### The testability of maximum magnitude estimates







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#### 1) We would like to test estimates of $M_{\text{max}}$ .

Earthquakes of unexpectedly large magnitude can cause major disasters.

Estimates of the maximum earthquake magnitude ( $M_{\rm max}$ ) in a region, or for a fault, are crucial for seismic hazard and risk assessment.

Can we statistically and rigorously test them?

They are commonly provided as single values, without a probability distribution. Also, they refer to the maximum possible magnitude ever expected, during an infinite period of time (for an overview, see Wheeler, R.L., 2009, USGS Open-File Report 2009-1018). However, a feasible test of  $M_{\rm max}$  can only be performed during a finite period.



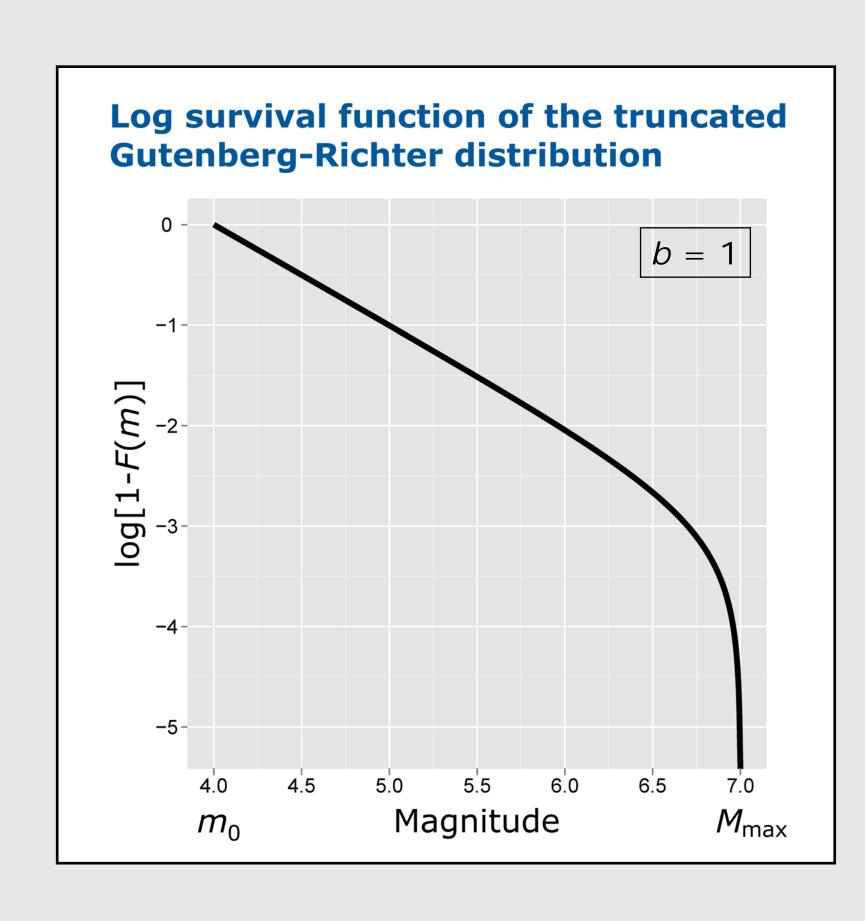
Tsunami
whirpool near
Oarai City,
Japan, after
the *M*=9
Tohoku
earthquake.

# 2) We will assume that earthquake magnitudes follow a doubly truncated Gutenberg-Richter distribution.

$$F(m) = \frac{e^{-\beta m_0} - e^{-\beta m}}{e^{-\beta m_0} - e^{-\beta \hat{M}_{ma}}}$$

#### Where:

- m is magnitude,
- β is blog(10), where b is the b-value of the Gutenberg-Richter distribution,
- $\cdot m_0$  is the magnitude of completeness, and
- $\hat{M}_{max}$  is the upper truncation, that is, a given estimate of  $M_{max}$ .



