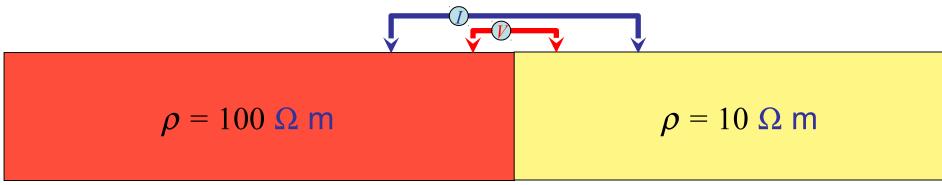
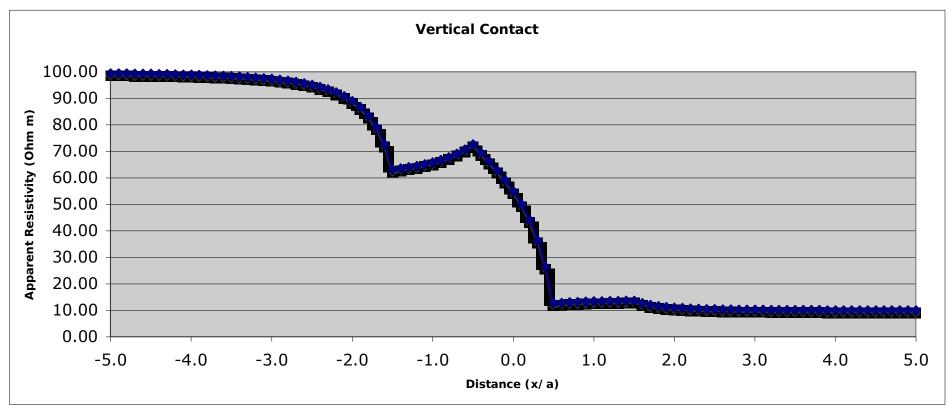
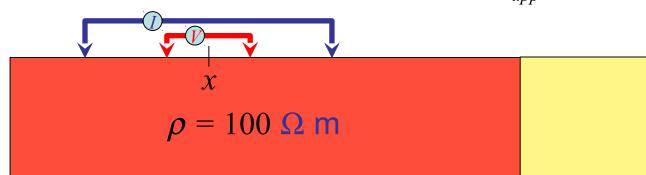
### **Profiling** maintains constant *a*-spacing while moving the array laterally. Wenner array $\rho_{app}$ for a vertical contact:

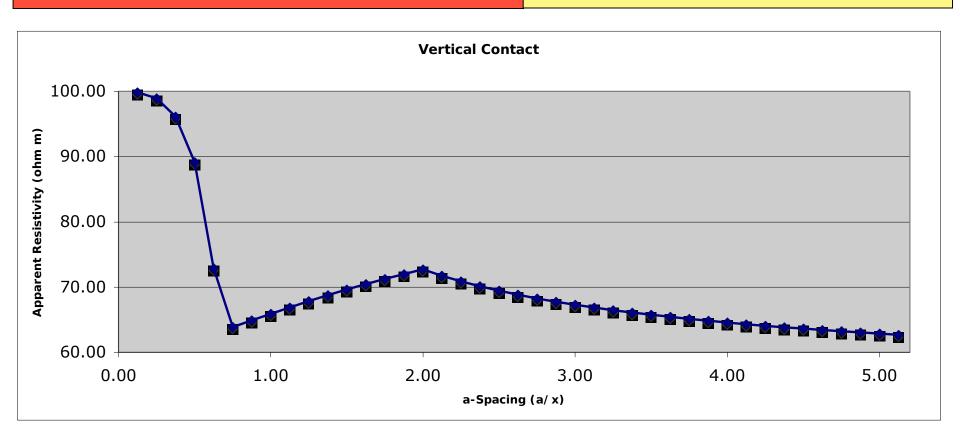




## **Sounding** maintains a constant center while increasing the a-spacing. Wenner array $\rho_{app}$ for a vertical contact:

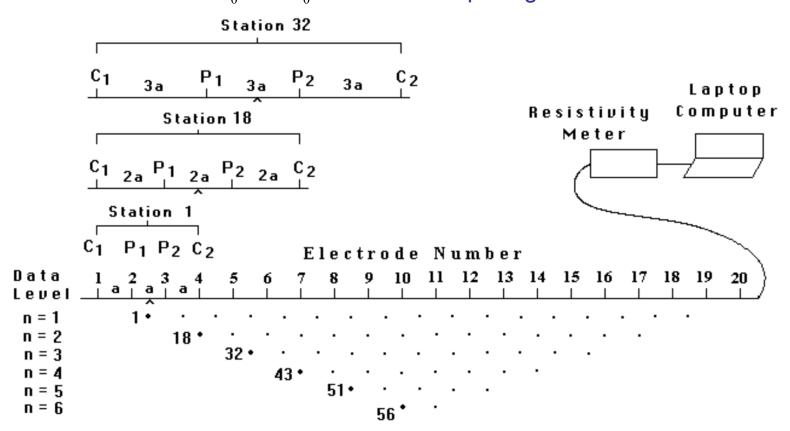


$$\rho = 10 \Omega \text{ m}$$



#### Resistivity "Pseudosection"

- Measurements are a combination of profiling (different x) & sounding (different a-spacings)
- Plot / contour  $\rho_{app}$  versus distance x and a-spacing increment n, where  $n = a/a_0$  and  $a_0$  is smallest a-spacing



One technique used to extend horizontally the area covered by the survey, particularly for a system with a limited number of electrodes, is the roll-along method. After completing the sequence of measurements, the cable is moved past one end of the line by several unit electrode spacings. All the measurements which involve the electrodes on part of the cable which do not overlap the original end of the survey line are repeated (Figure 6).

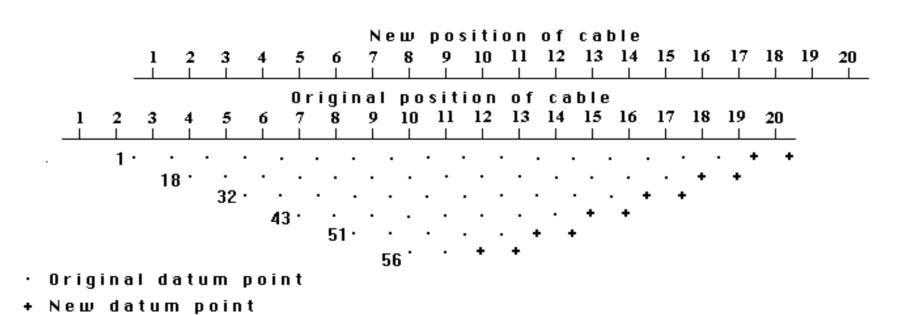
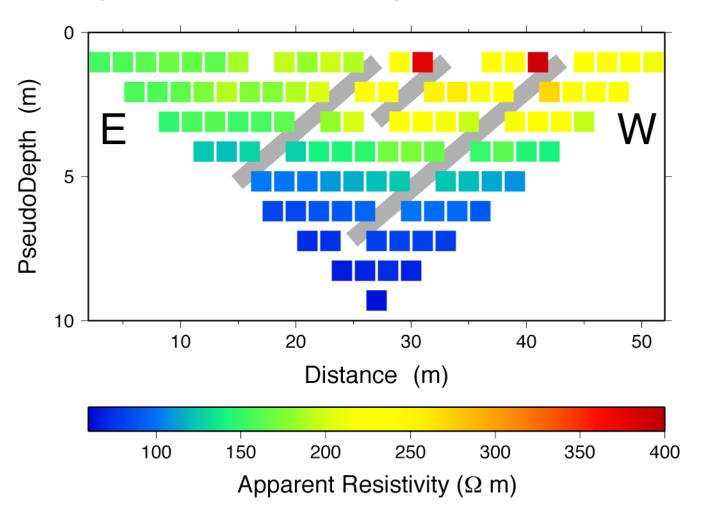


Figure 6. The use of the roll-along method to extend the area covered by a survey.

