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Dissolved Organic Matter Quality in a Shallow Aguifer of Bangladesh: ₂ Implications for Arsenic Mobility

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ABSTRACT: In some high arsenic (As) groundwater systems, correlations are observed between dissolved organic matter (DOM) and As concentrations, but in other systems, such relationships are absent. The role of labile DOM as the main driver of microbial reductive dissolution is not sufficient to explain the variation in DOM-As relationships. Other processes that may also influence As mobility include complexation of As by dissolved humic substances, and competitive sorption and electron shuttling reactions mediated



by humics. To evaluate such humic DOM influences, we characterized the optical properties of filtered surface water (n = 10)and groundwater (n = 24) samples spanning an age gradient in Araihazar, Bangladesh. Further, we analyzed fulvic acids (FA) isolated from large volume isolates (n = 6) for optical properties, C and N content, and 13 C NMR spectroscopic distribution. Old groundwater (>30 years old) contained primarily sediment-derived DOM and had significantly higher (p < 0.001) dissolved As concentration than groundwater that was younger than 5 years old. Younger groundwater had DOM spectroscopic signatures similar to surface water DOM and characteristic of a sewage pollution influence. Associations between dissolved As, iron (Fe), and FA concentration and fluorescence properties of isolated FA suggest that aromatic, terrestrially derived FAs promote As-Fe-FA complexation reactions that may enhance As mobility.

1. INTRODUCTION

30 In Bangladesh, it is estimated that ~45 million people are 31 exposed to groundwater arsenic (As) concentrations that are 32 above the World Health Organization guideline value of 10 μ g 33 L⁻¹. Although it is generally agreed that microbially mediated 34 reductive dissolution of As-bearing iron (Fe) minerals^{2,3} is the 35 dominant mechanism for enrichment of As in reducing 36 groundwater in the Bengal Basin, the role of dissolved organic 37 matter (DOM) remains enigmatic in part due to a lack of 38 characterization of DOM composition. Compared to the well-39 recognized role of labile DOM fueling microbial reduction, 40 demonstrated in several field studies of Bengal groundwater,4 41 less is known about the role of humic DOM in As-laden 42 groundwater. Although humic DOM is considered to be 43 recalcitrant to biodegradation, it is involved in both biological 44 and chemical reactions that may influence the mobility of As. 45 For instance, microbial reduction of oxidized humic quinones 46 has been shown to be an important step in the cascade that 47 results in Fe reduction 10 and may consequently contribute to 48 As mobilization. Quinone moieties in dissolved and solid-phase

humic substances are able to shuttle electrons and accelerate $_{\rm 49}$ microbial reduction of $Fe^{\rm 10-13}$ and other terminal electron $_{\rm 50}$ acceptors and are able to regenerate and continue to serve as 51 electron shuttles.1

In addition to the biological role for humic DOM, evidence 53 for humic DOM influencing As mobility through abiotic 54 mechanisms such as competitive sorption and complexation has 55 emerged in laboratory studies. It was demonstrated that 56 competition with As for sorption sites on hematite 15 and 57 goethite 16 minerals resulted in As desorption. Humic DOM has 58 also been shown to abiotically oxidize and reduce As from soils 59 and sediments. 16,17 Of particular interest is the formation of 60 complexes between humic substances, Fe, and As, which act to 61 keep As in solution under reducing conditions. Direct 62 complexation was initially suggested between dissolved humics 63

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64 and arsenate^{18,19} and arsenite.^{20,21} More recently, Fe-bridging 65 for ternary complex formation involving humic substances, Fe, 66 either arsenate^{22,23} or arsenite,²⁴ has been invoked as the 67 mechanism responsible. Indeed, X-ray absorption spectroscopy 68 results demonstrate that ternary As—Fe—humic DOM com-69 plexes form by inner-sphere binding of As(V) to Fe(III)-humic 70 DOM complexes.²⁵

Field observations also support that the quality of DOM 72 exerts an influence on As mobility. DOM quality refers to the 73 chemical characteristics of DOM that influence its biological 74 reactivity such as how labile it is for microbial degradation and 75 its chemical reactivity, which may be enhanced by reactive 76 humic moieties and N- and S-containing functional groups. 77 First, field observations of significant positive correlations, 78 significant negative correlations, 27 or no correlations at all 28 79 suggest that the role of DOM is more complex than previously 80 thought due to the heterogeneous nature of DOM. Humic 81 DOM, including dissolved humic and fulvic acids, participate in 82 a potential suite of reactions, such as complexation, 83 competition, and electron shuttling that likely also influence 84 As mobility. Second, studies in the Bengal Basin have 85 implicated the sediments including dispersed organic matter 86 or peaty strata as a likely source of DOM in reducing 87 groundwater. ^{29–32} Further, optical spectroscopic analyses of 88 Bangladesh groundwater demonstrated that the DOM in 89 groundwater with elevated As concentrations was largely 90 derived from aromatic and lignaceous organic compounds 91 already in the sediments.³³ Taken together, the role of the 92 humic DOM pool in reducing aquifers of the Bengal Basin 93 merits further exploration.

