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# Arsenic and Manganese Contamination of Drinking Water Resources in Cambodia: Coincidence of Risk Areas with Low Relief Topography

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Arsenic contamination of groundwater has been identified in Cambodia, where some 100,000 family-based wells are used for drinking water needs. We conducted a comprehensive groundwater survey in the Mekong River floodplain, comprising an area of 3700 km<sup>2</sup> (131 samples, 30 parameters). Seasonal fluctuations were also studied. Arsenic ranged from 1 to 1340 µg L<sup>-1</sup> (average 163 µg L<sup>-1</sup>), with 48% exceeding 10 µg L<sup>-1</sup>. Elevated manganese levels (57% > 0.4 mg L<sup>-1</sup>) are posing an additional health threat to the 1.2 million people living in this area. With 350 people km<sup>-2</sup> potentially exposed to chronic arsenic poisoning, the magnitude is similar to that of Bangladesh (200 km<sup>-2</sup>). Elevated arsenic levels are sharply restricted to the Bassac and Mekong River banks and the alluvium braided by these rivers (Kandal Province). Arsenic in this province averaged 233 µg L<sup>-1</sup> (median 100 µg L<sup>-1</sup>), while concentrations to the west and east of the rivers were < 10 µg L<sup>-1</sup>. Arsenic release from Holocene sediments between the rivers is most likely caused by reductive dissolution of metal oxides. Regions exhibiting low and elevated arsenic levels are co-incident with the present low relief topography featuring gently increasing elevation to the west and east of a shallow valley—understood as a relict of pre-Holocene topography. The full georeferenced database of groundwater analysis is provided as Supporting Information.

## Introduction

Arsenic is a persistent contaminant in groundwater and drinking water in countries such as Bangladesh, Vietnam, Argentina, and the United States (1–3). The Bengal Delta is the most prominent region because some 43 million people have been drinking arsenic-rich water for 20–30 years (1). Health problems are reported to occur after 10–15 years of chronic exposure to elevated As levels around 50 µg L<sup>-1</sup>. The World Health Organization (WHO) recommends 10 µg L<sup>-1</sup> As as drinking water guideline.

Drinking water supplies in Cambodia are dependent on groundwater resources. Although surface water is still used as drinking water in some areas, family-based groundwater tube-wells were becoming popular during the last 10 years. In Kandal Province, some 1 million people have stopped using surface water due to bacterial diseases that are at least partly responsible for a high infant mortality (71 deaths/1000 live births) (4).

Arsenic concentrations above 100 µg L<sup>-1</sup> were for the first time identified in Cambodia in 2000 through a small-scale drinking water quality screening in hand-pumped tube-wells (5). Four publications dealing with arsenic in the Mekong Delta (including Vietnamese part) are to date available in the international literature (6–9). However, they provide a very coarse picture of the situation, and published risk maps are misleading in certain areas. For example, the Takeo and Prey Vêng Provinces were assigned a high risk of arsenic contamination (6, 7), whereas we show that these provinces are low arsenic regions. Moreover, our study demonstrates that groundwater arsenic contamination in the Mekong Delta is a serious and lasting problem, but by no means short-lived (ephemeral) as was stated by Stanger et al. (7).

The Mekong Delta is located in southern Vietnam and neighboring Cambodia between 8°30' to 11°30' N and 104°40' to 106°50' E and is confined by the South China Sea in the southeast, the Gulf of Thailand in the west, the Vamcodong River in the northeast, and a well-defined late Pleistocene terrace to the north. The Mekong River is 4300 km long and has a catchment area of 520,000 km<sup>2</sup>. It originates in the Tibetan Plateau, and flows through China, Myanmar, Laos, Thailand, Cambodia, and Vietnam. Close to Phnom Penh, the Mekong divides into two branches: the Mekong to the southeast and the Bassac River to the south. To the north, the Tonle Sap River is connecting the Tonle Sap Lake with the Mekong River. The depositional environment in Phnom Penh is largely limited to a linear-trending valley that is fault controlled along the Bassac and limited by Pleistocene uplands adjacent to the Mekong. The Mekong River in Cambodia is a broad river that becomes tidal upstream to the northeast of Phnom Penh, near Kampong Cham (6). The delta plain has an area of about 62,000 km<sup>2</sup>, with 10,000 km<sup>2</sup> belonging to Cambodia and the rest located in southern Vietnam.

