Earth and Planetary Sciences (ES1101)

(Minerals: Building Blocks of Rocks) (Autumn 2020 by Gaurav Shukla)

Book: 1) Understanding Earth by Grotzinger & Jordan (Text Book)

- 2) Earth: An introduction to Physical Geology by Tarbuck & Lutgens
- 3) The Solid Earth: An introduction to global geophysics by Fowler

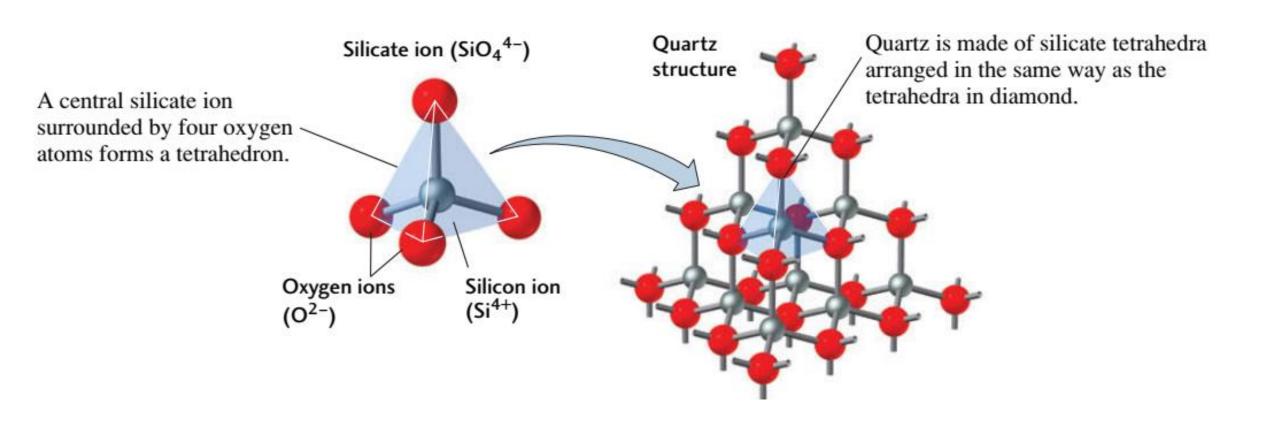
What are minerals?

- •Minerals are building blocks of rocks.
- •Rock is a solid aggregate or mass of minerals.
- •Understanding the behavior of minerals in response to changes in physical and chemical environment help us decipher the records of geological history.



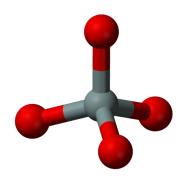
Definition of Minerals:

A mineral is a naturally occurring crystalline (inorganic) solid with a definite, but not necessarily a fixed, chemical composition.

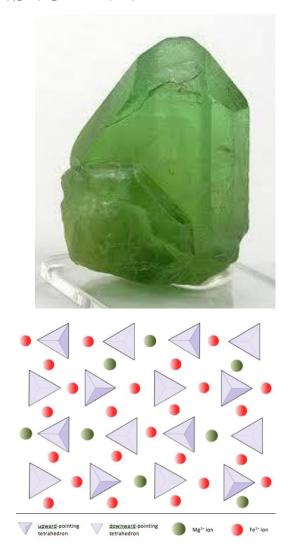


Definition of Minerals:

Definite BUT NOT NECESSARILY FIXED CHEMICAL COMPOSITION



OLIVINE (M_2SiO_4) ; M : Fe²⁺, Mg²⁺ M:Si:O = 2:1:4



Definition of Minerals:

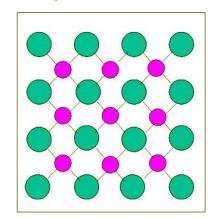
Volcanic GLASS (obsidian)

Criteria to be fulfilled

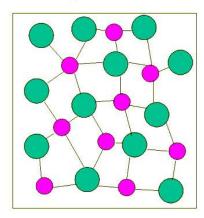
- Naturally occurring
- Regular arrangement of atoms X
- Solid
- Definite chemical composition X



Crystalline solid

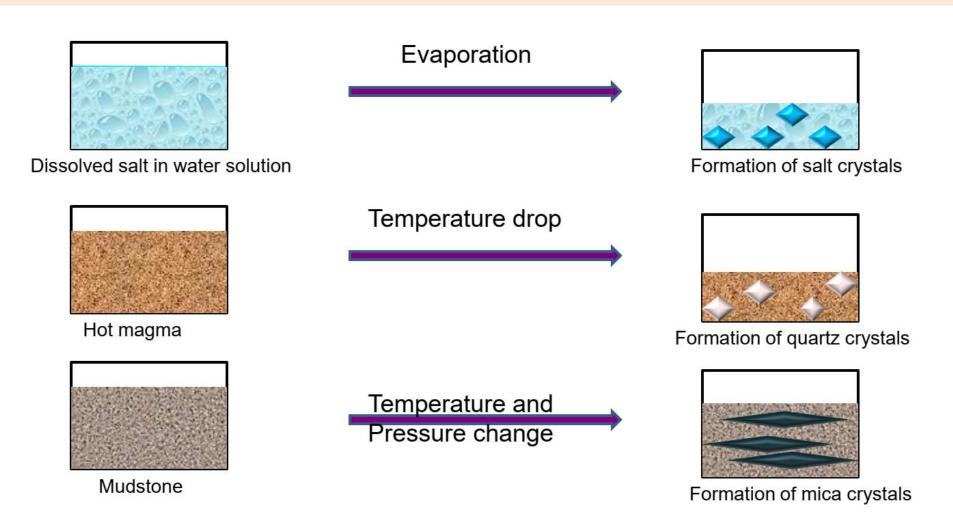


Amorphous solid



How do minerals form?

