

Paleoclimate



source: NASA

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Yesterday's Summary



Yesterday's Summary

- Paleoclimatology is very interdisciplinary
- many different archives and proxies, but data patchy and often uncertain
- long term climate determined by:
insolation, albedo, and greenhouse gases
- Early Earth climate has changed completely
- Life and Evolution have shaped Earth's chemistry



Lecture Progress

Today we finish 15 min early!

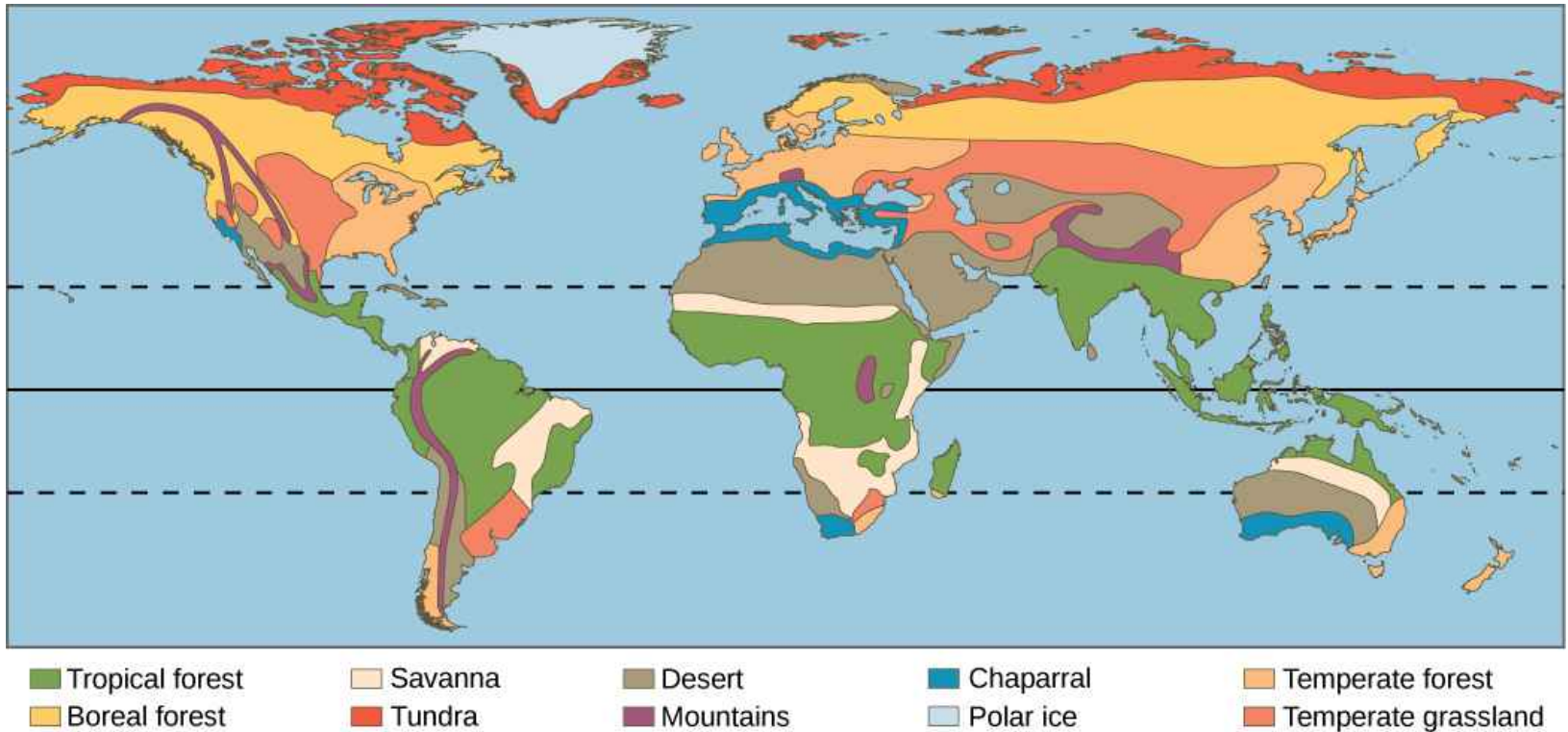
Monday	Introduction	Earth History
Tuesday	Proxies I	Cenozoic Hot & Warm House
Wednesday	Specific Climate System components	Pleistocene G-IG climate
Thursday	Proxies II & Climate System Interactions	Abrupt Climate Change
Friday	Current Climate Change	Future & Synthesis





modern biomes

lumenlearning.com
Environmental Biology



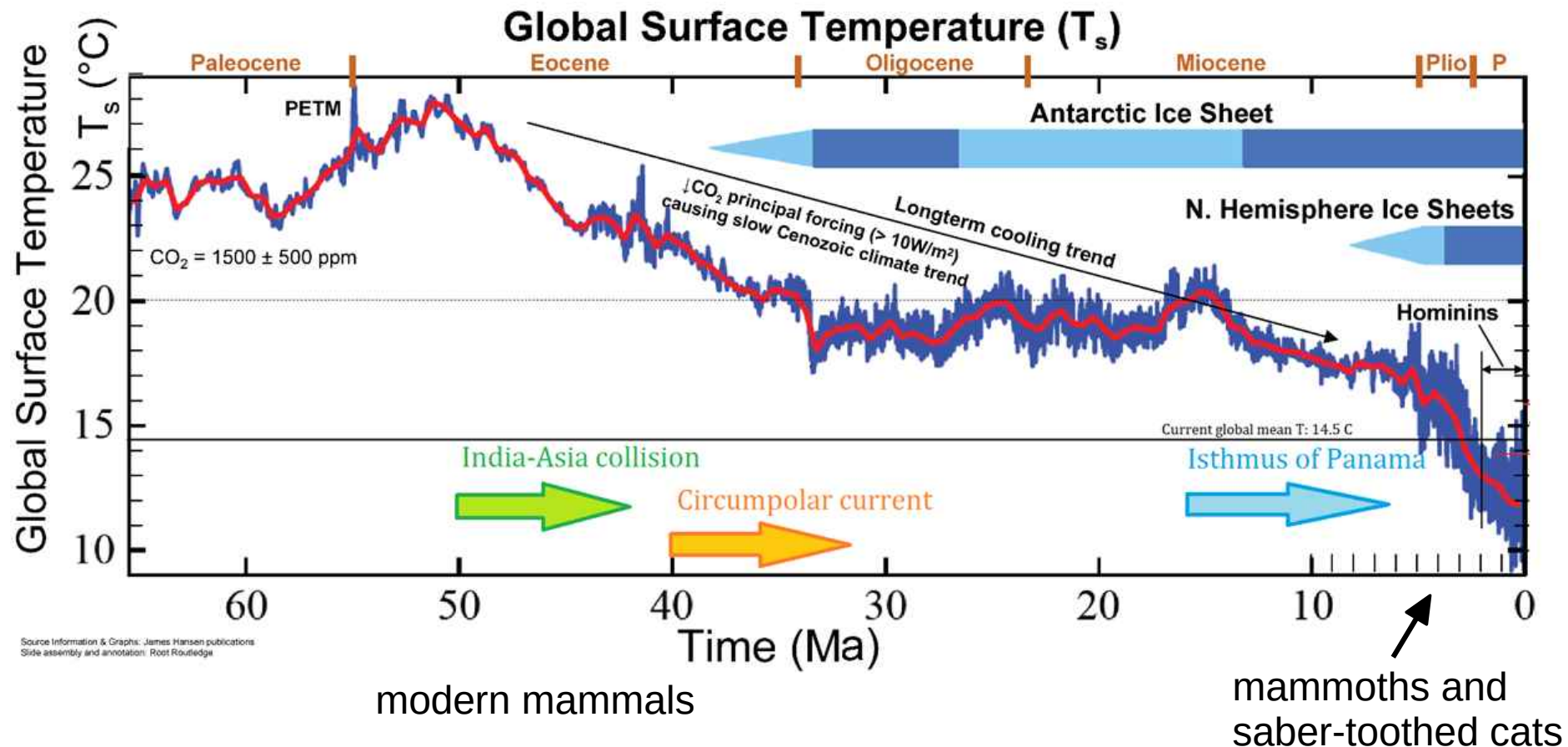
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Day 2 : Overview

- Overview of Cenozoic Climate
 - Marine Isotope Records
- Basics of Isotope Geochemistry
 - Oxygen Isotopes in Paleoclimatology
 - Clumped C-O Isotopes
 - Mg/Ca paleothermometer
 - TEX86 paleothermometer
- Hothouse “Equable Climate”
 - PETM hyperthermal
- Climate Sensitivity
- Mid-Late Cenozoic cooling

Cenozoic Climate



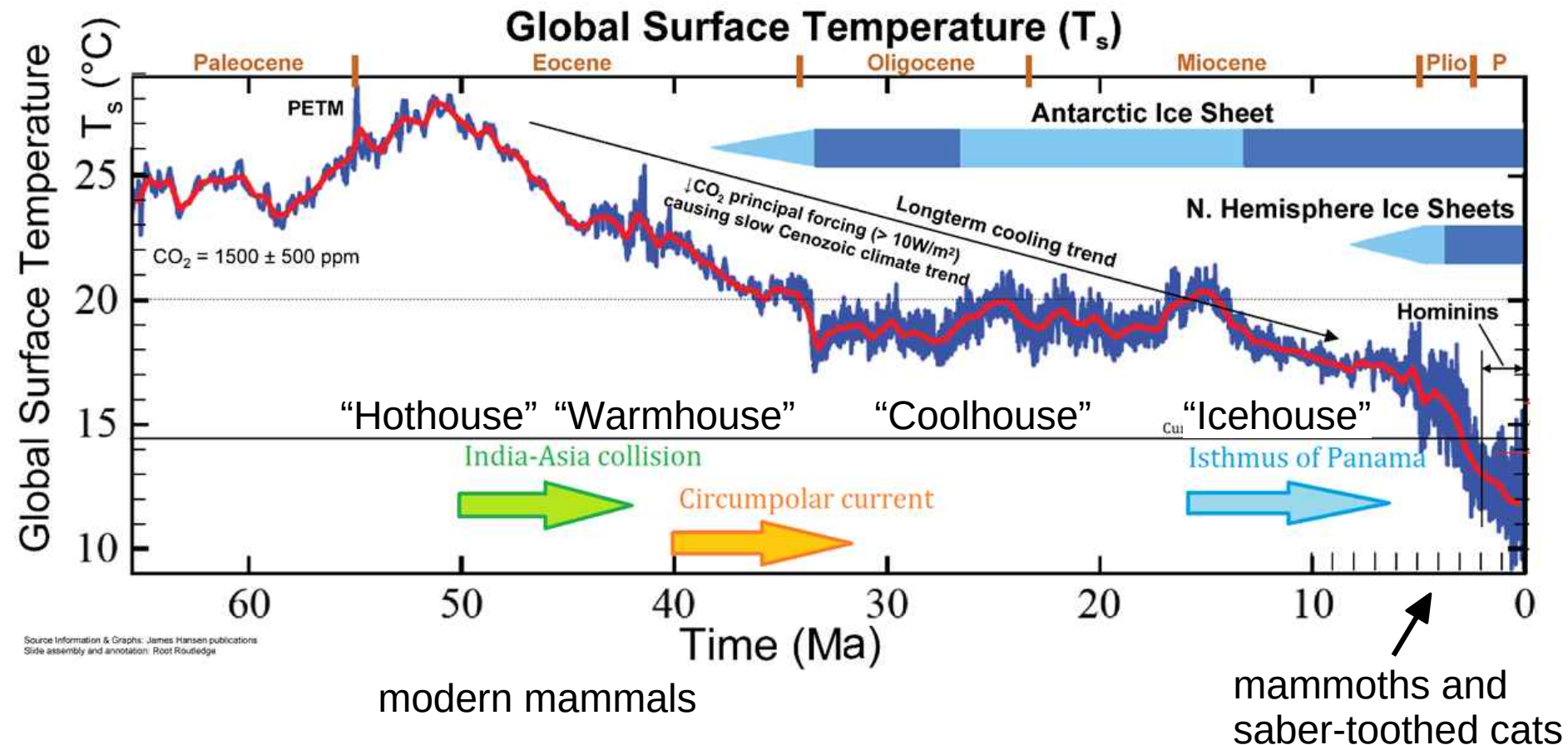
Source Information & Graphs: James Hansen publications
Slide assembly and annotation: Root Routledge

Earle (2016), opentextbc.ca
after James Hansen and Root Routledge

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Cenozoic Climate



modern mammals

mammoths and
saber-toothed cats

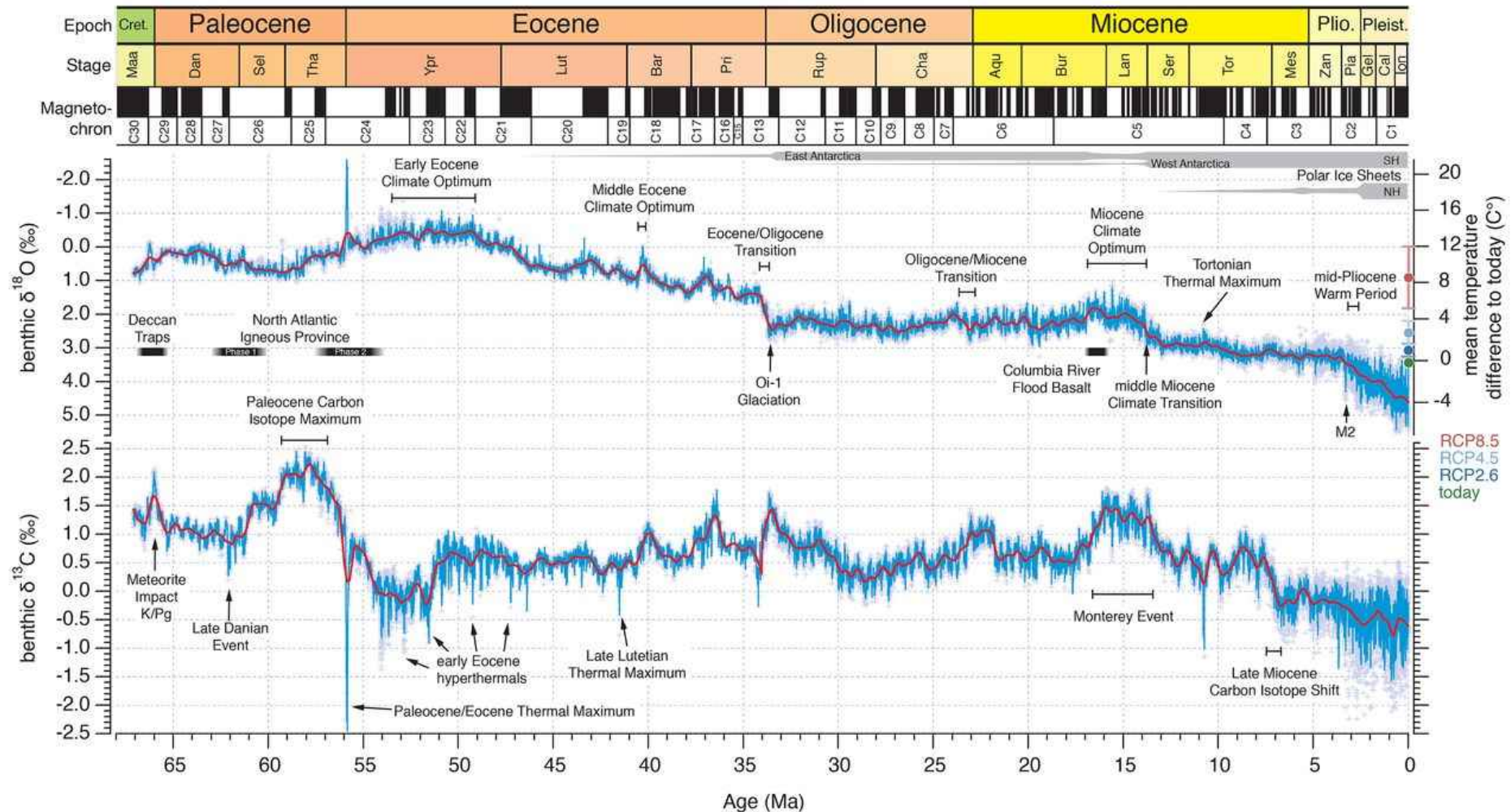
better records with plenty marine sediment cores

Earle (2016), opentextbc.ca
after James Hansen and Root Routledge

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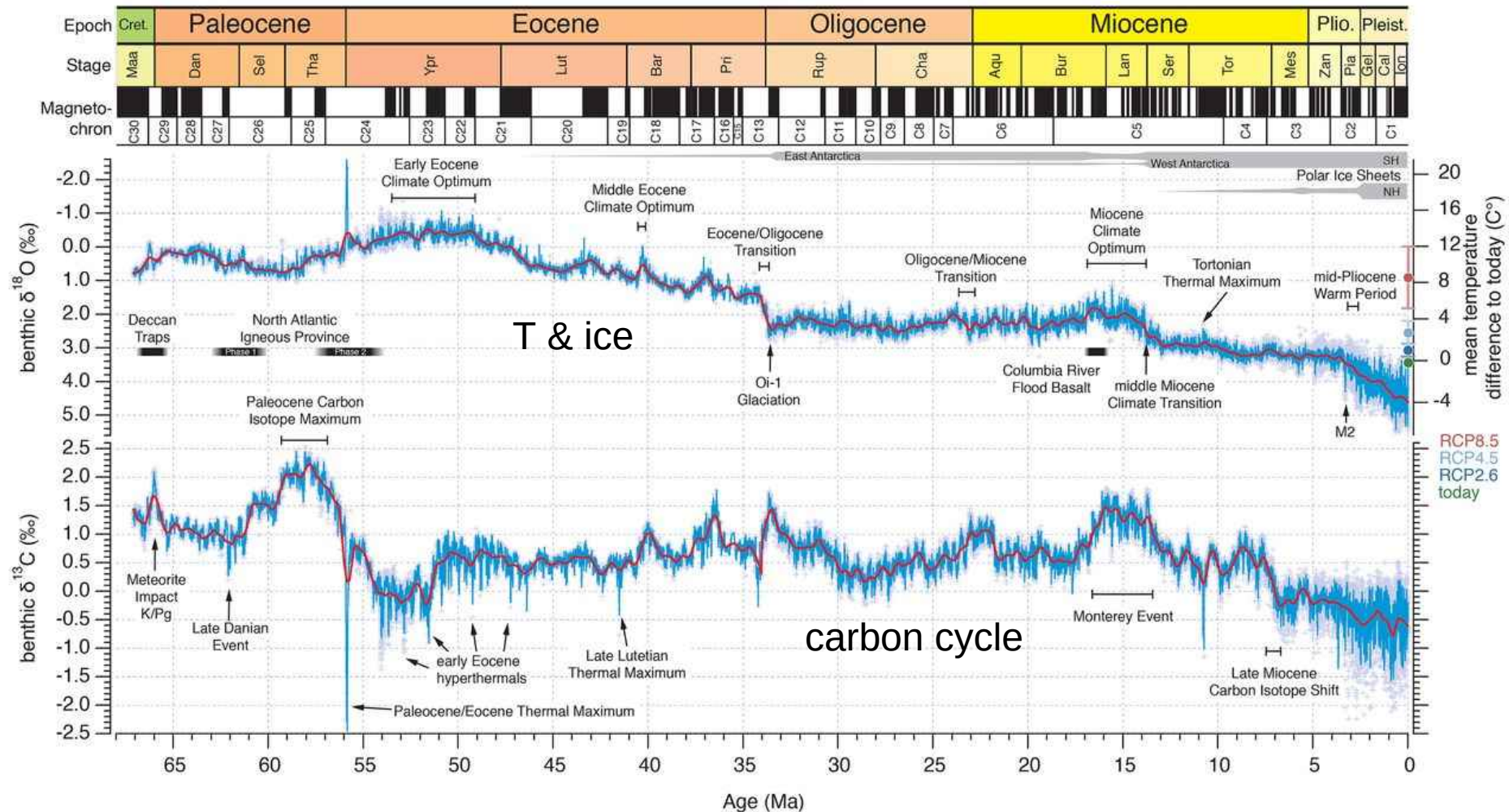
Cenozoic Climate



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Cenozoic Climate



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Isotope Geochemistry

Isotope Geochemistry

Isotope fractionation

^{16}O – 99.76 %

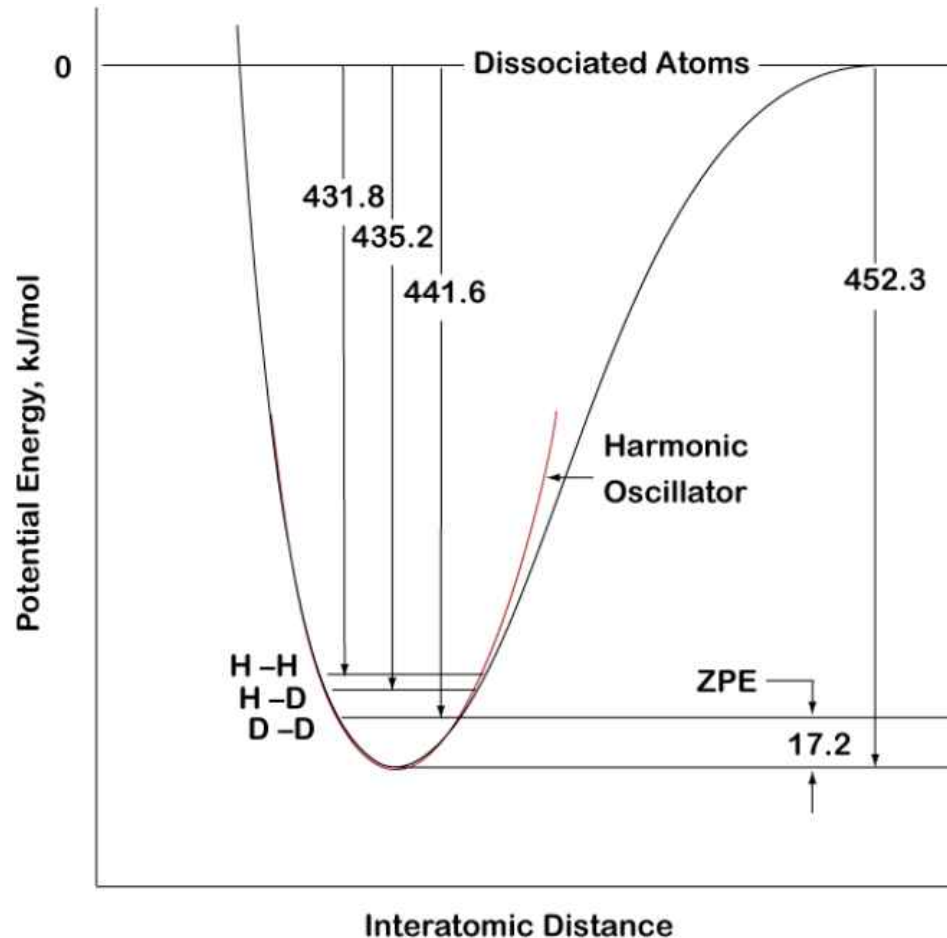
^{17}O – 0.04 %

^{18}O – 0.2 %

$$\delta^{18}\text{O} = \left[\frac{(^{18}\text{O}/^{16}\text{O})_{\text{sam}} - (^{18}\text{O}/^{16}\text{O})_{\text{SMOW}}}{(^{18}\text{O}/^{16}\text{O})_{\text{SMOW}}} \right] \times 10^3$$

Isotope Geochemistry

Isotope fractionation



vibrational energy
of hydrogen

Cornell University,
after O'Neil (1986)