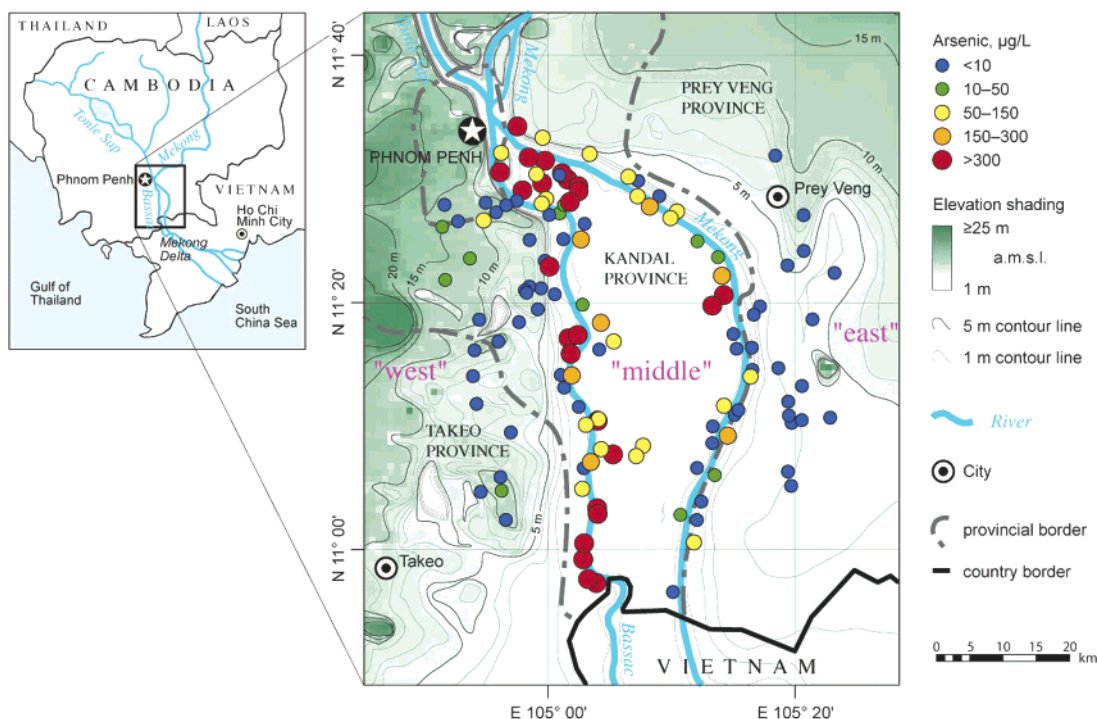
The Mekong River transports great volumes of sediments (160 million t yr<sup>-1</sup>) to the South China Sea and contributes up to 30% of the world's ionic input to the ocean (10, 11). The delta consists almost entirely of young alluvial soils (12). Sediment composition and groundwater characteristics vary in a complicated way and are known only in limited areas of the Vietnamese part of the Mekong Delta (13). The climate is humid and tropical, with average temperatures of 27–30 °C. The rainy season lasts from April to November (14).

For the study presented here, we collected 131 groundwater samples from private tube-wells in the Cambodian Mekong River floodplain (Kandal, Takeo, and Prey Vêng Provinces) for the determination of 28 hydrogeochemical parameters. To our knowledge, this is the first comprehensive published database of regional as well as temporal groundwater composition. The goals were (i) to study the magnitude and distribution of arsenic and manganese contamination, (ii) to identify hydrogeological features of the risk areas, (iii) to elucidate mechanisms of arsenic release, and (iv) to study the temporal trends of groundwater contamination. The main triggers leading to arsenic as well as manganese release are discussed with respect to risk assessment.

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**FIGURE 1.** Arsenic concentrations measured in Cambodian groundwater collected in November/December 2004 and March 2006. The low relief topography is visualized by green shading as well as 1 and 5 m contour lines. Note 1: The highest elevation on this map is only 25 m asl and the terrain is rather flat with very gentle slopes of 0.2–2 m km<sup>-1</sup>. This very small slope is not recognized by eye. Note 2: Only a few samples could be collected in the inner floodplain between the Mekong and Bassac Rivers because this area is swampy, scarcely populated, and almost inaccessible.

## Materials and Methods

**Study Area and Temporal Fluctuations.** In November and December 2004 and March 2006, family-based tube-wells were selected over an area of 3700 km<sup>2</sup> with a density of ~1 sample per 30 km<sup>2</sup>. A total of 131 groundwater samples were collected, of which 25 originated from the west of Bassac River in Kandal and Takeo Provinces (group “west”), 90 were from between the Bassac and Mekong River in Kandal Province (group “middle”), and 16 were from east of Mekong River in Prey Veng Province (group “east”) (see Figure 1). Seven wells located west of and very close to Bassac River (north of 11°47' N) were attributed to the “middle” group due to the meandering structures observed which were otherwise only found between the rivers. Moreover, four river water samples (Tonle Sap River, Bassac, Mekong south and north of Phnom Penh) were analyzed. In order to study fluctuations in groundwater composition, another 26 wells located in a special study area of the “middle” region (Kandal Province) were sampled in April 2004, August 2004, and January 2005 (for sampling design, see Figure SI 1 in the Supporting Information).

**Sample Collection.** Groundwater was collected at the tube by a hand pump or electrical pump. Samples were taken after 10 min of pumping, i.e., after the oxygen concentration in the water reached a stable value. Redox potential against SHE, pH, temperature, oxygen, and conductivity were recorded on-site by a portable system YSI 556 and a WTW Multi 340i (John Morris Scientific Pty Ltd). The samples were filled in polypropylene bottles (rinsed with 1% HNO<sub>3</sub> and 3 times with distilled water before shipping and 3 times with well water before taking the sample). An aliquot (60 mL) for the analysis of metals, ammonium, and phosphate was 0.45 µm filtered (cellulose nitrate filter, Schleicher & Schuell, Germany) and acidified with approximately 1 mL of concentrated nitric acid (65%, Fluka, Switzerland) to reach a pH

<2. Anions, alkalinity, and DOC were determined in non-acidified and nonfiltered water (120 mL). The samples were shipped to Switzerland by express mail and stored at 4 °C in the dark until analysis. Control samples transported to Cambodia and back did not show any impact of transport on the analytical results.

