



A vapor–liquid phase equilibrium model for binary CO₂–H₂O and CH₄–H₂O systems above 523 K for application to fluid inclusions

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ABSTRACT

Accurate prediction of both volumetric and vapor–liquid phase equilibria of binary CO₂–H₂O and CH₄–H₂O mixtures with a single equation of state proves to be difficult. In this study we use an activity–fugacity model to predict the vapor–liquid phase equilibria above 523 K and adapt a Helmholtz model to calculate volumetric properties of these two binary systems. The average deviations of water content in the vapor phase from experimental data are 3.25% and 3.19% for the CO₂–H₂O and CH₄–H₂O mixtures, respectively, and the average deviations of gas solubility in liquid phase from experimental data are 4.29% and 3.50%, respectively. The model can find wide applications, and an example is given for the analysis of fluid inclusions in geochemistry.

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1. Introduction

CO₂–H₂O and CH₄–H₂O mixtures are typical geological fluids in the Earth's crust. Vapor–liquid phase equilibria [1–4] and volumetric (PVTx) properties [5–7] of the two systems are fundamental in the quantitative interpretation of boiling, immiscibility, gas solubility, and fluid migration, and also in the studies of fluid inclusions [8–12].

Due to highly non-ideal mixing properties of the binary CO₂–H₂O and CH₄–H₂O fluid mixtures, it is difficult to predict both the phase equilibria and volumetric properties simultaneously with an equation of state within experimental uncertainty over a large temperature–pressure region, such as 273–647 K for vapor–liquid equilibria, or 273–1273 K for volume. However, it is possible to use an equation of state to calculate the volume of these binary mixtures, and use another model to predict the vapor–liquid phase equilibria with different parameters for the vapor and liquid phases. Previously, Duan and Sun [2] developed a model for calculating CO₂ solubility in pure water from 273 to 533 K and from 1 to 2000 bar and Duan and Mao [3] presented a model for calculating the vapor–liquid phase equilibria of the CH₄–H₂O system from 273 to 523 K and from 1 to 2000 bar. However, high temperature (above 523 K) phase equilibrium models of these two systems are still lacking.

In this study, phase equilibrium models at high temperatures between 523 and 623 K and pressures up to 1500 bar (or P_c , the critical pressure, lower than 1500 bar) are established. First, competitive volumetric (PVTx) models for the CO₂–H₂O and CH₄–H₂O mixtures are briefly reviewed. Then it is shown that by combining the vapor–liquid phase equilibrium models reviewed and volumetric models that homogenization pressure, homogenization volume (or density) and isochores at given homogenization temperature and composition can be applied to the interpretation of microthermometric observations of fluid inclusions in minerals containing CO₂–H₂O and CH₄–H₂O.

2. Vapor–liquid phase equilibria of the binary CO₂–H₂O and CH₄–H₂O mixtures at high temperatures

2.1. Review of experimental data above 523 K

There are many experimental data on the vapor–liquid phase equilibria of the binary CO₂–H₂O and CH₄–H₂O mixtures below 523 K, and they are reviewed by Duan and Sun [2] and Duan and Mao [3]. However, there are a few experimental data above 523 K. In this study, we focus on the temperature range from 523 to 647 K.

2.1.1. CO₂–H₂O mixture

Over 270 measurements of the CO₂ solubility in pure water have been reported [13–21], covering a temperature range of 523–623 K. A few data points of Drummond [19] measured during the pres-

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Table 1
Parameters of Eq. (5).

	CO ₂ –H ₂ O	CH ₄ –H ₂ O
c_1	−0.14103545D+05	0.20222024D+05
c_2	0.16583148D+02	−0.24120165D+02
c_3	0.53326776D+07	−0.75058093D+07
c_4	−0.73088779D−02	0.10747756D−01
c_5	−0.75551470D+09	0.10421822D+10
c_6	−0.32657539D−01	−0.16556694D−01
c_7	0.16982222D−04	0.16641012D−04
c_8	0.12515020D+02	0.47915976D+01
c_9	0.49580878D−09	−0.18148280D−09

sure increasing process are not consistent with those measured in the pressure decreasing process. The deviation between them is 8–15%. The CO₂ solubilities from Todheide and Franck [17] are overestimated due to systematic analytical error, as pointed out by Blencoe [21], and their data are inconsistent with others. Therefore, all CO₂ solubility data points listed above are used in the parameterization, except those of Drummond [19] and Todheide and Franck [17].

About 180 measurements of water content in the vapor phase of the CO₂–H₂O system have been reported [13,16,17,20,21]. However, the data points from Todheide and Franck [17] are questionable due to systematic analytical error, and their data are excluded from our parameterization, which covers a temperature range from 523 to 623 K and pressure up to 1500 bar.

2.1.2. CH₄–H₂O mixture

Above 523 K, there are only two data sets for the vapor–liquid phase equilibrium of CH₄–H₂O system [22,23]. For the CH₄ solubility in water, the data of Sultanov et al. [22] are not consistent with those of Price [23]. However, the data points of Price [23] at low temperatures are consistent with other measurements. Experimental data can be plotted on the same isobaric composition–temperature or isothermal composition–pressure figure, from which consistency of experimental data can be seen. Therefore, the data points of Price [23] are used in the parameterization with temperature up to 627 K and pressure to about 2000 bar. As for the water content of the CH₄–H₂O mixture at high temperatures, only Sultanov et al. [22] made experimental measurements (48 data points), and their data are included in the parameterization of our model.

Table 2
Calculated composition deviations from experimental data at vapor–liquid phase equilibria.

