Downscaling of global solar irradiation in complex areas in R

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Abstract

A methodology for downscaling solar irradiation from satellite-derived databases is described using R software. Different packages such as raster, parallel, solaR, gstat, sp and rasterVis are considered in this study for improving solar resource estimation in areas with complex topography, in which downscaling is a very useful tool for reducing inherent deviations in satellite-derived irradiation databases, which lack of high global spatial resolution. A topographical analysis of horizon blocking and sky-view is developed with a digital elevation model to determine what fraction of hourly solar irradiation reaches the Earth's surface. Eventually, kriging with external drift is applied for a better estimation of solar irradiation throughout the region analyzed including the use of local measurements. This methodology has been implemented as an example within the region of La Rioja in northern Spain. The mean absolute error found using the methodology proposed is $91.92 \ kWh/m^2$, vs. $172.62 \ kWh/m^2$ using the original satellite-derived database (a striking 46.75% lower). The code is freely available without restrictions for future replications or variations of the study at https://github.com/EDMANSolar/downscaling.

Keywords: Solar irradiation, R, raster, solaR, digital elevation model, shade analysis, downscaling.

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1 Introduction

During the last few years the development of photovoltaic energy has flourished in developing countries with both multi-megawatt power plants and micro installations. However, the scarcity of long-term, reliable solar irradiation data from pyranometers in many of these countries makes it necessary to estimate solar irradiation from other meteorological variables or satellite images [Schulz et al., 2009; Polo et al., 2014; Vindel et al., 2013]. In such cases, meteorological or satellite derived models need to be validated via nearby pyranometer records, since they lack spatial generalization. Thus, in some regions in which there are no pyranometers nearby these models are ruled out as an option and irradiation data must be obtained from satellite estimates. Although satellite-derived irradiation databases such as NASA's Surface meteorology and Solar Energy (SSE)¹, the National Renewable Energy Laboratory (NREL)², INPE³, SODA⁴ and the Climate Monitoring Satellite Application Facility (CM SAF)⁵ provide wide spatial coverage, only NASA and some CM SAF climate data sets give global coverage, albeit at a reduced spatial resolution (Table 1).

Satellite estimates tend to average solar irradiation and omit the impact of topography within each cell, which generally are in the range of kilometers. As a result, intra-cell variations can be significant in areas with local micro-climatic characteristics and in areas with complex topography (which are often one and the same) [Bosch et al., 2010]. In this case, the irradiation data might not be accurate enough to enable a photovoltaic installation to be designed. [Perez et al., 1994] analyze the spatial behavior of solar irradiation and conclude that the threshold distance from satellite estimates is in the order of 7 km. [Antonanzas-Torres et al., 2013a] reject ordinary kriging as a spatial interpolation method for solar irradiation in Spain with stations more than 50 km apart in mountainous regions, as a result of the high spatial variability of solar radiation in such areas. The NASA-SSE and CM SAF SIS Climate Data Sets provide global horizontal irradiation (GHI) with global coverage with resolutions of 1x1° and 0.25x0.25° (Table 1), which in most latitudes implies a grosser resolution than the previously mentioned 40-50 km.

The influence of terrain has not been widely implemented in satellite-derived solar irradiation databases although it is well known that complex topography attenuates solar irradiation due to shadowing, sky-view (portion of visible sky) and ground reflectance [Ruiz-Arias et al., 2010]. The incoming solar irradiation not blocked by terrain also varies at different solar elevations due to the atmospheric attenuation by aerosols and water vapor (being the influence higher at lower elevations). As a result, complex topography affects micro-climate and makes solar irradiation estimation more complex than in flat areas. For this reason, the analysis of solar irradiation with high spatial resolution is very interesting in areas with complex terrain.

¹http://maps.nrel.gov/SWERA

²http://www.nrel.gov/gis/solar.html

³http://www.inpe.br

⁴http://www.soda-is.com/eng/index.html

⁵http://www.cmsaf.eu

One of the alternatives for obtaining higher spatial resolution of solar irradiation is the downscaling of satellite estimates. Irradiation downscaling can be based on interpolation kriging techniques when pyranometer records are available, with the implementation of continuous irradiation-related variables such as elevation, sky-view-factor [Alsamamra et al., 2009; Batlles et al., 2008] and other meteorological variables (i.e. temperature gradient and rainfall) as external drifts [Antonanzas-Torres et al., 2013b]. Downscaling is generally based on digital elevation models (DEM) with satellite-derived irradiation data to account for the effect of complex topography. It has previously been applied in mountainous areas such as the Mont Blanc Massif (France) [Corripio, 2003] and Sierra Nevada (Spain) [Bosch et al., 2010; Ruiz-Arias et al., 2010] with image resolutions of 3.5x3.5 km. Influence of topography for shade analysis has also been implemented in geographical information systems for solar irradiation modeling, such as r.sun [Suri and Hofierka, 2004] with simplified atmospheric parametrizations, which limit accuracy [Ruiz-Arias et al., 2009].

