

Expectations for Exercises

Exercises are an important part of the Geophysics lecture. They will treat some aspects of the lecture in more detail, but also cover new ground. We expect that you work on the exercises at home and we will discuss questions and solutions interactively together. The joint meetings will start with randomly chosen students presenting their approach. It is ok if the full solution is not available at this stage and there will be no interrogation. However, please don't show up unprepared because this will inevitably be awkward.

1 Exercises for Gravimetry

1.1 Determining the mean density of the Earth with your own gravimeter

Group Work

Self-organize in groups with 5-6 team members. Choose a group name and a team captain that communicates with the instructors. It is ideal if this group stays together throughout the term for the applied exercises. Make sure that you are inclusive during the group formation.

The vertical gravitational acceleration can be determined by measuring the traveltime of a freely falling object for a known distance. Assuming that the radius of the Earth and the gravitational constant are known ($G \approx 6.674 \cdot 10^{-11} \text{ m}^3\text{kg}^{-1}\text{s}^{-2}$, $R_E \approx 6370$) the mass of the Earth can be determined with the basic principles derived in class.

- Your task is to design a home-built free-fall gravimeter that does the job. You will quickly realize that the time measurement (i.e. the time from the start to the end of the fall) is one critical aspect in the system design. A helper tool that we suggest is the *acoustic stopwatch* that can be accessed via a smartphone and the *phyphox* app. Figure out an efficient way how the traveltime of a freely falling object can be determined with this type of acoustic trigger.
- Collect a dataset of traveltimes and derive the mass of the earth. Provide an error estimate. Can you identify a measurement bias? How does your result compare to literature values?
- Given your result, what is the mean density of the Earth? How does that compare to rock densities found at the Earth's surface? What is a main conclusion that you can draw from this?
- Post your result, your dataset and a picture of your system setup in the ILIAS forum.
- Extra: Are you a tinkerer/Bastler? If so, feel free to improve the system design, e.g., by using light barriers in combination with a raspberry pi nano. We are more than happy to buy material for you, the only constraint is that it works reliably and that it remains cheapish (let's say <100 Euro). **If you succeed, you will win a gift certificate of that can be used in a Tübingen pub of your choice to celebrate your victory with your peers.** Moreover, your system will then be used for eternity for the following classes giving you much honor within the geo- and environmental student community.

Solutions

I made a video with a possible setup here:

[Link to Video](#)

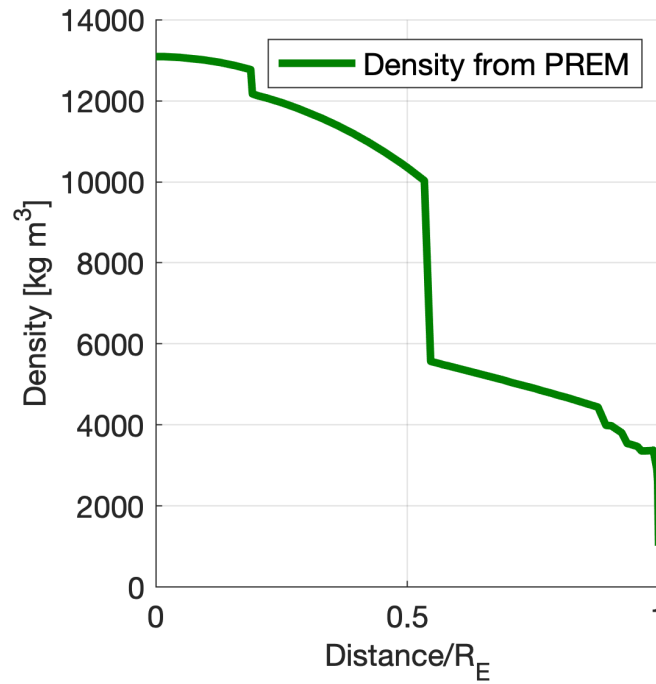
1.2 The shell theorem

[Consider including a sub-step. Quite technical.]

1.3 Gravitational acceleration inside the Earth

Knowing the gravitational acceleration *inside* the Earth is important, e.g., for understanding processes related to mantel convection.

- Use the consequences of the shell theorem to predict the gravitational acceleration $g(r)$ inside a spherical Earth. First assume that the Earth's density is constant, then that it decreases linearly with distance from the center. Assume that the density in the core is approximately $\rho_{core} = 13 \frac{g}{cm^3}$ and in the crust approximately $\rho_{crust} = 2.7 \frac{g}{cm^3}$.
- Draw the results on a x-y graph on paper or using a software of your choice.
- The PREM model provides an observationally constrained estimate of the density distribution inside the Earth. Knowing the shape of this profile, what is your guess of the gravitational acceleration inside the Earth? Why is it more difficult for you to calculate this quantitatively compared to the previous cases of constant and linearly varying density?
- Extra: Calculate the gravitational acceleration inside the Earth based on the PREM model using Matlab/Python/Excel. The data can be found on Ilias.



