



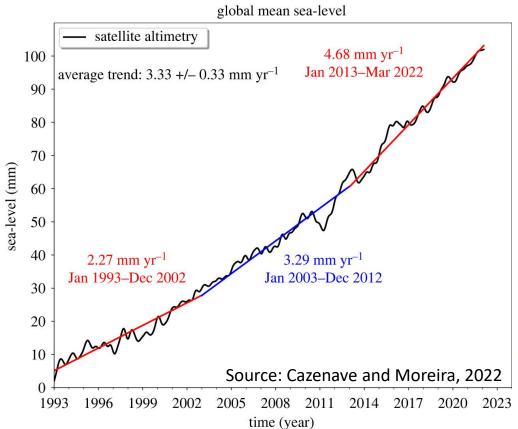
Adjoint sensitivities: Mechanisms for sea level variability in the Southern Ocean Indian sector under strong warming scenario

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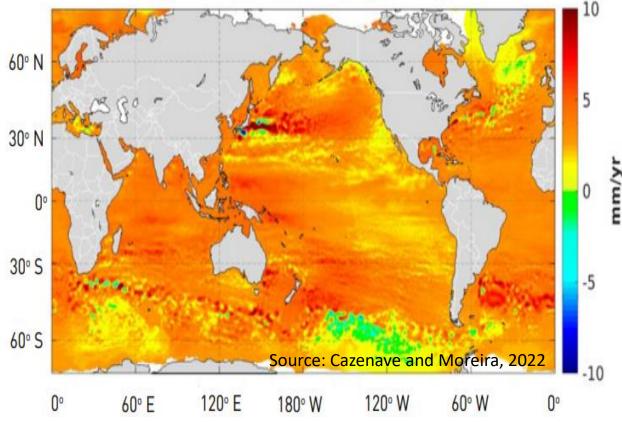
Institute of Oceanograhy, Universität Hamburg, January 13, 2022

Introduction



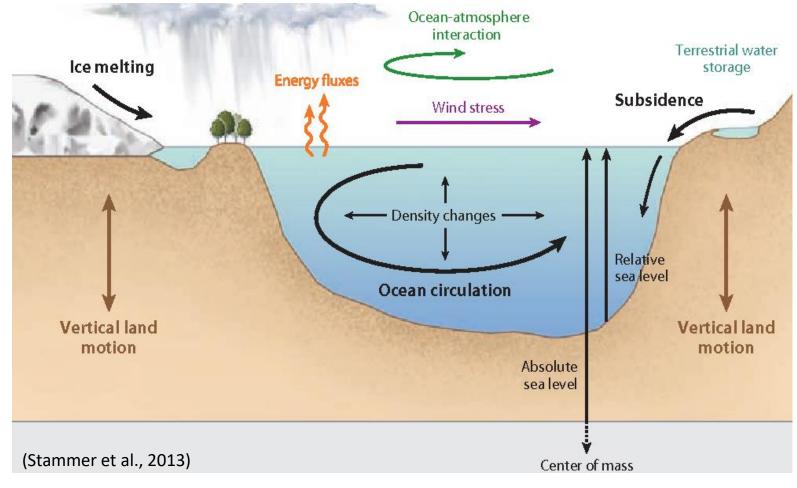
- Stabardiearisies level is rising
- GMSL rise has accelerated due to global warming
- Imporant!





- Sea level trend is not uniform spatially
- Different regions different mechanisms

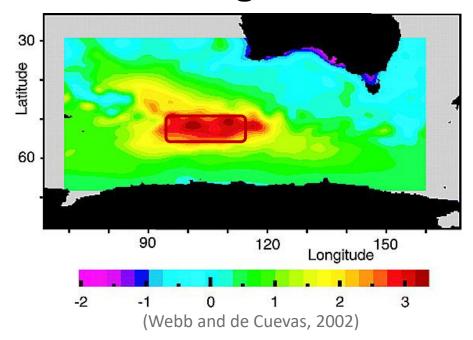
Processes that influence regional sea level



- (1) Dynamic Sea Level
- Ocean circulation
- Density changes
- **Wind Stress**
- Heat Flux
- Freshwater Flux
- (2) Other processes
- Land motion
- Solid earth changes in the long term

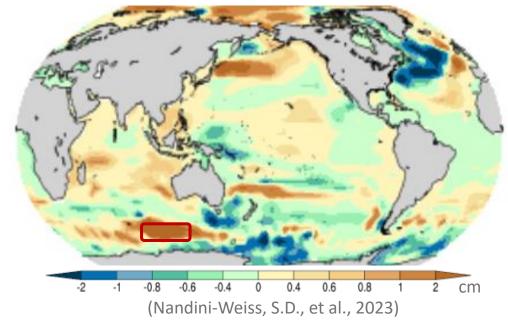
- Need to consider the specific region and timescale
- In most places, dynamic sea level is dominant on interannual to decadal timescale!