However, the NASA-SSE and CM SAF SIS Climate Data Sets are based on much lower resolutions and are the only irradiation datasets in numerous countries where there has been recent interest in solar energy. In this paper, a downscaling methodology of global solar irradiation is explained by means of R software and studied in the region of La Rioja (a mountainous region in northern Spain).

In a first step, hourly data from the CM SAF with 0.03x0.03° resolution is down-scaled to a higher resolution (200x200 m). In a second step, *kriging with external drift*, also referred to as *universal kriging*, is applied to interpolate data from 6 on-ground pyranometers in the region, and this downscaled CM SAF data is considered as an explanatory variable. The evaluation of the proposed method is performed with on-site measured GHI. Finally, a downscaled map of annual global solar radiation throughout this region is obtained.

2 Data

The CM SAF was funded in 1992 as a joint venture of several European meteorological institutes, with the collaboration of the European Organization for the Exploitation of Meteorological Satellites (EUMETSAT) to retrieve, archive and distribute climate data to be used for climate monitoring and climate analysis [Trentmann et al., 2011; Schulz et al., 2009]. CM-SAF provides satellite-based operational products and climate data (using the nomenclature by the CM SAF). Operational products are built on data validated with on-ground stations and provided in near-to-present time and climate data are long-term series for evaluating inter-annual variability. This study is built on hourly surface incoming solar radiation and direct irradiation climate data, denoted as SIS and SID by CM SAF respectively, for the year 2005. These climate data are derived from Meteosat first generation satellites (Meteosat 2 to 7, 1982-2005) and validated using on-ground records from the Baseline Surface Radiation Network (BSRN) as a reference. The validation threshold monthly mean absolute bias of SIS and SID is 15 and 20 W/m^2 [Posselt et al., 2012], providing a maximum spatial resolution of $0.03 \times 0.03^\circ$. The high

reflectivity of bright surfaces (i.e. snow covered areas, deserts, and salty areas) leads to a higher uncertainty in the calculation of SIS. The CM SAF uses an algorithm based on fuzzy logic, which specifically improves the estimation in these bright areas [Trentmann et al., 2011]. In the study, SIS and SID data are selected with spatial resolution of $0.03 \times 0.03^{\circ}$. Data is freely accessible via FTP through the CM SAF website. Hourly GHI records from SOS Rioja⁶, taken from 6 meteorological stations (shown in Figure 1 and Table 2) in 2005 serve as complementary measurements for downscaling within the region studied. These stations have First Class pyranometers (according to ISO 9060) with uncertainty levels of 5% in daily totals. These data are filtered from spurious (not available data and values with atmospheric transmittivity- relation between GHI and extra-terrestrial solar irradiation higher than 0.85), assuming when relevant the average between the previous and following hourly measurements. The digital elevation model (DEM) is also freely obtained from product MDT-200 by the ©Spanish Institute of Geography⁷ with a spatial resolution of 200x200 m.

Four meteorological stations, 1, 2, 4 and 6, (corresponding to Ezcaray, Logroño, San Roman and Yerga) are used in the downscaling and data from stations 3 and 5 (Moncalvillo and Ventrosa) are used for testing the downscaling proposed.

3 Method

This section has been divided into three subsections based on the irradiation decomposition, sky view factor and horizon blocking and finally, the post-processing with kriging with external drift.

3.1 Irradiation decomposition

Initially, diffuse horizontal irradiation (DHI) is obtained from the difference between global horizontal irradiation (GHI) and beam horizontal irradiation (BHI) rasters, previously obtained from CM SAF. DHI and BHI are firstly disaggregated from the original gross resolution ($0.03 \times 0.03^{\circ}$) into the DEM resolution (200×200 m) using the bilinear method. In a second step, DHI is divided in two components: isotropic diffuse irradiation (DHI_{iso}), and anisotropic diffuse irradiation (DHI_{ani}) as per the model by Hay & Mckay [Hay and Mckay, 1985] (Equation 1). This model is based on the anisotropy index (k_1), defined as the ratio of the beam irradiance (B(0)) to the extra-terrestrial irradiance (B(0)), as shown in Equation 2. High k_1 values are typical in clear sky atmospheres, while low k_1 values are frequent in overcast atmospheres and those with a high aerosol density.

$$DHI = DHI_{iso} + DHI_{ani} \tag{1}$$

^{6/}http://www.larioja.org/npRioja/default/defaultpage.jsp?idtab=442821

⁷http://www.ign.es

$$k_1 = \frac{B(0)}{B_0(0)} \tag{2}$$

The DHI_{iso} accounts for the incoming diffuse irradiation portion from an isotropic sky, being higher with higher cloudiness (Equation 3).

$$DHI_{iso} = DHI \cdot (1 - k_1) \tag{3}$$

 DHI_{ani} , also denoted as circumsolar diffuse irradiation, considers the incoming portion from the circumsolar disk, which is area surrounding the solar disk. It can be analyzed as beam irradiation [Perpiñán-Lamigueiro, 2013] (Equation 4).

