Infrared- and Raman-Active Phonons of Magnetite, Maghemite, and Hematite: A Computer Simulation and Spectroscopic Study

Irina Chamritski*,†,‡ and Gary Burns†

School of Chemical and Physical Sciences, Victoria University of Wellington, P.O. Box 600, Wellington, New Zealand, and Materials Performance Technologies, Lower Hutt, New Zealand

Received: March 21, 2004; In Final Form: November 29, 2004

Magnetite, maghemite, and hematite have been the subject of numerous studies using vibration spectroscopy to determine their infrared- and Raman-active phonons. However, no complete and unambiguous set of experimentally observed optically active phonons has yet been reported for these iron oxides. The use of atomistic simulation methods with a transferable Buckingham potential provides new data for the phonon densities of states of magnetite and the two associated phases, hematite and maghemite.

1. Introduction

The unusual properties of magnetite have spurred a continued effort toward elucidating the processes of its formation as a passive film,¹ the occurrence of magnetite on earth and planetary surfaces,² and the nature of the Verwey transition at 120 K.³

Central to most of these studies is the use of vibration spectroscopy to determine the phonon density of states. There have been numerous studies of the vibration spectra of magnetite, with the majority employing infrared absorption and Raman scattering^{4–9} to assign the four T_{1u} infrared-active phonons and the A_{1g} , E_g , and $3T_{2g}$ Raman-active modes, respectively.

There is some confusion over the assignments of the phonons for magnetite, particularly the Raman-active phonons. De Faria et al. have shown that laser power is an important consideration when interpreting Raman data. Heating or moderate laser power converts magnetite to hematite (α -Fe₂O₃) and maghemite (γ -Fe₂O₃), and some of the earlier Raman data shows evidence of attributing phonons of hematite or maghemite impurity to the parent magnetite phase. A recent study of the effects of laser-induced thermal effects on magnetite has confirmed the findings of de Faria et al.

Absence of a complete and unambiguous set of observed optically active phonons has led us to use atomistic simulation methods to obtain the phonon density of states for magnetite and the two associated phases, hematite and maghemite.

2. Atomistic Simulation

The simulation techniques have been described comprehensively elsewhere¹¹⁻¹⁴ and have been previously applied to a range of mixed metal oxides.¹⁵⁻¹⁹ The calculations were carried out using energy minimization procedures embodied in the GULP code. The interatomic potentials are based on the Born model of the solid, which includes a long-range Coulombic

interaction and a short-range term to model overlap repulsions and van der Waals forces

$$V_{ii}(r_{ii}) = [-Z_i Z_i e^2/r] + [A_{ii} \exp(-r/\rho_{ii}) - C/r_{ii}^6]$$
 (1)

The electronic polarizability of the ions is described by the shell model, which is effective in simulating the dielectric and lattice dynamical properties of metal oxides. It should be stressed that employing such a potential model does not necessarily mean that the electron distribution corresponds to a fully ionic system and that the validity of the potential model is assessed primarily by its ability to reproduce observed crystal properties. In practice, it is found that models based on formal charges work well even for compounds such as silicates and zeolites in which there is a significant degree of covalency.

The simulations of the iron oxide lattices were performed using two-body potentials for Fe³+-O, Fe²+-O, and O²--O²- interatomic interactions that have been developed by Catlow et al. $^{20-22}$ The shell model parameters interactions were refined to achieve the best fit of the predicted wavenumbers of infrared-and Raman-active vibrations to experimentally observed wavenumbers. The short-range potential parameters for each ion—ion interaction and shell model parameters for Fe₃O₄, γ -Fe₂O₃, and α -Fe₂O₃ used in this study are listed in Tables 1, 4, and 7. The calculated lattice parameters and atomic positions after energy minimization to generate a relaxed equilibrium structure for studied iron oxides are reported in Tables 2, 5, and 8. The data show excellent agreement with the experimentally observed structures.

