# Structure and Desorption Energetics of Ultrathin D<sub>2</sub>O Ice Overlayers on Serine- and Serinephosphate-Terminated Self-Assembled Monolayers

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This paper reports on the structure and desorption dynamics of thin  $D_2O$  ice overlayers (0.2–10 monolayers) deposited on serine- and serinephosphate- (with  $H^+$ ,  $Na^+$ ,  $Ca^{2+}$  counterions) terminated self-assembled monolayers (SAMs). The  $D_2O$  ice overlayers are deposited on the SAMs at  $\sim$ 85 K in ultrahigh vacuum and characterized with infrared reflection absorption spectroscopy (IRAS). Reflection absorption (RA) spectra obtained at sub-monolayer  $D_2O$  coverage reveal that surface modes, e.g. free dangling OD stretch, dominate on the serine SAM surface, whereas vibrational modes characteristic for bulk ice are more prominent on the serinephosphate SAMs. Temperature programmed desorption mass spectrometry (TPD-MS) and TPD-IRAS are subsequently used to investigate the energetics and the structural transitions occurring in the ice overlayer during temperature ramping.  $D_2O$  ice ( $\sim$ 2.5 monolayers) on the serine SAMs undergoes a gradual change from an amorphous- to a crystalline-like phase upon increasing the substrate temperature. This transition is not as pronounced on the serine phosphate SAM most likely because of reduced mobility due to strong pinning to the surface. We show also that the energy of desorption for a sub-monolayer of  $D_2O$  ice on serinephosphate SAM surfaces with a  $Na^+$  and  $Ca^{2+}$  counterions is equally high or even exceeds previously reported values for analogous high-energy SAMs.

#### Introduction

Phosphates play a crucial role in biology, both in the form of energy-rich groups in molecules such as ATP, as a part of the backbone structure of DNA, as well as an important component in bone together with calcium. This particular material, calcium hydroxyapatite (HA), has been studied extensively within the biomaterials science community, because of its importance for the design and development of new implantable devices, joints, and prostheses. It is, for example, of interest to develop coating processes that can improve the adhesion and integration between the biomaterial and bone. A large number of studies of the mineralization processes of calcium phosphates including work in suspension<sup>1-4</sup> and on solid supports<sup>5-7</sup> have been published. Moreover, Küther et al.<sup>8</sup> utilized self-assembled monolayers to investigate the nucleation and growth of CaCO<sub>3</sub> crystals. We recently synthesized a range of organic phosphates for self-assembly on gold.<sup>9</sup> The purpose was to develop surface coating that could be used as a template for mineralization, e.g. of HA, and to produce a set of biomimetic surfaces for fundamental studies of reactions where phosphates are involved. 10,11

We have for a long time been interested in the structure and properties of water, in particular, interfacial water. This type of water is analogous to the water present in hydration layers around ions and organic structures both in solution and at surfaces. <sup>12–14</sup> Interfacial water also plays an important role during mineralization, <sup>15,16</sup> and the structure of water close to HA, collagen, and a mixture thereof has been investigated. <sup>17</sup>

The aim of the present study is to investigate interfacial water in the form of ultrathin ice overlayers on organic surfaces. Such studies have been undertaken before at our laboratory as well as by others. <sup>18–24</sup> The main emphasis is on the phase behavior and energetics of ice on surfaces that mimic the structure of HA. These surfaces have been prepared on gold using thiols with serine and serinephosphate tail groups, <sup>9</sup> Scheme 1. The ice overlayers are studied at low temperature (85–250 K) on the serine and serinephosphate (H<sup>+</sup>, Na<sup>+</sup>, and Ca<sup>2+</sup>) SAMs using temperature programmed desorption infrared reflection spectroscopy (TPD-IRAS) and mass spectrometry (TPD-MS).