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Isotope Geochemistry

Isotope fractionation

R_X – isotope ratio in reservoir X

α, ε – fractionation factor

$$\alpha_{A-B} = R_A / R_B \quad \text{e.g. 1.0098 for } ^{18}\text{O} \text{ in evap. @ } 20^\circ\text{C}$$

$$\varepsilon = (\alpha - 1) * 10^3$$

$$\varepsilon_{A-B} \sim \delta_A - \delta_B$$

Isotope Geochemistry

Isotope fractionation

R_X – isotope ratio in reservoir X

α, ϵ – fractionation factor

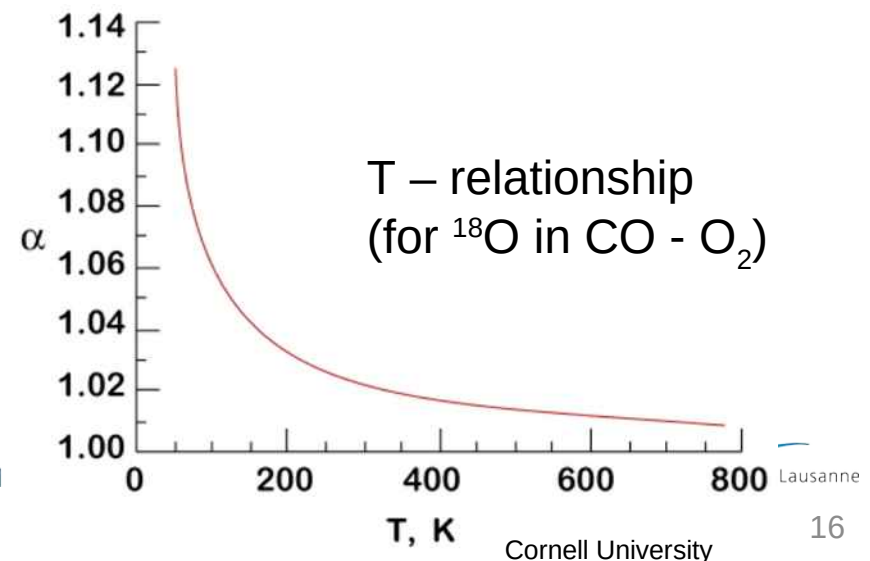
$$\alpha_{A-B} = R_A / R_B$$

e.g. 1.0098 for ^{18}O in evap. @ 20°C

$$\epsilon = (\alpha - 1) * 10^3$$

$$\epsilon_{A-B} \sim \delta_A - \delta_B$$

$$\alpha_{A-B} \sim 1/T^2$$



Isotope Geochemistry

Equilibrium fractionation:

- slow complete equilibration
(e.g. condensation @ 100% humidity)
- heavier isotopes enriched in colder phase

Kinetic fractionation:

- fast or incomplete reactions
(e.g. condensation @ < 100% humidity or with immediate rain)
- unidirectional
- more complex

Isotope Geochemistry

Rayleigh fractionation:

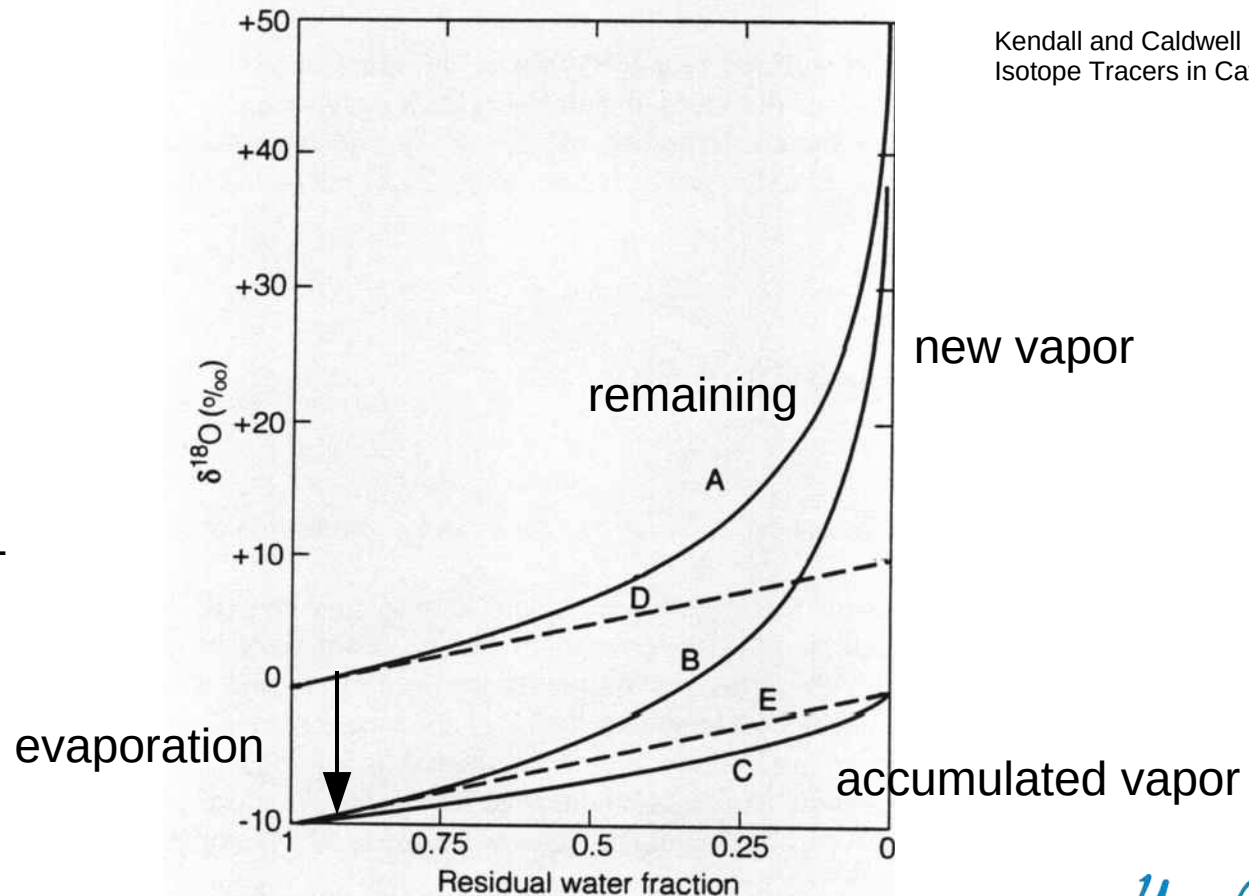
- equilibrium fractionation with removal of product
- reservoir decreases in size
- e.g.: raining clouds

$$R = R_0 f^{(\alpha-1)} \quad (f = \text{fraction remaining})$$

Isotope Geochemistry

Rayleigh fractionation during water evaporation

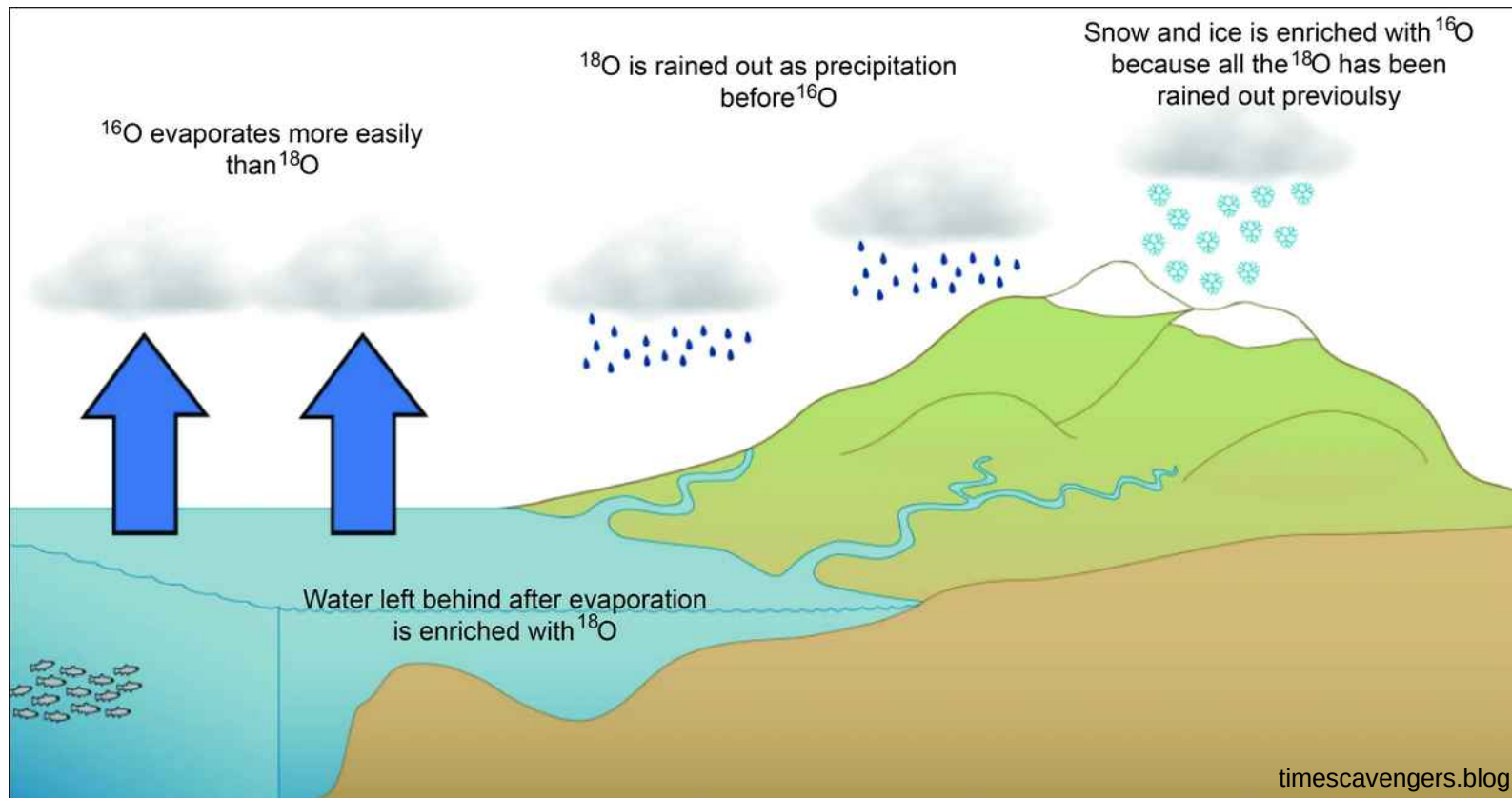
$$\alpha = 1.01$$



Oxygen Isotope Geochemistry

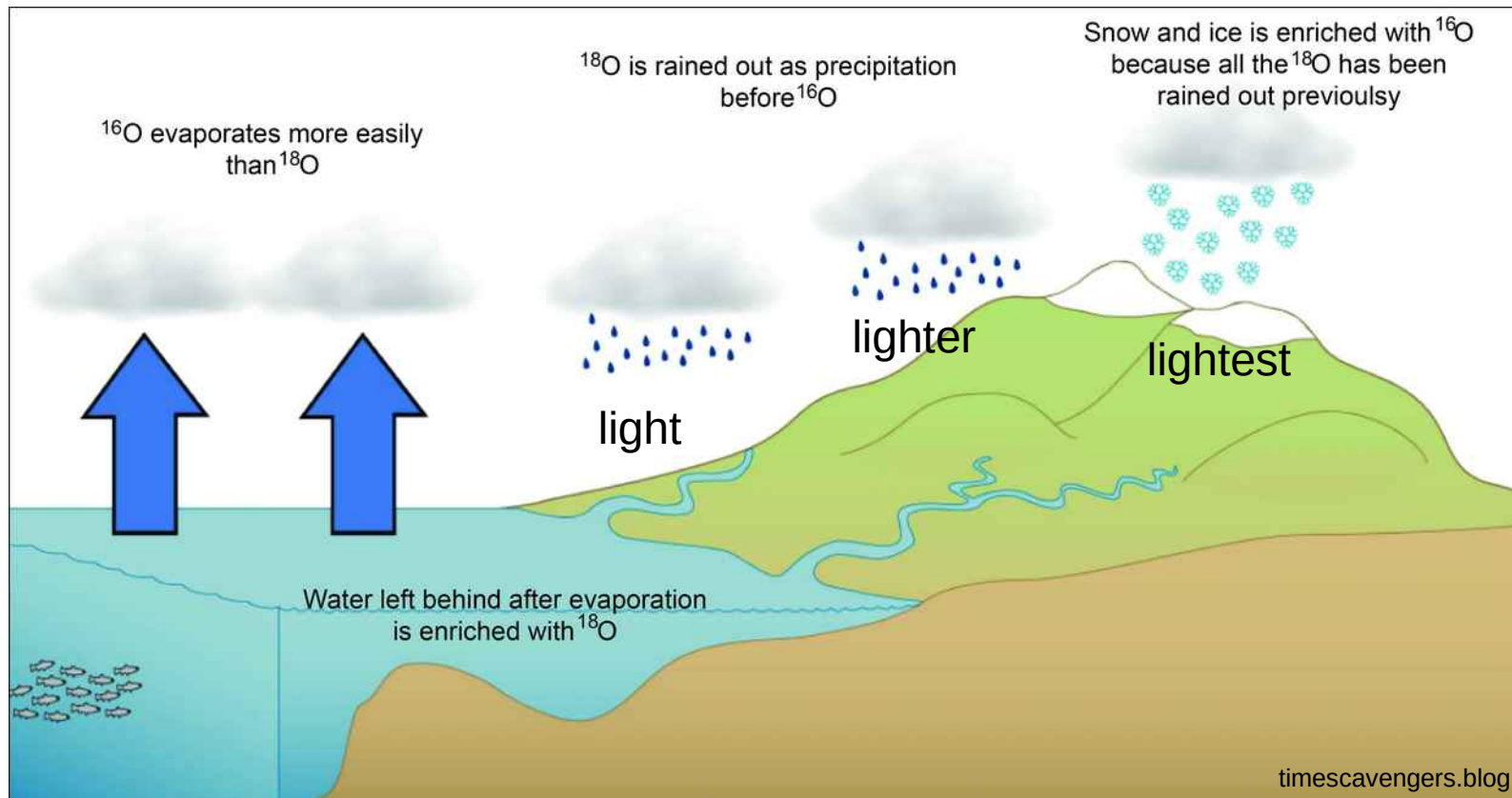
Oxygen Isotope Geochemistry

Rayleigh fractionation during water evaporation



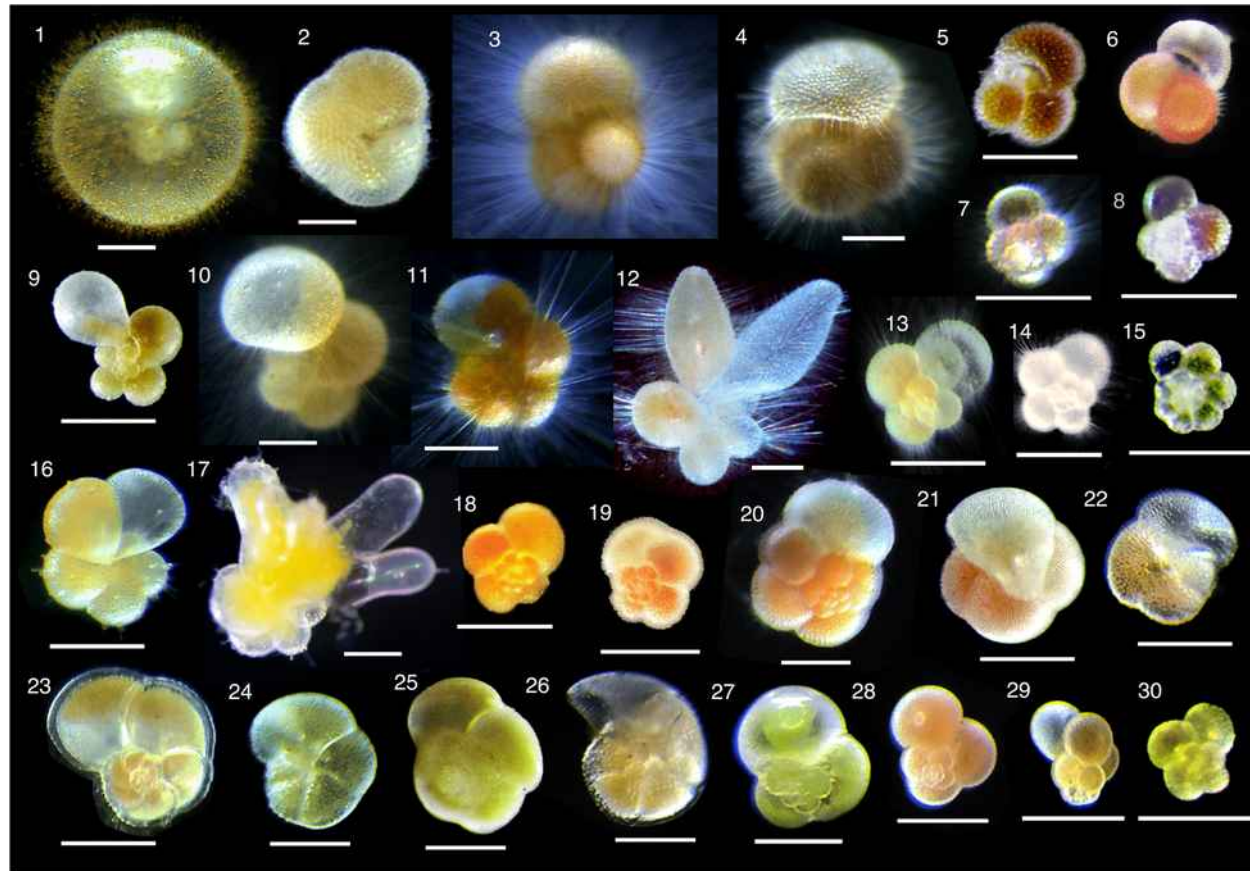
Oxygen Isotope Geochemistry

Rayleigh fractionation during water evaporation



Oxygen Isotopes in Carbonate

foraminifera – protists with CaCO_3 shells



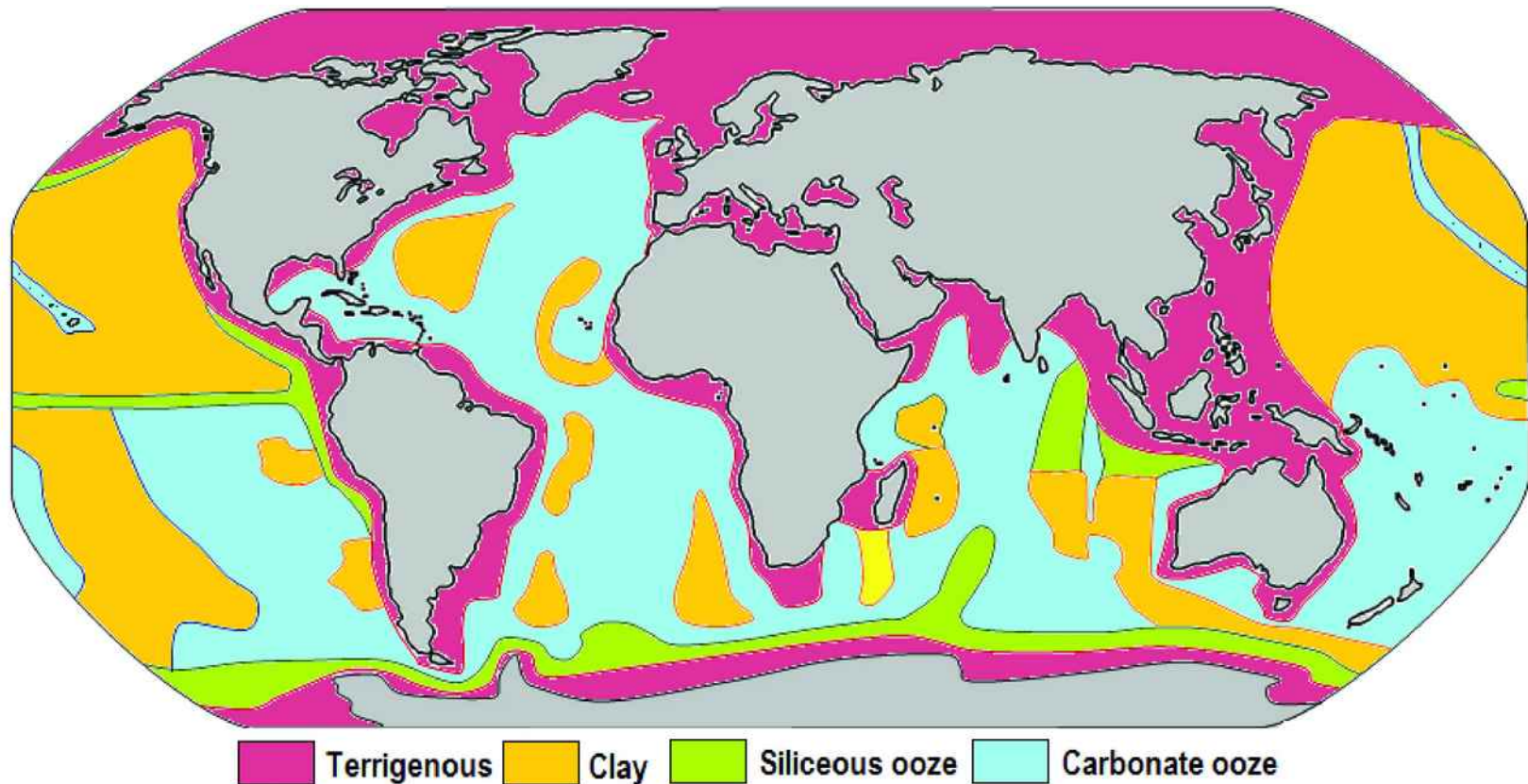
Oxygen Isotopes in Carbonate

foraminifera – protists with CaCO_3 shells

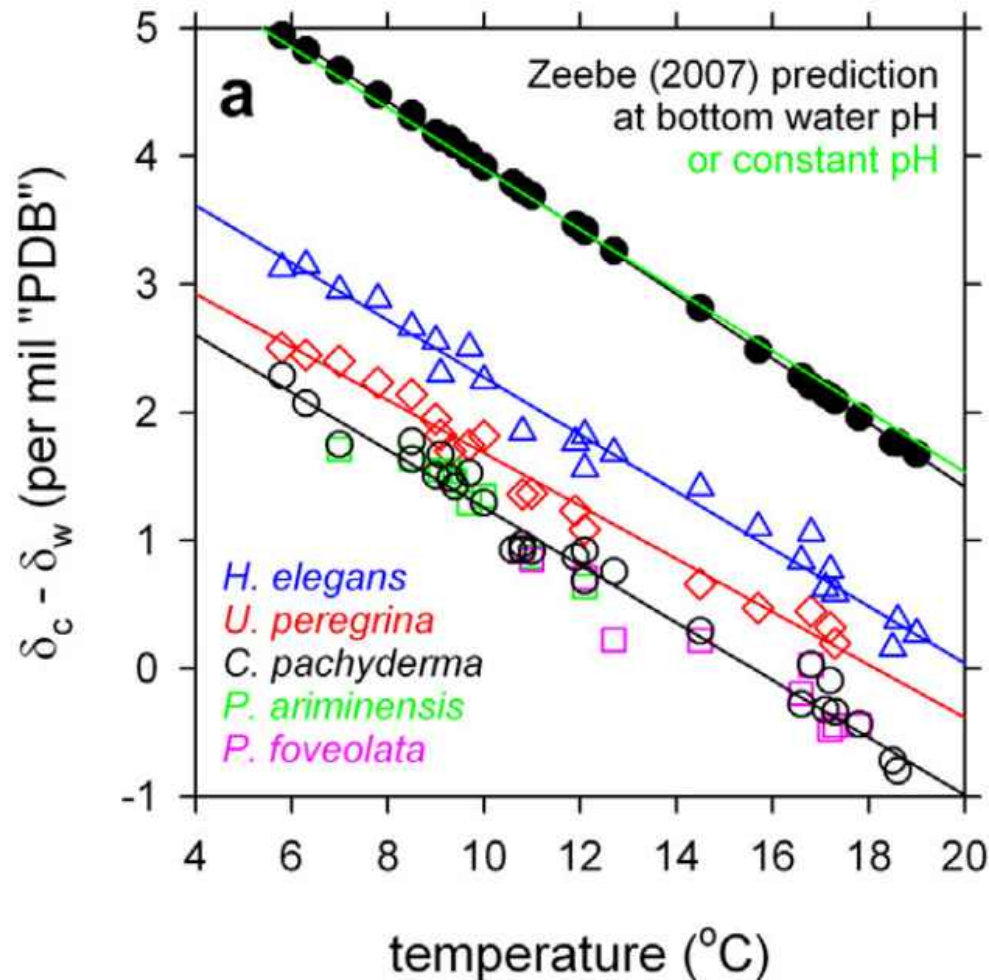


Oxygen Isotopes in Carbonate

foraminifera – protists with CaCO_3 shells



Oxygen Isotopes in Carbonate



$\sim 0.20 - 0.25 \text{ ‰}$
per $^{\circ}\text{C}$

uncertainty
 $\sim 0.1 \text{ ‰}$

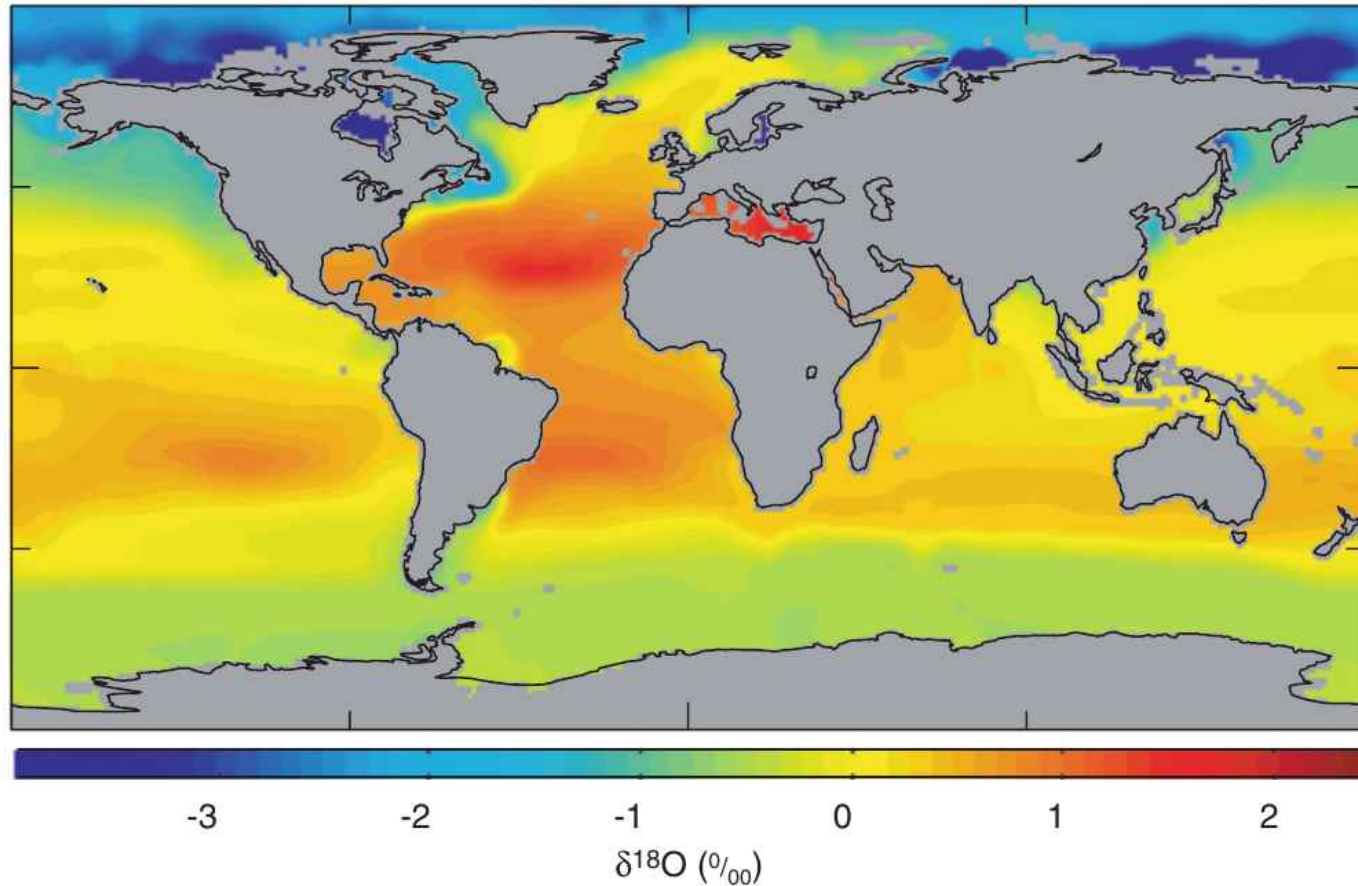
Marchitto et al. (2014)
Geochimica et Cosmochimica Acta