The phonon calculations presented here are based on standard methodology involving the Cartesian second derivative matrix. The vibrational frequencies are the square roots of the eigenvalues of the dynamical matrix, which consists of the mass-weighted second derivatives. Such methods have been used^{23–25} to calculate the phonon dispersion curves of oxides, such as $\alpha\text{-SiO}_2$, UO2, and La2CuO4, and to interpret the vibrational spectra of the important class of proton insertion materials based on $\lambda\text{-LiMn}_2\text{O}_4$. 26

3. Discussion

3.1. The Infrared- and Raman-Active Phonons for Magnetite. Gasporov et al.³ provide the most complete infrared and

^{*} Author to whom correspondence should be addressed. Present address: Measurement Standard Laboratory, P.O. Box 31-310, Lower Hutt, New Zealand. Fax: 64-4-9313117. E-mail: i.chamritskaia@irl.cri.nz.

[†] Victoria University of Wellington.

[‡] Materials Performance Technologies.

TABLE 1: Interatomic Potentials for Magnetite

Fe3+

 $\mathrm{Fe^{2+}}$

Short Range				
interaction	A (eV)	ρ (Å)	$C (\text{eV Å}^6)$	
Fe ³⁺ •••O ²⁻	1102.4	0.3299	0	
Fe^{2+} ···O ²⁻	694.1	0.3399	0	
Fe^{3+} ···· O^{2-}	1102.4	0.3299	0	
O ²⁻ ··· O ²⁻	22764	0.149	15	
Shell Model				
	<i>Y</i> (e)	k (eV Å ⁻²)		

TABLE 2: Calculated and Experimental Crystal Properties for Magnetite

304.7

10.92

	Unit Cell Parameters			
a (exp) Å	a (calcd) Å			
8.39	8.41			
Atomic Coordinates Fd3m				
	()	(aslad)		

species	xyz (exp)	xyz (calcd)
Fe ²⁺ Fe ³⁺ (16d)	0.625	0.625
Fe^{3+} (8a)	0.375	0.381
O^{2-} (32e)	0	0

TABLE 3: Calculated and Observed Wavenumbers for Magnetite with Intensity of Infrared-Active Phonons Shown in Parentheses

obs ³ /cm ⁻¹	$calcd/cm^{-1}$	sym species	activity	ions involved
210	231(0.05)	$T_{1u}(1)$	IR	A, B, O
193	241	$T_{2g}(1)$	Raman	A, O
308	296	$T_{2g}(2)$	Raman	O
270	326(2.0)	$T_{1u}(2)$	IR	A, B, O
350	360(0.06)	$T_{1u}(3)$	IR	A, B, O
	505	E_{1g}	Raman	B, O
540	581	$T_{2g}(3)$	Raman	A, O
560	491(0.6)	$T_{1u}(4)$	IR	A, B, O
670	666	A_{1g}	Raman	O

Raman measurements of magnetite in a study of the metal—insulator and structural transition (Verwey transition) at 120 K. Their temperature-dependent infrared and Raman data enable phonons attributable to the low-temperature orthorhombic insulator phase to be separated from those of the high-temperature spinel phase. They report polarized Raman spectra with the polarizations of the incident and scattered light parallel and perpendicular that discriminate clearly between the A_{1g} and the three T_{2g} modes. However, they do not observe the Raman-active E_g mode. Verble assigned a 420 cm $^{-1}$ band measured at 77 K as E_g , but it is more likely attributable to a band of the low-temperature orthorhombic phase. The infrared data of Gasparov et al. confirms earlier assignments for the two most intense infrared-active T_{1u} modes but does not identify the remaining two T_{1u} modes.

The phonon calculations for magnetite are based on the potential given in Table 1 that reproduces the cell parameters satisfactorily. The phonon calculations based on this potential are compared with the data of Gasparov et al. in Table 3. Agreement is generally satisfactory, confirming the assignments of Gasparov et al. for the A_{1g} and three T_{2g} modes and therefore giving some confidence in the calculated E_g phonon at 505 cm $^{-1}$.