## **Experimental Section**

Materials and Sample Preparation. The serine and serinephosphate derivatives were synthesized following recently published protocols.9 The substrates used for incubation were 2000 Å thick electron-beam evaporated gold films on top of freshly cleaned and precut (100) silicon slides that were precoated with a 25 Å adhesive layer of titanium. All evaporations took place in a Balzers UMS 500 P system operating at a base pressure of  $\sim 10^{-9}$  Torr and an evaporation pressure of  $\sim 10^{-7}$  Torr. The evaporation rate was 10 Å s<sup>-1</sup> for gold and 1 Å s<sup>-1</sup> for titanium. Prior to incubation in the ethanol thiol solutions, the samples were thoroughly cleaned in a 5:1:1 mixture of distilled water:25% hydrogen peroxide:30% ammonia for 10 min at 80 °C. The samples were incubated for a minimum of 15 h in 1 mM ethanol solutions of the thiols. The samples were rinsed with ethanol after incubation, ultrasonicated in ethanol for 3 min to remove physisorbed thiols, rinsed in ethanol again, and finally blown dry in N<sub>2</sub>. In some cases the H<sup>+</sup> counterion was replaced with a different counterion (Na<sup>+</sup>, Ca<sup>2+</sup>); this procedure is described elsewhere. 9 Before insertion in the introduction chamber of the ultrahigh-vacuum system was the

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SCHEME 1: Molecules and Counterions Used for Creating the Self-Assembled Monolayer Substrates

sample surface blown dry with nitrogen gas. The outcome of the self-assembly process in terms of monolayer thickness and wetting characteristics was immediately checked on twin samples using a Rudolph Research AutoELIII ellipsometer and a Ramé-Hart NRL 100 goniometer.

**Ultrahigh-Vacuum System and TPD.** The TPD-MS and temperature programmed desorption infrared reflection—absorption spectroscopy (TPD-IRAS) measurements were undertaken in a home-built ultrahigh-vacuum system described in detail elsewhere. The base pressure in the measurement chamber was  $5 \times 10^{-10}$  Torr and in the preparation chamber  $5 \times 10^{-9}$  Torr. The TPD-MS traces were obtained with a Hiden HAL 301 positive ion counting quadropole analyzer mounted in front of the sample in the measurement chamber. The sample temperature (85–250 K) was monitored with a Pt 100 element and controlled with a Eurotherm controller/programmer, Type 818. Deuterated water (Merck) was deposited onto the sample surface at 85 K via a capillary tube facing the sample surface. The heating rate used in the TPD experiments was 0.33 K/s.

The infrared measurements were made with a Bruker IFS 22 spectrometer, aligned at an angle of incidence of 83° and equipped with a narrow band liquid nitrogen cooled MCT detector. The reflection absorption (RA) spectra were recorded at a spectral resolution of 2 cm<sup>-1</sup>. A low noise background spectrum was always recorded at 85 K before each deposition of  $D_2O$  by averaging 500 interferograms ( $\sim$ 15 min acquisition time). After the D<sub>2</sub>O deposition, but prior to the TPD-MS, another RA spectrum was recorded in order to determine the coverage and structure of the deposited D<sub>2</sub>O ice overlayer at 85 K. A previously developed infrared method was used to calculate the equivalent thickness and coverage (monolayers) of the deposited D<sub>2</sub>O ice overlayer.<sup>21</sup> A new RA spectrum was recorded every 20 s (20 scans/spectrum) during the TPD-MS scan. Thus, each RA spectrum represents approximately a 7 K interval. These dynamic measurements were performed to monitor the structural transitions occurring prior to and during the desorption event. The TPD-MS scan was terminated at  $\sim$ 250 K for each measurement, after which the sample was cooled to 85 K, and a RA spectrum was recorded to check whether the surface was irreversibly affected by the adsorption/desorption of the D<sub>2</sub>O ice layer. If the surface was unaffected, another measurement cycle was performed at another coverage using the same surface.

#### **Results and Discussion**

Structure of Deposited  $D_2O$  Ice. The  $D_2O$  molecule displays three normal modes of vibration—the asymmetric and symmetric O-D stretching and the symmetric, scissoring, O-D deforma-

