Earth and Planetary Sciences (ES1101)

(Interior of the Earth)
(Autumn 2021 by Gaurav Shukla)

- **Book: 1) Understanding Earth by Grotzinger & Jordan (Text Book)**
 - 2) Earth: An introduction to Physical Geology by Tarbuck & Lutgens
 - 3) The Solid Earth: An introduction to global geophysics by Fowler

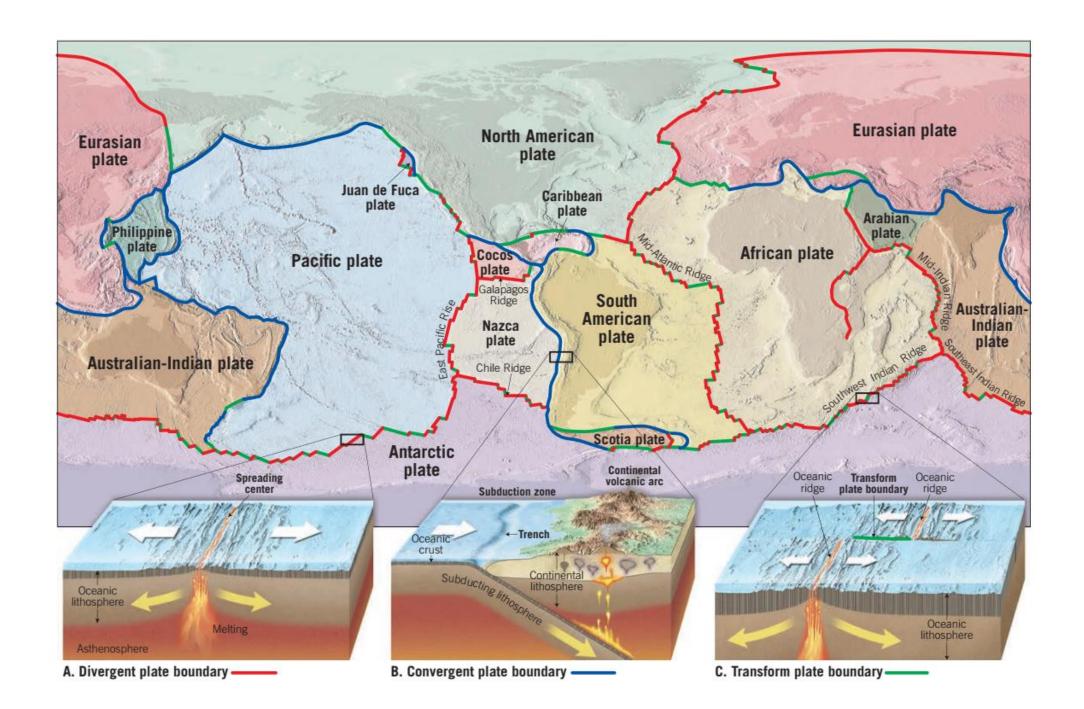
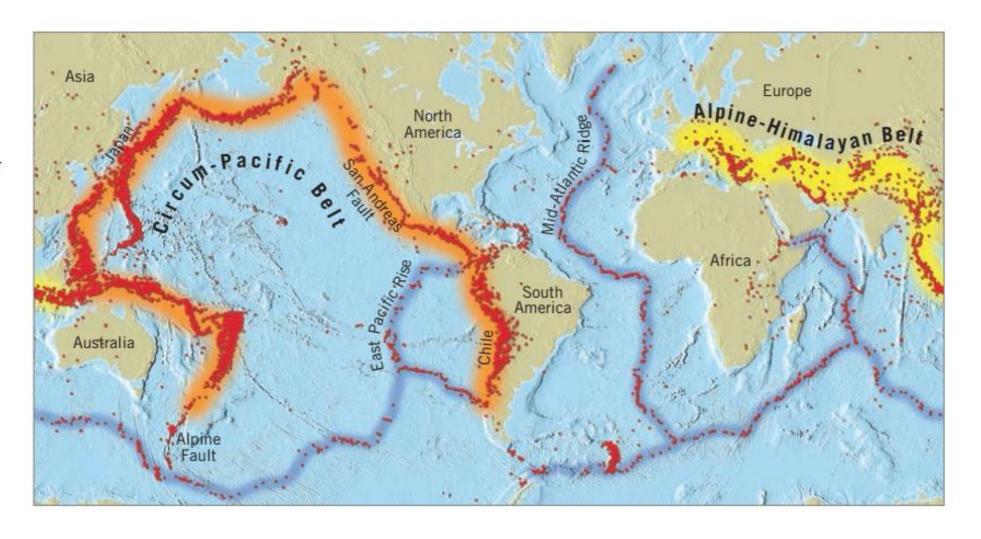