This study first elucidates the potential role that dissolved 95 humic substances, namely, fulvic acids (FAs), the predominant 96 humic fraction in the DOM pool of natural waters, play in As 97 mobility through a characterization of DOM chemistry in 98 groundwater of varying ages and As concentrations. Ground-99 water-dissolved fulvic acids have not previously been studied in 100 As-rich reducing groundwater environments, although extracts 101 of sediment organic compounds have recently been analyzed 102 and provide some insight into sedimentary DOM character. In 103 several regions with elevated groundwater As concentration, 104 Reza et al. 34,35 performed absorbance and fluorescence analyses 105 and ¹H- NMR, ¹³C NMR, and Fourier transformed infrared 106 spectral analyses of organic compounds extracted from 107 sediments using an acid-alkanine method. These studies 108 demonstrated that the extracted sedimentary organic com-109 pounds from the Meghna and Brahmaputra floodplains had 110 higher aromaticity and fluorescence intensity than those from 111 the Ganga floodplain and could be involved in complexation 112 reactions with As. These studies of sedimentary organic 113 compounds provide additional motivation for investigating 114 the role of humic DOM in groundwater. Noting, however, that 115 groundwater of the Bengal Basin likely has diverse DOM 116 sources in addition to sedimentary organic matter, there is a 117 need to evaluate the influence that both humic DOM, isolated directly from groundwater, as well as nonhumic DOM may 119 have on As mobility.

Therefore, the goals of this study are to better characterize the DOM fractions and investigate the interactions between DOM fractions, As, and Fe in groundwater. Here, we combine FA isolation, elemental analysis, ¹³C NMR spectroscopy, UV—124 vis absorbance spectroscopy, and fluorescence spectroscopy to characterize dissolved humic substances in groundwater of the varying ages and varying dissolved As concentrations that

represent the range of hydrogeological and biogeochemical 127 conditions typical of the Holocene shallow aquifer of 128 Bangladesh. The association between fulvic acid content, Fe, 129 and As was also investigated to evaluate if the presence of 130 humic DOM could help explain variability in DOC and As 131 concentration correlations noted in previous field studies.

2. METHODS

Groundwater and surface water samples (Tables S1 and S2) 133 were from Araihazar upazila, Bangladesh, located approximately 134 30 km northeast of the capital city Dhaka. It lies within the 135 floodplain of the Old Brahmaputra River, an abandoned river 136 channel that has been filled through recent sedimentation³⁶ and 137 reduced to a small stream (Figure S1). The study area was 138 chosen because the redox condition of the shallow Holocene 139 aquifer, and in turn, groundwater As level and age, spans nearly 140 the entire range found in Bangladesh. 37-41 The groundwater 141 flow in the shallow aquifer is localized, with flow direction 142 toward local rivers or streams except during the monsoon. The 143 shallow aguifer comprises fine to medium sand sediments with 144 less than 0.15% organic carbon, although the silty surficial 145 sediment or dispersed clay lenses in aquifer displayed organic 146 carbon concentration between 0.2% to 0.6%. 42,43 Only one peat 147 layer containing 8.6% organic carbon has been identified, at 148 nearby site A,44 but there is no evidence of peat at Site K.

Groundwater samples were collected from monitoring wells $_{150}$ installed at Site K described in Radloff $_{45}$ and Site B described in $_{151}$ Zheng et al. $_{42}$ The two sites are approximately 2 km apart and $_{152}$ are in a primarily rural setting with numerous ponds and $_{153}$ irrigated rice and vegetable fields.

To isolate FA, a total of six large volume water samples 155 (~215 L) were collected in October 2009: one from a small 156 stream next to the field where monitoring wells were installed 157 (Figure S1), three from wells at Site K, and two from wells at 158 site B. The three samples from Site K represented young 159 groundwater with an ${}^{3}H/{}^{3}He$ age 45 of younger than 5 years old 160 (K12 and K10) and old groundwater with an age of greater 161 than 30 years old according to ³H/³He dating (K8). Two 162 samples were also collected from Site B (B-11 and B-14) for 163 which the groundwater was known to be about 19 years old³⁹ 164 to ensure that the age range of 5-30 years was represented. 165 The wells were pumped for 10 to 30 min, until readings of 166 temperature, conductivity, pH, 45 and Eh 45 on multiprobes in a 167 flow-through cell stabilized. Then, about 215 L of groundwater 168 was collected into cubitainers and driven to the University of 169 Dhaka, Bangladesh, where they were then acidified and filtered. 170 Large volume filtered samples were shipped to the University of 171 Colorado, where they were pumped through columns packed 172 with cleaned XAD-8 resin. 46 After rinsing loaded columns with 173 nanopure water, the FA was isolated by back-eluting with base 174 and run through a cation exchange column. These "large 175 volume FA isolates" were freeze-dried and analyzed for C and 176 N content with an elemental analyzer and C functional group 177 distribution with solid state ¹³C-nuclear magnetic resonance ¹⁷⁸ (13C NMR) spectroscopy. The surface water sample was 179 collected and processed in the same way. Prior to isolation, an 180 aliquot of the large volume isolates was reserved and analyzed 181 for total dissolved As and Fe concentrations, DOC 182 concentration, and UV-visible absorbance and fluorescence 183 spectroscopy (see Supporting Information for details).