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Oxygen Isotopes in the ocean

$\delta^{18}\text{O}$ in seawater: \sim salinity



Oxygen Isotopes in the ocean

$\delta^{18}\text{O}$ in seawater: \sim salinity

$\delta^{18}\text{O}$ in carbonate: \sim salinity & temperature

$\rightarrow \sim$ density

Oxygen Isotopes in the ocean

$\delta^{18}\text{O}$ in seawater: \sim salinity

$\delta^{18}\text{O}$ in carbonate: \sim salinity & temperature

$\rightarrow \sim$ density

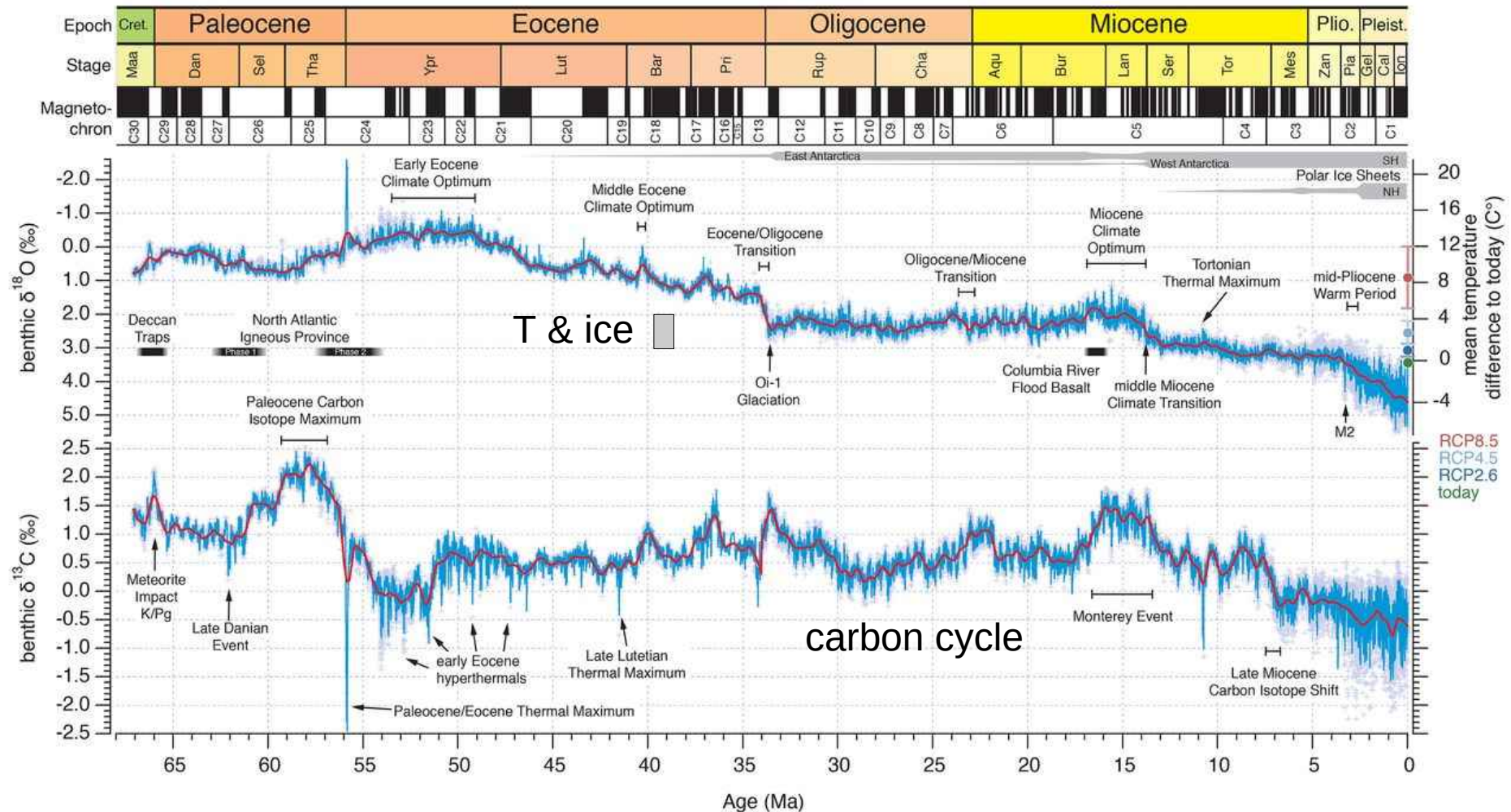
BUT salinity in paleoceanography depends on global ice volume!

$\sim +1 \text{ ‰}$ per 100m sea level as ice

Today's continental ice $\sim 0.6 \text{ ‰}$

Last Glacial Maximum ice $\sim +1 \text{ ‰}$

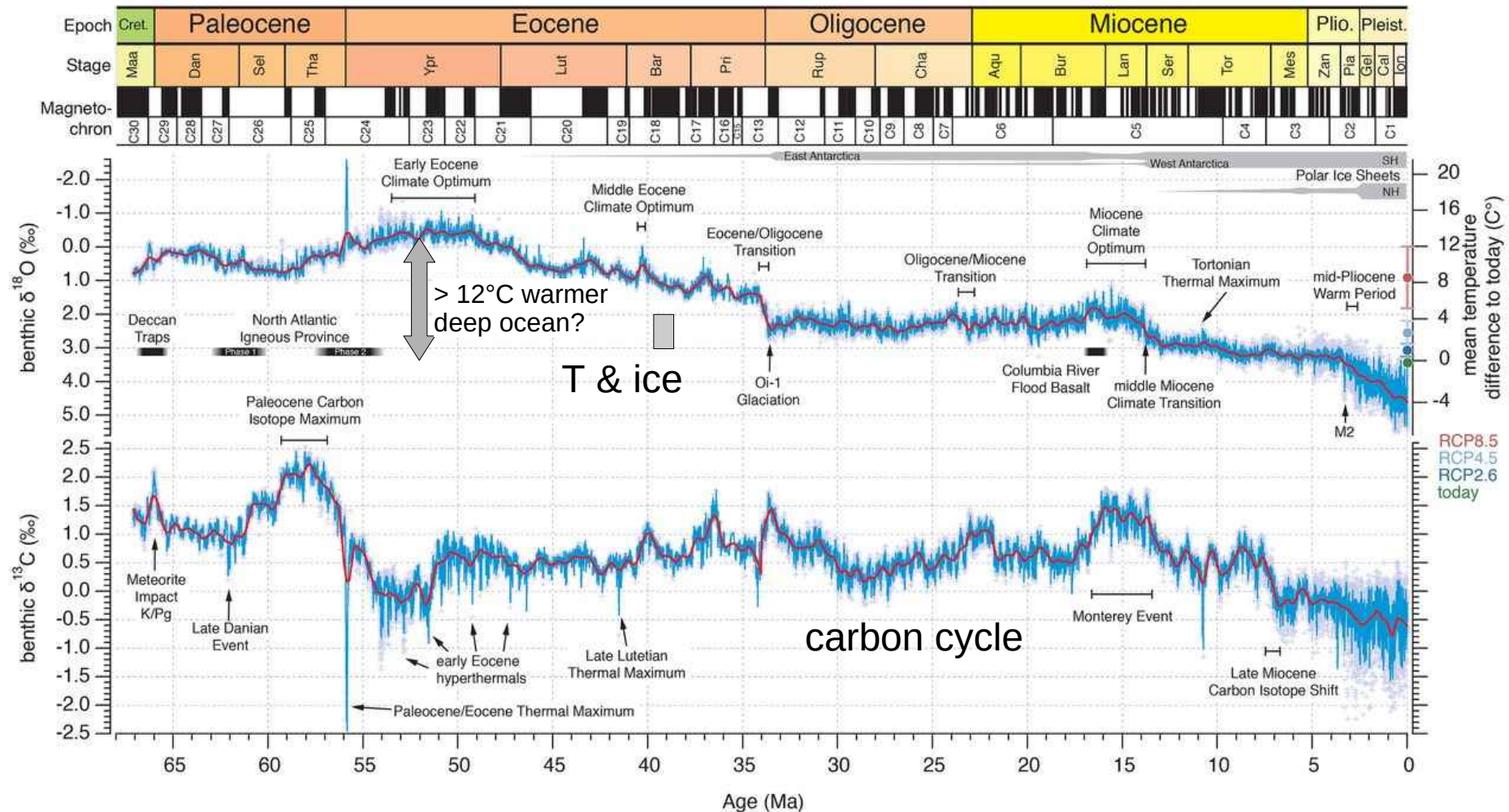
Cenozoic Climate



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Cenozoic Climate



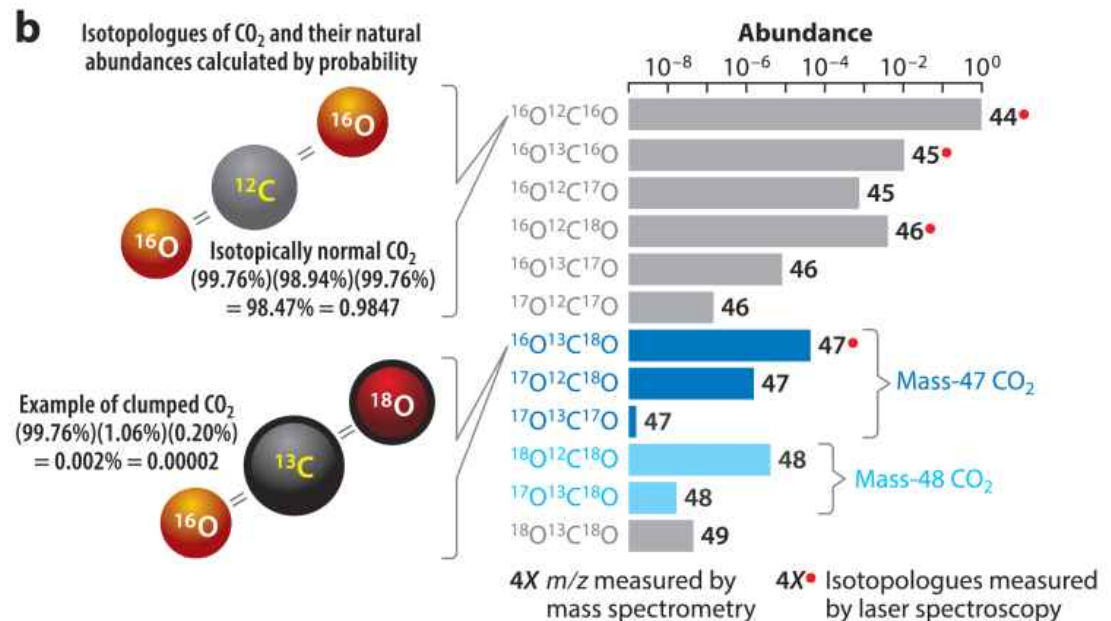
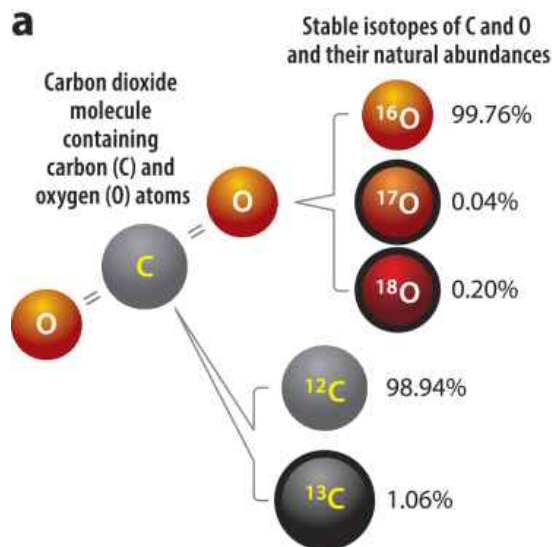
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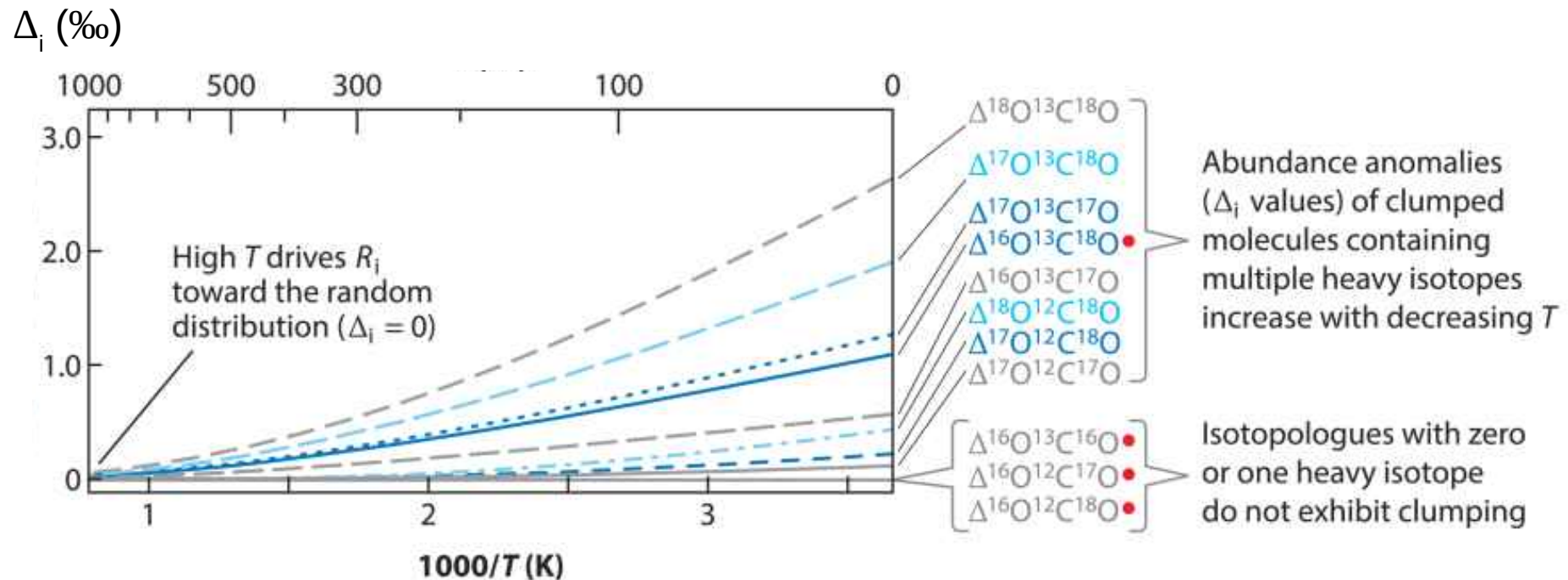
Clumped Isotope Thermometer

Clumped Isotopes

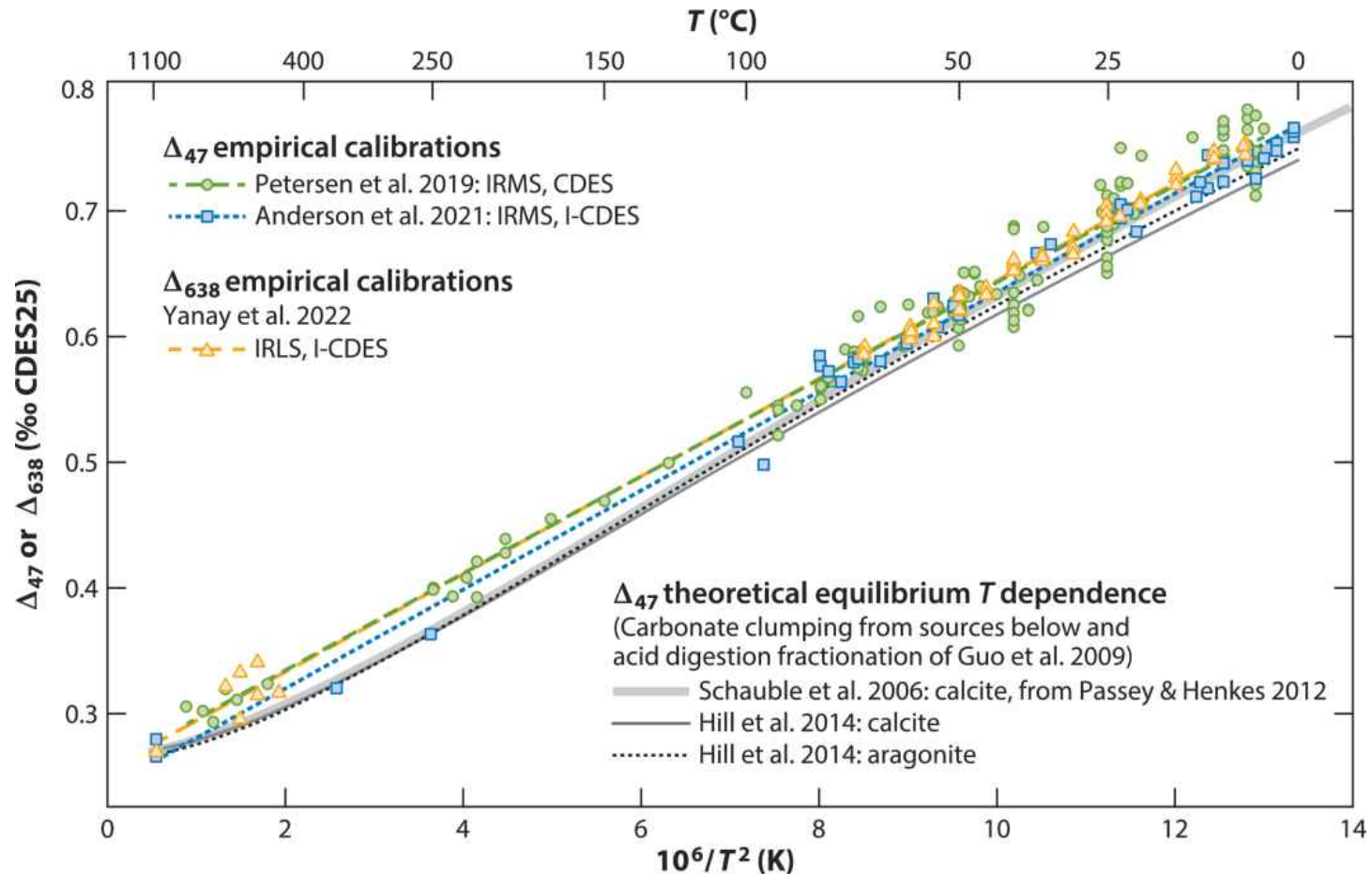
Clumping in carbon dioxide (CO₂)



Clumped Isotopes



Clumped Isotopes



Huntington KW, Petersen SV. 2023
Annu. Rev. Earth Planet. Sci. 51:611–41

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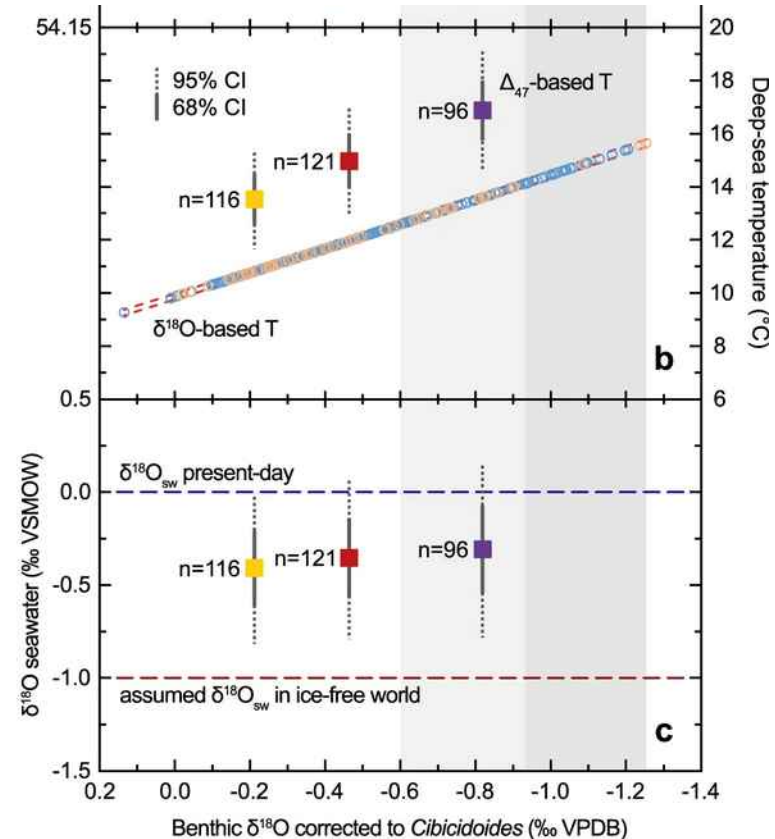
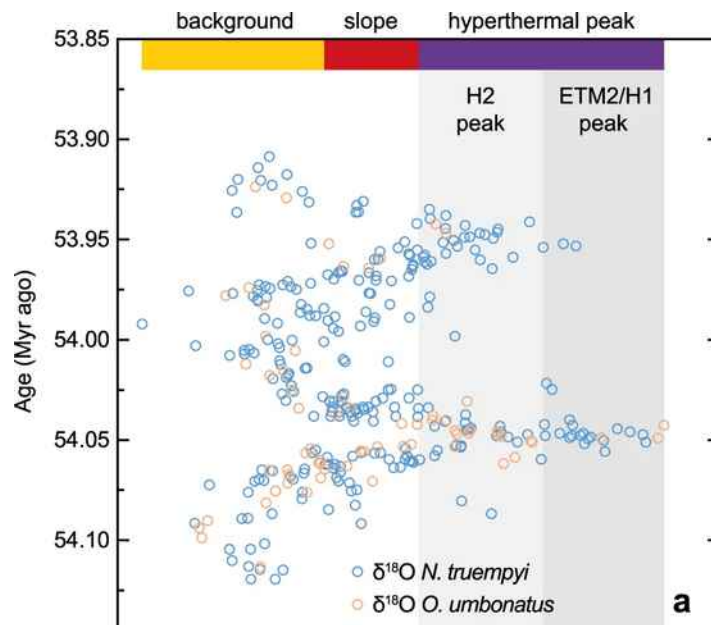
Clumped Isotopes

