

Spatio-temporal variability of zooplankton standing stock in
eastern Arabian Sea inferred from ADCP backscatter
measurements

Ranjan Kumar Sahu, P. Amol, D.V. Desai, S.G. Aparna, D. Shankar

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Abstract

The study focuses on the zooplankton standing stock in the eastern Arabian sea (EAS) and aims to understand its spatio-temporal variation using ADCP(acoustic Doppler current profiler) backscatter measurements. The ADCP moorings were deployed at multiple locations on the continental slope of the west coast of India; of which we have used data from October 2017 to January 2023. The ADCP (operating frequency 153.3 kHz) uses backscatter from or sediments or organisms such as copepods, ctenophores, salps and amphipods greater than 1 cm to calculate current profile. The conversion from backscatter to biomass is based on volumetric zooplankton sampling at the respective locations. Analysis of the data over 25-140 m shows that the backscatter and zooplankton biomass decrease from the upper ocean (215 mg m^{-3} biomass contour) to the lower depths. Seasonal variation is noticed in the monthly climatology zooplankton standing stock (integral of

¹² the biomass over the 20-140m water column) along with change as we move to northward slope
¹³ moorings in EAS. Complementary parameters (mixed layer depth, net primary production, Chl-a,
¹⁴ sea surface temperature) is used to explain the processes leading to growth or decay in zooplankton
¹⁵ biomass and on their migratory behaviour. Additionally, we have studied the effect of wind induced
¹⁶ vertical mixing events. The findings of this research will contribute to a better understanding of
¹⁷ the zooplankton dynamics in the EAS and provide valuable insights into the seasonal and annual
¹⁸ cycles of zooplankton standing stock.

¹⁹ **1 Introduction**

²⁰ **1.1 Background**

²¹ Zooplankton plays a vital role in food web of pelagic ecosystem by enabling the hierarchical trans-
²² port of organic matter from primary producers to higher trophic levels impacting the fish population
²³ and the carbon pump of the deep ocean [Ohman and Hirche, 2001, Le Qu et al., 2016]. They are
²⁴ presumably the largest migrating organisms in terms of biomass [Hays, 2003] which occurs in Diel
²⁵ Vertical Migration (DVM). Zooplanktons depend not only on phytoplankton but other environ-
²⁶ mental parameters (e.g. Mixed layer depth, insolation, Oxygen, thermocline, nutrient availability,
²⁷ chlorophyll concentration and daily primary production). The biological productivity of the ocean
²⁸ is essentially connected with physics and chemistry [Subrahmanyam, 1959, Ryther et al., 1966,
²⁹ Qasim, 1977, Nair, 1970, Banse, 1995, McCreary et al., 2009, Vijith et al., 2016]. The dynamic
³⁰ ocean results in varying physico-chemical properties, leading to bloom and growth of planktons
³¹ in favourable conditions. The changes are strongly influenced by the seasonal cycle in the North
³² Indian Ocean (NIO; north of 5 °N of Indian Ocean). The eastern boundary of Arabian Sea contains
³³ the West India Coastal Current (WICC; [Patil et al., 1964, Ramamirtham and Murty, 1965, Banse,
³⁴ 1968, Shetye et al., 1991, McCreary Jr et al., 1993, Shankar and Shetye, 1997, Shetye and Gouveia,
³⁵ 1998, Maheswaran et al., 2000, Amol et al., 2014, Chaudhuri et al., 2020, 2021]) which reverses
³⁶ seasonally, flowing poleward (equatorward) during November to February (June to September).
³⁷ The direct consequence of this reversal is the seasonal cycle of thermocline, oxycline and thick-
³⁸ ness of mixed Layer Depth (MLD) induced by upwelling favourable conditions in summer and down-
³⁹ welling favourable conditions in winter in eastern Arabian Sea (EAS). Further, the phytoplankton
⁴⁰ biomass and chlorophyll concentration changes with the season [Subrahmanyam and Sarma, 1960,