tion. For a free D<sub>2</sub>O molecule in a vacuum these modes appear at 2789, 2666, and 1179 cm<sup>-1</sup>, respectively.<sup>25</sup> However, upon interaction with other D<sub>2</sub>O molecules, e.g., in liquid water or ice, the vibrational modes are heavily affected by hydrogen bond interactions, and they shift to lower frequencies and appear as broad features. Bulk D<sub>2</sub>O absorbs in two regions. One peak appears around 1200 cm<sup>-1</sup> and another one in the 2250–2750 cm<sup>-1</sup> region. This latter peak contains features that can be used to extract information about the phase behavior and surface chemistry of the ice overlayers, and these topics have been thoroughly discussed in the literature. 19,20,24,26-31 In a perfect crystal of hexagonal D<sub>2</sub>O ice, for example, each molecule forms hydrogen bonds with four other molecules. This tetragonal symmetry is broken at the surface of a crystal or at imperfections resulting in a reduction of the coordination number and the appearance of new vibrations, so-called surface modes. Four surface modes have been observed of which three are relatively easy to locate at  $\sim$ 2725–2740 cm<sup>-1</sup> (free O–D stretching),  $\sim$ 2640 cm<sup>-1</sup> (free O orbital lobe), and  $\sim$ 2580 cm<sup>-1</sup> (four coordinated, distorted D<sub>2</sub>O), respectively.<sup>32</sup> Moreover, previous reports show that small clusters of D<sub>2</sub>O ice (2-10 molecules per cluster) display a similar set of narrow surface modes.<sup>33,34</sup> It is also observed that small clusters with a large surface-tovolume ratio tend to shift the surface modes toward high frequencies. 33,35,36 The exact frequencies of such surface modes can also be tuned by coating the ice cluster with molecules<sup>28,31</sup> or by depositing ultrathin ice overlayers onto chemically modified surfaces, e.g., on surfaces of varying wettability properties. 19 Thus, it can be concluded that surface modes appear at higher frequencies as compared to those observed for fully four-coordinated hexagonal  $(I_h)$  D<sub>2</sub>O ice, and as a rule of thumb, D<sub>2</sub>O ice modes appearing above 2520 cm<sup>-1</sup> are attributed to surface modes and those below 2520 cm<sup>-1</sup> belong to bulk (3D) modes.32,37

The serine thiol, Scheme 1, was used as a reference molecule in our experiment, since the molecule has the same backbone as the phosphate derivative. The molecule has been shown to form high-quality monolayers on gold with a thickness of 9 Å, approximately half of the thickness of a monolayer of phosphate analogue, 18 Å. This large difference in thickness is due to the bulkiness of the phosphate group. The contact angles with water for the serine and serinephosphate SAMs are 26 and <10°, respectively. The phosphate can exist as a charged or a neutral group, as well as a mixture thereof, on the SAM surface, depending on the nature of solvents and buffers used. We therefore equilibrated the serinephosphate SAMs prior to D<sub>2</sub>O deposition in solutions containing an excess of counterions H<sup>+</sup>, Na<sup>+</sup>, and Ca<sup>2+</sup>. The different counterions affect the appearance of the RA spectra mainly in the phosphate (P-O) stretching region, and these interactions have been discussed before.<sup>9</sup>

The  $D_2O$  molecules are deposited onto the substrate surface at  ${\sim}85$  K, a temperature that severely limits the lateral mobility. Thus, the  $D_2O$  molecules are assumed to stick to the substrate at the point where they hit the surface. The structure of the  $D_2O$  ice layer at low coverage is therefore mainly affected by interaction with the SAM surface. Upon increasing the surface coverage, the probability for a  $D_2O$  molecule to adsorb at an empty site on the substrate decreases while the probability of  $D_2O$  adsorbing on top of other  $D_2O$  molecules already residing to the surface increases. This gives a hint that the main effects on the  $D_2O$  ice overlayer will occur in the submonolayer regime. However, earlier studies suggest that the effect of the surface may extend into several adlayers of  $D_2O$ , depending on the nature and strength of the  $D_2O$ —surface interaction.  $^{18,24}$ 

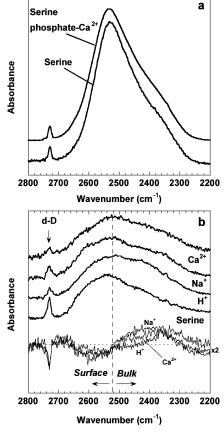
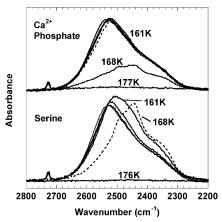