References	T (K)	P (bar)	Nd	AAD (%)	MAD (%)
CO ₂ solubility in water in CO ₂ –H ₂ O mixture					
[13]	523.15–603.15	98–490	47	3.06	8.06
[14]	534.15–607.17	71.49–198.70	9	2.61	3.90
[15]	523.15–623.15	100–1400	17	5.25	16.94
[16]	523.15–623.15	100–1500	71	5.10	19.33
[20]	573.15	276.1–515.2	3	6.57	16.08
[21]	623.15	209.7–296	4	8.33	21.71
Water content in CO ₂ –H ₂ O mixture at phase equilibria					
[20]	573.15	516.5–566.1	3	5.93	11.7
[16]	523.15–623.15	100–1500	71	3.47	16.22
[13]	523.15–603.15	196–490	30	2.87	9.05
[21]	623.15	187–298.9	6	1.36	3.74
CH ₄ solubility in water in CH ₄ –H ₂ O mixture					
[23]	507–627	81.1–1972.6	45	3.50	10.85
Water content in CH ₄ –H ₂ O mixture at phase equilibria					
[22]	523.2633.2	98.1–1132.7	48	3.19	10.14

AAD: average absolute deviations calculated from this model; MAD: maximal absolute deviations calculated from this model; Nd: number of data points.

2.2. Phase equilibrium model as a function of temperature and pressure

2.2.1. Gas solubility in water

CO₂ (or CH₄) solubility in aqueous solutions depends on the balance between the chemical potential of CO₂ (or CH₄) in the liquid phase μ_i^l (i =CO₂ or CH₄ of the binary CO₂–H₂O or CH₄–H₂O mixtures, respectively) and that in the vapor phase μ_i^v . The chemical potential can be written in terms of fugacity in the vapor phase and activity in the liquid phase:

$$\begin{aligned}\mu_i^v(T, P) &= \mu_i^{v(0)}(T) + RT \ln f_i(T, P) \\ &= \mu_i^{v(0)}(T) + RT \ln y_i P + RT \ln \varphi_i(T, P)\end{aligned}\quad (1)$$

$$\begin{aligned}\mu_i^l(T, P) &= \mu_i^{l(0)}(T, P) + RT \ln a_i(T, P) \\ &= \mu_i^{l(0)}(T, P) + RT \ln m_i + RT \ln \gamma_i(T, P)\end{aligned}\quad (2)$$

where $\mu_i^{l(0)}$, the standard chemical potential of i in liquid, is defined as the chemical potential in hypothetically ideal solution of unit molality [24] and $\mu_i^{v(0)}$, the standard chemical potential in vapor, is the hypothetical ideal gas chemical potential when the pressure is set to 1 bar, y_i is mole fraction of i in the vapor phase, and φ_i is the fugacity coefficient of i in vapor phase.

For phase equilibrium $\mu_i^l = \mu_i^v$, and one obtains:

$$\ln \frac{y_i P}{m_i} = \frac{\mu_i^{l(0)}(T, P) - \mu_i^{v(0)}(T)}{RT} + \ln \gamma_i(T, P) - \ln \varphi_i(T, P) \quad (3)$$

where the fugacity coefficient of i in gaseous mixtures is generally calculated from an equation of state (EOS). An EOS is not only a function of temperature, pressure and volume or density, but also a non-linear function of composition. Obtaining accurate H₂O content in vapor phase is difficult both experimentally and theoretically (solving the non-linear EOS for vapor composition). Because the fugacity coefficient of i in gaseous mixtures differs little from that of pure i in the studied region and the little difference from this approximation is canceled in the later parameter fitting, therefore, $\ln \varphi_i$ can be approximated from the EOS of pure CO₂ or CH₄ [25] (see Appendix). y_i , the mole fraction of i in the vapor phase, is calculated from

$$y_i = 1 - y_{\text{H}_2\text{O}} = \frac{P - P_{\text{H}_2\text{O}}}{P} \quad (4)$$

