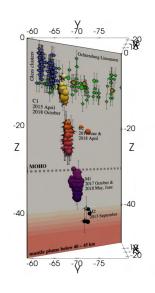
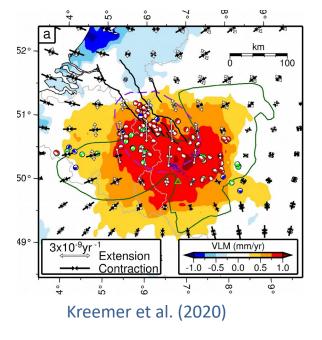
# LaMEM short course

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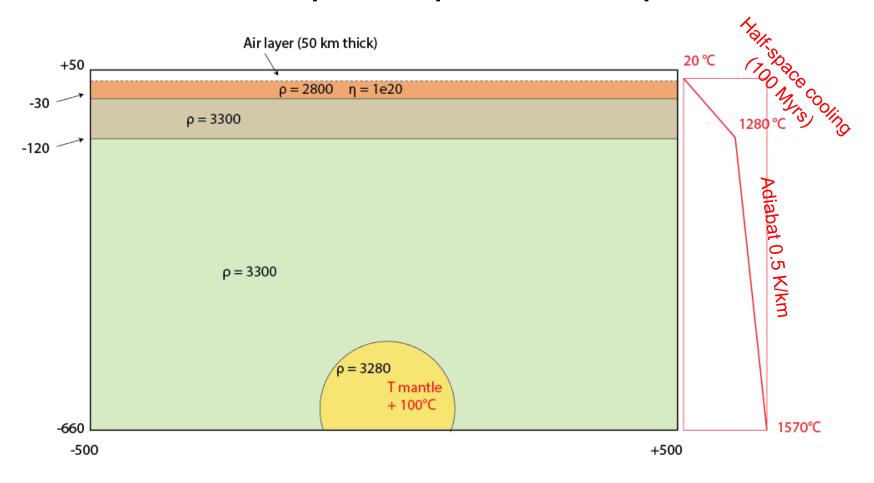


#### The "Eifel" anomaly



Plume emplacement in a single LID

 Study the topography effect of plume emplacement at the base of the lithosphere



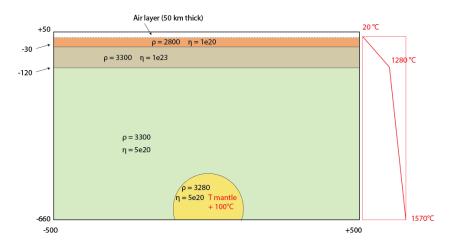
- Create a simplified plume setup using the sketch (start with constant density and viscosity).
- How much uplift can we expect from the plume emplacement?

#### TIPS:

- Start from the falling sphere setup with temperature and adapt it
- You can create a plate using the AddBox! function (similar to AddEllipsoid!)

```
# add single plate using Addbox!
AddBox!(model; xlim = (minx, maxX),
               ylim = (minY, maxY),
               zlim = (minZ, maxZ),
               Origin = nothing, StrikeAngle=0, DipAngle=0,
                       = LithosphericPhases(
                                                  Layers=[30 90],
               phase
                                                  Phases=[1 2 3])
                      = HalfspaceCoolingTemp(
               Т
                                                  Tsurface
                                                             = Tair,
                                                  Tmantle
                                                             = Tmantle,
                                                             = 100
                                                  Age
```

- LithosphericPhases() creates 3 layers and attribute phase 1, 2 and 3 to them
- HalfspaceCoolingTemp() applies a cooling age of 100 Myrs to phase 1 and 2

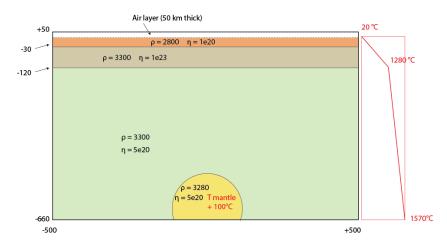


#### TIPS:

 The plume phase can be added either using AddEllipsoid!(), setting ConstanTemp(1650)

#### Or

• Using the equation of a circle with center = [x = 0, z = -600] and radius 100 km, filtering the particle coordinates inside the circle and applying a  $\Delta T$  +100 and a phase



#### TIPS:

 Linear gradient can be applied after defining lithosphere temperature profile such as:

```
Z = model.Grid.Grid.Z;
model.Grid.Temp = model.Grid.Temp -Z.*Adiabat;
```

Where Adiabat = 0.5 K/km for the mantle (see Turcotte and Schubert for more information)