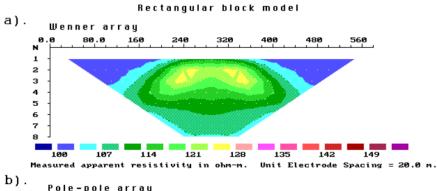
## (This image from the West Cache fault, collected six years ago, is an example of an apparent resistivity pseudo-section):

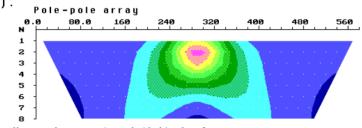


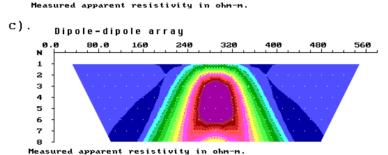
#### Depth-3.0 -2.0 -1.5 -0.5 0.5 1.5 2.0 2.5 3.0 -2.5 Wenner Array 1.5 2.5 -128 -32.0 -8.00 -2.00 2.00 8.00 32.0 Array parameters a=1.0 n=3 Plotting Point Sensitivity value x 100 Wenner-Schlumberger array 2D sensitivity function plot Depth $-4.5 \quad -4.0 \quad -3.5 \quad -3.0 \quad -2.5 \quad -2.0 \quad -1.5 \quad -1.0 \quad -0.5 \quad 0.0 \quad 0.5 \quad 1.0 \quad 1.5 \quad 2.0 \quad 2.5 \quad 3.0 \quad 3.5 \quad 4.0 \quad 4.5 \quad -1.0 \quad -1.5 \quad -1.0 \quad -1.$ Schlumberger Array 1.5 2.5 3.0 –128 –32.0 –8.00 –2.00 2.00 8.00 32.0 128 Sensitivity value x 188 \_ Pseudosection Array parameters a=1.0 n=3 Plotting Point Dipole-dipole array 2D sensitivity function plot Depth **Dipole-Dipole Array** 1.5 2.5 –128 –32.0 –8.00 –2.00 2.00 8.00 32.0 128 Sensitivity value x 188 + Pseudosection Array parameters a=1.8 n=3 Plotting Point

Figure 8. The sensitivity patterns for the (a) Wenner (b) Wenner-Schlumberger and (c) dipole-dipole arrays.

Wenner array 2D sensitivity function plot







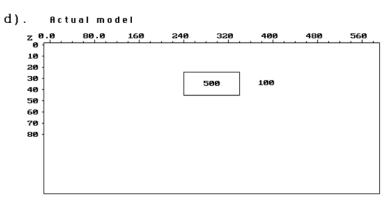
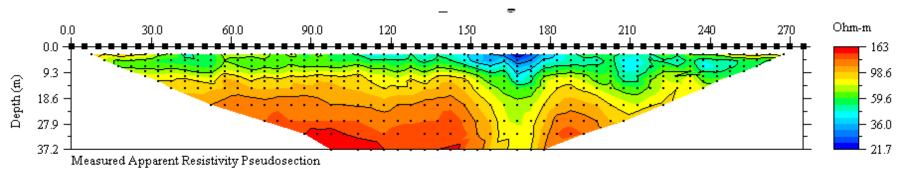
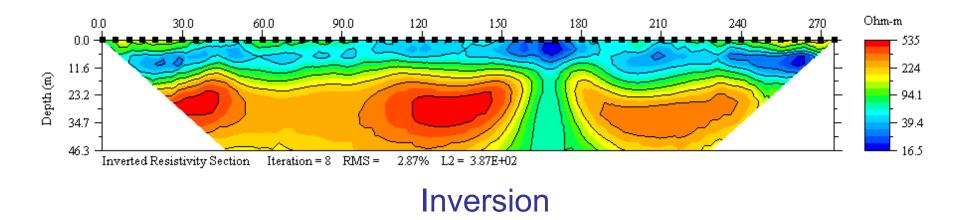


Figure 7. The apparent resistivity pseudosections from 2-D imaging surveys with different arrays over a rectangular block.