- very accurate
- few secondary effects

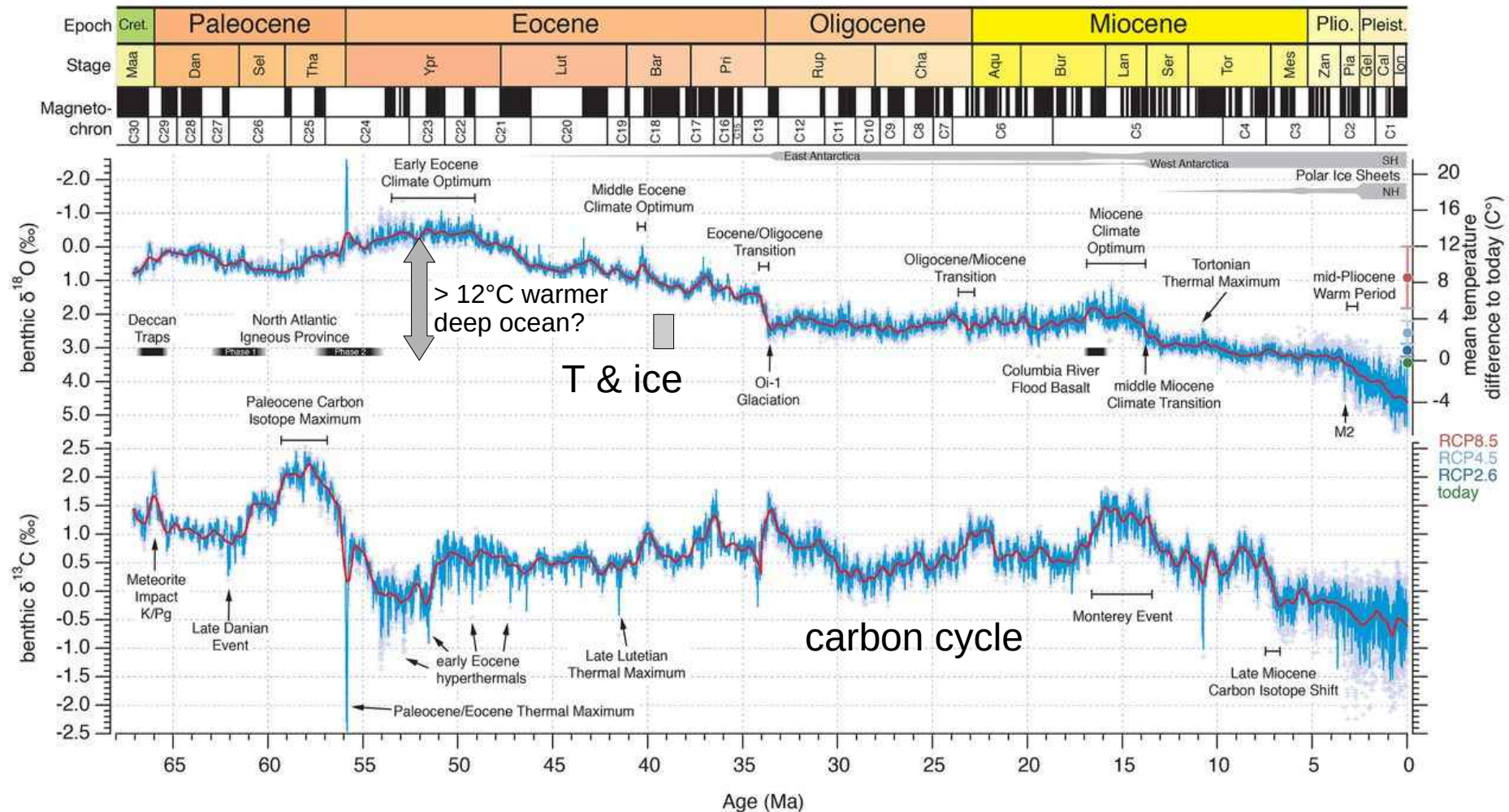
BUT

- low precision → many replicates
- large samples (few mg)
 - costly sample analysis

Clumped isotope temperatures



Cenozoic Climate



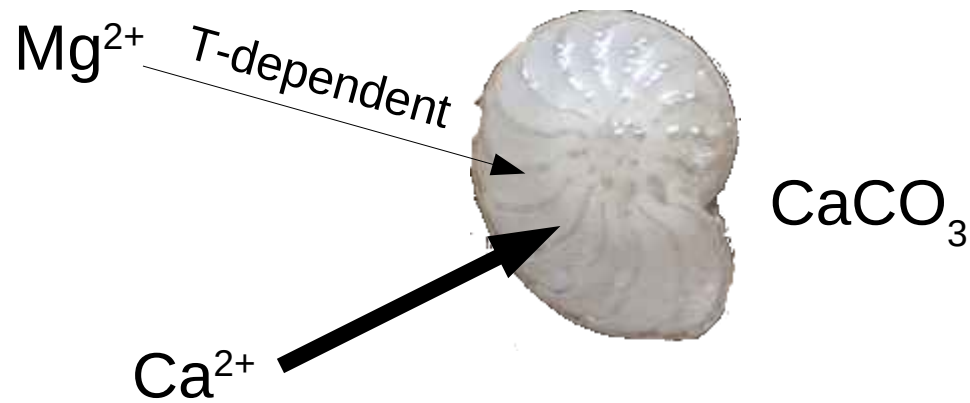
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Mg/Ca paleothermometer

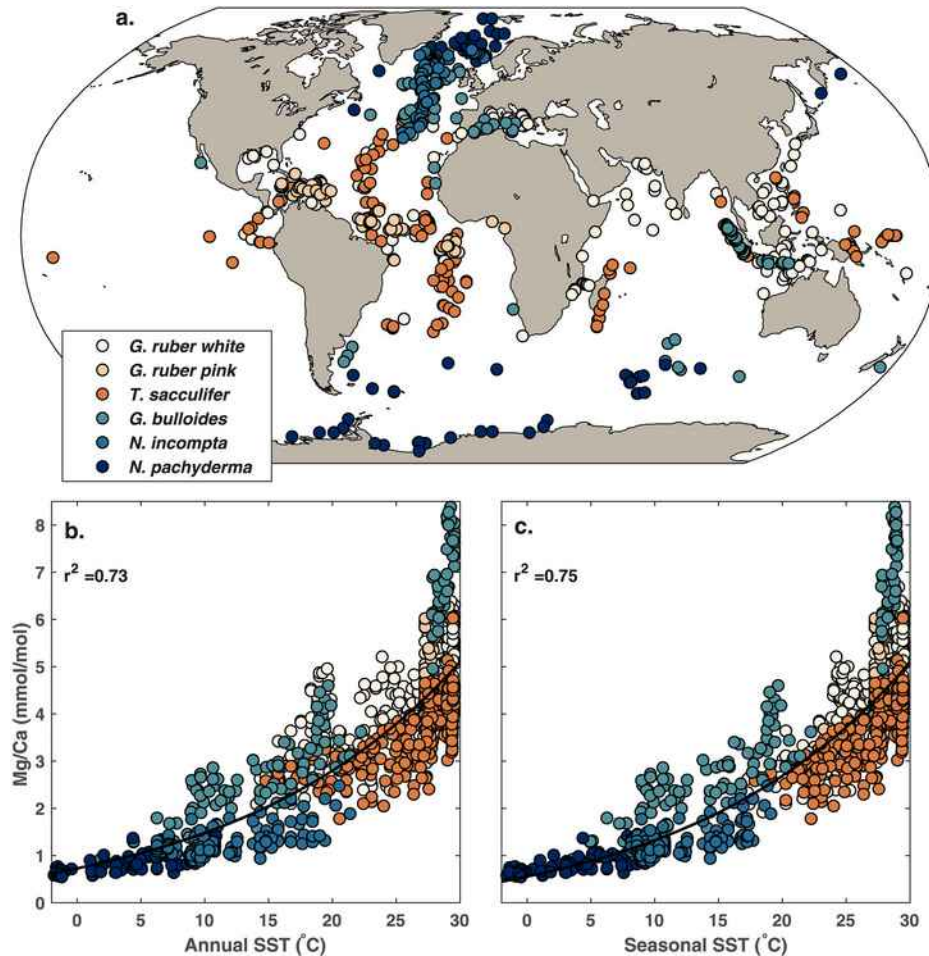
Mg/Ca paleothermometer

Mg/Ca in foraminifera



Mg/Ca paleothermometer

Mg/Ca in foraminifera



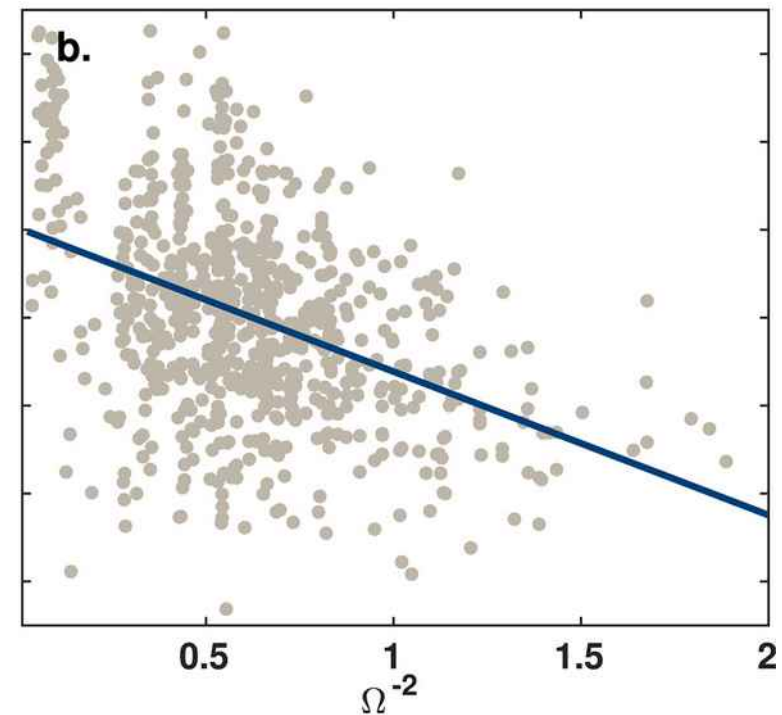
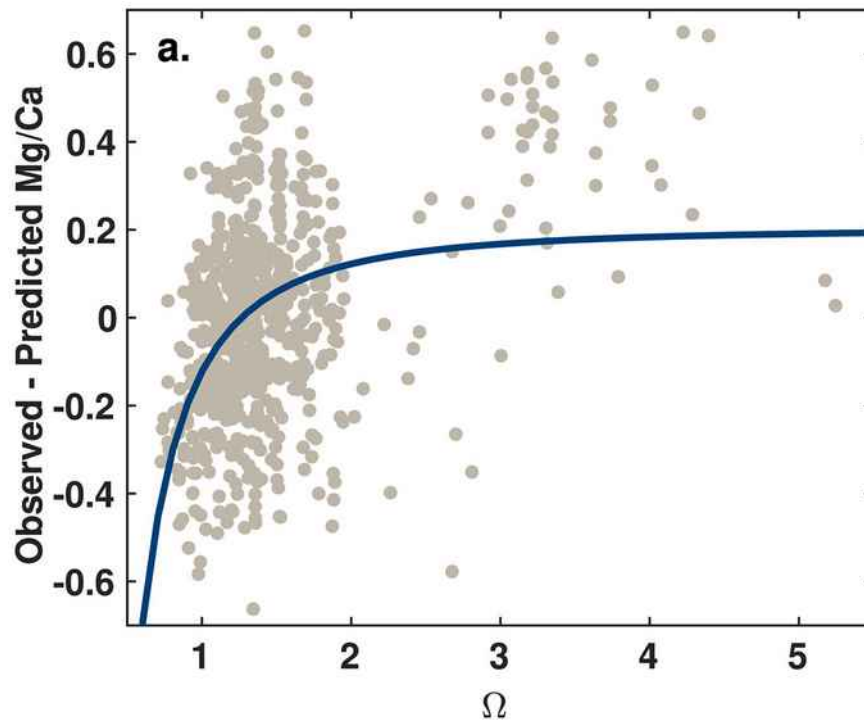
Tierney et al. (2019)
Paleoceanography and
Paleoclimatology

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Mg/Ca paleothermometer

Mg/Ca in foraminifera



Tierney et al. (2019)
Paleoceanography and
Paleoclimatology

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Mg/Ca paleothermometer

Mg/Ca in foraminifera

- records past water temperature
- species-specific calibrations
- Mg/Ca in seawater – dependent
- Ω – dependent
- precision typically $\sim 1^\circ\text{C}$

TEX86 paleothermometer

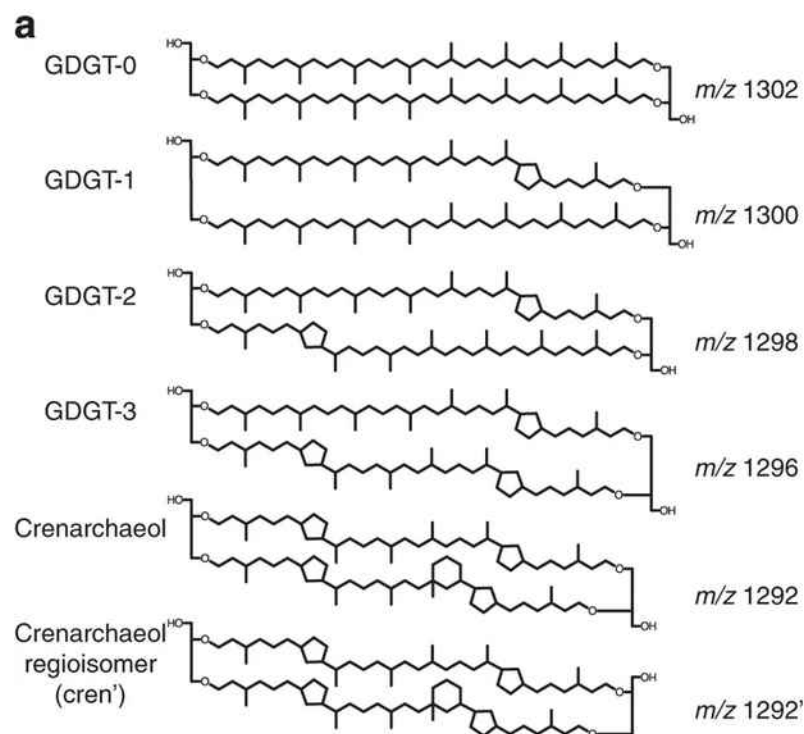
TEX86 paleothermometer

TEX86 proxy for SST

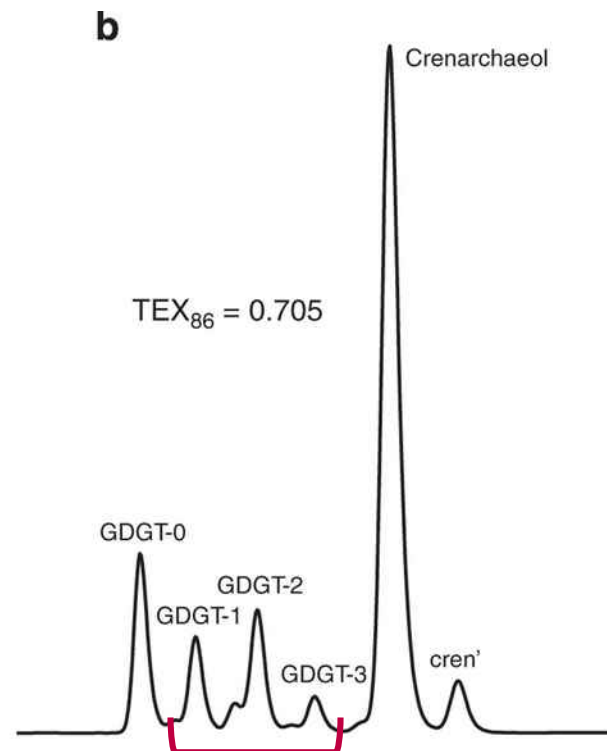
- based on specific organic molecules from sediments formed by archaea (similar to bacteria)
- the abundance ratio of certain molecules depends on ambient seawater temperatures
- mainly records near-sea surface T

TEX86 paleothermometer

TEX86 proxy for SST



GDGT = Glycerol dialkyl glycerol tetraether (lipids)



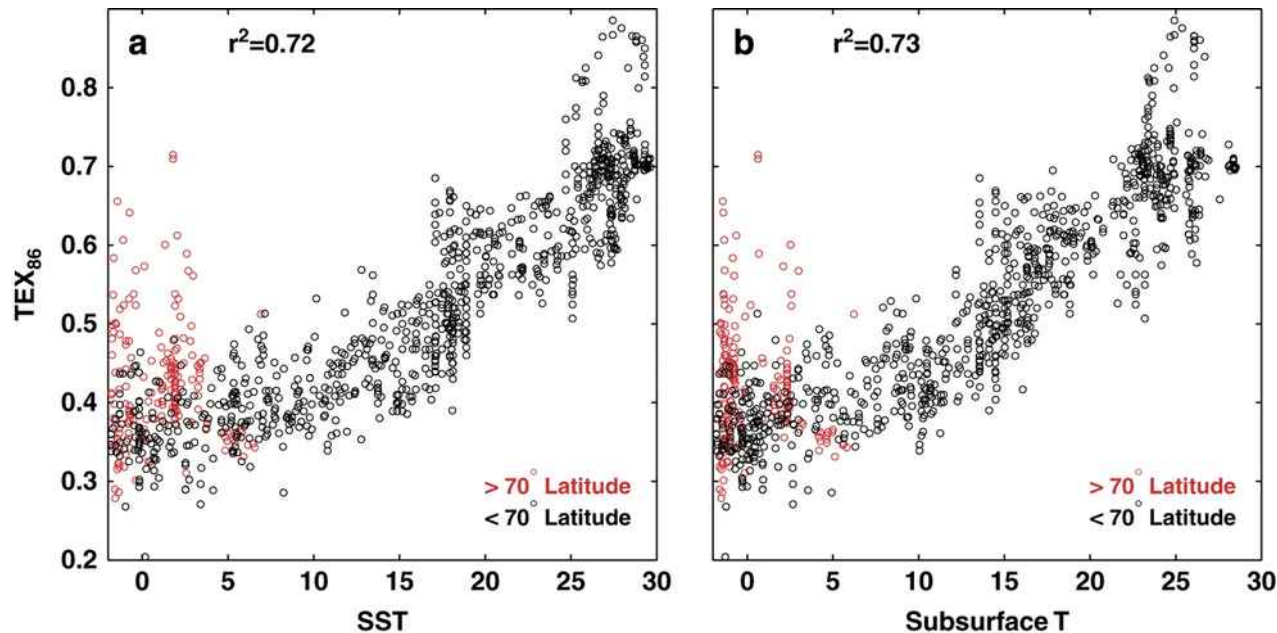
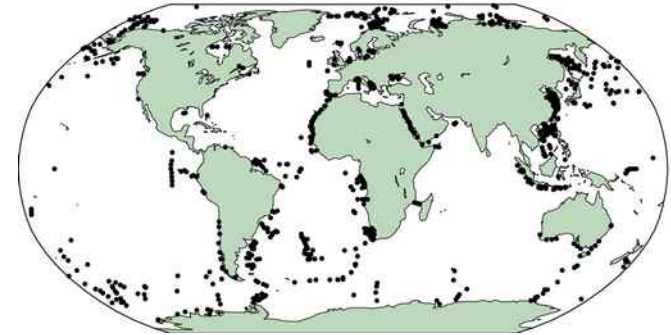
Tierney & Tingley (2015)
Nature Scientific data

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TEX₈₆ paleothermometer

TEX₈₆ proxy for SST



Tierney & Tingley (2015)
Nature Scientific data

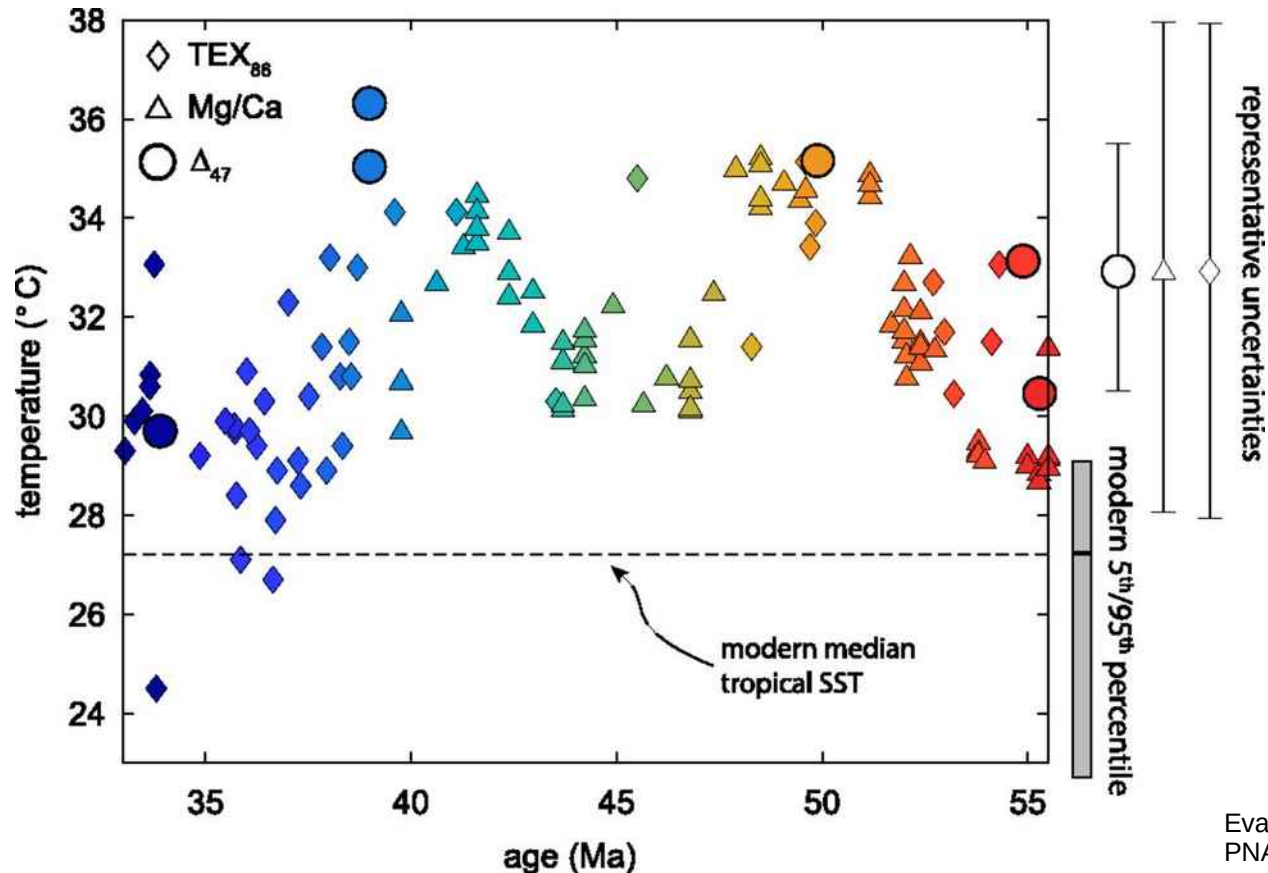
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Hothouse Climate