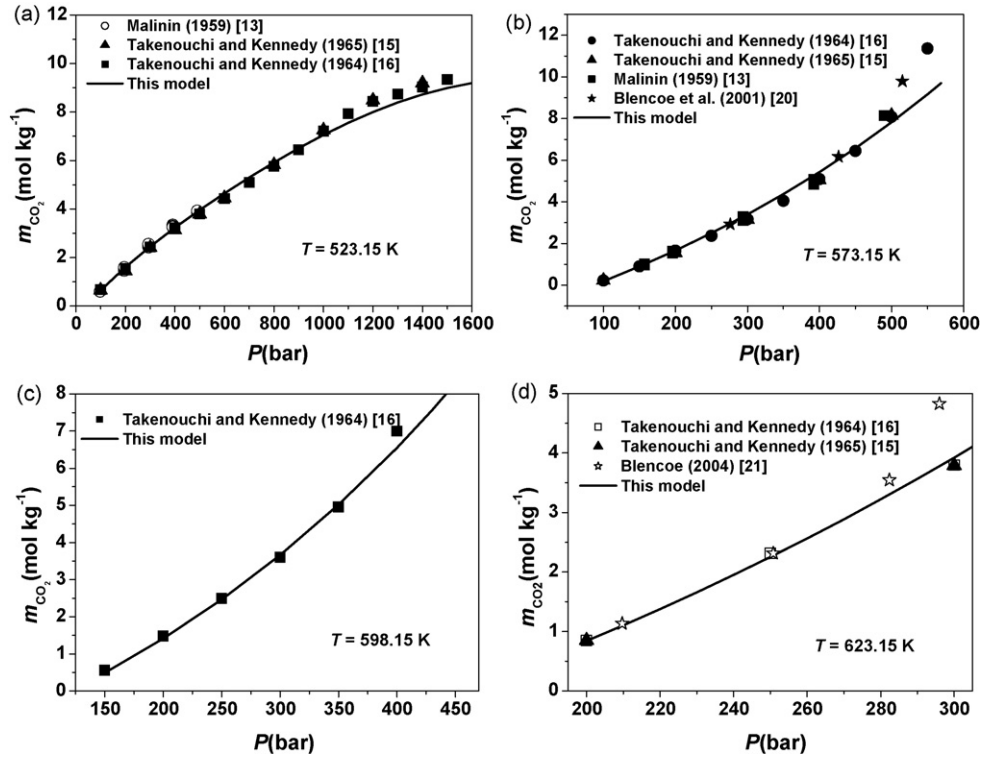


Fig. 1. CO₂ solubility in pure water at high temperatures: according to Blencoe [21], the critical pressures at 573.15, 598.15 and 623.15 K are 568.69, 469.49 and 301.92 bar, respectively.

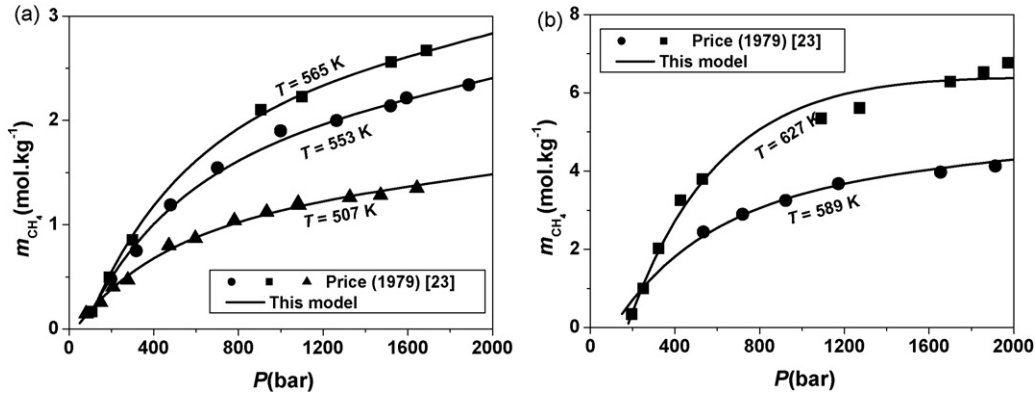


Fig. 2. CH₄ solubility in pure water at high temperatures.

where the partial pressure of water in vapor, $P_{\text{H}_2\text{O}}$, is approximated as the saturated pressure of pure water as that in our previous studies [26,27], and this approximation will lead to errors (up to 5%) for $\mu_i^{(0)}/RT$ and $\ln \gamma_i^l$. However, these errors tend to cancel out to a large extent in the parameterization and have little effect on the calculated gas solubility in water because in the calculation of gas solubility, only the chemical potential difference between liquid and vapor is important. In fitting to the experimental solubility

data, the error in the vapor phase can be transferred to the empirical parameters of the liquid phase. And the subsequent error in the liquid phase calculation is negligible. The saturated pressure of pure water here is from the equation of Wagner and Pruss [28].

$(\mu_i^{(0)}(T, P) - \mu_i^{v(0)}(T))/RT + \ln \gamma_i^l(T, P)$ in Eq. (3) is a function of temperature and pressure and can be calculated from the following equation:

$$\begin{aligned} \frac{\mu_i^{(0)}(T, P) - \mu_i^{v(0)}(T)}{RT} + \ln \gamma_i^l(T, P) = & c_1 + c_2T + \frac{c_3}{T} + c_4T^2 + \frac{c_5}{T^2} \\ & + c_6P + c_7PT + \frac{c_8P}{T} + c_9P^2T \end{aligned} \quad (5)$$

where c_1 – c_9 (Table 1) are obtained by regressing the reliable gas solubility data in the liquid phase for the binary CO₂–H₂O and CH₄–H₂O systems, as described in Section 2.1.

With these parameters, the CO₂ or CH₄ solubility in pure water above 523 K can be calculated; the average deviations from exper-

Table 3
Parameters of Eq. (7).

	CO ₂ –H ₂ O	CH ₄ –H ₂ O
a_1	0.15884702D – 01	0.47258673D – 01
a_2	–0.22882000D – 05	0.12938130D – 05
a_3	–0.13944627D – 04	–0.46620644D – 04
a_4	–0.61510488D + 01	–0.12663219D + 02
a_5	0.18428359D – 02	–0.95871477D – 04