Solutions

(a) At location $l < R$ inside the Earth, we do not need to worry about the mass distribution for distances $> l > R$ as these will cancel out due to the spherical symmetry. Also the acceleration will only have a radial component. Hence,

$$|\vec{g}(l)| = G \frac{M}{l^2} \quad (1)$$

where M is mass contained in the sphere with radius l . How does M change as we change l ? In this spherical symmetry it is best to adopt spherical coordinates so that we can write:

$$M = \int \rho dV \quad (2)$$

$$= \int_0^{\frac{\pi}{2}} d\theta \sin(\theta) \int_0^{2\pi} d\phi \int_0^l dl \rho l^2 dV \quad (3)$$

this is a fairly complicated way of writing something that you know by heart (i.e. the density times the volume of a sphere), but this is a good opportunity to make use of spherical coordinates. Most importantly, it is worthwhile to remember that the volume element $dV = l^2 \sin(\theta) dr d\theta d\phi$ is quite different from the cartesian coordinates that you are used to otherwise. Also take a moment and understand why the integration limits are the way they are. Solving the integral step by step for the case of a constant density:

$$M = \rho \int_0^{\frac{\pi}{2}} d\theta \sin(\theta) \int_0^{2\pi} d\phi \int_0^l l^2 dl \quad (4)$$

$$= \rho 2\pi \int_0^{\frac{\pi}{2}} d\theta \sin(\theta) \int_0^l l^2 dl \quad (5)$$

$$= \rho 4\pi \int_0^l l^2 dl \quad (6)$$

$$= \rho \frac{4}{3} \pi l^3 \quad (7)$$

Plugging (7) into (1) then results for $l < R$:

$$|\vec{g}(l)| = G\rho \frac{4}{3} \pi \frac{l^3}{l^2} = G\rho \frac{4}{3} \pi l \quad (8)$$

so the gravitational acceleration increases linearly with distance from the Earth's center as long as $l < R$ and ρ is constant. A linear decrease of density with increasing distance can be parameterized as:

$$\rho(l) = \frac{\rho_{crust} - \rho_{core}}{R}(l - R) + \rho_{crust} = \frac{\rho_{crust} - \rho_{core}}{R}l + \rho_{core} \quad (9)$$

As the density now changes with l we cannot take it out of the integration. In particular for eq. (6) this means:

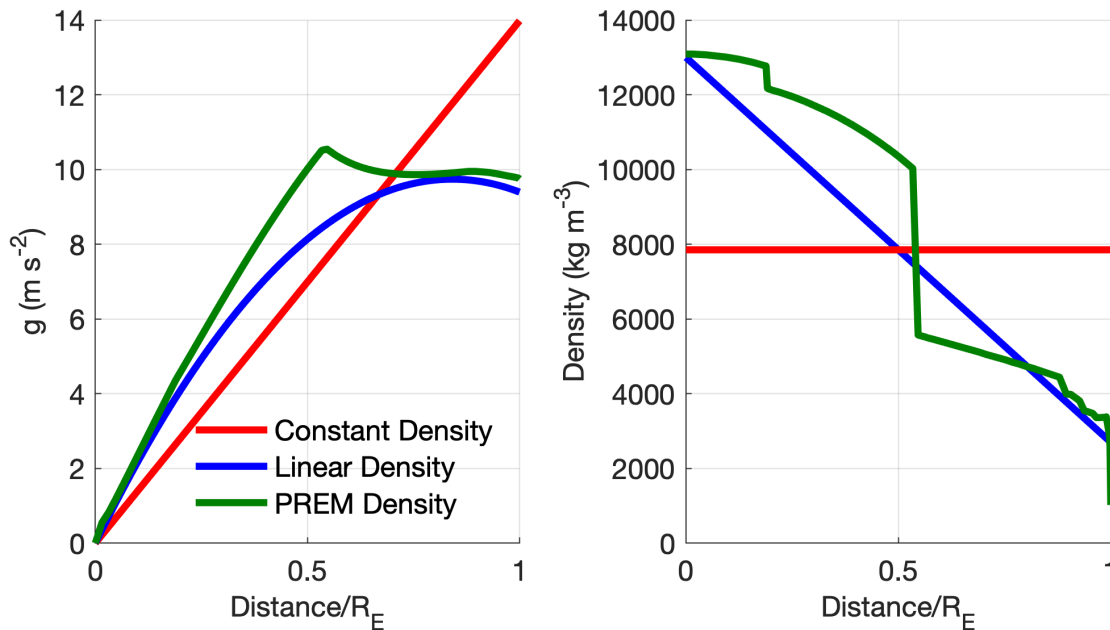
$$M = 4\pi \int_0^l dl \left(\frac{\rho_{crust} - \rho_{core}}{R} l^3 + \rho_{core} l^2 \right) \quad (10)$$