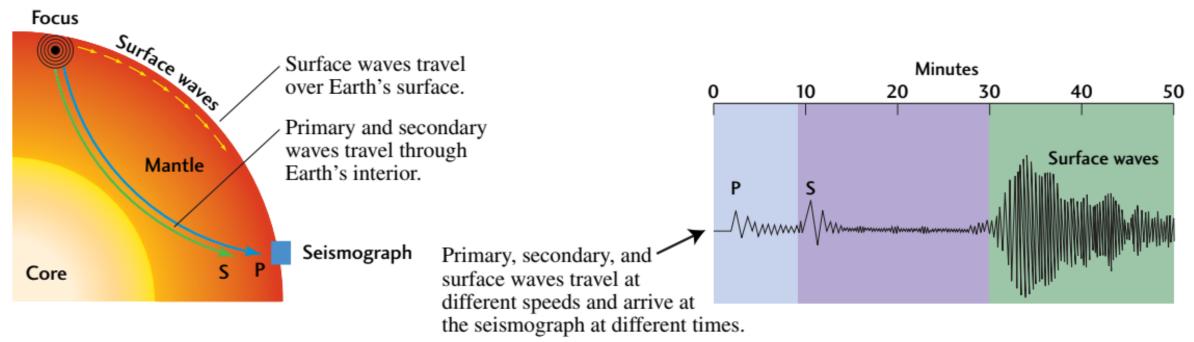
Figure 11.33

Global earthquake

belts Distribution of nearly 15,000 earthquakes with magnitudes equal to or greater than 5 for a 10-year period. (Data from USGS)



(a) Seismic waves generated at an earthquake focus arrive at a seismograph far from the earthquake.

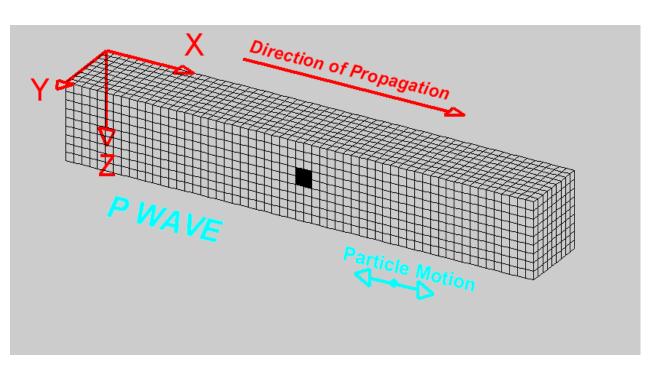


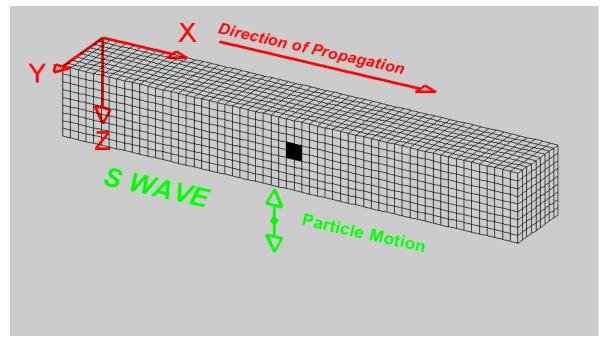
Body Waves: P-wave or Compressional wave S-wave or Shear wave