⁴¹ Banse, 1968, Lévy et al., 2007, Vijith et al., 2016]. Upwelling in summer monsoon leads to maxi-
⁴² mum chlorophyll growth in the entire EAS [Banse, 1968, Banse and English, 2000, McCreary et al.,
⁴³ 2009, Hood et al., 2017]. During winter monsoon, the convective mixing induced winter mixed layer
⁴⁴ [Shetye et al., 1992, Madhupratap et al., 1996b, Lévy et al., 2007, Vijith et al., 2016, Shankar et al.,
⁴⁵ 2016, Keerthi et al., 2017] results in winter chlorophyll peak in northern EAS (NEAS) while the
⁴⁶ downwelling Rossby waves modulate chlorophyll along the southern EAS (SEAS) albeit limited to
⁴⁷ coast and islands [Amol et al., 2020]. (For a detailed description on EAS division, please refer figure
⁴⁸ 1 of [Shankar et al., 2019].

⁴⁹ The zooplankton grazing peak is instantaneous with no time delay from peak phytoplankton
⁵⁰ production [Li et al., 2000], but its population growth lags [Rehim and Imran, 2012, Almén and
⁵¹ Tamelander, 2020] depending on its gestation period and other limiting aspects. While some studies
⁵² suggest that the peak timing of zooplankton may not change in parallel with phytoplankton blooms
⁵³ [Winder and Schindler, 2004], others indicate that lag exists between primary production and the
⁵⁴ transfer of energy to higher trophic levels [Brock and McClain, 1992, Brock et al., 1991]. A recent
⁵⁵ work [Aparna et al., 2022] had shown that peak zooplankton population may never occur even
⁵⁶ with a bloom in phytoplankton such as in SEAS, leading to the collapse of ecological models and
⁵⁷ succeeding food webs of higher trophic levels.

⁵⁸ The conventional zooplankton measurements, where only few snapshot/s of the event is cap-
⁵⁹ tured gives an incoherent or incomplete understanding in terms of spatio-temporal variation of
⁶⁰ zooplankton [Ramamurthy, 1965, Piontkovski et al., 1995, Madhupratap et al., 1992, 1996a] as
⁶¹ much information is revealed by later studies [Jyothibabu et al., 2010, Vijith et al., 2016, Shankar
⁶² et al., 2019, Aparna et al., 2022] using high resolution data. Calibrated acoustic instruments such as
⁶³ Acoustic Doppler Current Profiler (ADCP) along with relevant data can be utilised to understand

64 small scale variability [Nair et al., 1999, Edvardsen et al., 2003, Smith and Madhupratap, 2005,
65 Smeti et al., 2015, Kang et al., 2024], the complex interplay between the physico-chemical parame-
66 ters and ecosystem [Jiang et al., 2007, Potiris et al., 2018, Shankar et al., 2019, Aparna et al., 2022,
67 Nie et al., 2023], the zooplankton migration [Ursella et al., 2018, 2021] and their seasonal to annual
68 variation [Jiang et al., 2007, Hobbs et al., 2021, Liu et al., 2022, Aparna et al., 2022].