$$= 4\pi \left(\frac{\rho_{crust} - \rho_{core}}{4R} l^4 + \frac{1}{3} \rho_{core} l^3 \right) \quad (11)$$

so that:

$$|\vec{g}(l)| = 4G\pi \frac{1}{l^2} \left(\frac{\rho_{crust} - \rho_{core}}{4R} l^4 + \frac{1}{3} \rho_{core} l^3 \right) \quad (12)$$

$$= 4G\pi \left(\frac{\rho_{crust} - \rho_{core}}{4R} l^2 + \frac{1}{3} \rho_{core} l \right) \quad (13)$$



so that the gravitational acceleration now changes quadratically as shown in the Figure above. Is this realistic? We will have to compare it with a realistic dataset such as the PREM model.

(c) Because we don't have an analytical expression for the PREM density model, we cannot pursue an analytical integration as done in the previous two cases. Hence we do it numerically:

```

1 clear all
2 close all
3
4 %Gravitational constants in SI units
5 G_SI = 6.674e-11;
6 %Density of core and crust in kg m-3
7 rho_crust = 2.7*1003/1000;
8 rho_core = 13.0 * 1003/1000;
9 %Mean density in this model (ouch, too high)
10 rho_mean = (rho_crust+rho_core)/2;
11 %Radius of a spherical Earth in m
12 R = 6371*1000;l=0:1:R;
13
14
15 %% Made up linear density
16 LinearDensity = (rho_crust-rho_core)/R*(1-R)+rho_crust;
17 %% PREM Density
18 PREM = dlmread('PREM.txt',' ',4,0);
19 PREM_Depth = PREM(:,2)*1000;PREM_Density = PREM(:,5)*1003/1000;PREM_dz = (PREM(2,2)
    -PREM(3,2))*1000;
20 [PREM_Depth,ia,ic] = unique(PREM_Depth);PREM_Density=PREM_Density(ia);PREM_Depth(1)
    =1e-3;
21
22 %% Get the corresponding gs
23 gConstantDensity = G_SI*rho_mean*4.0/3.0*pi*l;
24 gLinearDensity = 4*pi*G_SI*((rho_crust-rho_core)/(4*R)*l.^2+1.0/3.0*rho_core*l);
25

```

```

26 %% Numerical integration using the trapezoidal rule
27 for kk=2:length(PREM_Depth)
28     gPREM(kk) = 4*pi*G_SI.*1./((PREM_Depth(kk).^2).*trapz(PREM_Depth(1:kk),
29         PREM_Density(1:kk).*PREM_Depth(1:kk).^2);
30 end
31
32 fig=figure(1)
33 subplot(1,2,1)
34 plot(1/R,gConstantDensity,'r-','LineWidth',3);hold on
35 plot(1/R,gLinearDensity,'b-','LineWidth',3);hold on
36 plot(PREM_Depth/max(PREM_Depth),gPREM,'-','LineWidth',3,'color',[0 0.5 0])
37 ylabel('g (m s^{-2})')
38 xlabel('Distance/R.E')
39 legend('Constant Density','Linear Density','PREM Density','Location','SouthEast')
40 box off;grid on;
41 legend('box','off')
42 subplot(1,2,2)
43 plot(1/R,LinearDensity,'b-','LineWidth',3);hold on
44 plot(1/R,LinearDensity*0+rho_mean,'r-','LineWidth',3);
45 plot(PREM_Depth/max(PREM_Depth),PREM_Density,'-','LineWidth',3,'color',[0 0.5 0])
46 ylabel('Density (kg m^{-3})')
47 xlabel('Distance/R.E')
48 box off;grid on;
49
50 %Export to a png. (This can be done much better.)
51 set(gcf,'PaperUnits','centimeters','PaperPosition',[0 0 20 10])
52 set(findall(fig,'-property','FontSize'),'FontSize',12)
53 print('-dpng','-r300','../..../Exercises/Figures/Gravimetry/
54     Gravimetry01-GravityInsideEarth.png')
55
56 fig=figure(2)
57 plot(PREM_Depth/max(PREM_Depth),PREM_Density,'-','LineWidth',3,'color',[0 0.5 0])
58 box off;
59 grid on;
60 legend('Density from PREM');ylabel('Density [kg m^3]');xlabel('Distance/R.E')
61 set(gcf,'PaperUnits','centimeters','PaperPosition',[0 0 10 10])
62 set(findall(fig,'-property','FontSize'),'FontSize',12)
63 print('-dpng','-r300','../..../Exercises/Figures/Gravimetry/Gravimetry01-PREM.png')