$$DHI_{ani} = DHI \cdot k_1 \tag{4}$$

3.2 Sky view factor and horizon blocking

Topographical analysis is performed accounting for the visible sky sphere (sky view) and horizon blocking. The DHI_{iso} is directly dependent on the sky-view factor (SVF), which computes the proportion of visible sky related to a flat horizon. The sky-view factor is considered in earlier irradiation assessments [Ruiz-Arias et al., 2010; Corripio, 2003]. It is calculated in each DEM pixel by considering 72 vectors (separated by 5° each) and evaluating the maximum horizon angle (θ_{hor}) over 20 km in each vector (Equation 5). The θ_{hor} stands for the maximum angle between the altitude of a location and the elevation of the group of points along each vector, related to a horizontal plane on the location. Locations without horizon blocking have SVFs close to 1, which means a whole visible semi-sphere of sky.

$$SVF = 1 - \int_0^{2\pi} \sin^2 \theta_{hor} d\theta \tag{5}$$

Eventually, the downscaled DHI_{iso} ($DHI_{iso,down}$) is computed with Equation 6.

$$DHI_{iso,down} = DHI_{iso} \cdot SVF \tag{6}$$

Horizon blocking is analyzed by evaluating the solar geometry in 15 minute samples, particularly the solar elevation (γ_s) and the solar azimuth (ψ_s). Secondly, the mean hourly γ_s and ψ_s (from those 15 minute rasters) are calculated and then disaggregated as explained above for DHI and BHI rasters. The θ_{hor} corresponding to each ψ_s is compared with the γ_s . As a consequence, if the γ_s is lower than the θ_{hor} , then there is horizon blocking on the surface analyzed and therefore, BHI and DHI_{ani} are blocked. Finally, the sum of $DHI_{ani,down}$, $DHI_{iso,down}$ and $BHI_{iso,down}$ constitutes the downscaled global horizontal irradiation GHI_{down} .

3.3 Post-processing: kriging with external drift

The fact that this downscaling accounts for the irradiation loss due to horizon blocking and the sky-view factor leads us to introduce a trend in estimates (lowering them) compared to the original data (gross resolution data). However, satellite-derived irradiation data implicitly considers shade, as a consequence of the lower albedo recorded in these zones, but it is later averaged over the pixel. GHI_{down} is used as the raster layer with the shading behavior on solar irradiation within the region studied. *Universal kriging* or *kriging with external drift* (KED) includes information from exhaustively-sampled explanatory variables in the interpolation. As a result, GHI_{down} is considered as the explanatory variable for interpolating measured irradiation data from on-ground calibrated pyranometers, which is denoted as *post-processing*. GHI_{down} is correlated with the DEM as a consequence of the major influence of horizon blocking with topography, estimations can be derived by separating the deterministic and stochastic components (Equation 7).

$$\hat{z}(\mathbf{s}_{\theta}) = \underbrace{\sum_{k=0}^{p} \hat{\beta}_{k} q_{k}(\mathbf{s}_{\theta})}_{\text{deterministic}} + \underbrace{\sum_{i=1}^{n} \lambda_{i} \varepsilon(\mathbf{s}_{i})}_{\text{stochastic}}$$
(7)

where $\hat{z}(\mathbf{s}_{\theta})$ is the estimated value in \mathbf{s}_{θ} , $\hat{\beta}_k$ are the estimated coefficients of the deterministic model, $q_k(\mathbf{s}_{\theta})$ are the auxiliary predictors obtained from the fitted values of the explanatory variable at the new location, λ_i are the kriging weights determined by the spatial dependence structure of the residual, and $\epsilon(\mathbf{s}_i)$ are the residual at location \mathbf{s}_i [Antonanzas-Torres et al., 2013a].

The semivariogram is a function defined as Equation 8 based on a constant variance of ϵ and also on the assumption that spatial correlation of ϵ depends on the distance amongst instances (**h**) rather than on their position [Pebesma, 2004].

$$\gamma(\mathbf{h}) = \frac{1}{2} \mathbf{E}[\epsilon(\mathbf{s}) - \epsilon(\mathbf{s} + \mathbf{h})]^2$$
 (8)

Given that the above sample variogram only collates estimates from observed points, a fitting model of this variogram is generally considered to extrapolate the spatial behavior of observed points to the area studied. Different variogram functions are commonly defined such as the exponential, Gaussian or spherical models. Along these equations, different parameters such as the sill, range and nugget of the model must be adjusted to best fit the sample variogram [Hengl, 2009]. The nugget effect, generally associated with intrinsic micro-variability and measurement error, models the discontinuity of the variogram at the source. It must be highlighted that when the nugget effect is recorded, estimates are different from measured values in the stations. The variogram model of solar horizontal irradiation is evaluated in Spain, and the conclusion reached is that a pure nugget fitting behaves best, which implies no spatial auto-correlation on residuals [Antonanzas-Torres et al., 2013a].

