presents a detailed PSHA for the Central Nepal Himalaya, using an updated
 75 earthquake catalog and state-of-the-art methods. We aim to quantify the current
 76 level of hazard (in terms of PGA for a 10% exceedance in 50 years) for the region
 77 and compare it with previous assessments. In the following sections, we describe

⁷⁸ the data and methodology, present the resulting seismic hazard model and maps,
⁷⁹ discuss their implications in the context of prior studies, and finally conclude with
⁸⁰ recommendations for earthquake risk management in Nepal.

2 METHODS

⁸¹ Our probabilistic seismic hazard assessment follows the standard PSHA proce-
⁸² dure, which involves: (1) assembling a seismic source model from earthquake oc-
⁸³ currence data, (2) selecting ground-motion prediction models to estimate shaking
⁸⁴ from those earthquakes, and (3) performing a probabilistic calculation to obtain
⁸⁵ hazard levels for specified exceedance probabilities. Below, we detail each of these
⁸⁶ steps as applied to the Central Nepal Himalayan region.

87 2.1 Study Region and Earthquake Catalog

⁸⁸ The study region spans approximately 80°E–88°E in longitude and 26°N–30.5°N in
⁸⁹ latitude, covering central Nepal and adjoining areas of the Himalayan frontal fault
⁹⁰ system. We compiled an earthquake catalog for this region from multiple sources,
⁹¹ including the United States Geological Survey (USGS) and Nepal’s National Seis-
⁹² mological Centre, incorporating both historical events and modern instrumental
⁹³ records. The catalog extends back to the year 1255 A.D. for major historical earth-
⁹⁴ quakes and is comprehensive for instrumental events roughly since the early 20th
⁹⁵ century. To ensure consistency, all earthquakes magnitudes were converted to mo-
⁹⁶ ment magnitude (M_w) using appropriate empirical relationships for different orig-
⁹⁷ inal magnitude scales. The raw compiled catalog contained on the order of a few
⁹⁸ thousand events down to small magnitudes.

⁹⁹ A key aspect of preparing the catalog for PSHA is declustering, which removes
¹⁰⁰ dependent events (aftershocks and foreshocks) so that the remaining events repre-
¹⁰¹ sent a Poissonian, independent occurrence process. We applied a standard declus-
¹⁰² tering algorithm (the Gardner and Knopoff window method) to eliminate after-
¹⁰³ shocks from the catalog. This reduced the total event count significantly, from 2250
¹⁰⁴ events to 1271 events in the final declustered catalog (for the magnitude range con-
¹⁰⁵ sidered). The declustered catalog is assumed to represent the activity of the princi-
¹⁰⁶ pal seismogenic sources in the region without double-counting clusters of shocks
¹⁰⁷ from single earthquake sequences. We then assessed the completeness of the cat-
alog to determine a magnitude threshold above which the data can be considered

complete for the time period of interest. By examining the rate of earthquake occurrences over time, we identified a magnitude of completeness around $Mc \simeq 4.4$, meaning that all earthquakes of about 4.4 and greater are reliably recorded in the catalog over the past several decades. We therefore base our recurrence analysis on events with $M \geq 4.4$ in the declustered catalog.