To provide the geochemical context for the FA isolate 185 chemistry data, a total of 24 small volume groundwater samples 186 were also collected from Site K, along with a total of 10 surface 187

Table 1. Characteristics of Small Volume Samples, Including Age, Depth, Solute Chemistry and DOM Optical Spectroscopic Properties^c Collected from Araihazar, Bangladesh, in March 2008^a

	surface water $n = 10^d$	<5 year old groundwater $n = 10$	5-30 year old groundwater $n = 6$	>30 year old groundwater $n = 8$	t test results ^e
age (years)	_	1.9 ± 1.5	10.5 ± 5.0	34.7 ± 1.6	_
depth (m)	_	8.9 ± 2.2	11.1 ± 3.6	10.5 ± 3.2	G
$As_T (\mu g L^{-1})$	24 ± 14	37 ± 29	158 ± 107	281 ± 112	B-F
% As(V)	66 ± 30	16 ± 33	1 ± 1	4 ± 4	G
$Fe_T (mg L^{-1})$	0.11 ± 0.12	6.7 ± 5.3	10.3 ± 5.5	10.8 ± 2.0	E
$SO_4 (mg L^{-1})$	19 ± 15	15 ± 23	4.4 ± 6.4	0.05 ± 0.07	C, E
$Ca (mg L^{-1})$	38 ± 4.0	39 ± 25	41 ± 3.0	57 ± 17	C, E, F
Mn (μ g L ⁻¹)	149 ± 117	1885 ± 1054	2145 ± 1036	1252 ± 970	A-C
$TDN (mg L^{-1})$	1.37 ± 1.11	0.62 ± 0.60	1.20 ± 0.62	1.76 ± 0.58	A, E
$DOC (mg L^{-1})$	6.30 ± 2.98	2.78 ± 0.50	3.51 ± 0.38	3.65 ± 0.69	A-E
SUVA (L mg ⁻¹ m ⁻¹)	1.97 ± 1.22	1.18 ± 0.64	1.89 ± 0.45	2.06 ± 0.48	D, E
FI	1.69 ± 0.24	1.47 ± 0.10	1.38 ± 0.04	1.40 ± 0.03	A-E
RI	0.35 ± 0.07	0.50 ± 0.04	0.51 ± 0.02	0.48 ± 0.02	A, B, C, F
AA-like (%)	0.084 ± 0.034	0.10 ± 0.07	0.05 ± 0.01	0.04 ± 0.01	В, С, Е

"Means, standard deviations, and T-tests comparing young (< 5 yr), old (5-30 yr), and very old (> 30 yr) groundwater collected from site K are shown. b Total dissolved arsenic (As_T), % arsenic as As(V), total dissolved Fe (Fe_T), and dissolved organic carbon (DOC); n = sample size; dash = not measured or performed. FI and RI are dimensionless; AA-like = amino acid-like fluorescence. Sample size is six for As, %As(V), Fe, SO₄, Ca, and Mn values (Table S1). et test results show significant difference between (A) surface water and <5 year old groundwater, (B) surface water and 5-30 year old groundwater, (C) surface water and >30 year old groundwater, (D) <5 year old and 5-30 year old groundwater, (E) <5 year old groundwater and >30 year old groundwater, and (F) 5-30 year old and >30 year old groundwater. G represents no statistical difference in any sample sets.

188 water samples representing the Old Brahmaputra River and 189 local streams and ponds in Araihazar in March 2008 (Table 190 S1). These "small volume" whole water samples were analyzed 191 for anions and cations, dissolved As and Fe concentrations and 192 speciation, DOC concentration, and UV-visible absorbance 193 and fluorescence spectroscopy (see Supporting Information for

For both FA isolates and small volume water samples, key 195 196 metrics from UV-vis absorbance and fluorescence spectroscopy were evaluated. Specific UV absorbance (SUVA)⁴⁷ was utilized to provide information about the aromaticity of whole waters, whereas for FAs the aromatic C content was measured directly with solid state 13C NMR. Three dimensional fluorescence excitation emission matrices (EEMs) were acquired and instrument specific corrections, inner-filter correction, Raman normalization, and blank subtraction were applied. EEMs were fit to the parallel factor analysis (PARAFAC) model of Cory and McKnight, 48 and individual 206 component loadings are presented in Raman units (RU) and as 207 percentages of total fluorescence. The relative amount of amino 208 acid-like fluorescence, taken as the sum of tyrosine-like and 209 tryptophan-like PARAFAC components (C8 + C13 in Cory 210 and McKnight⁴⁸), which are described as the more biologically 211 labile components in the fluorescence literature, was tracked in samples. The fluorescence source index (FI) was also calculated⁴⁸ to provide information about the DOM sources in whole waters and FAs. The FI has provided consistent and 215 reliable information on the sources of DOM and humic 216 substances in natural waters that were consistent with 217 characterization using other techniques. For example, Mlade-218 nov et al. isolated FAs from wetland surface water and found 219 that spatial changes in the FA content, C:N ratios, C functional 220 groups, and FI reflected an influx of plant-derived organic 221 compounds during flooding. 49 The redox index (RI) has been 222 used to track the presence of reduced quinone-like compounds 223 in surface water and groundwater samples⁴⁹ and is the ratio of reduced quinone-like PARAFAC components to total quinone- 224 like PARAFAC components.⁵⁰

A multivariate analysis of variance (MANOVA; vegan 226 package, R project) was employed to test whether groundwater 227 chemistry (As, Fe, sulfate, DOC, total dissolved nitrogen 228 (TDN), SUVA, FI, and percent amino acid-like fluorescence) 229 was significantly different among surface water (n = 10) and 230 groundwater samples in the age groups of younger than 5 years 231 (n = 10), 5–30 years (n = 6), and greater than 30 years old (n = 232)8). Additionally, an analysis of variance (ANOVA) was also 233 used to determine significance in As concentration difference 234 between data sets using a two-tailed t test (R project).