Figure 1. RA spectra of (a)  $\sim$ 2.5 monolayers D<sub>2</sub>O adsorbed to the serine analogue surface and to the serinephosphate analogue surface with  $Ca^{2+}$  as a counterion (b)  $\sim 0.2$  monolayers  $D_2O$  adsorbed to the serine analogue surface and serinephosphate with H<sup>+</sup>, Na<sup>+</sup>, and Ca<sup>2+</sup> as a counterion, respectively. For clarity, the difference spectra, where the spectrum has been subtracted from the H<sup>+</sup>, Na<sup>+</sup>, and Ca<sup>2+</sup> spectra are also displayed. The d-D (free O-D stretch) peak is marked as well as the approximate boundary between surface and bulk modes (---); see text for details.

In this study are the RA spectra for thick D<sub>2</sub>O overlayers (>2 monolayers) close to independent of the nature of surfaces, cf. the RA spectra obtained for  $\sim$ 2.5 monolayers of  $D_2O$  ice on the serine and serinephosphate-Ca<sup>2+</sup> SAMs, Figure 1a. A comparison with previous modeling of RA spectra suggests that amorphous-like ice is formed.<sup>21</sup> At low D<sub>2</sub>O coverage (less than a full monolayer), however, there is a clear difference in the D<sub>2</sub>O ice structure on the serine and serinephosphate surfaces. Figure 1b displays the RA spectra of  $\sim$ 0.2 monolayers ( $\sim$ 0.8 Å) D<sub>2</sub>O deposited onto the four different surfaces at 85 K. The difference is especially pronounced when comparing the spectra of serine and serinephosphate, whereas the difference is less pronounced when comparing the spectra of the serinephosphate with different counterions. The RA spectra of D<sub>2</sub>O deposited onto the serinephosphate surfaces display relatively less intense surface modes than D<sub>2</sub>O deposited onto serine. This is clearly displayed in the spectra as an overall increase in the intensity below 2520 cm<sup>-1</sup> and a decrease in the intensity above 2520 cm<sup>-1</sup>, where the decrease of the non-hydrogen-bonded O-D vibration at 2738 cm<sup>-1</sup> is particularly apparent; see also the difference spectra (lower panel) in Figure 1b. The difference is probably due to the presence of the highly polar phosphate entity that possesses both hydrogen bond donating and accepting groups. These groups can readily replace the missing hydrogen bonds in the ice overlayer, resulting in an enhanced intensity of the bulk D<sub>2</sub>O modes (below 2520 cm<sup>-1</sup>) and a reduced



**Figure 2.** Dynamic IRAS of  $\sim$ 2.5 monolayers D<sub>2</sub>O adsorbed to the Ca<sup>2+</sup> serine phosphate analogue (top) and serine analogue substrate (bottom). Each spectrum is averaged over a 7 K temperature span, since the surface temperature increases linearly during measurement.

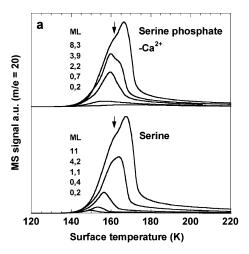
intensity of the vibrations originating from noncoordinated D<sub>2</sub>O bonds (surface modes). In the case of D<sub>2</sub>O being deposited to the serinephosphate with Ca<sup>2+</sup> as a counterion, the free O-D vibration is the weakest for the four surfaces, suggesting that also the counterion plays an important role for the coordination (hydration) of D<sub>2</sub>O molecules to serinephosphate groups. Thus, it is from this way of reasoning plausible to assume that the D<sub>2</sub>O ice overlayer interacts more vigorously with the serinephosphate SAMs than with the corresponding serine SAM (see below).