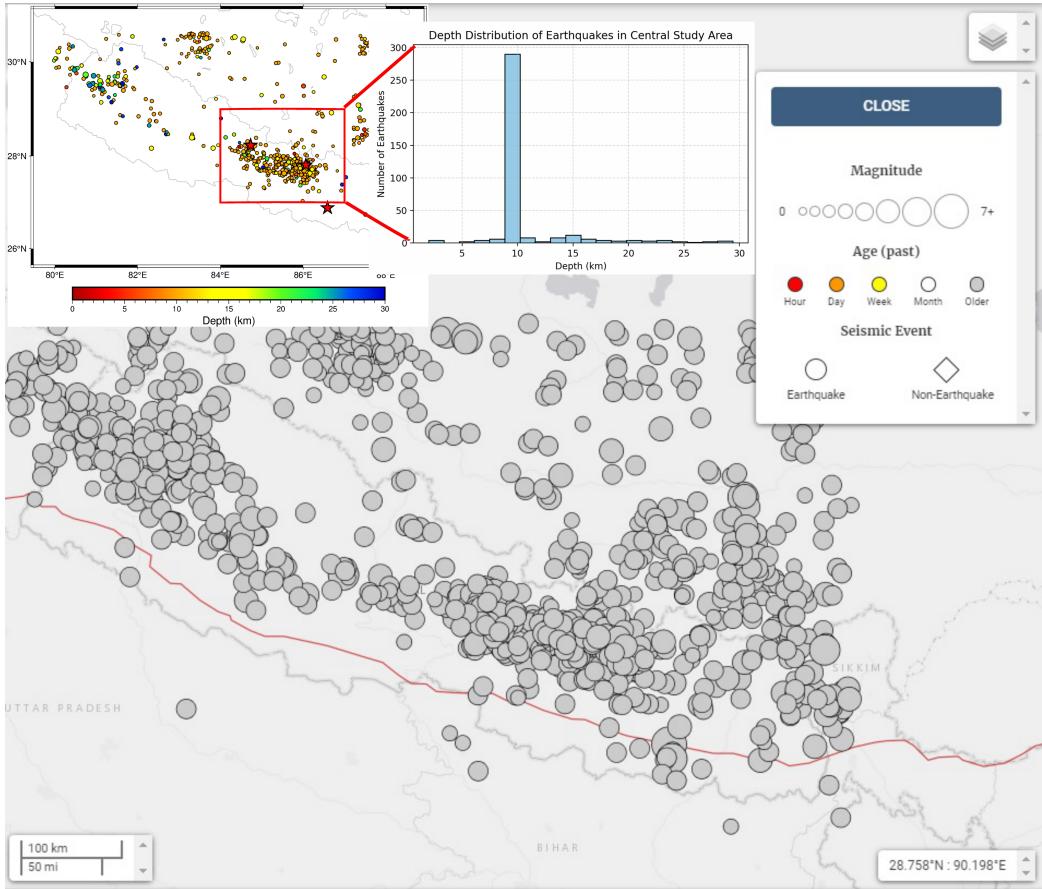


Figure 2: Spatial distribution of earthquakes in the declustered catalog for central Nepal (circles denote earthquake epicenters). This map, generated from the USGS earthquake database, shows 1,271 independent events (after aftershock removal) in the study region. The density of epicenters is highest along the Himalayan belt, delineating the active Main Himalayan Thrust and associated structures. The red line represents the Main Frontal Thrust (southern border of the Himalayas).

2.2 Earthquake Recurrence Model

We characterized the seismic source for PSHA using an area-source model covering the central Nepal Himalaya. Within this source, earthquakes are assumed

117 to follow the Gutenberg-Richter frequency–magnitude relationship, which is com-
118 monly expressed as:

$$\log_{10} N(M) = a - bM \quad (1)$$

119 where $N(M)$ is the cumulative annual frequency of earthquakes with magni-
120 tude $\geq M$, b is the slope (the Gutenberg-Richter b-value), and a is the productivity
121 constant representing overall activity rate. We estimated the parameters a and b
122 from the processed earthquake catalog. The Gutenberg-Richter b-value was cal-
123 culated using the maximum-likelihood method of Aki (1965), which provides an
124 unbiased estimator even for incomplete data bins. In this method, b is given by
125 (Aki, 1965; Bender, 1983; Utsu, 1965):

$$b = \frac{\log_{10}(e)}{\bar{M} - \left(M_c - \frac{\Delta M}{2}\right)} \quad (2)$$

126 where \bar{M} is the mean magnitude of events above the completeness threshold
127 M_c , and ΔM is the binning interval (here 0.1, so $\Delta M/2 = 0.05$). For our catalog
128 (with $M_c = 4.4$), the maximum-likelihood computation yielded $b \approx 0.90$, with a
129 standard error on the order of 0.03. The statistical uncertainty for maximum likeli-
130 hood b-value estimates was determined using Shi & Bolt (1982).

$$\sigma_b = \frac{b^2}{\log(e)} \sqrt{\frac{\sum_{i=1}^n (M_i - \bar{M})}{n(n-1)}}, \quad (3)$$

131 This b-value indicates that the relative frequency of large to small earthquakes in
132 central Nepal is slightly below unity, which is in line with typical active tectonic
133 regions (a b-value around 1.0 is often observed globally). The Gutenberg-Richter
134 a-value was determined from the rate of events above M_c ; in our case, the catalog
135 data suggest an annual rate of roughly 5.41 earthquakes of $M \geq 4.4$, corresponding
136 to an a-value (intercept) of about 6.6 (in the log10 scale). The resulting regional
137 recurrence relationship can be written as $\log_{10}N = 6.6 - 0.90M$, which was used as
138 the seismic source model for the PSHA calculations.