```

../Src/Exercises/Gravimetry/Gravimetry03-GravityInsideSphere.m

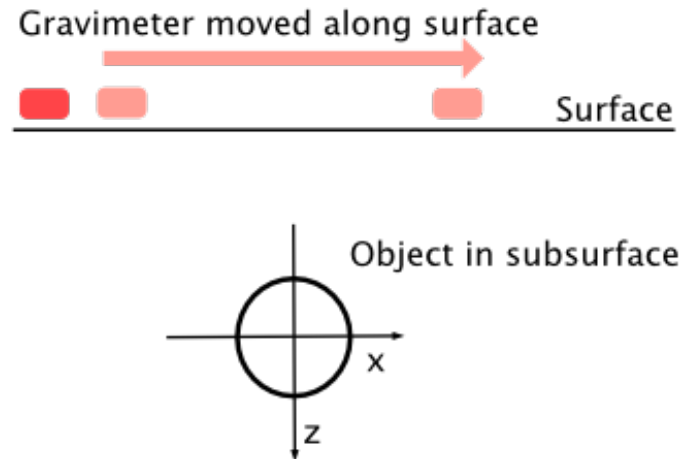


Figure 1: Sketch for problem 1.4. It is easiest to choose a coordinate system with the origin inside the subsurface object.

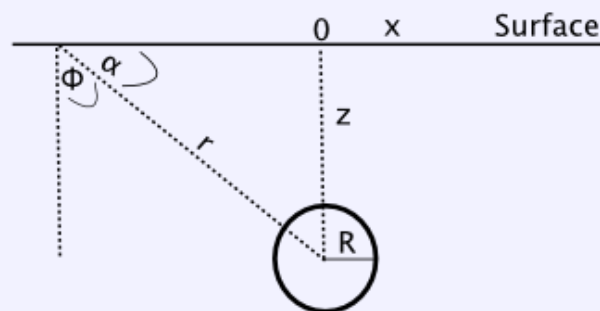
1.4 Detection of a spherical object in the sub-surface

(a) Consider a spherical object with radius R located at depth z below the surface, and a gravimeter that is moved along the surface across the anomaly (Fig.1) Derive an analytical expression for the expected *vertical* anomaly g_z as a function of the lateral distance x and depth z . Calculate the maximum anomalies that you would expect for some realistic settings (e.g., a limestone cave.)

(b) Use a piece of software of your choice (e.g., Excel, SciDAVis, Matlab, Python) and plot the expected vertical gravity anomaly for a specific setting as a function of lateral distance x . Visualize how this profile changes as you vary, e.g., the depth of the object. Label your axis and post a picture in the forum alongside with a comment which software you used.

(c) We now investigate how we can estimate the depth of the object from an anomaly as derived in (a,b). For this you have to derive an expression that links the half-width of the anomaly (i.e. $g_z = \frac{1}{2}g_{z,max}$ for $x = x_{1/2}$) to the depth z of the object. Derive this relationship which has the form $z = cx_{1/2}$. Why is such a relationship be useful?

Solutions



Solution: Because we have a spherical object, we can use the shell theorem and treat it as a point

mass. The gravitational force exhibited by the sphere is:

$$\vec{g} = -G \frac{M}{r^2} \hat{r}$$

where \hat{r} is the unit vector pointing from the center of mass to our gravimeter at the surface. The minus sign makes sure that the force is attractive (and not repulsive). The mass anomaly M formulated in a relative sense of the sphere is given by:

$$M = \frac{4}{3} \pi R^3 \Delta \rho$$

At distance x the distance to the sphere is $r = \sqrt{x^2 + z^2}$. The gravimeter only measures the vertical component of \vec{g} :

$$g_z = -|\vec{g}| \cos \phi = -|\vec{g}| \frac{z}{r}$$

which brings us to:

$$\begin{aligned} g_z &= -\frac{4}{3} \pi R^3 \Delta \rho \frac{G}{r^2} \frac{z}{r} \\ &= -\frac{4}{3} \pi R^3 \Delta \rho \frac{G}{r^3} z \\ &= -\frac{4}{3} \pi R^3 \Delta \rho G \frac{z}{(x^2 + z^2)^{\frac{3}{2}}} \end{aligned}$$