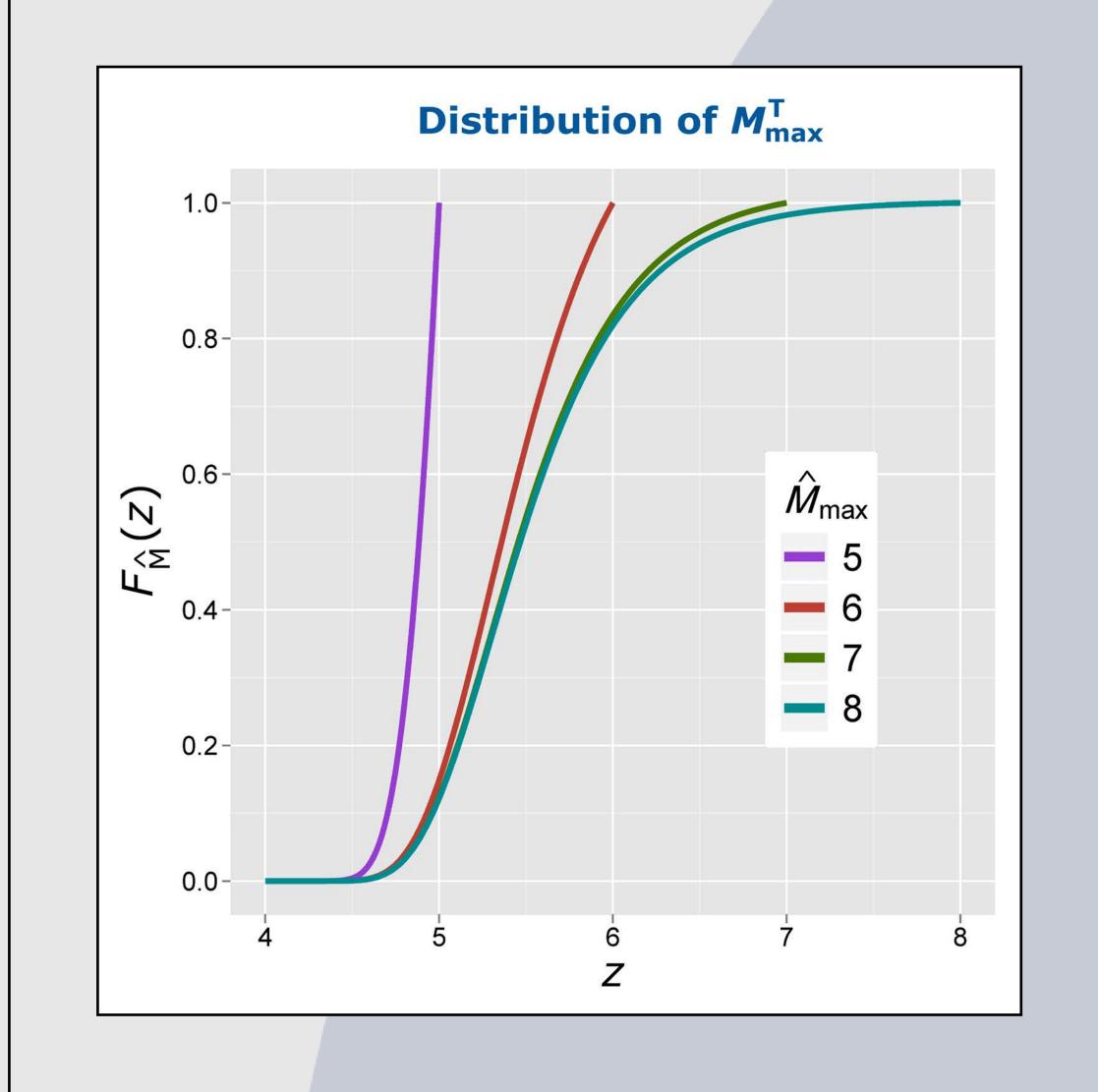
## 3) This implies that we can expect to observe certain $M_{\text{max}}$ values in a finite time period...

Let us say that during a finite time period, T, a total of n earthquakes are observed above  $m_0$ .

The magnitude of the largest one is a random variable, denoted  $M_{\text{max}}^T$ . The longer the T, the larger n can be, and the more likely it is to observe a larger  $M_{\text{max}}^T$ .

The corresponding probability distribution of  $M_{\text{max}}^T$  is given, generically, by

