



Introduction to Geophysics  
R. Drews

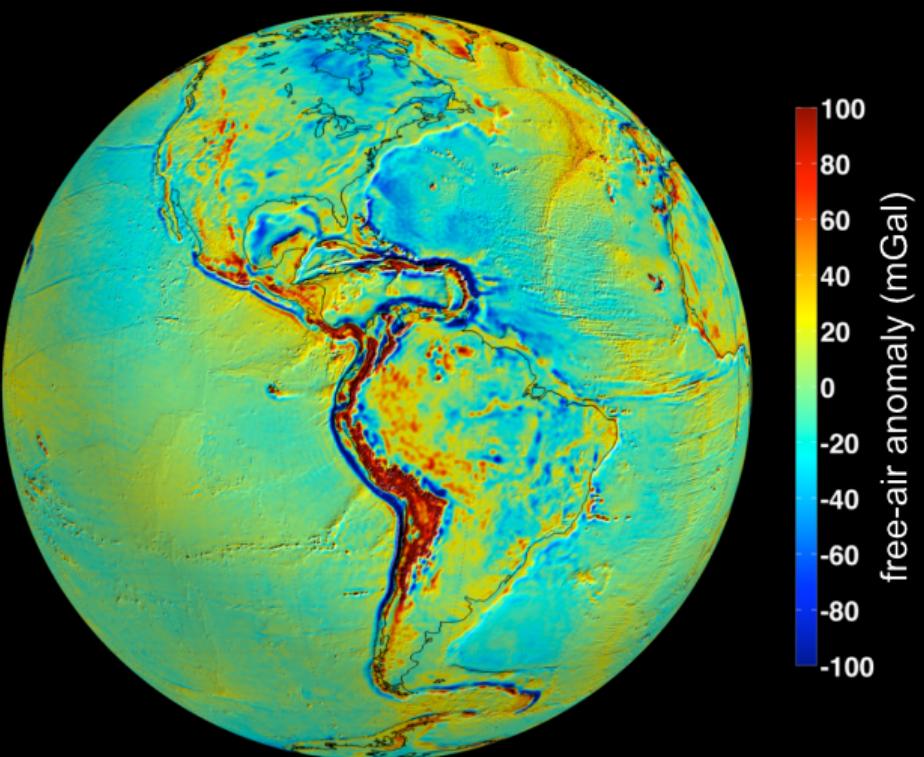
# Gravimetry



## Learning goals today:

- ▶ The gravitational force, its potential field, and how to measure it.
- ▶ The gravitational field of the Earth.
- ▶ ....