69 **1.2 ADCP backscatter and zooplankton biomass**

70 At present, there are two types of acoustic samplers: non-calibrated single frequency acoustic pro-
71 filer such as ADCP or calibrated multi and mono frequency acoustic profilers such as zooplankton
72 acoustic profiler (ZAP) and Tracor acoustic profiling system(TAPS). The use of acoustics as a proxy
73 for zooplankton biomass estimation can be traced to [Pieper, 1971, Sameoto and Paulowich, 1977]
74 and earlier studies which used echograms to approximate the large-scale horizontal extents [Barr-
75 aclough et al., 1969], and small scale vertical extent [McNaught, 1968]. The relationship between
76 backscatter and the abundance and size of zooplankton was described by [Greenlaw] wherein it was
77 pointed out that single frequency backscatter can be used to estimate abundance if mean zooplank-
78 ter size is known. This paved the way for use of single frequency acoustic profiler. A drastic increase
79 in study temporal and spatial variation of zooplankton biomass using backscatter-proxy came in
80 1990s by introduction of high frequency echosounders, with studies [Flagg and Smith, 1989, Wiebe
81 et al., 1990, Batchelder et al., 1995, Greene et al., 1998, Rippeth and Simpson, 1998] methodically
82 showing acoustic backscatter estimated zooplankton biomass in various shelf and slope locations
83 around North Atlantic, North pacific location. The foundation for further research that investigated
84 the potential of acoustic backscatter from ADCPs and multi frequency echosounders in assessing
85 zooplankton biomass and comprehending zooplankton dynamics in diverse maritime habitats was

⁸⁶ established by these initial explorative experiments.

⁸⁷ Acoustic backscatter and zooplankton biomass have been better understood as a result of tech-
⁸⁸ nological and methodological developments over time. Net sampling augmented ADCP backscatter
⁸⁹ have been used to study DVM and the spatial and temporal variability of zooplankton biomass by
⁹⁰ [Cisewski et al., 2010, Smeti et al., 2015, Guerra et al., 2019] in various marine regions, such as the
⁹¹ Southwestern Pacific, the Lazarev Sea in Antarctica and the Corsica Channel in the north-western
⁹² Mediterranean Sea. The zooplankton biomass variation in the Arabian sea has been studied during
⁹³ JGOFS programme in 1990s [Herring et al., 1998, Nair et al., 1999, Fielding et al., 2004, Smith and
⁹⁴ Madhupratap, 2005]. However, their studies were limited to the cruise duration as vessel mounted
⁹⁵ ADCPs were predominantly used; hence long-term data was sparsely produced. The first such
⁹⁶ study to fully exploit the immense potential of ADCPs in EAS was carried out by [Aparna et al.,
⁹⁷ 2022] using ADCP moorings deployed on continental slopes off the Indian west coasts [Amol et al.,
⁹⁸ 2014, Chaudhuri et al., 2020].

⁹⁹ **1.3 Objective and scope of the manuscript**

¹⁰⁰ A network of ADCPs has been installed off the continental slope and shelf on the west coast of
¹⁰¹ India. This ADCPs have enabled a rigorous view of intraseasonal to seasonal scale variability [Amol
¹⁰² et al., 2014, Chaudhuri et al., 2020]. Initially a network of 4 ADCPs (off Mumbai, Goa, Kollam and
¹⁰³ Kanyakumari) on continental slope, it has been extended to include 3 more moorings (off Okha from
¹⁰⁴ 2018, Jaigarh and Udupi from 2017). In the recent study [Aparna et al., 2022] have used ADCP
¹⁰⁵ moorings off Mumbai, Goa and Kollam to explain the temporal variability of zooplankton biomass.
¹⁰⁶ The study showed that the zooplankton peaks (and troughs) is not only non-uniform in latitude but
¹⁰⁷ also heavily influenced by the oxygen minimum zone, MLD and the seasonal upwelling/downwelling

108 conditions. Stark contrast in the phytoplankton bloom and subsequence growth of zooplankton or
109 the lack thereof was observed in the EAS regimes.

110 We build upon the existing work by extending to include the newly incorporated ADCPs so as to
111 have a better understanding in the latitudinal variation of zooplankton biomass in EAS. The paper
112 is organized as follows, datasets and methods employed are described in detail in Secion 2. Section
113 3 describes the observed seasonal cycle of zooplankton biomass and standing stock. The role of
114 mixed layer depth, net primary production, sea surface temperature, wind forcing and circulation
115 in determining the biomass is discussed in results section 4, with conclusion in section 5.

