

Announcement

- No class on Monday (holiday)
- A make-up class on Thursday (here at Hebb 100, at 13:00 pm)
- Survey: your feedback (best over the long weekend)

https://ubc.ca1.qualtrics.com/jfe/form/SV_1ZCDtHXC0sLIIVA



Lecture 13.

Energy balance and earth climate

- Wien's law

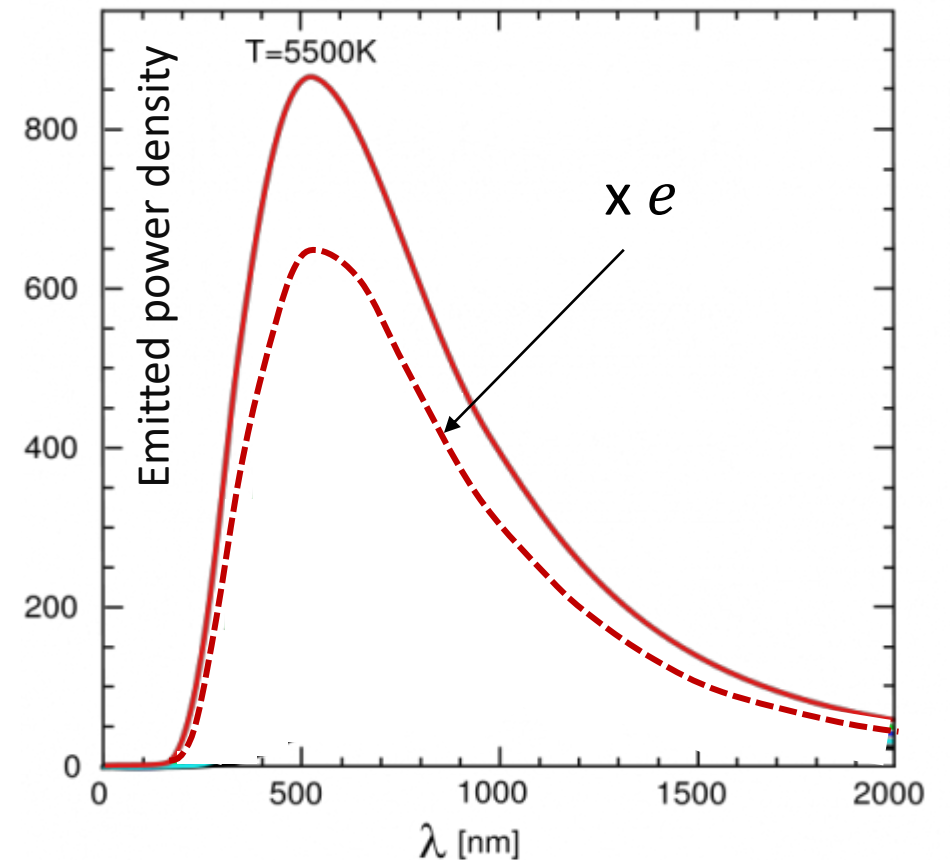
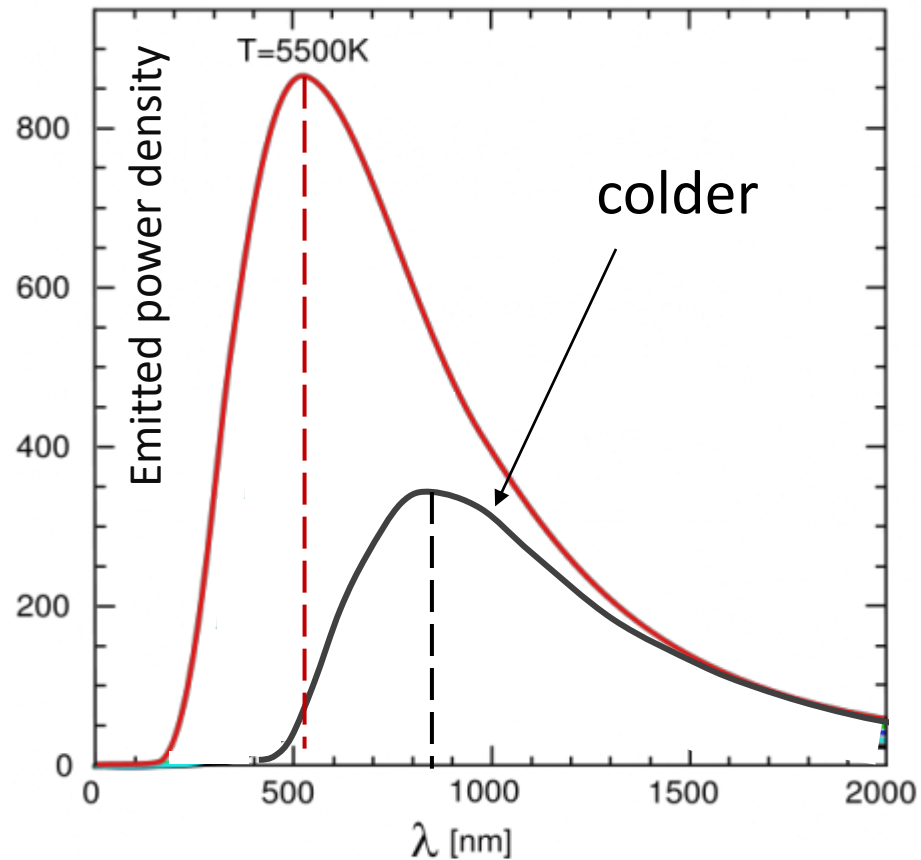
$$\lambda_{max} = \frac{b}{T}$$

- Stefan-Boltzmann law

$$H = \frac{Q}{\Delta t} = Ae\sigma T^4$$

Last Time

- Emissivity



Q: Yoltar heats their little planet (far from any stars) with a 1GW heater. If they wish to double the equilibrium surface temperature of their planet, they should increase the power of their heater to:

$$H_{in} = H_{out}$$

$$H = Ae\sigma T^4$$

- A. 1.21 GW
- B. 2 GW
- C. 4 GW
- D. 8 GW
- E. 16 GW

Hint: where does the energy from the heater go?



$$H = Ae\sigma T^4$$



Q: Yoltar heats their little planet (far from any stars) with a 1GW heater. If they wish to double the equilibrium surface temperature of their planet, they should increase the power of their heater to:

Steady state:
power from heater = power radiated

To double T, need 16 x P

- A. 1.21 GW
- B. 2 GW
- C. 4 GW
- D. 8 GW
- E. 16 GW



*Hint: where does
the energy from
the heater go?*



A harder problem, but *really* interesting!