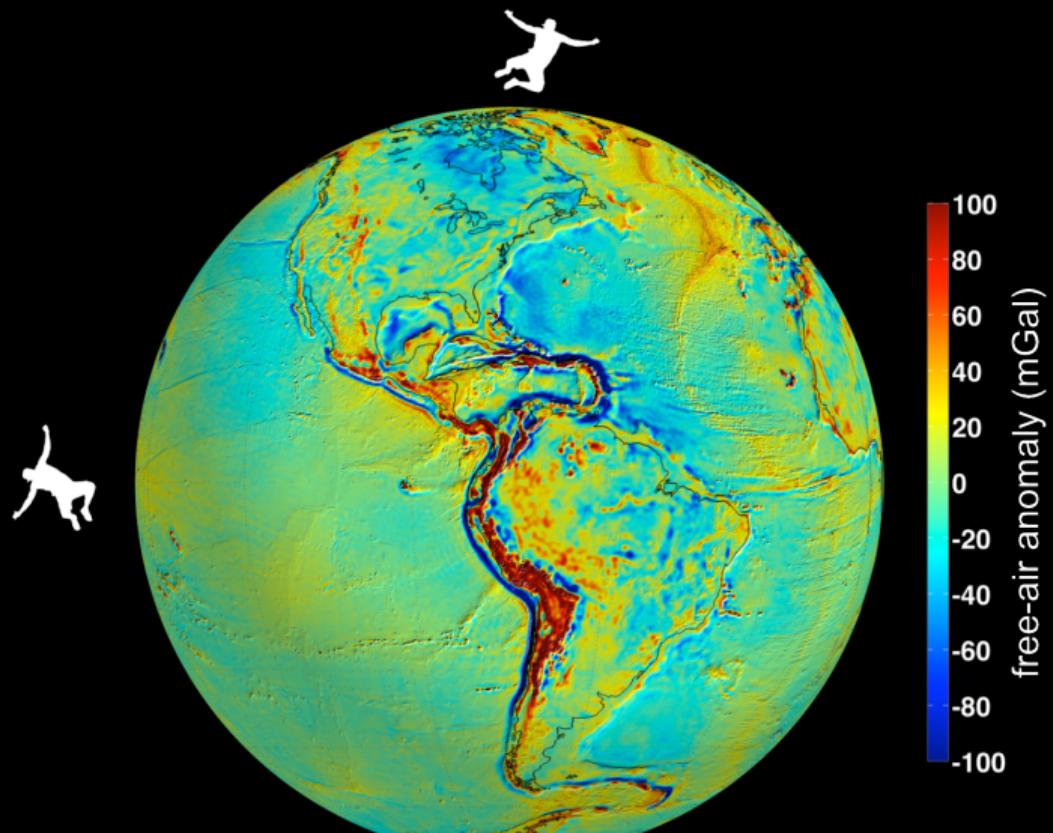
# Example: Global variability



# Example: Global variability



Your mass is constant but your weight is not.



# What is a force?



[Newton (1642-1726) / G. Johnson.]



$$\vec{F} = m\vec{g}$$

$\vec{F}$  : Force (N;  $\text{kg m s}^{-2}$ )

$\vec{a}$  : Acceleration ( $\text{m s}^{-2}$ )

m : Mass (kg)