Figure 2 displays the method diagram using red ellipses and lines for data sources, blue ellipses and lines for derived rasters (results), and black rectangles and lines for operations.

The mean absolute error (MAE), root mean square error (RMSE) and mean bias error (MBE), described in Equations 9, 10 and 11, are used as indicators of models deviation.

$$MAE = \frac{\sum_{i=1}^{n} |x_{est} - x_{meas}|}{n} \tag{9}$$

$$RMSE = \sqrt{\frac{\sum_{i=1}^{n} (x_{est} - x_{meas})^{2}}{n}}$$
 (10)

$$MBE = \frac{\sum_{i=1}^{n} x_{est} - x_{meas}}{n} \tag{11}$$

where n is number of stations and x_{est} and x_{meas} the annual estimated and measured irradiation, respectively.

4 Implementation

The method proposed is applied in the region of La Rioja (northern Spain). Figure 3 shows the corresponding annual global horizontal irradiation from CM SAF with resolution $0.03 \times 0.03^{\circ}$.

4.1 Packages

The downscaling described in this paper has been implemented using the free software environment R [R Development Core Team, 2013] and various contributed packages:

- raster [Hijmans and van Etten, 2013] for spatial data manipulation and analysis.
- solaR [Perpiñán-Lamigueiro, 2012] for solar geometry.
- gstat [Pebesma and Graeler, 2013] and sp [Pebesma et al., 2013] for geostatistical analysis.
- parallel for multi-core parallelization.
- rasterVis [Perpiñán-Lamigueiro and Hijmans, 2013] for spatial data visualization methods.

```
R> library(sp)
```

R> library(raster)

R> rasterOptions(todisk=FALSE)

R> rasterOptions(chunksize = 1e+06, maxmemory = 1e+07)

R> library(maptools)

R> library(gstat)

```
R> library(lattice)
R> library(latticeExtra)
R> library(rasterVis)
R> library(solaR)
R> library(parallel)
```

4.2 Data

Satellite data can be freely downloaded after registration from CM SAF⁸ by going to the data access area, selecting *web user interface* and *climate data sets* and then choosing the hourly climate data sets named *SIS* (Global Horizontal Irradiation)) and *SID* (Beam Horizontal Irradiation) for 2005. Both rasters are projected to the UTM projection for compatibility with the DEM.

```
R> projUTM <- CRS('+proj=utm_+zone=30')
R> projLonLat <- CRS('_+proj=longlat_+ellps=WGS84')
R> listFich <- dir(pattern='SIShm2005')
R> stackSIS <- stack(listFich)
R> stackSIS <- projectRaster(stackSIS,crs=projUTM)
R> listFich <- dir(pattern='SIDhm2005')
R> stackSID <- stack(listFich)
R> stackSID <- projectRaster(stackSID, crs=projUTM)</pre>
```

We compute the annual global irradiation, which will be used as a reference for subsequent steps.

```
R> SISa2005 <- calc(stackSIS, sum, na.rm=TRUE)</pre>
```

The Spanish Digital Elevation Model can be obtained after registration from the ©Spanish Institute of Geography⁹ by going to the *free download of digital geographic information for non-commercial use* area, and then cropping to the region analyzed (La Rioja). As stated above, this DEM uses the UTM projection.

```
R> elevSpain <- raster('elevSpain.grd')
R> elev <- crop(elevSpain, extent(479600, 616200, 4639600, 4728400))
R> names(elev)<-'elev'</pre>
```

4.3 Sun geometry

The first step is to compute the sun angles (height and azimuth) and the extraterrestrial solar irradiation for each cell of the CM SAF rasters. The function calcSol from the solaR package calculates the daily and intradaily sun geometry. By means of this function and overlay from the raster package, three multilayer raster objects are

 $^{^8}$ www.cmsaf.eu

⁹http://www.ign.es

generated with the sun geometry needed for the next steps. For the sake of brevity we show only the procedure for extraterrestrial solar irradiation. The sun geometry is calculated with the spatial resolution of CM SAF (0.03x0.03°). First, it is defined a function to extract the hour for aggregation, choose the annual irradiation raster as reference, and define a raster with longitude and latitude coordinates.

```
R> hour <- function(tt)as.POSIXct(trunc(tt, 'hours'))
R> r <- SISa2005
R> latlon <- stack(init(r, v='y'), init(r, v='x'))
R> names(latlon) <- c('lat', 'lon')</pre>
```

The extraterrestrial irradiation is calculated with 5-min samples. Each point is a column of the data frame locs. Its columns are traversed with lapply, so for each point of the raster object a time series of extraterrestrial solar irradiation is computed. The result, B05min, is a RasterBrick object with a layer for each element of the time index BTi, which is aggregated to an hourly raster with zApply and transformed to the UTM projection.

```
R> BTi <- seq(as.POSIXct('2005-01-01_00:00:00'),
+ as.POSIXct('2005-12-31_23:55:00'), by='5_min')

R> B05min <- overlay(latlon, fun=function(lat, lon){
+ locs <- as.data.frame(rbind(lat, lon))
+ b <- lapply(locs, function(p){
+
+ hh <- local2Solar(BTi, p[2])
+ sol <- calcSol(p[1], BTi=hh)
+ Bo0 <- as.data.frameI(sol)$Bo0
+ Bo0 })
+ res <- do.call(rbind, b)})

R> B05min <- setZ(B05min, BTi)
R> names(B05min) <- as.character(BTi)</pre>
R> B0h <- zApply(B05min, by=hour, fun=mean)
R> projectRaster(B0h,crs=projUTM)
```