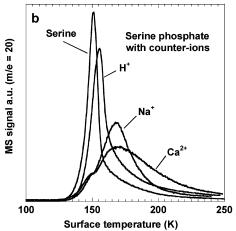
Desorption Dynamics. The structure of the D<sub>2</sub>O ice overlayers may also be temperature-dependent, which can be explored with dynamic IRAS. The sampling time during dynamic IRAS is relatively short, and the signal strength of a submonolayer of D<sub>2</sub>O is weak. This makes it difficult to monitor the structural differences for D2O coverage of less than 1 monolayer because of high noise levels in the RA spectra. But for a  $D_2O$  coverage exceeding  $\sim$ 2 monolayers the S/N ratio is acceptable and allows for dynamic investigations. Figure 2 displays the dynamic RA spectra for approximately 2.5 monolayers of D<sub>2</sub>O deposited onto the serine and serinephosphate-Ca<sup>2+</sup> SAMs, respectively. When comparing these spectra to the low-coverage spectra in Figure 1b, the difference in the surface to bulk mode intensity ratios is apparent, especially the intensity of the d-D peak at  $\sim 2730$  cm<sup>-1</sup> compared to the main peak. As the sample temperature is increased the D<sub>2</sub>O layer deposited onto the serine SAM changes gradually its structure from amorphous- to crystalline-like ice before it finally disappears at about 168-176 K. Worth noticing is that the shift in the spectrum toward lower frequencies is consistent with the formation of a 3D bulk ice structure, i.e., hexagonal  $I_h$  ice.<sup>24</sup>

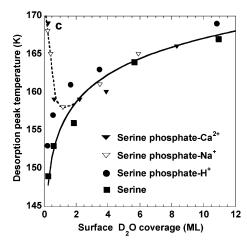
Moreover, the free O-D stretching peak disappears as the amorphous ice undergoes transition to crystalline-like ice. Thus, our data suggest that the interaction with the substrate is relatively weak, allowing the ice to rearrange and form a new structure. The D<sub>2</sub>O ice overlayer on the serinephosphate—Ca<sup>2+</sup> SAM surface does not undergo the same gradual change in structure as observed for serine. The only visible change in the RA spectrum appears very late during the desorption process, i.e., when a substantial amount of the ice layer already has desorbed. The delay in the phase transition is most likely a direct consequence of the stronger interaction between the D2O ice and the serinephosphate-Ca<sup>2+</sup> surface. This interaction influences the lateral mobility of the D2O molecules in intimate contact with the serinephosphate—Ca<sup>2+</sup> and prevents indirectly the subsequent adlayers of D2O from changing structure into crystalline-like  $I_h$  ice. Thus, pinning to the substrate influences to a large extent the phase behavior of thin ice layers, a phenomenon that has been observed before.<sup>21,37</sup>

To increase the understanding of the interaction between D<sub>2</sub>O and the SAM surfaces, TPD-MS traces were recorded in parallel to the dynamic RA spectra. The binding strength (energy of desorption) of the D2O molecules deposited at the surface at low coverages can thereby be determined. TPD-MS can also be utilized to identify different modes of adsorption on the SAM surfaces. Figure 3a shows the MS signal for D<sub>2</sub>O desorbing from serine and serinephosphate SAM, here with Ca<sup>2+</sup> as a counterion, for  $D_2O$  coverage between 0.2 and  $\sim 11$  monolayers. As the coverage increases (>1 monolayer), the peak desorption temperature increases continuously with coverage, suggesting that desorption of D<sub>2</sub>O from both the serine and serinephosphate surfaces follows zero-order kinetics. The mass trace obtained at high coverage on both the serine and serinephosphate—Ca<sup>2+</sup> displays a single broad peak most likely originating from a multitude of energetically narrow sites on the surface. At high coverage there is also a slight irregularity (shoulder) in the MS trace, a phenomenon that has been attributed to reflect the transition from amorphous- to polycrystalline-like ice. 38,39 The shoulder in the traces most likely appears because crystalline D<sub>2</sub>O ice has a lower desorption rate than amorphous D<sub>2</sub>O.<sup>40</sup> The transition from amorphous- to crystalline-like D<sub>2</sub>O ice is also clearly visible in the dynamic RA spectra for D<sub>2</sub>O deposited onto serine for a  $D_2O$  coverage of  $\sim 2.5$  monolayers, Figure 2. This transition is barely visible in the RA spectra for lower coverage (data not shown), suggesting that the D<sub>2</sub>O ice layers must exceed a critical thickness (~2 monolayers/6 Å) in order to adopt a structure characteristic for crystalline ice.