**Chemical Analysis.** The chemical constituents in the groundwater samples were quantified in triplicates. The full georeferenced database of measured concentrations is available as Supporting Information (Table SI 1). Arsenic concentrations were measured in parallel by atomic fluorescence spectroscopy (AFS, Millennium Excalibur, PS Analytical, UK) and inductively coupled-plasma mass spectrometry (ICP-MS, Element 2, Thermo Electron, Bremen, Germany). Cross-evaluation of these methods agreed within 5% (Figure SI 2). Fe, Mn, Na, K, Ca, Mg, and Ba concentrations were measured by inductively coupled-plasma optical emission spectroscopy (ICP-OES, Spectro Ciros CCD, Kleve, Germany); Co, Ni, Cu, Zn, Pb, Cr, Cd, and Ba were measured by ICP-MS; ammonium and phosphate were analyzed by photometry; nitrate, sulfate, and chloride were determined by ion chromatography (Dionex, Switzerland); alkalinity was measured by titration; and dissolved organic carbon (DOC) was measured with a TOC 5000 A analyzer (Shimadzu, Switzerland). Details on quality assurance are given in the SI.

**Statistical Analysis.** For the three regions “west”, “middle”, and “east”, statistical analysis of variance (ANOVA) was conducted using Systat 11 (Table SI 2). Log-transformations were performed for all parameters that did not have a normal distribution of residuals (residual analysis using the Tukey–Anscombe plot and the Q.–Q. plot). Parameters including zero-values were log-transformed after performing a correction after Stahel (15). Pairwise comparison probabilities (*p*) were calculated for each pair of regions using the Bonferroni adjustment.

**TABLE 1. Average Concentrations (Arithmetic Mean), Medians, and Ranges of Groundwater Parameters Analyzed in the “West”, the “Middle”, and the “East” Regions (for Region Definition, See Section Study Area and Seasons)**

parameter	unit	West region			Middle region			East region		
		average	median	range	average	median	range	average	median	range
As <sub>total</sub> <sup>a</sup>	μg L <sup>-1</sup>	9	3	<1–100	233	100	1–1340	3	3	1–5
Fe	mg L <sup>-1</sup>	1.0	<0.05	<0.05–10	2.8	1.6	<0.05–16	0.1	<0.05	<0.05–1.3
Mn	mg L <sup>-1</sup>	1.1	1.0	<0.1–2.6	0.6	0.4	<0.1–3.1	0.3	0.2	0.1–0.9
Na	mg L <sup>-1</sup>	130	93	11–700	79	44	9–560	40	36	6–92
K	mg L <sup>-1</sup>	2.8	2.2	0.7–7.9	2.9	2.2	0.4–24	2.5	2.4	0.8–4.9
Ca	mg L <sup>-1</sup>	52	35	1–210	44	40	7–220	18	12	2–77
Mg	mg L <sup>-1</sup>	36	22	0.6–150	21	18	5.2–130	14	6.8	1.8–82
Ba	μg L <sup>-1</sup>	160	100	20–690	414	260	20–4200	103	77	12–350
PO <sub>4</sub> <sup>3-</sup> –P	mg L <sup>-1</sup>	0.2	<0.2	<0.2–1.4	0.66	0.40	<0.2–3.14	<0.2	<0.2	<0.2–0.44
H <sub>4</sub> SiO <sub>4</sub> –Si	mg L <sup>-1</sup>	24.1	24.7	15.9–31.2	17.2	17.0	5.0–35.3	28.6	30.6	17.3–35.5
HCO <sub>3</sub> <sup>-</sup>	mg L <sup>-1</sup>	351	343	47–838	364	343	121–652	166	124	34–483
Cl <sup>-</sup>	mg L <sup>-1</sup>	152	39.3	3.5–1180	63.4	13.5	1.9–709	25.7	15.9	0.6–110
SO <sub>4</sub> <sup>2-</sup>	mg L <sup>-1</sup>	82	26	<5–590	21	<5	<5–1000	24	5	<5–310
NO <sub>3</sub> <sup>-</sup> –N	mg L <sup>-1</sup>	1.5	<0.25	<0.25–22.0	0.1	<0.25	<0.25–1.9	<0.25	<0.25	<0.25–<0.25
NH <sub>4</sub> <sup>+</sup> –N	mg L <sup>-1</sup>	0.8	0.2	<0.01–5.4	6.8	2.7	<0.01–52	0.2	0.2	<0.01–0.5
DOC	mg L <sup>-1</sup>	1.9	2.5	<1.5–8.1	3.4	2.7	<1.5–15	1.7	2.3	<1.5–3.1
total hardness	mmol L <sup>-1</sup>	3.1	2.0	0.3–12	2.1	2.0	0.4–11	1.1	0.6	0.2–5.3
pH (field)		6.83	6.77	5.74–7.62	7.03	7.02	6.55–7.65	6.41	6.57	5.42–6.85
E <sub>c</sub> (field)	μS cm <sup>-1</sup>	1300	744	92–6150	870	741	243–3597	420	300	78–1220
E <sub>h</sub> (field)	mV	0	0	–128–96	–78	–80	–408–45	14	20	–62–72
T (field)	°C	30.0	30.1	29.1–30.8	29.4	29.4	28.2–30.2	29.9	29.9	29.3–30.6
well depth	m	26	25	9–40	40	40	17–65	33	33	22–48
Co	μg L <sup>-1</sup>	1.4	0.9	0.1–5.7	0.3	0.0	0.1–2.6	2.3	0.6	0.1–17
Ni	μg L <sup>-1</sup>	3.3	2.3	1.0–23	2.5	1.9	0.4–11	4.8	3.6	1.1–23
Cu	μg L <sup>-1</sup>	7.8	6.1	2.6–31	6.2	6.2	0.4–21	8.5	7.8	4.8–18
Cr	μg L <sup>-1</sup>	2.1	0.4	<0.1–14	0.4	0.3	<0.1–2	0.3	0.3	<0.2–0.6
U	μg L <sup>-1</sup>	5.8	3.1	<0.1–32	1.1	0.1	<0.1–9.3	1.1	0.1	<0.1–8.1
Cd	μg L <sup>-1</sup>	0.2	0.1	0.1–1.0	0.1	0.1	0.1–2.3	0.2	0.1	0.1–0.6

