## **Paleoclimate**



source: NASA

Introduction Paleoclim. Methods Planet Earth Earth History

### Yesterday's Summary

- Paleoclimatology is very interdisciplinary
- many different archives and proxies, but data patchy and often uncertain
- long term climate determined by: insolation, albedo, and greenhouse gases
- Early Earth climate has changed completely
- Life and Evolution have shaped Earth's chemistry



## **Lecture Progress**

#### Today we finish 15 min early!

Monday	Introduction	Earth History
Tuesday	Proxies I	Cenozoic Hot & Warm House
Wednesday	Specific Climate System components	Pleistocene G-IG climate
Thursday	Proxies II & Climate System Interactions	Abrupt Climate Change
Friday	Current Climate Change	Future & Synthesis

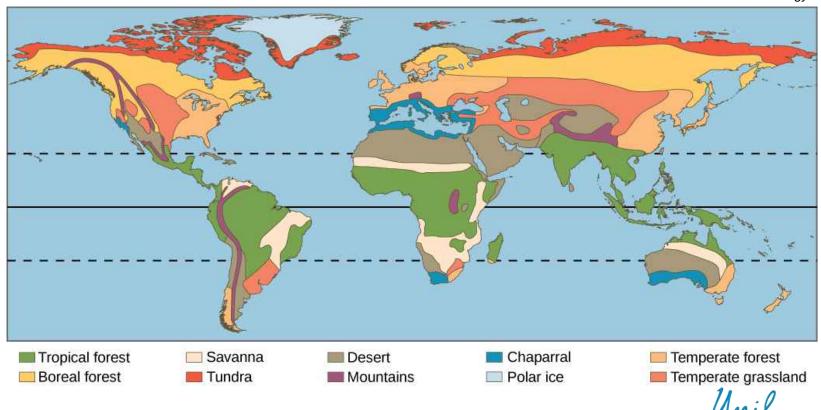
Mul.





#### modern biomes

lumenlearning.com Environmental Biology

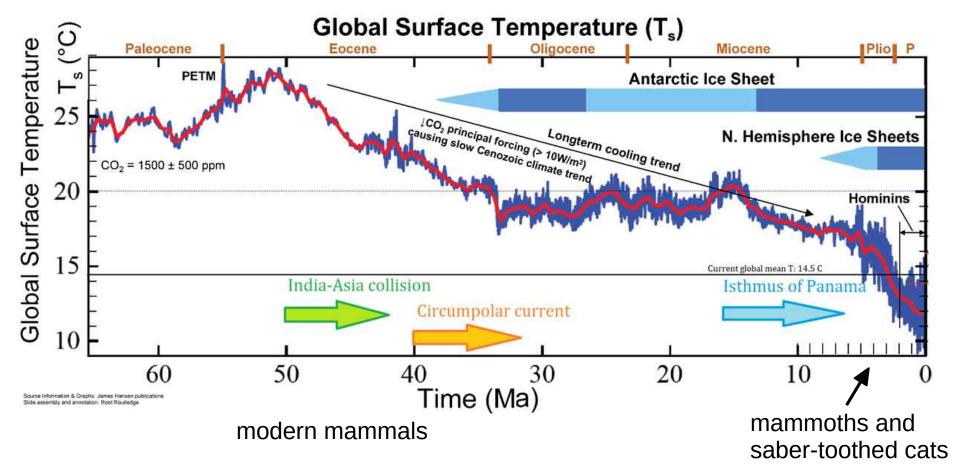


#### Day 2 : Overview

- Overview of Cenozoic Climate
  - Marine Isotope Records
- Basics of Isotope Geochemistry
  - Oxygen Isotopes in Paleoclimatology
  - Clumped C-O Isotopes
  - Mg/Ca paleothermometer
  - TEX86 paleothermometer
- Hothouse "Equable Climate"
  - PETM hyperthermal
- Climate Sensitivity
- Mid-Late Cenozoic cooling

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#### **Cenozoic Climate**

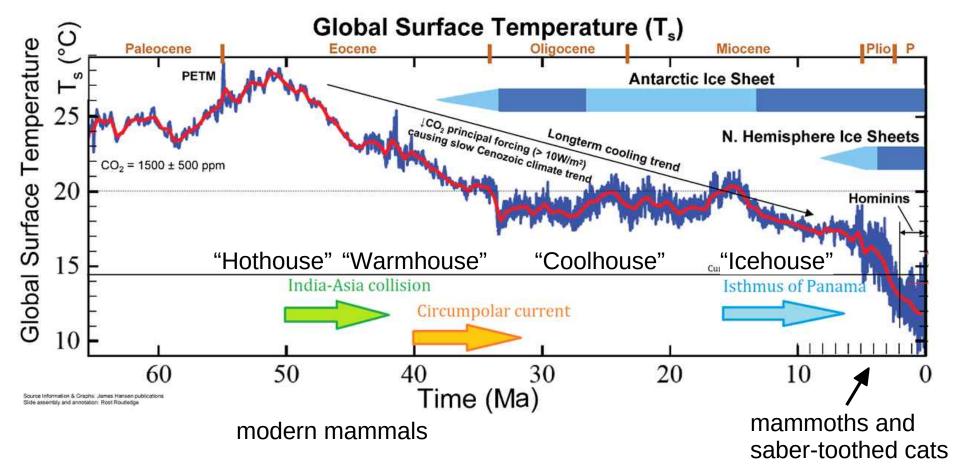


better records with plenty marine sediment cores

Earle (2016), opentextbc.ca after James Hansen and Root Routledge



#### **Cenozoic Climate**

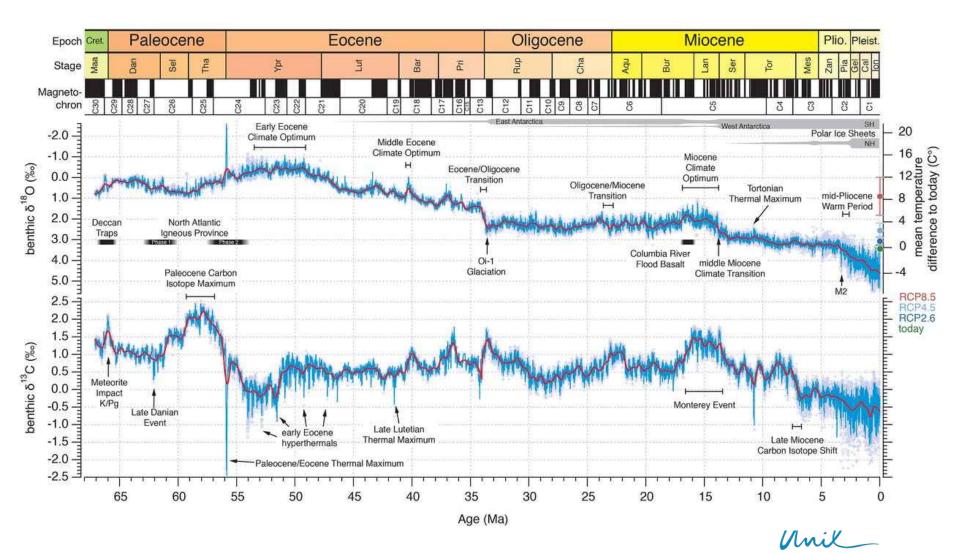


better records with plenty marine sediment cores

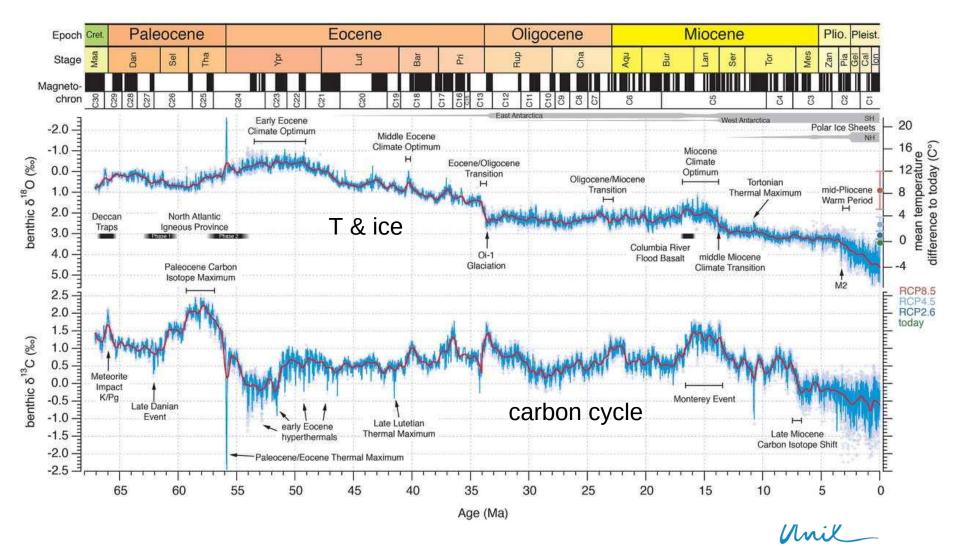
Earle (2016), opentextbc.ca after James Hansen and Root Routledge

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#### **Cenozoic Climate**



#### **Cenozoic Climate**



#### **Isotope Geochemistry**



# **Isotope Geochemistry**

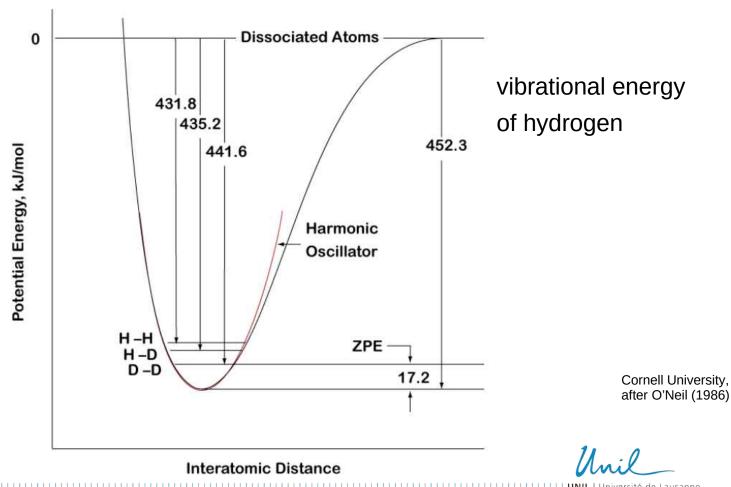
#### Isotope fractionation

$$\delta^{18}O = \left[ \frac{(^{18}O/^{16}O)_{sam} - (^{18}O/^{16}O)_{SMOW}}{(^{18}O/^{16}O)_{SMOW}} \right] \times 10^{3}$$