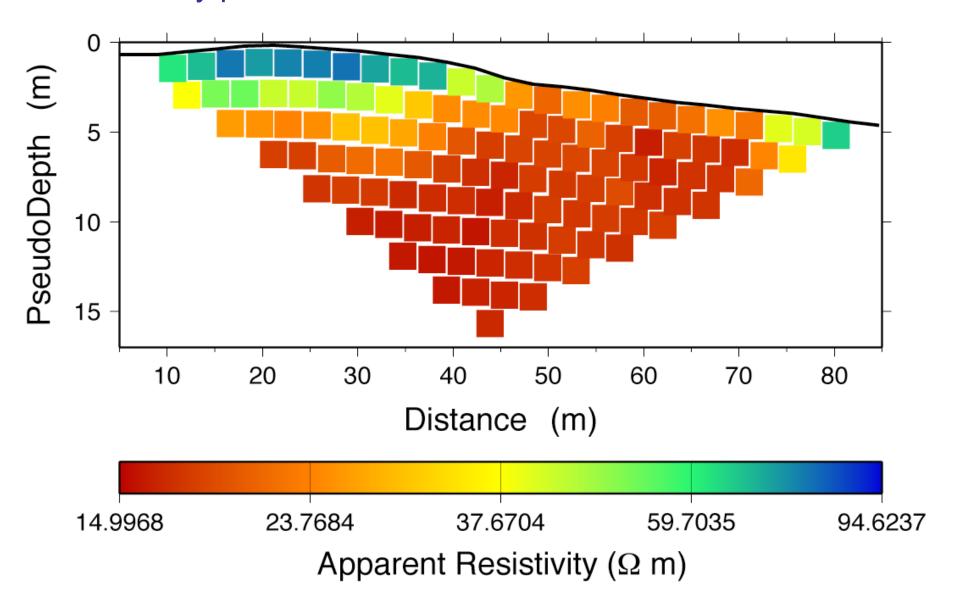


True inversion of the data will yield a much more reliable image (for much the same reasons that depth migration does in the seismic case!)