Minerals form through crystallization. The molecules, or ions chemically bond to form an orderly structure.



Hardness:

- Hardness of a mineral primarily depends on the crystal structure.
- It is classified based on the Mohs scale of relative hardness.

Habit/Shape

- Determined by crystal structure
- Depends on the speed and direction of crystal growth
- Types:
 - I. Bladed
 - II. Platy
 - III. Prismatic
 - IV. Banded
 - V. Fibrous
 - VI. Botryoidal

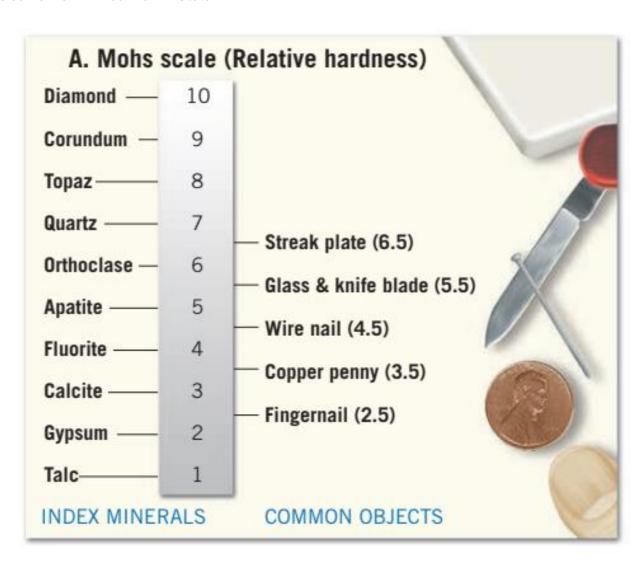
Cleavage: These are weak planes in the crystal structures of the minerals.

Specific gravity: Ratio of the weight of mineral with the equal volume of pure water at 4°C.

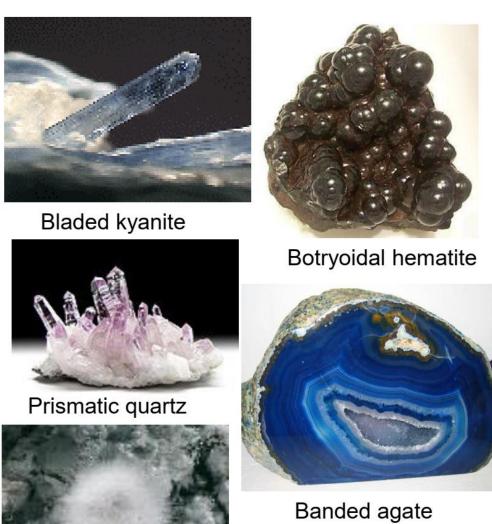
Optical Properties:

- Color
- Streak
- Luster

Hardness: Mohs scale of hardness



Shape



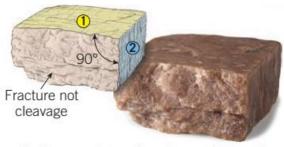
Fibrous okenite

Cleavage

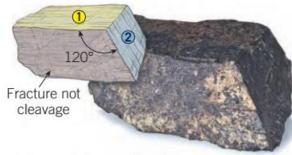
- Preferred plane of breakage.
- Creates a smooth plane
- Generated by crystal arrangements.



A. Cleavage in one direction. Example: Muscovite



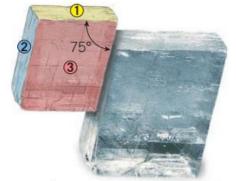
B. Cleavage in two directions at 90° angles. Example: Feldspar



C. Cleavage in two directions not at 90° angles. Example: Hornblende



D. Cleavage in three directions at 90° angles. Example: Halite



E. Cleavage in three directions not at 90° angles. Example: Calcite



F. Cleavage in four directions.
Example: Fluorite

Mineral Properties (Optical)

Color



Mineral Properties (Optical)

Streak: The color of powdered mineral



FIGURE 3.20 Hematite may be black, red, or brown, but it always leaves a reddish brown streak when scraped along a ceramic streak plate. [Breck P. Kent.]

Mineral Properties (Optical)

Luster: The appearance of light reflected from mineral surface.

Type:

- Metallic
- Non-metallic: Earthy/Dull, Glassy, Pearly, Silky



Metallic luster of pyrite



Glassy luster of quartz



Earthy luster of limonite



Pearly luster of talc



Silky luster of siderite

TABLE 3-4	Physical Properties of Minerals			
Property	Relation to Composition and Crystal Structure			
Hardness	Strong chemical bonds result in hard minerals. Covalently bonded minerals are generally harder than ionically bonded minerals.			
Cleavage	Cleavage is poor if bonds in crystal structure are strong, good if bonds are weak. Covalent bonds generally give poor or no cleavage; ionic bonds are weaker and so give good cleavage.			
Fracture	Related to distribution of bond strengths across irregular surfaces other than cleavage planes.			
Luster	Tends to be glassy for ionically bonded crystals, more variable for covalently bonded crystals.			
Color	Determined by ions and trace elements. Many ionically bonded crystals are colorless. Iron tends to color strongly.			
Streak	Color of fine mineral powder is more characteristic than that of massive mineral because of uniformly small size of grains.			
Density	Depends on atomic weight of atoms or ions and their closeness of packing in crystal structure.			
Crystal habit	Depends on the planes of a mineral's crystal structure and the typical speed and direction of crystal growth.			