# The gravitational force



$$\vec{F} = G \frac{mM}{r^2} \hat{r}$$

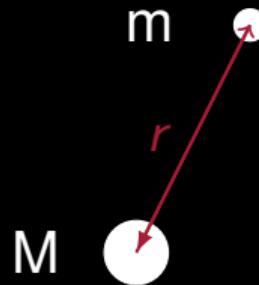
$$G = 6.674 \cdot 10^{-11} \text{ (m}^3\text{kg}^{-1}\text{s}^{-2}\text{)}$$

$\hat{r}$  : unit vector

$r$  : distance between point masses



[Newton (1642-1726) / G. Johnson.]



# Example: The gravitational constant

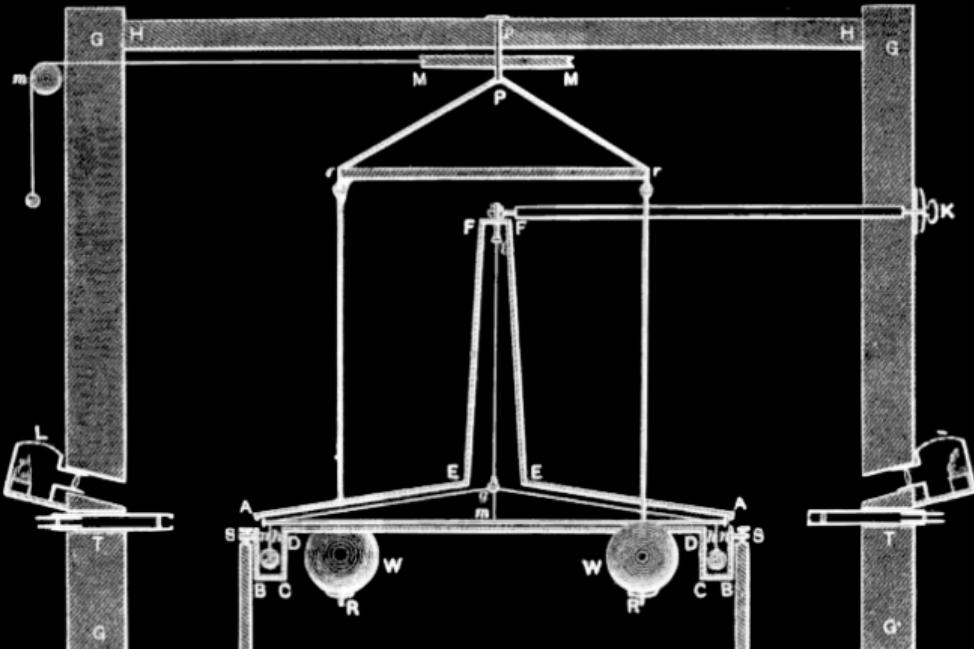
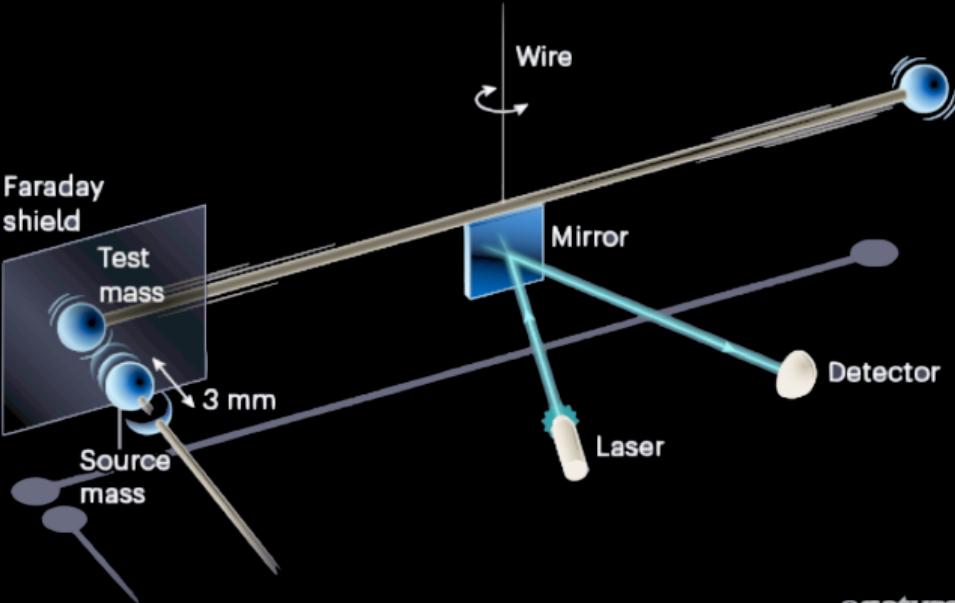