Motivation/Region/Previous studies



Amplitude of 1st mode of SSH (normalized to have unit variance when averaged over the region)

- A region of high SSH variablity
- Wind-induced Ekman pumping



Difference in standard deviation of monthly SSHa, RCP8.5(2081-2100)-Historical period (1986-2006)

 Increased SSH variability under global warming in Southern Ocean Indian sector, linked with wind stress

Knowledge gap: Mechanisms of low-frequency SSH variability (>3yr) in this region is to be explored

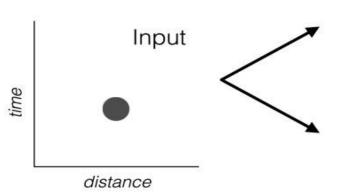
Research question:

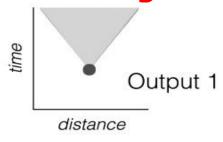
How sensitive of large-scale low-frequency dynamic sea level to surface forcing and whats are associated mechanisms?

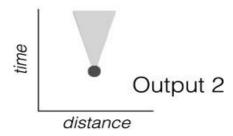
- 1. When/where/which forcing is able to generate long term SSH anomaly (a typical signal of low-frequency variability)?
- 2. Which forcing has the most contribution? Wind stress or Buoyancy fluxes?
- 3. What ocean dynamics are involved to produce SSH anomaly?

Methodology: sensitivity analysis

(a) Perturbation experiment

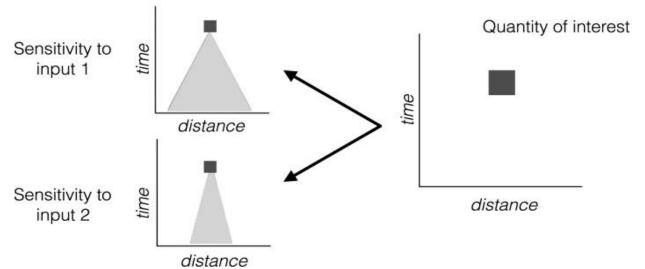






- 1. Perturbe one input
- 2. Run forward model
- 3. Obtain response in all outputs
- Ensemble of perturbations
- Expensive, re-run for every change applied

(b) Adjoint sensitivity experiment



- 1. Perturb one output (e.g SSH)
- 2. Run adjoint model ("backward" model)
- 3. Obtain linearized sensitivity to all inputs
- √ Efficient! One time simulation
- ✓ Identify physical processes and pathways via dynamic chain

Schematic of (a) traditional forward perturbation experiment and (b) an adjoint sensitivity experiment. (Jones et al., 2018)

What is adjoint sensitivity? & how to derive it?

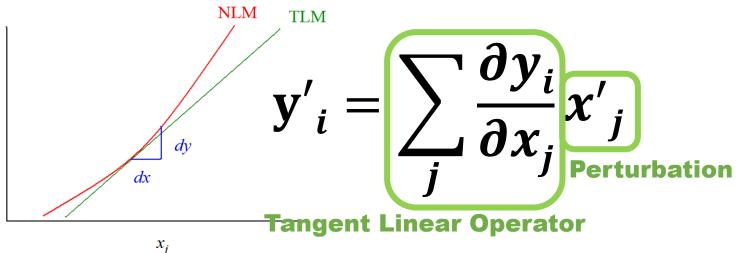
Nonlinear Forward Mode!

Maps inout x to output y

y = M(x)A graphical TLM schematic
Nonlinear Operator

Tangent Linear Model

First order approximation of nonlinear model



Adjoint Model

A transformation of tangent linear model

Cost Function
$$J$$

A function of the mode state $J = J(y) = J[M(x)]$

Adjoint sensitivity $\frac{\partial J}{\partial x}$

Reveal how the changes in x can affect J

$$\frac{\partial J}{\partial x_j} = \sum_{i} \frac{\partial y_i}{\partial x_j} \frac{\partial J}{\partial y_i}$$
Gradient

Adjoint Operator

Cost Function J

last 3 yr mean dynamic SL anomaly averaged over target region

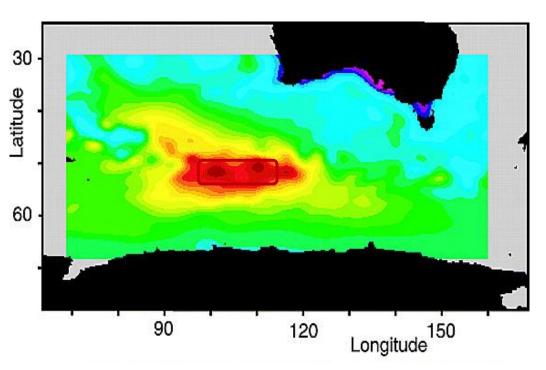
$$J = rac{1}{(t_2 - t_1)A} \int_{t_1}^{t_2} \int_{A} \eta \ \mathrm{d} A \, \mathrm{d} t$$