Surface Waves: Love wave

Rayleigh wave

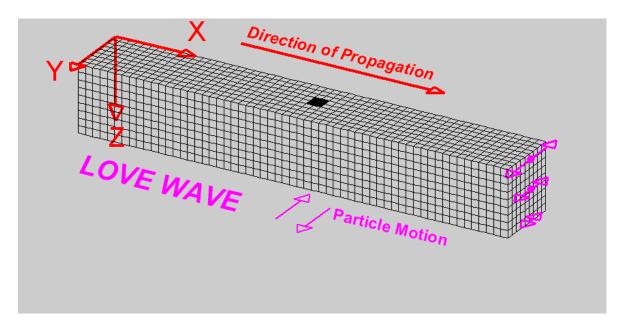
Body waves: Compressional (P-wave) Shear (S-wave) wave

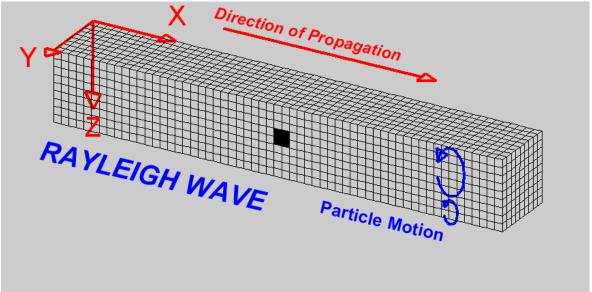




http://www.geo.mtu.edu/UPSeis/waves.html

Surface waves: Love wave and Rayleigh wave





http://www.geo.mtu.edu/UPSeis/waves.html

Note the time interval (about 5 minutes) between the arrival of the first P wave and the arrival of the first S wave.