139 2.3 Ground Motion Prediction Models

140 To compute ground shaking at sites due to potential earthquakes, we adopted
141 empirically-based ground motion prediction equations (GMPEs) appropriate for

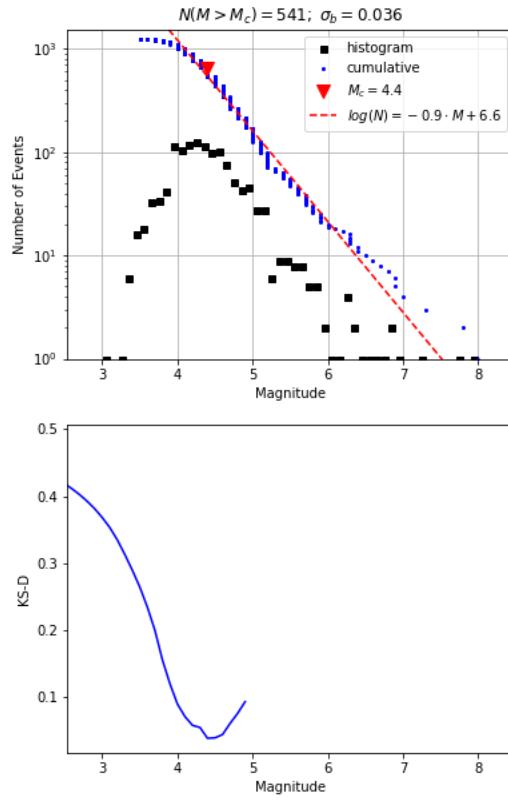


Figure 3: Magnitude-frequency analysis of the central Nepal earthquake catalog. (Top panel): Logarithmic plot of the earthquake magnitude distribution. Black squares represent the histogram of events, while blue dots show the cumulative number of events. The red inverted triangle indicates the magnitude of completeness, $M_c = 4.4$, above which the Gutenberg–Richter law is applied. The red dashed line is the best-fit Gutenberg–Richter relation, $\log_{10} N = -0.9M + 6.6$, derived using the maximum likelihood method. The catalog includes 541 independent events above M_c , with an estimated b-value uncertainty of $\sigma_b = 0.036$. (Bottom panel): Kolmogorov–Smirnov (K–S) statistic as a function of magnitude, showing the goodness-of-fit between the observed and modeled distributions. The minimum K–S distance at $M = 4.4$ supports the choice of completeness magnitude.

the Himalayan region. Previous PSHA studies in Nepal have used attenuation relations developed for similar tectonic environments (for example, Youngs et al. 1997 for subduction-zone interface earthquakes). For this study, we selected a modern GMPE that has been validated against strong-motion data in active continental collision zones comparable to Nepal. In particular, we utilized a ground-

147 motion model that accounts for the magnitude, distance, and site conditions to es-
148 timate peak ground acceleration (PGA) on rock sites. The chosen GMPE provides a
149 median PGA value and associated standard deviation (log-normal dispersion) for
150 a given earthquake scenario (magnitude and distance). To capture epistemic un-
151 certainty in ground motion estimates, one could in principle use multiple GMPEs;
152 however, for simplicity, our base case uses a single representative GMPE while we
153 later comment on the potential range of results. All sites in the region were as-
154 sumed to be rock or firm soil (reference site condition for the GMPE) to produce a
155 regional hazard map on bedrock; this can later be adjusted for local soil conditions
156 if needed.

157 2.4 PSHA Calculation

158 Using the defined seismic source model—characterized by a Gutenberg–Richter
159 recurrence relationship within a uniform area source—and the selected ground
160 motion prediction equation (GMPE), we conducted probabilistic seismic hazard
161 calculations over a grid of sites covering the central Nepal Himalayan region. We
162 assumed a *Poissonian model* for earthquake occurrence, which treats earthquakes
163 as statistically independent events in time. This assumption is standard in proba-
164 bilistic seismic hazard analysis (PSHA), as it simplifies the mathematical treatment
165 and aligns with the use of declustered seismic catalogs.