- More than 4000 minerals have been identified, and several new ones are identified each year.
- Common minerals that make up most of the Earth's crust are only a few dozens and known as rock-forming minerals.

Most abundant elements in the crust

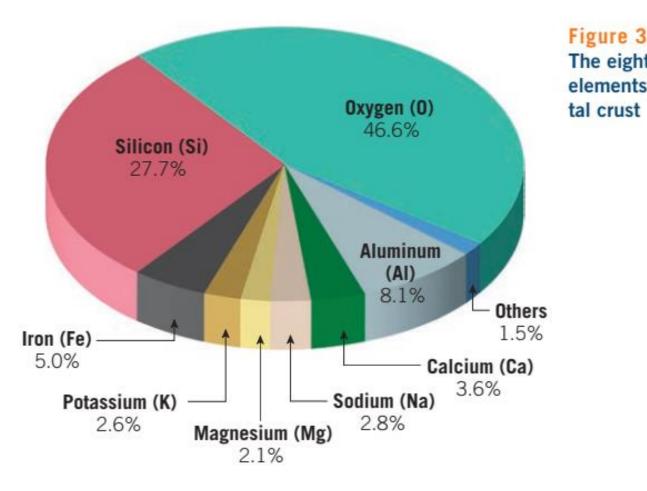
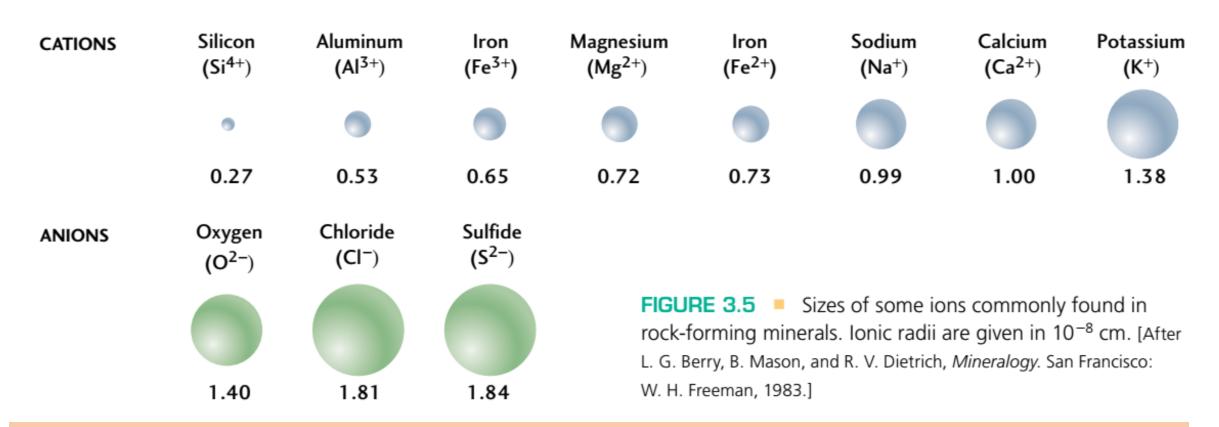


Figure 3.30
The eight most abundant elements in the continen-

Radii of the most relevant cations and anions



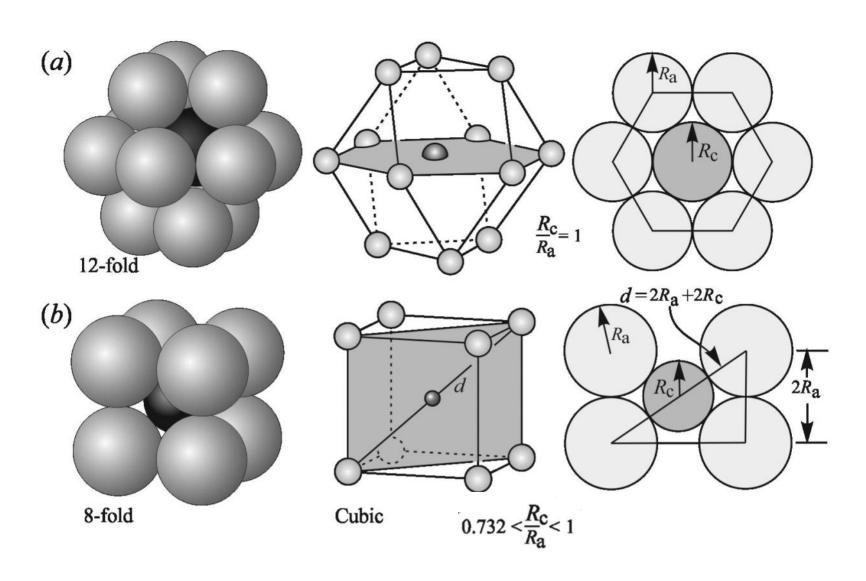
These radii shown here are averaged out rough estimates. You may follow the link for the expected values of radii: http://abulafia.mt.ic.ac.uk/shannon/ptable.php