We will encounter the maximum value of the anomaly at $x = 0$ (directly above the target):

$$g_{z,max} = \frac{4}{3} \pi R^3 \Delta \rho G \frac{1}{z^2}$$

which for the limestone cave example is approximately -0.35 mGal (see below).

```

1 clear all
2 close all
3
4 %This is vector x for the lateral distance
5 x = -150:0.01:150;
6 %Depth of the sphere
7 z = 50;
8 %Radius of the sphere
9 R = 25;
10 %Constants for the limestone cave scenario
11 %(SI stands for System International units)
12 DensityLimestone_InGramsPerCentimeterCubed = 2.0;
13 DensityLimestone_SI = DensityLimestone_InGramsPerCentimeterCubed * 1/1000 * 100^3;
14 DenistyAir_SI = 1.2;
15 %Density Contrast in kg/m^3
16 DeltaRho_SI = DenistyAir_SI - DensityLimestone_SI;
17 %Gravitational constants
18 G_SI = 6.674e-11;
19
20
21 %This is the formula we want to plot

```

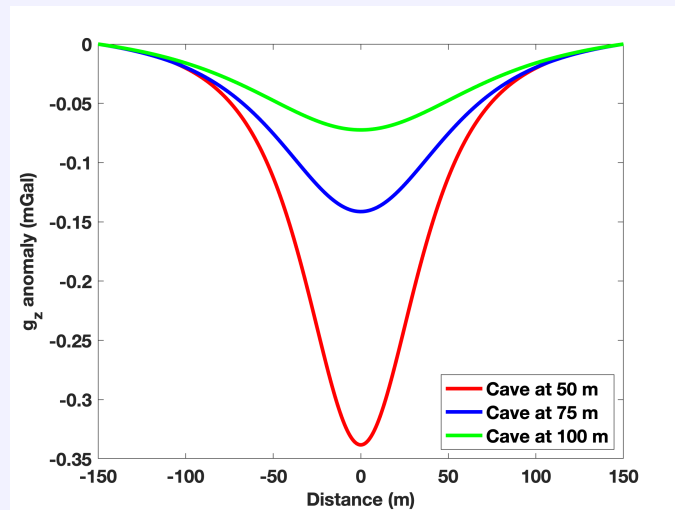


```

22 gz_SI = 4/3*pi*R^3*DeltaRho_SI*G_SI*z./((z.^2+x.^2).^(3/2));
23 gz_SI_max = 4/3*pi*R^3*DeltaRho_SI*G_SI*1/(z^2);
24 gz_mGal = gz_SI*100*1000;
25
26 fig=figure(1)
27 plot(x,gz_mGal-gz_mGal(1),'r-','LineWidth',3);hold on
28
29 %And for a different depths
30 z=75;
31 gz_SI = 4/3*pi*R^3*DeltaRho_SI*G_SI*z./((z.^2+x.^2).^(3/2));
32 gz_mGal = gz_SI*100*1000;
33 plot(x,gz_mGal-gz_mGal(1),'b-','LineWidth',3);hold on
34
35 z=100;
36 gz_SI = 4/3*pi*R^3*DeltaRho_SI*G_SI*z./((z.^2+x.^2).^(3/2));
37 gz_mGal = gz_SI*100*1000;
38 plot(x,gz_mGal-gz_mGal(1),'g-','LineWidth',3);hold on
39
40 legend('Cave at 50 m', 'Cave at 75 m', 'Cave at 100 m', 'Location', 'SouthEast')
41 xlabel('Distance (m)')
42 ylabel('g_z anomaly (mGal)')
43 %Export to a png. (This can be done much better.)
44 set(findall(fig, '-property', 'FontSize'), 'FontSize', 15, 'fontWeight', 'bold')
45 print('-dpng', '-r300', 'GravityAnomaly.png')

```

../Src/Exercises/Gravimetry/Gravimetry01_SphereVisualization.m



This requires some arithmetic manipulation to isolate z .

$$\begin{aligned}
 g_z &= \frac{1}{2}g_{z,max} \\
 \rightarrow \frac{z}{(x_{1/2}^2 + z^2)^{\frac{3}{2}}} &= \frac{1}{2z^2} \\
 \rightarrow \frac{z^3}{(x_{1/2}^2 + z^2)^{\frac{3}{2}}} &= \frac{1}{2} \\
 \rightarrow \frac{1}{(\frac{x_{1/2}^2}{z^2} + 1)^{\frac{3}{2}}} &= \frac{1}{2} \\
 \rightarrow z = \frac{x_{1/2}}{\sqrt{2^{\frac{2}{3}} - 1}} &\approx 1.305x_{1/2}
 \end{aligned}$$

Those expressions can be useful to estimate the depth of an object from an observed anomaly without using a full forward model. However, this only works if the object is close to spherical (which we usually don't know necessarily ahead of time.) Similar estimates exist for other shapes (e.g., horizontal and vertical cylinders).