3. RESULTS AND DISCUSSION

Surface water DOM and FA Isolate. Surface water 236 samples (Table S1), consisting of the old Brahmaputra River 237 (OBR) channel and local streams and ponds near Site K, 238 displayed high average DOC concentrations, greater than 239 groundwater of any age (Table 1). Because small volume water 240 tl samples were obtained in March when the groundwater tables 241 are substantially higher than the surface water tables, 41 the river 242 and stream waters are primarily groundwater fed, with the water 243 chemistry subject to "redox trapping" during groundwater 244 discharge. 51,52 "Redox trapping" means that redox-sensitive 245 elements such as As, Fe, and Mn have a tendency to be 246 removed from solution and "trapped" in sediments under 247 oxidizing conditions, and therefore, their concentrations are 248 lower than when reducing conditions prevail. The low surface 249 water concentrations of As, Fe, and Mn compared to those in 250 old groundwater and similar concentrations of Ca, which is not 251 redox sensitive, in surface water and groundwater are consistent 252 with this discharge process (Table 1). Additionally, because the 253 surface water environment is oxic, sulfate concentrations in 254 surface water are higher than in old groundwater. Similarly, 255 young groundwater, which is influenced by oxic surface water 256 inputs, is known to have substantial variability in concentrations 257 of these redox-sensitive elements (As, Fe, and Mn; Table 1).53 258

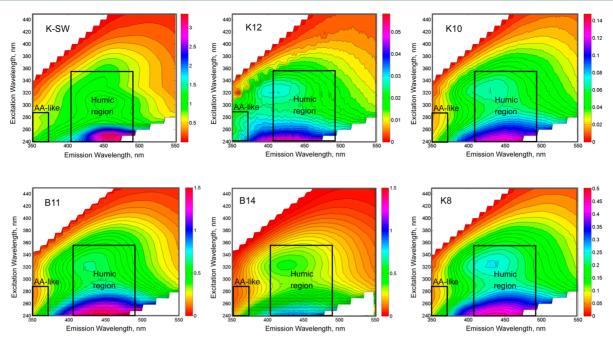


Figure 1. EEMs of surface water (K-SW) and Site K and Site B groundwater whole waters collected for fulvic acid isolation (samples described in Table 2). Fluorescence intensities (*z*-axis) are in Raman units (RU). Regions of amino acid-like fluorescence (proxy for labile DOM) and humic-like fluorescence are marked.

259 The low RI of surface water DOM (Table 1) also reflects this 260 oxic state of surface water and the presence of more oxidized quinone-like moieties⁵⁰ than in groundwater. Because surface water DOC concentrations are two to three times higher than those in old groundwater, additional allochthonous inputs from the surrounding landscape³³ and autochthonous biological inputs (e.g., refs 53 and 54) contribute to the high DOC concentrations. The FI value is highest in surface water samples (Table 1), indicating a microbial source, which would support autochthonous inputs such as bacteria and algae. This is not surprising considering the surface water quality of rural Bangladesh, where wastewater is frequently disposed of directly 271 into ponds and other surface water bodies. 54 The EEMs of the small volume surface water samples showed a peak in the region of amino acid-like fluorescence near excitation/emission of 275/350 nm (Figure 1) and provide further evidence of a microbial DOM source in this water type. However, there is also substantial humic fluorescence in the Peaks A and C regions (Figure 1 K-SW), which reflects contributions from soil and terrestrial plant sources to the DOM pool in surface water. Thus, the DOM of surface water is characterized by protein-like 280 as well as humic DOM.

Although we have only one fulvic acid isolate of a local streamwater (K-SW) sample, the characteristics of this surface water FA isolate are illustrative. The DOC of this surface water sample comprised greater than 40%, fulvi acid, which was greater than the FA fraction of all other large volume groundwater samples (Table 2). Further, the combination of high SUVA, reflecting terrestrial DOM contributions, and high FI and amino acid-like fluorescence, reflecting autochthonous microbial contributions, would also suggest that, like the small volume water sample, this fulvic acid was derived from a combination of microbial constituents as well as soil and vascular plant sources in the area surrounding the river.