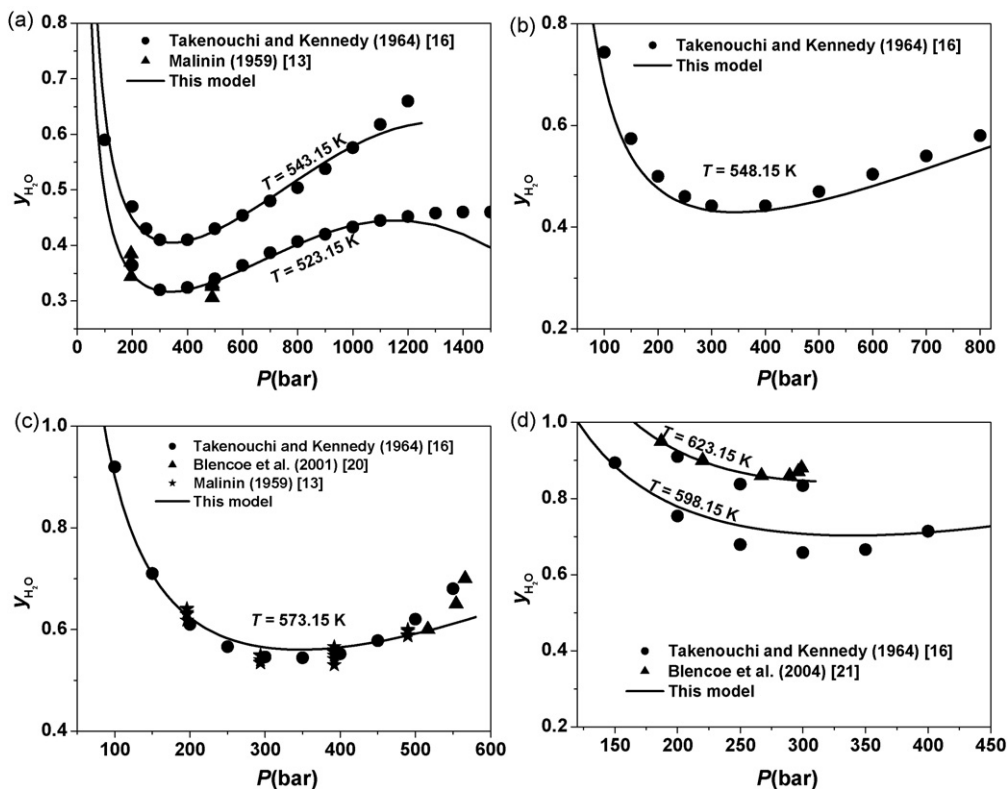


Fig. 3. Water content of the vapor phase for $\text{CO}_2\text{-H}_2\text{O}$ system at high temperatures: $y_{\text{H}_2\text{O}}$ is mole fraction of water in vapor phase; according to Blencoe [21], the critical pressures at 543.15, 548.15, 573.15, 598.15 and 623.15 K are 1253.00, 995.11, 568.69, 469.49 and 301.92 bar, respectively.

imental data are 4.29% and 3.50% for the binary $\text{CO}_2\text{-H}_2\text{O}$ and $\text{CH}_4\text{-H}_2\text{O}$ mixtures, respectively, and they are within experimental uncertainties. Table 2 shows the average and maximal deviations of the model from each data set. Figs. 1 and 2 show the comparisons

between the experimental results and model predictions. As can be seen in Figs. 1 and 2, the molality of CO_2 or CH_4 in water increases with pressures at given temperatures, and the experimental gas solubility data are accurately reproduced by Eq. (5) except in the

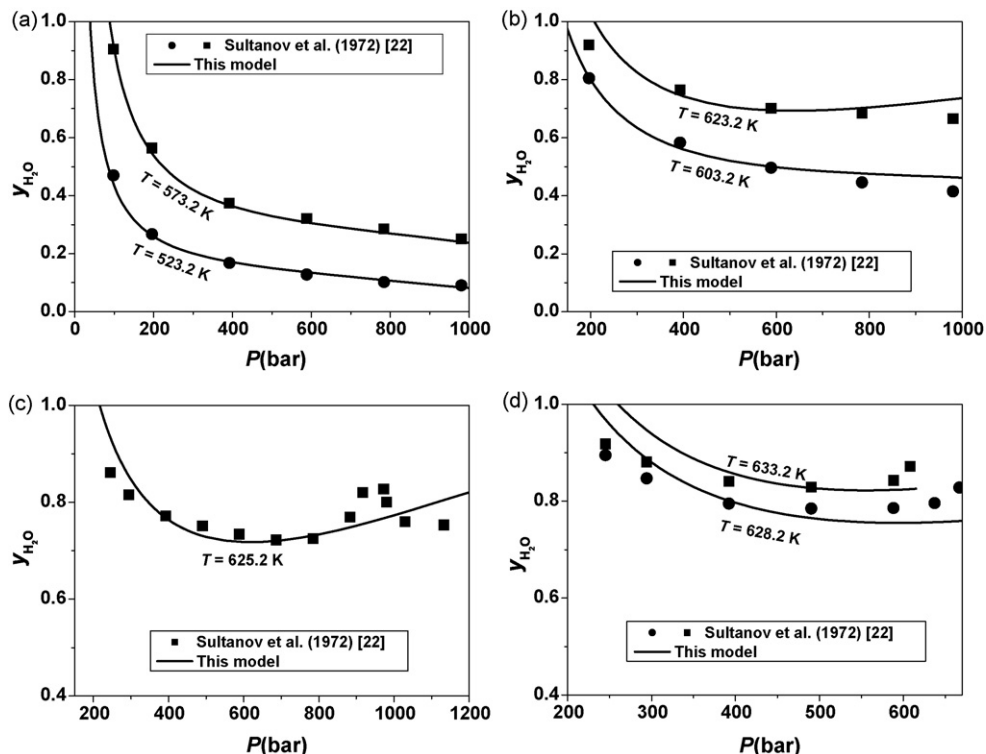


Fig. 4. Water content of the vapor phase for $\text{CH}_4\text{-H}_2\text{O}$ system at high temperatures: $y_{\text{H}_2\text{O}}$ is mole fraction of water in vapor phase.

Table 4Models adopted for the CO₂–H₂O and CH₄–H₂O fluid inclusions.

	CO ₂ –H ₂ O	T (K)	P (bar)	CH ₄ –H ₂ O	T (K)	P (bar)
Vapor–liquid phase equilibria models	[2]	273–523	<2,000	[3]	273–523	<2,000
	This model	523–623	<1,500	This model	523–623	<2,000
Volumetric models	[5]	273–1273	<10,000	[6]	273–1273	<10,000
	[43]	1273–2573	<100,000			

critical region, where deviations increase to over 10%. Eq. (5) can be used to satisfactorily calculate homogenization pressure of fluid inclusions.