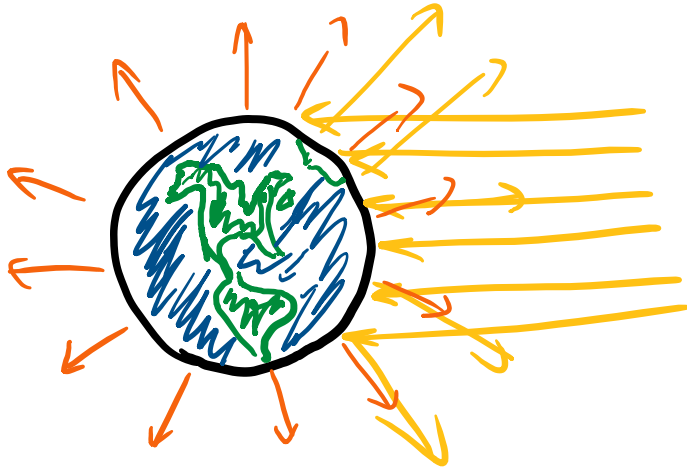
A planet with radius $r = 6400$ km lies at a distance $R = 150,000,000$ km from a yellow star with temperature $T = 5700\text{K}$ and radius $R_s = 695,000$ km.

Which planet are we talking about?

The earth!

Let's use our knowledge of thermodynamics to estimate the surface temperature of our planet!

Energy balance for the earth



- Model:



- Key relation for steady state heat flow:

$$H_{in} = H_{out}$$



- H_{in} = absorbed sunlight = ??

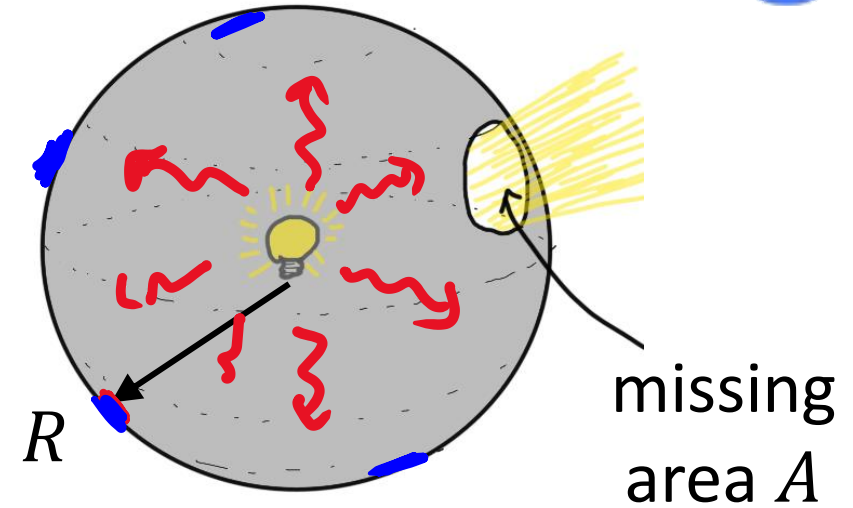
- H_{out} = IR radiation = $Ae\sigma T^4$

➤ The emissivity e of earth is close to 1.

What is H_{in} ?

Q: A light bulb producing 100 W of radiation is placed at the center of a sphere of perfectly absorbing material, with radius R . A hole is cut into the sphere, removing an area A of material. Assume the light from the bulb spreads out uniformly in all directions.

What is the rate of energy flow through the hole?



$$P_{\text{bulb}} \frac{A_{\text{catch}}}{A_{\text{sphere}}}$$

distributed

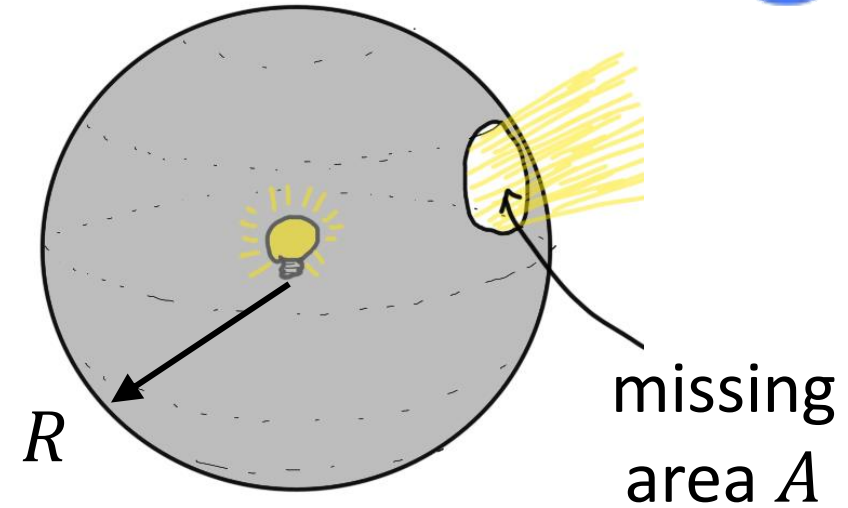
- A. 100W
- B. $100\text{W} \times A$
- C. $100\text{W} \times A/R^2$
- D. $100\text{W} \times A/(\pi R^2)$
- ☒ E. $100\text{W} \times A/(4\pi R^2)$

Remember: Area of circle: πR^2
Area of sphere: $4\pi R^2$



Q: A light bulb producing 100 W of radiation is placed at the center of a sphere of perfectly absorbing material, with radius R . A hole is cut into the sphere, removing an area A of material. Assume the light from the bulb spreads out uniformly in all directions.

What is the rate of energy flow through the hole?