Fig. 1

Cavendish, PNAS, 1798

# Example: The gravitational constant



©nature

Westphal et al., Nature, 2021

G is the worst known constant in physics.  
Why?

# Example: Measuring acceleration



$$\vec{F} = m\vec{g}$$

$$\vec{F} = G \frac{mM}{r^2} \hat{r}$$

$$\rightarrow \vec{g} = G \frac{M}{r^2} \hat{r}$$

$$\rightarrow \frac{d^2\vec{x}}{dt^2} = G \frac{M}{r^2} \hat{r}$$

This is a differential equation.

# Example: Measuring acceleration



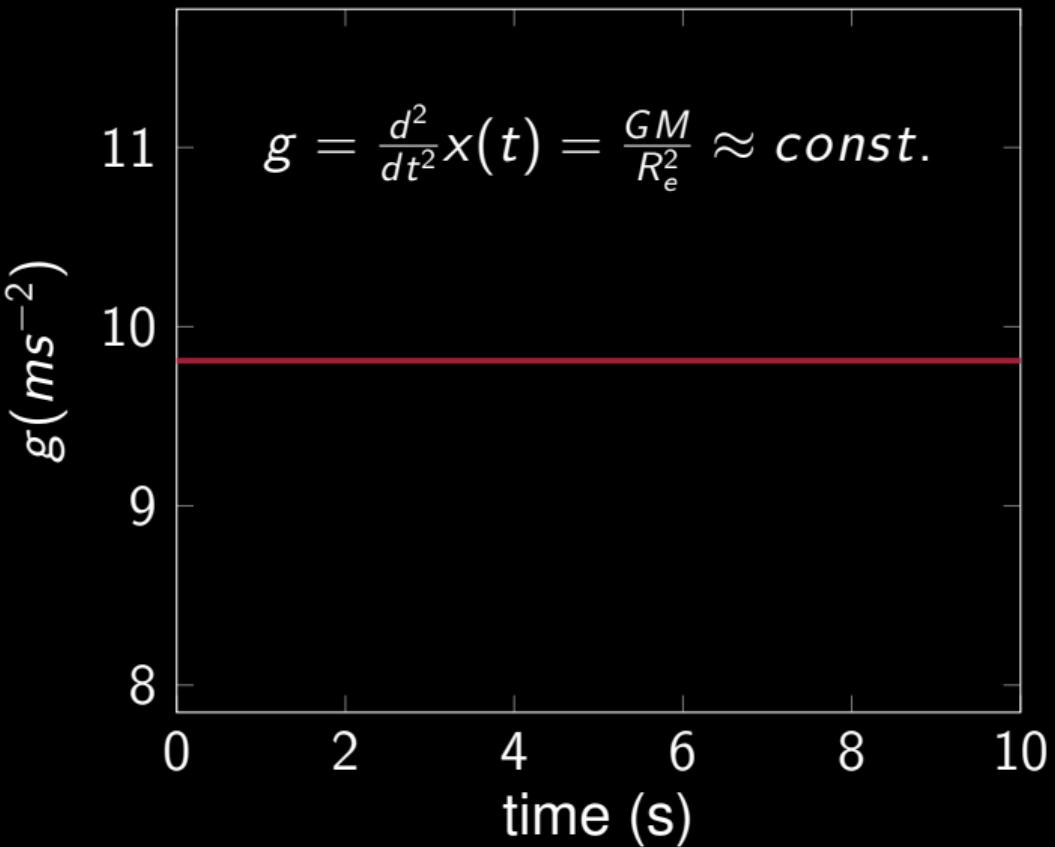
$$\frac{d^2 \vec{x}}{dt^2} = G \frac{M}{R_E^2} \approx \text{const.}$$