Radii of the most relevant cations and anions

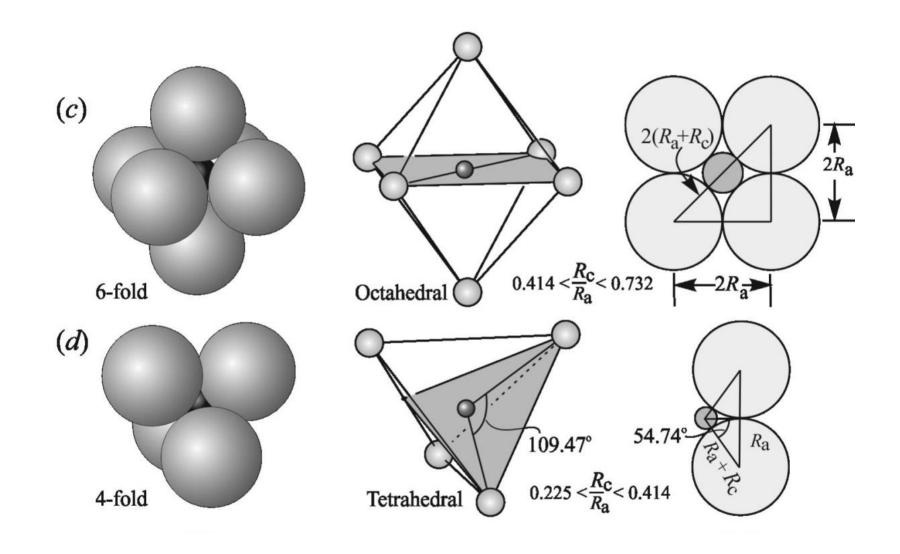
Table 4.2 Cation Sizes Appropriate for Regular Coordination Polyhedra When Coordinating with O²⁻ (~1.26 Å)

Coordination	Radius Ratio		~Maximum Radius (Å)	Common Cations
12	~1.00	1.26	N/A	K+, Ca2+, Na+
8	0.732-1.00	0.92	1.26	Fe ²⁺ , Ca ²⁺ , Na ⁺ , Mg ²⁺
6	0.414-0.732	0.52	0.92	Al^{3+} , Fe^{2+} , Fe^{3+} , Mg^{2+}
4	0.225-0.414	0.28	0.52	Si ⁴⁺ , Al ³⁺ , S ⁶⁺ , P ⁵⁺
3	0.155-0.225	0.20	0.28	C ^{4+ a}
2	< 0.155	N/A	0.20	None

Radius ratio and coordination number (contd)



Radius ratio and coordination number (contd)