116 **2 Data and methods**

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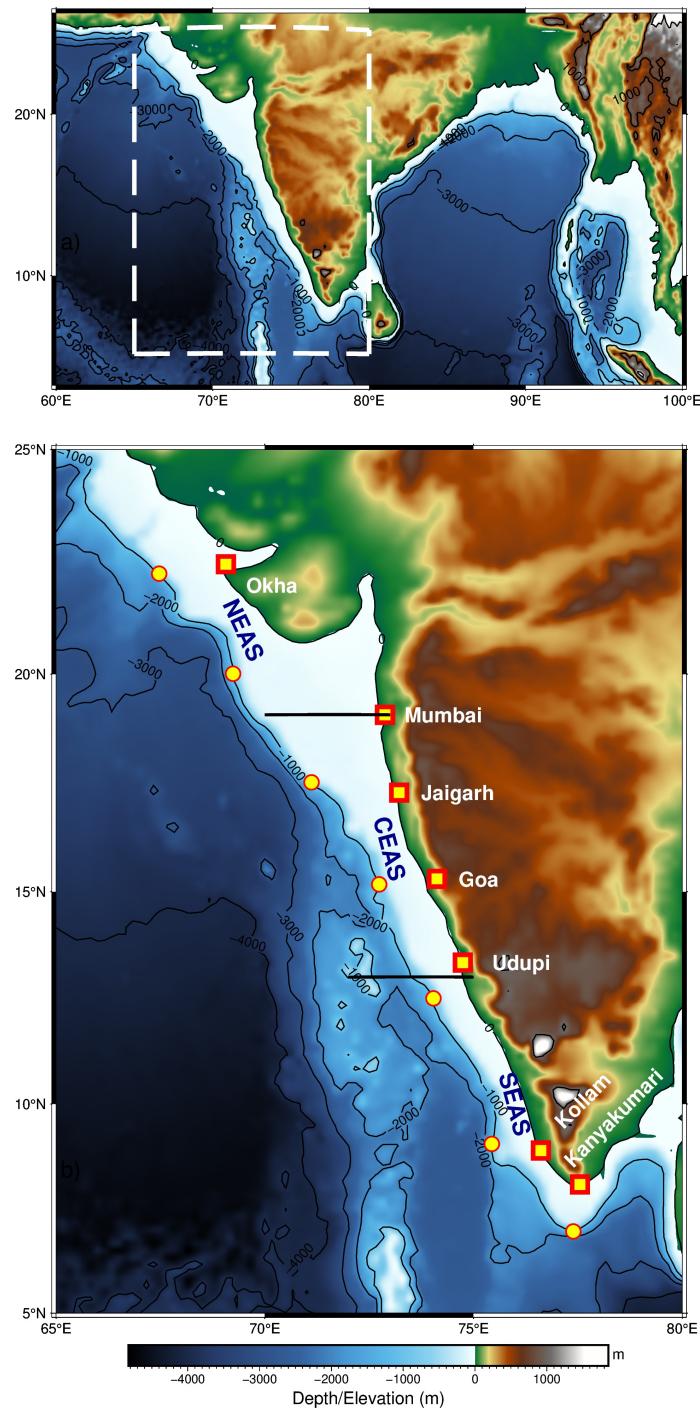


Figure 1: Map showing region of interest. The slope moorings are deployed at 1000 m depth.