Hothouse Climate

Eocene low latitude sea surface T



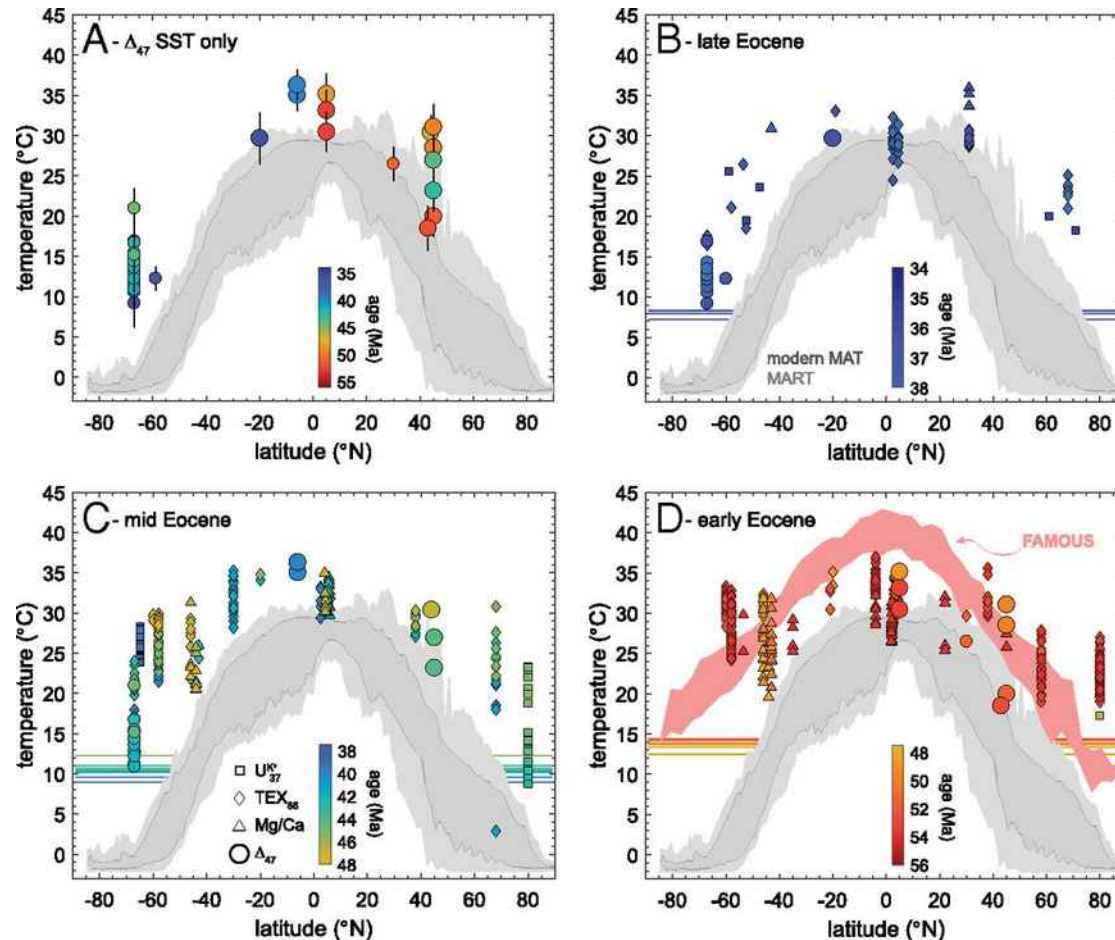
Evans et al. (2018)
PNAS

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Equable Climate

Eocene latitudinal sea surface T



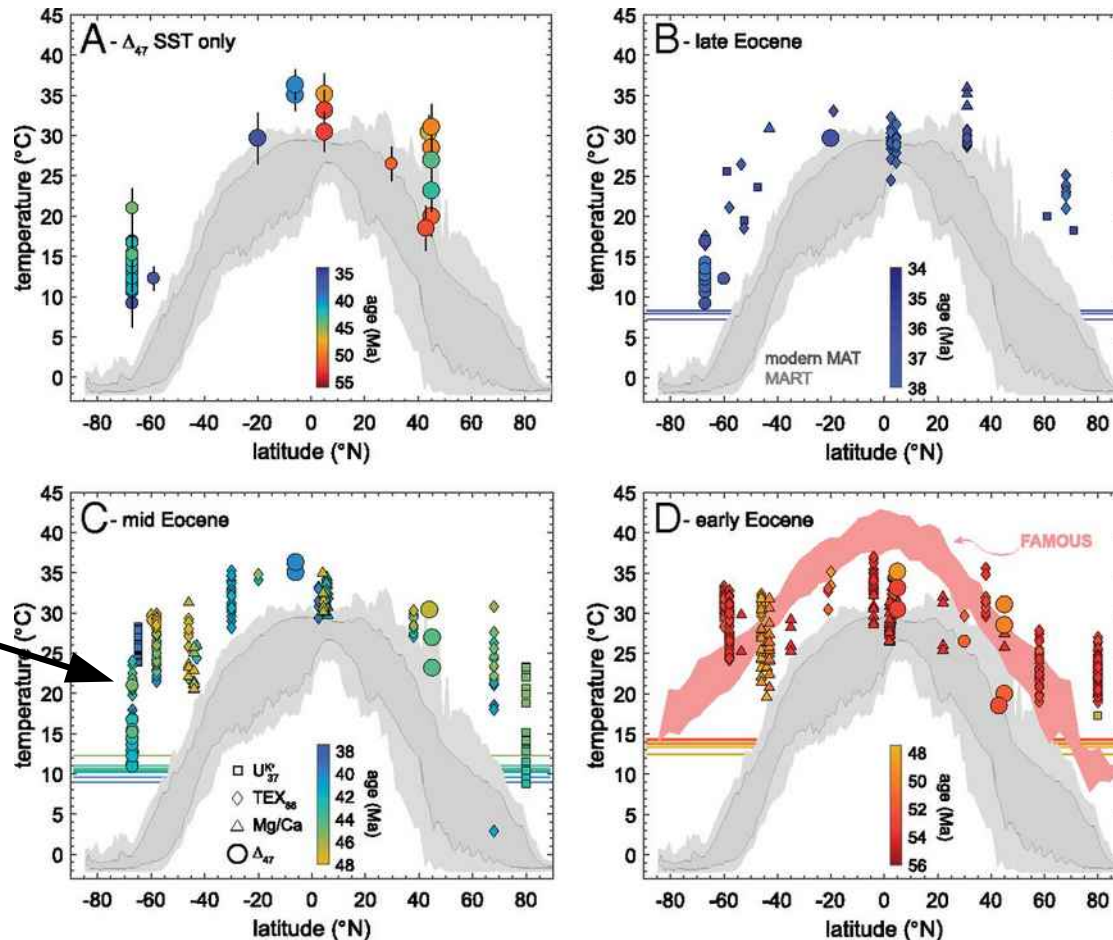
Evans et al. (2018)
PNAS

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Equable Climate

Eocene latitudinal sea surface T



warm
poles!

“Equable
Climates”

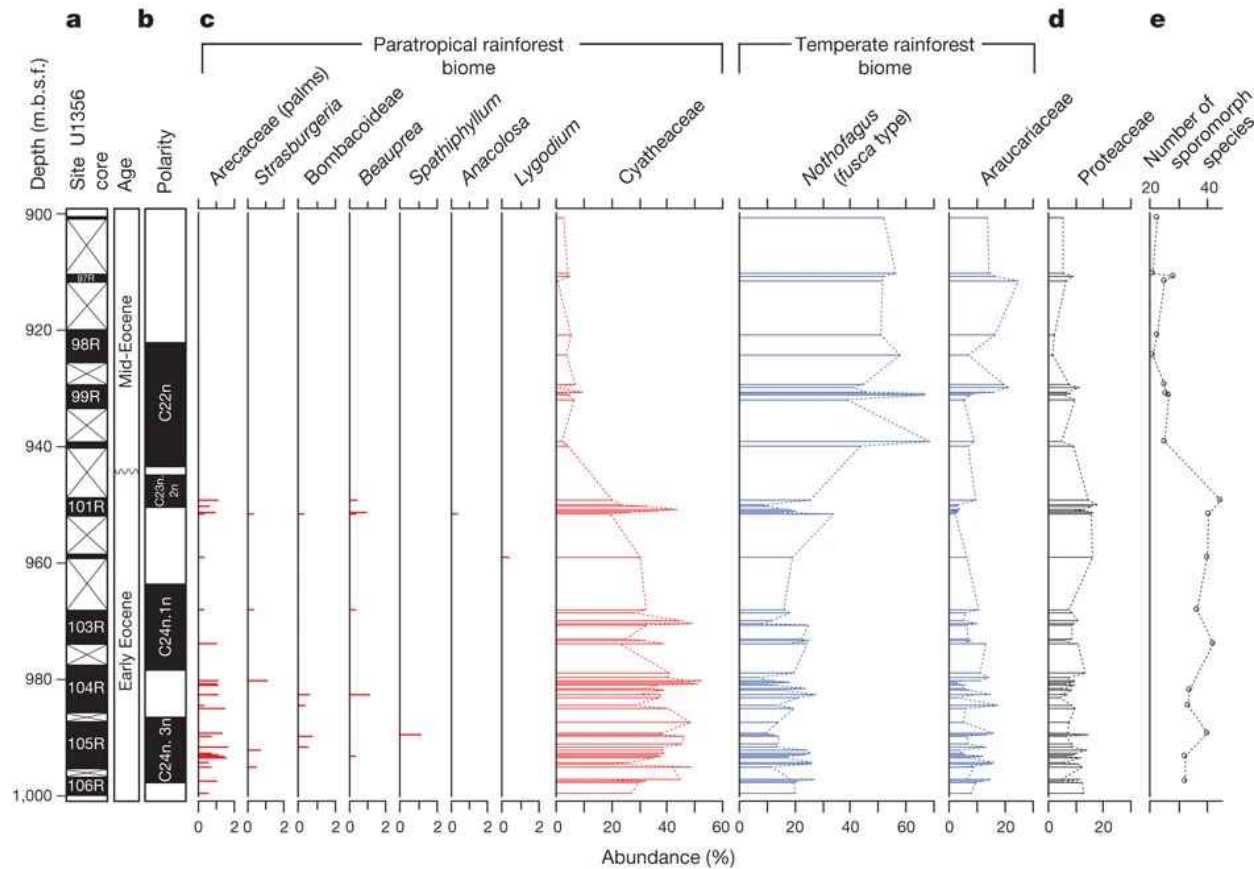
Evans et al. (2018)
PNAS

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Equable Climate

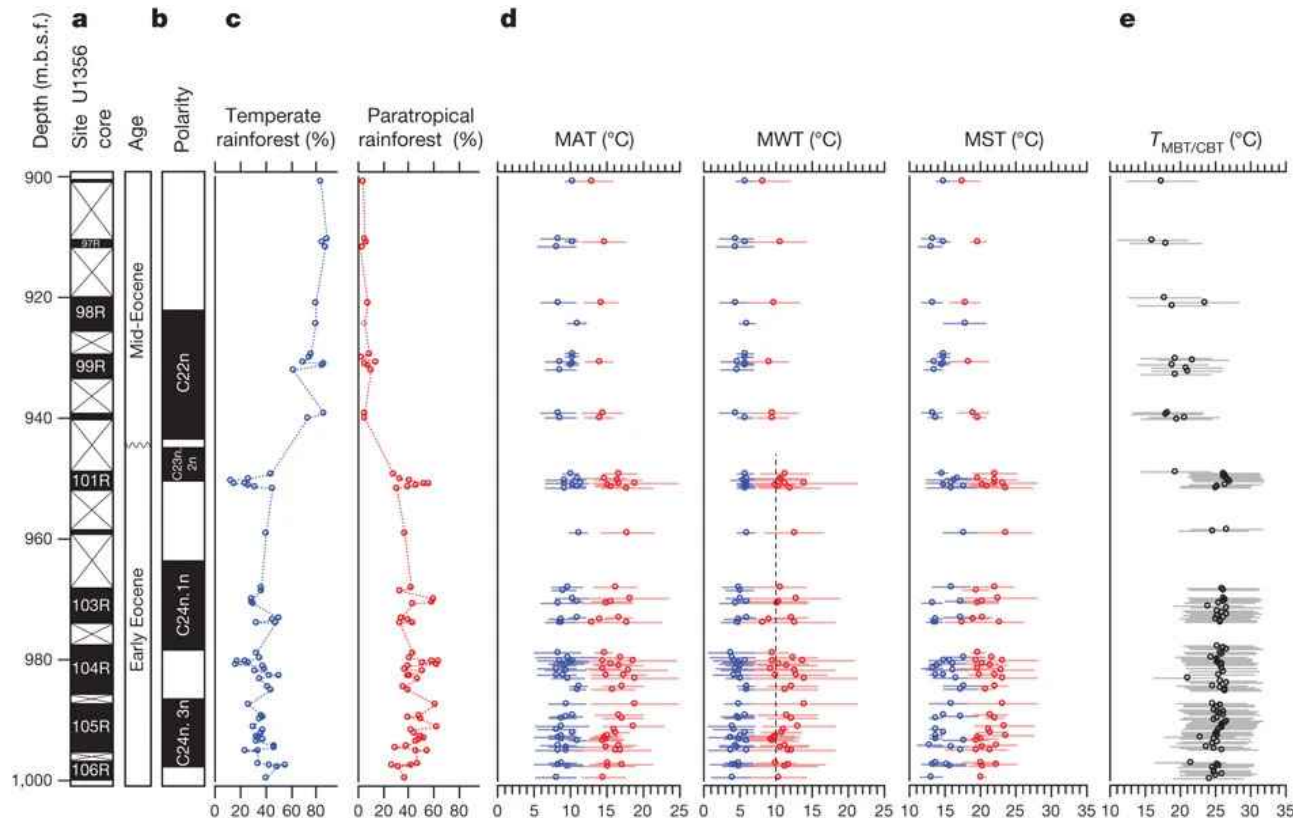
Eocene East Antarctic climate from pollen



Pross et al. (2012)
Nature

Equable Climate

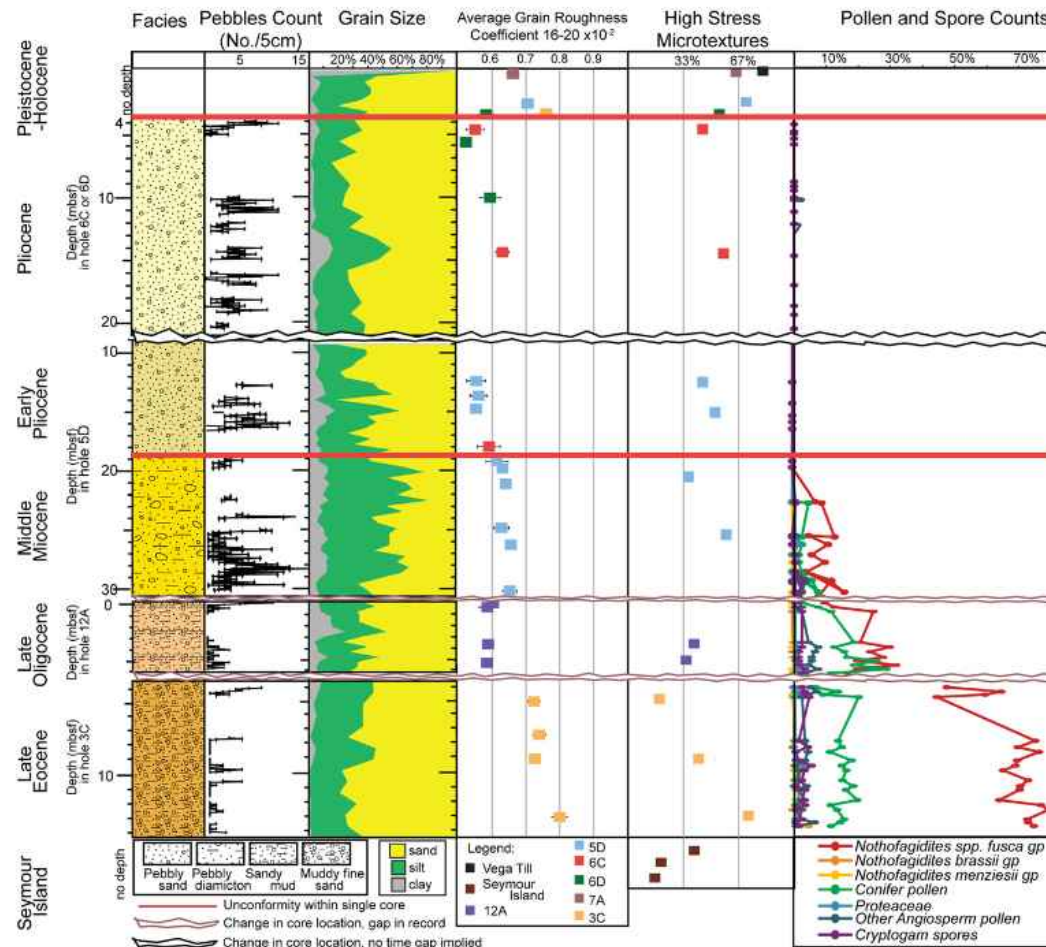
Eocene East Antarctic climate from pollen



Pross et al. (2012)
Nature

Equable Climate

Cenozoic West Antarctic climate from pollen



Beech

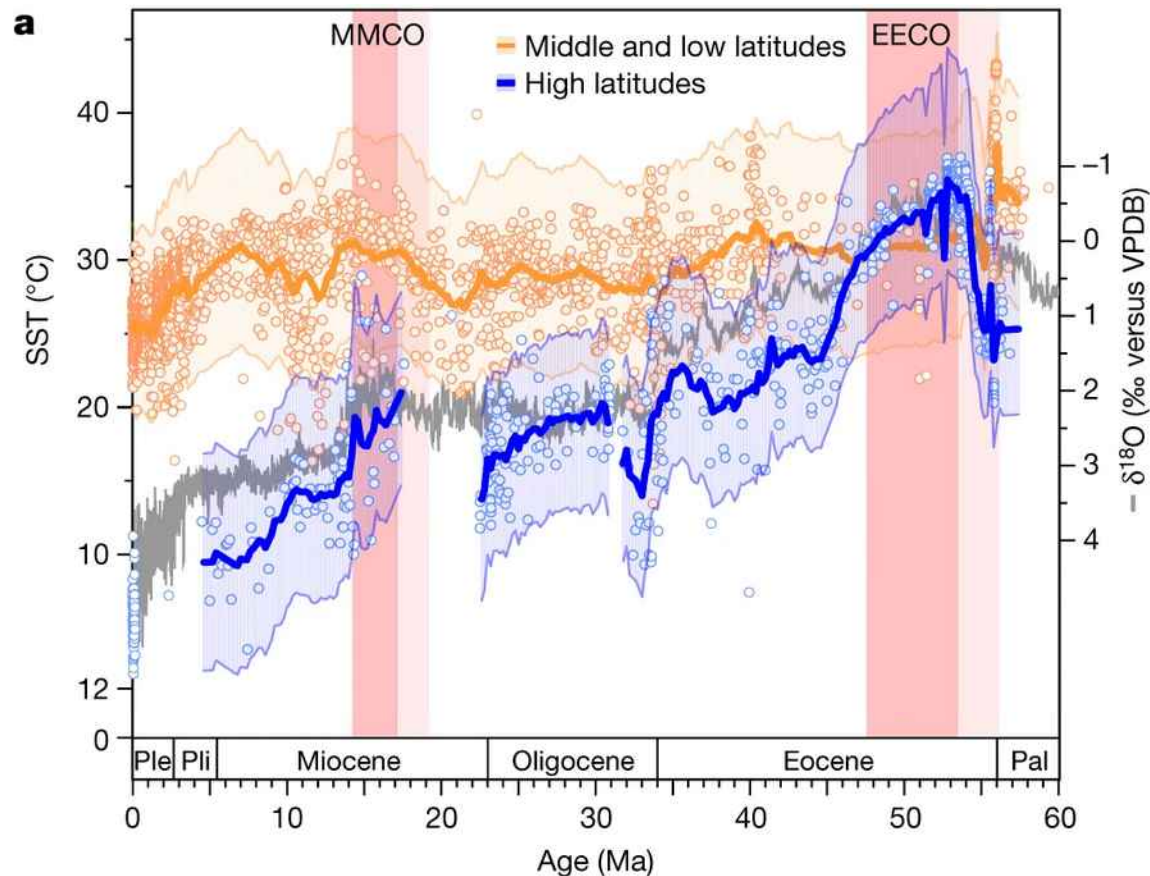
Anderson et al. (2011)
PNAS

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Equable Climate

Evolution of latitudinal temperature gradient

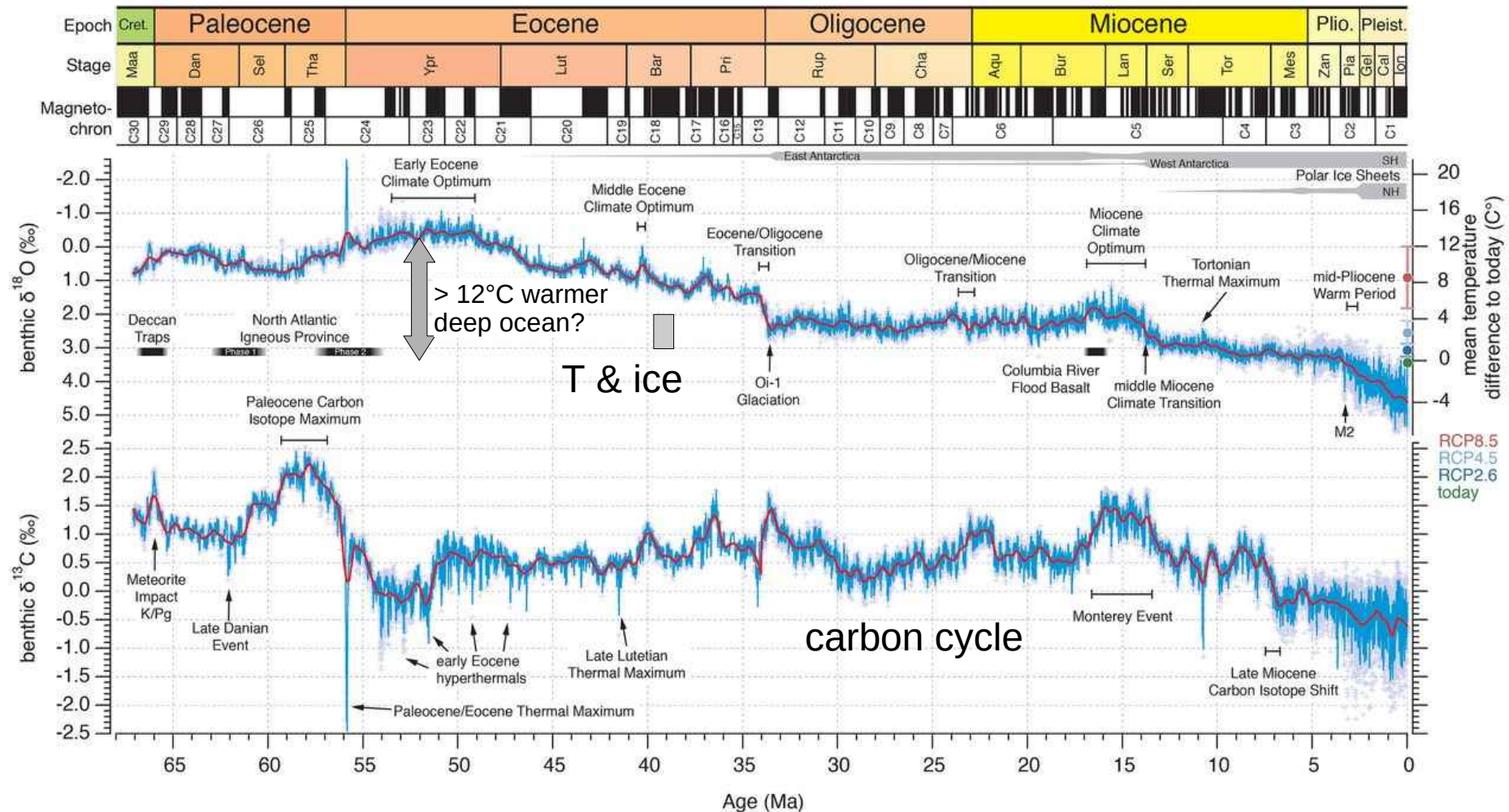


Auderset et al. (2022)
Nature

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Equable Climate



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Equable Climate

Causes

- high altitude cloud cover?