293 **Groundwater DOM and FA Isolates.** The extensive 294 spatial survey of groundwater from different wells and depths

showed that there were significant differences in DOC 295 concentration as well as in the quality of DOM among 296 groundwater of different ages. Differences in DOC concen- 297 tration, SUVA, FI, and %AA-like fluorescence of small volume 298 samples were significant between young, younger than 5 year 299 old groundwater, and older groundwater (Table 1). All 300 groundwater samples have comparable RI values consistent 301 with reducing conditions indicated by elevated levels of Fe and 302 Mn concentrations in groundwater compared to surface water 303 (Table 1). Further, the results showing significantly (p < 3040.00001) higher concentrations of As, Fe, Ca, and TDN, mostly 305 as ammonia (Table S2) and lower sulfate concentrations in 306 older groundwater than in young groundwater (Table 1) are 307 consistent with progressively greater reducing conditions as 308 groundwater ages and subsequent occurrence of mineral 309 weathering (calcite dissolution) and redox reactions. 39,44 310 Additionally, significant differences, determined with MAN- 311 OVA ($R^2 = 0.38$; p < 0.0001), were also observed for DOC, 312 SUVA, FI, and percent amino acid-like fluorescence between 313 groups of groundwater. The lower FI, lower AA-like 314 fluorescence, and higher SUVA in the two groups of older 315 groundwater samples suggest inputs of "terrestrial" sources, 316 such as lignaceous and aromatic C compounds, common in 317 dispersed sedimentary organic matter. One explanation for the 318 more terrestrial DOM signatures in the older groundwater is 319 that DOM is mobilized from sediments as a consequence of 320 microbial Fe reduction and Fe mineral dissolution. This 321 scenario would be consistent with other studies showing that 322 sedimentary organic matter contributes substantially to the $_{323}$ groundwater DOM pool. 55,56 DOM mobilization from sedi- $_{324}$ ments is supported by the results of sediment incubations from 325 our earlier study conducted also at Araihazar, which showed an 326 increase in DOC and Fe concentrations and DOM terrestrial 327 fluorescence signatures after incubation with native ground- 328 water.33

Table 2. Characteristics of Six Large Volume Whole Waters, Including Site Description, Solute Chemistry (total dissolved arsenic (As_T), % Arsenic as As(III), Total Dissolved Fe (Fer.), Dissolved Organic Carbon (DOC), and Fulvic Acid (FA) content and concentration), and DOM Optical Spectroscopic Properties (specific UV absorbance (SUVA) Fluorescence index (FI), and Amino Acid-like (AA-like) fluorescence), Collected from Site K and Site B Groundwater and Site K Surface Water (K-SW) in October 2009

			_		_			
optical spectroscopic properties	AA-like (% of fluorescence)	5.7	24	8.0	4.0	3.3	4.1	
pectroscop	HI	1.65	1.57	1.46	1.49	1.57	1.57	
optical s	$\frac{\mathrm{SUVA}^c}{(\mathrm{L\ mg}^{-1}\ \mathrm{m}^{-1})}$	2.66	1.60	1.72	2.24	2.52	2.73	
	FA conc. (mg FA L ⁻¹)	4.63	0.04	0.26	1.28	0.54	1.70	
	FA content (% of DOC)	42.7	2.80	8.30	25.0	13.3	34.4	
solute chemistry	$\frac{\mathrm{DOC}^b}{(\mathrm{mg~C~L}^{-1})}$	4.80	0.59	0.92	0.93	2.08	2.51	
	$({ m mg} \ { m L}^{-1})$		0.5			9.6		
	%As (III)	40	06	100	94	98	98	
	$(\mu g \stackrel{\mathrm{As_{T}}}{\mathrm{L}^{-1}})$	23.3	2.0	69	363	213	340	
	Age ^a (yrs)	N.A.	<>	<>	> 30	5-30	5-30	
	depth Age ^a (m) (yrs)	0	7.50	11.0	14.8	11.4	14.0	
site description	longitude (degrees)	90.6101	90.6284	90.6281	90.6280	90.640	90.640	
	latitude (degrees)	23.7916	23.7948	23.7941	23.7932	23.780	23.780	
	sample	K-SW	K12.1	K10.2	K8.3	B11	B14	

^aWater age category based on 3H/3He dating reported in Stute et al.³⁹ and Radloff.⁴⁵ ^bDOC concentrations are lower in large volume whole waters collected in October 2009 post-monsoon than in small volume samples collected in March 2008 during dry season (Table S1). EHigh SUVA values (>5.0) may indicate additional Fe absorbance in the UV-range (Weishaar et a Similarly, the higher fulvic acid content of large volume 330 isolates B11, B14, and K8, representative of older groundwater 331 (>5 years old), compared to the younger K10 and K12 332 groundwater (Table 2) indicates that the older groundwater has 333 higher amounts of dissolved humic substances, most likely 334 originating from organic material in the sediments. Also, the 335 FAs isolated from large volume isolates of older groundwater 336 had higher aromatic C concentration, lower FI values, and 337 lower amino acid-like fluorescence than FA isolates of young 338 groundwater (Table 3), consistent with the trends observed in 339 t3 the small volume samples that point to a terrestrial, 340 sedimentary DOM source. The lower amino acid-like 341 fluorescence was also evident in EEM spectra of the large 342 volume isolates of old groundwater (Figure 1), which were 343 instead dominated by fulvic acid-like fluorescence.