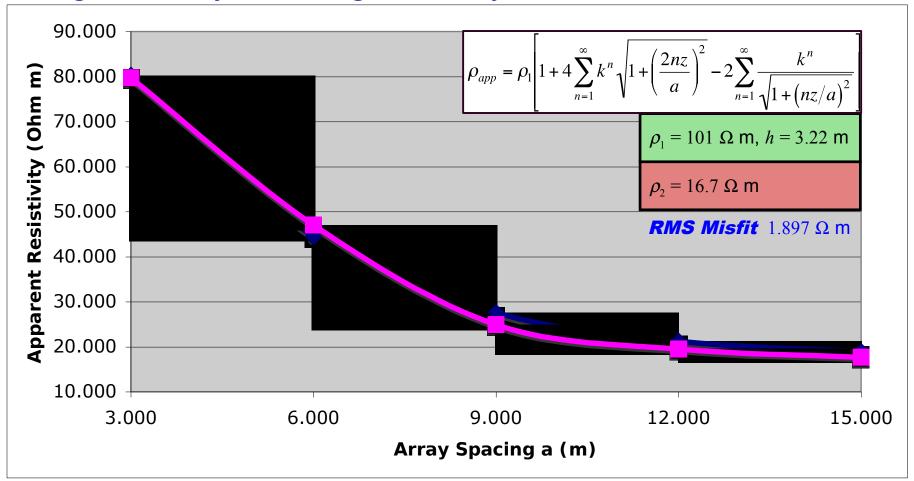


#### **East Cache Valley fault zone near Paradise**

Resistivity pseudosection from the western trench location



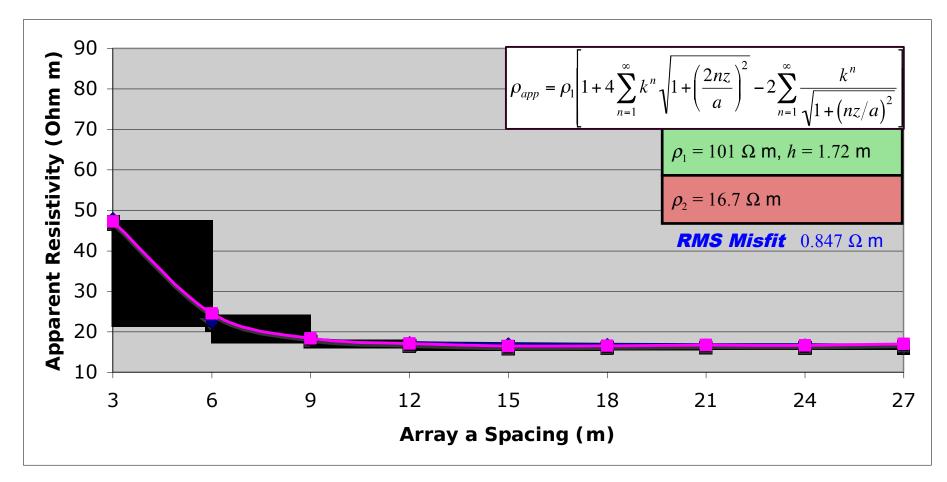
#### Using resistivity sounding & two layer model:



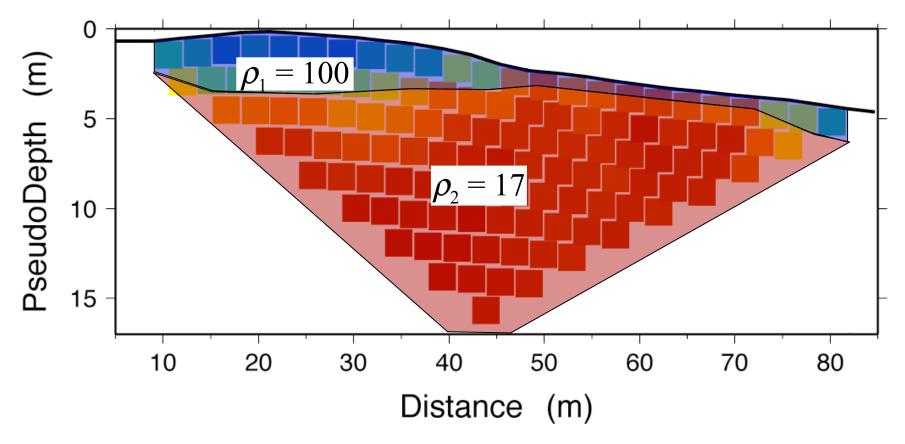
Vertical sounding at 25.7 m. Pink line is measured; Blue line modeled using Excel spreadsheet based on Burger's Table 5-4. Modeling with *RESIST* (with interpolated *a*-spacing) differs in resulting estimates of thickness, but the differences are small.

#### Can model equally well by using

- constant thickness with varying the resistivities, OR
- using constant resistivities with varying the thickness !!