- A scaler value calculated from forward model
- A signal of large-scale sea level variability
- A forcing that drives adjoint model backward

Timescale:

(t2 – t1) is the last 3 year over a 10-year simulation

 η : Dynamic Sea Level (sea level anomaly above GMSL due to ocean dynamics)



Region of interest A: the box covering the first EOF mode [52-56S, 96-112E] (1x4 grids box)

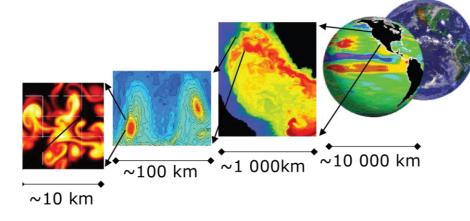
Model

MITgcm and its adjoint

- Navier-Stokes equations
- Adjoint code via automatic differentiation

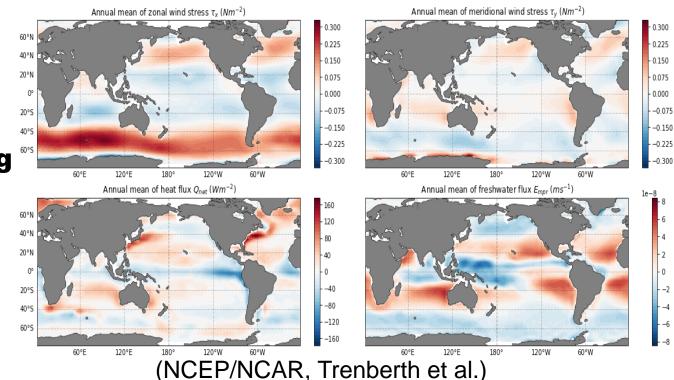
2
$\frac{\partial \mathbf{v}}{\partial t} + (f + \zeta)\hat{\mathbf{k}} \times \mathbf{v} + \nabla_{z} \times KE + w \frac{\partial \mathbf{v}}{\partial z} + g \nabla_{z} \cdot \eta + \nabla_{h} \Phi'$
$= D_{z^*,v} + D_{\perp,v} + \mathcal{F}_v,$
$\frac{\partial \Phi'}{\partial z} = g \frac{\rho'}{\rho_c},$
$\frac{1}{H}\frac{\partial \eta}{\partial t} + \nabla_{z^*}(s^*v) + \frac{\partial w}{\partial z^*} = s^*\mathcal{F},$
$\frac{\partial (s^*\theta)}{\partial t} + \nabla_{z^*}(s^*\theta v_{\text{res}}) + \frac{\partial (\theta w_{\text{res}})}{\partial z^*}$
$= s^*(\mathcal{F}_{\theta} + D_{\sigma,\theta} + D_{\perp,\theta}),$
$\frac{\partial (s^*S)}{\partial t} + \nabla_{z^*}(s^*Sv_{\text{res}}) + \frac{\partial (Sw_{\text{res}})}{\partial z^*}$
$= s^*(\mathcal{F}_S + D_{\sigma,S} + D_{\perp,S}),$

MIT general circulation model



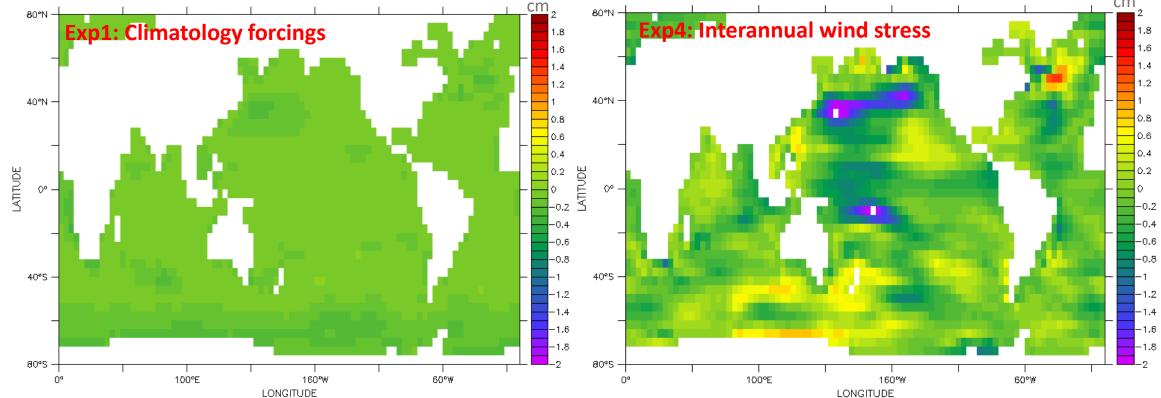
configuration

- 4°x4° spherical polar grid
- Quasi-global (80N-80S)
- 15 vertical layers (50m 690m)
- · Climatological monthly mean forcing
- Initiated from climatology
- Spin-up: 100 year
- Run forward 10 years
- Timestep: 12 hours
- SST/SSS relaxation
- Monthly snapshot fields are saved