There is also satisfactory agreement between the observed and calculated data for the infrared-active T_{1u} modes. The lowest wavenumber mode is calculated to be at 231 cm⁻¹, a value well above the assignment of Degiorgi et al., a result that supports the conclusion of Gasparov et al. that the two lowest wave-

TABLE 4: Interatomic Potentials for Maghemite

Short Range					
interaction	A (eV)	ρ (Å)	$C \text{ (eV Å}^6)$		
Fe ³⁺ •••O ²⁻	1102.4	0.329	0		
$Fe^{2+} \cdot \cdot \cdot O^{2-}$	694.1	0.340	0		
$Fe^{3+} \cdots O^{2-}$	1102.4	0.329	0		
$O^{2-} \cdots O^{2-}$	22764	0.143	43		
Shell Model					
Y (e) k (eV Å ⁻²)					
Fe ³⁺	4.97 805		805		
Fe^{2+}	2	2 301			
O^{2-}	-2.239		60		

TABLE 5: Calculated and Experimental Crystal Properties for Maghemite

Unit Cell Parameters			
a (exp) Å	a (calcd) Å		
8.351	8.455		
	Atomic Coordinates Pm3m		

species xyz (exp) x y z (calcd) Fe³⁺ 0.125 0.125 0.366 0.374 -0.1160.875 Fe3+ 0.6250.625 0.625 0.625 0.625 0.625 0.6250.625 Fe2+ 1.0 0.99 0.991.0 1.0 0.99 O²⁻ 0.386 0.391 0.386 0.391 0.386 0.391 O^{2-} 0.108 0.117 0.108 0.13 0.383 0.391

number T_{1u} modes observed by Degiorgi et al. are due to phonons of the orthorhombic phase.

There is a weak band observed in the 300 K optical conductivity data of Gasparov et al. that matches the calculated infrared-active phonon at 231 cm⁻¹, with a very small calculated intensity.

The two most intense T_{1u} modes observed at 350 and 560 cm⁻¹ can be matched with the phonons calculated at 326 and 491 cm⁻¹ with relative intensities of 2.1 and 0.6. The weak T_{1u} mode at 360 cm⁻¹ would not be resolvable at temperatures above the Verwey transition and forms a part of the envelope of $T_{1u}(2)$. Variation in the published infrared peak positions is probably due to this weak unresolved T_{1u} mode.

3.2. The Infrared- and Raman-Active Phonons for Maghemite. The magnetite Buckingham potential was used with the A and ρ parameters unchanged but with an increase in the oxygen C parameter. The shell model spring constants for Fe³⁺, Fe²⁺, and O²⁻ were adjusted, but the charges on the three ions were held at the values found for magnetite. The potential shown in Table 4 reproduces the a cell parameter for maghemite (Table 5) within less than 0.5%.

The more complex unit cell of maghemite compared with magnetite results in many more Raman- and infrared-active phonons. The observed and calculated phonon data are shown in Table 6. The inverse spinel lattice of maghemite is observed to be iron-deficient, and this results in observed infrared and Raman spectra that are dependent on the method of sample preparation. De Faria et al.⁹ report three broad bands in the

TABLE 6: Calculated and Observed Wavenumbers for Maghemite

obs ^{2,9,25} /cm ⁻¹	calcd/cm ⁻¹	sym species	activity	ions involved
	212	T ₂ (1)	IR	Fe+O
350	356	T_1	Raman	Fe+O
				broad structure
319	362	$T_2(2)$	IR	Fe+O
440	440	$T_2(3)$	IR	Fe+O
500	493	E	Raman	Fe+O
				broad structure
553	523	$T_2(4)$	IR	Fe+O
700	695	A_1	Raman	Fe+O
				broad structure

TABLE 7: Interatomic Potentials for Hematite

Short-Range					
interaction	A (eV)	ρ (Å)	C (eV Å ⁶)		
Fe ³⁺ ····O ²⁻ O ²⁻ ···· O ²⁻	Fe ³⁺ O^{2-} 1102.4 0.329 O^{2-} O^{2-} 22764 0.149		0 40		
Shell Model					
Y (e) k (eV Å $^{-2}$)					
Fe ³⁺ O ²⁻	4.97 -2.239	60.7 200			

TABLE 8: Calculated and Experimental Crystal Properties

	Unit Cell Parameters	
	exp	calcd
a (Å)	5.04	5.09
b (Å)	5.04	5.09
c (Å)	13.77	12.96
α (deg)	90	90
β (deg)	90	90
γ (deg)	120	120
	Atomic Coordinates \bar{R}	3c
	, ,	
species	xyz (exp)	xyz (calcd)
Fe ³⁺	<i>xyz</i> (exp) 0	xyz (calcd)
•		
•	0	
•	0 0	0 0
Fe ³⁺	0 0 0.355	0 0 0.359

Raman spectrum for their sample of maghemite, at \sim 350 cm⁻¹, \sim 500 cm⁻¹, and \sim 700 cm⁻¹. The calculated Raman-active T₁, E, and A_1 phonons are in good agreement with this observed data. The calculated infrared-active phonons at 442 and 524 cm⁻¹ are in good agreement with the observed infraredactive phonons reported by Gehring et al.²⁷ and Bell et al.² at 440 and 553 cm⁻¹. The lowest energy T₂ modes are calculated at 212 and 362 cm⁻¹.