By contrast, several lines of evidence suggest that DOM in 345 younger groundwater (<5 years old) is influenced by microbial 346 sources. The average FI value of small volume samples of young 347 groundwater is similar to surface water, and the average value of 348 %AA-like fluorescence is the highest among all water types 349 (Table 1), likely reflecting the influence of sewage-polluted 350 surface water on young groundwater. This influence was 351 identified at the same study site⁵⁷ in young groundwater, which 352 has been shown to contain high counts of the fecal indicator 353 bacterium E. coli.⁵⁸ The lowest average SUVA value in young 354 groundwater suggests that the young groundwater is the least 355 influenced by terrestrially derived DOM among all water types 356 (Table 1). Similarly the large volume isolates of young 357 groundwater from sites K10 and K12 had low fulvic acid 358 content (Table 2), and the FA isolated from those sites had 359 higher FI and amino acid-like fluorescence and lower %C 360 (Table 3) than older groundwater.

Nonhumic DOM. In addition to humic DOM in Araihazar 362 groundwater, nonhumic DOM is present in both young and old 363 groundwater. Although nonhumic fractions were not measured 364 directly, the influence of nonhumic DOM can be observed in 365 the optical spectroscopic properties of whole water samples in 366 this study. For the five large volume isolates described above, 367 fluorescence properties were measured of the initial whole 368 water, prior to isolation. The whole water samples comprise 369 multiple fractions of organic matter, including fulvic acids as 370 well as nonhumic and potentially more labile compounds, such 371 as amino acids, carbohydrates, low molecular weight organic 372 acids, bases, and neutrals. The FI and percent amino acid-like 373 fluorescence of whole water samples prior to FA isolation 374 (Table S1) were higher than for the FA isolates (Figure 2) 375 f2 because the nonhumic fraction was still present in those whole 376 water samples. The nonhumic fraction, therefore, reflects 377 contributions to the DOM pool from microbial and other 378 nonhumic DOM sources. FI values of whole waters are known 379 to be higher when compared to FA isolates, in part, because of 380 the presence of other fluorescent microbial compounds, such as 381 extracellular quinone-containing exudates, that would act to 382 raise the FI.⁵⁹ These microbial sources may include microbial 383 exudates or microbial products from lysed microbial cells, 384 which may derive from sewage-influenced surface water drawn 385 laterally or to depth or from the sediments themselves. Other 386 microbial consortia, including bacterial taxa capable of Fe, As, 387 and humics reduction, which are key members of the 388 community in Araihazar sediments, 43 may also contribute to 389 the amino acid-like fluorescence, which is known to represent 390 biodegradable DOM, ⁶¹ that was observed in the whole water ³⁹¹ DOM. Therefore, this nonhumic fraction is also potentially 392

Table 3. Elemental Content, 13C-NMR Distribution of C Functional Groups of Fulvic Acid Isolates, and Optical Spectroscopic Characteristics of Fulvic Acids Isolated from Site K and Site B Groundwater and Site K Surface Water (K-SW)

	elem	elemental content	ntent		13C N.	MR distribution (%	¹³ C NMR distribution (%) and C functional group distribution ^a	oup distribution ^a			optical spect	roscopic	optical spectroscopic characteristics
sample	∪ <u>%</u>	C N C:N (%) ratio	C:N ratio	aliphatic (60–0)	aliphatic (60–0) heteroaliphatic (90–60)	acetal (110–90)	acetal (110–90) aromatic (165–110) carboxyl (190–165) carbonyl (220–190)	carboxyl (190–165)	carbonyl (220–190)	Ar:Al ratio	$ \begin{array}{c c} \text{Al} & \text{SUVA} \\ \text{io} & (\text{L mg}^{-1} \text{ m}^{-1}) & \text{F} \end{array} $	H	AA-like fluorescence (%)
-SW	44.3	44.3 2.9 15.28	15.28	48	15.62	0.01	19.43	13.61	3.27	0.40	4.36	1.69	11.0
K12.1	45.2	45.2 1.41	32.1	46.1	8.57	1.92	24.7	16.1	2.58	0.54	3.86	1.48	4.9
K10.2	28.8 1.30	1.30	22.2	63.2	11.4	0.12	14.1	10.6	0.59	0.22	5.16	1.45	4.1
K8.3	15.4	0.77	19.9	47.1	5.76	3.34	30.2	11.9	1.75	0.64	4.71	1.40	3.7
B11	51.3	1.57	32.7	53.1	9.30	0.56	19.6	16.0	1.47	0.37	3.96	1.43	3.5
B14	50.6 1.50	1.50	33.7	45.3	10.4	2.98	25.8	14.1	1.50	0.57	4.15	1.41	3.3

labile and important for setting in motion reductive dissolution 393 of Fe minerals.