Mark Piana
Harvard University
Equable Climate Dynamics

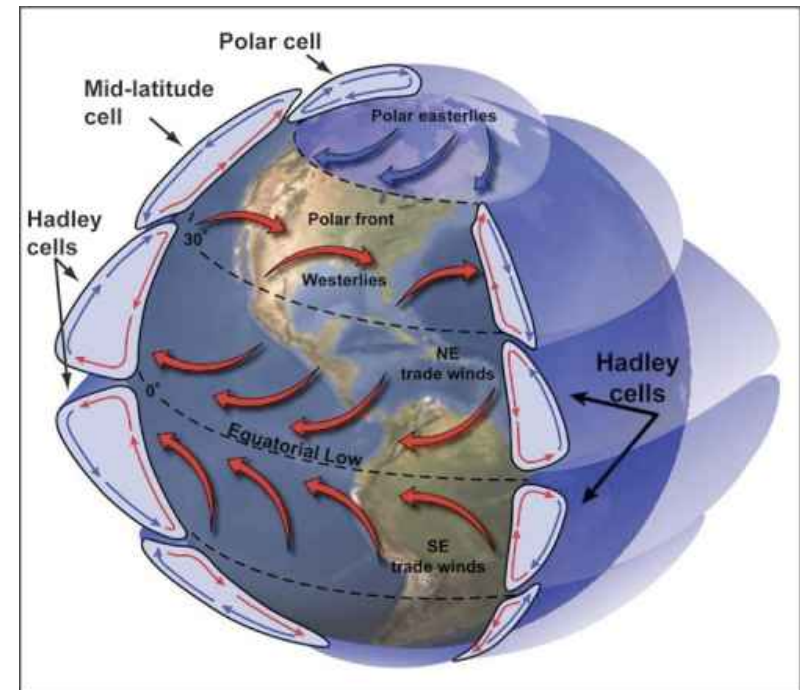
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Equable Climate

Causes

- high altitude cloud cover?
- atmospheric cell change?



Mark Piana
Harvard University
Equable Climate Dynamics

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Equable Climate

Causes

- high altitude cloud cover?
- atmospheric cell change?
- polar stratospheric clouds?

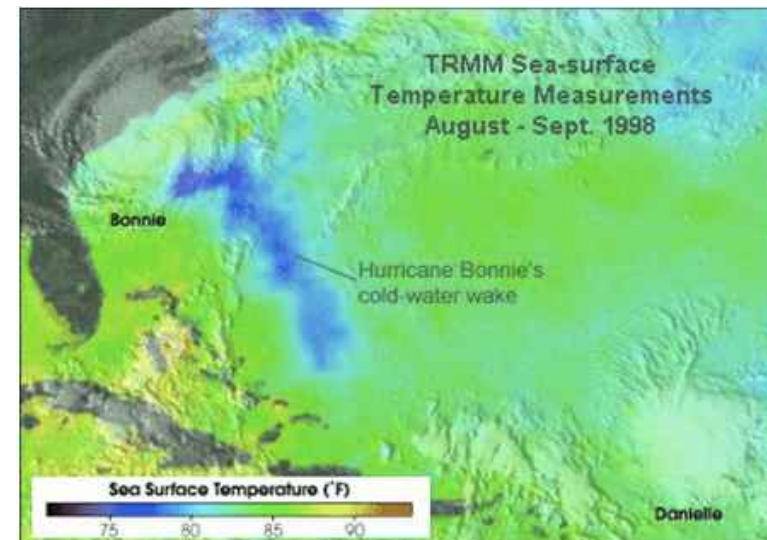


Mark Piana
Harvard University
Equable Climate Dynamics

Equable Climate

Causes

- high altitude cloud cover?
- atmospheric cell change?
- polar stratospheric clouds?
- cyclone ocean mixing?

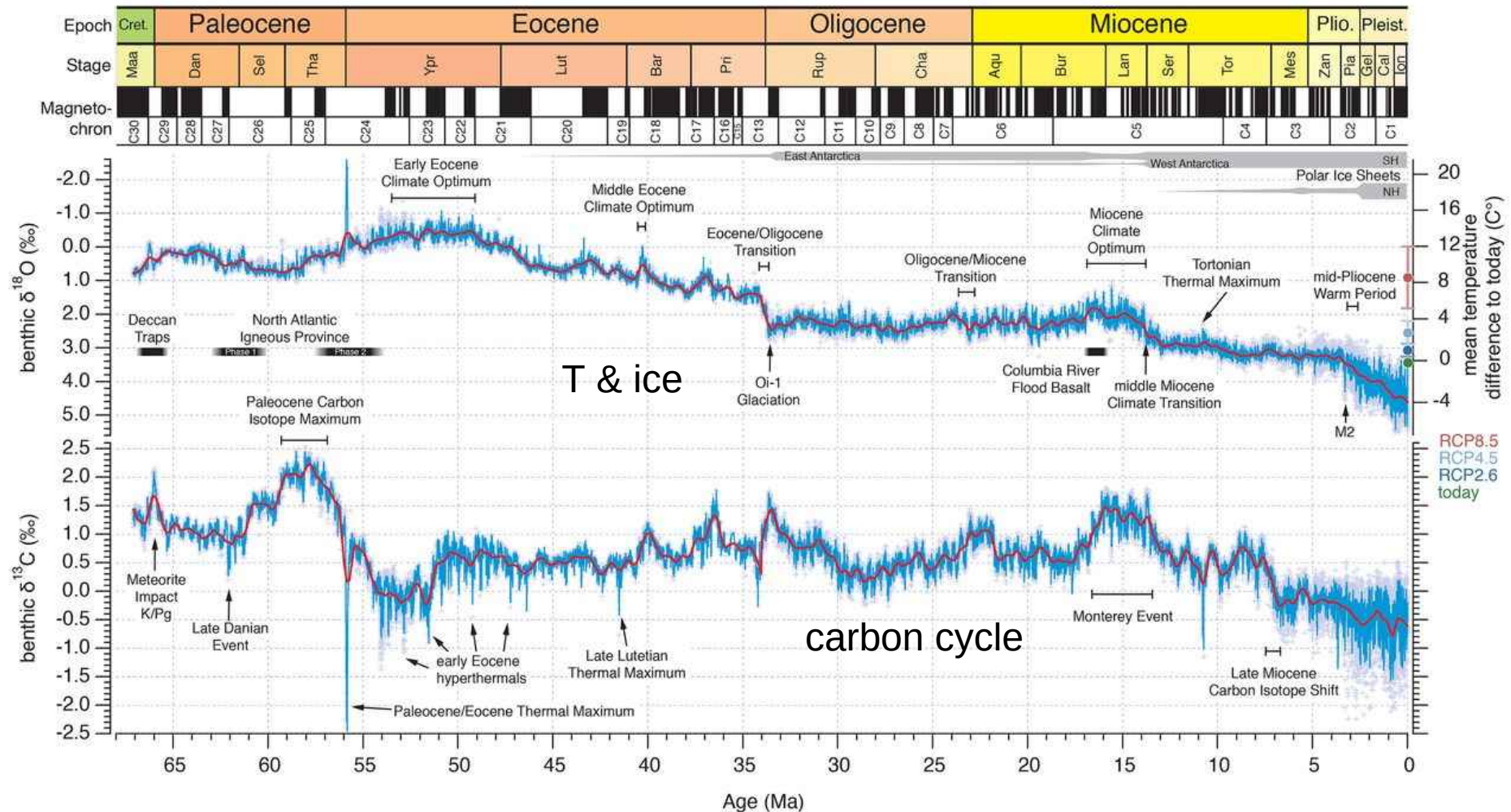


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Cenozoic Climate



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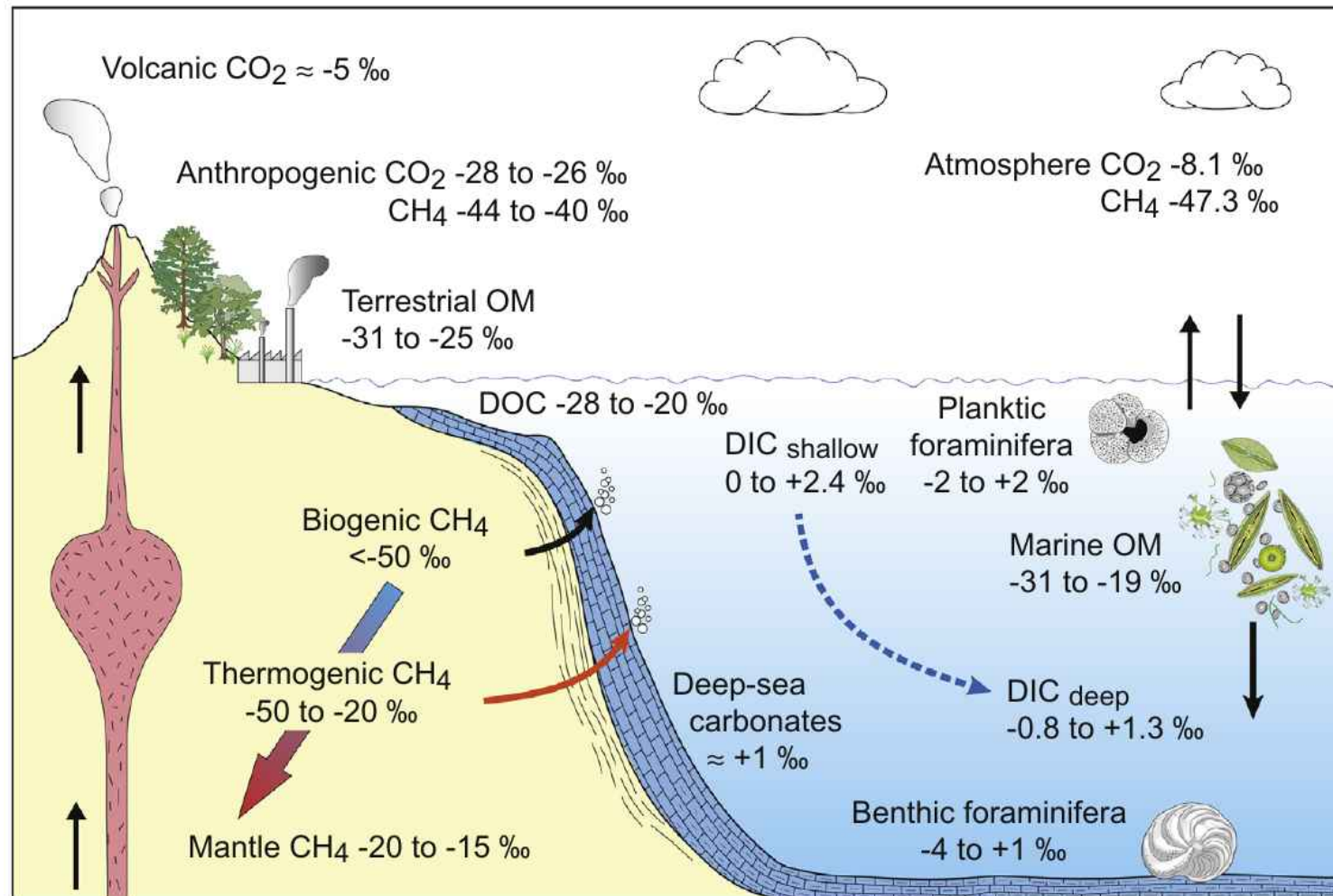
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Carbon Isotopes

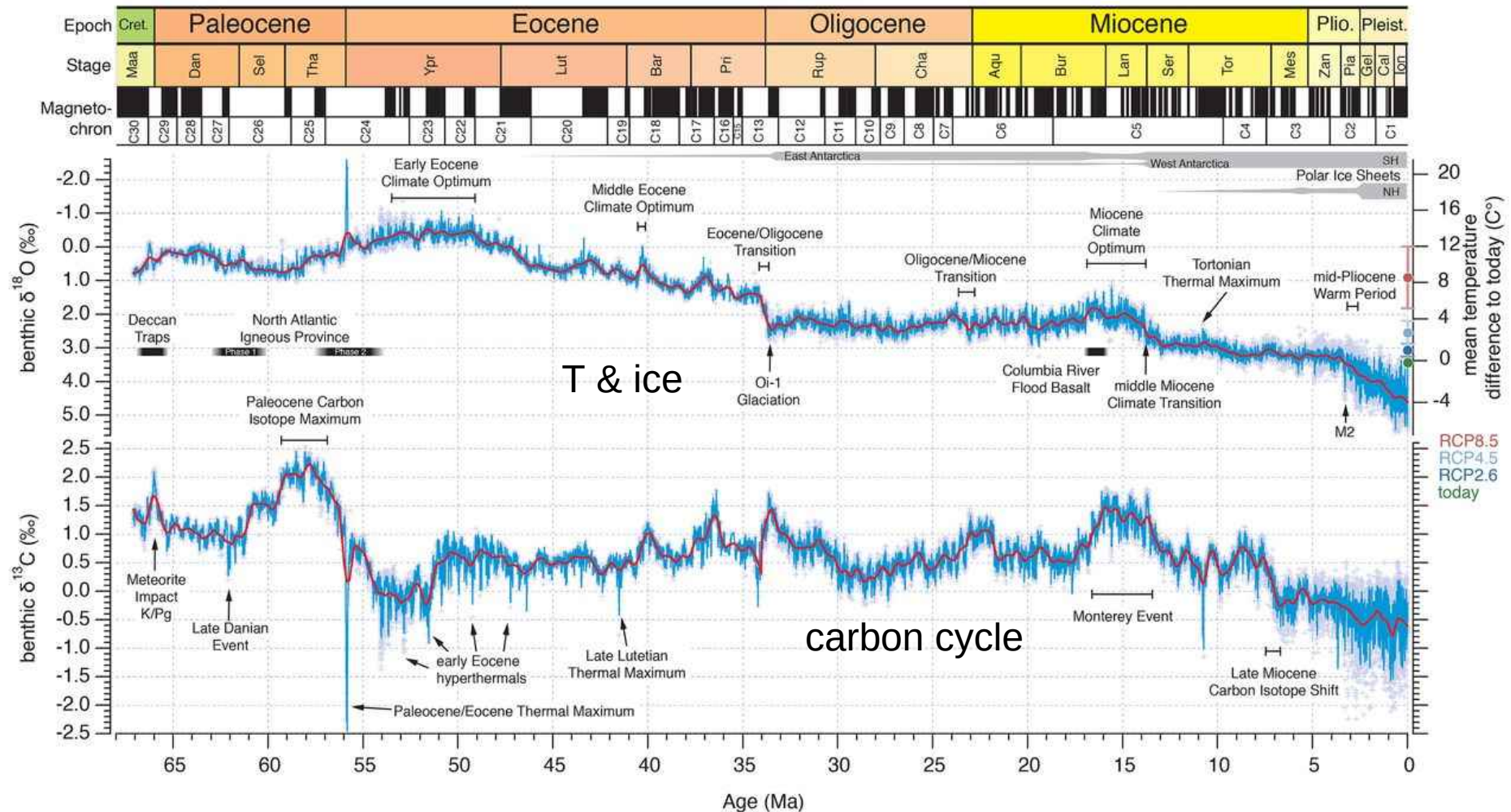
Carbon Isotopes

- C & O isotopes can be measured from CaCO_3
- C also in organics or gas
- ^{12}C – 98.9 %
 ^{13}C – 1.1 %
- photosynthesis discriminates against ^{13}C with $\alpha \sim 1.25$

Carbon Isotopes



Cenozoic Climate



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PETM

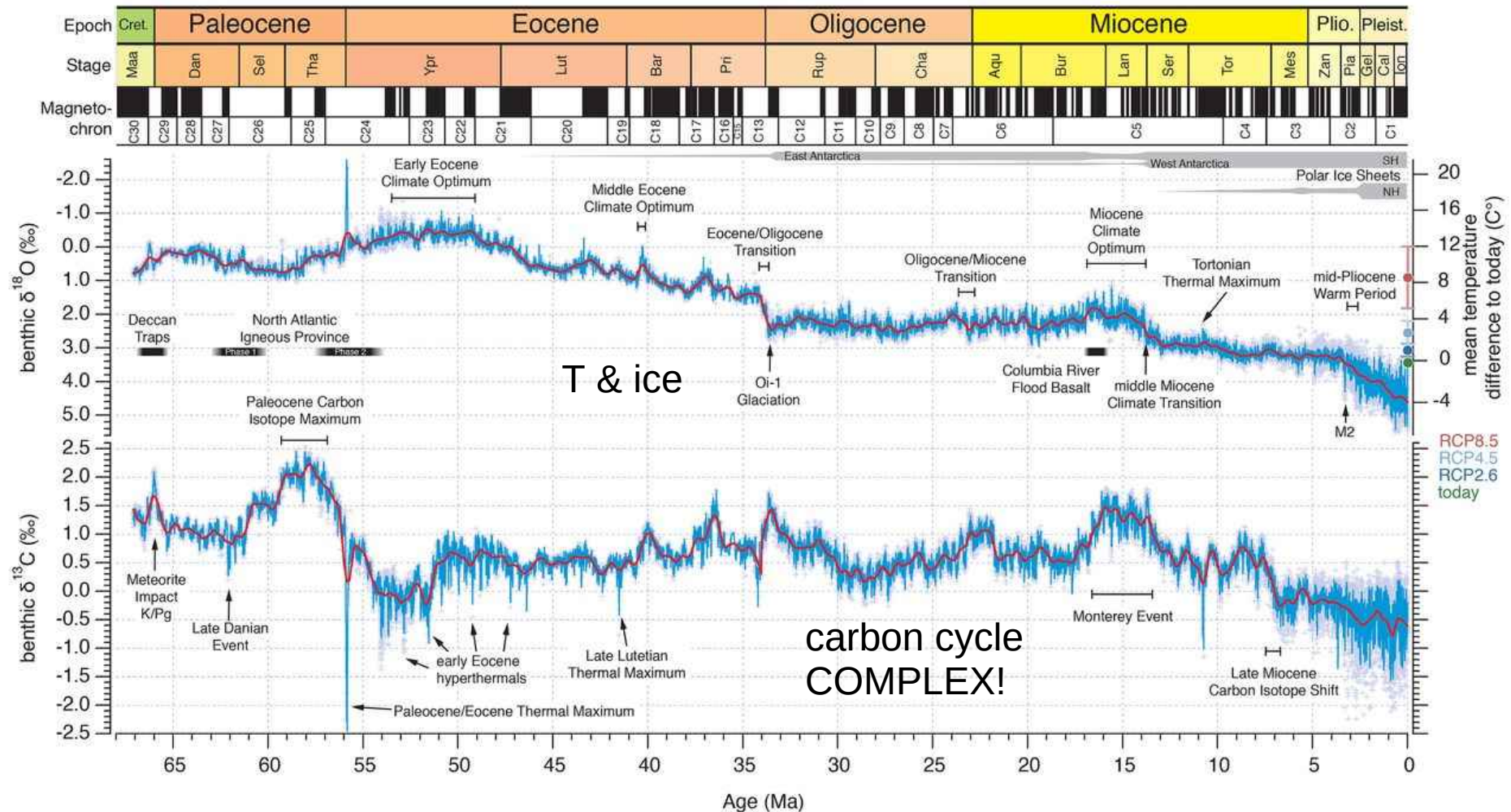
Paleocene–Eocene Thermal Maximum

caused by massive input of greenhouse gases (CO₂ and/or CH₄)

possible causes:

- submarine methane hydrates
- uplift and weathering of marine shelves
- warming-induced death of tropical plants (due to photorespiration)
- North Atlantic volcanism
- permafrost thaw

Cenozoic Climate



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Boron Isotopes

Boron Isotopes

boron species in seawater:

$\text{B}(\text{OH})_3$ und $\text{B}(\text{OH})_4^-$

^{10}B – 19.65 %

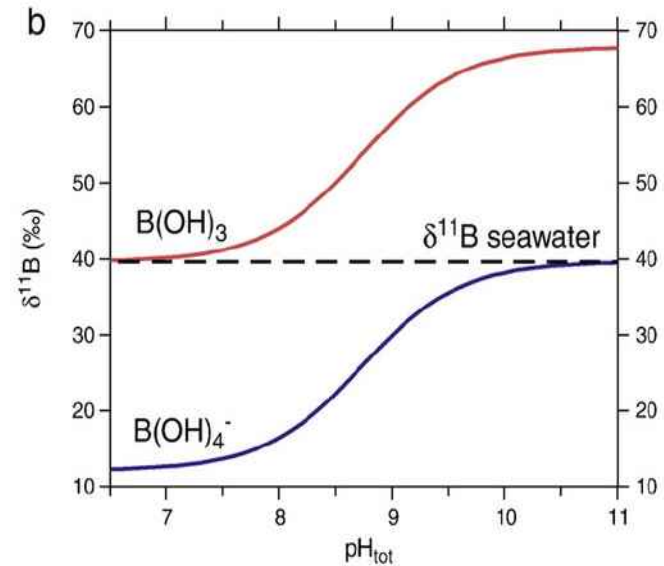
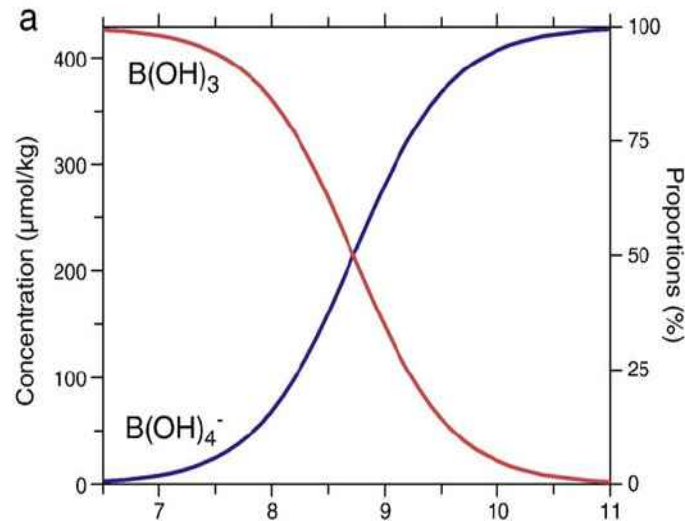
^{11}B – 80.35 %

$$\delta^{11}\text{B}(\text{‰}) = \left[\left(\frac{^{11}\text{B} / ^{10}\text{B}_{\text{sample}}}{^{11}\text{B} / ^{10}\text{B}_{\text{NIST951}}} \right) - 1 \right] \times 1000.$$

- B is fractionated between the two species
- $\text{B}(\text{OH})_4^-$ is built into the shells of foraminifera

Boron Isotopes

boron species in seawater:



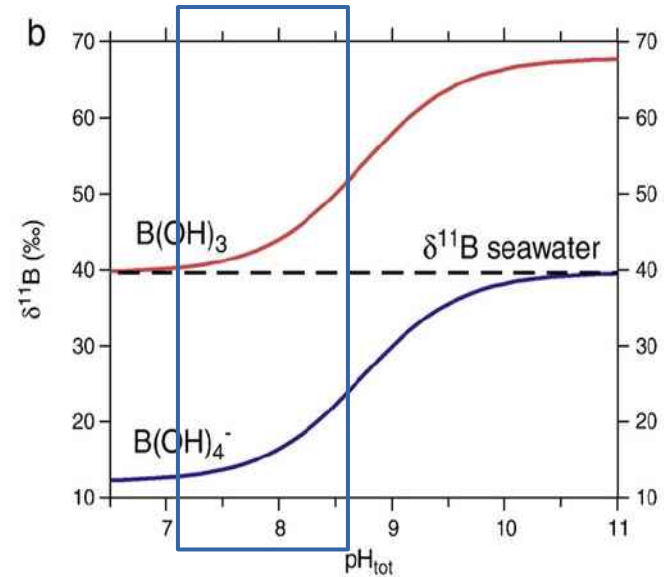
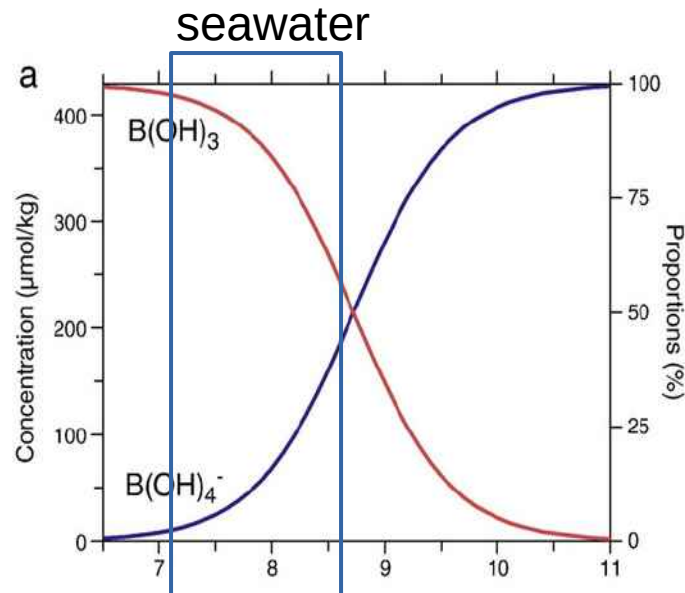
Rae et al. (2011)
Earth and Planetary Science Letters

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Boron Isotopes

boron species in seawater:



Rae et al. (2011)
Earth and Planetary Science Letters

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Boron Isotopes

boron isotopes in benthic foraminifera

Epifaunal

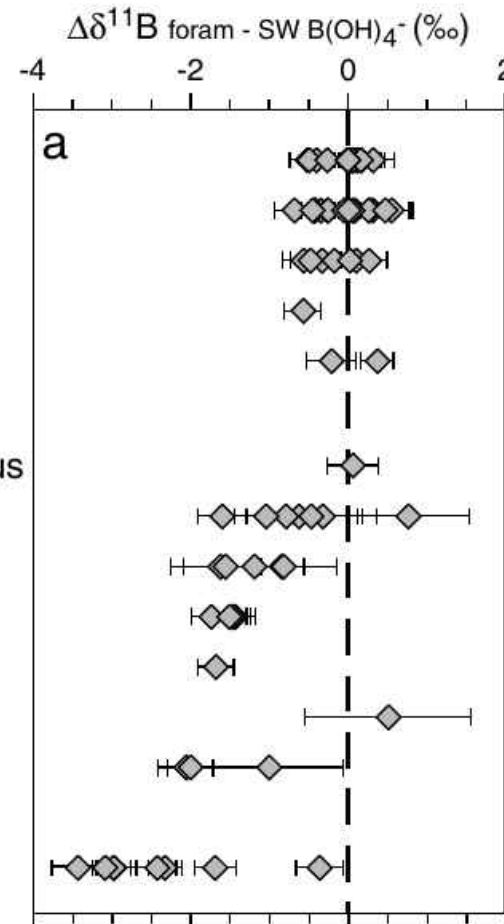
Cibicidoides wuellerstorfi
Cibicidoides mundulus
Planulina ariminensis
Cibicidoides lobatus
Cibicidoides ungerianus

Infaunal

Cibicidoides robertsonianus
Oridorsalis umbonatus
Gyroidina soldanii
Lenticulina vortex
Ammonia beccarii
Melonis zaandamae
Uvigerina peregrina

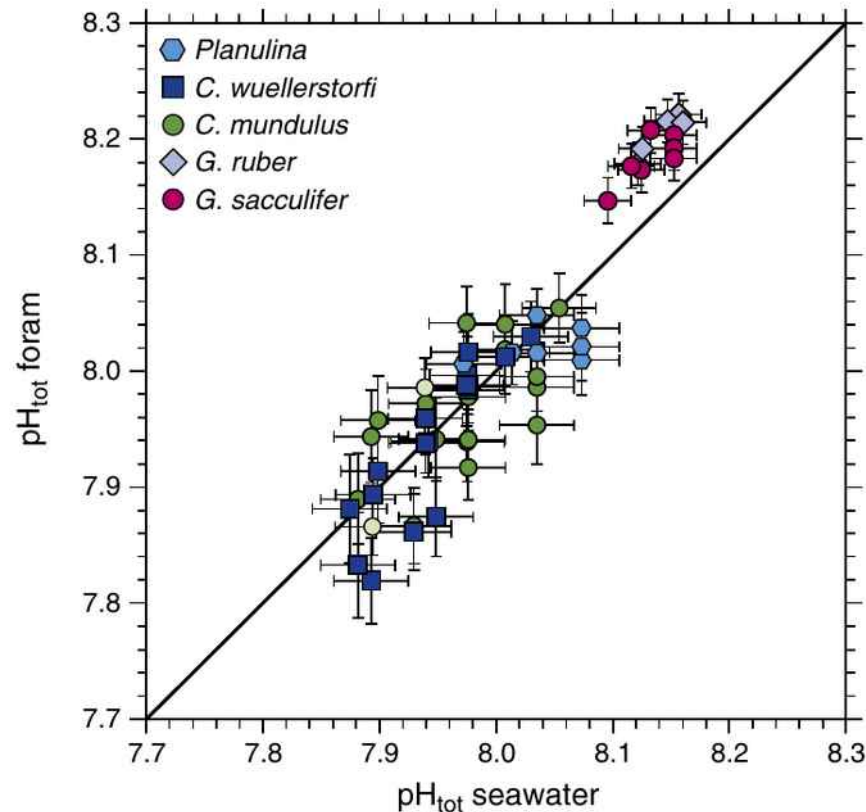
Aragonite

Hoeglundina elegans



Boron Isotopes

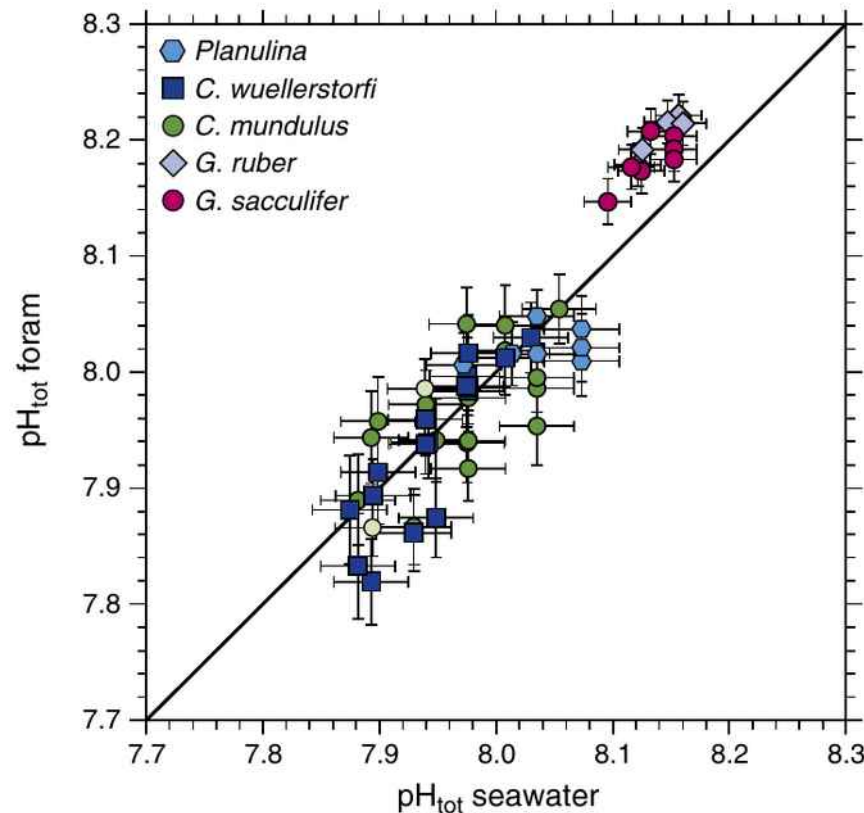
pH reconstructed from foraminifera



Rae et al. (2011)
Earth and Planetary Science Letters

Boron Isotopes

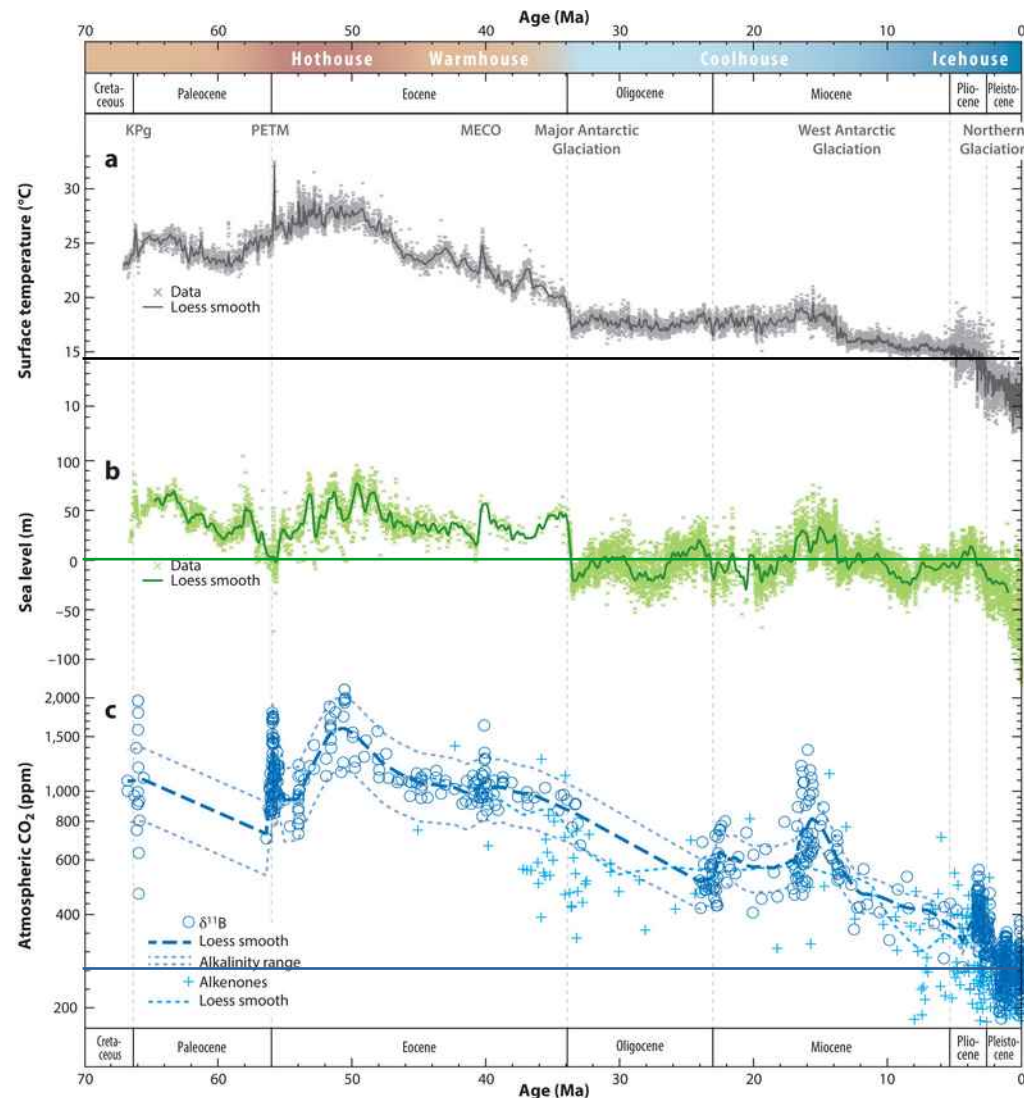
pH reconstructed from foraminifera



together with further assumptions, ocean pH traces (long term) atmospheric CO_2

Rae et al. (2011)
Earth and Planetary Science Letters

Cenozoic Climate

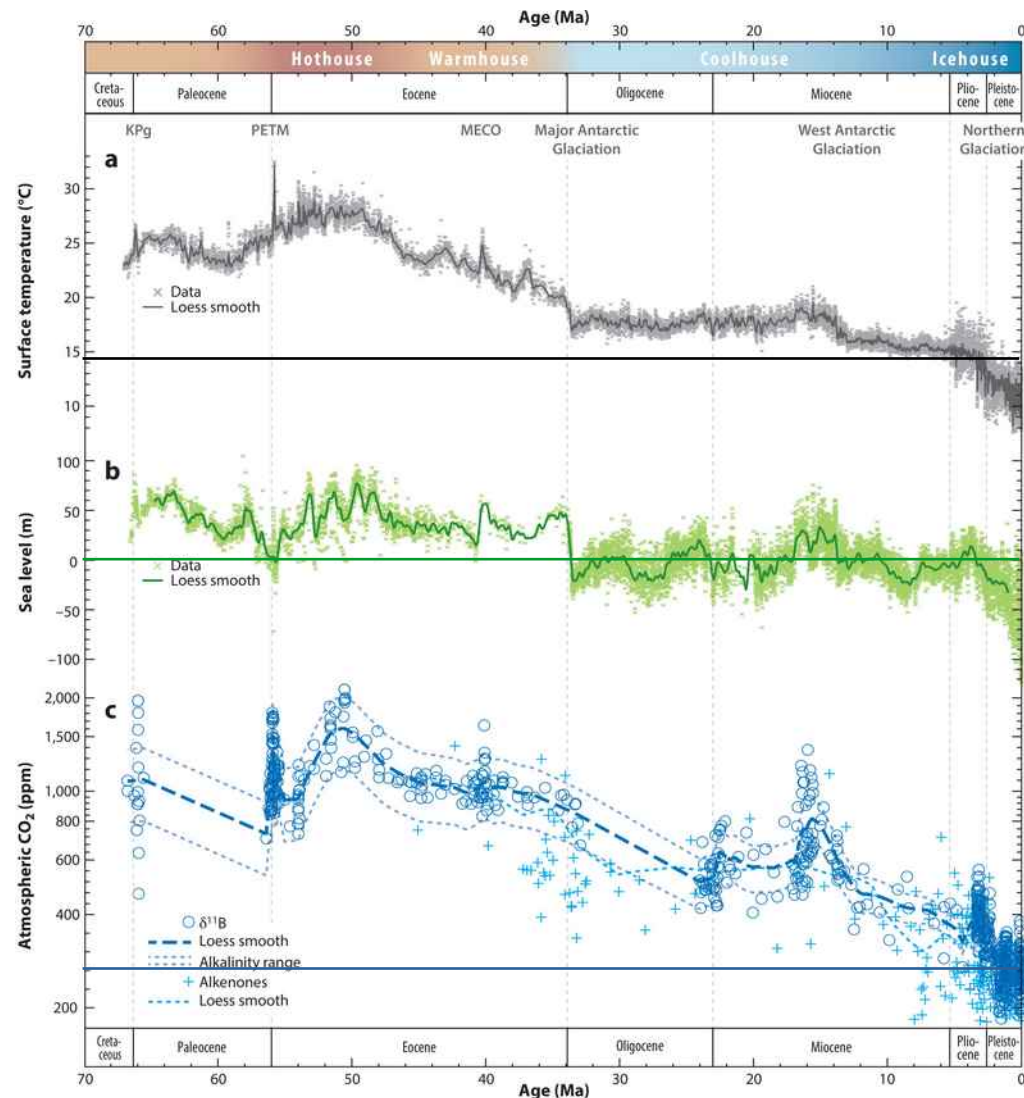


Rae JWB, et al. 2021
Annu. Rev. Earth Planet. Sci. 49:609–41

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Cenozoic Climate



Rae JWB, et al. 2021
Annu. Rev. Earth Planet. Sci. 49:609–41

T and CO₂
 parallel?

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Climate Sensitivity

How much does Earth warm with increasing CO₂?

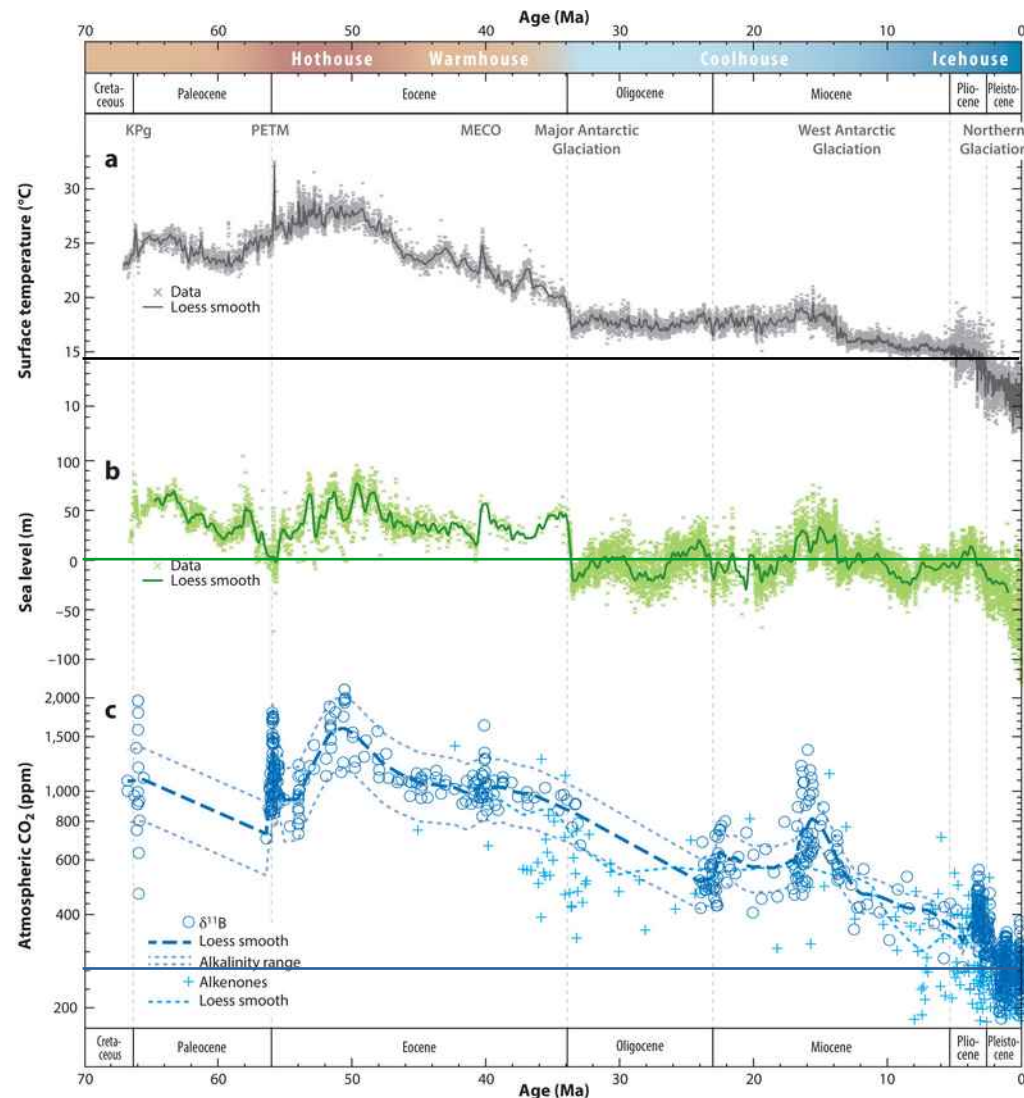
Equilibrium Climate Sensitivity (ECS)

- long-term, including geologic feedbacks
- usually referenced to doubling of CO₂

Transient Climate Response (TCR)

- short term (~ 20 years) climate response
- including fast feedbacks
- often used for models

Cenozoic Climate



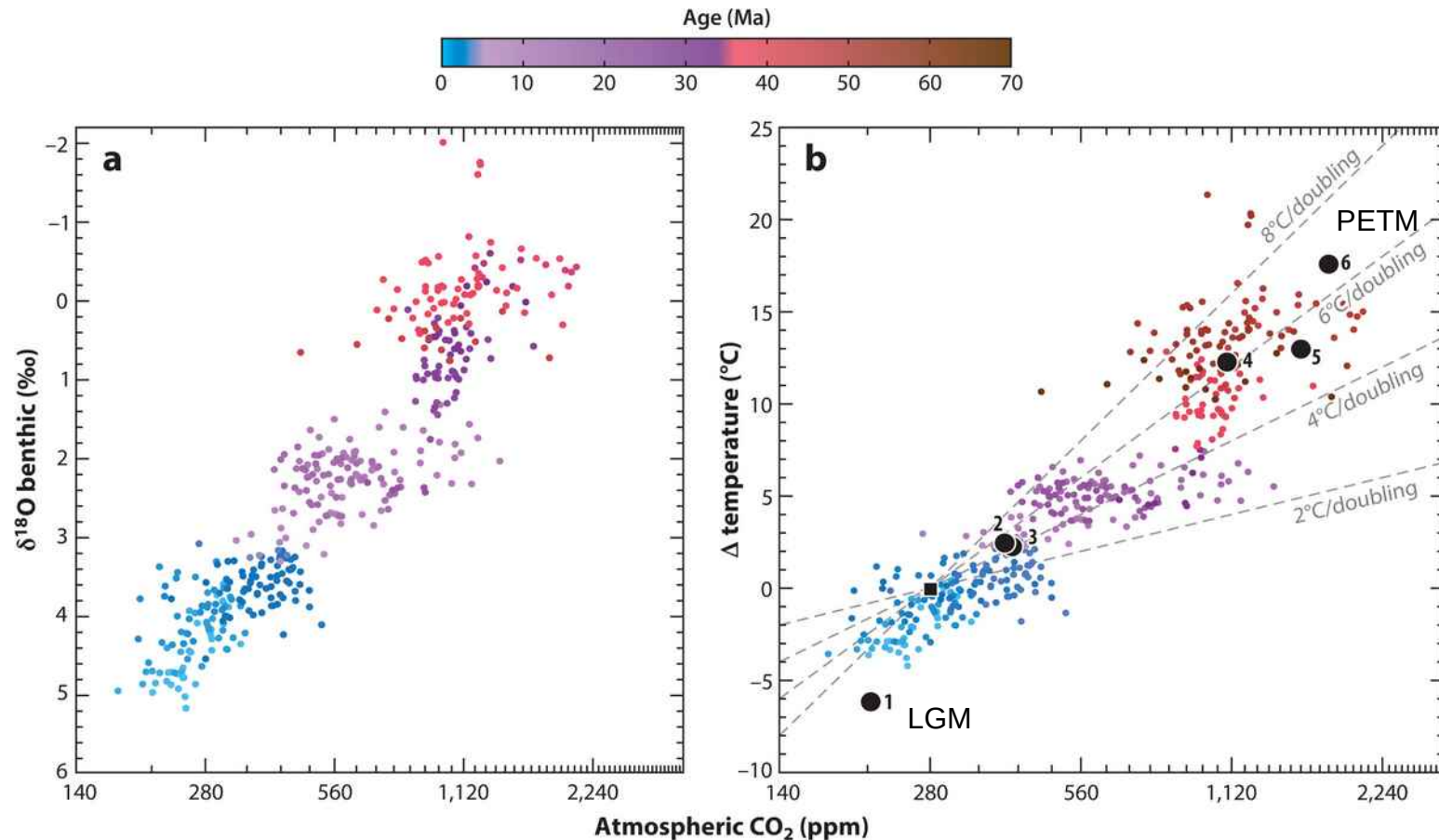
Rae JWB, et al. 2021
Annu. Rev. Earth Planet. Sci. 49:609–41

T and CO₂
 parallel?

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Cenozoic Climate Sensitivity



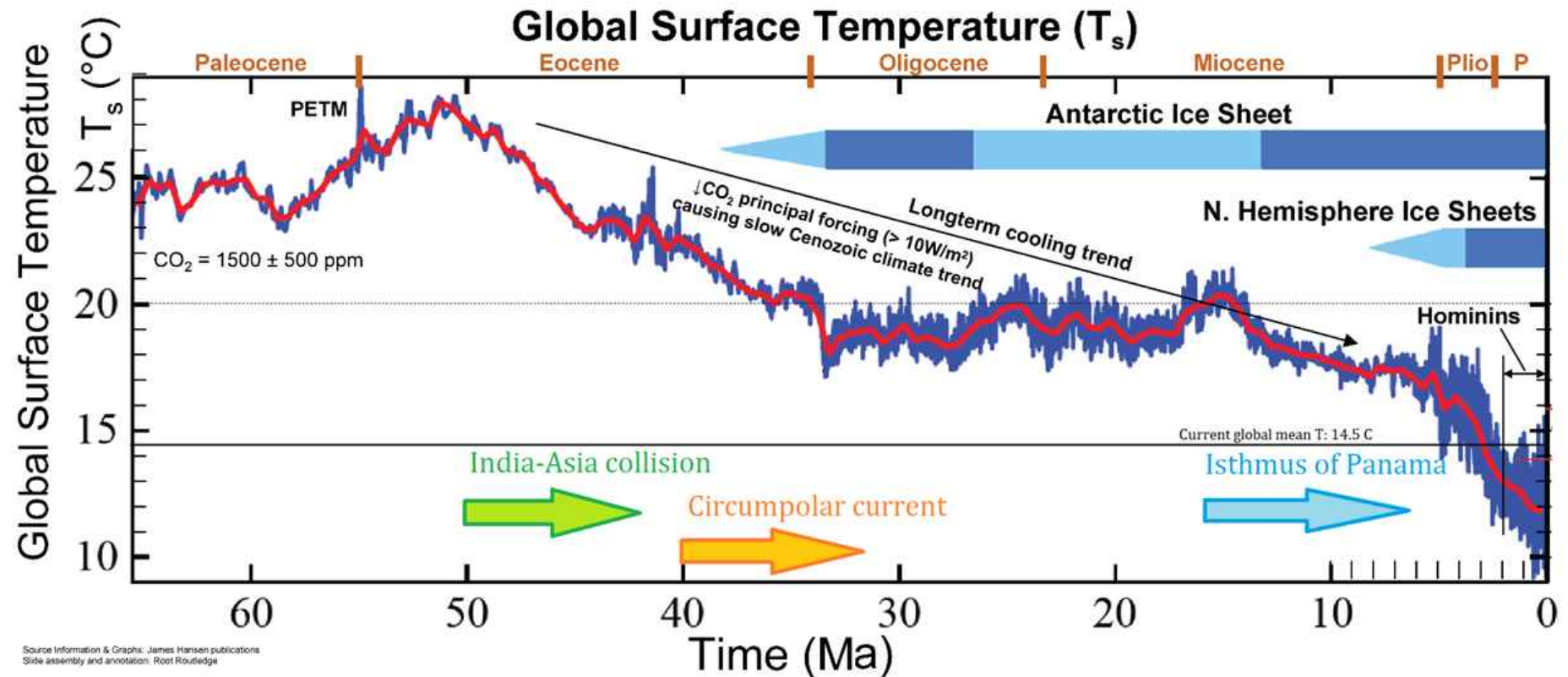
Rae JWB, et al. 2021
Annu. Rev. Earth Planet. Sci. 49:609–41

Black circles indicate independent proxy-derived estimates of surface temperature: ● Last Glacial Maximum (Tierney et al. 2020b), ● Pliocene (de la Vega et al. 2020), ● Pliocene (McClymont et al. 2020), ● late Paleocene, ● Early Eocene Climatic Optimum, and ● Paleocene-Eocene Thermal Maximum (Inglis et al. 2020). Dashed lines denote different degrees of temperature change per CO_2 doubling, providing an estimate of Earth system sensitivity.