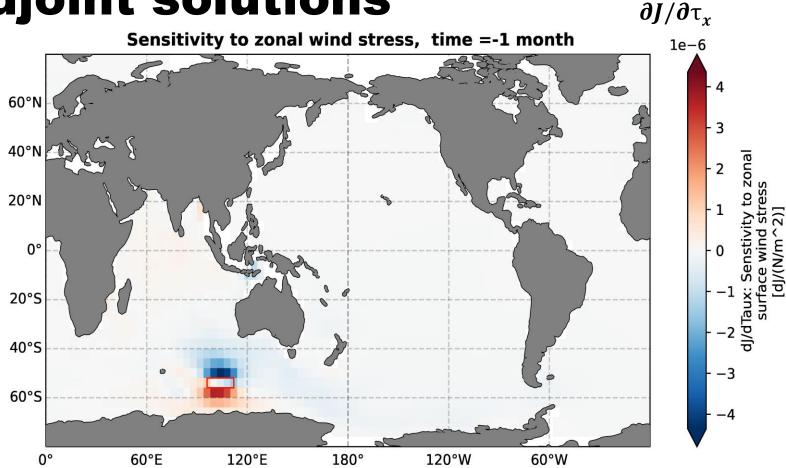
The forward perturbation experiment

- 1. Driven by climatological zonal & meridional wind stress, heat & freshwater fluxes
- 2. Replace one forcing with interannual-varying forcing (with interannual variability)
- 3. Compare changes in SSH standard deviation (RCP 8.5 scenario historical scenario)



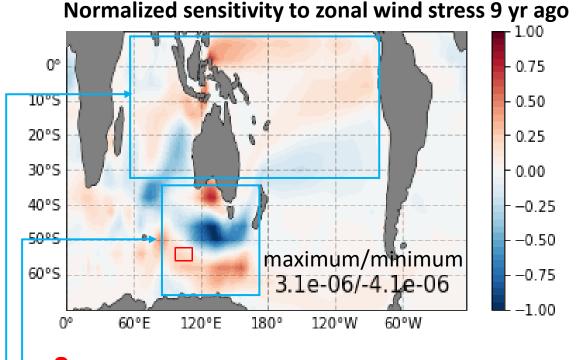
- Wind stress seems be to the most important to influence SSH variability
- But wehre and when wind stress is important?
- Can we assess of contribution of each forcing quantitatively?

The Adjoint solutions



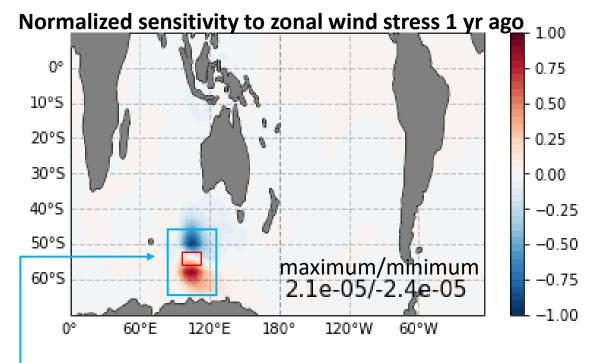
- Pinpoint sensitive regions with time backwards
- Red colors show areas where positive zonal wind stress anomaly at that time lag increase SSHa in the target region over the last 3-year mean period
- Blue colors show areas where positive zonal wind stress anomaly at that time lag decrease SSHa in the target region over the last 3-year mean period

sensitivity to zonal wind stress



9 years ago

- Most sensitive to local zonal wind stress between Australia & Antarctica
- Less sensitive to non-local wind stress over equator and subtropical

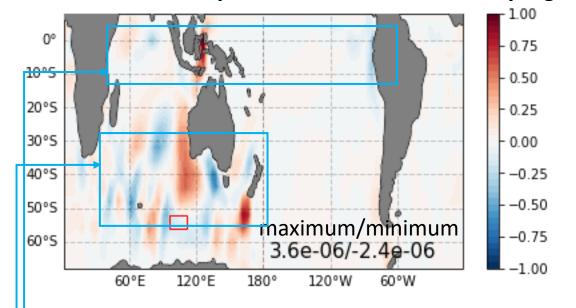


1 year go

Sensitive to local zonal wind stress
ONLY north & south of the target region

sensitivity to meridional wind stress

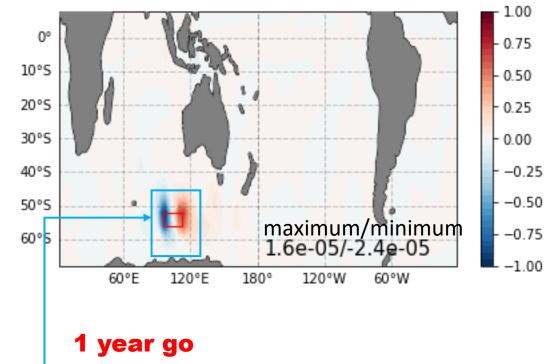
Normalized sensitivity to meridional wind stress 9 yr ago