Radius ratio and coordination number

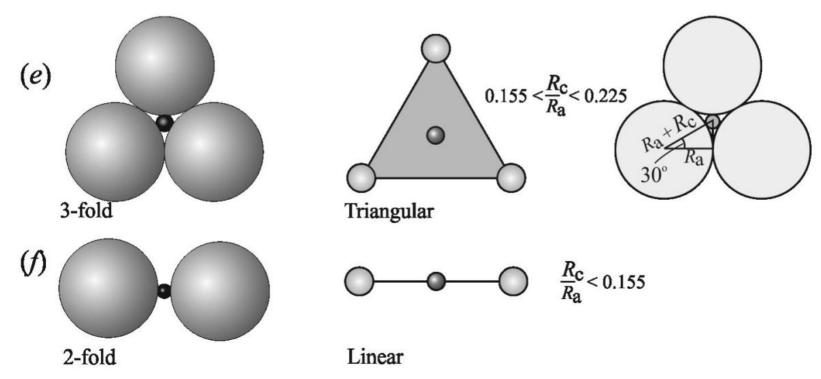
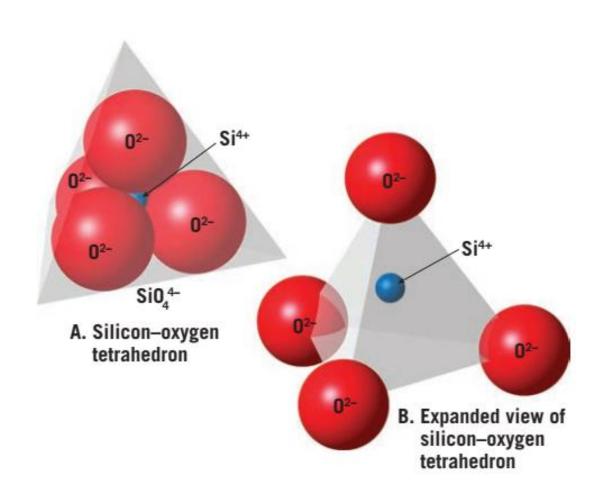
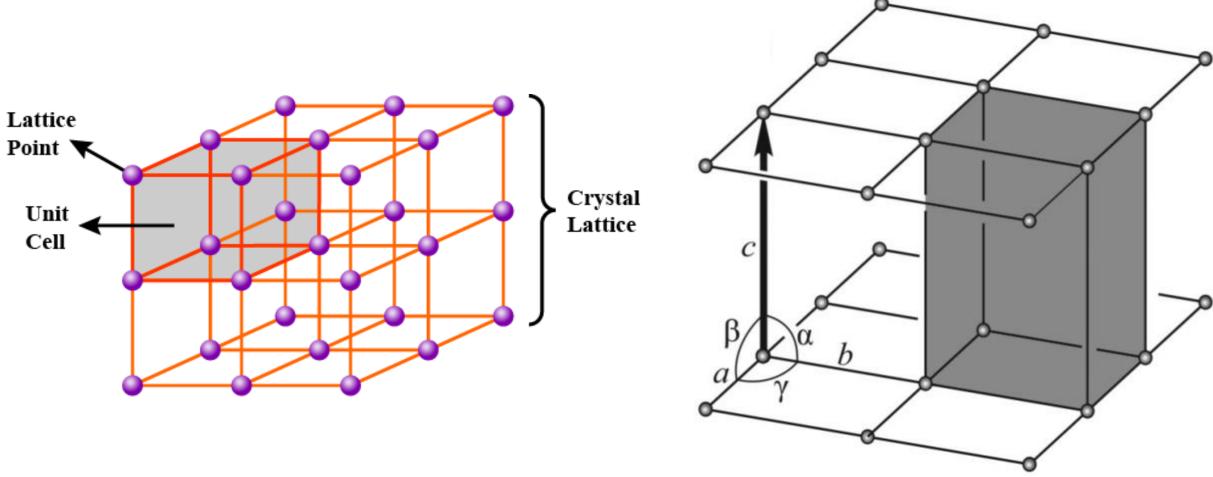
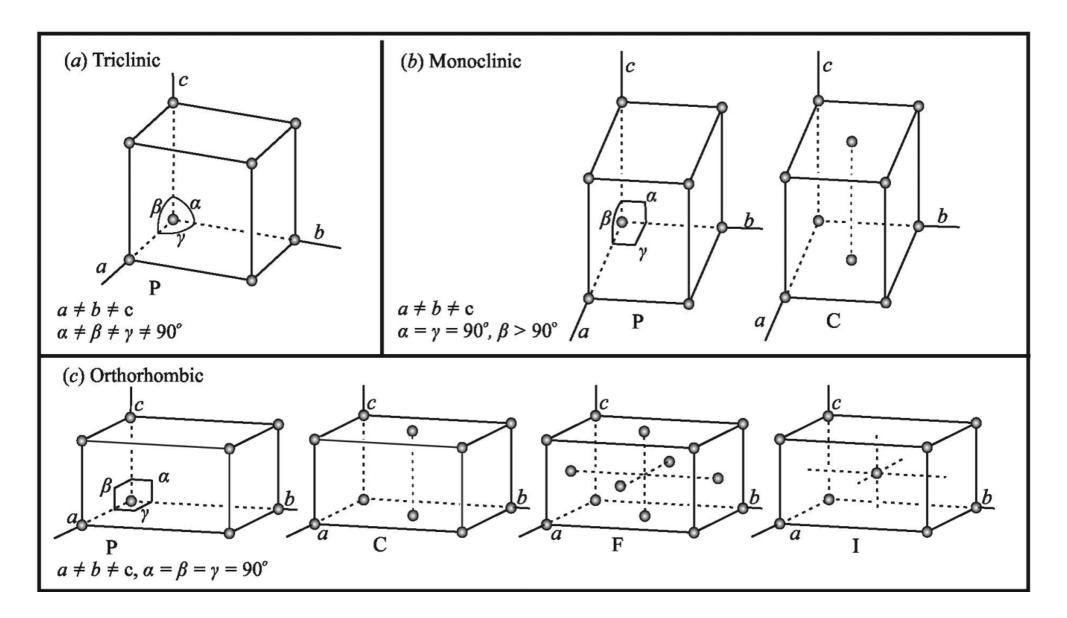


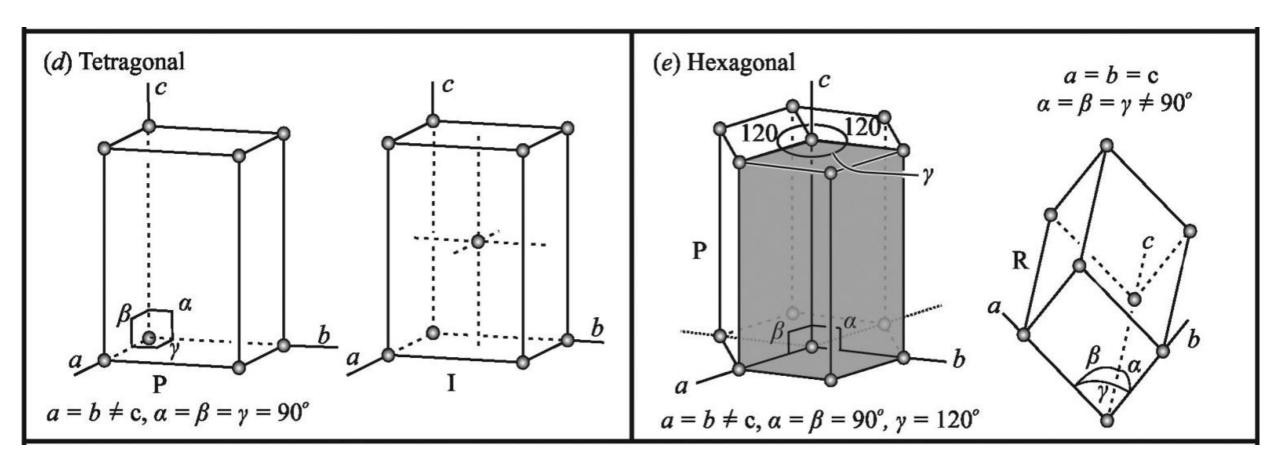
Figure 4.4 Coordination polyhedra. Anions with radius R_a are shown with light shading, cations with radius R_c with dark. Left view shows cation and anions drawn to scale. Center view shows the coordination polyhedra. Right view shows the plane through the polyhedron (shaded in center view) from which the radius ratio is calculated. See text for additional discussion. (a) 12-fold coordination (based on cubic closest packing). The coordination polyhedron is not a regular shape. (b) 8-fold, or cubic, coordination. (c) 6-fold, or octahedral, coordination. (d) 4-fold, or tetrahedral, coordination. (e) 3-fold, or triangular, coordination. (f) 2-fold, or linear, coordination.