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# **Isotope Geochemistry**

#### Isotope fractionation



# **Isotope Geochemistry**

Isotope fractionation

 $R_x$  – isotope ratio in reservoir X

 $\alpha, \epsilon$  – fractionation factor

$$\alpha_{A-B} = R_A / R_B$$

e.g. 1.0098 for <sup>18</sup>O in evap. @ 20°C

$$\varepsilon = (\alpha - 1) * 10^3$$

$$\varepsilon_{A-B} \sim \delta_A - \delta_B$$

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# **Isotope Geochemistry**

Isotope fractionation

 $R_x$  – isotope ratio in reservoir X  $\alpha, \epsilon$  – fractionation factor

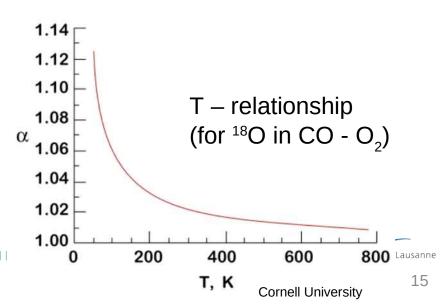
$$\alpha_{A-B} = R_A / R_B$$

e.g. 1.0098 for <sup>18</sup>O in evap. @ 20°C

$$\varepsilon = (\alpha - 1) * 10^{3}$$

$$\varepsilon_{A-B} \sim \delta_{A} - \delta_{B}$$

$$\alpha_{\text{A-B}} \sim 1/T^2$$



# **Isotope Geochemistry**

#### Equilibrium fractionation:

- slow complete equilibration (e.g. condensation @ 100% humidity)
- heavier isotopes enriched in colder phase

#### Kinetic fractionation:

- fast or incomplete reactions (e.g. condensation @ < 100% humidity or with immediate rain)
- unidirectional
- more complex

## **Isotope Geochemistry**

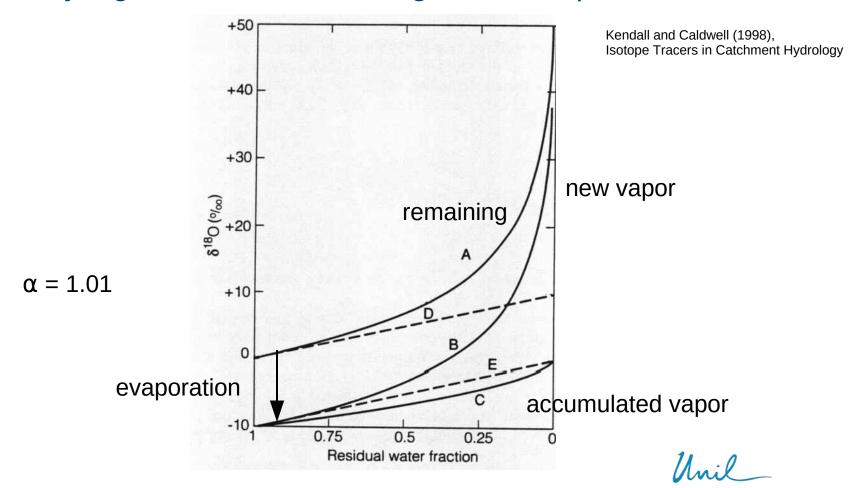
#### Rayleigh fractionation:

- equilibrium fractionation with removal of product
- reservoir decreases in size
- e.g.: raining clouds

$$R = R_0 f^{(\alpha-1)}$$
 (f = fraction remaining)

# **Isotope Geochemistry**

#### Rayleigh fractionation during water evaporation



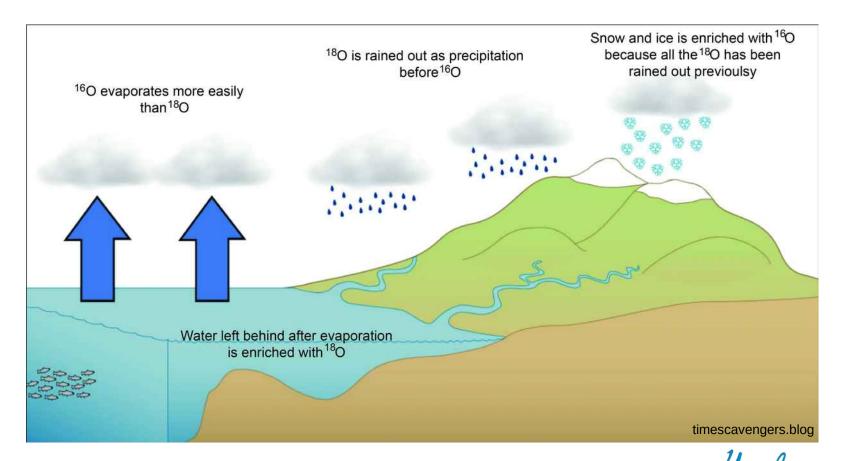
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#### **Oxygen Isotope Geochemistry**



# **Oxygen Isotope Geochemistry**

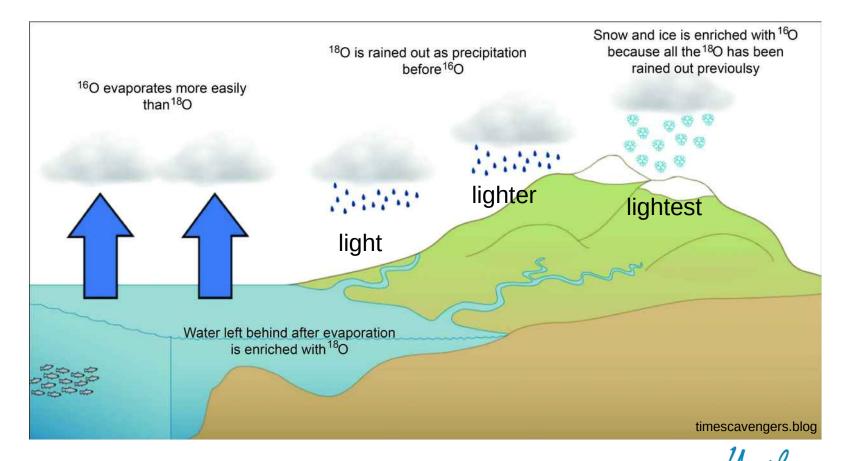
#### Rayleigh fractionation during water evaporation



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# **Oxygen Isotope Geochemistry**

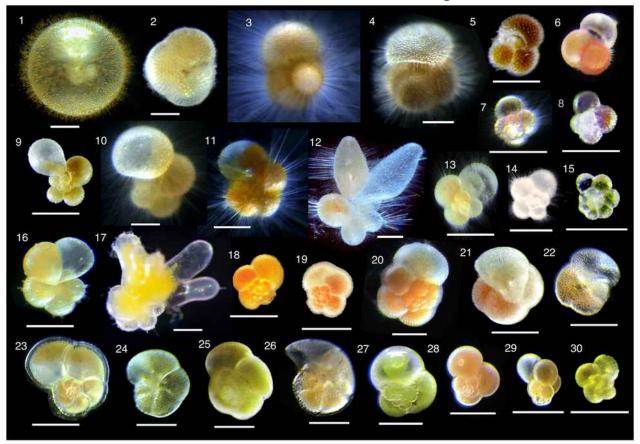
#### Rayleigh fractionation during water evaporation



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## Oxygen Isotopes in Carbonate

foraminifera – protists with CaCO<sub>3</sub> shells





### Oxygen Isotopes in Carbonate

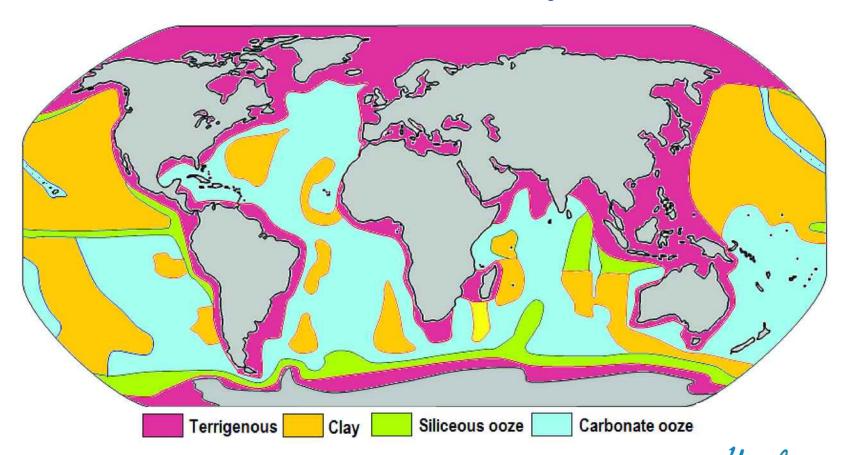
foraminifera – protists with CaCO<sub>3</sub> shells