The largest effects on the desorption behavior are seen in the TPD-MS traces recorded at low D2O coverage on the different SAM surfaces. The TPD-MS traces, obtained at a constant D<sub>2</sub>O coverage of 0.2  $\pm$  0.02 monolayer (0.8  $\pm$  0.08 Å), are displayed in Figure 3b. As can be seen, changing the tail group and/or the counterion affects the desorption behavior of the D<sub>2</sub>O ice overlayer to a varying degree. For example, the desorption peak temperature and shape are not very different when comparing the trace of D<sub>2</sub>O desorbing from serinephosphate, in the form of a free base (H<sup>+</sup>), with that obtained for serine. Moreover, each desorption trace consists of a sharp and asymmetric peak with a high-temperature tail. This tail originates partly from the limited pumping capacity of the vacuum system, which gives rise to a slight increase in the background partial pressure of D<sub>2</sub>O in the chamber. Exchange of counterions to Na<sup>+</sup> or Ca<sup>2+</sup> increases the desorption peak temperature as well as the width of the MS peaks. The broadening of the MS peak is especially pronounced in the case of D<sub>2</sub>O desorbing from the serinephosphate—Ca<sup>2+</sup> surface. Thus, it seems reasonable to conclude that this broad peak is due to desorption from several unresolved and narrow sites on the serinephosphate-Ca<sup>2+</sup> surface. One can of course argue that the high-temperature tail in the Ca<sup>2+</sup> phosphate trace is due to the limited pumping capacity of the ultrahigh-vacuum system. However, we find this explanation less likely because the high-temperature tail is far more pronounced than expected for such a low coverage. If the tail solely appears because of the limited pumping capacity of the system, then we also expect to observe a tail of similar shape in the trace obtained from the Na+ phosphate surface (same coverage and same desorption peak temperature). This is definitely not the case. Instead we observe a symmetric desorption peak, without high-temperature tail, for the Na<sup>+</sup> sample, Figure 3b. Thus, we conclude that the high-temperature







**Figure 3.** MS traces of (a)  $D_2O$  of different monolayer (ML in the figure) coverage desorbing from the  $Ca^{2+}$  serinephosphate analogue and serine analogue substrate. The arrow points out the irregularity originating from the amorphous to polycrystalline effect discussed in the text. (b)  $\sim 0.2$  monolayer  $D_2O$  desorbing from serine and serine-phosphate analogues with  $H^+$ ,  $Na^+$ , and  $Ca^{2+}$  as a counterion, respectively. The  $Ca^{2+}$  and serine trace are displayed also in a. (c) MS peak temperature plotted as a function of surface coverage in monolayers. Lines are inserted as a guide to the eye.

tail in the Ca<sup>2+</sup> trace appears because of specific interaction(s) with the surface (e.g., with yet unidentified sites). One possible way to shed some light on the nature of the "unidentified" sites is to vary the temperature ramping rate. This type of experiment, however, is beyond the scope of the present contribution. Interestingly, the shape of the trace resembles that obtained for another system of strongly bound water on carboxylic acid