Humic DOM and As Mobility. Having illustrated the 395 differences between two pools of groundwater DOM, humic 396 and nonhumic, we now attempt to shed light on the variable 397 relationship, either positive, negative or a lack thereof, between 398 groundwater As and DOC concentrations observed in several 399 field studies.^{26–28} There was no significant correlation between 400 whole water DOC concentration and As in this set of shallow 401 groundwater samples from Bangladesh (Figure S2) and only a 402 weak correlation for the five large volume isolates (Figure 2). 403 Instead, the FA concentration of groundwater had significant 404 positive relationships with As and Fe (Figure 2). We 405 acknowledge that the sample size is small given the extensive 406 sampling and preparation time needed for FA isolation (see 407 Supporting Information); however, FA isolation is a highly 408 representative technique. 46 Moreover, the change in DOM 409 quality supports the transition of DOM source from nonhumic, 410 microbially derived material in young water to humic, 411 terrestrially derived material as the groundwater ages. Although 412 FA content would be expected to increase as a result of 413 microbial degradation of labile fractions that may leave behind 414 more recalcitrant components, the FI of such microbially 415 derived FAs is typically higher (>1.50). 48,62 Instead the isolates 416 have a lower FI (\sim 1.40 to 1.43 (Figure 2), which suggests that 417 the original source of fulvic acids in older groundwater is 418 probably not labile microbial organic matter transported from 419 other zones but rather sedimentary material from the local 420

We interpret the correlation between the chemical character- 422 istics (aromaticity, FI, % amino acid-like fluorescence) of FAs 423 and both As and Fe (Figure 2) as indicating the involvement of 424 humic DOM in reactions that maintain As in solution. Prior 425 laboratory and field studies suggest a likely mechanism through 426 formation of As-Fe-DOM complexes. 22-24 Although humic 427 DOM, As, and Fe may all be present as free molecules as a 428 consequence of reductive dissolution, the stoichiometric 429 relationship showing proportionately greater humic DOM 430 content in water with greater dissolved As or Fe concentrations 431 (Figure 2) implies potential complex formation among these 432 species. Indeed the aromatic carbon content of humic 433 substances has been shown to have a strong positive correlation 434 with Fe(III).⁶³ Multiple laboratory studies have also shown that 435 As is able to form ternary complexes with Fe and DOM. 25 436 Using Pahokee Peat Humic Acid as their DOM source, Sharma 437 et al.²³ demonstrated that at neutral pH As(V) formed colloids 438 and complexes with Fe-DOM but not with DOM in the 439 absence of Fe. Formation of complexes between As and DOM 440 in the presence of metals but not in their absence was also 441 demonstrated for DOM extract from compost.⁶⁴ Fulvic acids 442 readily complex Fe at neutral pH, and As can bind strongly to 443 such complexes via bridging mechanisms²⁴ to stay in solution. 444 Therefore, it is not unexpected that humic DOM correlates well 445 with As in the presence of Fe in the neutral pH groundwater of 446 our study. Using a molecular weight of 470 Da for groundwater 447 FA from a similar environment, 65 we calculated that molar 448 Fe:As:FA ratios were about 47:1.3:1 (see slopes of linear 449 regressions in Figure 2). The average molar ratios of Fe to As 450 for large volume isolates from groundwater younger than 5 451 years old ranged from 40 to 68 in our study, which is slightly 452 higher than the 20 to 55 range reported in laboratory studies of 453 As-Fe-DOM complexes under controlled conditions. 23,24

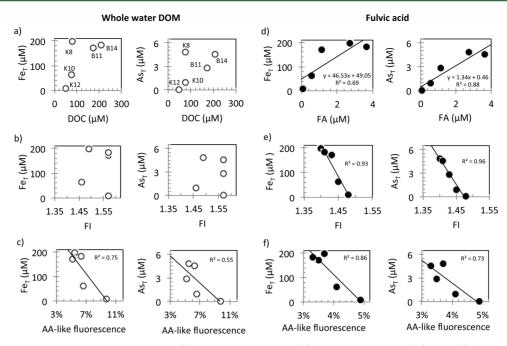


Figure 2. Scatterplots of total dissolved As and Fe vs (a) DOC concentration, (b) fluorescence index (FI), and (c) amino acid-like (AA-like) fluorescence of large volume whole water samples prior to isolation at Site B and Site K. (d) Fulvic acid concentration of large volume samples. (e) Fluorescence index and (f) amino acid-like fluorescence of fulvic acids isolated from the large volume samples at Site B and Site K (sample labels shown in panel a). Regression lines shown only for significant relationships (p < 0.05).

Organic matter that passed through the approximately 0.9 456 µm Ballston filter, which may consist of colloids and small 457 particles, also appears to have undergone associations with As 458 and Fe. Prior to isolation of the FAs in this study, we recovered 459 orange-colored organic matter on the filters after filtering the 460 large volume water samples that were acidified to pH 2 to prevent Fe precipitation. This orange-colored retentate (material retained on filters) for three samples, K12, K8, and B11, was redissolved in 0.1 N HCl, and its FA fraction was also isolated. 46 The FA content of the filter retentates was substantially higher (FA of 18%, 28%, and 37% for samples K12, K8 and B11, respectively) than in the corresponding large volume isolates (Table 2), suggesting that organic matter in larger molecules was more humic than DOM in groundwater. The FAs of filter retentates also displayed the most terrestrially derived signatures with the much lower mean FI values (FI values of 1.24, 1.23, and 1.29 for samples K12, K8 and B11, 472 respectively) than dissolved FA isolates of groundwater from the same three sites (Table 3; mean FI = 1.54).