9 years ago

- Mostly sensitive to meridional wind stress over broad Southern Indian Ocean
- Also sensitive to far wind stress over indonesian passage and Peru coast

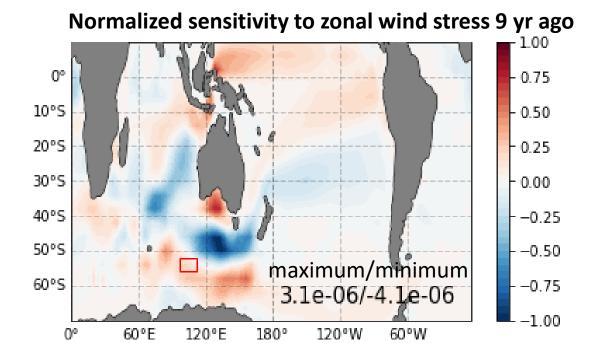
Normalized sensitivity to meridional wind stress 1 yr ago



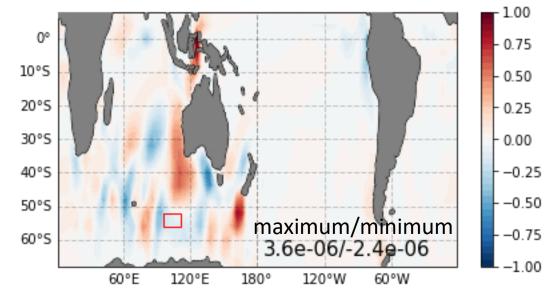
Sensitive to meridional wind stress ONLY west & east of the target region

Why sensitive in these regions? What does the pattern mean?

Mechanisms indicated by structure of sensitivity map



Normalized sensitivity to meridional wind stress 9 yr ago

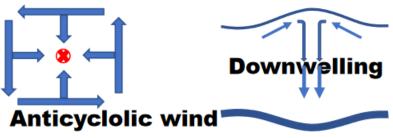


Strip structure →

- waves, gyres?
- Northeast-southwest tilt
 → Beta effect

$\textbf{Dipole structure} \rightarrow$

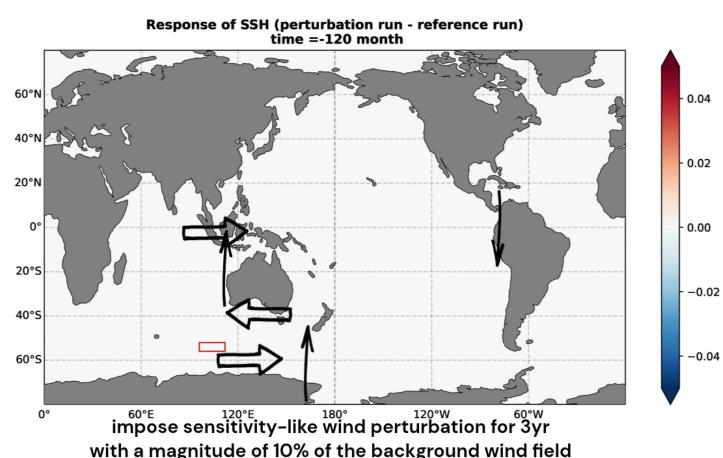
- Convergence/divergence
- Ekman pumping/suction