3.3. The Infrared- and Raman-Active Phonons for Hematite. The Buckingham potential used for magnetite and maghemite was retained for hematite, but the spring constants used for Fe³⁺ and O²⁻ have been adjusted compared to the magnetite and maghemite values. The potential shown in Table 7 reproduces the a cell parameter to better than 1% and the c parameter to $\pm 6\%$ (Table 8).

The observed²⁸ and calculated Raman-active phonons for hematite are shown in Table 9. Agreement is good for the two A_{1g} and four lowest energy E_g modes but is poor for the highest wavenumber E_g mode calculated at 467 cm⁻¹ compared to the observed band at 613 cm⁻¹. The Buckingham potential parameters for magnetite do not adequately account for the hematite

TABLE 9: Calculated and Observed Wavenumbers for Hematite

${\rm obs^{26,27}/cm^{-1}}$	calcd/cm ⁻¹	sym species	activity	ions involved
225	228	A _{1g}	Raman	Fe,O
229	290	Eu	IR	Fe,O
247	200	E_{g}	Raman	Fe,O
293	294	E_{g}	Raman	Fe,O
299	327	E_{g}	Raman	Fe,O
310-335	310	$E_{\rm u}$	IR	Fe,O
310-335	339	A_{2u}	IR	Fe,O
380-400	368	E_{u}	IR	Fe,O
380-400	378	A_{2u}	IR	Fe,O
412	391	E_{g}	Raman	Fe,O
440 - 470	454	$E_{\rm u}$	IR	Fe,O
498	447	A_{1g}	Raman	Fe,O
613	466	E_{g}	Raman	Fe,O

lattice interactions along c, and this results in the anomalously low value for the calculated E_g phonon.

There is good agreement for the calculated infrared-active phonons when compared to the infrared data of Serna et al.²⁹

4. Conclusions

Through use of atomistic simulation, phonon energies and their eigenvectors have been calculated for the infrared- and Raman-active phonons of Fe₃O₄, γ -Fe₂O₃, and α -Fe₂O₃. Lowering the symmetry in the sequence $Fe_3O_4 > \gamma - Fe_2O_3 > \alpha - Fe_2O_3$ leads to an increase in the number of calculated and observed Raman-active phonons. This increase in the number of Raman fundamentals can explain the change in the Raman spectrum of Fe₃O₄ under heating when the following phase transitions have been shown to occur.

$$Fe_3O_4 \xrightarrow{200 \text{ °C}} \gamma$$
- $Fe_2O_3 \xrightarrow{400 \text{ °C}} \alpha$ - Fe_2O_3

Our calculations confirm the Raman results obtained by de Faria et al., postulating that the Raman spectra of Fe₃O₄ and γ -Fe₂O₃ are each unique. According to our calculations, the observed Raman phonons at similar wavenumber for Fe₃O₄ and γ -Fe₂O₃ have a different atomic origin. For example, each of the given iron oxides has a fundamental in the region 670-700 cm⁻¹. The calculated Cartesian displacements show that the A_{1g} mode of Fe₃O₄ (670 cm⁻¹) is a mode of oxygen ion displacement only, whereas the A_1 mode of γ -Fe₂O₃ (700 cm⁻¹) is characterized by displacements of both iron and oxygen.

Disagreement between Raman spectra reported for magnetite, maghemite, and hematite may in large part be attributed to the heating effect of the incident laser radiation. Secondary effects will be present due to the nonstoichiometry of iron oxides.

Acknowledgment. The authors wish to thank Dr. S. Islam for assistance with the potentials used for the calculations and Dr. J. Gale for allowing us to use the GULP code. The Foundation for Research, Science and Technology is thanked for a Bright Futures Scholarship awarded to Irina Chamritski.