<sup>a</sup> Determined by AFS.

The fluctuations over time (three times: April 2004, August 2004, and January 2005) were also variance analyzed in a mixed model. Here, two factors were used: time (fixed factor) and well-number (random factor). The interaction term was omitted using the General Linear Model (GLM). Probabilities are given in Table SI 3 and average concentrations with standard errors for selected parameters are given in Figure SI 3a and b.

To identify parameter associations for the three regions studied, principal component analysis (PCA) was performed. A detailed description and the results are given in the SI (Figure SI 4).

## Results and Discussion

**Arsenic Contamination and Redox Parameters.** Figure 1 shows the arsenic distribution in the studied area. Very high arsenic concentrations were observed along the rivers and in the floodplain between Mekong and Bassac River (maximum 1340 μg L<sup>-1</sup>, average 233 μg L<sup>-1</sup>, median 100 μg L<sup>-1</sup>). The classification of three groups of well samples (“west”, “middle”, and “east”, see section Study Area and Temporal Fluctuations) and comparison of their groundwater characteristics (Table SI 1) provides insight into the conditions promoting As release in Cambodia (see below). The concentration ranges of all species measured in this study are summarized in Table 1 for the three regions. Out of the 131 wells sampled, 63 (48%) had concentrations above the WHO guideline value of 10 μg L<sup>-1</sup> with 60 wells located between the two rivers. Moreover, 48 wells (36%) exceeded the 50 μg L<sup>-1</sup> level which is the drinking water limit in many developing countries.

The average arsenic concentration in the “middle” region is significantly higher than those in the “west” and “east” regions (see Table SI 2 for statistical verification). Furthermore, groundwater in the “middle” region has significantly lower average E<sub>h</sub>, and higher DOC, NH<sub>4</sub><sup>+</sup>, HCO<sub>3</sub><sup>-</sup>, PO<sub>4</sub><sup>3-</sup>, and

Fe(II) concentrations (Figure 2a, Table SI 2). Thus, this groundwater is reducing in nature, supported by the fact that almost all inorganic N is present in the reduced NH<sub>4</sub><sup>+</sup> species (only 7 of 90 samples had traces of NO<sub>3</sub><sup>-</sup> ranging from 0.2 to 1.9 mg L<sup>-1</sup>). A few wells in the “middle” region have low As concentrations despite E<sub>h</sub> values that are below –150 mV. This might be due to As<sub>2</sub>S<sub>3</sub> precipitation under sulfate reducing conditions (16).