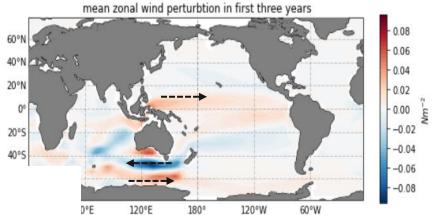




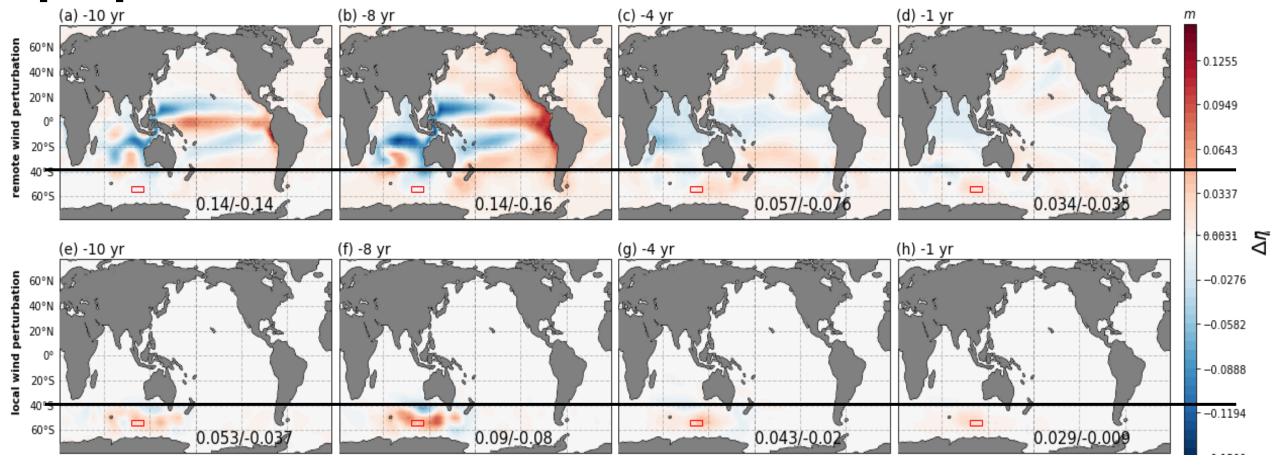
Mechanisms revealed by sensitivity-based perturbation experiments

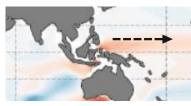
- 1. Impose time-evolving wind stress perturbation for 3 years over the background wind field
- 2. Patterns resemble its sensitivity with a magnitude of 10% of background wind field
- 3. Compare the SSH with reference run





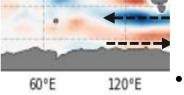
Split perturbation into local & nonlocal





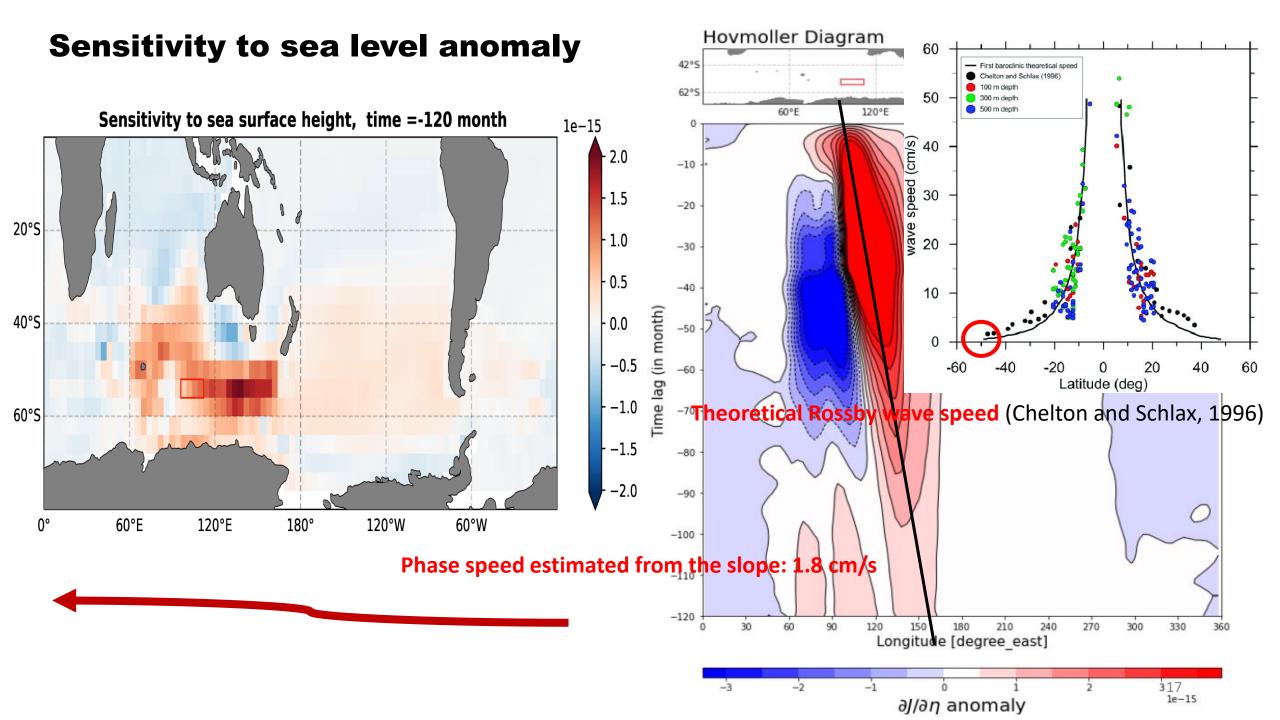
Strengthened Equtorial zonal wind?