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Cenozoic Climate



Earle (2016), opentextbc.ca
after James Hansen and Root Routledge

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Cenozoic cooling

Cenozoic cooling

Long term cooling trend from hot-house to ice-house

Causes debated and likely complex

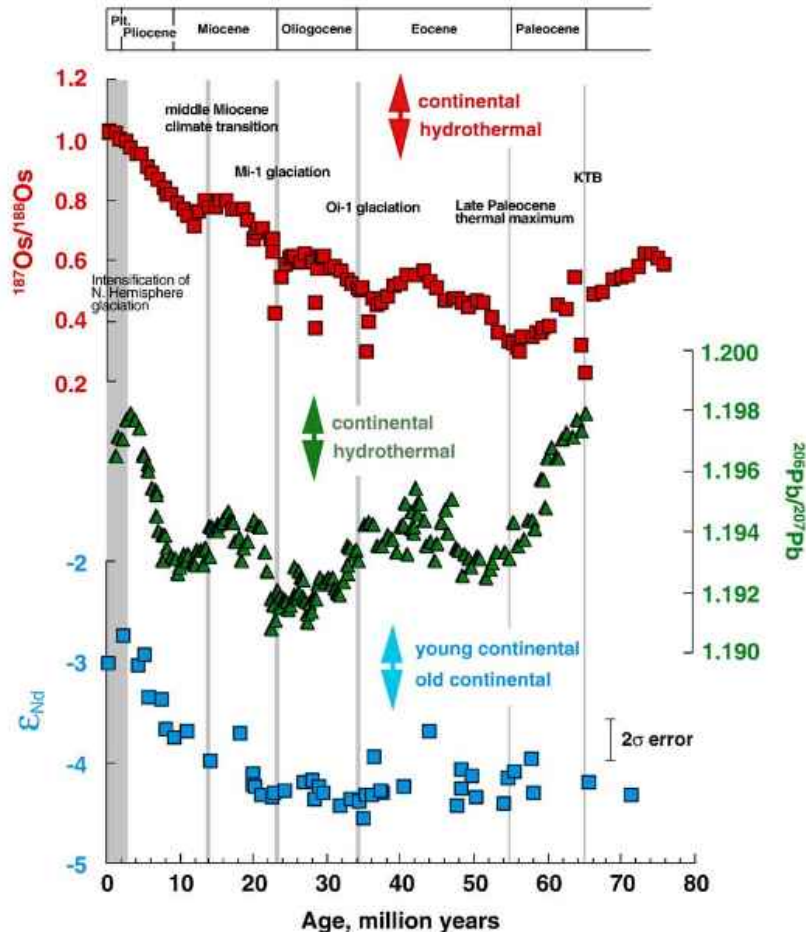
- weathering?
- isolation of Antarctica?
- faunal changes?

Cenozoic cooling

Long term cooling trend from hot-house to ice-house

Causes debatec

- weathering?



Burton (2006)
Journal of Geochemical
Exploration

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Cenozoic cooling

spread of extensive grass lands during Miocene
favoured e.g. by seasonal aridity

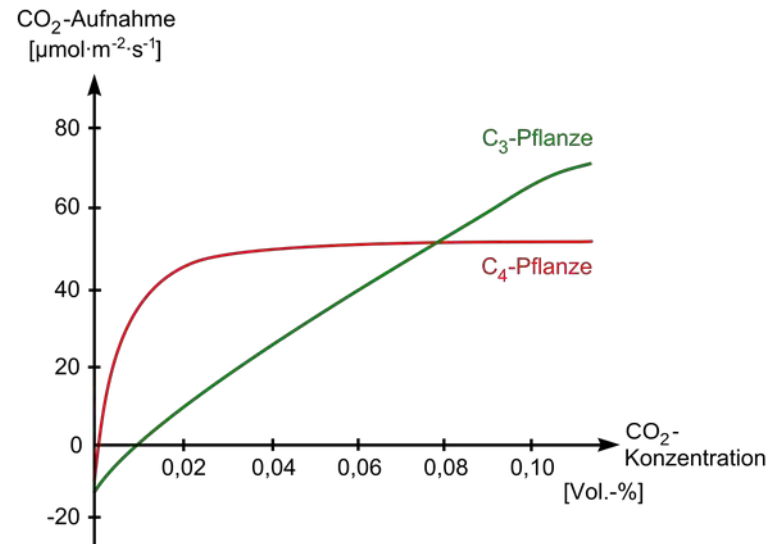
development of C4 photosynthesis at ~ 10 Ma

- developed multiple times
- fixes C in molecule containing 4 C atoms
- deals better with aridity and low CO₂

Cenozoic cooling

spread of extensive grass lands during Miocene
favoured e.g. by seasonal aridity

deals better with aridity and low CO_2



Wikipedia

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Cenozoic cooling

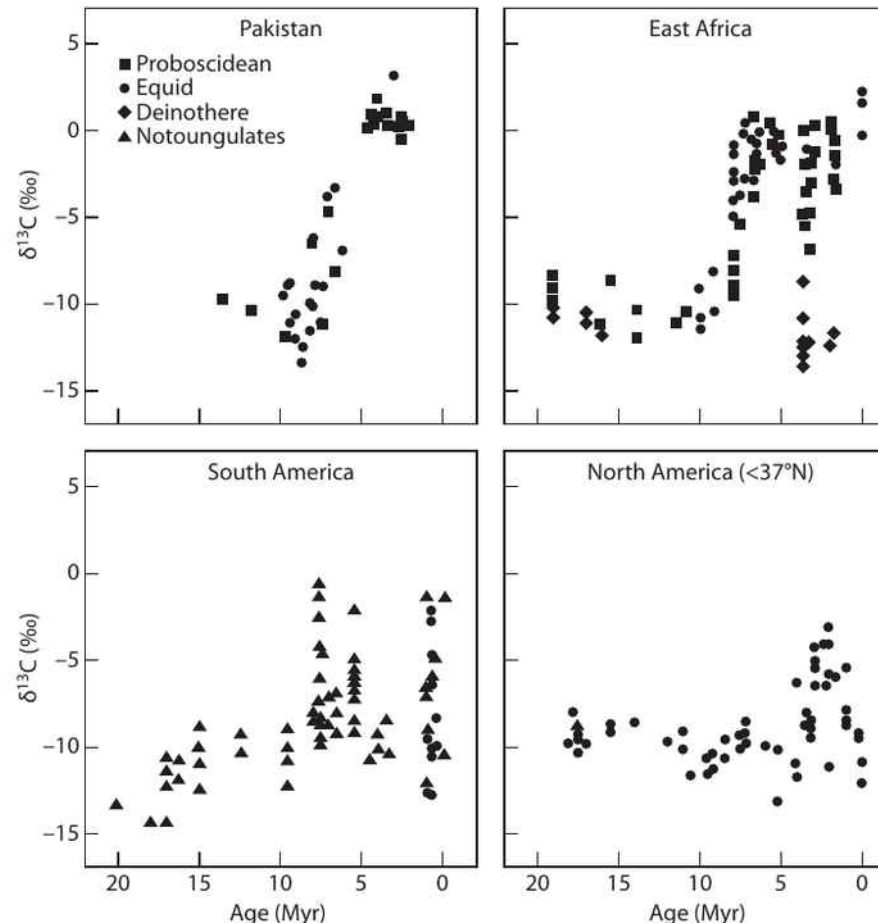
spread of extensive grass lands during Miocene
favoured e.g. by seasonal aridity

development of C4 photosynthesis at ~ 10 Ma

- developed multiple times
- fixes C in molecule containing 4 C atoms
- deals better with aridity and low CO₂
- today ~ 25% of plants, mostly grasses
- global food production depends on C4 plants
- fractionates ¹³C less than C3 plants

Cenozoic cooling

spread of extensive grass lands during Miocene



$\delta^{13}\text{C}$ in fossil teeth documents global spread of C4 grasses

Michael Bender
Paleoclimate

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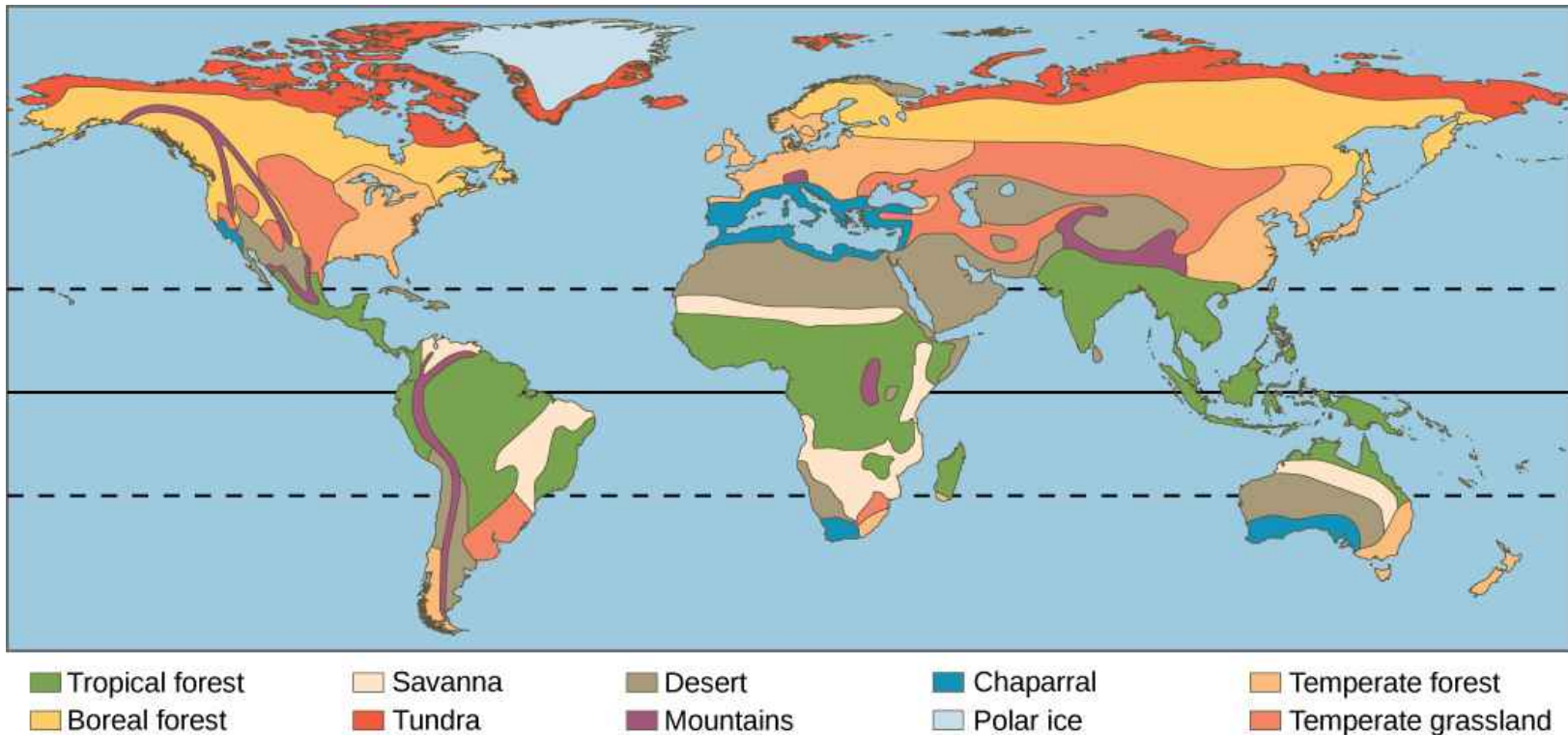
Cenozoic cooling

spread of extensive grass lands during
Miocene



lumenlearning.com
Environmental Biology

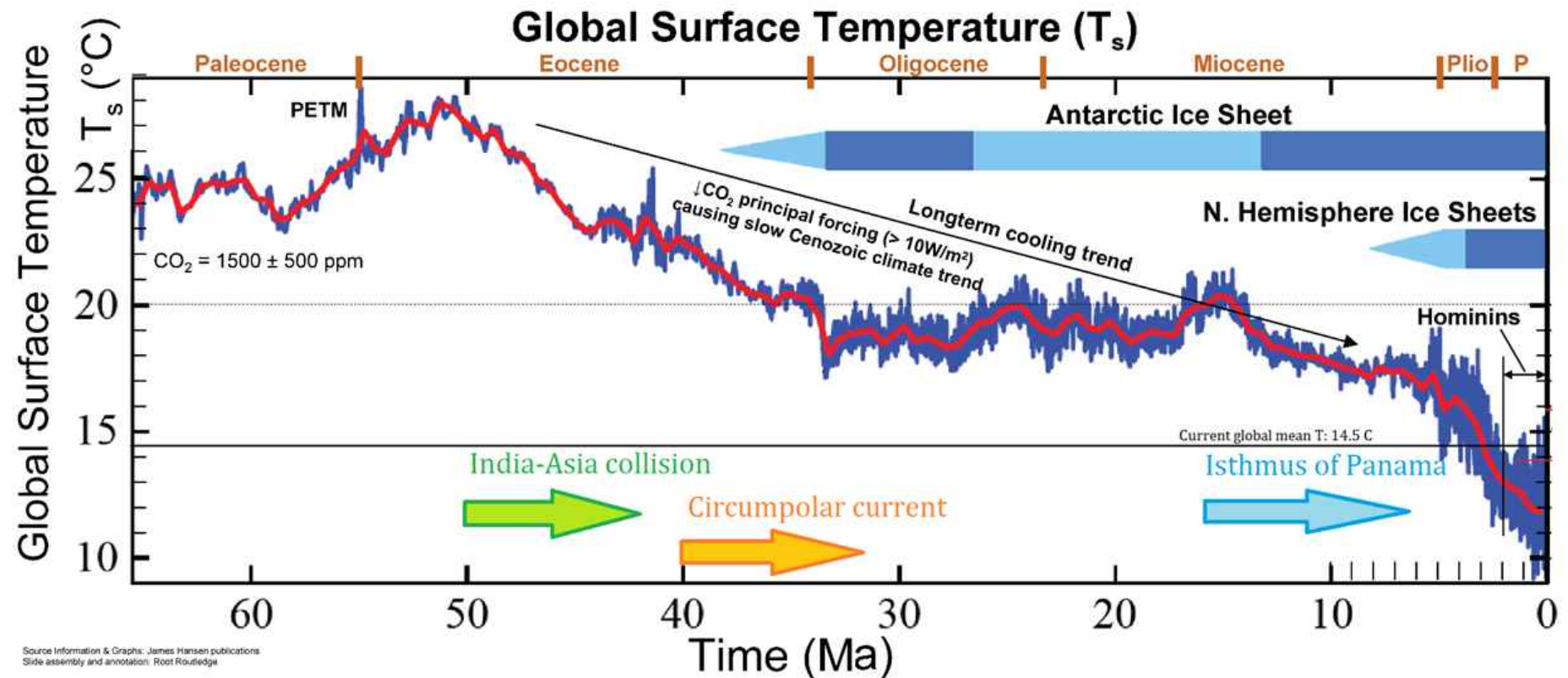
modern biomes



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Cenozoic Climate

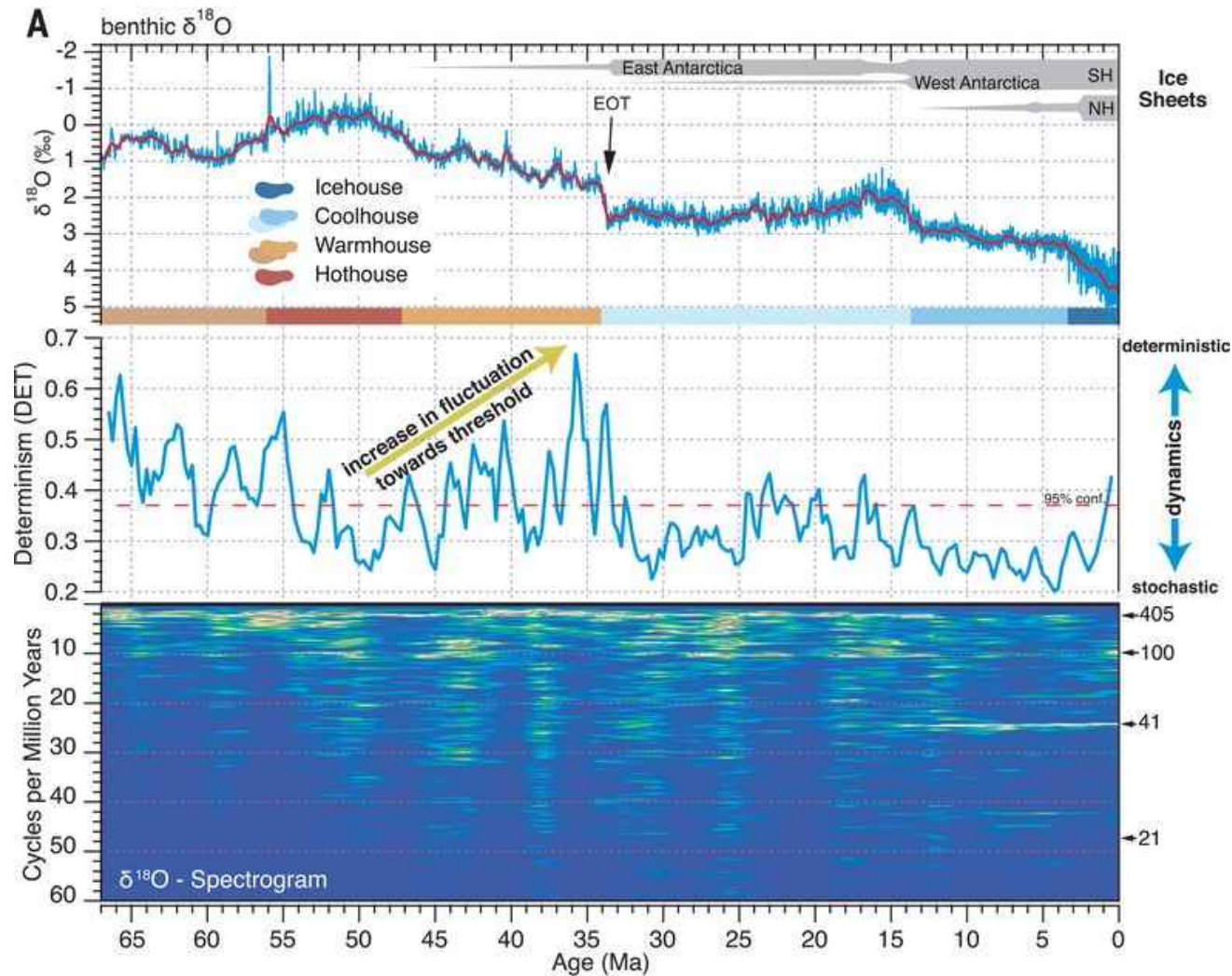


Earle (2016), opentextbc.ca
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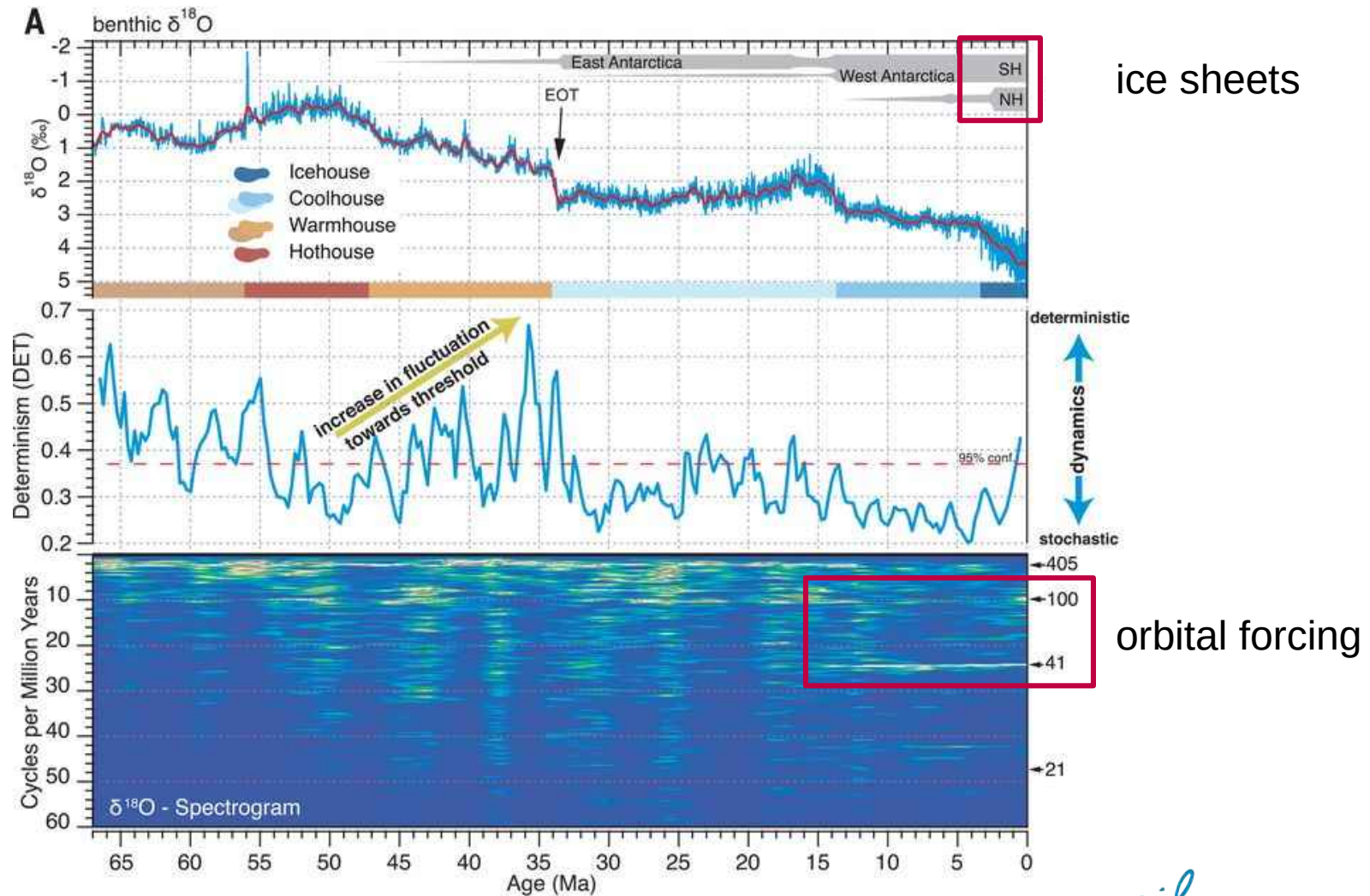
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Cenozoic Climate



Cenozoic Climate



Today's Summary

- Eocene Hothouse was very hot
- Equable climate led to warm poles
- PETM was extreme warm event caused by GHG
- Cenozoic climate dominated by CO₂
- Cooling was accompanied by CO₂ reduction and changes in weathering and fauna
- Temperature proxies: $\delta^{18}\text{O}$, $\Delta 47$, Mg/Ca, TEX86
- Carbon proxies: $\delta^{13}\text{C}$ & $\delta^{11}\text{B}$

Outlook

Today we finish 15 min early!

Monday	Introduction	Earth History
Tuesday	Proxies I	Cenozoic Hot & Warm House
Wednesday	Specific Climate System components	Pleistocene G-IG climate
Thursday	Proxies II & Climate System Interactions	Abrupt Climate Change
Friday	Current Climate Change	Future & Synthesis