Don't forget to overwrite the air temperature after applying the adiabat

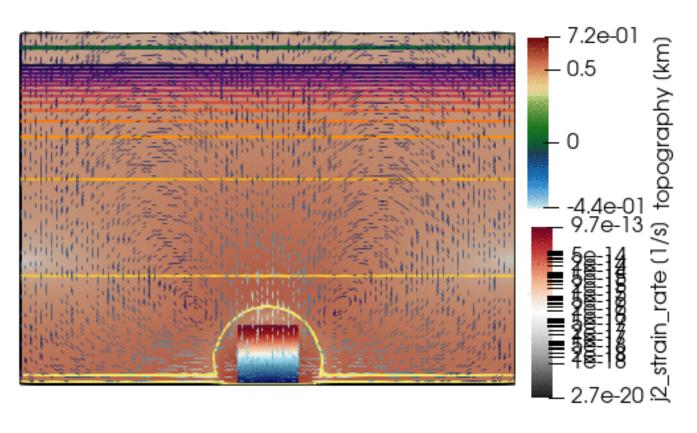
```
model.Grid.Temp[Z.>0.0] .= 20.0;
```

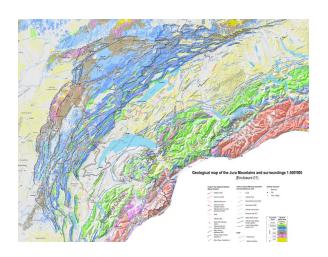
Let's add creep laws. Use quartzite for the crust, dry peridotite for the mantle

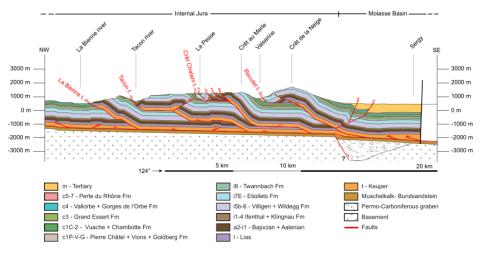
- Perform the simulation for < 50 timesteps, what is happening?</li>
- Correct plume density, use a lower value such as 3280 and increase the number of timesteps to 600
- Compare the strain-rate field with simulation using isoviscous rheology
- What difference does it make to use non-linear creep laws?

```
Solver( SolverType = "direct",
DirectSolver = "mumps",
PETSc_options = [ "-snes_rtol 1e-2", "-snes_max_it 100"] )
```

### 0 Myrs



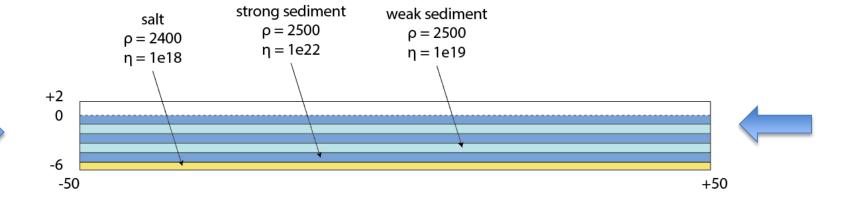




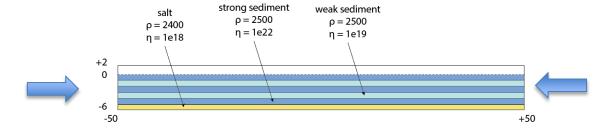
**Jura Mountains** 

### Fold and thrust belt

- How are folds generated?
- What is the relation between folds and thrusts?



 $\dot{\varepsilon}_{xx}$  = 1e-14 (compression)



#### **TIPS**

- Copy and past the falling sphere setup with free surface
- Change the resolution to the following

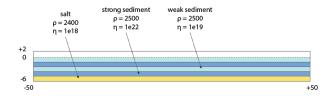
```
Grid( x = [-50.0,50.0], y = [-0.25,0.25], z = [-6.0,2.0], nel = (256,1,48))
```

Notice that the size in the Y direction is chosen to be 0.5, so that the cell size is as close as possible to a cube.

Update boundary condition to use prescribed xx strain-rate

```
exx_num_periods = 1,
exx_strain_rates= [-1e-14],
noslip = [0, 0, 0, 0, 0, 0]),
```