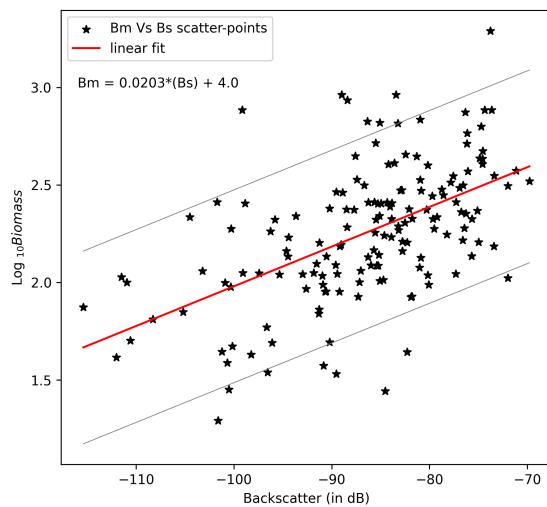


Figure 2: The linear fit line of Biomass (taken in log of biomass) and Backscatter. The linear fit line is within the error range of previous result of [Aparna et al., 2022] onto which latest zooplankton volumetric sample data is added.

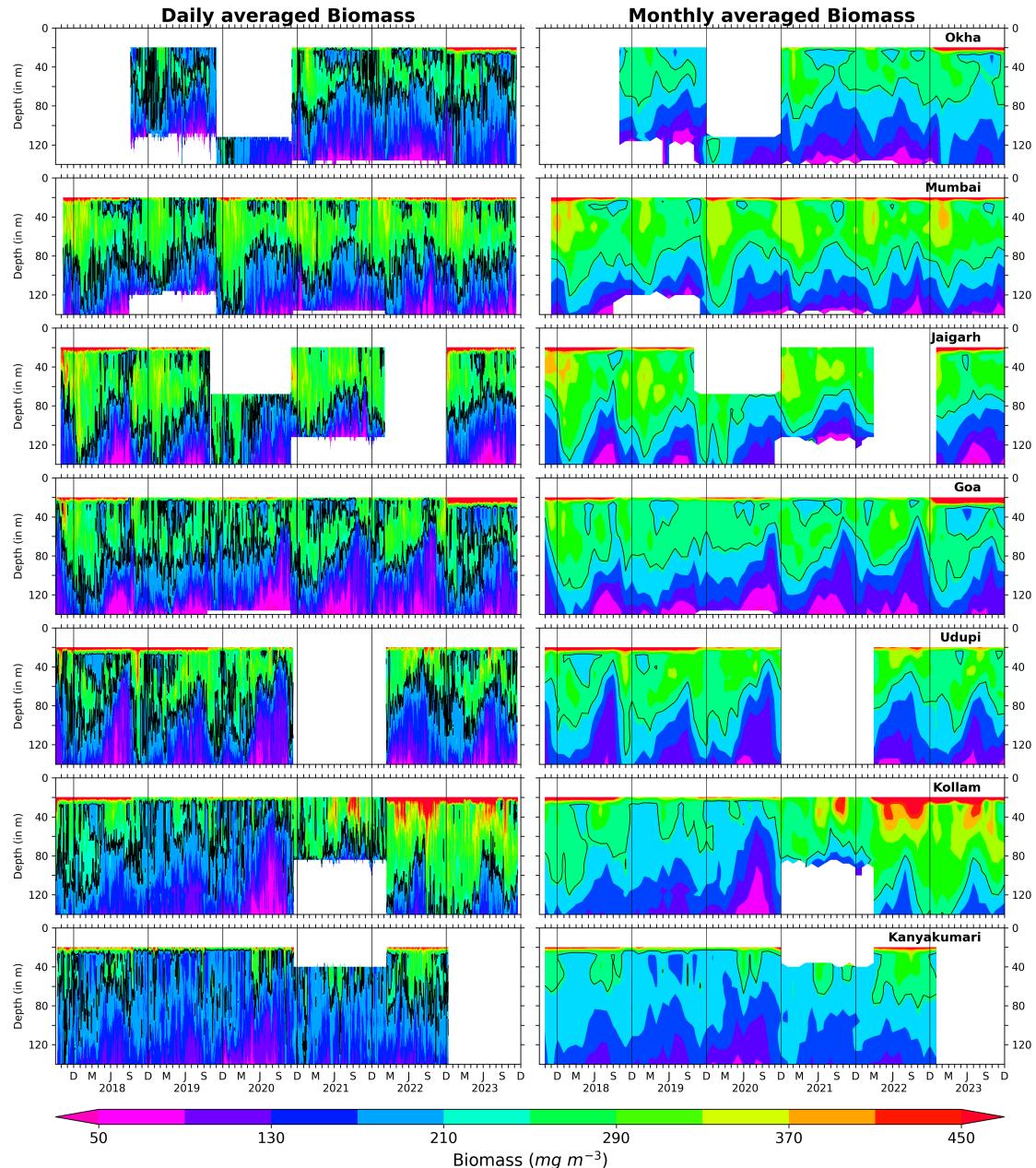


Figure 3: The Daily and monthly averaged biomass for EAS moorings, north (top) to south (bottom). The dark contours are marking 215 mg m^{-3} .