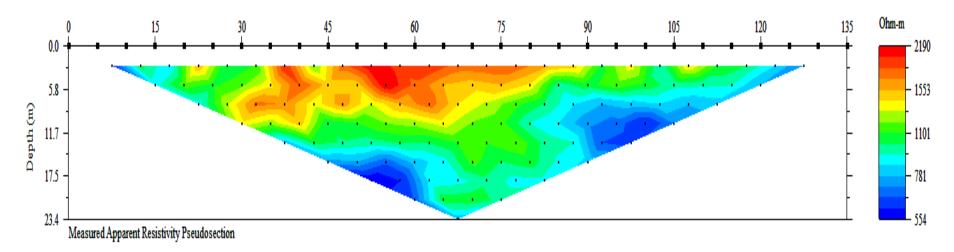


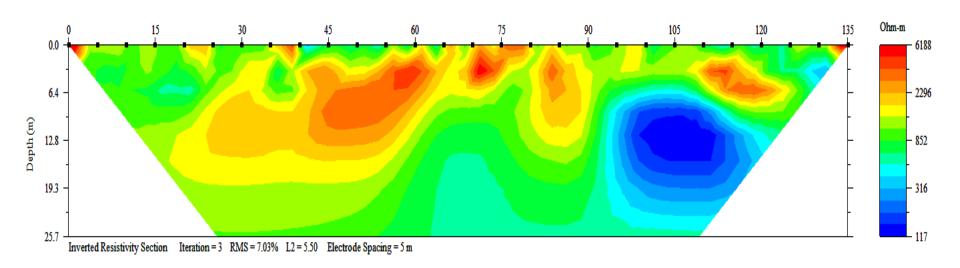
Because these data did not go to a small enough a-spacing to nail down  $\rho$  in the upper layer.



Layer-over-halfspace modeling suggested a variable-thickness upper layer with  $\rho \sim 100~\Omega-m$  (Bonneville deposits) over  $\rho=17$  (Great Salt Lake fm). The data imply the trenched geomorphology represents a beach burm deposit.

#### East Cache fault in Green Canyon

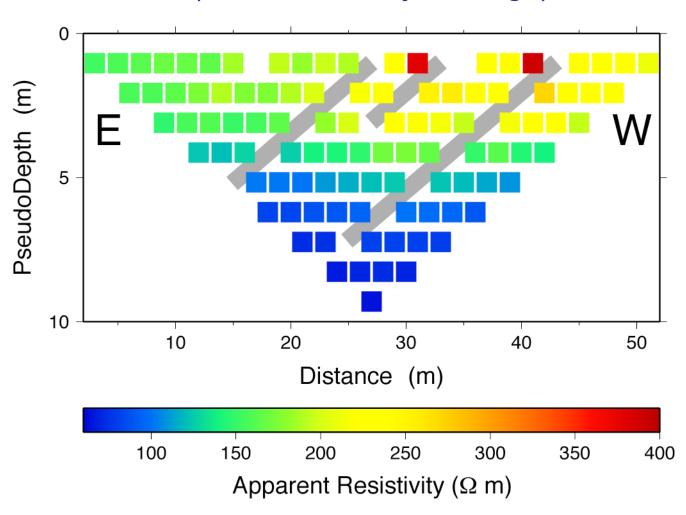


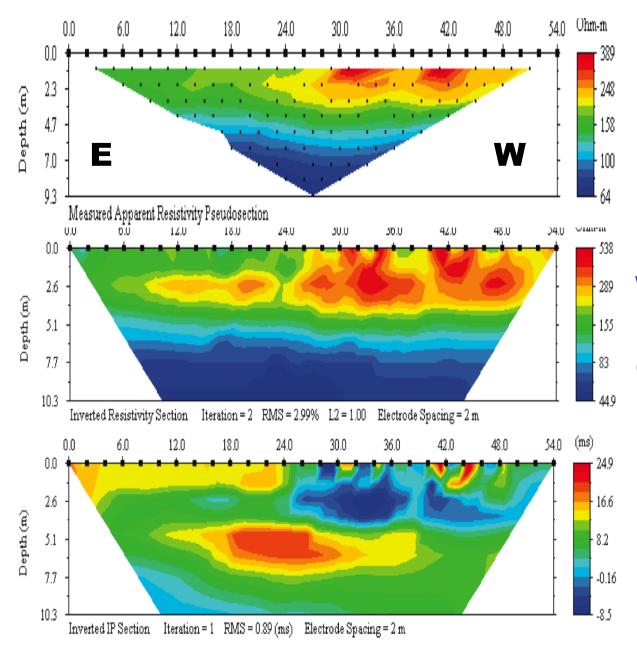


Pseudosection and inversion can be very similar though ...