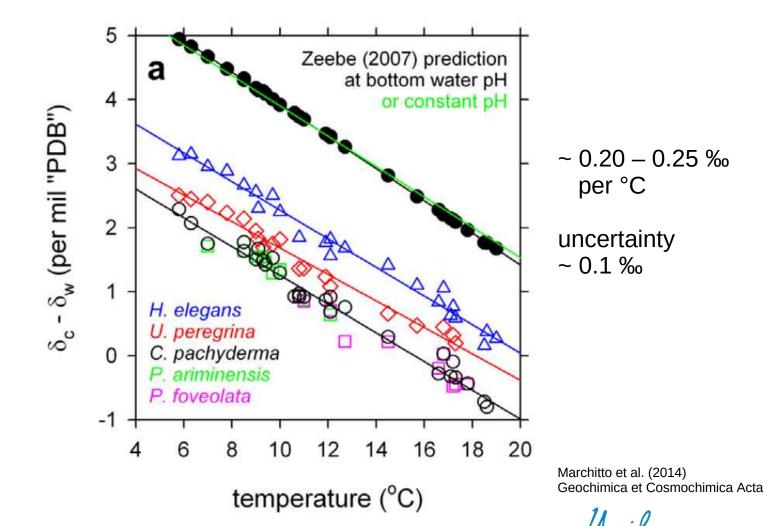


## Oxygen Isotopes in Carbonate

foraminifera – protists with CaCO<sub>3</sub> shells



#### Oxygen Isotopes in Carbonate

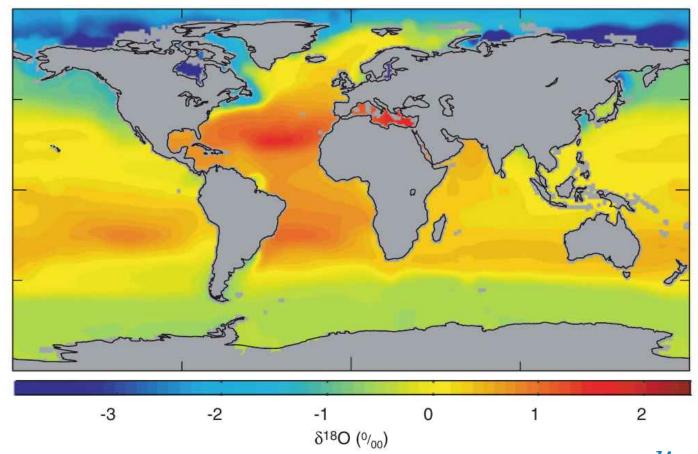


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Isotopes Cen. Cooling Cenozoic Clim. Hothouse Climate Sens.

# Oxygen Isotopes in the ocean

 $\delta^{18}$ O in seawater: ~ salinity



26

### Oxygen Isotopes in the ocean

 $\delta^{18}$ O in seawater: ~ salinity

 $\delta^{18}$ O in carbonate: ~ salinity & temperature

→ ~ density

### Oxygen Isotopes in the ocean

δ<sup>18</sup>O in seawater: ~ salinity

 $\delta^{18}$ O in carbonate: ~ salinity & temperature

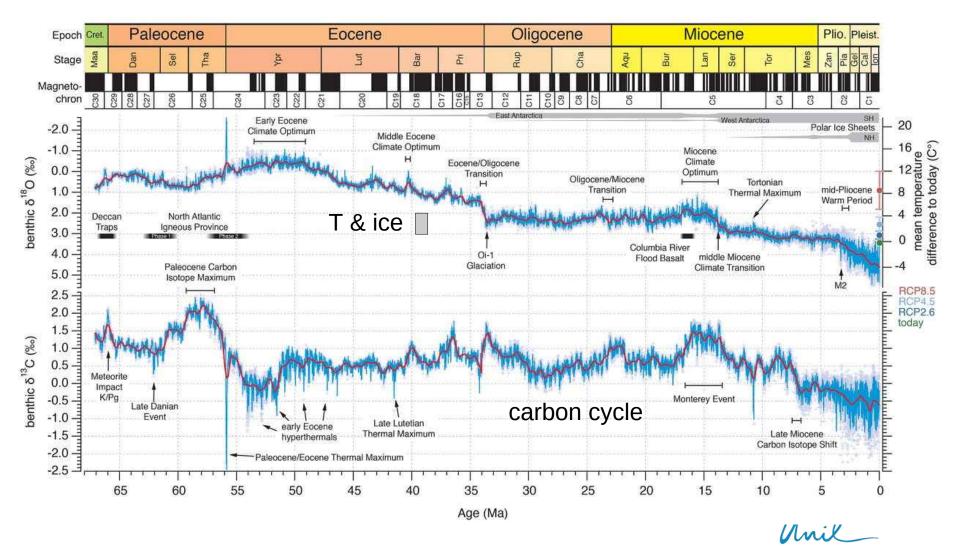
→ ~ density

BUT salinity in paleoceanography depends on global ice volume!

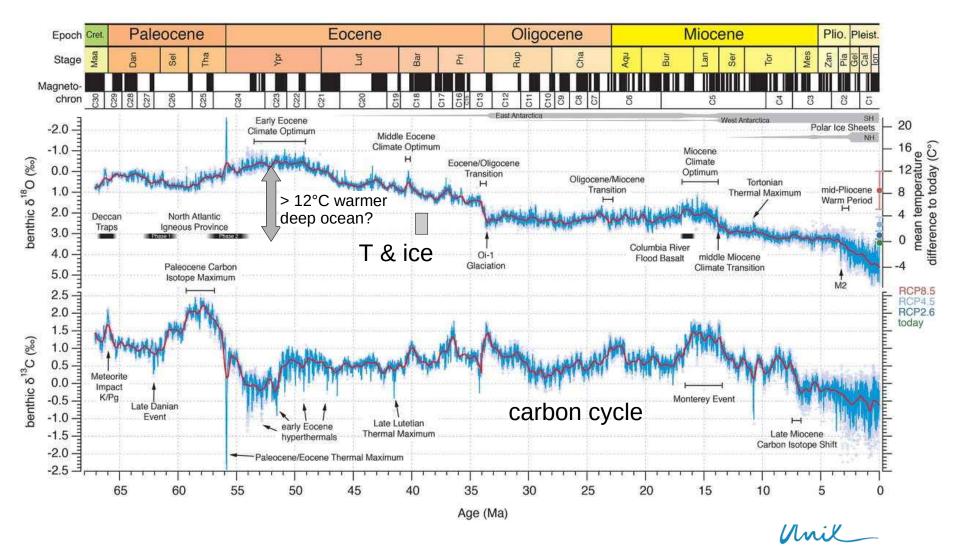
~ +1 ‰ per 100m sea level as ice Today's continental ice ~ 0.6 ‰ Last Glacial Maximum ice ~ + 1 ‰



#### **Cenozoic Climate**



#### **Cenozoic Climate**



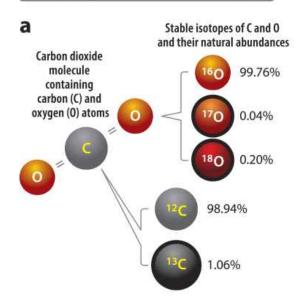
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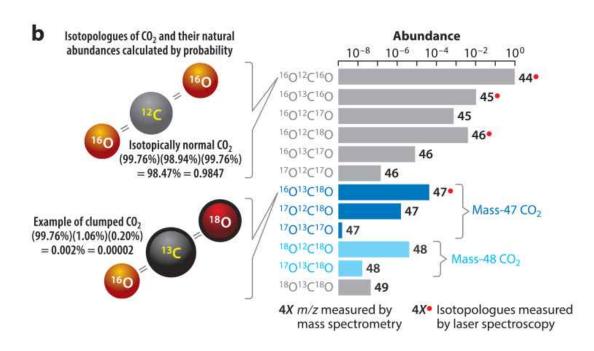
### **Clumped Isotope Thermometer**



### **Clumped Isotopes**

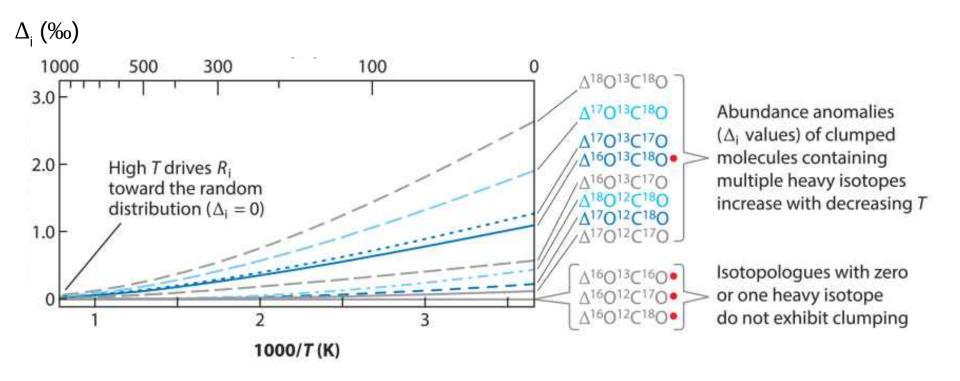
#### Clumping in carbon dioxide (CO<sub>2</sub>)



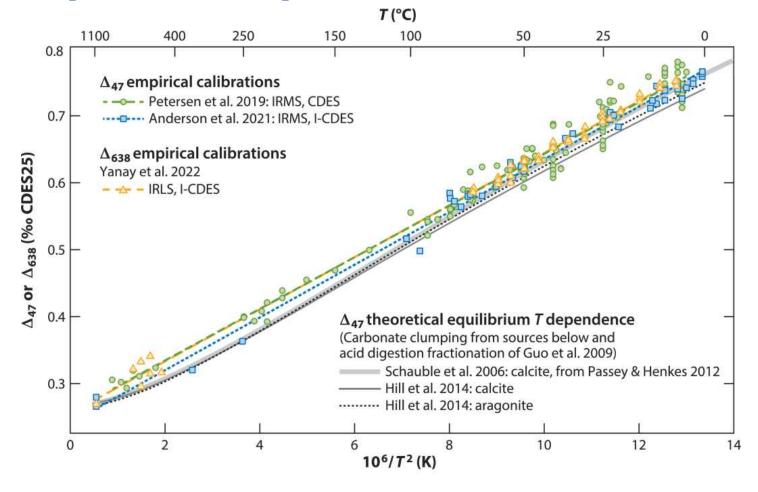


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### **Clumped Isotopes**



#### **Clumped Isotopes**



Huntington KW, Petersen SV. 2023 Annu. Rev. Earth Planet. Sci. 51:611–41

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## **Clumped Isotopes**

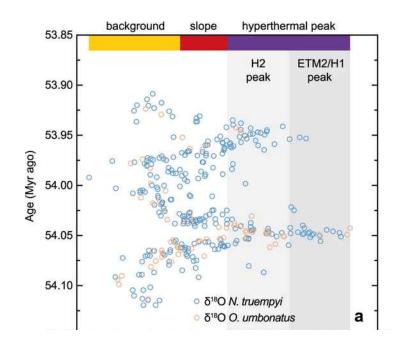
- very accurate
- few secondary effects

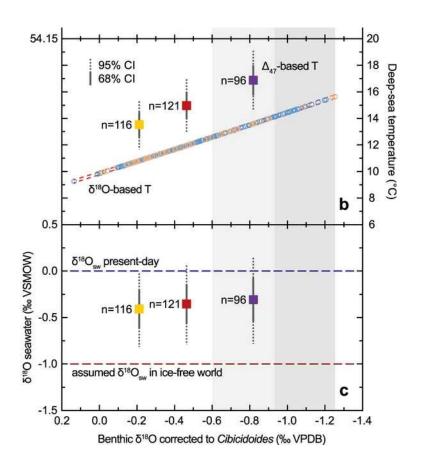
#### **BUT**

- low precision → many replicates
- large samples (few mg)
  - → costly sample analysis