166 At the heart of the PSHA is the estimation of the likelihood that ground motion
167 at a site will exceed a certain threshold within a given time frame due to earth-
168 quakes of various magnitudes and distances. This is achieved using the *total prob-
169 ability theorem*, originally formulated for seismic applications by Cornell & Toro
170 (1970). According to this framework, the annual rate of exceedance $\lambda(Y > y)$ of
171 a ground motion level y is obtained by integrating over all possible earthquake
172 magnitudes and source-to-site distances, weighted by their respective occurrence
173 probabilities and ground motion exceedance probabilities:

$$\lambda(Y > y) = \int_{m_{\min}}^{m_{\max}} \int_{r_{\min}}^{r_{\max}} \nu(m) \cdot f_R(r|m) \cdot P(Y > y | m, r) dr dm \quad (4)$$

174 where:

- 175 • $\lambda(Y > y)$ is the annual frequency of exceedance of ground motion level y ,
- 176 • $\nu(m)$ is the magnitude-dependent annual occurrence rate, derived from the

177 Gutenberg–Richter recurrence law,

- 178 • $f_R(r|m)$ is the probability density function of source-to-site distance r for a
179 given magnitude m ,
- 180 • $P(Y > y | m, r)$ is the conditional probability of exceeding ground motion y
181 given an earthquake of magnitude m at distance r , as modeled by the GMPE.

182 This integral quantifies the contribution of all possible earthquake scenarios to
183 the seismic hazard at a specific site. In practice, the integral is evaluated numer-
184 ically by discretizing both magnitude and distance bins. The magnitude range
185 considered spans from the completeness threshold ($M_c = 4.4$) up to a maximum
186 credible magnitude $M_{\max} = 8.5$, which reflects the upper bound for Himalayan
187 seismicity based on both historical records (e.g., the 1934 M8.1–8.4 Nepal–Bihar
188 earthquake) and geological fault constraints.

189 The ground motion metric used in this study is Peak Ground Acceleration
190 (PGA), a commonly adopted intensity measure in earthquake-resistant design and
191 building code formulations. The GMPE provides median PGA values along with
192 variability (standard deviation) for each earthquake scenario, taking into account
193 magnitude, distance, and standard site conditions (assumed here as rock or firm
194 soil).

195 For each site on the computation grid, we calculated a hazard curve—a func-
196 tion describing the annual probability of exceeding various levels of PGA. From
197 this hazard curve, we extracted the PGA value corresponding to a 10% probability
198 of exceedance in 50 years, which is equivalent to a 475-year return period. This
199 level of exceedance is a conventional benchmark in engineering seismology and is
200 widely used in the development of seismic building codes and design spectra.

201 The resulting hazard values across the study region were spatially interpolated
202 to generate a probabilistic seismic hazard map for central Nepal. This map reflects
203 the spatial distribution of expected shaking intensities under a common design
204 scenario and provides a basis for comparing with prior national assessments, such
205 as those conducted by Pandey et al. (2002), as well as for updating zoning param-
206 eters in Nepal’s building codes.

207 In summary, our methodology adheres to established PSHA standards, inte-
208 grating a rigorously processed earthquake catalog, statistically robust recurrence
209 modeling, and regionally appropriate ground motion predictions. This approach

210 provides a scientifically grounded estimate of seismic hazard and supports efforts
211 in earthquake risk mitigation and resilient infrastructure planning across central
212 Nepal.

3 RESULTS

213 3.1 Seismicity and Recurrence Characteristics

214 The processed earthquake catalog for central Nepal provides insight into the re-
215 gion's seismicity and forms the basis for hazard quantification. The spatial distri-
216 bution of past earthquakes (Figure 2) shows that seismicity is concentrated along
217 the Himalayan arc, especially beneath the higher Himalaya north of the foothills.
218 This pattern aligns with the Main Himalayan Thrust (MHT) fault system, which
219 is the source of large megathrust earthquakes. The declustering process retained
220 mainly mainshock events, including several significant historical earthquakes in
221 the catalog. The magnitude–frequency analysis of these events confirms that the
222 Gutenberg-Richter law is a reasonable representation for central Nepal's seismic-
223 ity. Figure 3 illustrates the frequency–magnitude distribution and the fitted re-
224 currence law. The plot appears linear above the completeness threshold (M4.4),
225 supporting the use of a power-law recurrence model.