1.5 Potential of an infinite plate (Bouguer plate)

1.6 Forward modelling and non-uniqueness in potential field methods

Matlab (or Python)

Basic programming (Matlab/Python/R) will likely be part of your study experience when you move to MSc level courses. It is a useful skill to have, but here we do not cover any introduction. What we do is that we start with codes that need little user interaction to give you a feel for what programming can be about. In order to run this exercise you should have a working Matlab version on your Computer, please follow the installation instructions provided by the ZDV. Alternatively, we can also give you a laptop for the joint meeting.

In order to predict how any kind of object will appear in a gravity survey, we need to solve the volume integral:

$$\vec{g}(r) = G \int \frac{1}{r^3} \rho(r) \vec{r} dV$$

which simplifies slightly to:

$$g_z(r) = G \int \frac{1}{r^2} \cos(\phi) \rho(r) dV = G \int \frac{z}{r^3} \rho(r) dV$$

because often only the vertical component is of interest (Ex. cf. 1.4). However, the problem remains complicated as the integration bounds depend on the object's geometry and the integral needs to be solved for every r along the gravimetry profile. Some solutions for special shapes you already know (e.g. sphere, bouger plate). Here we use the solution for a rectangular prism which fortunately others have already calculated for us (*Naggy 1966, Geophysics VOL. XXX, SO. 2*). Using this solution, we can build up more complicated shapes out of individual prisms.

In the specific model applied individual prisms are defined with their widths in the horizontal (w_x, w_y) and the vertical (w_z), together with the positions in the subsurface. The key is that the position coordinates ($dx_1, dx_2, dy_1, dy_2, dz_1, dz_2$) need to be prescribed relative to the measurement position which changes along the profile. The expected anomaly is then calculated based on the analytical solution.

- (a) This exercises uses Matlab. However, only minimal Matlab skills are required to follow along. Download the files *Gravimetry02_ForwardModelling.m* and *gravprism.m* into the same folder on your computer. Check out case 1 which simulates a rectangular object in the subsurface. Change it's location and size so that you know what is going on.
- (b) Switch to case 2. This one treats the combined effect of two prisms. See what's different compared to case 1. Play around with positions to see what is going on.
- (c) Switch to case 3. This one treats the individual effects of two prisms meaning that it doesn't sum them up. This one will not run until you fill out the parts marked with XXX. Use this case to illustrate that multiple situations in the sub-surface (e.g. a shallow prism with low density contrast vs. a lower prism with larger density contrast) can result in similar anomalies. This is an important finding. Forward models are often not unique, and therefore your interpretation won't be either. This situation occurs in many geophysical situations. Remember that.

Solutions

```

1 clear all;
2 close all;
3
4 %% This code quantifies the gravitational potential of a rectangular
5 %% Prism located in the sub-surface using an analytical solution of
6 %% Nagy 1966 (Geophysics)
7
8
9
10 %%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%
11 % Case Numbers
12 % RectangleCentered: 1 ;
13 % Two Rectangles: 2;
14 % Ambiguities: 3
15 %%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%
16 CaseNumber = 1;
17
18 switch CaseNumber
19     case 1
20         display('Calculating centered rectangle.')
21         %This is the density contrast
22         drho = 400;
23
24         % Location and geometry of Prisms.
25         % -----
26         % Width height and depth of the prism.
27         wx = 10;wy = 100;wz = 0.5;
28         % Offset in depth z and lateral direction x.
29         % The y-dimension is (but doesn not have to)
30         % is symetric to the profile direction
31
32         offsetz=1;offsetx=0;
33
34         % Sample Points along profile in x-direction.
35         % Coordinates are symetric with origin in center.
36         % -----
37         dx=0.5;xp = -20:dx:20;yp = xp*0;zp=xp*0;
38
39         % Coordinates of the prism relative to sample points
40         % -----
41         dx1=flipr (min(xp)-wx/2:dx:max(xp)-wx/2);dx2=dx1+wx;
42         dy1=dx1*0-wy/2;dy2=dy1+wy;
43         dz1=dx1*0+offsetz;dz2=dz1+wz;
44
45         % This applies the analysitcal solution of Nagy 1966.
46         % -----
47         dg = gravprism(drho,dx1,dx2,dy1,dy2,dz1,dz2);
48
49         % Here we visualize the results.
50         fig = figure()
51         subplot(3,1,1)
52         plot(xp,dg)
53         ylabel('Gravity Anomaly (mGal)');box off;set(gcf,'color','none');set(gca,