Silicon-Oxygen Tetrahedra









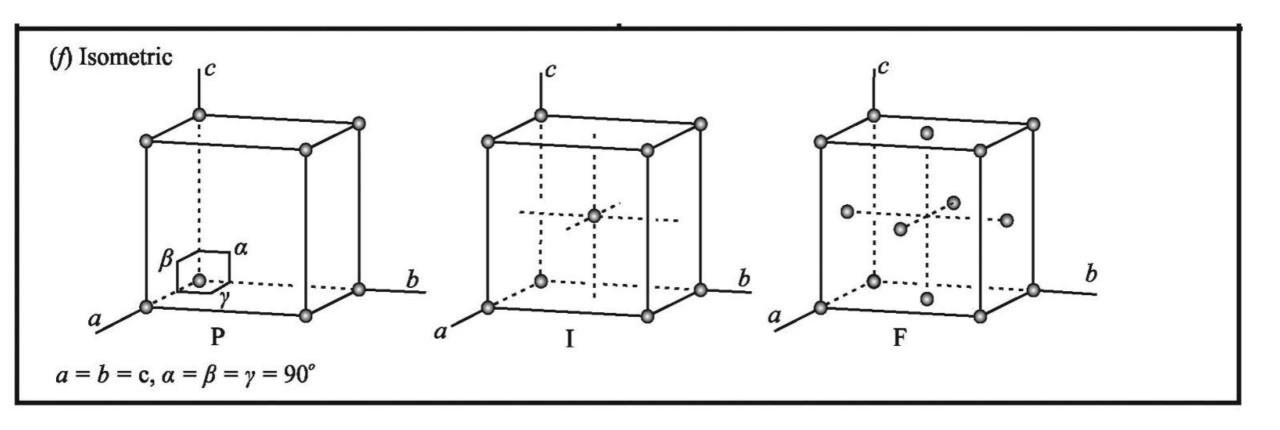


Figure 2.10 The 14 Bravais lattices define six different three-dimensional volumes (a-f) that correspond to the unit cells of the six crystal systems. The lengths of the three unit cell axes are a, b, and c, and the angles between them are α , β , and γ . In the notations, the \neq sign indicates that equality of the axis lengths or angles is not required, although occasionally equality may occur by chance. The hexagonal (R) lattice shown in (e) is based on the rhombohedral axes shown in Figure 2.8c. A unit cell with the same geometry as the hexagonal (P) lattice shown in Figure 2.8a is far more commonly used.

- More than 4000 minerals have been identified, and several new ones are identified each year.
- Common minerals that make up most of the Earth's crust are only a few dozens and known as rock-forming minerals.
- As we have seen that the oxygen and silicon are the most common elements in the Earth's crust, so the silicate minerals account for more than 90% of the crust.

TABLE 3-1	Some Chemical Classes of Minerals	
Class	Defining Anions	Example
Native elements	None: no charged ions	Copper metal (Cu)
Oxides	Oxygen ion (O ²⁻)	Hematite (Fe ₂ O ₃)
Halides	Chloride (Cl $^-$), fluoride (F $^-$), bromide (Br $^-$), iodide (I $^-$)	Halite (NaCl)
Carbonates	Carbonate ion (CO ₃ ²⁻)	Calcite (CaCO ₃)
Sulfates	Sulfate ion (SO ₄ ²⁻)	Anhydrite (CaSO ₄)
Silicates	Silicate ion (SiO ₄ ⁴⁻)	Olivine (Mg,Fe) ₂ SiO ₄
Sulfides	Sulfide ion (S ²⁻)	Pyrite (FeS ²)

Nonsilicate Minerals

Common Nonsilicate Mineral Groups				
Mineral Group (key ion(s) or element(s))	Mineral Name	Chemical Formula	Economic Use	Examples
Carbonates (CO ²⁻ ₃)	Calcite Dolomite	CaCO ₃ CaMg(CO ₃) ₂	Portland cement, lime Portland cement, lime	B. Dolomite A. Calcite
Halides (Cl ¹⁻ , F ¹⁻ , Br ¹⁻)	Halite Fluorite Sylvite	NaCI CaF ₂ KCI	Common salt Used in steelmaking Used as fertilizer	C. Halite D. Fluorite
Oxides (O²-)	Hematite Magnetite Corundum Ice	Fe ₂ O ₃ Fe ₃ O ₄ Al ₂ O ₃ H ₂ O	Ore of iron, pigment Ore of iron Gemstone, abrasive Solid form of water	E. Hematite F. Magnetite
Sulfides (S²-)	Galena Sphalerite Pyrite Chalcopyrite Cinnabar	PbS ZnS FeS ₂ CuFeS ₂ HgS	Ore of lead Ore of zinc Sulfuric acid production Ore of copper Ore of mercury	H. Chalcopyrite G. Galena
Sulfates (SO ² ₄ -)	Gypsum Anhydrite Barite	CaSO ₄ •2H ₂ O CaSO ₄ BaSO ₄	Plaster Plaster Drilling mud	J. Anhydrite I. Gypsum
Native elements (single elements)	Gold Copper Diamond Graphite Sulfur Silver	Au Cu C C S Ag	Trade, jewelry Electrical conductor Gemstone, abrasive Pencil lead Sulfadrugs,chemicals Jewelry, photography	K. Copper L. Sulfur