- Kelvin wave → coastal waves → reflected Rossby wave → passing through Tasman Sea
- Takes several years to arrive

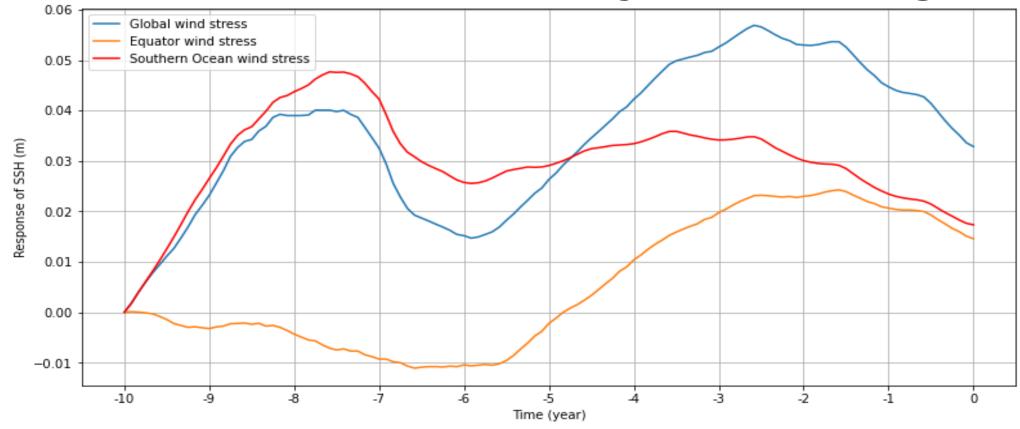


Strengthened Southern Ocean wind stress anomaly?

- Westward Rossby wave (advected with ACC)
- Immediate response

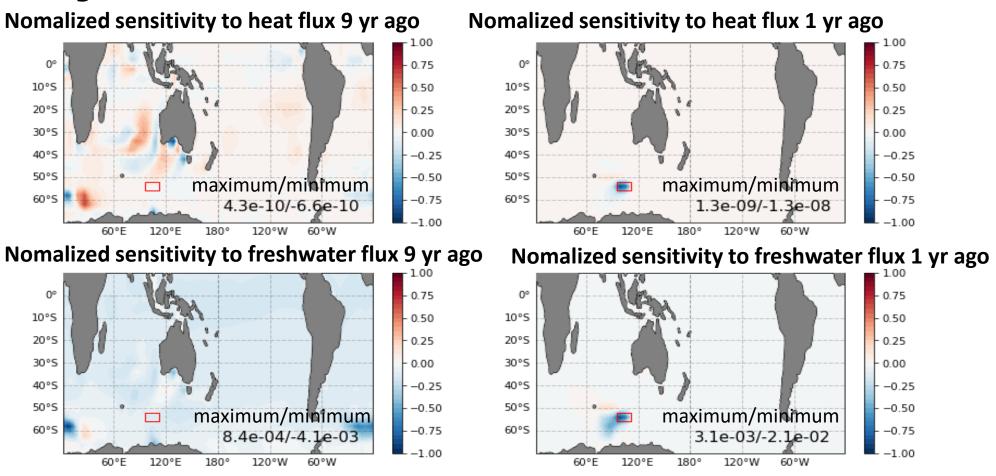


Time evolution of mean SL anomaly in the box region



- Red Curve: Southern Ocean wind stress perturbation has immediate effect on sea level in the target region
- Orange Curve: Equatorial wind stress perturbation causes negative SSHa at first and takes ~5 yr to produce positive SSHa in the target region
- Blue curve: the reponse of global wind perturbation is almost = the combined response of other two perturbation experiments
- Response is linear/ Large-scale surface circulation is linear

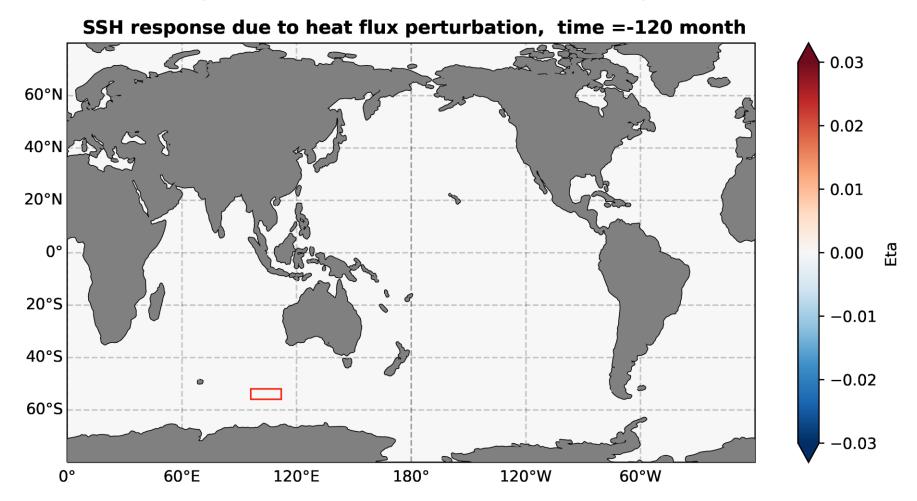
sensitivity to surface heat and freshwater flux



- Similar teamporal and spatial patterns
- Sensitivities moving along ACC indicates ACC propogates temperature&salinity anomaly
- Sensitivity to heat flux sugguests ocean surface heating/cooling
- Sensitivity to freshwater flux sugguests ocean freshening/salinification

Heat flux-perturbation experiments

What if impose sensitivity-like heat flux perturbation for 3 yr?