#### West Cache fault

Apparent resistivity pseudo-section (collected four years ago)

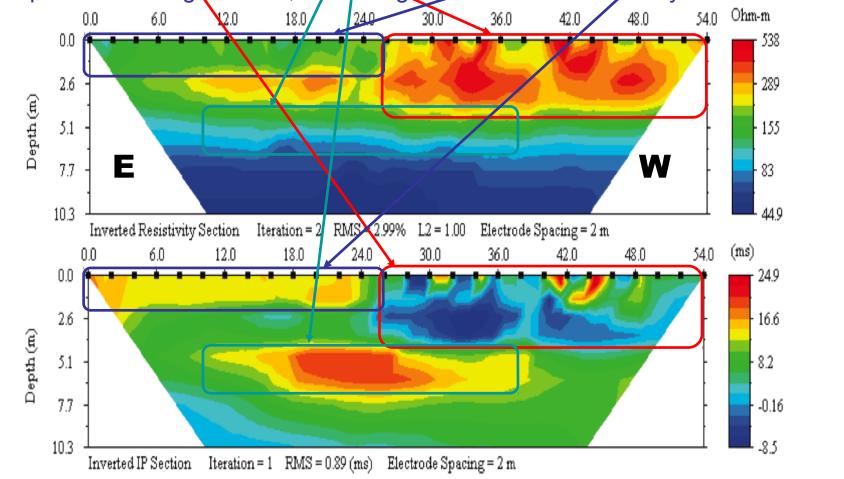




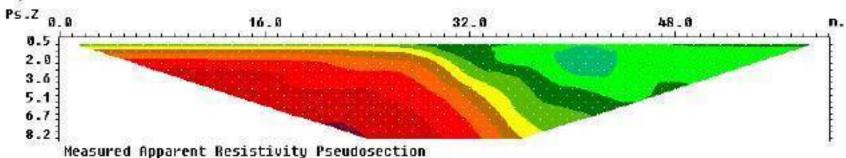
Measurements at WCF included both DC resistivity and IP (induced potential) ... Here, inverted.

Results plus conditions suggested that the DC resistivity/IP inverted models are not imaging lithology variations but wetting!

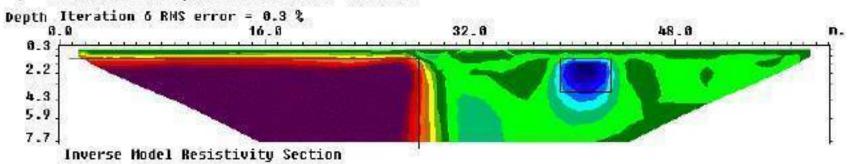
Then the west-side high resistivity could be low soil moisture, resulting from retardation of snow-melt in the forested region, while the high- $\rho$  mid-layer on the east side might sit between the water table and a late spring wetting event, rather than a 20-m long (1) wedge of colluvium that is lithologically different from hanging-wall loess. The way to test this possibility would be to repeat the experiment during summer, after the ground has had time to dry out a bit!



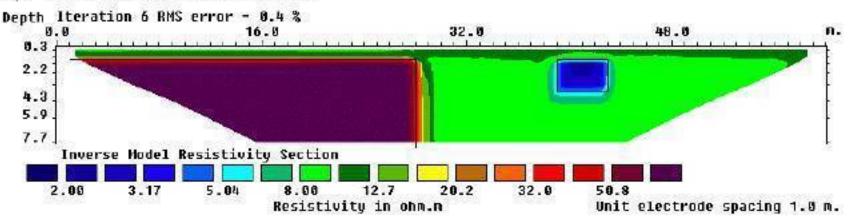




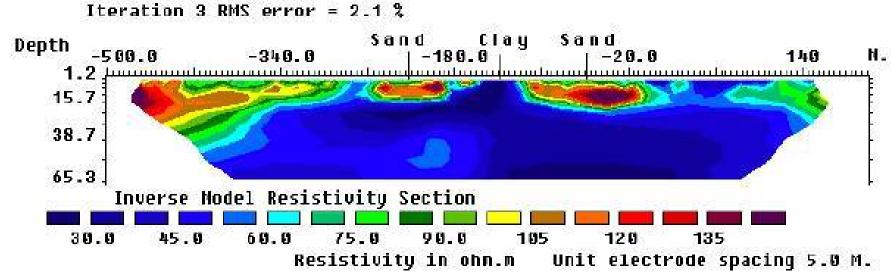
#### b). Standard least-squares smoothness-constrain