References and Notes

- (1) Hendy, S.; Laycock, N.; Ryan, M.; Walker, B. Phys. Rev. B 67, **2003**, 085407.
- (2) Bell, J. F., III; Roush, T. L.; Morris, R. V. J. Geophys. Res. 1995, 100, 5297.
- (3) Gasparov, L. V.; Tanner, D. B.; Romero, D. B.; Berger, H.; Margaritondo, G.; Forro, L. Phys. Rev. B 2000, 62, 7939.
 - (4) Verble, J. L. Phys. Rev. B 1974, 9, 5236.
 - (5) Dunnwald, J.; Otto, A. Corros. Sci. 1989, 29, 1167.
- (6) Degiorgi, L.; Blatter-Morke, I.; Wachter, P. Phys. Rev. B 1987, 11, 5421.

- (7) Oblonsky, L.; Devine, T. Corros. Sci. 1995, 37, 17.
- (8) Watanabe, H.; Brockhouse, B. Phys. Lett. 1962, 1, 189.
- (9) De Faria, D. L. A.; Venancio Silva, S.; de Oliveira, M. T. J. Raman Spectrosc. 1997, 28, 873.
 - (10) Shebanova, O. N.; Lazor, P. J. Raman Spectrosc. 2003, 34, 845.
- (11) (a) Catlow, C. R. A. In Solid State Chemistry Techniques; Cheetham A. K., Day, P., Eds.; Clarendon Press: Oxford, 1987; Chapter 7. (b) Catlow, C. R. A.; Price, G. D. Nature 1990, 347, 243.
 - (12) Dick, B. G.; Overhauser, A. W. Phys. Rev. 1958, 112, 90.
- (13) Mott, N. F.; Littleton, M. J. *Trans. Faraday Soc.* **1938**, *34*, 485. (14) (a) Gale, J. D. *GULP* (General Utility Lattice Program); Royal Institution of GB and Imperial College: London, 1991-1996. (b) Gale, J. D. J. Chem. Soc., Faraday Trans. 1997, 93, 629.
- (15) Cherry, M.; Islam, M. S.; Catlow, C. R. A. J. Solid State Chem. **1995**, 118, 125.
- (16) Cherry, M.; Islam, M. S.; Gale, J. D.; Catlow, C. R. A. J. Phys. Chem. 1995, 99, 14614.
- (17) Islam, M. S.; Cherry, M.; Winch, L. J. J. Chem. Soc., Faraday Trans. 1996, 92, 479.
- (18) Wu, J. W.; Tepesch, P. D.; Ceder, G. Philos. Mag. B 1998, 77, 1039.

- (19) Balducci, G.; Kaspar, J.; Fornasiero, P.; Graziani, M.; Islam, M. S.; Gale, J. D. J. Phys. Chem. B 1997, 101, 1750.
 - (20) Lewis, G. V.; Catlow, C. R. A. J. Phys. C 1985, 18, 1149.
- (21) Donnerberg, H.; Exner, M.; Catlow, C. R. A. Phys. Rev. B 1993, 47, 14.
- (22) Woodley, S. M.; Battle, P. D.; Gale, J. D.; Catlow, C. R. A. Phys. Chem. Chem. Phys. 1999, 1, 2535.
 - (23) Schroder, K.-P.; Sauer, J. J. Phys. Chem. 1996, 100, 11043.
- (24) Jackson, R. A.; Murray, A. D.; Harding, J. H.; Catlow, C. R. A. Philos. Mag. A 1986, 53, 27.
- (25) Islam, M. S.; Leslie, M.; Tomlinson, S. M.; Catlow, C. R. A. J. Phys. C 1988, 21, L109.
- (26) Ammundsen, B.; Burns, G.; Saiful Islam, M.; Kanoh, H.; Roziere, J. J. Phys. Chem. B 1999, 103, 5175.
 - (27) Gehring, A.; Hofmeister, A. Clays Clay Miner. 1994, 42, 409.
 - (28) Beattie, I. R.; Gilson, T. R. J. Chem. Soc. 1970, A5, 980.
- (29) Serna, C. J.; Rendon, J. L.; Iglesias, J. E. Spectrochim. Acta 1982, *38*, 797.