Choose nstep\_max = 150



#### **TIPS**

To create a layer, you can use the AddBox!() function

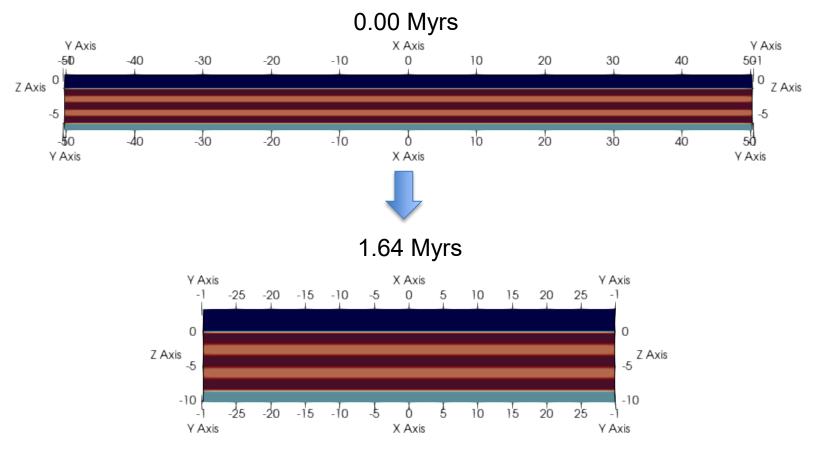
```
AddBox!(model; xlim=(model.Grid.coord_x[1], model.Grid.coord_x[2]),
    ylim=(model.Grid.coord_y[1], model.Grid.coord_y[2]),
    zlim=(-5.0, 0.0),

Origin = nothing,
    StrikeAngle = 0,
    DipAngle = 0,
    phase = ConstantPhase(2),
    T = nothing)
```

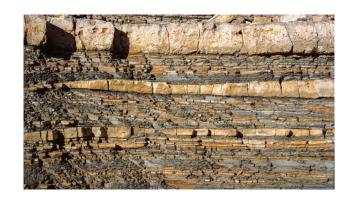
Notice minX = model.Grid.coord\_x[1] and maxX = model.Grid.coord\_x[2]

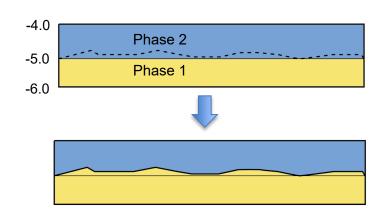
 Note that you need only 4 phases: Air, strong sediment, weak sediment and salt

#### Results



Why do you think there is no fold forming? How to fix that?





- Natural sedimentary layers are irregular
   → we need randomness!
- How to add randomness? Make use of the subarray filter with point wise operation and add randomness for each sedimentary layer interface.

```
Z = model.Grid.Grid.Z;
model.Grid.Phases[Z .> val1 .&& Z .<= val2 ] .= val3;</pre>
```

... and the random function: "rand()" that sends back a random value between 0.0 and 1.0

Note that "rand()" can also be used in a pointwise manner as "rand.()"

lso-viscous folding

0.00 Myrs



Results from one simulation are different, why?

- Now that we have folding, let's add plasticity (brittle deformation)
- Adding brittle deformation is quite straightforward, simply add "ch" and "fr" to the phase() definition

However, plasticity largely increases computational time!

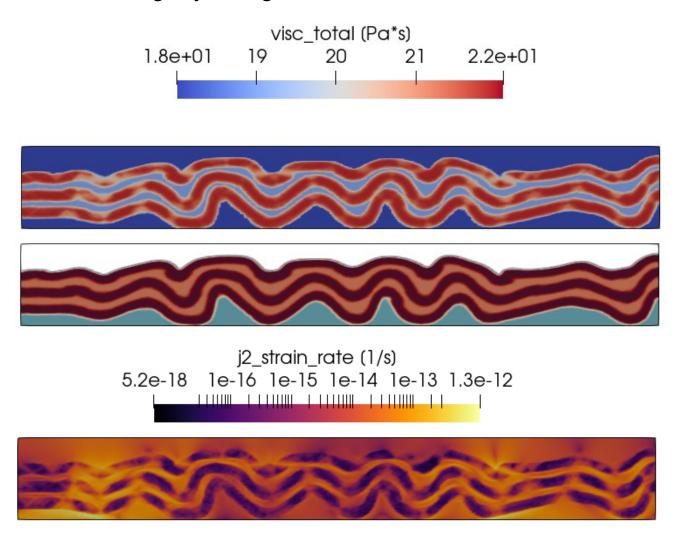
ch = cohesion, fr = frictional angle

Change solver options to the following (speeds up calculation)

```
Solver( SolverType = "direct",
DirectSolver = "mumps",
PETSc_options = [ "-snes_rtol 1e-2", "-snes_max_it 100"] )
```

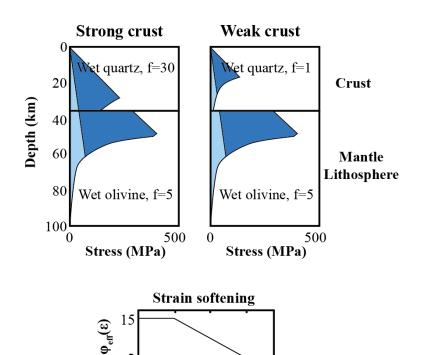
- Perform the simulation
- How much does that change the results? Why?
   (Check viscosity, and strain-rate invariant fields)

Results with Drucker-Prager yielding



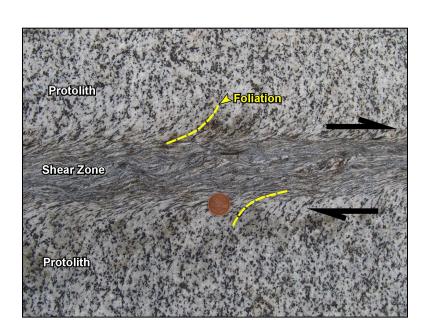
No much localization: fault/shear zone are well defined!