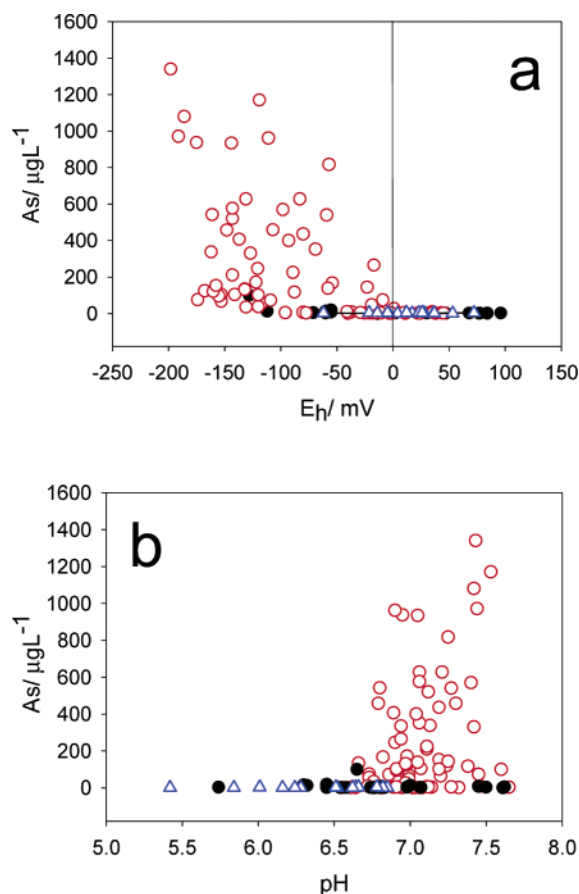
Concentrations of NH<sub>4</sub><sup>+</sup> and DOC show a weak, but positive correlation with As ( $r^2 = 0.46$  and  $0.44$ , respectively, Figure SI 5a and b). Considering these facts, As release seems to be triggered by microbially induced reductive dissolution of As-coated iron oxides and MnO<sub>2</sub> surfaces. Principal Component Analysis supports this hypothesis (see Statistical Analysis in Materials and Methods Section).

As can be seen in Figure 2b, very high As concentrations (>1000 μg L<sup>-1</sup>) were accompanied by pH values >7.3. Neutral to high pH conditions favor As release by promoting desorption processes compared to the predominantly acidic conditions found in the “west” and “east” regions (1, 3, 17). High arsenic concentrations in the studied area of Cambodia are clearly triggered by reducing conditions, and pH ≥ 7 might possibly enhance the mobilization of As at some locations.

Sediment coring and analysis was beyond the scope of this study and omitted because the composition of sediments is not necessarily indicative of the aqueous phase composition (18, 19). Or in the words by Meharg et al.: “...the solid-to-solution mass transfers of As required to produce contaminated groundwater are small in relation to the total amount of As in the sediments. This makes it difficult to use mineral mass balances to deduce the mobilization and precipitation mechanisms responsible for the high-As groundwaters” (20).

**Manganese Contamination.** The “west” region exhibits suboxic conditions (E<sub>h</sub> ≈ 0) with low As concentrations (9 μg L<sup>-1</sup>), but it contains significantly higher average Mn levels (1.1 mg L<sup>-1</sup>) compared to the “middle” (0.6 mg L<sup>-1</sup>) and the

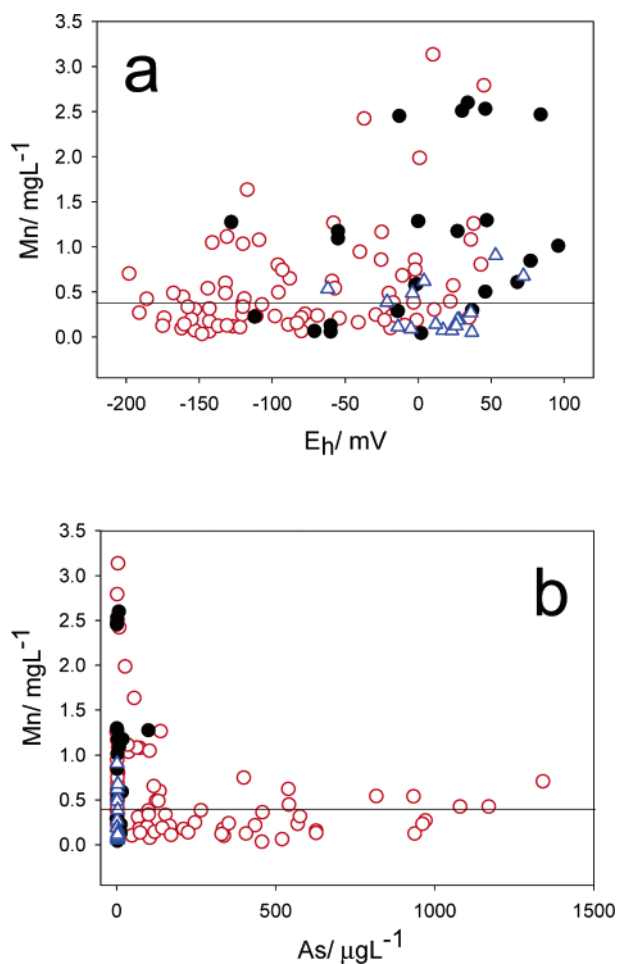




**FIGURE 2.** a, b. Arsenic concentration as a function of (a) redox potential,  $E_h$ , and (b) pH, in the “west” (●), “middle” (○), and “east” (△) regions.