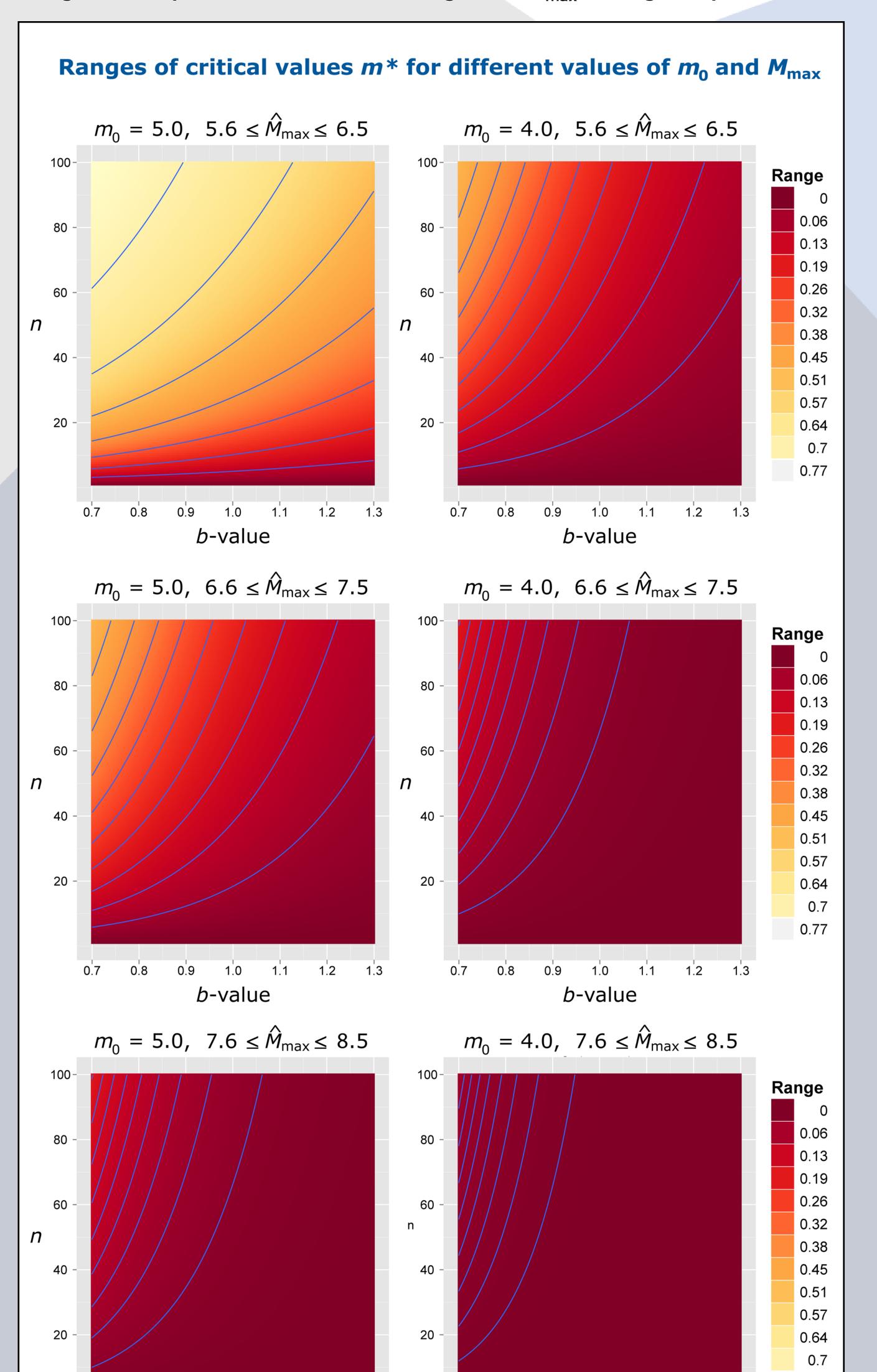
$$F_{\hat{M}}(z) = P(M_{\max}^T < z) = [F(m)]^n.$$



### 4) ...which are not very sensitive, in the short term, to the long-term estimate of $M_{\text{max}}$ .

Let  $m^*$  be the critical value, i.e., the value at which  $P(M^T_{\text{max}} < m^*) \le 0.05$ . We measure the change in  $m^*$  with respect to  $\hat{M}_{\text{max}}$ , for all combinations of  $0.7 \le b \le 1.3$ ,  $5.6 \le \hat{M}_{\text{max}} \le 8.5$ ,  $1 \le n \le 100$ , and  $m_0 \in \{4,5\}$ .

 $F_{\hat{M}}(z)$  is most sensitive in regions of high earthquake frequency and low b-values, because in these regions there is a greater chance of observing larger earthquakes (closer to the long-term  $M_{\text{max}}$ ) in a given period.