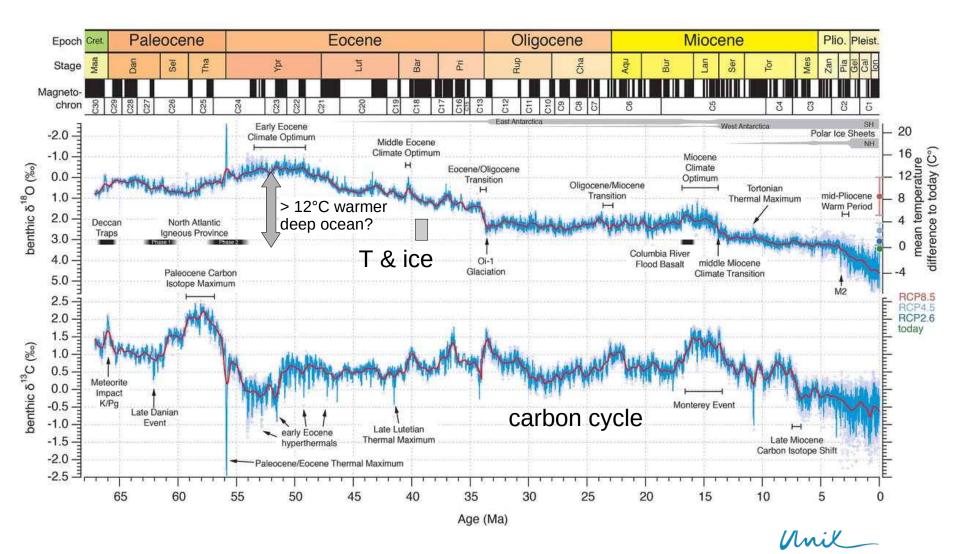


#### Clumped isotope temperatures





### **Cenozoic Climate**

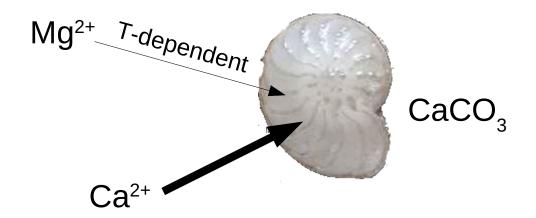


### Mg/Ca paleothermometer



### Mg/Ca paleothermometer

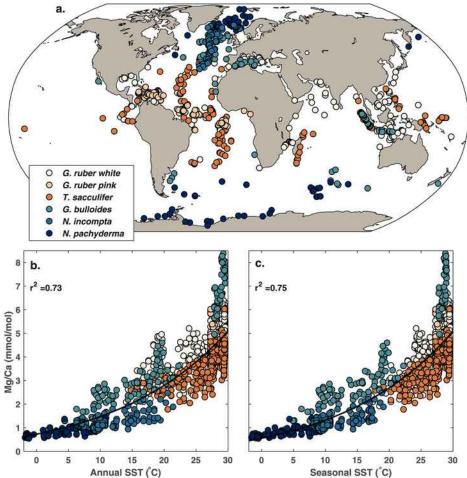
#### Mg/Ca in foraminifera





# Mg/Ca paleothermometer

Mg/Ca in foraminifera

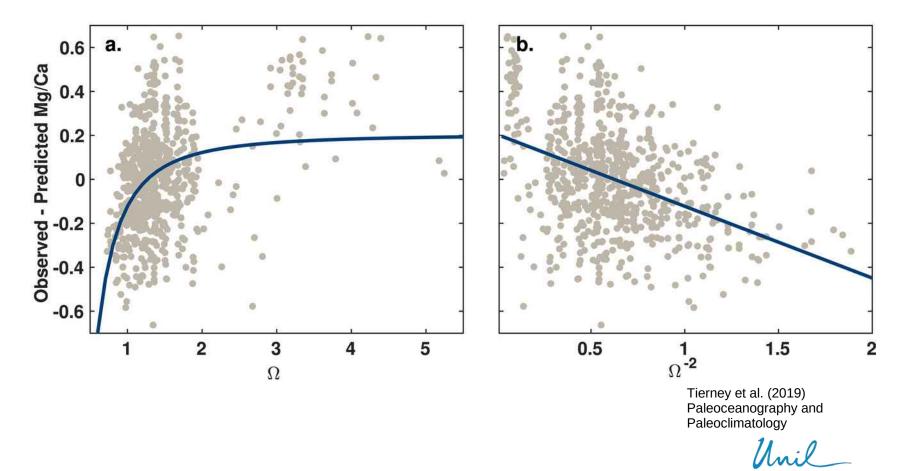


Tierney et al. (2019) Paleoceanography and Paleoclimatology

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## Mg/Ca paleothermometer

#### Mg/Ca in foraminifera



Hothouse Cenozoic Clim. Isotopes Climate Sens. Cen. Cooling

## Mg/Ca paleothermometer

#### Mg/Ca in foraminifera

- records past water temperature
- species-specific calibrations
- Mg/Ca in seawater dependent
- $\Omega$  dependent
- precision typically ~ 1°C



### **TEX86** paleothermometer



### **TEX86** paleothermometer

#### TEX86 proxy for SST

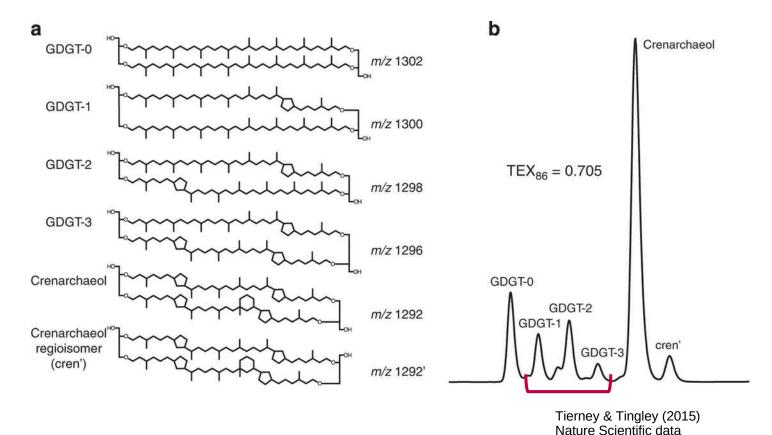
- based on specific organic molecules from sediments formed by archaea (similar to bacteria)
- the abundance ratio of certain molecules depends on ambient seawater temperatures
- mainly records near-sea surface T



Cenozoic Clim. Hothouse Climate Sens. Isotopes Cen. Cooling

### **TEX86** paleothermometer

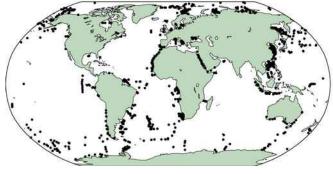
#### TEX86 proxy for SST

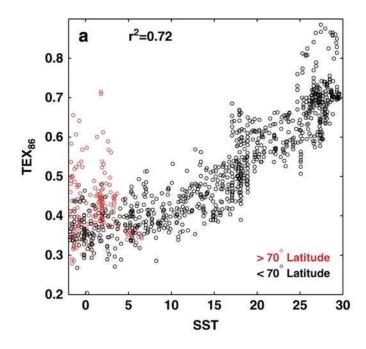


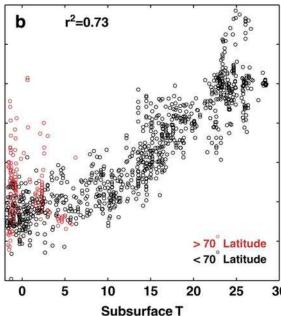
GDGT = Glycerol dialkyl glycerol tetraether (lipids)

**TEX86** paleothermometer

TEX86 proxy for SST







Tierney & Tingley (2015) Nature Scientific data

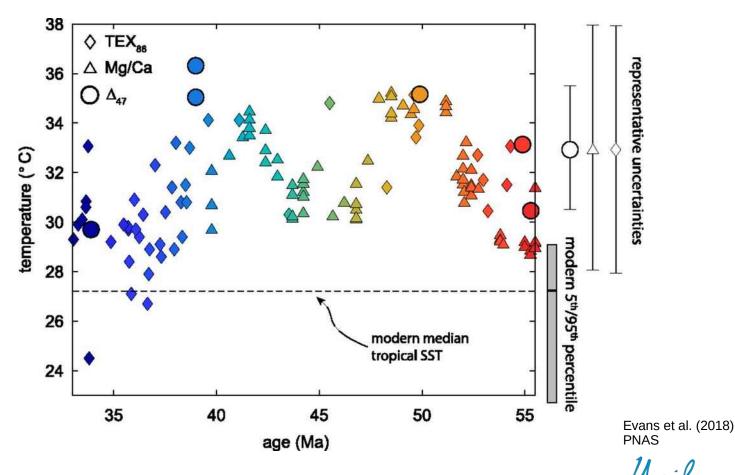
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### **Hothouse Climate**