4.4 Irradiation components

The CM SAF rasters must be transformed to the higher resolution of the DEM (UTM 200x200 m). Because of the differences in pixel geometry between DEM (square) and irradiation rasters (rectangle) the process is performed in two steps.

The first step increases the spatial resolution of the irradiation rasters to a larger pixel size than the DEM with disaggregated data, where sf is the scale factor. The second step post-processes the previous step by means of a bilinear interpolation which

resamples the raster layer and achieves the same DEM resolution (resample). This twostep disaggregation prevents the loss of the original values of the gross resolution raster that would be directly interpolated with the one-step disaggregation.

```
R> sf <- res(stackSID)/res(elev)
R> SIDd <- disaggregate(stackSID, sf)
R> SIDdr <- resample(SIDd, elev)
R> SISd <- disaggregate(stackSIS, sf)
R> SISdr <- resample(SISd, elev)</pre>
```

The diffuse irradiation is obtained from the global and beam irradiation rasters. The two components of the diffuse irradiation, isotropic and anisotropic, can be separated with the anisotropy index, computed as the ratio between beam and extraterrestrial irradiation.

```
R> Difdr <- SISdr - SIDdr
R> B0hd <- disaggregate(B0h, sf)
R> B0hdr <- resample(B0hd, elev)
R> k1 <- SIDdr/B0hdr
R> Difiso <- (1-k1) * Difdr
R> Difani <- k1 * Difdr</pre>
```

4.5 Sky view factor and horizon blocking

4.5.1 Horizon angle

The maximum horizon angle required for the horizon blocking analysis and to derive the SVF is obtained with the next code. The alpha vector is visited with mclapply (using parallel computing). For each direction angle (elements of this vector) the maximum horizon angle is calculated for a set of points across that direction from each of the locations defined in xyelev (derived from the DEM raster and transformed in the matrix locs visited by rows).

```
R> xyelev <- stack(init(elev, v='x'),
+ init(elev, v='y'),
+ elev)
R> names(xyelev) <- c('x', 'y', 'elev')

R> inc <- pi/36
R> alfa <- seq(-0.5*pi,(1.5*pi-inc), inc)
R> locs <- as.matrix(xyelev)</pre>
```

Separations between the source locations and points along each direction are defined by resD, the maximum resolution of the DEM, d, maximum distance to visit, and consequently in the vector seps.

```
R> resD <- max(res(elev))

R> d <- 20000
R> seps <- seq(resD, d, by=resD)</pre>
```

The elevation (z1) of each point in xyelev is converted into the horizon angle: the largest of these angles is the horizon angle for that direction. The result of each apply step is a matrix, which is used to fill in a RasterLayer (r). The result of mclapply is a list, hor, of RasterLayer which can be converted into a RasterStack with stack. Each layer of this RasterStack corresponds to a different direction.

```
R> hor <- mclapply(alfa, function(ang){
+ h <- apply(locs, 1, function(p){
+ x1 <- p[1]+cos(ang)*seps
+ y1 <- p[2]+sin(ang)*seps
+ p1 <- cbind(x1,y1)
+ z1 <- elevSpain[cellFromXY(elevSpain,p1)]
+ hor <- r2d(atan2(z1-p[3], seps))
+ maxHor <- max(hor[which.max(hor)], 0)
+ })
+ r <- raster(elev)
+ r[] <- matrix(h, nrow=nrow(r), byrow=TRUE)
+ r}, mc.cores=8)
R> horizon <- stack(hor)</pre>
```

This operation is very time-consuming as it is necessary to work with high resolution files. Computation time can be decreased by increasing the sampling space (200 m) or the sectoral angle (5 $^{\circ}$) or by reducing the maximum distance (20 km).

4.5.2 Horizon blocking

Horizon blocking is analyzed by evaluating the solar geometry in 15 minute samples, particularly the solar elevation and azimuth angles from the original irradiation raster. Secondly, the hourly averages are calculated, disaggregated and post-processed as explained above for the irradiation rasters. The decision to solve the solar geometry with low resolution rasters enables a significant reduction to be obtained in computation time without penalizing the results.