Light spreads out uniformly

Power leaving bulb = power reaching sphere

Hole covers fraction $A/(4\pi R^2)$ of sphere

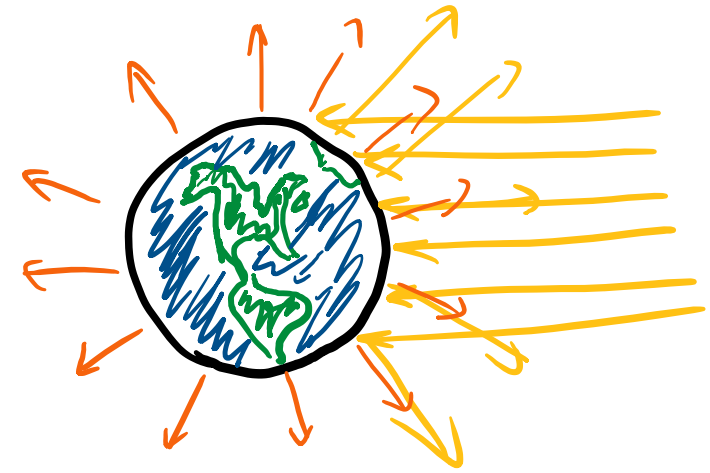
So power of light coming out is $100\text{W} \times A/(4\pi R^2)$

- A. 100W
- B. $100\text{W} \times A$
- C. $100\text{W} \times A/R^2$
- D. $100\text{W} \times A/(\pi R^2)$
- E. $100\text{W} \times A/(4\pi R^2)$ ✓

Remember:	Area of circle: πR^2 Area of sphere: $4\pi R^2$
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Albedo

- **Albedo** is the overall average **reflection coefficient** for solar radiation incident on an object
- Albedo of the Earth is about 0.3 which means that the Earth as a whole reflects 30% of solar radiation



Picture of earth taken by Apollo 17,
December 7, 1972



- What is the power H_{in} of solar radiation absorbed by the Earth? Answer in terms of H_{sun} , earth's albedo, a_e , (fraction of sunlight reflected) and the parameters r_e and R_{s-e} as shown.

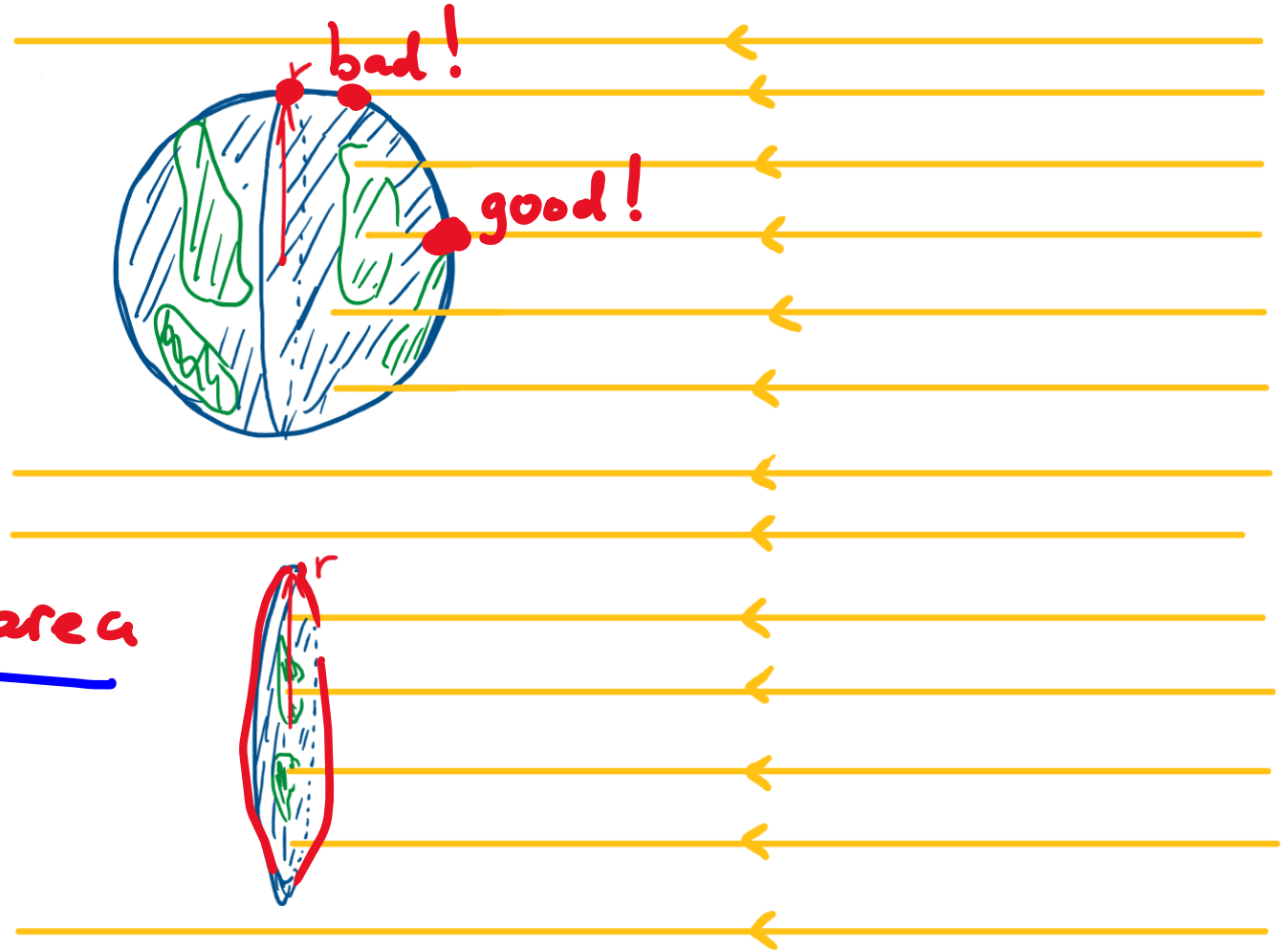
Hint: think about the earlier clicker question.