Strain softening



0.5

0.1



 Strain softening: as the rock is plastically deformed it becomes increasingly weaker which allow to localize deformation

- Create a new model setup by copying the previous isoviscous folding one
- Add a Softening() law

Several Softening laws can be added for different phases.

APS = accumulated plastic strain  $\rightarrow$  0.2 means 20% of deformation

APS1 is the value at which cohesion and frictional angle start to be reduced until APS2 is reached.

The reduction is defined by A  $\rightarrow$  0.99 means 99% of strength reduction with respect to the starting brittle strength of the phase.

Then the strain softening law can be added to the "Phase()" such as

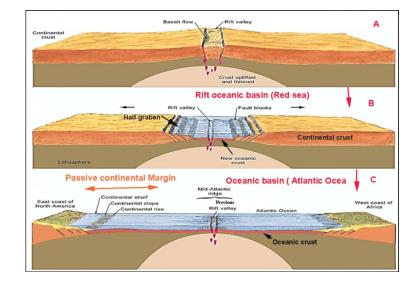
Then don't forget to add the softening law to the model:

```
add_phase!( model, air, salt, sediments1,sediments2 )
add_softening!( model, softening)
```

Perform the simulation

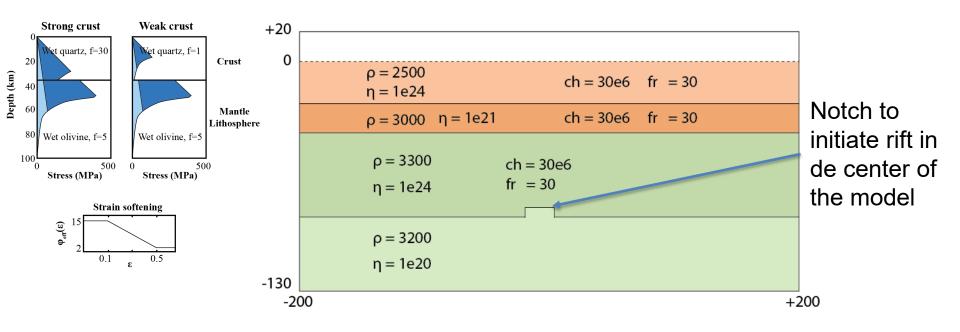
Example of modeling results with softening

0.00 Myrs



How do plates break apart?

### Proposed rifting setup



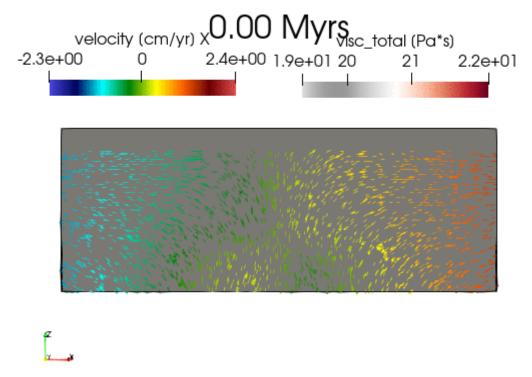
- Notice that to approximate warm asthenosphere we use low viscosity lower density.
- Create a simplified rifting setup using iso-viscous rheologies
- Start from the "fold and thrust" setup

```
\rho = 2500 \qquad ch = 30e6 \quad fr = 30
\rho = 3000 \quad \eta = 1e21 \qquad ch = 30e6 \quad fr = 30
\rho = 3300 \qquad ch = 30e6
\eta = 1e24 \qquad fr = 30
\rho = 3200
\eta = 1e20
-130
```

In Model() add a new section, to add random noise to the particles

Perform the simulation for 300 timesteps

Example of simulation result



- Exercise: update to a thermo-mechanical rifting model
  - Add a temperature profile (using similar approach as what we did for the plume, with adiabatic gradient for the asthenosphere)
  - Use viscous creep rheology (Dry Olivine for the mantle phases, Dry plagioclase for LC and wet quartz for the upper crust)

- Exercise 1: update to a thermo-mechanical rifting model
  - Add a temperature profile (using similar approach as what we did for the plume, with adiabatic gradient for the asthenosphere)
  - Use viscous creep rheology (Dry Olivine for the mantle phases, Dry plagioclase for LC and wet quartz for the upper crust)
- Exercise 2: Add density diagrams