Nonsilicate Minerals: Calcite



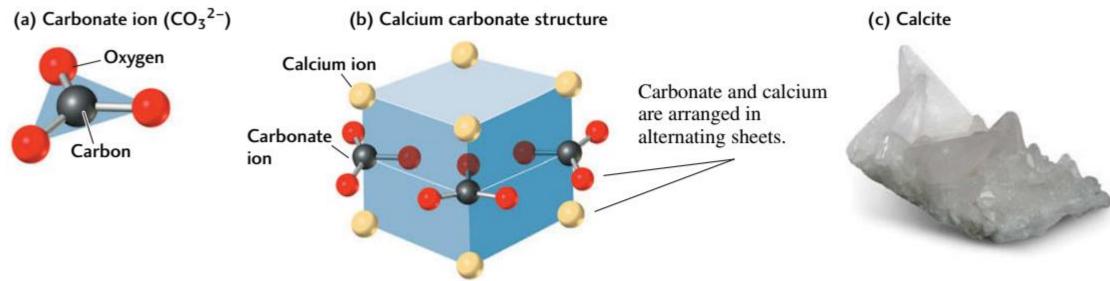


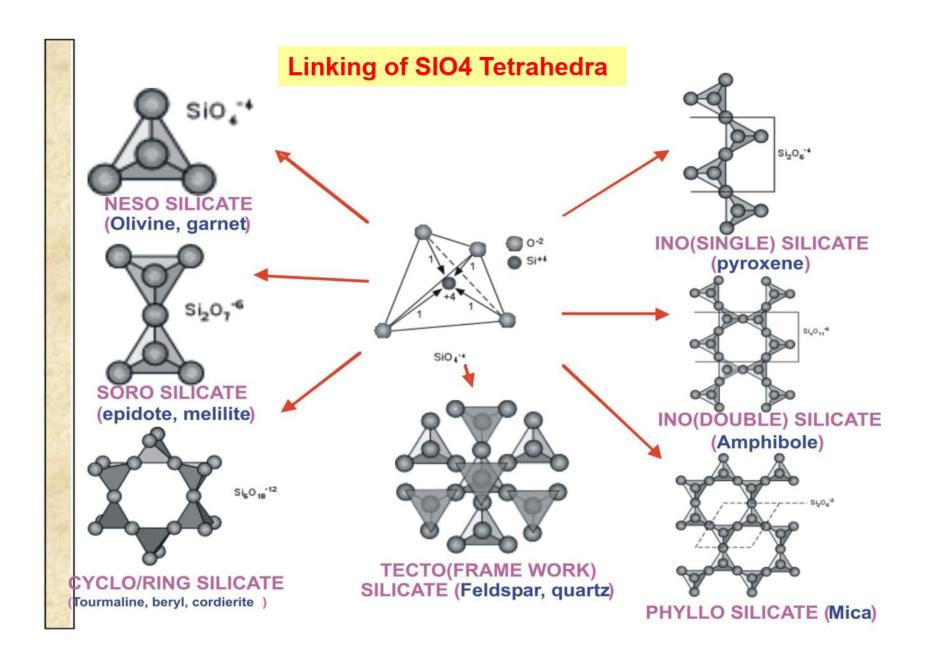
FIGURE 3.12 Carbonate minerals, such as calcite (calcium carbonate, CaCO₃), have a layered structure. (a) Top view of the carbonate ion, composed of a carbon ion surrounded by three oxygen ions in a triangle. (b) View of the alternating layers of calcium and carbonate ions in calcite. (c) Calcite. [Photo by John Grotzinger/Ramón Rivera-Moret/Harvard Mineralogical Museum.]

Nonsilicate Minerals: Pyrite



FIGURE 3.14 Pyrite, a sulfide mineral, is also known as "fool's gold." [John Grotzinger/Ramón Rivera-Moret/Harvard Mineralogical Museum.]

Minerals Classification: Silicate Minerals



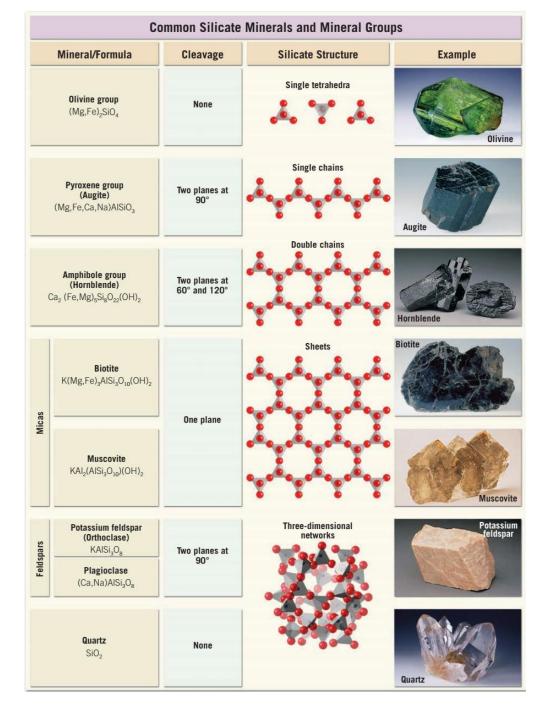
Minerals Classification: Silicate Minerals