- radiation absorbed by the Earth? Answer in terms of H_{sun} ,
(light reflected) and the parameters r_e and R_{s-e} as shown.
- er question.
-
- πR_e^2
- $H_{abs} = H_{sun} \cdot \frac{A_{catch}}{A_{spread}} \cdot (1 - \alpha)$
- $4\pi R_{s-e}^2$
- $$e_{sun} \approx 1$$
$$A_{sun} = 4\pi R_s^2$$

$$\begin{aligned} e_{sun} &\approx 1 \\ A_{sun} &= 4\pi R_S^2 \end{aligned}$$

Catching light

- Regular earth:



- We catch sunlight with a cross-section
- We emit with surface area

- Flat earth:

- Each blocks the same amount of sunlight
- What matters for **catching radiation** is the cross-sectional area: πr_e^2
- What matters for **emitting radiation** is the surface area: $4\pi r_e^2$



Q: Let us put these two pieces together:

- The power from the sun is: $H_{sun} \approx A_{sun} \sigma T_{sun}^4$
- Albedo (reflection coefficient) is α_e

What is the power H_{in} of solar radiation absorbed by the Earth? Answer in terms of H_{sun} , earth's albedo, α_e , (fraction of sunlight reflected) and the parameters r_e and R_{s-e} as shown.

Hint: think about the earlier clicker question.

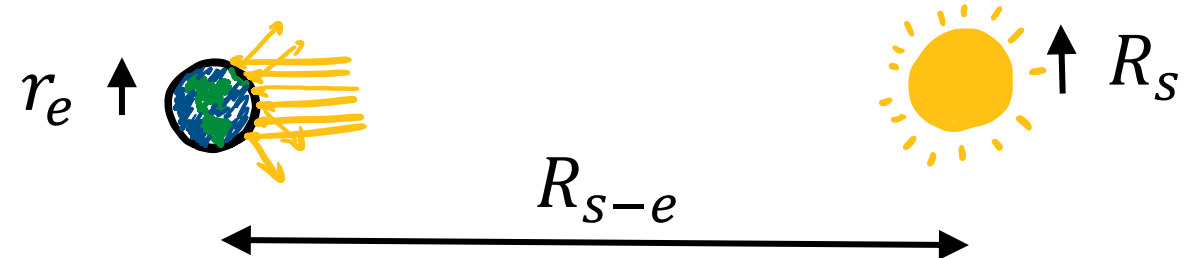
A. $H_{sun} \frac{\pi r_e^2}{4\pi R_{s-e}^2} a_e$

B. $H_{sun} \frac{\pi r_e^2}{4\pi R_{s-e}^2} (1 - a_e)$ ✓

C. $H_{sun} \frac{2\pi r_e^2}{4\pi R_{s-e}^2} a_e$

D. $H_{sun} \frac{2\pi r_e^2}{4\pi R_{s-e}^2} (1 - a_e)$

E. $H_{sun} \frac{4\pi r_e^2}{4\pi R_{s-e}^2} a_e$



Cross-sectional area of earth is πr_e^2

Power hitting earth is: $H_{sun} \frac{\pi r_e^2}{4\pi R_{s-e}^2}$

Fraction $(1 - a_e)$ is absorbed, so:

$$H_{in} = H_{sun} \frac{\pi r_e^2}{4\pi R_{s-e}^2} (1 - a_e)$$

$$\begin{aligned} e_{sun} &\approx 1 \\ A_{sun} &= 4\pi R_s^2 \end{aligned}$$

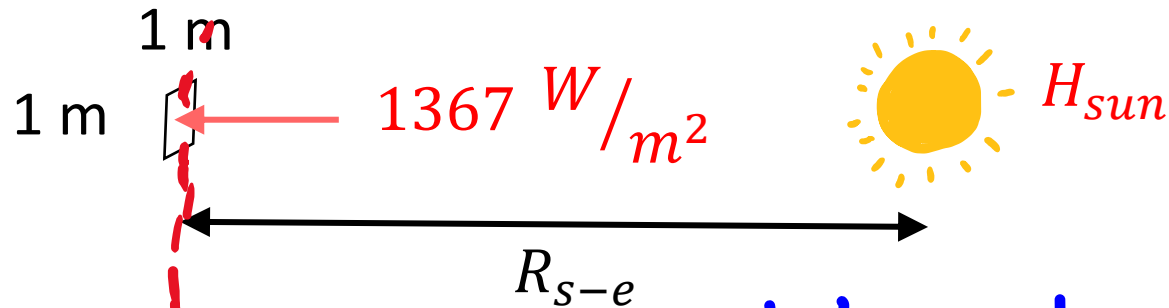
The solar constant

- At earth's orbit, the power per unit area, or **intensity**, of sunlight is:

$$I_{sc} = \frac{H_s}{4\pi R_{s-e}^2} = 1367 \text{ W/m}^2$$

- I_{sc} is called the “**solar constant**”

➤ Doubling the distance would mean 1/4 the intensity



- In terms of I_{sc} , the heat current into the earth due to sunlight is:

$$H_{in} \approx \pi r_e^2 (1 - a_e) I_{sc}$$

Handwritten annotations in blue ink: "catchment area" points to πr_e^2 ; "reflection" points to $(1 - a_e)$; "source intensity" points to I_{sc} , which is circled in blue.