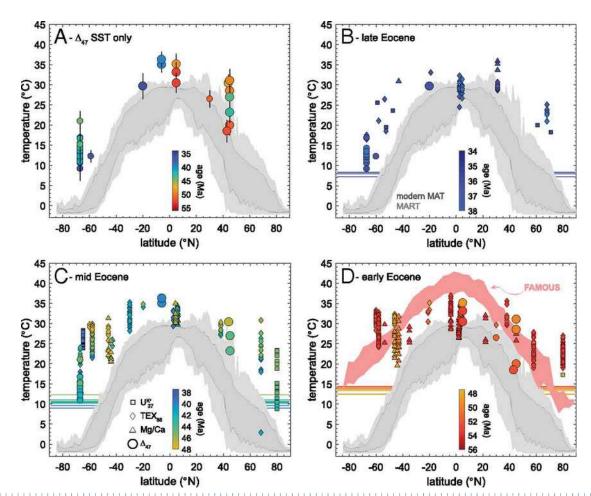
### **Hothouse Climate**

#### Eocene low latitude sea surface T



## **Equable Climate**

#### Eocene latitudinal sea surface T

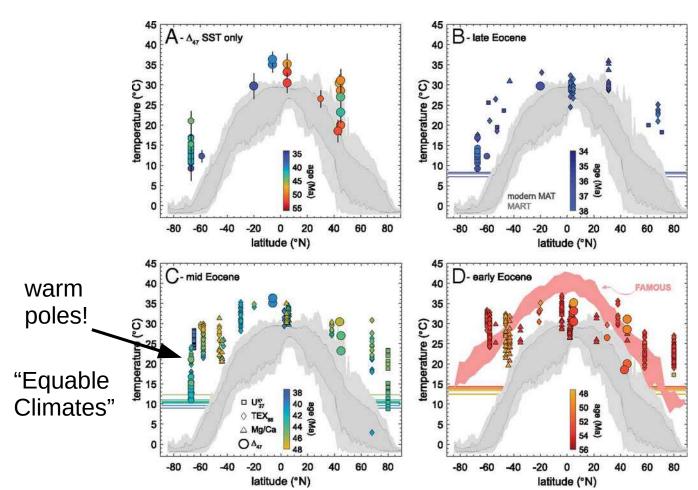


Evans et al. (2018) PNAS

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## **Equable Climate**

#### Eocene latitudinal sea surface T

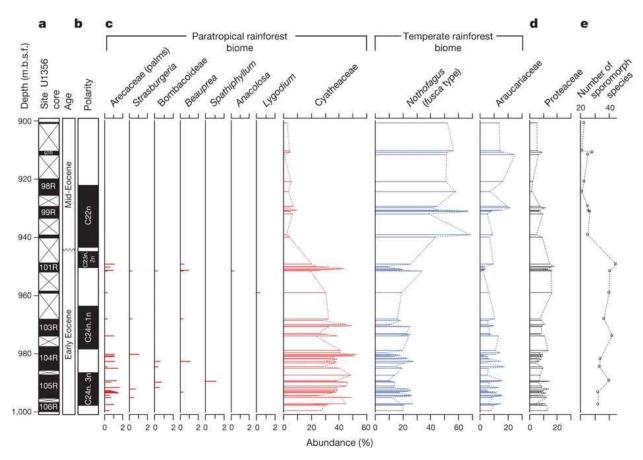


Evans et al. (2018) PNAS

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# **Equable Climate**

#### Eocene East Antarctic climate from pollen

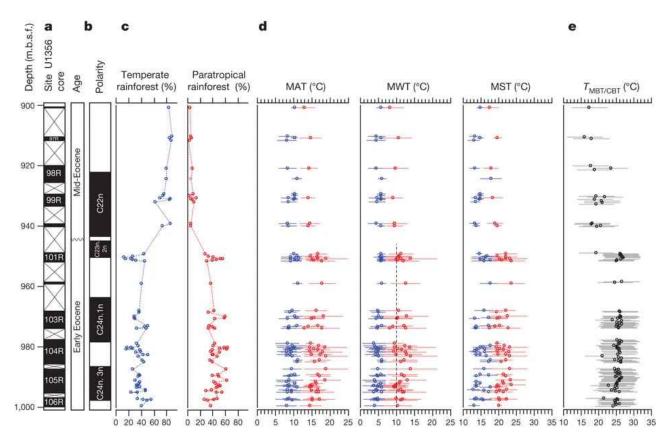


Pross et al. (2012) Nature

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## **Equable Climate**

#### Eocene East Antarctic climate from pollen

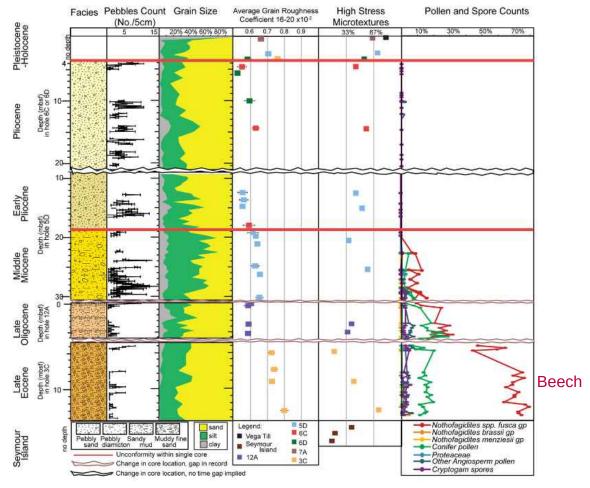


Pross et al. (2012) Nature

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## **Equable Climate**

#### Cenozoic West Antarctic climate from pollen



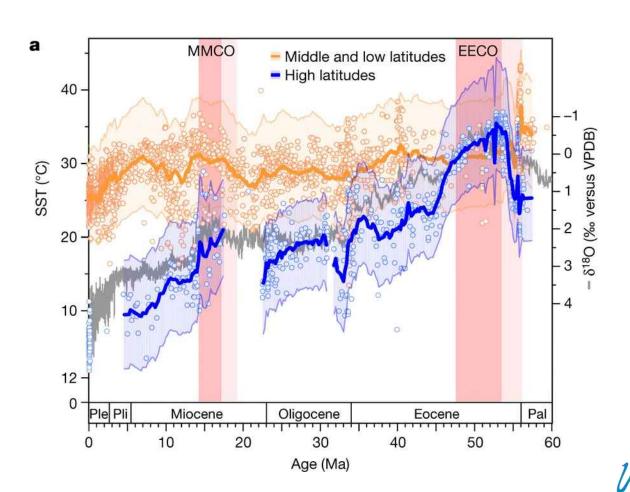
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**PNAS** 

Anderson et al. (2011)

## **Equable Climate**

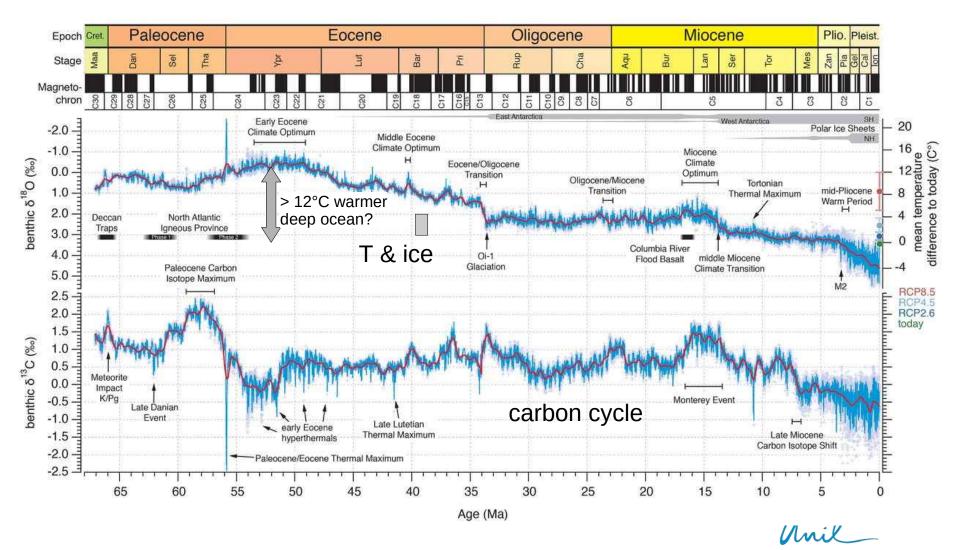
#### Evolution of latitudinal temperature gradient



Auderset et al. (2022) Nature

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## **Equable Climate**



# **Equable Climate**

#### Causes

high altitude cloud cover?

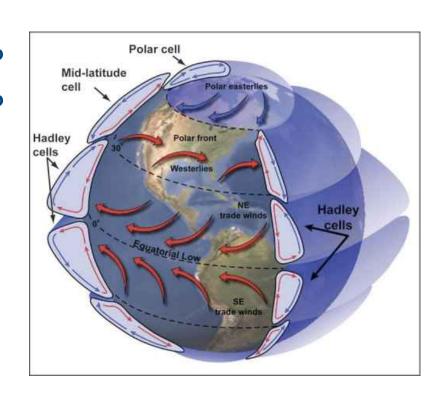




# **Equable Climate**

#### **Causes**

- high altitude cloud cover?
- atmospheric cell change?





# **Equable Climate**

#### Causes

- high altitude cloud cover?
- atmospheric cell change?
- polar stratospheric clouds?