First, the azimuth raster is cut into different classes according to the alpha vector (directions). The values of the horizon raster corresponding to each angle class are extracted using stackSelect.

```
R> idxAngle <- cut(AzShr, breaks=r2d(alfa))
R> AngAlt <- stackSelect(horizon, idxAngle)</pre>
```

The number of layers of AngAlt is the same as idxAngle and can therefore be used for comparison with the solar height angle, AlShr. If AngAlt is greater, there is horizon blocking (dilogical=0).

```
R> dilogical <- ((AngAlt-AlShr) < 0)
```

With this binary raster, beam irradiation and diffuse anisotropic irradiation can be corrected with horizon blocking.

```
R> Dirh <- SIDdr * dilogical
R> Difani <- Difani * dilogical
```

4.5.3 Sky view factor

The sky-view factor can be easily computed from the horizon object with the equation 5. This factor corrects the isotropic component of the diffuse irradiation.

```
R> SVFRuizArias <- calc(horizon, function(x) sin(d2r(x))^2)
R> SVF <- 1 - mean(SVFRuizArias)
R> Difiso <- Difiso * SVF</pre>
```

Finally, the global irradiation is the sum of the three corrected components, beam and anisotropic diffuse irradiation including horizon blocking, and isotropic diffuse irradiation with the sky view factor.

```
R> GHIh <- Difanis + Difiso + Dirh
R> GHI2005a <- calc(GHIh, fun=sum)
```

4.6 Kriging with external drift

The downscaled irradiation rasters can be improved by using kriging with external drift. Irradiation data from on-ground meteorological stations is interpolated with the downscaled irradiation raster as the explanatory variable. To define the variogram here we use the results previously published in [Antonanzas-Torres et al., 2013a].

```
R> load('Stations.RData')
R> UTM <- SpatialPointsDataFrame(Stations[,c(2,3)], Stations[,-c(2,3)],
+ proj4string=CRS('+proj=utm_+zone=30_+ellps=WGS84'))</pre>
```

The file Stations .RData contains information regarding annual measured and CM SAF GHI (GHI_{med} and GHI_{CMSAF} , respectively). For a better insight of this file it is provided at https://github.com/EDMANSolar/downscaling. The variogram is built using the spatial data of the SpatialPointsDataFrame object and the nugget model.

```
R> vgmCMSAF <- variogram(GHImed ~ GHIcmsaf, UTM)
R> fitvgmCMSAF <- fit.variogram(vgmCMSAF, vgm(model='Nug'))</pre>
```

```
R> gModel <- gstat(NULL, id='GOyKrig',
+ formula= GHImed ~ GHIcmsaf,
+ locations=UTM, model=fitvgmCMSAF)
R> names(GHI2005a) <- 'GHIcmsaf'
R> GOyKrig <- interpolate(GHI2005a, gModel, xyOnly=FALSE)</pre>
```

5 Results

This methodology leads to an improvement in estimates: the MAE is down by 46.75% and the RMSE by 43.38% compared to CM SAF. Table 3 shows the errors obtained with CM SAF and with the methodology proposed. In addition, the downscaling proposed obtains lower MBE (-91.92 kWh/m^2) than the validation threshold by CM SAF (131.4 kWh/m^2) [Posselt et al., 2012]. Nevertheless, this threshold value is surpassed in the case of CM SAF (175.62 kWh/m^2) in testing meteorological stations. Both, CM SAF and downscaling under-estimate GHI proved by negative MBE.

Figure 3 shows the annual GHI as per CM SAF with the gross resolution analyzed (0.03x0.03°). Figure 4 shows the downscaled maps (200x200 m), presenting lower values of GHI in mountainous areas with lower sky-view and a higher effect of shades.

In order to evaluate the behavior of the method proposed, relative differences evaluated with station measurements are shown in Figure 5. In this figure, stations used for the downscaling and for testing are shown. As can be deduced from this Figure, relative differences are lower in the *downscaling* than in CM SAF or GHI_{down} , at \pm 11.2%. Some stations (1 and 4) present higher relative error with the downscaling than with CM SAF and this might be explained due to a lower influence of sky-view and horizon blocking.

The intrapixel variability due to the downscaling procedure is indicative of the importance of the topography as an attenuator of solar irrradiation. As a result, this zonal variability is higher in pixels with complex topographies and downscaling is more useful. Figure 6 shows the relative difference between downscaling with KED and CM SAF. As might be deduced, CM SAF over-estimates GHI in this region by between 11 and 22%. Figures 7 and 8 display the standard deviations of the downscaled maps within each cell of the original CM SAF raster (0.03x0.03°). The zonal function from the raster library permits this calculation, explaining the intrinsic variability of solar radiation within gross resolution pixels. Consequently, in those pixels with higher standard deviations there will be greater variability. Figure 8 shows how the KED method smooths the deviation within pixels and also the range of solar irradiation in the region (Figure 4).

6 Concluding comments

A methodology for downscaling solar irradiation is described and presented using R software. This methodology is useful for increasing the accuracy and spatial resolution

of gross resolution satellite-estimates of solar irradiation.

It has been proved that areas whose topography is complex show greater differences with the original gross resolution data as a consequence of horizon blocking and lower sky-view factors, so downscaling is highly recommended in these areas.

Kriging with external drift with the gstat package has proved very useful in down-scaling solar irradiation when on-ground registers are available and an explanatory variable is provided.