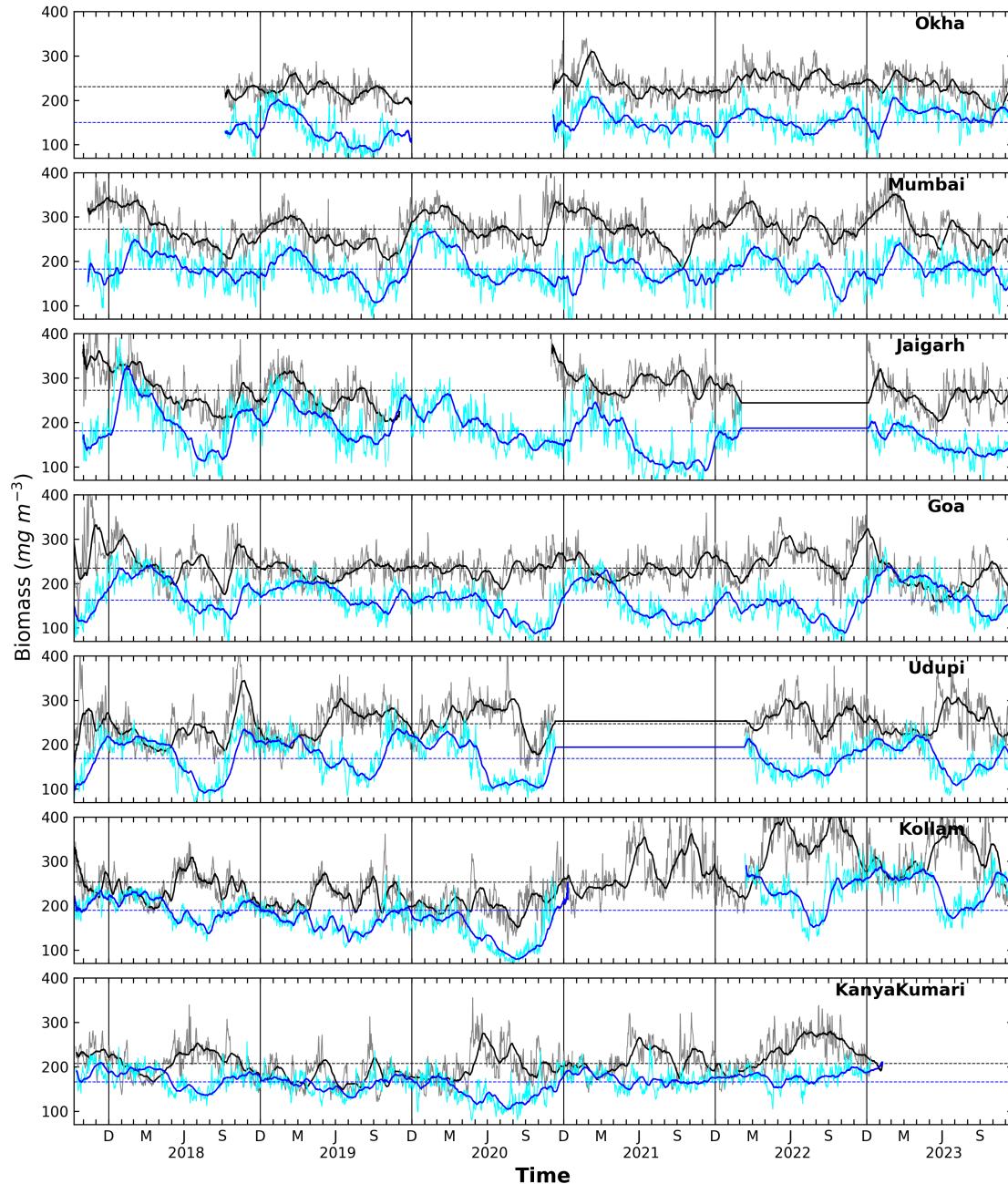


Figure 4: The daily biomass at depth of 40 m and 104 m for all locations shown by grey and