“east” ( $0.3 \text{ mg L}^{-1}$ ) regions (Table SI 2). This finding indicates that  $\text{MnO}_2$  is reduced and  $\text{Mn}^{2+}$  is released to the aqueous phase. The  $E_h$  values seem to be generally not low enough to entirely reduce sediment-bound iron(hydr)oxides (low dissolved iron levels) which therefore still offer strong binding sites for As and phosphate. The absence of nitrate in 20 of 24 wells supports the hypothesis of  $\text{MnO}_2$  reducing conditions. On the other hand, a weak but positive correlation is found for Mn concentrations with  $E_h$  (Figure 3a). This might be due to  $\text{MnCO}_3$  precipitation (the “west” region is slightly over-saturated with respect to  $\text{MnCO}_3$ ;  $Q_{sp} = 3.2 \cdot 10^{-11} > K_{sp} = 1.8 \cdot 10^{-11}$ ) and/or, to a lesser extent,  $\text{MnS}$  precipitation at lower  $E_h$  values (Table SI 4). As a result of such redox conditions, elevated Mn levels are strongly anti-correlated with As (the higher the As level, the lower the Mn level and vice versa) (Figure 3b).

Elevated Mn concentrations have an impact on plants and human beings. Manganese is one of the major factors limiting crop production on acidic soil (about 30% of the world’s total land area) (22). In plants, symptoms such as chlorosis (photooxidation of chlorophyll) and necrosis due to accumulation of phenolic compounds are observed. It has also been shown that maternal environmental exposure to Mn is associated with a reduced activity of the newborn’s erythrocyte Ca-pump (23). Furthermore, exposure to Mn in drinking water is associated with neurotoxic effects in children, for example with a diminished intellectual function (24). Consequently, high Mn concentrations in irrigation and drinking water are hazardous for plants and human beings. It is worth noting that out of 131 wells sampled, 75 (57%) exhibit Mn concentrations  $> 0.4 \text{ mg L}^{-1}$  (Figure 3a and b). Upon oxidation, Mn leads to clouding of drinking water by



**FIGURE 3.** a, b. Manganese concentration as a function of (a)  $E_h$  and (b) As concentration, in the “west” (●), “middle” (○) and “east” (△) regions. The line at  $0.4 \text{ mg L}^{-1}$  represents the WHO guideline (27).

precipitation of  $\text{MnO}_2$  which is, however, a rather slow process. Although in the west of Bassac River groundwater use for drinking water purposes might be regarded as chemically safe with respect to arsenic, average manganese concentrations are notably above the WHO guideline value.

**Coincidence of Low-Relief Topography with Areas of Elevated Arsenic Levels.** The low-relief topography depicted in Figure 1 features a flat area confined between the rivers (max. 1 m altitude variation) that is accompanied by tiny gradual slopes of only  $0.2\text{--}2 \text{ m km}^{-1}$  to the west and to the east. These slopes cannot be recognized by eye and thus seem totally flat to the observer, however, they become apparent in the shaded contour plot that we derived from the USGS digital elevation model GTOPO30 by interpolation with ArcGIS (Figure 1). Obviously, elevated groundwater As levels are exclusively present in the flat land embraced by Mekong and Bassac as well as on the adjacent river banks. To the best of our knowledge, such sharply confined risk areas have not been reported elsewhere.

We suggest that the present low relief of the studied area is indicative for the distribution of arsenic in groundwater in the studied area. Although our present knowledge is too limited to provide clear evidence for a causal link between topography and arsenic risk areas, there are a number of basic indications. The shallow valley was deeper after the Pleistocene period and was then filled with Holocene alluvial deposits. Satellite pictures reveal that former river channels (seen as oxbows and old meanders) are almost exclusively situated between the two rivers, but not to the east and west

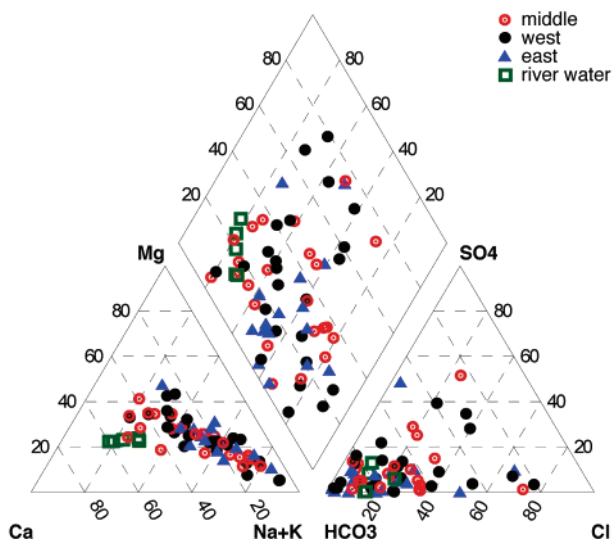


FIGURE 4. Piper diagram depicting the main hydrochemical conditions of the three groups of groundwater in this study, river water (Bassac, Tonle Sap River, Mekong north of and south of Phnom Penh (Table SI 1)) and of Mekong River water (11).

(Figure SI 6). It can further be noted that during the dry season only the “middle” region retains enough water to sustain vegetation and wetlands while the “slopes” in the east and the west lay barren (Figure SI 7). Wetlands such as the ones formed between the rivers are known to accumulate natural organic matter in the buried sediments (25).