This methodology is implemented as an example in the region of La Rioja in northern Spain, and striking reductions of annual 46.75% and 43.38% in MAE and RMSE are obtained compared to the original gross resolution database. MBE obtained with downscaling are lower the threshold value of CM SAF (131.4 kWh/m^2), proving the high influence of topography on solar irradiation.

The high repeatability of this methodology and the reduction in errors obtained with annual values might be also very useful in the downscaling of hourly and daily values of solar irradiation and also for different meteorological variables (omitting the sky-view and horizon blocking steps). The code is freely available without restrictions for future replications or variations of the study at https://github.com/EDMANSolar/downscaling.

Software information

The results discussed in this paper were obtained in a R session with these characteristics:

- R version 2.15.2 (2012-10-26), x86_64-apple-darwin9.8.0
- Locale: es_ES.UTF-8/es_ES.UTF-8/es_ES.UTF-8/C/es_ES.UTF-8/es_ES.UTF-8
- Base packages: base, datasets, graphics, grDevices, grid, methods, parallel, stats, utils
- Other packages: foreign 0.8-51, gstat 1.0-16, hexbin 1.26.0, lattice 0.20-15, latticeExtra 0.6-19, maptools 0.8-14, raster 2.1-16, rasterVis 0.20-01, RColorBrewer 1.0-5, rgdal 0.8-01, solaR 0.33, sp 1.0-8, zoo 1.7-9
- Loaded via a namespace (and not attached): intervals 0.14.0, spacetime 1.0-4, tools 2.15.2, xts 0.9-3

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Database	Product	Spatial coverage	Spatial resolution	Temporal coverage	Temporal resolution
CM SAF	SIS Climate Data Set (GHI)	Global	0.25x0.25°	1982-2009	Daily means
CM SAF	SIS Climate Data Set (GHI)	70S-70N, 70W- 70E	0.03x0.03°	1983-2005	Hourly means
CM SAF	SID Climate Data Set (BHI)	70S-70N, 70W- 70E	0.03x0.03°	1983-2005	Hourly means
SODA	Helioclim 3 v2 and v3 (GHI)	66S-66N,66W- 66E	5km	2005	15 minutes
SODA	Helioclim 3 v2 and v3 (GHI)	66S-66N,66W- 66E	5km	2005	15 minutes
NREL	GHI Moderate resolution	Central and South America, Africa, India, East Asia	40x40km	1985-1991	Monthly means of daily GHI
NASA	SSE	Global	1x1°	1983-2005	Daily means

Table 1: Summary of satellite-derivd solar irradiation databases

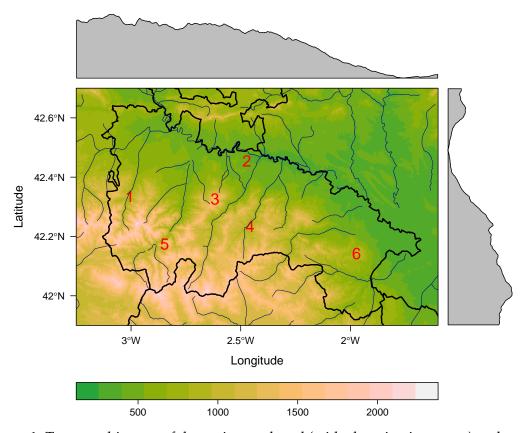


Figure 1: Topographic map of the region analyzed (with elevation in meters) and meteorological stations considered

#	Name	Net.	Lat.(°) Long.(o)Alt.	GHI_a
1	Ezcaray	SOS	42.33 -3.00	1000	1479
2	Logroño	SOS	42.45 -2.74	408	1504
3	Moncalvillo	SOS	42.32 -2.61	1495	1329
4	San Roman	SOS	42.23 -2.45	1094	1504
5	Ventrosa	SOS	42.17 -2.84	1565	1277
6	Yerga	SOS	42.14 -1.97	1235	1448

Table 2: Summary of the meteorological stations selected. GHI_a stands for the annual GHI in kWh/m^2

Figure 2: Methodology of downscaling: this figure uses red ellipses and lines for data sources, blue ellipses and lines for derived rasters (results), and black rectangles and lines for operations. GHI and BHI stand for global and direct horizontal irradiation, respectively. DHI stands for diffuse horizontal irradiation. Subscripts *dis* stand for disaggregated, *iso* for isotropic, *ani* for anisotropic, *down* for downscaled and *ground* for measured values. KED and DEM stand for kriging with external drift and digital elevation model, respectively