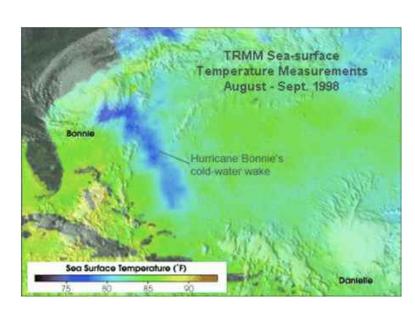




## **Equable Climate**

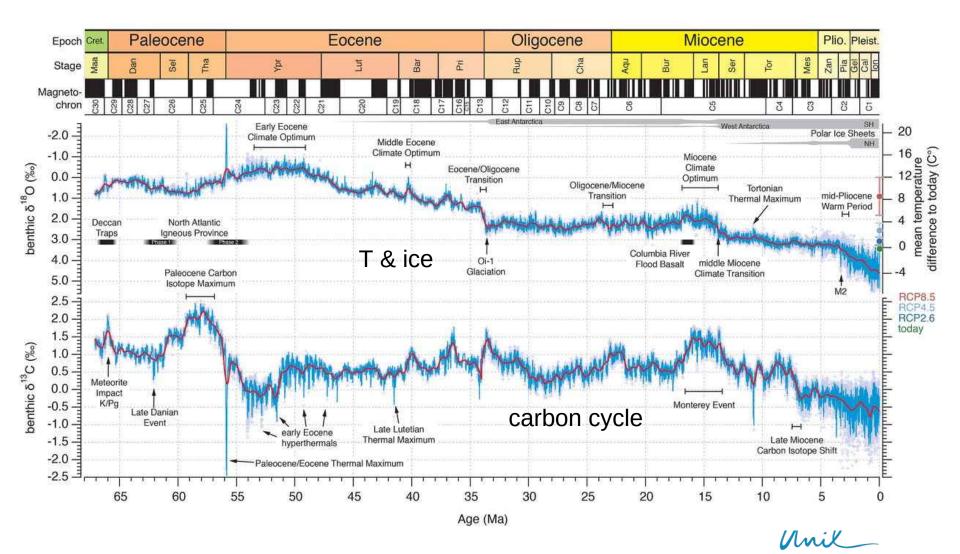
#### Causes

- high altitude cloud cover?
- atmospheric cell change?
- polar stratospheric clouds?
- cyclone ocean mixing?





### **Cenozoic Climate**



## **Carbon Isotopes**

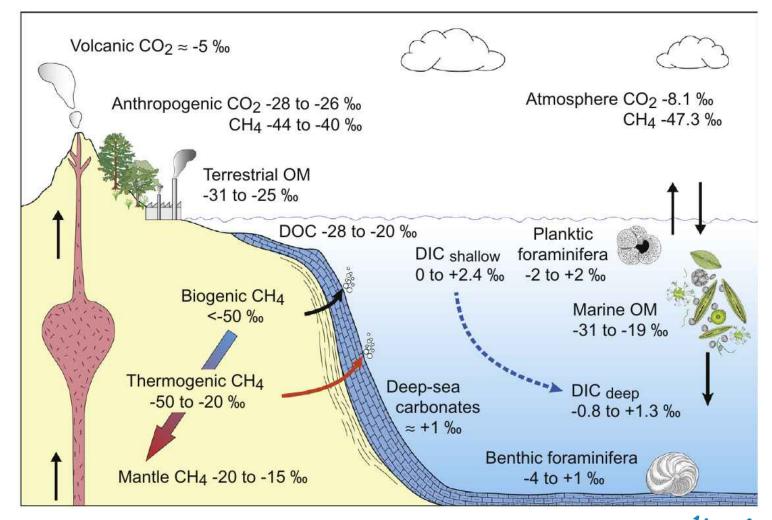


## **Carbon Isotopes**

- C & O isotopes can be measured from CaCO<sub>3</sub>
- C also in organics or gas
- <sup>12</sup>C 98.9 %
   <sup>13</sup>C 1.1 %
- photosynthesis discriminates against  $^{13}\text{C}$  with  $\alpha \sim 1.25$

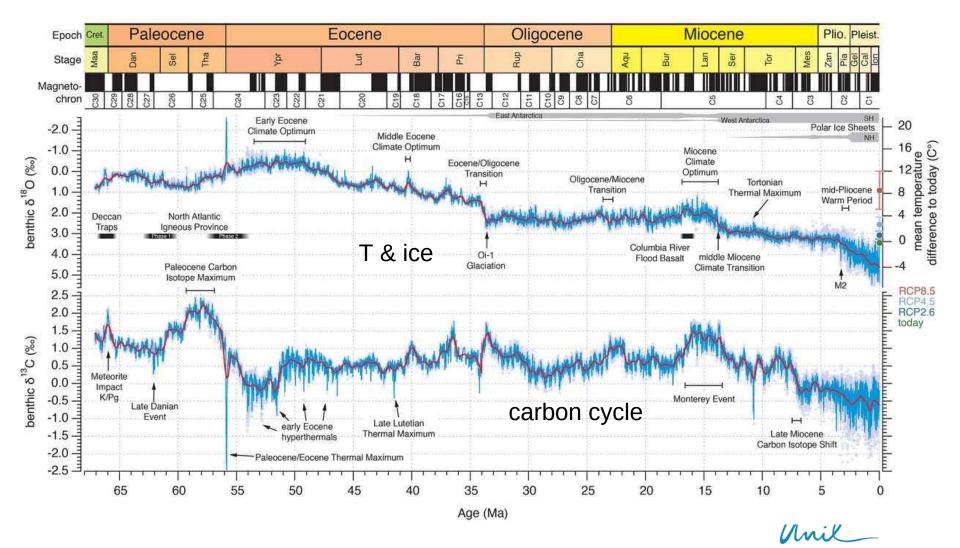


### **Carbon Isotopes**





### **Cenozoic Climate**



Cenozoic Clim. Hothouse Climate Sens. Isotopes Cen. Cooling

### **PETM**

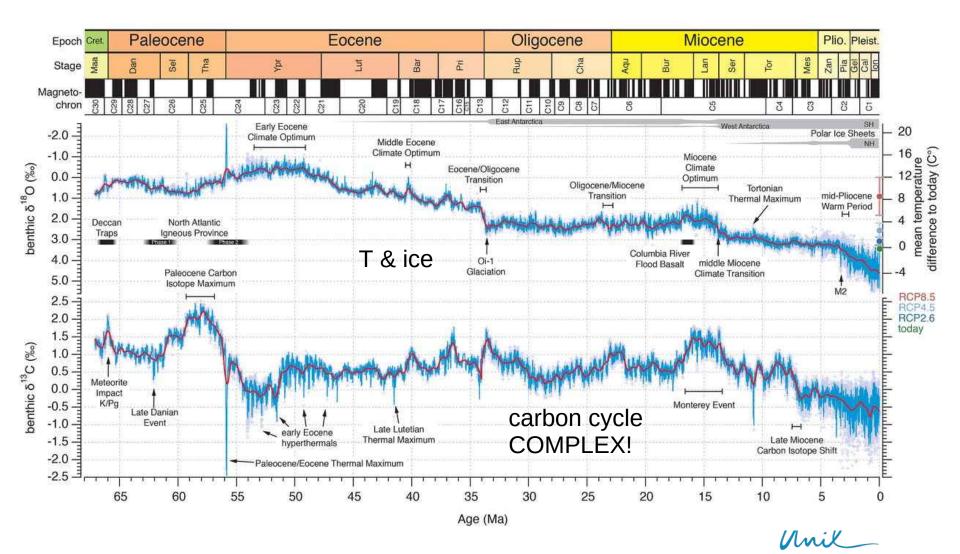
#### Paleocene-Eocene Thermal Maximum

caused by massive input of greenhouse gases (CO, and/or CH,

#### possible causes:

- submarine methane hydrates
- uplift and weathering of marine shelves
- warming-induced death of tropical plants (due to photorespiration)
- North Atlantic volcanism
- permafrost thaw

### **Cenozoic Climate**



### **Boron Isotopes**



Cenozoic Clim. Hothouse Climate Sens. Isotopes Cen. Cooling

## **Boron Isotopes**

boron species in seawater:  $B(OH_3)$  und  $B(OH)_4$ 

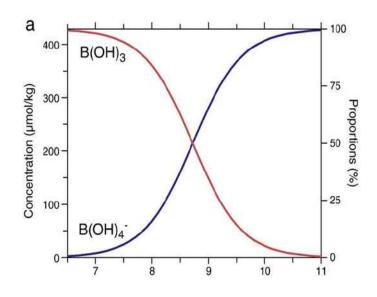
$$\delta^{11}B(\%) = \left[ \left( \frac{{}^{11}B/{}^{10}B_{sample}}{{}^{11}B/{}^{10}B_{NIST951}} \right) - 1 \right] \times 1000.$$

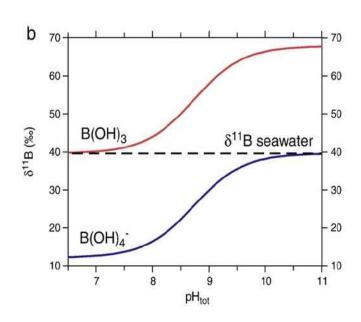
- B is fractionated between the two species
- B(OH)<sub>1</sub> is built into the shells of foraminifera

Hothouse Cenozoic Clim. Isotopes Climate Sens. Cen. Cooling

### **Boron Isotopes**

#### boron species in seawater:



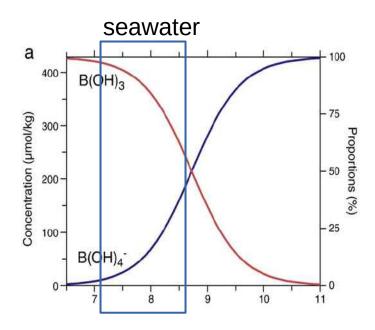


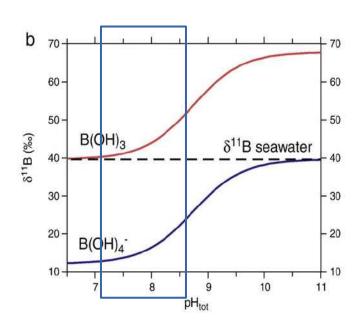
Rae et al. (2011) Earth and Planetary Science Letters



