



Testing the importance of explicit glacier dynamics for future glacier evolution in the Alps

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Volume-area scaling vs. shallow ice approximation

Even though the **Open Global Glacier Model (OGGM)** is a rather simple model (concerning the implemented physics), the **volume-area scaling (VAS) glacier model** originally used by Marzeion et al. (2012) is even more basic. While the OGGM implements shallow ice approximation to drive the model glacier, the VAS model relies solely on volume/area and volume/length scaling principles.

What am I doing?! Currently I'm implementing the original volume/area scaling model used by Ben (at least I'm trying to).

Why am I doing it?! More complex models generally come with higher computational costs. While this may be necessary for certain detailed analysis, not all scientific questions call for such a high degree of accuracy. In addition, alpine glaciers will most likely not advance much in the coming decades. Hence, ice dynamics play a secondary role compared to ice melt.

What do I want to achieve?! In a nutshell: how far can I dumb the dynamic model down, while still producing reasonable results on a regional scale. To do so I will:

- 1 implement glacier model(s) with different levels of complexity in the OGGM framework.
- 2 investigate the strengths and weaknesses of the VAS model approach (sensitivity analysis).
- 3 complete regional commitment runs for the alpine region (past and future).

Mass balance model

The mass balance model implements a the **temperature index model**.

$$B(t) = \left[\sum_{i=1}^{12} (P_i^{\text{solid}} - \mu^* \cdot \max(T_i^{\text{terminus}} - T_{\text{melt}}, 0)) \right] - \beta^* \quad (1)$$

The annual specific surface mass balance $B(t)$ is computed as the difference of monthly solid precipitation fallen onto the glacier surface P_i^{solid} (mass input) and monthly positive melting temperature at the glacier terminus T_i^{terminus} (energy input), summed over a year. The temperature sensitivity parameter μ^* needs to be calibrated for every single glacier. Additionally, a mass balance bias β^* can be added/subtracted.

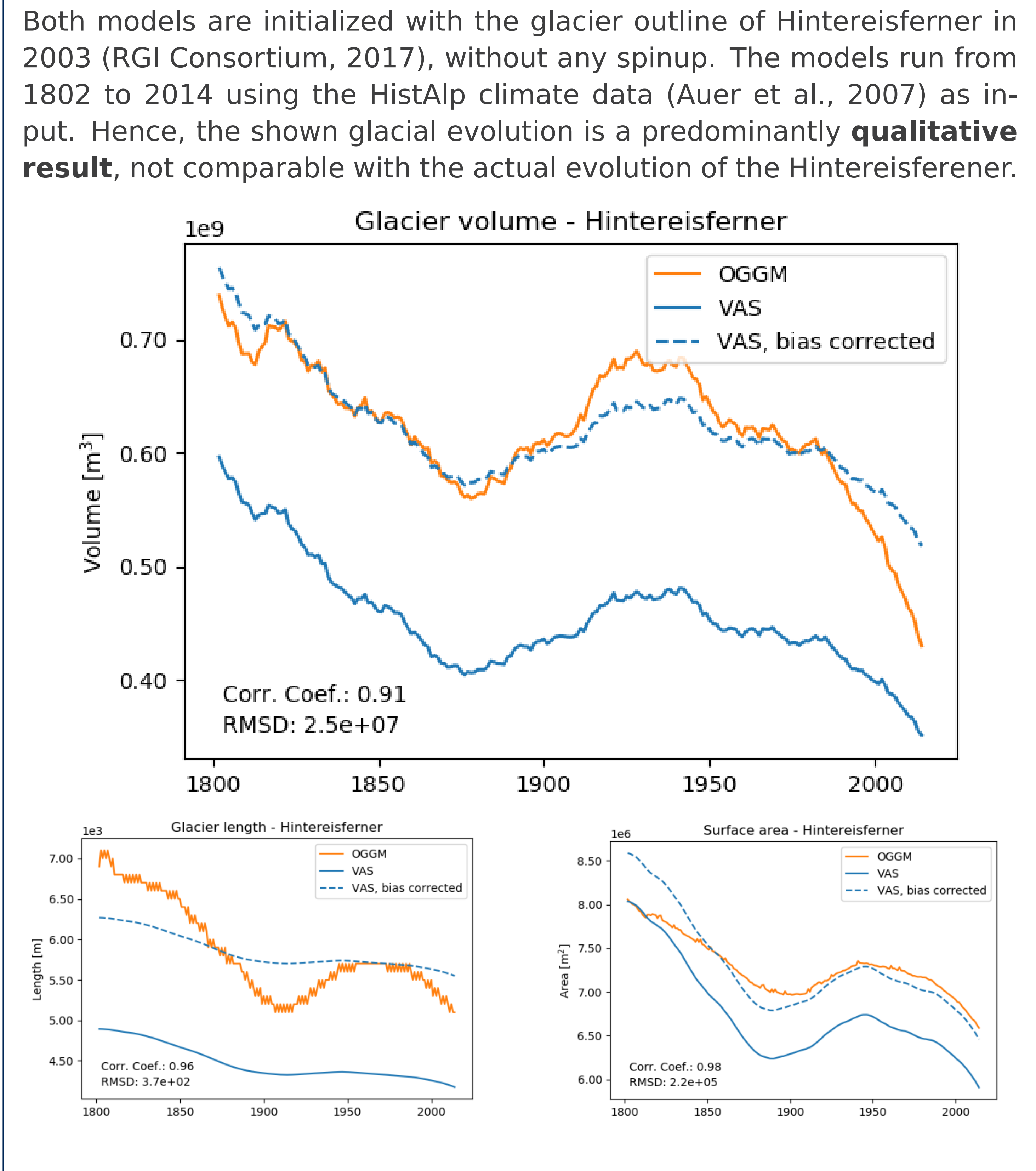
The scaling model

The used scaling laws are quite basics:

$$V(t) = c_A \cdot A(t)^\gamma, \quad V(t) = c_L \cdot L(t)^q \quad (2)$$

where c_A , c_L and γ , q are scaling parameters. For each year t the volume change $dV(t)$ is computed, given the specific mass balance $B(t)$ and the surface area $A(t)$. The resulting new volume $V(t+1)$ is used to compute the area change dA and length change dL , scaled by the corresponding response time τ_A and τ_L , resp. A change in glacier length $L(t+1)$ translates into a change in terminus elevation $z_{\text{term}}(t)$, which influences the specific mass balance $B(t+1)$.

Comparing the model performance on a single glacier



First small scale regional run - Rofental

Let's see if I produce something sensible here, until it has to be printed...

../plots/rofental.png

What next?!

- Given that this is just the first implementation step, the (qualitative) results are quite promising. However, some more work is required before deriving any conclusions. This includes the following tasks/questions:
- Does the length change have to be comparable?!
 - Finding appropriate/sensible start values...
 - Performance test against reference glaciers.
 - Regional (alpine) runs
 - Sensibility analysis

References:

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