Q: The diameter of Saturn is $\sim 10x$ that of Earth and it is $\sim 10x$ as far away from the Sun. The power of solar radiation hitting Saturn compared to that hitting Earth is:

A. 100x smaller

B. 10x smaller

☒ C. the same

D. 10x bigger

E. 100x bigger


$$H_{in, Sat} = H_{sun} \frac{A_{catch}}{A_{spread}} =$$
$$= H_{sun} \frac{\pi R_{sat}^2}{4\pi R_{sun-sat}^2} = H_{sun} \frac{\pi (\cancel{10} R_e)^2}{4\pi (\cancel{10} R_{sun-e})^2}$$

100

100



Q: The diameter of Saturn is $\sim 10x$ that of Earth and it is $\sim 10x$ as far away from the Sun. The power of solar radiation hitting Saturn compared to that hitting Earth is:

- A. 100x smaller
- B. 10x smaller
- C. the same 
- D. 10x bigger
- E. 100x bigger

Intensity of sunlight (power per unit area): $I = \frac{H_{sun}}{4\pi R^2}$

$$10x R \Rightarrow I_j = \frac{1}{100} I_e$$

Total power (area x Intensity) and $A = \pi r^2$

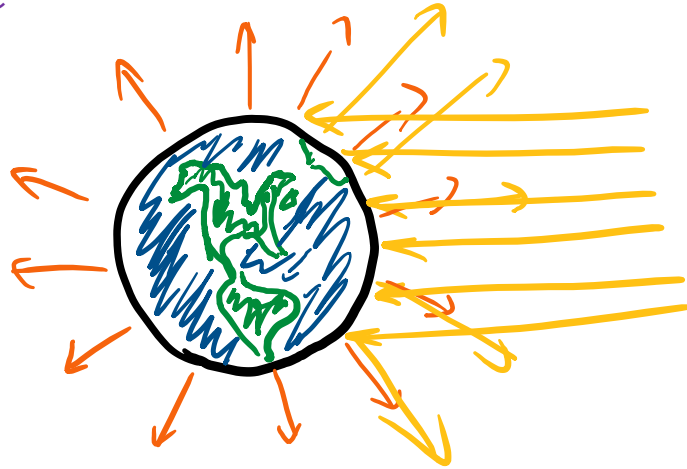
$$10x r \Rightarrow A_j = 100A_e$$

$$\text{So, } H_j = H_e$$

Here R = distance from the source to the planet (determines “spread area”), and r is the radius of the planet (determines “catching area”).



Q: Calculate the equilibrium surface temperature T_e in terms a_e , I_{sc} , r_e , R_{s-e} , σ , and the emissivity e_e



Our model:

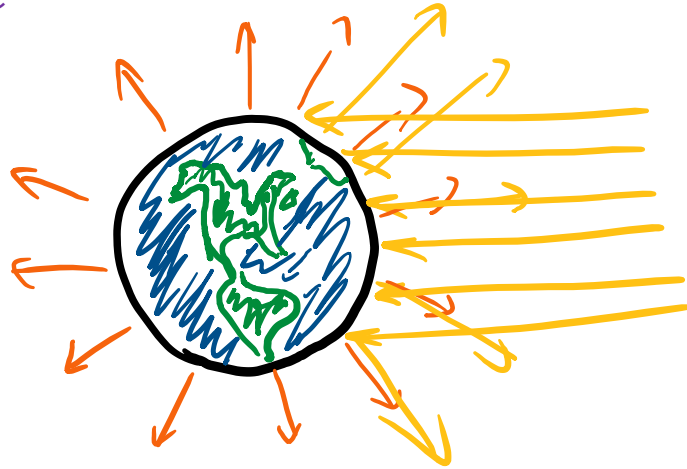
- Steady state: $H_{in} = H_{out}$



- A. $T_e = \left[\frac{a_e I_{sc}}{4e_e \sigma} \right]^{1/4}$
- B. $T_e = \left[\frac{(1-a_e) I_{sc}}{4e_e \sigma} \right]^{1/2}$
- C. $T_e = \left[\frac{(1-a_e) I_{sc}}{4e_e \sigma} \right]^{1/4}$
- D. $T_e = \left[\frac{\pi r_e^2}{4\pi R_{s-e}^2} \frac{(1-a_e) I_{sc}}{e_e \sigma} \right]^{1/4}$



Q: Calculate the equilibrium surface temperature T_e in terms a_e , I_{sc} , r_e , R_{s-e} , σ , and the emissivity e_e



Our model:



• Steady state: $H_{in} = H_{out}$

• H_{in} = Absorbed sunlight $\approx \pi r_e^2 (1 - a_e) I_{sc}$

• H_{out} = IR radiation = $A e \sigma T^4$
 \uparrow surf, earth

$$\pi r_e^2 (1 - a_e) I_{sc} = 4 \pi r_e^2 e_e \sigma T_e^4$$

$$T_e = \left[\frac{(1 - a_e) I_{sc}}{4 e_e \sigma} \right]^{1/4}$$

A. $T_e = \left[\frac{a_e I_{sc}}{4 e_e \sigma} \right]^{1/4}$

B. $T_e = \left[\frac{(1 - a_e) I_{sc}}{4 e_e \sigma} \right]^{1/2}$

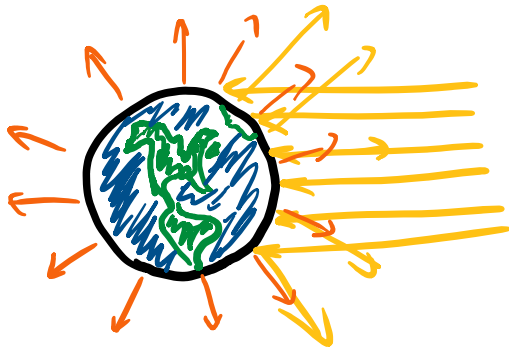
C. $T_e = \left[\frac{(1 - a_e) I_{sc}}{4 e_e \sigma} \right]^{1/4}$ ✓

D. $T_e = \left[\frac{\pi r_e^2}{4 \pi R_{s-e}^2} \frac{(1 - a_e) I_{sc}}{e_e \sigma} \right]^{1/4}$

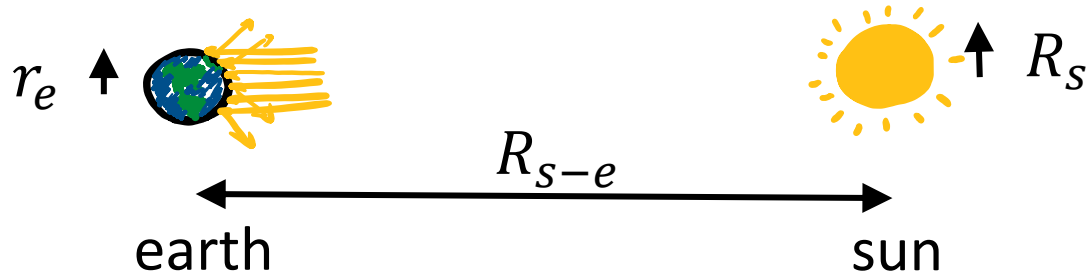
Equilibrium surface temperature of earth: Summary