```

```

'color', 'none');
54 subplot(3,1,2)%[x y w h]
55 rectangle('Position',[-wx/2,offsetz,wx,wz]);
56 set(gca,'XAxisLocation','top','YAxisLocation','left','ydir','reverse');
57 xlim([min(xp),max(xp)]);xlabel('Horizontal Distance x (km)');ylabel('Depth (
m)');ylim([0,2]);set(gcf,'color','none');set(gca,'color','none');
58 subplot(3,1,3)%[x y w h]
59 rectangle('Position',[-wx/2,-wy/2,wx,wy]);
60 xlim([min(xp),max(xp)]);xlabel('Horizontal Distance x (km)');ylabel('
Horizontal Distance y (km)');
61 %Export to a png. (This can be done much better.)
62 set(gcf,'color','none');set(gca,'color','none');
63 set(gcf,'PaperUnits','centimeters','PaperPosition',[0 0 10 20])
64 set(findall(fig,'-property','FontSize'),'FontSize',12)
65 print('-dpng','-r300','../..../LatexSlidesLectures/Figures/Gravity/Exported
/ForwardModelPrism.png')
66 case 2
67 display('Calculating the effect of multiple rectangles.')
68 % This is the density contrast
69 % -----
70 drho = [400];
71 % Width, height and depth of the two prisms.
72 % -----
73 wx = [10 8];
74 wy = [100 100];
75 wz = [0.5 0.25];
76
77 % Offset in depth z and lateral direction x.
78 % The y-dimension is (but doesn not have to)
79 % is symetric to the profile direction
80 % -----
81 offsetz=[1 0.25];offsetx=[-10 12]
82
83 % Number of prisms in this example
84 % -----
85 np = length(offsetz);
86
87 % Sample Points along profile in x-direction.
88 % Coordinates are symetric with origin in center.
89 % -----
90 dx=0.5;xp = -20:dx:20;yp = xp*0;zp=xp*0;
91
92 % Coordinates of the prisms relative to sample points
93 % -----
94 for kk=1:np
95 dx1(kk,:)=fliplr(min(xp)-wx(kk)/2:dx:max(xp)-wx(kk)/2)+offsetx(kk);dx2(
kk,:)=dx1(kk,:)+wx(kk);
96 dy1(kk,:)=dx1(kk,:)*0-wy(kk)/2;dy2(kk,:)=dy1(kk,:)+wy(kk);
97 dz1(kk,:)=dx1(kk,:)*0+offsetz(kk);dz2(kk,:)=dz1(kk,:)+wz(kk);
98 end
99 dg = gravprism(drho,dx1,dx2,dy1,dy2,dz1,dz2);
100
101 % Visualization of the combined effect
102 % -----

```

```

103 figure()
104 subplot(3,1,1)
105 % Here we show the combined effect by summing the effects if individual
prisms
106 plot(xp,sum(dg,1))
107 xlabel('Horizontal Distance x (km)');ylabel('Gravity Anomaly (mGal)');
108 subplot(3,1,2)
109 hold on;
110 for kk=1:np
111     rectangle('Position',[-wx(kk)/2+offsetx(kk),offsetz(kk),wx(kk),wz(kk)]);
112 end
113 set(gca,'XAxisLocation','top','YAxisLocation','left','ydir','reverse');
114 xlim([min(xp),max(xp)]);xlabel('Horizontal Distance x (km)');ylabel('Depth (
m)');
115 subplot(3,1,3)%[x y w h]
116 for kk=1:np
117     rectangle('Position',[-wx(kk)/2+offsetx(kk),-wy(kk)/2,wx(kk),wy(kk)]);
118 end
119 xlim([min(xp),max(xp)]);xlabel('Horizontal Distance x (km)');ylabel('
Horizontal Distance y (km)');
120 case 3
121     display('Showcasing ambiguity.')
122     % Those are density contrasts. We now choose two.
123     % -----
124     drho1 = 400;drho2=950;
125     % Width, height and depth of the two prisms.
126     % -----
127     wx = [10 8.7];
128     wy = [100 100];
129     wz = [0.5 0.25];
130     %Offset in depth z and lateral direction x.
131     %The y-dimension is (but doesn't have to)
132     %is symetric to the profile direction
133     offsetz=[1 2];offsetx=[0 0]
134
135     % Number of prisms
136     % -----
137     np = length(offsetz);
138
139     % Sample Points along profile in x-direction.
140     % Coordinates are symetric with origin in center.
141     % -----
142     dx=0.5;xp = -20:dx:20;yp = xp*0;zp=xp*0;
143
144     % Coordinates of prisms relative to sample points
145     % -----
146     for kk=1:np
147         dx1(kk,:)=fliplr(min(xp)-wx(kk)/2:dx:max(xp)-wx(kk)/2)+offsetx(kk);dx2(
kk,:)=dx1(kk,:)+wx(kk);
148         dy1(kk,:)=dx1(kk,:)*0-wy(kk)/2;dy2(kk,:)=dy1(kk,:)+wy(kk);
149         dz1(kk,:)=dx1(kk,:)*0+offsetz(kk);dz2(kk,:)=dz1(kk,:)+wz(kk);
150     end
151
152     % Now calculate the effects with variable densities.