2.2.2. Water content in vapor phase

Water content of the binary CO₂–H₂O or CH₄–H₂O mixtures in the vapor phase above 523 K can be estimated using the following equation:

$$y_{\text{H}_2\text{O}} = \frac{\gamma_{\text{H}_2\text{O}} x_{\text{H}_2\text{O}} P_{\text{H}_2\text{O}}^S}{\phi_{\text{H}_2\text{O}} P} \exp \left(\frac{v_{\text{H}_2\text{O}}^l (P - P_{\text{H}_2\text{O}}^S)}{RT} \right) \quad (6)$$

where $y_{\text{H}_2\text{O}}$ denotes the mole fraction of water in the vapor phase, $x_{\text{H}_2\text{O}}$ is the mole fraction of H₂O in the liquid phase, $\gamma_{\text{H}_2\text{O}}$ is the activity coefficient of water in the liquid phase, $P_{\text{H}_2\text{O}}^S$ the saturation pressure (bar) of water calculated from the correlation of Wagner and Pruss [28], $v_{\text{H}_2\text{O}}^l$ is the molar volume of liquid water (cm³ mol^{−1}) approximated as the saturated liquid phase volume of water calculated from the equation of Wagner and Pruss [28], and $\phi_{\text{H}_2\text{O}}$ refers to the fugacity coefficient of water in the vapor phase.

The remaining parameters of Eq. (6) for the calculation of water content in the vapor phase are $\phi_{\text{H}_2\text{O}}/\gamma_{\text{H}_2\text{O}} x_{\text{H}_2\text{O}}$, which are functions of temperature and pressure and can be calculated from the follow-

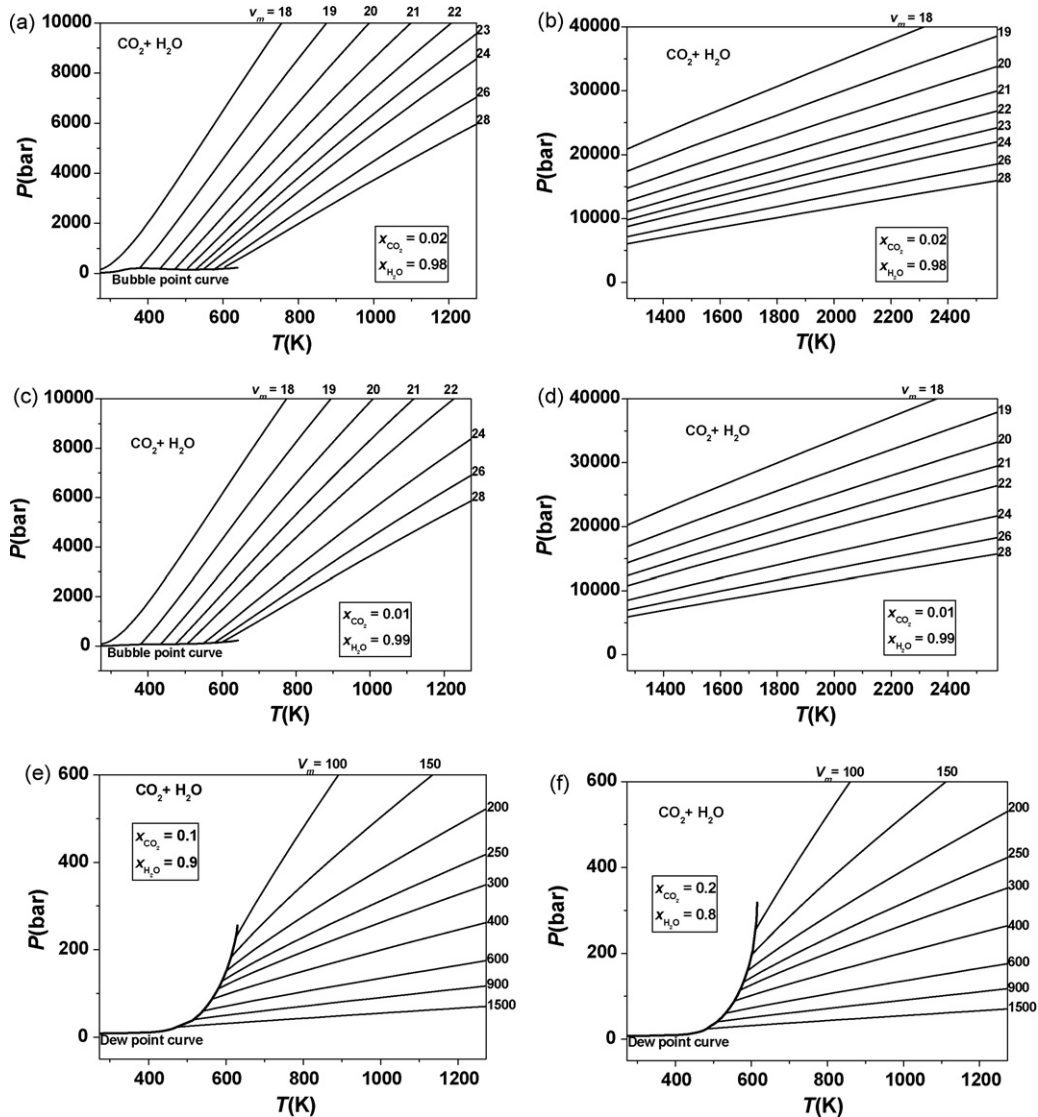


Fig. 5. Isochores of CO₂–H₂O system: bubble point curve and dew point curve above 523 K are from this model, and those below 523 K are from the model of Duan and Sun [2]; isochors between 273 and 1273 K are from the PVTx model of Paulus and Penoncello [5], and isochors between 1273 and 2573 K are from the volume model of Duan and Zhang [43]; unit of V_m is cm³ mol^{−1}.