226 From the Gutenberg-Richter fit, we obtained a b-value around 0.90 for the cen-
227 tral Nepal catalog. This b-value is slightly below 1.0, which suggests a somewhat
228 greater relative frequency of large earthquakes compared to an average tectonic
229 region (where b is often near unity). In physical terms, a lower b-value can imply
230 a region of higher tectonic stress or one dominated by large fault structures capa-
231 ble of generating major earthquakes. The $b \approx 0.9$ found here is consistent with
232 other studies of Himalayan seismicity - for example, previous estimates for Nepal
233 have found b-values in the range $\sim 0.8 - 1.0$. The total seismic activity rate given
234 by the a-value (~ 6.6 in the equation 1) indicates that, on average, the region expe-
235 riences an earthquake of about M5.0 or greater roughly every 2 - 3 months, and an
236 earthquake of M6.0 or greater roughly every few years (according to the recurrence
237 model). These recurrence characteristics reflect an active plate boundary with fre-
238 quent moderate earthquakes and occasional great earthquakes. They also set the
239 stage for calculating hazard: the relatively high rate of $M \geq 5$ events and the possi-
240 bility of $M \geq 8$ events in this region means that there is a significant probability of

241 strong ground shaking within any given 50-year period.

242 3.2 Probabilistic Hazard Analysis and Hazard Map

243 Using the recurrence parameters and ground-motion models described, we com-
244 puted the probabilistic seismic hazard for central Nepal. The primary result is
245 the estimated peak ground acceleration (PGA) values for a 10% probability of ex-
246 ceedance in 50 years. Figure 4 presents the resulting seismic hazard map for the
247 region. Each contour or color band on the map represents the level of PGA (as a
248 fraction of gravity, g) that has a 10% chance of being exceeded at least once in a
249 50-year interval at that location.

Nepal Himalayan Region 10% in 50y PGA Hazard Map

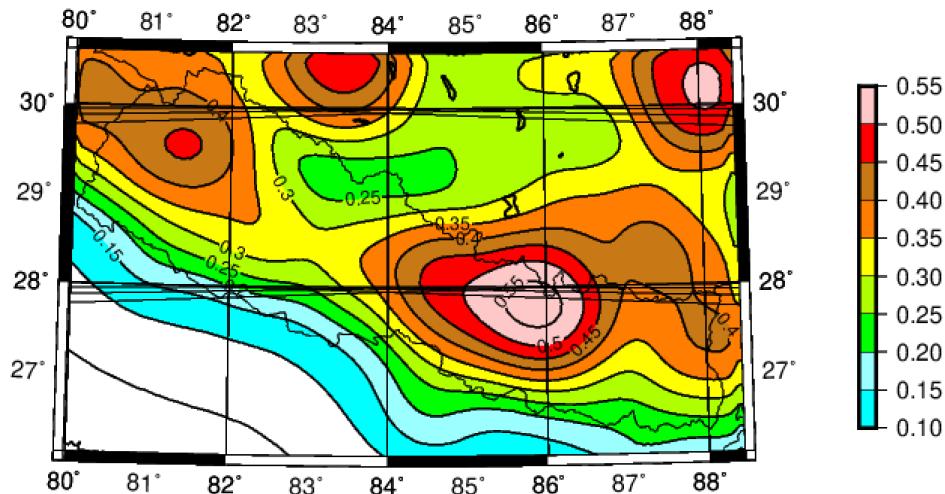


Figure 4: Probabilistic seismic hazard map for the central Nepal Himalayan region, showing the peak ground acceleration (PGA) with 10% chance of exceedance in 50 years (approximately a 475-year return period). The map is in units of gravity (g). Higher hazard levels (warm colors) are predicted along the Himalayan range, indicating where strong shaking is most likely. Notably, parts of central and eastern Nepal exhibit PGA values on the order of 0.4–0.5g (red shades), while areas further south into the Ganges plain have significantly lower values (blue-green shades).

250 As shown in Figure 4, the hazard is not uniform across the region. PGA values
251 range from about 0.10–0.15g in the southernmost Terai plains (at the foothills of
252 the Himalayas) to over 0.50g in some northern areas closer to the high Himalaya.

253 Broadly, the highest hazard is concentrated in a belt running west-east along cen-
254 tral Nepal, roughly paralleling the main seismic sources. There are two notable
255 high hazard concentrations: one in central-western Nepal and another in east-
256 central Nepal. In these zones, the 10%-in-50yr PGA reaches around 0.45g to 0.55g,
257 which is extremely high and comparable to design level shaking in the world's
258 most earthquake-prone regions. Kathmandu valley, located in central Nepal, falls
259 within a high hazard area – our results indicate a design-level PGA on the order of
260 0.35–0.40g for stiff soil or rock sites in Kathmandu. This is consistent with the ex-
261 pectation that the capital region, having experienced major shaking in events like
262 1934 and 2015, remains one of the zones of highest seismic hazard. By contrast, the
263 far-western Nepal Himalaya (e.g., around 80°–81°E longitude) shows somewhat
264 lower hazard in our model (PGA generally under 0.3g). The lesser hazard in the
265 far west from our analysis may reflect the lower frequency of large earthquakes in
266 that segment over the catalog period; however, we note that some other studies
267 have found high hazards in the far-west if considering the potential for an over-
268 due great earthquake there. The southernmost parts of central Nepal (toward the
269 Gangetic plain) have the lowest hazard in the region, with PGA generally below
270 0.2g, due to their greater distance from the Himalayan seismogenic sources and
271 the attenuative thick sediments of the Ganga basin.