```

```

153 % For this we calculate all prisms for all densities.
154 % It is a bit silly, but it works.
155 % -----
156 dg1 = gravprism(drho1,dx1,dx2,dy1,dy2,dz1,dz2);
157 dg2 = gravprism(drho2,dx1,dx2,dy1,dy2,dz1,dz2);
158
159 % Visualization
160 % -----
161 figure()
162 subplot(3,1,1)
163 hold on;
164 plot(xp,dg1(1,:), 'r') %This is prism 1 with density 1
165 plot(xp,dg2(2,:), 'm') %This is prism 2 with density 2
166 xlabel('Horizontal Distance x (km)'); ylabel('Gravity Anomaly (mGal)');
167 subplot(3,1,2)
168 hold on;
169 plot(xp,0*xp, 'b-x')
170     kk=1
171     rectangle('Position',[-wx(kk)/2+offsetx(kk),offsetz(kk),wx(kk),wz(kk)], '
172     FaceColor','r');
173     kk=2
174     rectangle('Position',[-wx(kk)/2+offsetx(kk),offsetz(kk),wx(kk),wz(kk)], '
175     FaceColor','m');
176
177 set(gca,'XAxisLocation','top','YAxisLocation','left','ydir','reverse');
178 xlim([min(xp),max(xp)]); xlabel('Horizontal Distance x (km)'); ylabel('Depth (
179 m)');
180 subplot(3,1,3)%[x y w h]
181 kk=1;
182 rectangle('Position',[-wx(kk)/2+offsetx(kk),-wy(kk)/2,wx(kk),wy(kk)], '
183 FaceColor','r');
184 kk=2;
185 rectangle('Position',[-wx(kk)/2+offsetx(kk),-wy(kk)/2,wx(kk),wy(kk)], '
186 FaceColor','m');
187 xlim([min(xp),max(xp)]); xlabel('Horizontal Distance x (km)'); ylabel('
188 Horizontal Distance y (km)');
189 end

```

../Src/Gravimetry/PrismForwardModel/GravForwardModelPrismRD.m

2 Some general questions to reflect on

1. Why does the mean sea level follow the shape of the Geoid?

Solutions

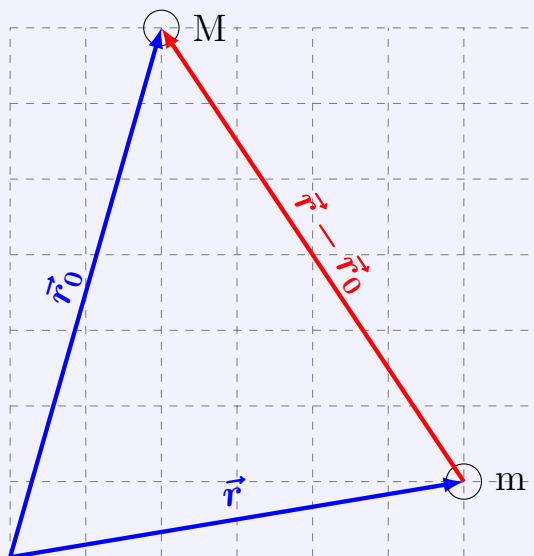
Fluids cannot sustain shear stresses. If shear stresses exist the fluid surface will adjust so that the net force is normal to the surface. Without flow the gravitational field will therefore be perpendicular to the ocean surface, which is the definition of an equipotential line. The gravitational force along

that line may vary.

2. How can you describe the gravitational attraction between two point masses if none of them is located in the origin of the coordinate system applied?

Solutions

$$\vec{F}(\vec{r}) = -GmM \frac{1}{\|\vec{r} - \vec{r}_0\|^2} \frac{\vec{r} - \vec{r}_0}{\|\vec{r} - \vec{r}_0\|} = -GmM \frac{1}{\|\vec{r} - \vec{r}_0\|^3} \vec{r} - \vec{r}_0$$



3. How does and equipotential line change by crossing an area of (a) mass deficit, and (b) mass excess?

Solutions

For a mass deficit the gravitational vectors will point away from the anomaly. Therefore the corresponding equipotential lines is curved downwards. The opposite holds for a mass excess. I didn't figure out how to draw that yet in tikz.