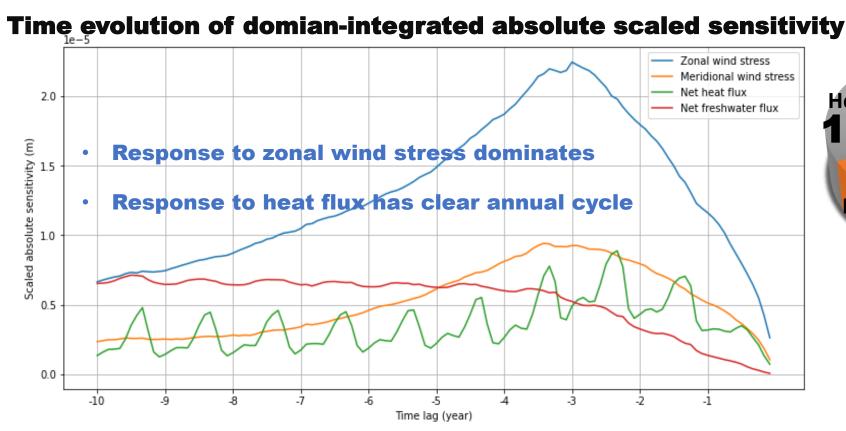
- Generate positive SSH anomaly propogating along ACC mean flow
- Sea level response is less significant

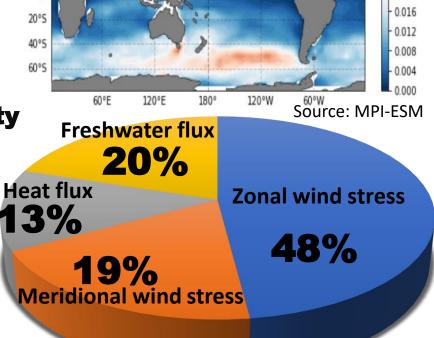
Relative contribution under a global warming scenario

$$J(x_{i}) = J(x_{i0}) + \frac{\partial J}{\partial x_{i}}\Big|_{x_{0}} \cdot (x_{i} - x_{i0}) + O(\|x_{i} - x_{i0}\|^{2})$$

Cost Function Response ≈ Sensitivity x Forcing Perturbation

- Strong sensitivity & high forcing variability →enhanced response
- Strong sensitivity & low forcing variability → diminished response



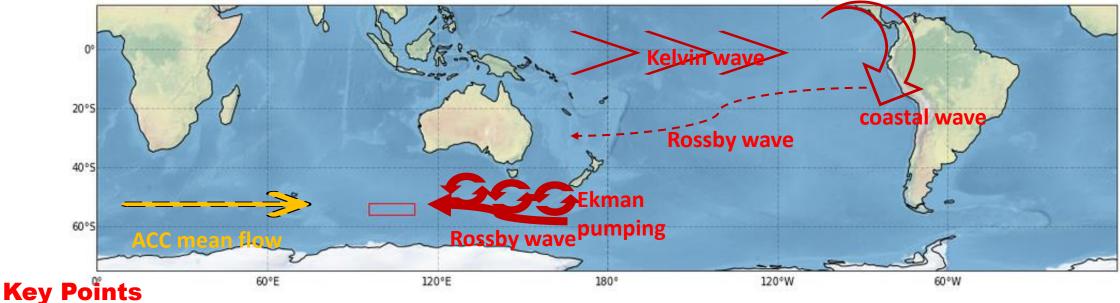


Relative contribution

0.032

Summary

 I use the adjoint sensitivity analysis to reveal how local and remote atmospheric forcing anomalies contribute to low-frequency Sea-Level variability in the chosen region



- 1. Large-scale circulation and SSH variability is essentially forced by wind stress
- 2. Quantity of interest is mainly driven by wind-induced Ekman pumping and Rossby wave between Australia and Antarctica
- 3. Local and regional winds is the leading factor while remote winds take years
- 4. Buoyancy fluxes modulate sea level variation via ACC mean flow
- 5. Momentum and buoyancy-driven mechanisms have opposite pathways in Southern Ocean

Application

 If forcing anomaly is observed, the associated mechanism is expected to be excited (sensitivity-based observation system design)