Consequently, the present-day topography appears to reflect the boundary of organic-rich Holocene sediments deposited between the rivers. This is supported by the findings of Polizzotto et al. (personal communication) stating “The Holocene stratigraphy between the rivers is featuring 10–20 m of clay overlaying fine gray sand which persists down to ~45–60 m. The clay layer becomes anoxic below the groundwater table, with abundant organic matter, often in the form of large pieces of wood. Many variations from this general characterization are present”.

**Major Cations and Anions.** Total water hardness is 2–3 times higher in the “west” region compared to the “middle” and “east” regions (Table 1). The Na/Cl ratios being  $>1$  in most of these samples suggests ion exchange by Ca on clays (Figure SI 8a and b). The stoichiometric ratios of  $\text{Ca}/\text{Mg} = 1:1.1$  (ideally 1:1 for  $\text{CaMg}(\text{CO}_3)_2$ ) and  $\text{Ca}/\text{HCO}_3 = 0.3:1$  (ideally 0.25:1) found in the water samples imply dissolution of  $\text{CaMg}(\text{CO}_3)_2$  (Figure SI 8c and d). Thus, the origin of the alkalinity is predominantly based on the dissolution of dolomite.

The pronounced gradient of total hardness from the “west” to the “middle” and “east” indicates that groundwater in the “west” and the “east” is influenced by sources other than river water infiltration. Obviously, the sharply confined arsenic risk areas exhibit different total hardness. Again, the low relief topography coincides with the groundwater composition as discussed above.

According to the relative molar portion of the dissolved ionic species, groundwater in the “middle” region is of Ca–Mg–(Na)– $\text{HCO}_3$  type (40%), with the rest dominated by Mg–Ca–Na– $\text{HCO}_3$  and Na–Ca–Mg– $\text{HCO}_3$  (Figure 4). The “west” and “east” regions are mainly of the Na–(Mg)– $\text{HCO}_3$ –(Cl) type. There is a mixing trend toward more sodium rich water with distance from the rivers Mekong and Bassac. The few “middle” region samples that show a clear Ca–Mg prevalence are situated very close to the rivers, and might therefore be predominantly infiltrated river water. This is supported by the fact that Mekong River water composition shows also a clear Ca–Mg prevalence (Figure 4) (11).

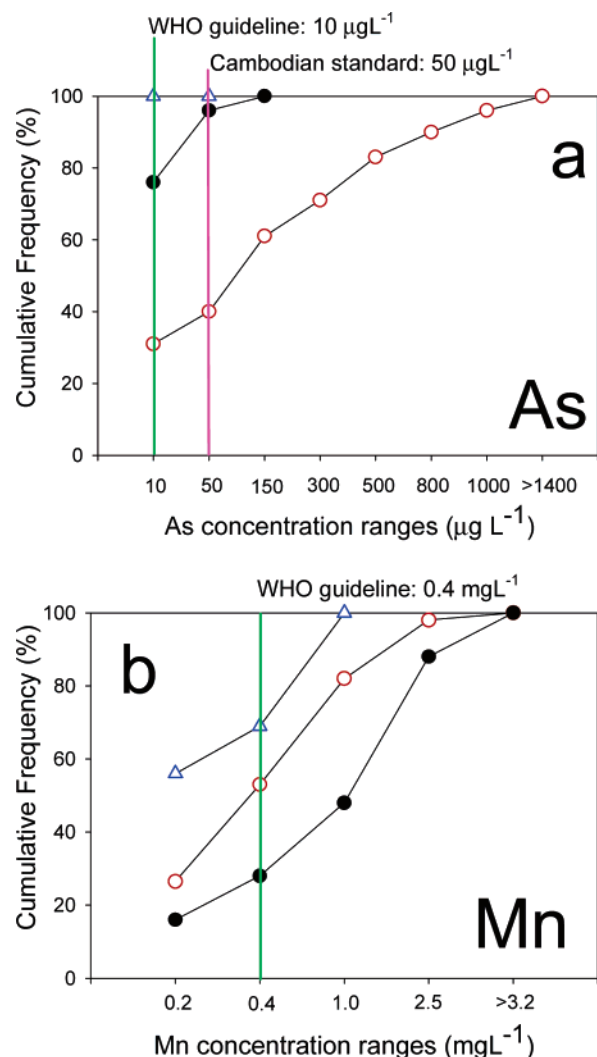
**Well Depths.** Although our study revealed As levels between 1 and  $460 \mu\text{g L}^{-1}$  in shallow wells (15–20 m), concentrations were generally higher (up to  $>1000 \mu\text{g L}^{-1}$ )

between 25 and 45 m depth (Figure SI 9). The deepest well of our study was 65 m. The wells in the “west” and “east” regions are, on average, not as deep as the ones in the “middle” region. There is, however, no correlation between region, well depth, and As concentration. In other words, although the “east” and “west” regions have numerous wells with depths of 25–40 m, they all have very low As levels.