- Steady state: $H_{in} = H_{out}$



- H_{in} = Absorbed sunlight $\approx \pi r_e^2 (1 - a_e) I_{sc}$
- H_{out} = IR radiation = $A e \sigma T^4$



$$I_{sc} = \frac{H_s}{4\pi R_{s-e}^2} = 1367 \text{ W/m}^2$$

Handwritten note: $A_{\text{sun surf}} e \sigma T_{\text{sun}}^4$ with an arrow pointing to H_s

- Plug in and get: $\pi r_e^2 (1 - a_e) I_{sc} = 4\pi r_e^2 e_e \sigma T_e^4 \Rightarrow$

$$T_e = \left[\frac{(1 - a_e) I_{sc}}{4e_e \sigma} \right]^{1/4}$$

Does this make sense?

$$T_e = \left[\frac{(1 - a_e) I_{sc}}{4e_e \sigma} \right]^{1/4}$$

- Surface of earth has $e_e \approx 1$ for IR radiation and

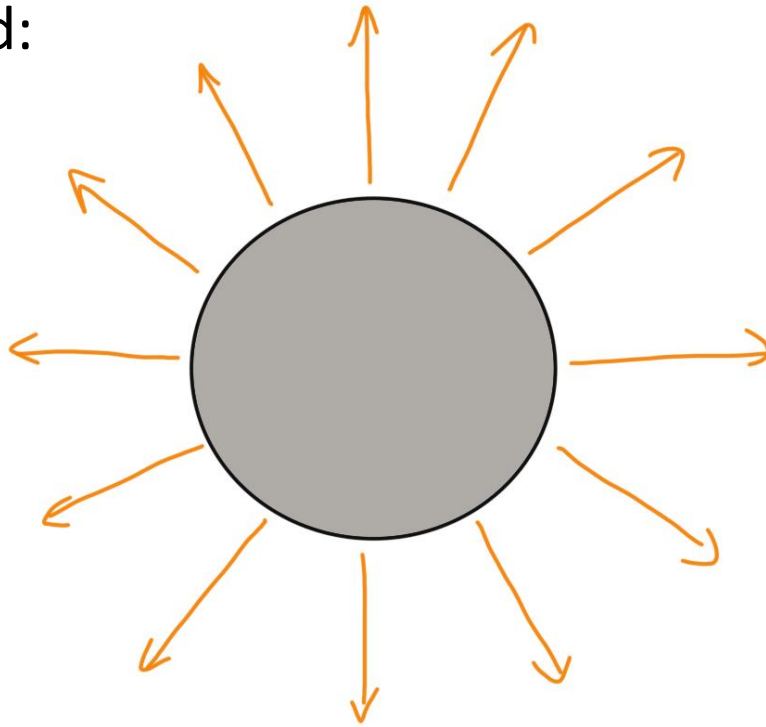
$$I_{sc} = 1367 \text{ W/m}^2 \quad a_e = 0.3 \quad \sigma = 5.67 \times 10^{-8} \frac{\text{W}}{\text{m}^2 \text{K}^4}$$

- These give: $T_e \cong -18^\circ \text{C}$
- Something is off.....
- What have we neglected?
 - Heating from earth's core (like Yoltar)?
 - Nope; that amounts to 47 TW whereas $H_{in} \sim 173,000 \text{ TW} !!$
- There is something else going on...

Actual surface temperature is larger due to the greenhouse effect

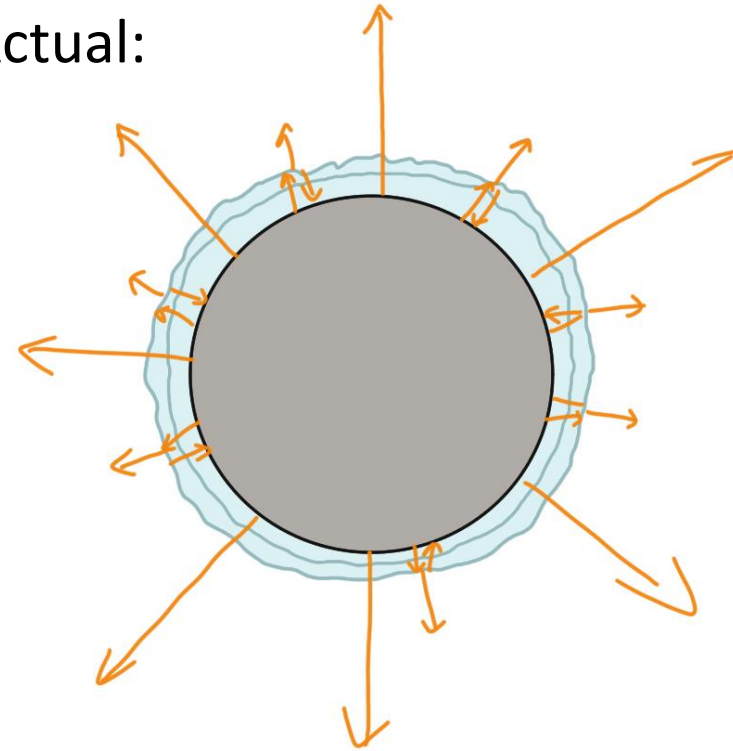
- **Greenhouse effect:** Some IR radiation is absorbed by “greenhouse gases” and re-emitted back to earth