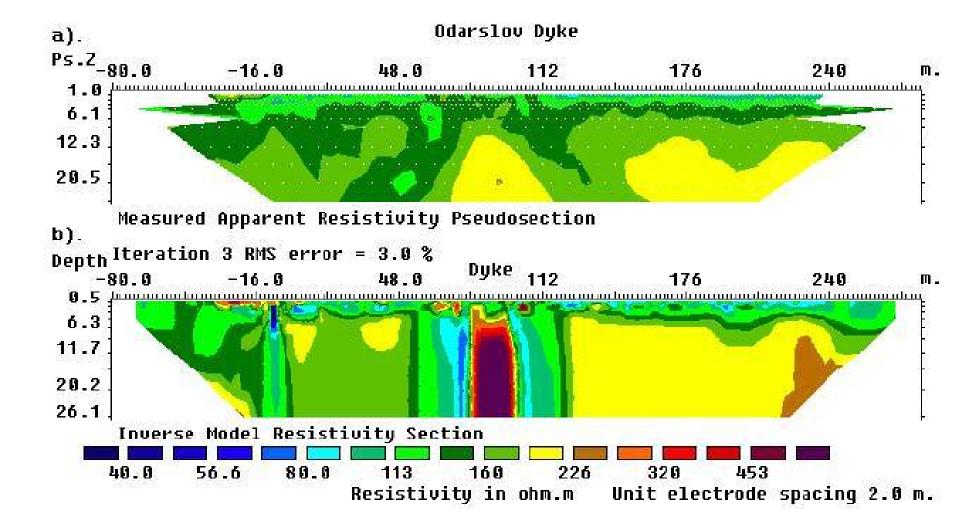
#### c). Robust inversion model constrain



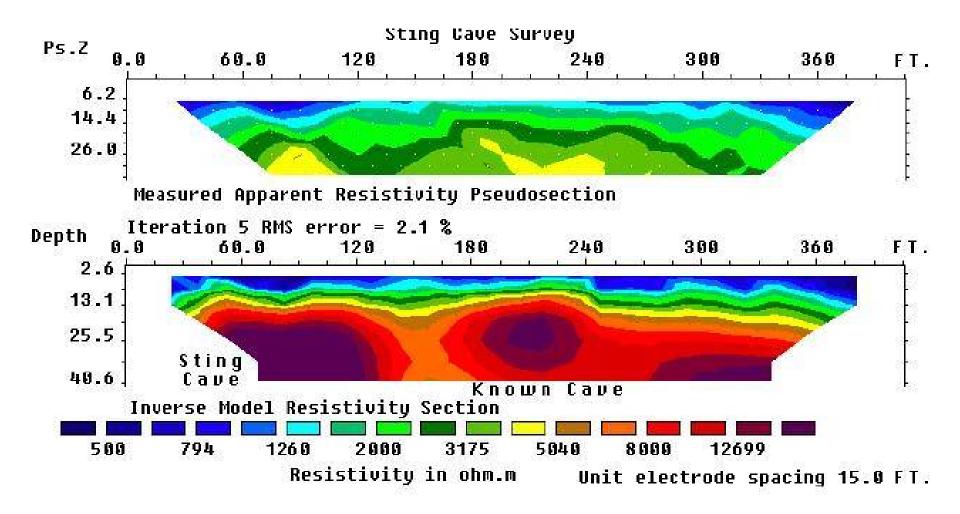
# GRUNDFOR SURVEY LINE 2 Ps.2 -500.0 -340.0 -180.0 -20.0 140 H. 2.6 20.5 41.0 61.4 Measured Apparent Resistivity Pseudosection INVERSION MODEL



Lithology application: Sand/clay in a study of agricultural groundwater contaminant plumes.



Mapping an intrusive dike feature...



Mapping cave locations in karst terrain.