**Fluctuations in Groundwater Composition.** To elucidate temporal fluctuations of groundwater constituents, 26 tube-wells were sampled in a special study area of Kandal Province 3 times over a time span of 9 months, i.e., in April 2004, August 2004, and January 2005 (for locations, see Figure SI 1). Average As, Mn, Fe,  $\text{PO}_4^{3-}$ ,  $\text{SO}_4^{2-}$ ,  $\text{NH}_4^+$ , and DOC concentrations for each season were not significantly different from each other taking only time as a factor in the statistical analysis of variance. However, ANOVAs considering each individual well as a second (random) factor showed that concentrations differed significantly with time (Table SI 3). Yet, compared to the differences found between the three regions, these fluctuations were small. The deviations from average ( $n = 3$ , maximum 25%) are shown for As in Figure SI 10. Although DOC,  $\text{NH}_4^+$ ,  $\text{PO}_4^{3-}$ , and Fe had higher average values in April compared to August or January (Figure SI 3a and b), these parameters were not accompanied by a lower  $E_h$  value. Due to altered directions of the groundwater flow caused by changing groundwater heads, the tube-wells possibly received water from shifted flow paths over the seasons. Nevertheless, it can be concluded that fluctuations of groundwater composition over time are small compared to regional differences. In contrast to the assertion by Stanger et al., the arsenic levels in groundwater are by no means ephemeral (7).

Concentrations of major cations and anions remained constant over the 9-month study period and so did concentrations of heavy metals, leading to a constant total hardness of these groundwater samples. No significant dilution was observed during the inundation period.

**Health Considerations.** Figure 5 shows the cumulative frequencies for As and Mn concentration ranges in the three regions studied. The high arsenic concentrations in the Kandal Province are posing an alarming health threat to the people who consume this water daily. In addition, the elevated Mn levels found in Prey Vêng Province must be considered harmful for infant children. Out of 131 wells, 48% had As concentrations exceeding the WHO guideline value ( $10 \mu\text{g L}^{-1}$ ) and 57% had Mn concentrations higher than  $0.4 \text{ mg L}^{-1}$  (WHO). While high Mn concentrations were often



**FIGURE 5.** Cumulative frequency distributions of (a) arsenic and (b) manganese concentration ranges of groundwaters in the "west" (●), "middle" (○) and "east" (△) regions, where 0.2 mg L<sup>-1</sup> Mn means the range 0–0.2 mg L<sup>-1</sup>.

coupled with low As concentrations and vice versa (Figure 3b), many wells exhibited an As problem, others had a Mn problem, and only 18% provided chemically safe drinking water. The long-term health impact for the 1.2 million people drinking this groundwater without treatment is profound if remedial actions are not taken. Simple and low-cost mitigation measures for the household level include sand filters for arsenic removal, ion exchange, activated alumina, modified coagulation/filtration, and lime softening. Sand filters were demonstrated to operate effectively if groundwater contains >1 mg L<sup>-1</sup> iron and <2 mg L<sup>-1</sup> phosphate (26). In the highly affected "middle" region, 52% of the wells had Fe >1 mg L<sup>-1</sup> and 26% had Fe >5 mg L<sup>-1</sup> (with As >10 μg L<sup>-1</sup> and phosphate <2 mg L<sup>-1</sup>). Compared to Vietnam (86% with Fe >1 mg L<sup>-1</sup>, *n* = 134) and Bangladesh (69% with Fe >1 mg L<sup>-1</sup>, *n* = 3284), on average, less dissolved iron is present in Cambodia restricting the potential use of arsenic removal by sand filters to a lower percentage (26).

Most of the tube-wells were built during the last 10 years. Incidents of chronic As related diseases such as arsenicosis were already reported (M. Sampson, personal communication). Because arsenicosis is known to occur after 10–15 years of chronic As ingestion (>50 μg L<sup>-1</sup>), people drinking contaminated water are exposed to a serious health risk. Skin lesions in children and significantly higher As levels in

human hair samples of people living in Kandal Province are proof that the problem must be taken seriously (27). In addition, the elevated Mn concentrations should be considered. Infant mortality in Cambodia is 71/1000 which is 18 times higher than that in central Europe (4). The health impact of Mn is known to be harmful for newborns and children (23, 24).

With 350 people km<sup>-2</sup> potentially exposed to chronic arsenic poisoning, the magnitude is similar to that in Bangladesh (200 km<sup>-2</sup> (1)), West Bengal (250 km<sup>-2</sup> (1)), or Vietnam (270 km<sup>-2</sup> (2)). This health threat justifies international recognition. Consequently, we urgently propose early mitigation actions in order to reduce the risk of chronic arsenic poisoning of a million people living in Kandal Province.

Besides a scientifically sound distinction of relevant geochemical factors, we have shown that the present topography (low relief) coincides with sharply confined high and low arsenic areas, and that it reflects the boundary of recent Holocene sediment deposits. We believe such topography is a worthy indicator to be considered in As groundwater studies.

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## Supporting Information Available

Full geo-referenced database of 30 parameters measured in the groundwater samples (Excell file, Table SI 1), details and results of quality assurance, statistical analyses of variance (ANOVA), principal component analysis (PCA), several plots of cross-correlation, temporal variations of arsenic, and satellite pictures of the study area. This material is available free of charge via the Internet at <http://pubs.acs.org>.

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