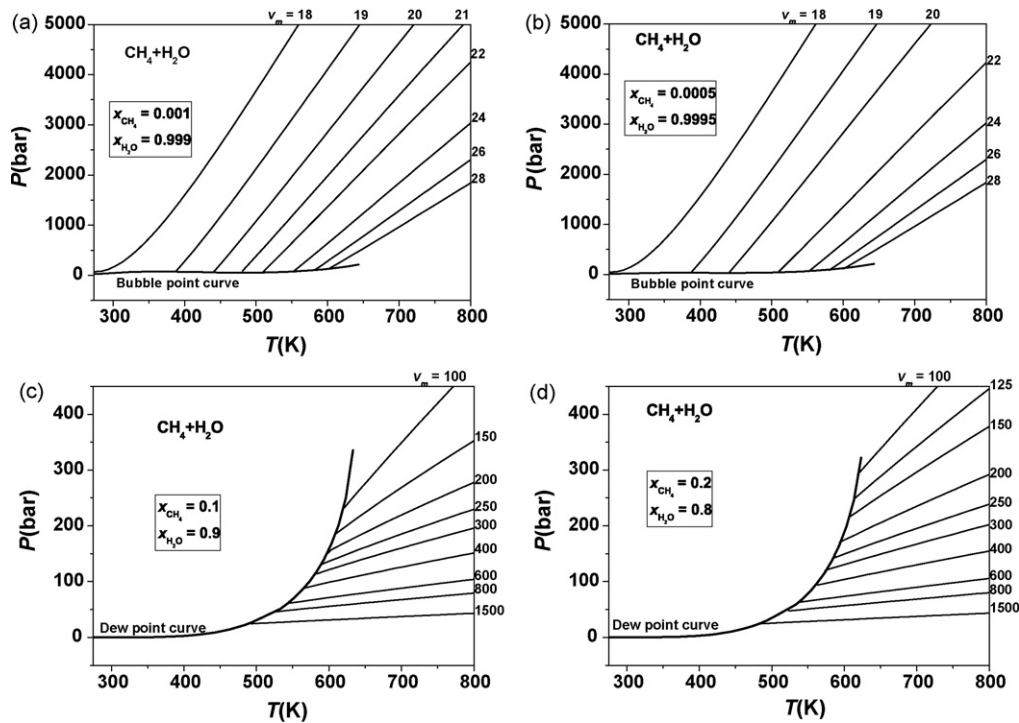


Fig. 6. Isochores of CH₄–H₂O system: bubble point curve and dew point curve above 523 K are from this model, and those below 523 K are from the model of Duan and Mao [3]; isochores are from the PVTx model of Kunz et al. [6]; unit of v_m is cm³ mol^{−1}.

ing semi-empirical equation:

$$\frac{\varphi_{H_2O}}{\gamma_{H_2O}x_{H_2O}} = \exp \left(a_1(P - P_{H_2O}^S) + a_2(P - P_{H_2O}^S)^2 + a_3(P - P_{H_2O}^S)T \right. \\ \left. + \frac{a_4(P - P_{H_2O}^S)}{T} + \frac{a_5(P - P_{H_2O}^S)^2}{T} \right) \quad (7)$$

where a_1 – a_5 (Table 3) are obtained by regressing the water content data in the vapor phase discussed above for the CO₂–H₂O or CH₄–H₂O systems. The water content in the vapor phase can be reproduced by Eqs. (6) and (7), as shown by Figs. 3 and 4. Table 2 also shows the average and maximal deviations of the water content from each experimental set. The average deviations of the water content from experimental data in the vapor phase are 3.25% and 3.19% for the CO₂–H₂O and CH₄–H₂O mixtures, respectively. When total pressure P approaches the saturation pressure of pure

water, y_{H_2O} approaches 1. As CO₂–H₂O or CH₄–H₂O fluid inclusions homogenize to the vapor phase at high temperatures, homogenization pressures can be obtained from Eqs. (6) and (7) by an iterative method.

3. Volumetric (PVTx) models for the binary CO₂–H₂O and CH₄–H₂O mixtures

There are five types of models for the calculation of volumetric properties of the CO₂–H₂O and CH₄–H₂O systems: (1) equations of state (EOS) [5–7,29–43]; (2) liquid density models [44–47]; (3) partial molar volume models [48–51]; (4) apparent molar volume models [45,52–56]; and (5) excess volume models [29,57–59]. Each model has its strength and weakness. For the binary CO₂–H₂O and CH₄–H₂O fluid mixtures, Helmholtz energy equations of state are the best among these five types of models. These equations of state can not only reproduce the existing

Table 5
Calculated results of the CO₂–H₂O and CO₂ fluid inclusions.