SAMs.<sup>22</sup> There is in both the Na<sup>+</sup> and Ca<sup>2+</sup> serinephosphate traces a weak shoulder at  $\sim$ 150 K. This peak arises from D<sub>2</sub>O weakly bound to other D<sub>2</sub>O molecules and not directly to the phosphate surface, and it grows in intensity with increasing D<sub>2</sub>O coverage and becomes eventually the dominant feature of the MS trace; see Figure 3a. Figure 3c summarizes the desorption peak temperatures as a function of coverage (0–12 monolayers) for the serine and serinephosphate surfaces displayed in Figure 3a,b. The desorption peak temperature increases continuously with coverage for all surfaces at coverages above ~2 monolayers. It is also interesting to note that the peak desorption temperatures obtained for each coverage (>2 monolayers) are essentially the same ( $\pm 3$  K) regardless of the chemical nature of the surface. Furthermore, the peak desorption temperatures obtained for the serinephosphate—H<sup>+</sup> is almost identical to those observed for the serine (OH) surface for all investigated coverages. However, the desorption peak temperature of these two surfaces, at a coverage of 0.2 monolayer, differ substantially from the serinephosphate surfaces having Na<sup>+</sup> and Ca<sup>2+</sup> as counterions. Both the Na<sup>+</sup> and Ca<sup>2+</sup> phosphate display desorption peaks at about 170 K for a coverage of ~0.2 monolayers, a value that is  $\sim 20$  K ( $\Delta E_{\rm des} = \sim 6$  kJ/mol) higher than those obtained for the other two surfaces. The energy of desorption,  $E_{\rm des}$ , is estimated by assuming a first-order desorption model at these low coverages and by setting the preexponential factor to  $10^{13}$  s<sup>-1</sup> and the heating rate  $\beta$  to 0.33 K/s.<sup>41</sup> It is also interesting to note that the obtained values of  $E_{des}$  obtained here for the Na<sup>+</sup> and Ca<sup>2+</sup> phosphate surfaces of 45 and 46 kJ/mol, respectively, are considerably higher that those reported for analogous self-assembled monolayer systems, except for the  $E_{\rm des}$ reported for D<sub>2</sub>O desorbing from carboxylic acid SAM, another substrate also displaying strong hydrogen bonding properties.<sup>22</sup> The corresponding values for serinephosphate—H<sup>+</sup> and serine are 41 and 40 kJ/mol, respectively. This supports our TPD-IRAS observation that the D<sub>2</sub>O ice layer is strongly bound to the serinephosphate surfaces with Na<sup>+</sup> and Ca<sup>2+</sup> counterions and that this strongly bound layer influences the lateral mobility of the D2O molecules and thereby also the phase behavior of thicker ice layers, cf. those in Figure 2. The desorption peak temperatures obtained for the Na<sup>+</sup> and Ca<sup>2+</sup> phosphate surfaces decrease rapidly with increasing coverage and approach the desorption peak temperatures obtained for the two other surfaces in the 1-2 monolayers regime. Lateral interactions with other D<sub>2</sub>O molecules obviously become very prominent at these coverages. Although the D<sub>2</sub>O-D<sub>2</sub>O interactions are weaker than the D<sub>2</sub>O-phosphate interactions, they tend to dominate and become the main part of the MS signal at a coverage of  $\sim$ 2 monolayers, after which the MS traces of D<sub>2</sub>O desorbing from phosphate (H<sup>+</sup>, Na<sup>+</sup>, and Ca<sup>2+</sup>) and serine behave similarly.

### Conclusions

We have studied the interaction and phase behavior of D<sub>2</sub>O ice overlayers (0.2-11 monolayers) onto a number of model phosphate surfaces using infrared spectroscopy and mass spectrometry. These surfaces are, as a zero-order model, aimed at mimicking the properties of bone mineral, i.e., hydroxyapatite. Self-assembled monolayers with serine and serinephosphate tail groups with three different counterions H<sup>+</sup>, Na<sup>+</sup>, and Ca<sup>2+</sup> have been prepared on gold. The serinephosphate SAMs with Na+ and Ca<sup>2+</sup> as a counterion behave qualitatively the same for ice coverage/film thickness ranging from 0.2 to 11 monolayers (0-40 Å). Ultrathin  $D_2O$  ice overlayers ( $\sim 0.2$  monolayer) are strongly bound to these surfaces, possibly at several energetically narrow sites, and the energy of desorption is generally higher

than previously published values obtained for analogous highenergy SAMs. Ultrathin D<sub>2</sub>O ice deposited onto serine and serine phosphate with H<sup>+</sup> as a counterion is less strongly bound to the SAM surface. These differences in interaction energy also influence the lateral mobility of D<sub>2</sub>O molecules, a behavior that is clearly visible in dynamic infrared studies of the phase properties of thick  $D_2O$  ice overlayers ( $\sim 2.5$  monolayers). For example, the D<sub>2</sub>O ice overlayer on the serine surface undergoes a gradual change from amorphous to crystalline-like ice over a fairly broad range of temperatures, whereas the same transition occurs abruptly at high temperatures on the serine phosphate SAMs with Na<sup>+</sup> and Ca<sup>2+</sup> counterions. We believe that this type of fundamental studies of interfacial water and ice may contribute to an improved understanding of chemistry and growth of biomineral surfaces, e.g. hydroxyapatite, and in the long run to the design and development of novel implantable biomaterials.

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