Therefore, the influence of colloidal and particulate organic matter and Fe on As mobility also merits further study. Ultrafiltration of As- and DOM-rich groundwater from Hetao Plain, Inner Mongolia has demonstrated stronger association between As and smaller organic colloids than with larger Fe colloids. Additionally, SEM images, EDS analysis, and synchrotron XRF analyses in their study confirmed the association of As with natural organic matter (NOM) with molecular weights of 5–10 kDa.

Implications for Groundwater As Enrichment. Deltaic 484 regions with elevated As concentrations in reducing ground-485 water tend to be rich in sedimentary organic compounds 486 derived from the deposition of plant and animal biomass 487 detrital compounds over time. In Bengal Deltaic sediments, 488 Meharg et al. postulated that As and organic matter were co-489 deposited during the Holocene, 67 especially in productive

coastal wetland ecosystems that developed at the time. Similar 490 co-deposition of As and organic matter may be true for other 491 deltaic or depositional environments with elevated As in 492 groundwater, such as the Red River Delta, Mississippi River 493 Delta, Okavango Delta, and Pearl River Delta; however, 494 previous studies have focused more on relationships with 495 total organic matter content of sediments or DOC of whole 496 water. Our results suggest that the relationships between As, Fe, 497 and humic DOM that may influence As mobility in reducing 498 groundwater environments deserve further exploration. With 499 respect to the linear relationships between As, DOM, and Fe 500 observed in our study, it is possible that As-DOM-Fe 501 complexes were sequestered in the sediments simultaneously 502 during deposition, thousands of years ago. Co-deposition of 503 organic matter and As with Fe-containing minerals, as proposed 504 by Meharg et al.,67 and later ternary complex formation after sos reductive dissolution is a plausible scenario to explain why 506 groundwater with elevated As contains more terrestrially 507 derived, humic DOM than groundwater with low As 508 concentrations. Regardless of whether the association with As 509 is via dissolved complex or colloidal complex, our results 510 suggest that the humic fraction of DOM, in particular, 511 undergoes reactions with As. Differentiating the humic, 512 chemically reactive, from the nonhumic, biologically reactive, 513 fraction of DOM in reducing groundwater of deltaic aquifers is 514 a first step to better elucidating that role.

It is also important to note that the labile, biologically 516 reactive nature of DOM in young groundwater at our study site 517 should stimulate microorganisms capable of Fe reductive 518 dissolution leading to As mobilization. However, instead of 519 finding high As concentration in young groundwater with labile 520 DOM, our results showed that As concentration was 521 significantly higher in old groundwater, which contained higher 522 fulvic acid content, DOC concentrations, and terrestrially or 523 sediment-derived DOM sources, compared to young ground-

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525 water. These results indicate that the nature of the As-Fe-526 DOM relationship is more complex than the simple picture of 527 labile DOM as a driver for As mobilization. The involvement of 528 humic DOM in reactions with As and/or Fe, such as 529 complexation, competitive sorption, or electron shuttling, may 530 help explain why As concentrations are elevated in old 531 groundwater. Field research and experimental studies directly 532 investigating the role of humic DOM for As mobilization under 533 environmentally relevant conditions are needed.

In addition to its potential influence on As mobility in reducing aquifers, the chemical reactivity of humic DOM influences treatment and remediation strategies used for dealing with high As concentrations in drinking water. For example, the presence of colloidal As—Fe—DOM associations may hamper filtration efforts, and competition with humic or fulvic acids inhibits several treatment processes, such as As sorption on Fe minerals, As(V) adsorption during coagulation with ferric chloride, and As sorption on Fe nanoparticles.

Despite the important role of organic matter in As cycling 544 and the clear value of elucidating its chemical structure to 545 understand and predict biogeochemical reactions, much still 546 remains to be understood regarding the quality of DOM found 547 in groundwater of different age and As concentrations. This 548 study provides new information on the role of DOM fractions 549 as being both biologically and chemically reactive, with the 550 humic DOM acting to maintain As in solution in reducing 551 groundwater. In particular, we have shown that DOM in older 552 groundwater (>5 years old) is characterized by higher FA 553 content, more terrestrial, sedimentary sources, and less 554 microbial spectroscopic signatures than DOM in younger 555 groundwater or surface water. The highly significant relation-556 ships we observed between As and Fe concentrations and FA 557 content, the FI, and amino acid-like fluorescence of FA isolates 558 further suggest that humic As-Fe-DOM interactions may 559 have occurred long ago during the precipitation of As-560 containing Fe minerals or recently as complexation reactions 561 after free As was mobilized from sediments. The presence of 562 humic DOM in groundwater and the heterogeneous distribu-563 tion of dispersed organic matter in sediments may, therefore, be 564 additional factors contributing to the large spatial variability in 565 As concentrations over short distances. In addition to 566 improving our understanding of the heterogeneity of elevated 567 groundwater As across the landscape, accounting for the quality 568 of DOM in groundwater of different ages is important to the 569 development remediation and treatment strategies.

ASSOCIATED CONTENT

571 S Supporting Information

572 The Supporting Information is available free of charge on the 573 ACS Publications website at DOI: 10.1021/acs.est.5b01962.

Detailed methods, supporting tables, and supporting figures (PDF)

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Notes
The authors declare no competing financial interest.

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