cyan curves. The black and blue lines shows the 30 day rolling averaged biomass.

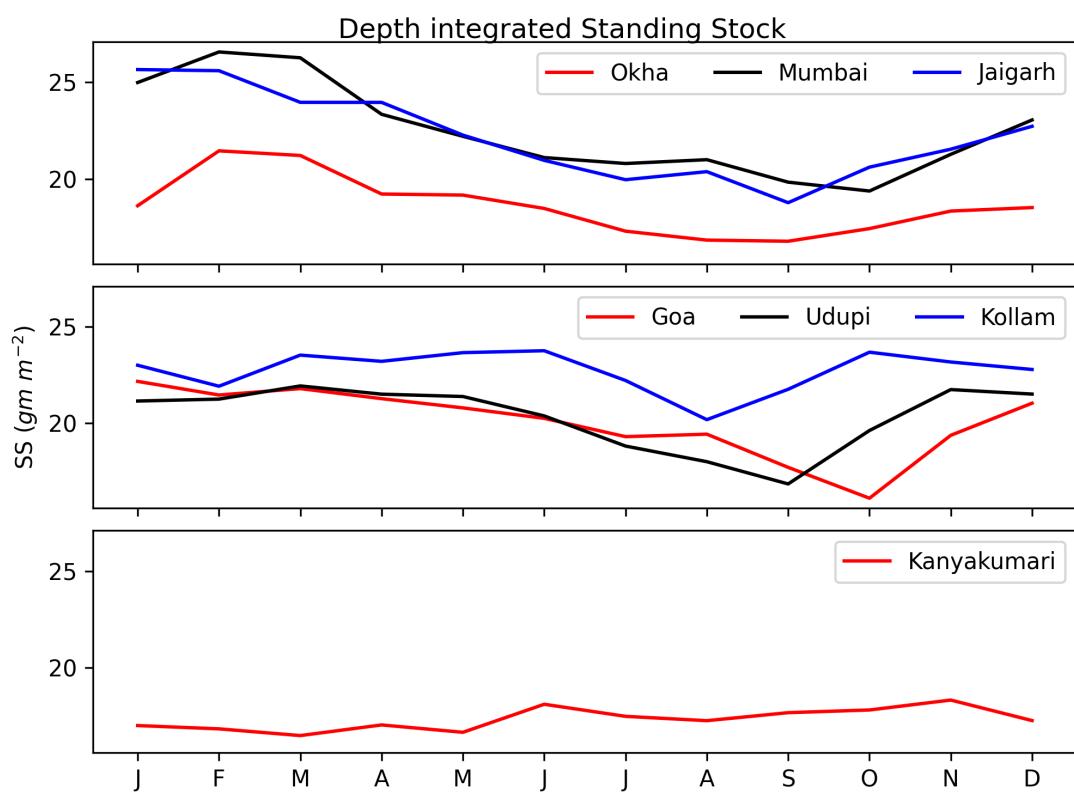


Figure 5