272 Overall, the 10%/50yr PGA values we obtained for central Nepal are on the
273 order of 0.3–0.5g in the high hazard areas. These values are in line with or slightly
274 higher than the previous national-scale assessment by Pandey et al. (2002), which
275 gave 0.10–0.45g across Nepal. In particular, our estimate of 0.4–0.5g in parts of cen-
276 tral/eastern Nepal slightly exceeds Pandey et al.'s maximum of 0.45g, which could
277 be due to the inclusion of the 2015 Gorkha earthquake data and updated ground
278 motion models that predict somewhat higher motions. It is also noteworthy that
279 our hazard map highlights the central Nepal region (including Gorkha, Lamjung,
280 and Kathmandu areas) as having very high hazard, whereas Pandey's map had
281 indicated the highest contours more towards eastern Nepal. Our results show a
282 broad high-hazard region that extends from central to eastern Nepal, reflecting
283 that both the central (e.g., 2015 rupture area) and eastern (e.g., 1934 rupture area)
284 segments of the Himalayan fault system are capable of producing severe shaking
285 in the future. To interpret these results: a PGA of 0.5g at 10%/50yr means that
286 there is a 10% chance that earthquake shaking will exceed 50% of gravity at that
287 location at least once in the next 50 years. Such shaking would be extremely dam-

aging, likely causing collapse of poorly constructed buildings. Even 0.3g shaking (found over wide areas in Figure 4) can cause serious damage without adequate engineering. Therefore, the hazard map quantitatively confirms that central Nepal faces a very high seismic risk. The map also provides a basis for more detailed risk analysis – for example, combining these hazard levels with exposure (buildings, population) would allow estimation of expected losses under the 475-year return period event. It should be noted that the hazard values carry uncertainties. The map represents the mean estimates given our model assumptions. If alternative reasonable GMPEs or slightly different b-values were used, the PGA levels might shift by a few tenths of g. Nonetheless, the overall pattern of highest hazard along the Himalayan front is robust. The results are thus a best-estimate of the current state of seismic hazard in central Nepal, suitable for informing building code updates and disaster preparedness planning.

4 DISCUSSION

The PSHA results for central Nepal presented above have important implications and are generally consistent with our understanding of regional seismic risk, though there are some notable points to discuss in comparison to previous studies and in the context of uncertainties.

4.1 Comparison with Previous Hazard Assessments

Our hazard map broadly agrees with earlier assessments in identifying the Himalaya of central and eastern Nepal as a zone of very high hazard. The national seismic zoning in the 1994 Nepal building code (based on the 1993 study) had peak accelerations of about 0.08–0.12g in central Nepal, which is considerably lower than what we find – this is expected, as the 1993 study used a shorter catalog and perhaps more conservative assumptions, and building code values often incorporate safety margins and older attenuation models. The subsequent 2002 study by Pandey et al. raised the estimated hazard levels to as high as 0.45g in eastern Nepal. Our study, incorporating data from the last two decades (including the Gorkha earthquake), suggests hazard levels that are equal or slightly higher in some locales (up to ~0.5g). This could indicate that the central-eastern Nepal segment is at least as hazardous as previously thought, if not more so. On the other hand, our model yielded somewhat lower hazard in far-western Nepal than Thapa

319 and Wang (2013) reported. Thapa and Wang identified the far-west as another high
320 hazard zone (they used a different methodology involving morphostructural zon-
321 ing and a Chinese GMPE). The difference underscores how sensitive hazard results
322 can be to assumptions about seismic source characterization – if one assumes the
323 500 km western Nepal segment is capable of rupturing in a single great earthquake
324 ($M_w \sim 8.5+$, as some geological studies suggest a large earthquake deficit there),
325 the hazard in the far-west would rise. Our study implicitly assumes the past cat-
326 alog (which lacks an event of that size in the far-west) is indicative of the future;
327 if that is wrong, hazard there could be underestimated. This points to the need
328 for careful consideration of M_{max} and fault segmentation in PSHA. Future work
329 integrating paleoseismic findings (e.g., evidence of great earthquakes prior to 20th
330 century) could refine the source model and possibly raise the hazard in segments
331 that have been quiet in recent centuries.