- We assumed:



$$e \approx 1$$

- Actual:



$$e \approx 0.6$$

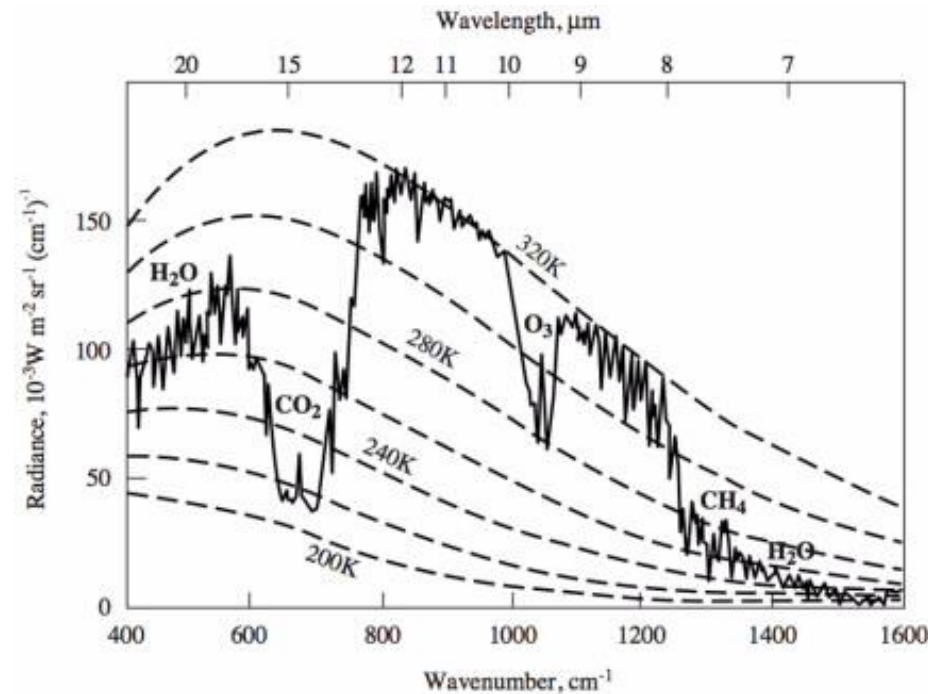
Greenhouse Effect

$$T_e = \left[\frac{(1 - a_e) I_{sc}}{4e_e \sigma} \right]^{1/4}$$

- Lower e \Rightarrow higher T
- Greenhouse effect: $e \approx 0.6 \Rightarrow T \cong 14.5^\circ C$
- Actual average surface temperature of earth $T \cong 13.9^\circ C$
 - Our simple calculation is amazingly close to the actual value!
- But e can decrease due to, e.g., increasing CO_2 concentration in atmosphere
 - Resulting in **Global Warming...**

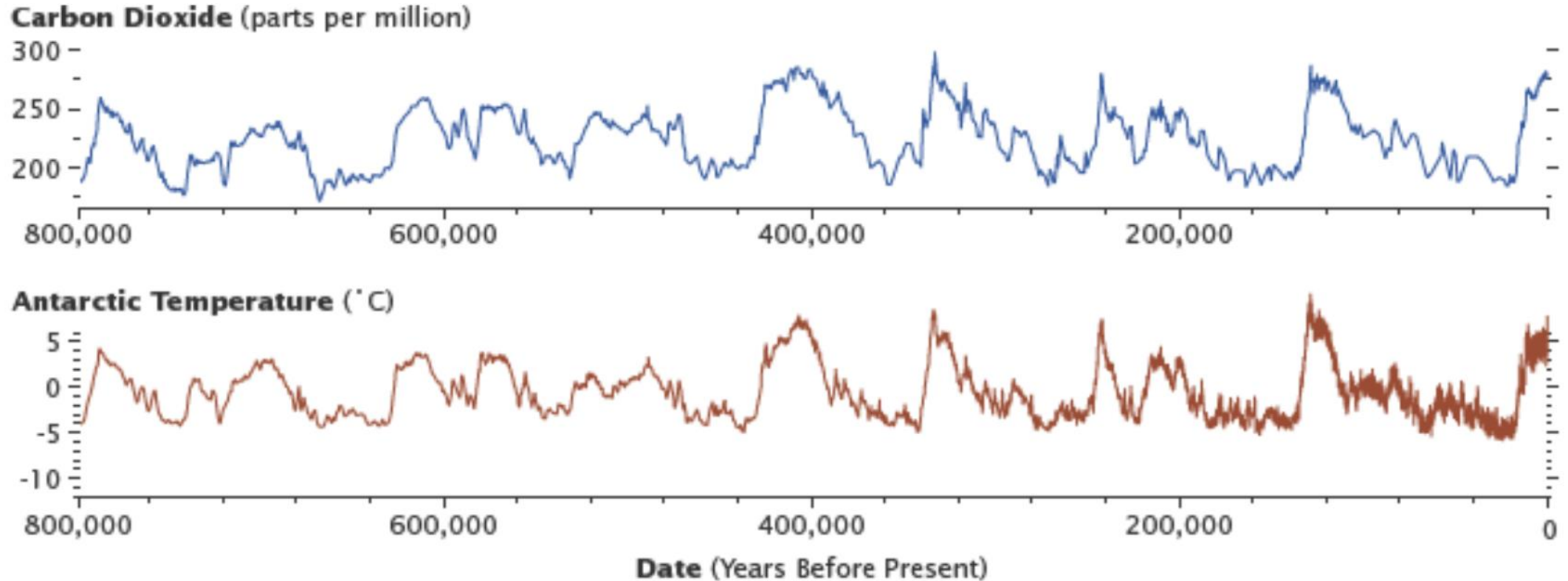
Global Warming

- Greenhouse gases, especially CO_2 , absorb IR radiation emitted by the earth and re-emit it back to earth, effectively reducing the earth's emissivity