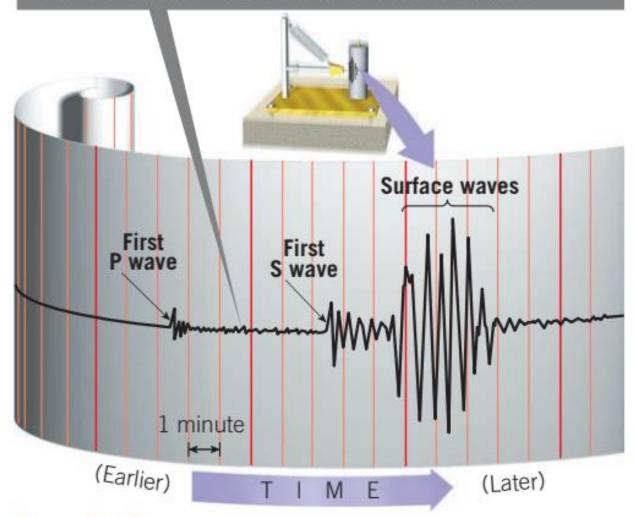


Figure 11.15
Typical seismogram

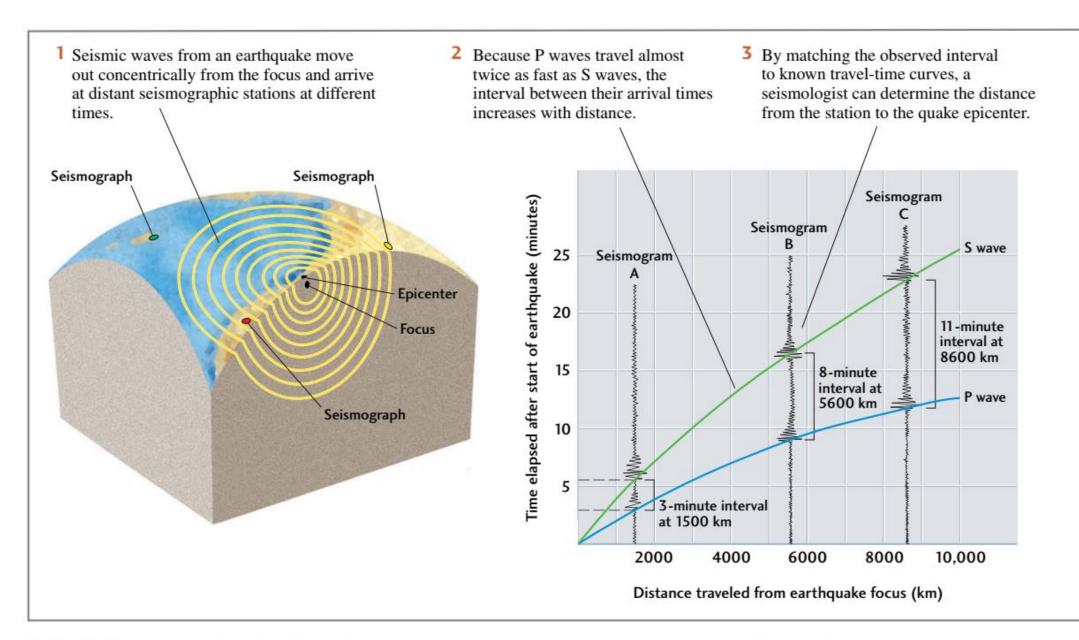


FIGURE 13.10 Readings from three or more seismographic stations can be used to determine the location of an earthquake's focus.

- Propagation of seismic wave depends on the physical properties of the material through which they propagate.
- The specific properties are
 - **Bulk modulus (K):** The ability of material to resist being <u>compressed</u>.
 - \triangleright Shear modulus (μ): The ability of material to resist being sheared.
 - \triangleright Density (ρ)

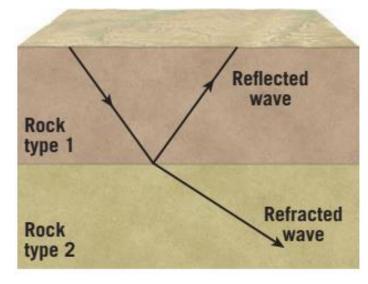
Body wave velocities: $V_P = \sqrt{\frac{\left(\frac{4}{3}\mu + K\right)}{\rho}}$ and $V_S = \sqrt{\frac{\mu}{\rho}}$

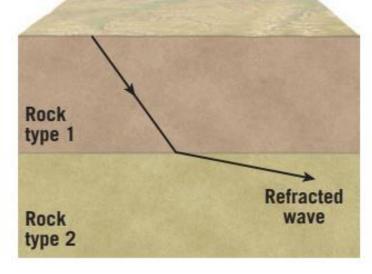
 Reflection and Refraction of the seismic waves as they pass through different material inside the Earth.

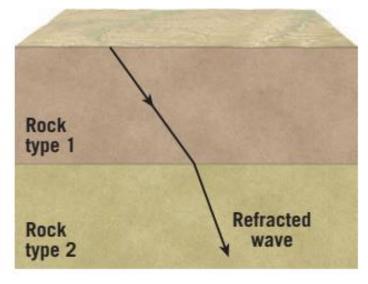
When seismic waves (rays) encounter a boundary between materials with different properties, such as air and water, the energy splits into reflected and refracted (bent) waves.

When the velocity of seismic waves increases as they pass from one layer into another, the waves refract (bend) toward the boundary separating the layers.

When the velocity of seismic waves decreases as they pass from one layer into another, the waves refract (bend) away from the boundary separating them.







P waves cannot reach the surface within the shadow zone because of the way they are refracted when they enter and leave the core. Although S waves reach the core, they cannot travel through its liquid outer region, and therefore never emerge beyond 105° from the focus.