0.7 0.8 0.9 1.0 1.1 1.2 1.3

*b*-value

0.7 0.8 0.9 1.0 1.1 1.2 1.3

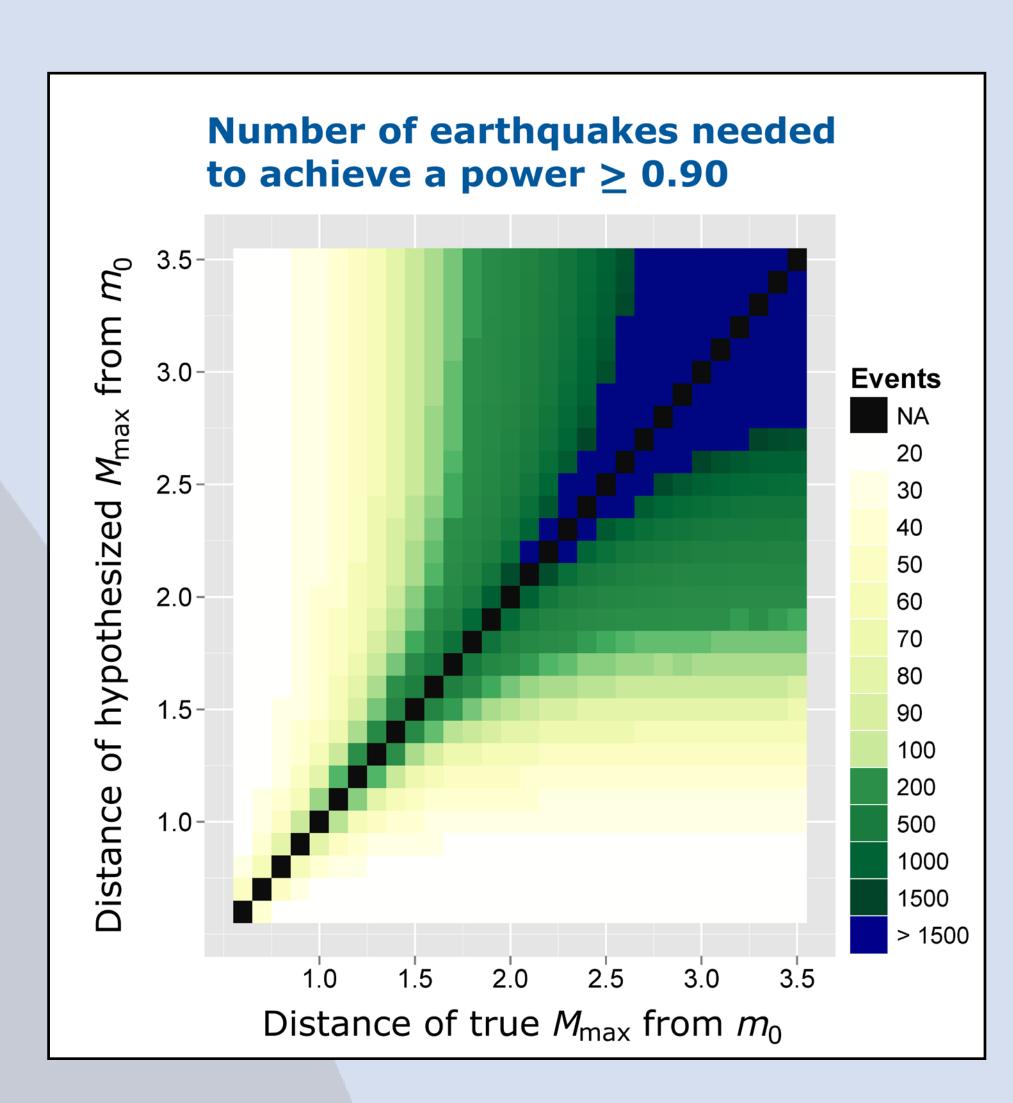
*b*-value

# 5) The minimum number of earthquakes required for a powerful test depends on the difference between $M_{\text{max}}$ and the completeness magnitude.

If the observed maximum magnitude is too unlikely, given the assumed earthquake magnitude distribution, we can reject it.

We want to know how many earthquakes must occur for our test to correctly reject a false  $\hat{M}_{\text{max}}$  with high probability (at least 0.90).

Using a simulation study, we notice that as the true  $M_{\rm max}$  gets further away from  $m_0$ , we will likely never have enough earthquakes in a reasonable time period to test  $\hat{M}_{\rm max}$ , unless  $\hat{M}_{\rm max}$  is grossly underestimated.



#### 6) Conclusions:

Testing  $M_{\text{max}}$  in a reasonable time period may not be possible in most regions.

Using our approach, estimates of  $M_{\rm max}$  that are provided without a probability distribution, and for long time periods, are testable only in certain ideal situations (high earthquake frequency, low b-value, and a relatively low  $M_{\rm max}$ ).

In the future, it would be beneficial to further examine the impact of  $M_{\rm max}$  on the expected ground motions, to see to what degree hazard assessment is affected by this apparently nontestable parameter.