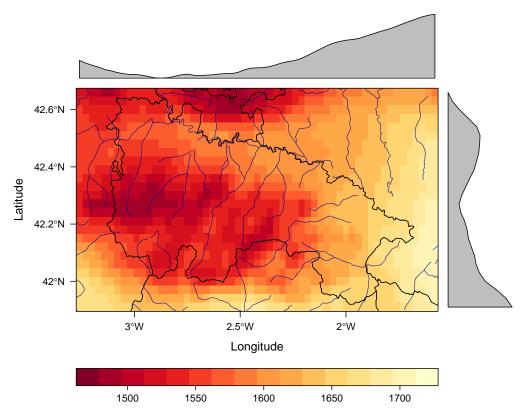


Figure 3: Annual GHI of 2005 from CM SAF estimates (0.03x0.03°) in La Rioja (kWh/m^2)

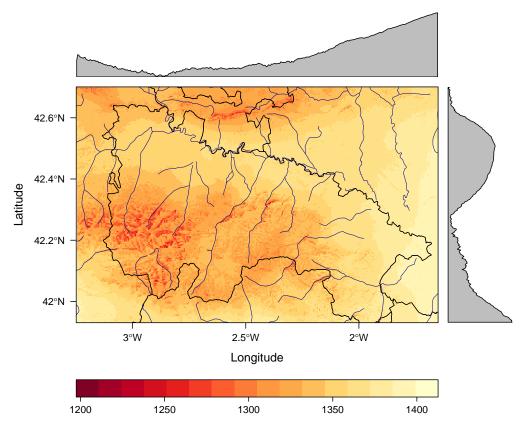


Figure 4: Annual $GHI_{down,ked}$ of 2005 (0.03x0.03°) in La Rioja (kWh/m^2)

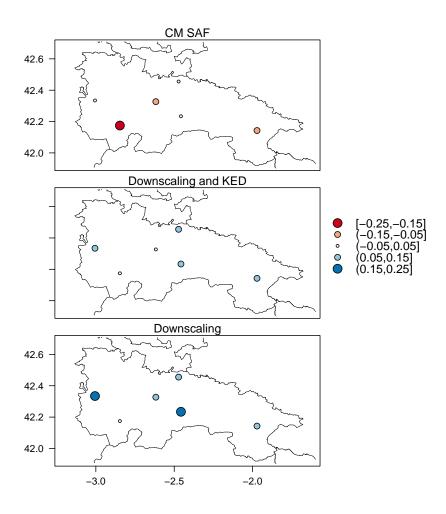


Figure 5: Annual relative differences in per one units evaluated with station measurements (1- $GHI_{estimated}/GHI_{ground}$). The x and y axis represents the latitude and longitude (°), respectively

	GHI_{CMSAF}	GHI _{down,ked}
MAE	172.62	91.92
RMSE	175.02	99.08
MBE	-172.62	-91.92

Table 3: Summary of testing errors obtained in kWh/m^2

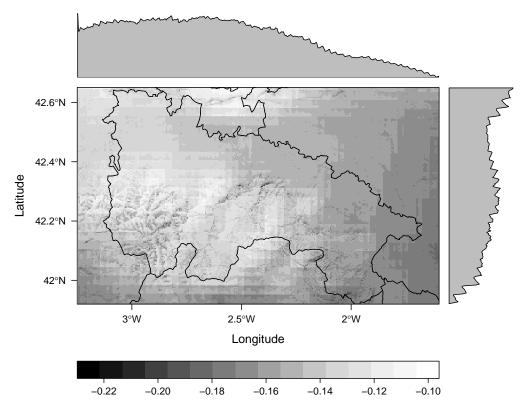


Figure 6: Relative difference of $GHI_{down,ked}$ and GHI_{dis} ($GHI_{down,ked}/GHI_{dis}$ -1) in per one units

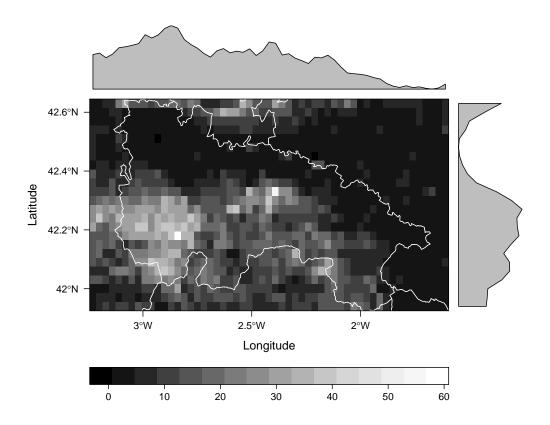


Figure 7: Difference of zonal standard deviations (kWh/m^2) of GHI_{down} and $GHI_{down,KED}$

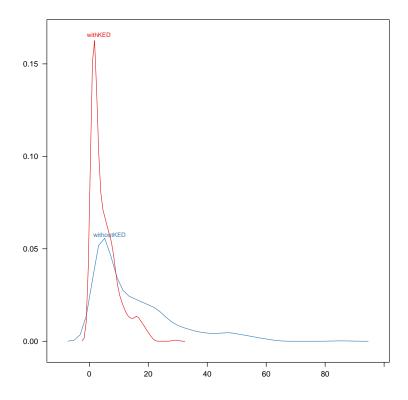


Figure 8: Density plot of zonal standard deviations between CM SAF and downscaling