## **Boron Isotopes**

#### boron species in seawater:





Rae et al. (2011) Earth and Planetary Science Letters



## **Boron Isotopes**

#### boron isotopes in benthic foraminifera

#### **Epifaunal**

Cibicidoides wuellerstorfi

Cibicidoides mundulus

Planulina ariminensis

Cibicidoides lobatus

Cibicidoides ungerianus

#### Infaunal

Cibicidoides robertsonianus

Oridorsalis umbonatus

Gyroidina soldanii

Lenticulina vortex

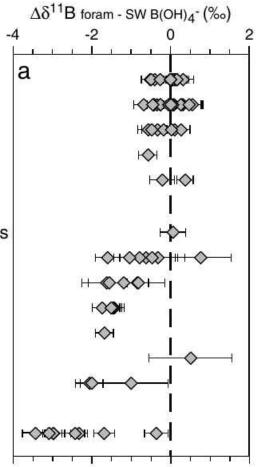
Ammonia beccarii

Melonis zaandamae

Uvigerina peregrina

#### Aragonite

Hoeglundina elegans

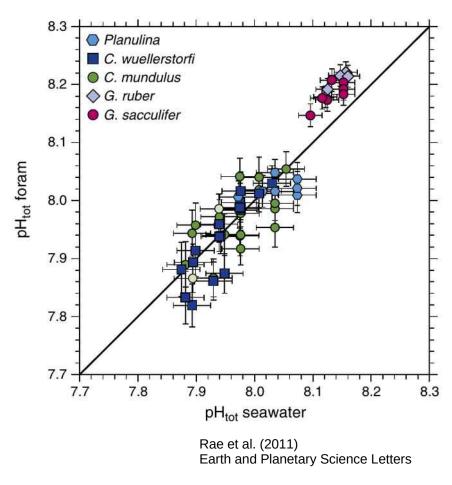




Rae et al. (2011)

### **Boron Isotopes**

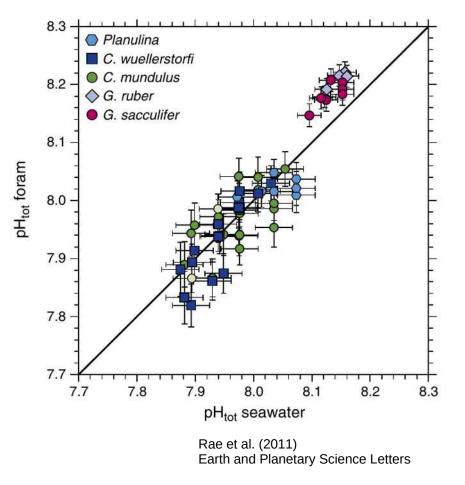
#### pH reconstructed from foraminifera



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### **Boron Isotopes**

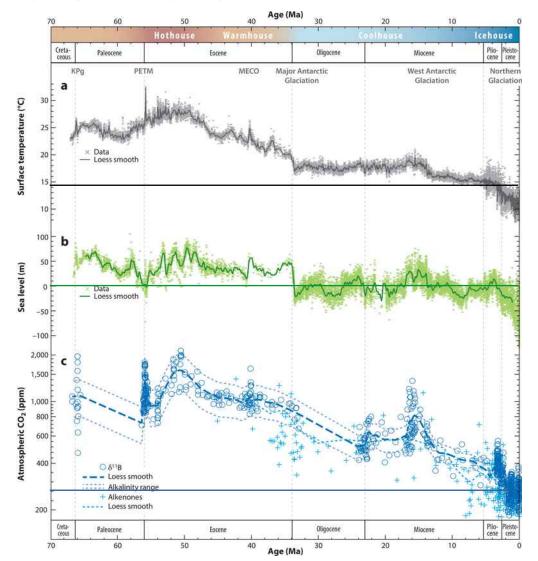
#### pH reconstructed from foraminifera



together with further assumptions, ocean pH traces (long term) atmospheric CO<sub>2</sub>

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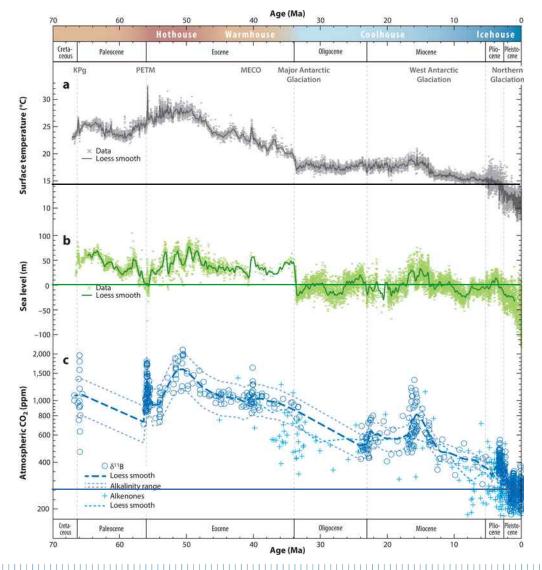
#### **Cenozoic Climate**





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#### **Cenozoic Climate**



Rae JWB, et al. 2021 Annu. Rev. Earth Planet. Sci. 49:609-41

T and CO<sub>2</sub> parallel?

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Cenozoic Clim. Hothouse Climate Sens. Isotopes Cen. Cooling

## **Climate Sensitivity**

How much does Earth warm with increasing CO<sub>2</sub>?

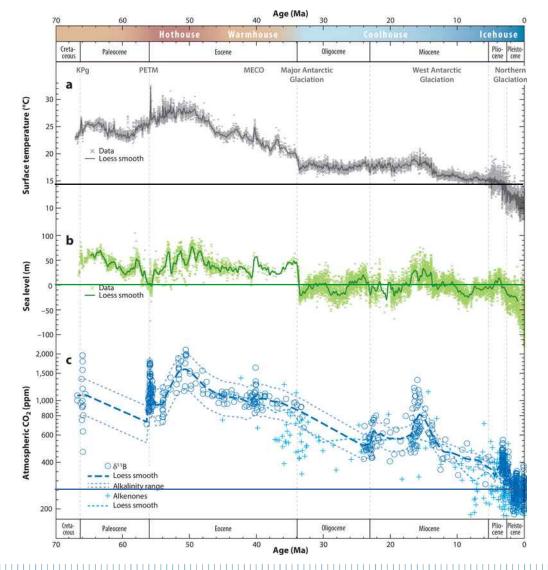
#### Equilibrium Climate Sensitivity (ECS)

- long-term, including geologic feedbacks
- usually referenced to doubling of CO<sub>2</sub>

#### Transient Climate Response (TCR)

- short term (~ 20 years) climate response
- including fast feedbacks
- often used for models

#### **Cenozoic Climate**

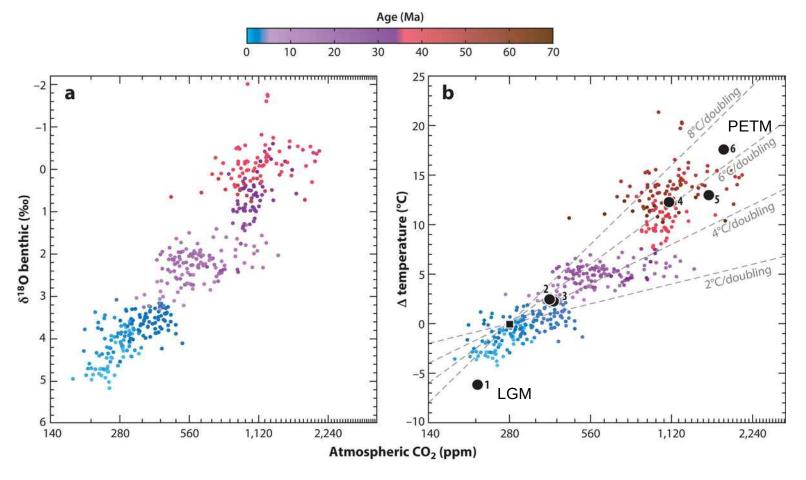


Rae JWB, et al. 2021 Annu. Rev. Earth Planet. Sci. 49:609-41

T and CO<sub>2</sub> parallel?

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### **Cenozoic Climate Sensitivity**

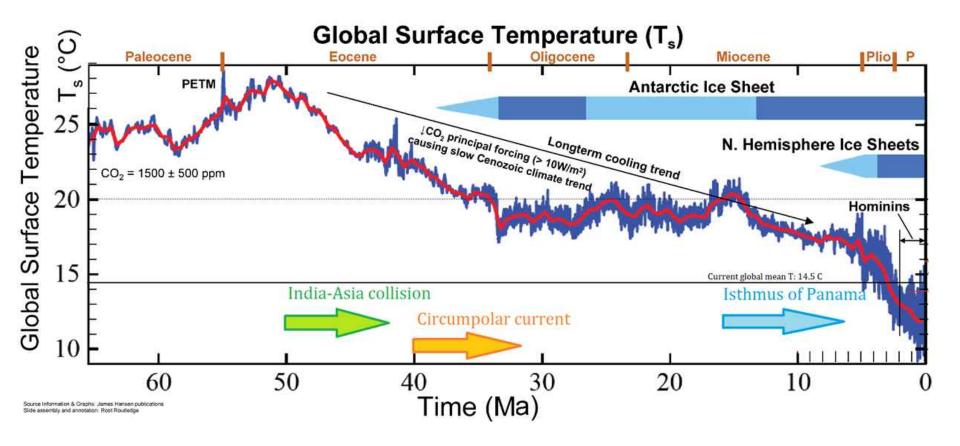




indicate independent proxy-derived estimates of surface temperature: • Last Glacial Maximum (Tierney et al. 2020b), • Pliocene (de la Vega et al. 2020), • Pliocene (McClymont et al. 2020), • late Paleocene, • Early Eocene Climatic Optimum, and • Paleocene–Eocene Thermal Maximum (Inglis et al. 2020). Dashed lines denote different degrees of temperature change per CO<sub>2</sub> doubling, providing an estimate of Earth system sensitivity.

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#### **Cenozoic Climate**



Earle (2016), opentextbc.ca after James Hansen and Root Routledge



Hothouse Cen. Cooling Cenozoic Clim. Isotopes Climate Sens.