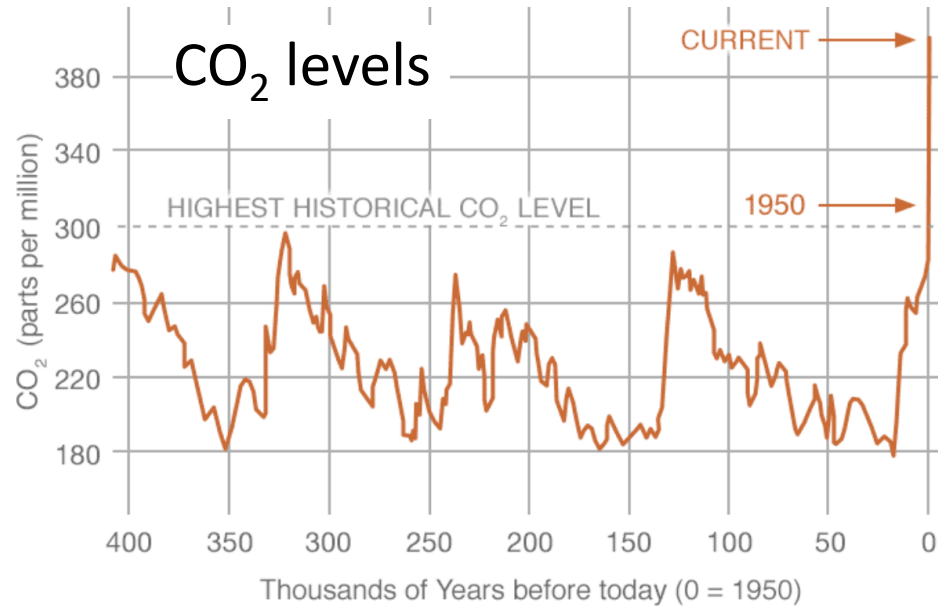
- In addition, as earth warms, white, sunlight reflecting ice in the Arctic disappears, lowering our albedo and further increasing the temperature

CO₂ correlates closely with temperature

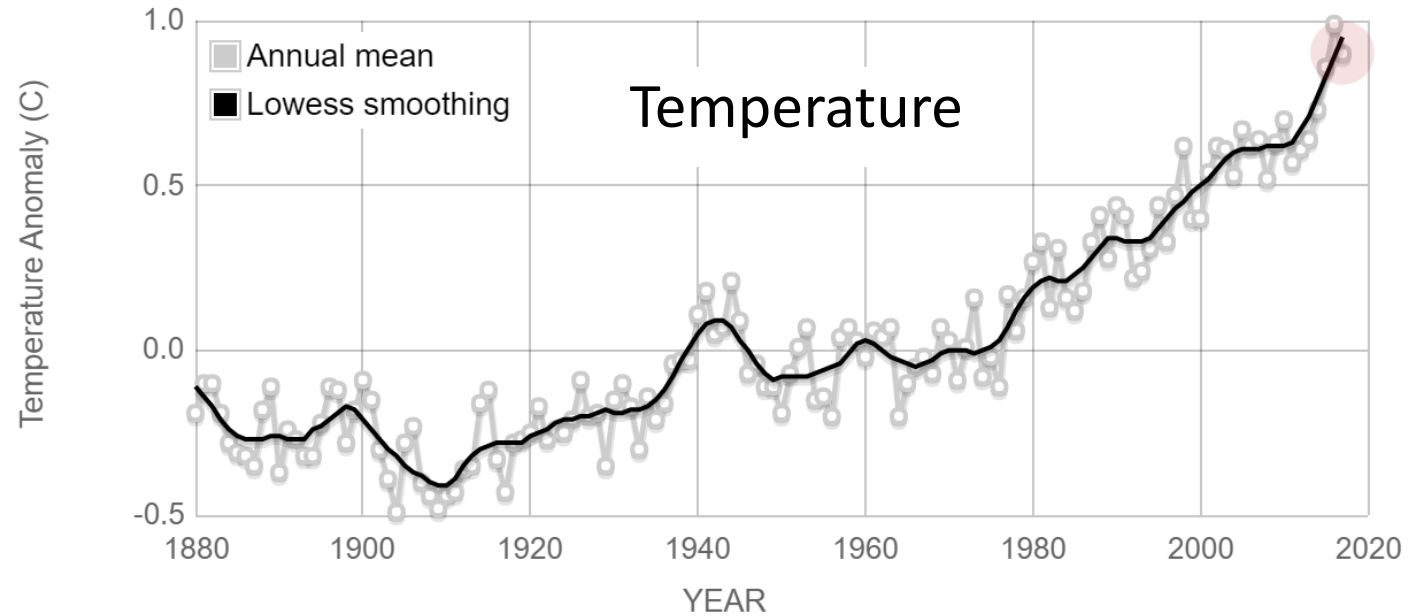


- There is nothing guaranteed about our current temperature of $\sim 14^{\circ}\text{C}$
- With plausible changes to our atmosphere (emissivity) and albedo, we can dial up pretty much any temperature we like, from freezing to boiling

CO₂ correlates closely with temperature



- Almost all climate scientists believe this rise due to human activity



Temperature of Venus...

- Venus has an albedo of 0.750 and let's assume it's emissivity is ~ 1 . The radius of Venus is 6052 km. The Solar constant at Earth is 1367 W/m^2 and the distance from Venus to the Sun is 0.72 times the Earth to Sun distance.
- What is the temperature of Venus?

A. Click A if you are done

B. Click B if you are stuck

Temperature of Venus...

- Venus has an albedo of 0.750 and let's assume it's emissivity is ~ 1 . The radius of Venus is 6052 km. The Solar constant at Earth is 1367 W/m^2 and the distance from Venus to the Sun is 0.72 times the Earth to Sun distance.

- What is the temperature of Venus?

We still have: $H_{in} = H_{out}$

$$\pi r_v^2 (1 - a_v) I_v = 4\pi r_v^2 e_v \sigma T_v^4$$

$$T_v = \left[\frac{(1 - a_v) I_v}{4e_v \sigma} \right]^{1/4}$$

I_v is different than I_{SC} : I_{SC}

$$I_v = \frac{H_s}{4\pi R_{S-v}^2} = \frac{(R_{S-e}^2)}{R_{S-v}^2} \left(\frac{H_s}{4\pi(R_{S-e}^2)} \right) = \frac{R_{S-e}^2}{R_{S-v}^2} I_{SC} = \left(\frac{1}{0.72} \right)^2 1367 = 2637 \frac{W}{m^2}$$

$$T_v = -41^\circ \text{ C}$$

- Actual temperature is $T_v = 464^\circ \text{ C}$
- What could we be missing?
- Atmosphere of Venus is dense and predominately CO_2
- \Rightarrow effective emissivity ~ 0.01 !!
- Huge greenhouse effect!!