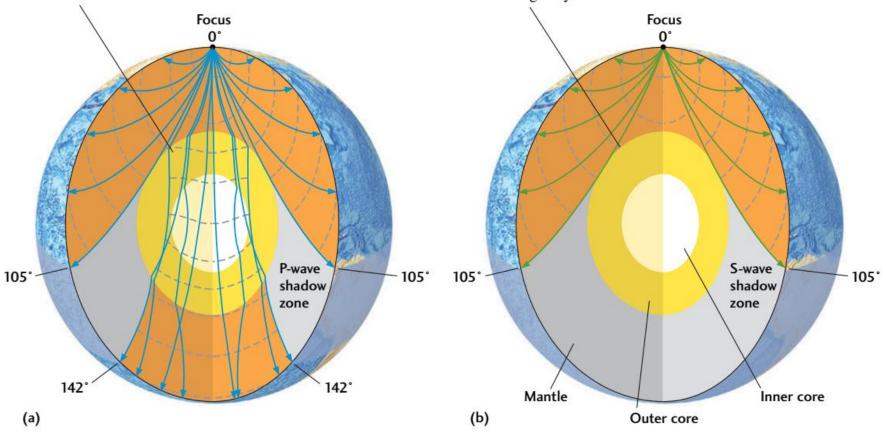


FIGURE 14.2 Earth's core creates P-wave and S-wave shadow zones. The ray paths of the seismic waves from an earthquake focus through Earth's interior are shown by solid lines (blue for P-waves, green for S-waves). The dashed lines show the progress of the waves at 2-minute intervals. Distances are measured in angular degrees from the earthquake focus. (a) The P wave shadow zone extends from 105° to 142°. (b) The larger S-wave shadow zone extends from 105° to 180°.

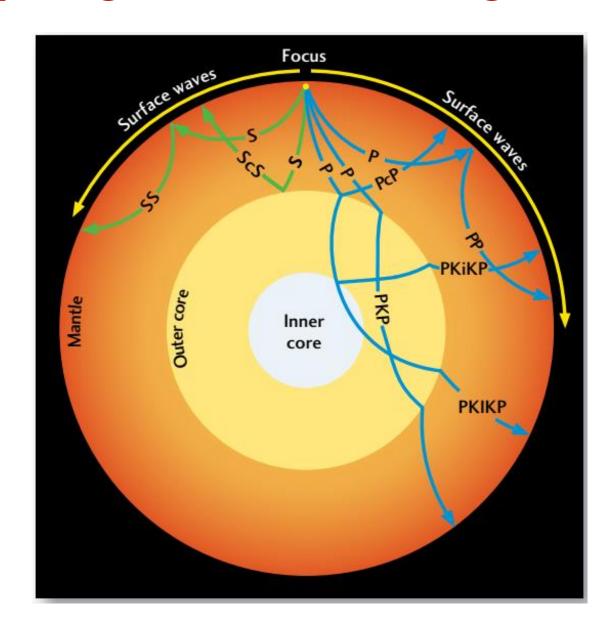


FIGURE 14.3 Seismologists use a simple labeling scheme to describe the various ray paths taken by seismic waves. PcP and ScS are compressional and shear waves, respectively, that are reflected by the core. PP and SS waves are internally reflected from Earth's surface. A PKP wave travels through the liquid outer core, a PKIKP wave travels through the solid inner core, and a PKiKP wave is reflected by the inner core. Surface waves propagate along Earth's outer surface, like waves on the surface of a pond.