	P–T relation					
	T (K)	P (bar)	T (K)	P (bar)	T (K)	P (bar)
CO ₂ –H ₂ O inclusion (V+L–L)						
Th = 573 K	573.00	143.14	823.15	2773.35	1073.15	5295.87
$x_{CO_2} = 0.015$	623.15	660.04	873.15	3292.64	1123.15	5777.38
Ph = 143.14 bar	673.15	1187.73	923.15	3804.93	1173.15	6251.42
Vh = 25.68 cm ³ mol	723.15	1718.60	973.15	4309.69	1223.15	6718.24
Dh = 0.7167 g cm ^{−3}	773.15	2248.09	1023.15	4806.67	1273.15	7178.10
CO ₂ inclusion (V+L–L)						
Th = 295 K	295	59.82	600	1359.52	950	2715.17
Ph = 59.82 bar	300	79.82	650	1562.64	1000	2898.03
Vh = 58.48 cm ³ mol	350	293.4	700	1762.28	1050	3078.71
Dh = 0.7526 g cm ^{−3}	400	511.73	750	1958.64	1100	3257.35
	450	728.53	800	2151.90	1150	3468.46
	500	942.41	850	2342.29	1200	3649.73
	550	1152.79	900	2529.98	1250	3829.62

Note: Th, Ph, Vh and Dh denote homogenization temperature, pressure, mole volume and density, respectively; “V+L=L” denotes homogenization to liquid phase; x_{CO_2} is mole fraction CO₂ in CO₂–H₂O mixture. Used models or equations of state are the same as those in Fig. 7.

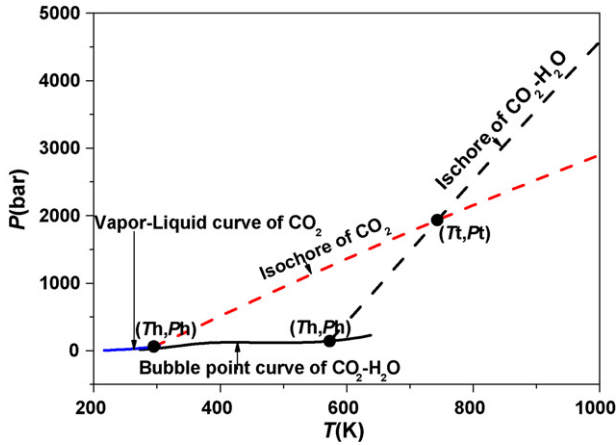


Fig. 7. Isochores of $\text{CO}_2\text{-H}_2\text{O}$ and CO_2 fluid inclusions in Sanja lode gold deposit, Jiaodong peninsula of eastern China: T_h and P_h denote homogenization temperature and pressure, respectively, values see Table 5; T_t and P_t denote trapped temperature and pressure, respectively. The two inclusions homogenize to liquid phase. Bubble point curve of $\text{CO}_2\text{-H}_2\text{O}$ fluid above 523 K is from this model, and that below 523 K is from the model of Duan and Sun [2]. Isochores of $\text{CO}_2\text{-H}_2\text{O}$ fluid are from the PVTx model of Paulus and Penoncello [5]. Vapor-liquid curve and isochores of CO_2 fluid are from equation of state of Span and Wagner [63]. Calculated $T_t = 743.32$ K and $P_t = 1932.57$ bar.

data, but also can be extrapolated to predict volumetric properties in the high temperature–pressure region with reasonable trend as pointed out by Span and Wagner [60], which can find wide applications in the study of fluid inclusions. Kunz et al. [6] and Paulus and Penoncello [5] respectively developed an equation of state in terms of Helmholtz energy for the volume of the $\text{CO}_2\text{-H}_2\text{O}$ fluid mixture. It should be noted that when CO_2 is dissolved in liquid water up to 473 K, the density of solutions will increase [61]. But the volumetric model of Kunz et al. [6] fails in this aspect. Therefore, the equation of state of Paulus and Penoncello [5] is chosen for calculating the volume of the $\text{CO}_2\text{-H}_2\text{O}$ fluid mixture, covering the T – P range of 273–1273 K and 1–10,000 bar, where equations of pure H_2O are from Wagner and Prüss [62] and equations of pure CO_2 are from Span and Wagner [63]. Above 1273 K, the equation of state of Duan and Zhang [43] based on the data of molecular dynamics simulations is used to calculate isochores of the $\text{CO}_2\text{-H}_2\text{O}$ fluid mixtures. The equation of state of Kunz et al. [6] is adopted for calculating the volume of the $\text{CH}_4\text{-H}_2\text{O}$ fluid mixture.

4. Application in fluid inclusions

The model presented here can find wide applications, such as predicting fluid boiling of hydrothermal fluids, calculating CO_2 solubility related to CO_2 geological sequestration, and applying to the study of fluid inclusions. Applications to fluid inclusions are discussed below.

The vapor–liquid phase equilibrium models from 273 to 623 K and volumetric models from 273 to 2573 K adopted for the study of fluid inclusions are listed in Table 4. Usually, temperature and composition of the fluid inclusions can be obtained from microthermometric observations, but homogenization density (volume), homogenization pressure, and isochores must be acquired from thermodynamic models, as demonstrated in this study by combining the vapor–liquid phase equilibrium and volumetric models (Figs. 5 and 6). From Figs. 5 and 6, it can be seen that isochores in liquid phase are a bit curved below 400 K and 1000 bar, but are almost linear at higher temperatures.

Here an actual example is given to illustrate how to use this model for the study of fluid inclusions. There are two kinds of original fluid inclusions in Sanja lode gold deposit, Jiaodong peninsula

of eastern China: one is $\text{CO}_2\text{-H}_2\text{O}$ inclusion with CO_2 mole fraction of 0.015, and total homogenization temperature $T_h(\text{CO}_2\text{-H}_2\text{O})$ is 573 K; the other is pure CO_2 inclusion, and homogenization temperature $T_h(\text{CO}_2)$ is 295 K. The two kinds of inclusion finally homogenize to liquid phase. Based on these microthermometric data, corresponding homogenization pressure (density or volume) and isochores of $\text{CO}_2\text{-H}_2\text{O}$ mixture is obtained from the phase equilibrium and volumetric models. Saturation curve and isochores of pure CO_2 fluid are from equation of state of Span and Wagner [63]. Calculated results of the $\text{CO}_2\text{-H}_2\text{O}$ and CO_2 fluid inclusions are listed in Table 5. Isochores are plotted in Fig. 7, from which trapped temperature (T_t) and pressure (P_t) of ore fluids are determined: $T_t = 743.32$ K and $P_t = 1932.57$ bar. Therefore, temperature–pressure conditions of the trapped ore fluids are obtained from the model.