Table 11.1 Silicate Classification^a

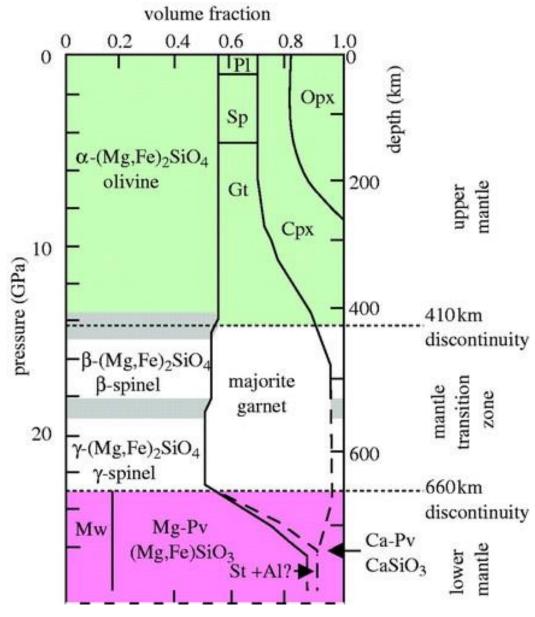
Silicate Class	Number of O ²⁻ Shared per Tetrahedron	Z:O Ratio	Structural Configuration
Orthosilicates	0	1:4	Isolated tetrahedra
Disilicates	1	2:7	Double tetrahedra
Ring silicates	2	1:3	Rings of tetrahedra
Chain silicates			Chains of tetrahedra
Single chain	2	1:3	
Double chain	2 or 3	4:11	
Sheet silicates	3	2:5	Sheets of tetrahedra
Framework silicates	4	1:2	Framework of tetrahedra

^aZ refers to the cation(s), usually Si⁴⁺, and also Al³⁺, that occupy the tetrahedral sites.

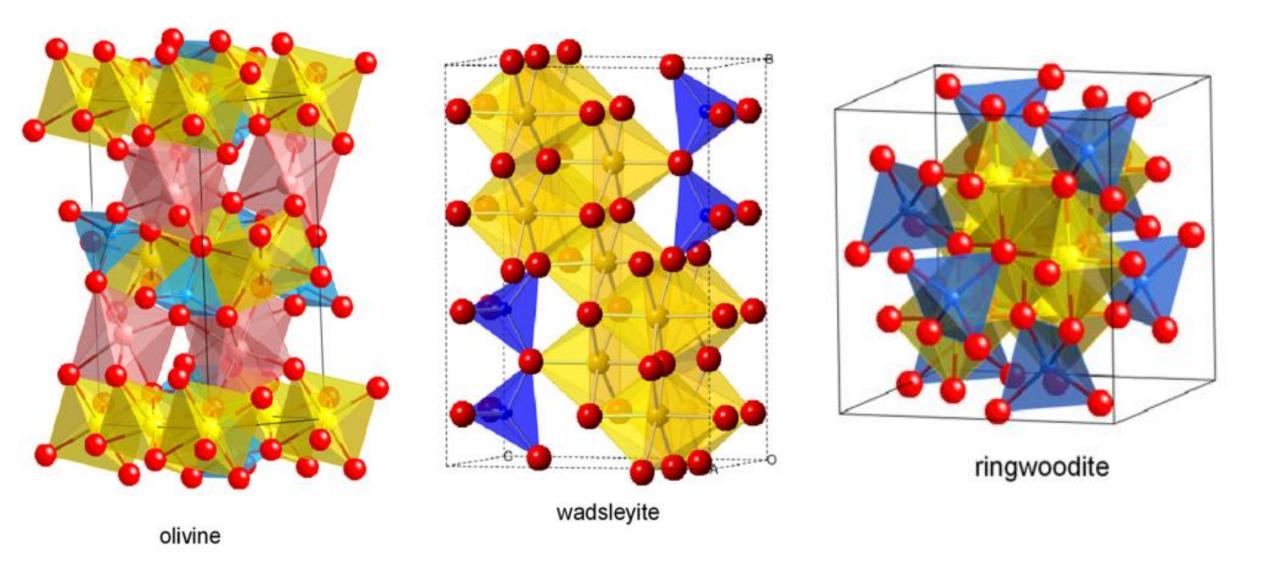
Minerals Classification: Silicate Minerals



Phase Transitions in Olivine

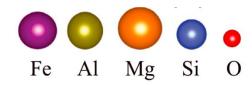


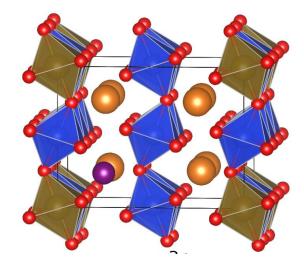
Phase Transitions in Olivine



Phase Transitions in Bridgmanite







Post-perovskite, Orthorhombic (expected to be in the D" region)