## Cenozoic cooling



## Cenozoic cooling

Long term cooling trend from hot-house to ice-house

Causes debated and likely complex

- weathering?
- isolation of Antarctica?
- faunal changes?

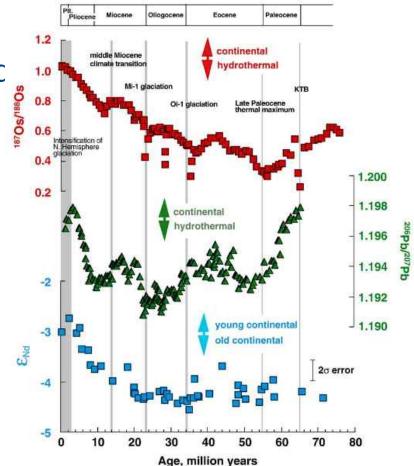


### Cenozoic cooling

Long term cooling trend from hot-house to ice-house

Causes debatec

• weathering?



Burton (2006) Journal of Geochemical Exploration

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### Cenozoic cooling

spread of extensive grass lands during Miocene favoured e.g. by seasonal aridity

development of C4 photosynthesis at ~ 10 Ma

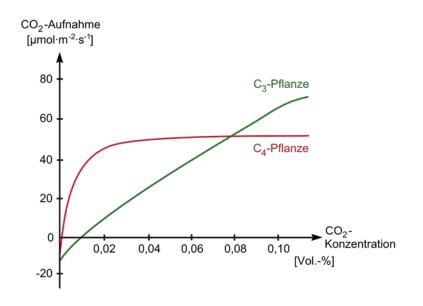
- developed multiple times
- fixes C in molecule containing 4 C atoms
- deals better with aridity and low CO<sub>2</sub>



## Cenozoic cooling

spread of extensive grass lands during Miocene favoured e.g. by seasonal aridity

deals better with aridity and low CO<sub>2</sub>



Wikipedia

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### Cenozoic cooling

spread of extensive grass lands during Miocene favoured e.g. by seasonal aridity

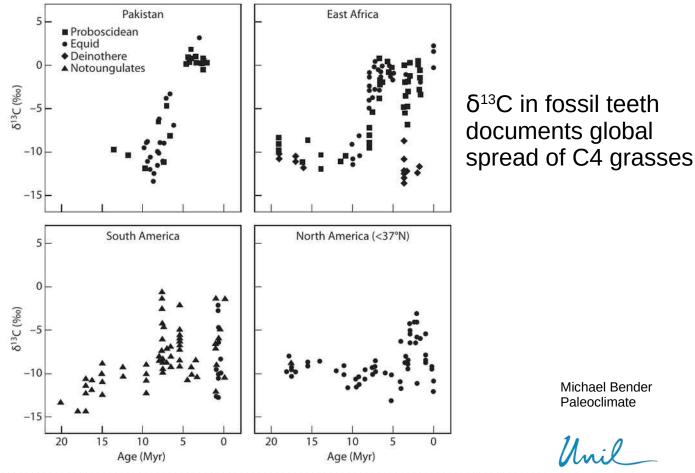
development of C4 photosynthesis at ~ 10 Ma

- developed multiple times
- fixes C in molecule containing 4 C atoms
- deals better with aridity and low CO<sub>2</sub>
- today ~ 25% of plants, mostly grasses
- global food production depends on C4 plants
- fractionates <sup>13</sup>C less than C3 plants



## Cenozoic cooling

#### spread of extensive grass lands during Miocene



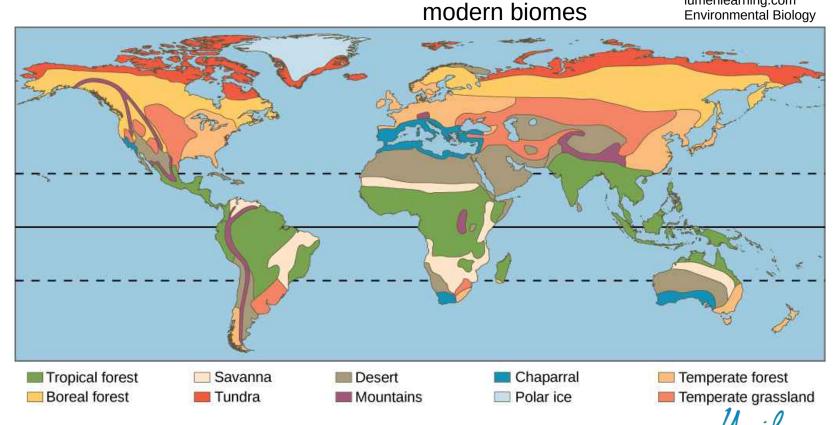
Cenozoic Clim. Hothouse Cen. Cooling Isotopes Climate Sens.

# Cenozoic cooling

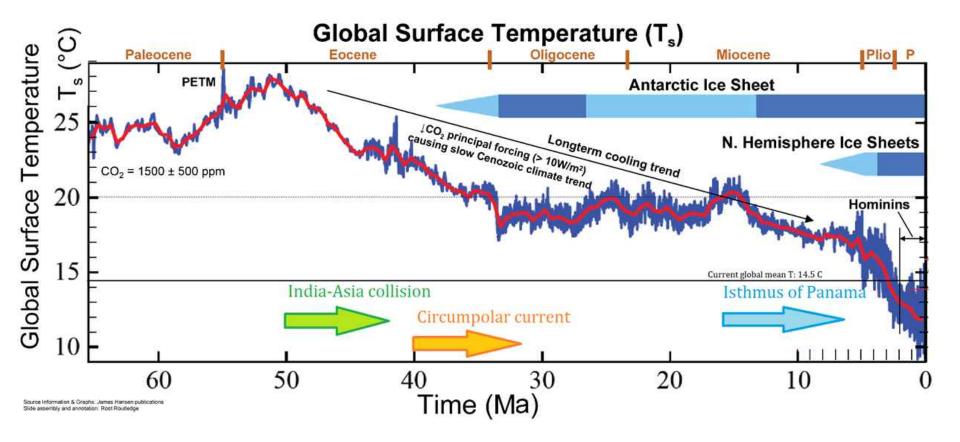
spread of extensive grass lands during Miocene



**Environmental Biology** 



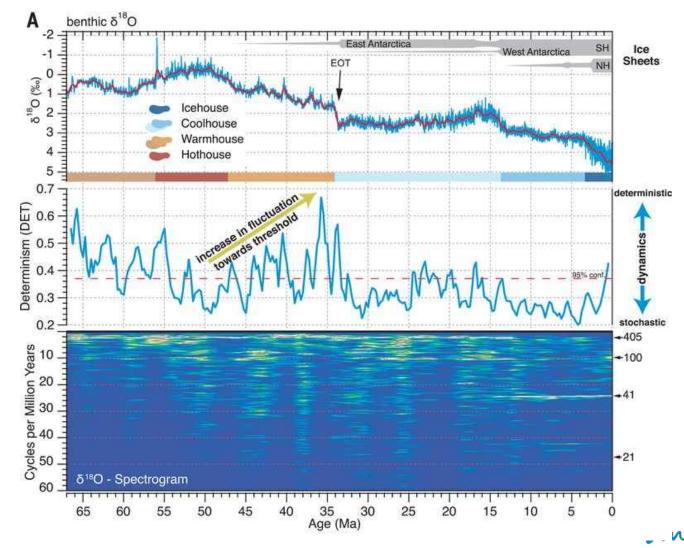
#### **Cenozoic Climate**



Earle (2016), opentextbc.ca after James Hansen and Root Routledge

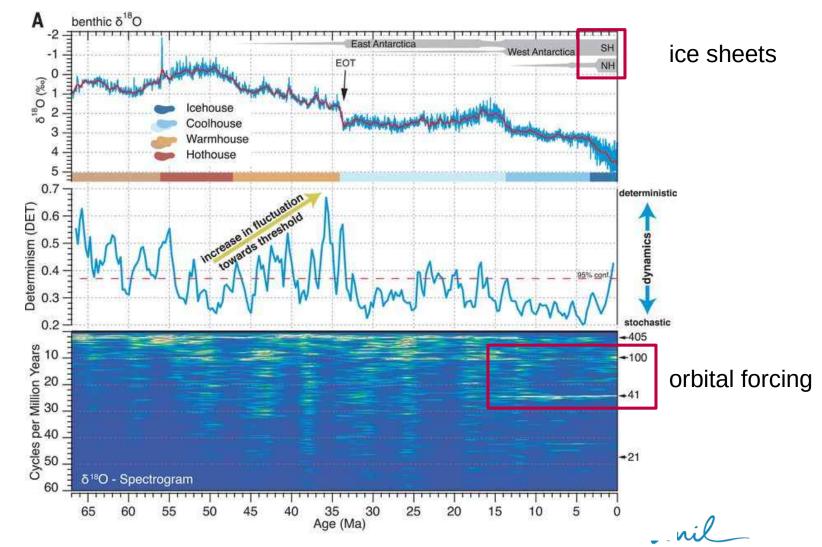


#### **Cenozoic Climate**



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#### **Cenozoic Climate**



Introduction Paleoclim. Methods Planet Earth Earth History

### **Today's Summary**

- Eocene Hothouse was very hot
- Equable climate led to warm poles
- PETM was extreme warm event caused by GHG
- Cenozoic climate dominated by CO<sub>2</sub>
- Cooling was accompanied by CO2 reduction and changes in weathering and fauna
- Temperature proxies:  $\delta^{18}$ O,  $\Delta 47$ , Mg/Ca, TEX86
- Carbon proxies:  $\delta^{13}$ C &  $\delta^{11}$ B

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### **Outlook**

#### Today we finish 15 min early!

Monday	Introduction	Earth History
Tuesday	Proxies I	Cenozoic Hot & Warm House
Wednesday	Specific Climate System components	Pleistocene G-IG climate
Thursday	Proxies II & Climate System Interactions	Abrupt Climate Change
Friday	Current Climate Change	Future & Synthesis

Mul.