332 4.2 Gutenberg-Richter Parameters and Seismotectonic Interpretation

333 The estimated b-value of ~ 0.9 for central Nepal is in line with other estimates in
334 active tectonic regions, but slightly on the lower side. A b-value less than 1 sug-
335 gests relatively more frequent large earthquakes, which could reflect the presence
336 of very large faults (the Himalayan megathrust) that dominate the seismic energy
337 release. It is known that b-values can vary spatially; for instance, some studies
338 have found higher b ($\sim 1.0\text{--}1.1$) in aftershock sequences and lower b ($\sim 0.7\text{--}0.8$)
339 in locked, highly stressed fault regions. Our regional b is an average; the 2015
340 aftershock zone itself had a somewhat higher b, whereas the locked segments
341 might have lower b. This nuance is somewhat averaged out in our area-source
342 approach. The relatively small uncertainty of our b estimate (± 0.03) is likely too
343 optimistic, as it's based on assuming completeness above $M4.4$; if the catalog com-
344 pleteness or magnitude homogeneity had issues, the true uncertainty could be
345 higher. Nonetheless, a b in the $0.8\text{--}1.0$ range is reasonable for Himalayan seismic-
346 ity, and our chosen value contributes to hazard mainly by controlling the frequency
347 of moderate vs. large events. We did consider $M_{max} \sim 8.5$; if a larger maximum
348 (say 8.7 or 9.0) were considered, the hazard might increase slightly at long return
349 periods, but such an event might be beyond what the Himalaya can generate in a
350 single rupture, according to current geological understanding.

351 **4.3 Uncertainties in Ground Motion Modeling**

352 One of the largest uncertainties in any PSHA is the choice of GMPE. We used a
353 single representative GMPE for the Himalayan region; however, different ground-
354 motion models can predict significantly different PGA values at a given distance-
355 magnitude. Parajuli et al. (2010) addressed this by using five attenuation relation-
356 ships and averaging. They found that Kathmandu's hazard could be quite high
357 (they reported 0.5g PGA for 475-year period on soft rock). Our result for Kath-
358 mandu ($\sim 0.35\text{--}0.4\text{g}$ on rock) might have been higher if we had included models
359 that predict higher shaking or if site amplification in the valley's soil was consid-
360 ered. We focused on rock site PGA; actual shaking in the basin could be amplified
361 by a factor of 1.5 to 2 at certain frequencies due to deep sediments, which is an
362 aspect beyond the scope of this study but critical for urban seismic risk. The use of
363 different GMPEs also influences the geographic distribution of hazard; some mod-
364 els might predict slower attenuation (thus higher hazard further south into the
365 plains). In future assessments, a logic-tree approach with multiple GMPEs would
366 provide a range (and median) of hazard estimates, increasing the robustness of
367 conclusions.

368 **4.4 Implications for Building Codes and Risk Mitigation**

369 Our updated hazard assessment for central Nepal carries significant implications
370 for engineering design and public policy. The current Nepal National Building
371 Code (NBC 105:1994) zone factors were based on older studies that gave PGA val-
372 ues of 0.08–0.12g in much of central Nepal. Our findings suggest that design lev-
373 els should be higher – in the range of 0.3g or more for important structures in
374 Kathmandu and surrounding high-hazard areas – to achieve a comparable level
375 of safety (10% probability of exceedance in 50 years). This corroborates calls by
376 previous researchers to revise the existing hazard estimates and code provisions
377 in Nepal. Incorporating new hazard maps into the building code will ensure that
378 structures are designed with appropriate lateral force levels for the actual seismic
379 threat. Moreover, microzonation within the Kathmandu Valley (accounting for
380 local site effects) should be undertaken using our hazard results as a base, to bet-
381 ter guide construction practices on soft sediments. Aside from engineering, the
382 hazard map is valuable for land-use planning: areas with extremely high hazard
383 might avoid critical infrastructure or require special reinforcement. Disaster pre-

384 paredness and response planning can also be informed by understanding which
385 areas are likely to experience the strongest shaking in future scenario earthquakes.