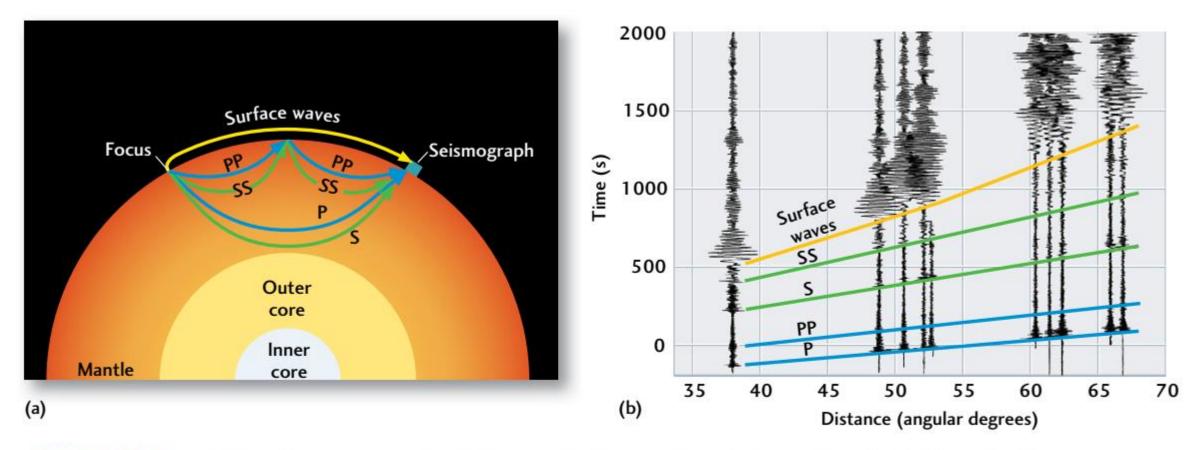


FIGURE 14.4 (a) P and S waves are refracted upward in the mantle and also can be reflected from Earth's surface. A seismic wave that has been reflected once from Earth's surface is labeled with a double letter (PP or SS). (b) Seismograms recorded at various distances from an earthquake focus in the Aleutian Islands, Alaska. The colored lines identify the arrival times of the P and S waves, the surface waves, and the PP and SS waves reflected from Earth's surface.

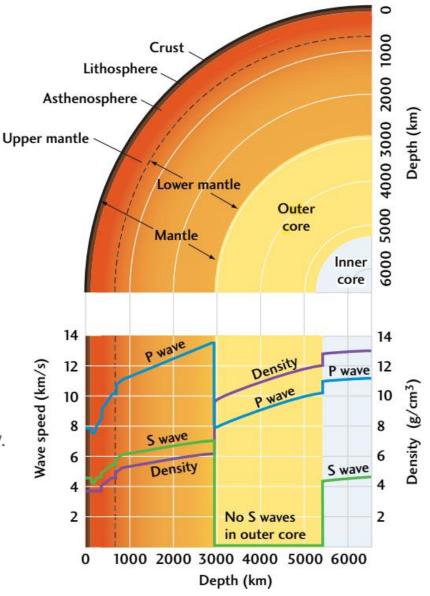


FIGURE 14.7 Earth's layering as revealed by seismology. The lower diagram shows changes in P-wave and S-wave velocities and rock densities with depth. The upper diagram is a cross section through Earth on the same depth scale, showing how those changes are related to the major layers (see also Figure 1.12).

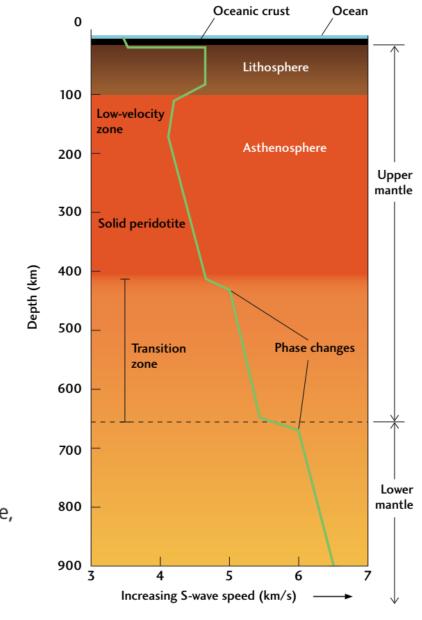


FIGURE 14.8 The structure of the mantle beneath old oceanic lithosphere, showing S-wave velocities to a depth of 900 km. Changes in S-wave velocity mark the strong, brittle lithosphere, the weak, ductile asthenosphere, and a transition zone, in which increasing pressure forces rearrangements of atoms into denser and more compact crystal structures (phase changes).