5. Conclusions

A phase equilibrium model of the binary $\text{CO}_2\text{-H}_2\text{O}$ and $\text{CH}_4\text{-H}_2\text{O}$ fluid mixtures between 523 and 623 K and pressures up to 1500 bar (or the critical pressure, lower than 1500 bar) are established within or about experimental uncertainties. Volumetric (PVTx) models for the $\text{CO}_2\text{-H}_2\text{O}$ and $\text{CH}_4\text{-H}_2\text{O}$ mixtures are reviewed, and the competitive volumetric models of the $\text{CO}_2\text{-H}_2\text{O}$ [5,43] and $\text{CH}_4\text{-H}_2\text{O}$ [6] fluid mixtures are adapted. Results of this study, together with previous phase equilibrium models of Duan and Sun [2] and Duan and Mao [3], and the volumetric models of the $\text{CO}_2\text{-H}_2\text{O}$ [5,43] and $\text{CH}_4\text{-H}_2\text{O}$ [6] mixtures are applied to the fluid inclusions, such that homogenization pressure, homogenization volume or density, and isochores at given homogenization temperatures and compositions can be obtained. Online calculations can be made on: <http://www.geochem-model.org/research/fluidinc>.

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Appendix A. Appendix

Equation of state for pure CO_2 and CH_4

$$Z = \frac{P_r V_r}{T_r} = 1 + \frac{a_1 + a_2/T_r^2 + a_3/T_r^3}{V_r} + \frac{a_4 + a_5/T_r^2 + a_6/T_r^3}{V_r^2} + \frac{a_7 + a_8/T_r^2 + a_9/T_r^3}{V_r^4} + \frac{a_{10} + a_{11}/T_r^2 + a_{12}/T_r^3}{V_r^5} + \frac{a_{13}}{T_r^3 V_r^2} \left(a_{14} + \frac{a_{15}}{V_r^2} \right) \exp \left(-\frac{a_{15}}{V_r^2} \right) \quad (\text{A1})$$

$$P_r = \frac{P}{P_c}; \quad T_r = \frac{T}{T_c} \quad (\text{A2})$$

$$V_r = \frac{V}{V_c}; \quad V_c = \frac{RT_c}{P_c} \quad (\text{A3})$$

where P_r , T_r and V_r are reduced pressure, temperature and volume, respectively; P_c and T_c are critical pressure and critical temperature, respectively; R is universal gas constant (83.14472 bar $\text{cm}^3 \text{K}^{-1} \text{mol}^{-1}$); V is molar volume. Note that V_c is not the real critical volume. The parameters of the EOS are referred to Duan et al. [25]. Critical properties of CO_2 are: $T_c = 304.2$ K;

Table A1
Fugacity coefficient comparison for pure CO₂ and CH₄.

T (K)	P (bar)			
	100	250	500	1000
CO ₂				
523	0.951 (0.947)	0.898 (0.890)	0.868 (0.860)	0.970 (0.962)
573	0.973 (0.968)	0.947 (0.938)	0.943 (0.934)	1.065 (1.058)
623	0.988 (0.984)	0.980 (0.972)	0.997 (0.988)	1.134 (1.129)
CH ₄				
523	1.008 (1.008)	1.034 (1.035)	1.117 (1.121)	1.413 (1.428)
573	1.016 (1.016)	1.051 (1.052)	1.139 (1.145)	1.426 (1.444)
623	1.021 (1.022)	1.061 (1.064)	1.153 (1.161)	1.428 (1.450)

Note: The first value of fugacity coefficient is calculated from equation of state of Duan et al. [25]; the second value of fugacity coefficient in parentheses is calculated from equation of state NIST recommended: equation of state of Span and Wagner [63] is for CO₂ and equation of state of Setzmann and Wagner [64] is for CH₄.

$P_c = 73.825$ bar; the critical properties of CH₄ are: $T_c = 190.6$ K; $P_c = 46.41$ bar. Fugacity coefficient of pure CO₂ or CH₄ can be derived from Eq. (A1):

$$\ln \varphi(T, P) = Z - 1 - \ln Z + \frac{a_1 + a_2/T_r^2 + a_3/T_r^3}{V_r} + \frac{a_4 + a_5/T_r^2 + a_6/T_r^3}{2 V_r^2} + \frac{a_7 + a_8/T_r^2 + a_9/T_r^3}{4 V_r^4} + \frac{a_{10} + a_{11}/T_r^2 + a_{12}/T_r^3}{5 V_r^5} + \frac{a_{13}}{2 T_r^3 a_{15}} \left[a_{14} + 1 - \left(a_{14} + 1 + \frac{a_{15}}{V_r^2} \right) \exp \left(- \frac{a_{15}}{V_r^2} \right) \right] \quad (\text{A4})$$

Fugacity coefficients of pure CO₂ or CH₄ from equation of state of Duan et al. [25] are compared with those from equations of state NIST recommended (Table A1): EOS of Span and Wagner [63] is for CO₂ and EOS of Setzmann and Wagner [64] is for CH₄. It can be seen that the calculated results from these equations are in good agreement. However, equation of state of Duan et al. [25] is simple and thus is used in this study.

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