386 **4.5 Limitations and Future Work**

387 While comprehensive, this study has limitations that should be acknowledged.
388 We treated the entire central Nepal Himalaya as a single area source with av-
389 eraged properties. In reality, the seismic source characterization could be made
390 more physically-based by incorporating individual fault segments (e.g., differenti-
391 ating the locked MHT segment that ruptured in 2015 vs. the adjacent segment that
392 ruptured in 1934, etc.) and assigning activity rates to each. Due to limited time
393 and data, we did not explicitly model fault-specific recurrence (e.g., characteris-
394 tic earthquakes on known faults), which could refine the hazard locally. We also
395 assumed a time-invariant (Poisson) model; however, after a large event like 2015,
396 there could be a temporary reduction in hazard on that fault segment and an in-
397 creased stress (and hazard) on the unruptured segments immediately to the west
398 and east. Time-dependent PSHA models (which were beyond our scope) could
399 explore how hazard might transiently decrease or increase after big earthquakes.
400 Additionally, our hazard assessment did not explicitly account for induced seis-
401 micity or earthquakes in the Indo-Gangetic plains (which are low, but possibly
402 not zero). Despite these simplifications, the broad results are robust for regional
403 planning purposes. Future work should aim to incorporate the latest research on
404 Himalayan earthquake recurrence intervals (for example, geological rupture evi-
405 dence) and to use a logic-tree PSHA approach that encompasses multiple models
406 of seismic source and ground motion. Regular updates to the hazard model are
407 recommended as more earthquake data are recorded and as seismic science ad-
408 vances. This will ensure that Nepal's seismic hazard assessments remain accurate
409 and useful for mitigating earthquake risk.

5 CONCLUSION

410 This study provided a detailed probabilistic seismic hazard analysis for the Cen-
411 tral Nepal Himalayan region, leveraging an updated earthquake catalog and con-
412 temporary methodologies. The analysis yields an estimated b-value of 0.9 for
413 the regional Gutenberg-Richter relation, indicating a slightly higher propensity for
414 larger-magnitude earthquakes relative to smaller ones in the seismicity of central

415 Nepal. Using these recurrence parameters along with appropriate ground motion
416 models, we calculated the expected levels of ground shaking for a 10% probability
417 of exceedance in 50 years. The resulting hazard map demonstrates that much of
418 central and eastern Nepal faces very high seismic hazard, with peak ground accel-
419 erations on the order of 0.4–0.5g possible in the Kathmandu area and surrounding
420 highlands over a 475-year return period. Even the more “moderate” hazard areas
421 in the region have design-level PGAs of 0.2–0.3g, which are significant enough to
422 warrant serious engineering consideration.

423 Comparing our results with earlier studies, we find overall agreement that the
424 Himalayan frontal region is exceptionally hazardous, though our updated model
425 suggests slightly higher hazard in central Nepal, likely due to the inclusion of re-
426 cent seismic data and improved ground-motion predictions. These findings rein-
427 force the necessity for updating building codes in Nepal to reflect current hazard
428 levels and for enforcing those codes to ensure structures can withstand the pre-
429 dicted shaking. Implementing our hazard findings in urban planning (for exam-
430 ple, avoiding critical facilities in the highest hazard zones, or retrofitting vulne-
431 rable structures) could greatly reduce future earthquake losses. The 2015 Gorkha
432 earthquake was a wake-up call that, despite being slightly less severe than the
433 worst-case scenarios, caused tremendous damage and loss of life; our hazard as-
434 sessment indicates that similar or stronger shaking is plausible and should be an-
435 ticipated in resilience planning.

436 In conclusion, the PSHA of central Nepal highlights a continued high risk from
437 earthquakes in the region. By synthesizing geological, seismological, and statis-
438 tical data, we have provided a comprehensive picture of the threat. This study
439 contributes to the scientific understanding of Himalayan seismic hazard and of-
440 fers practical inputs for risk reduction efforts. As new data emerge and methods
441 evolve, such hazard assessments should be periodically revised. Ongoing research
442 – including paleoseismic studies, dense seismic instrumentation, and advanced
443 modeling – will further refine these estimates. Nonetheless, the evidence is clear
444 that central Nepal requires rigorous earthquake preparedness. The combination of
445 dense populations and high hazard makes this region one where proactive mea-
446 sures (education, emergency planning, resilient construction) are urgently needed.
447 By adopting policies informed by studies like this, Nepal can improve its resilience
448 and reduce the potential impact of the inevitable future earthquakes in the Hi-
449 malaya.

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