At the Earth's surface ( $R_E$ )  $g$  is close to constant and only vertical. (Later we will see that none of this is not quite true).

# Example: Measuring acceleration



$$g = \frac{d^2}{dt^2}x(t) = \frac{GM}{R_e^2} \approx \text{const.}$$



# Example: Measuring acceleration



$$v = \int g dt = \frac{d}{dt} x(t) = \frac{GM}{R_e^2} t + c_1$$

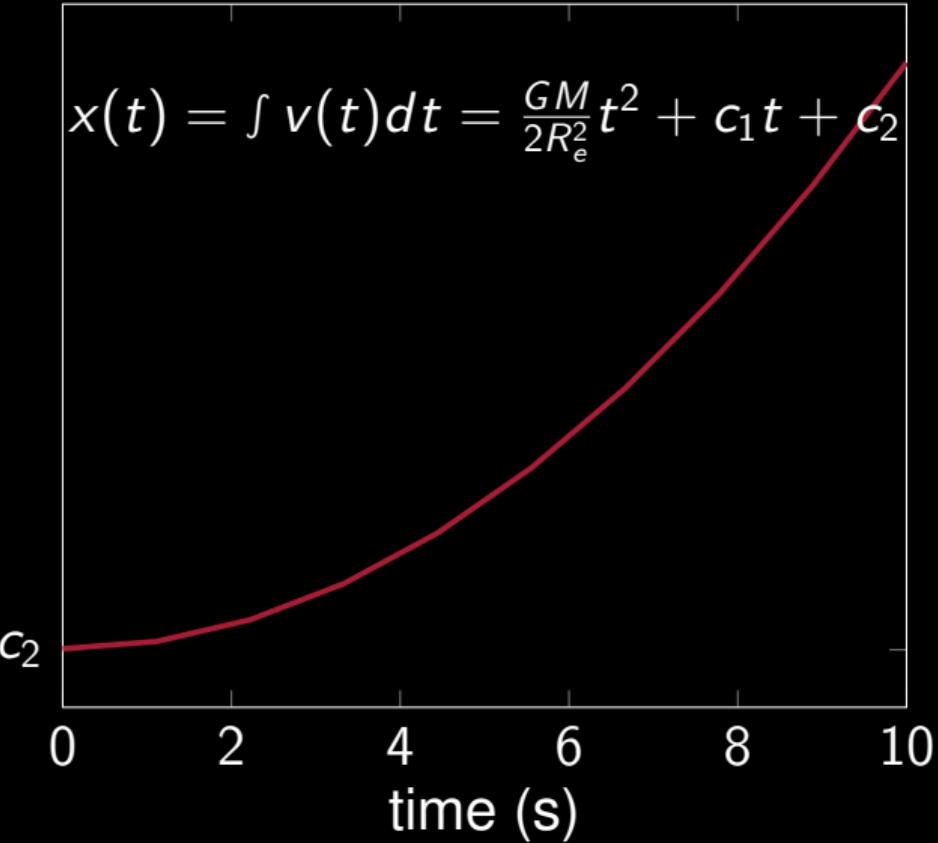


# Example: Measuring acceleration



$$x(t) = \int v(t)dt = \frac{GM}{2R_e^2}t^2 + c_1t + c_2$$

$x(m)$





$$x(t) = \frac{GM}{2R_e^2} t^2 + c_1 t + c_2$$

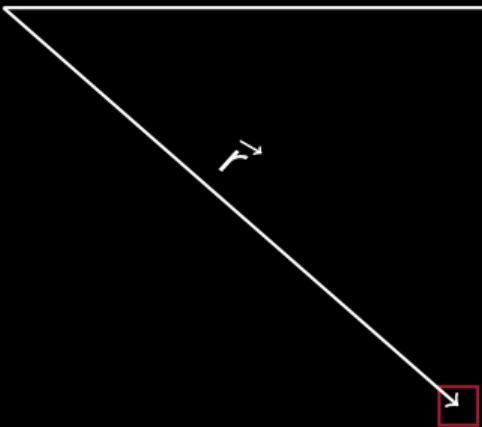
- ▶ Setting, e.g.,  $c_1 = 0$  (initial velocity) and  $c_2 = 0$  (initial position) is quite convenient.
- ▶ This is the principle of a free-fall gravimeter.



- ▶ Thanks to the Greeks we know the radius  $R_E$  for the Earth. However, its mass was unknown for a while.
- ▶ Go ahead and determine the mass of the Earth  $M$  with your Smartphone!
- ▶ There is an important first-order finding in Earth Sciences that you can (re-) discover. Which one?

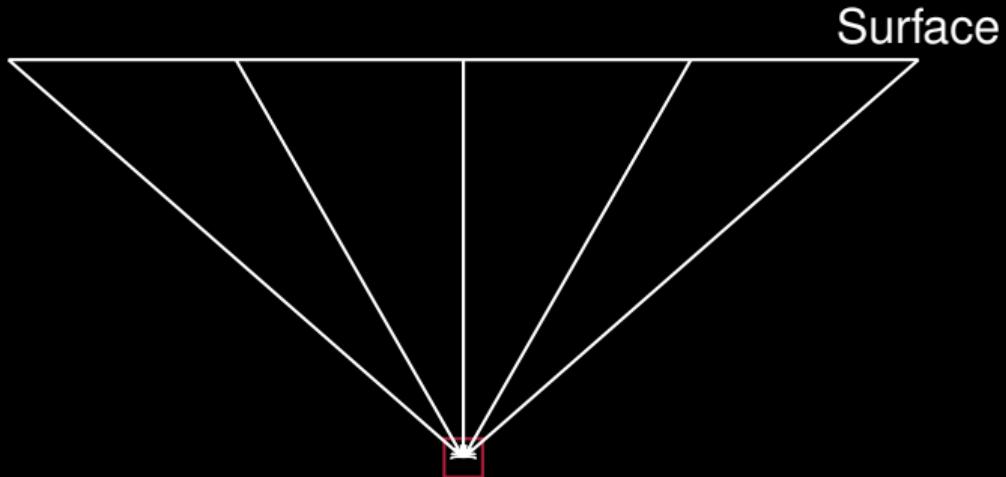


Surface



$$\vec{F} = G \frac{dM}{r^2} \hat{r}$$

For a small mass  $dM$  the point mass approximation holds.

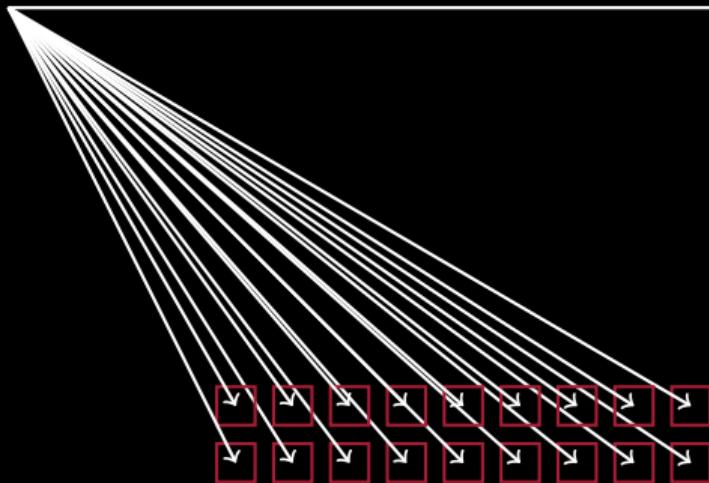


$$\vec{F} = G \frac{dM}{r^2} \hat{r}$$

Profiling across a sub-surface target results in a gravity anomaly ( $\rightarrow$  Exercises).



Surface



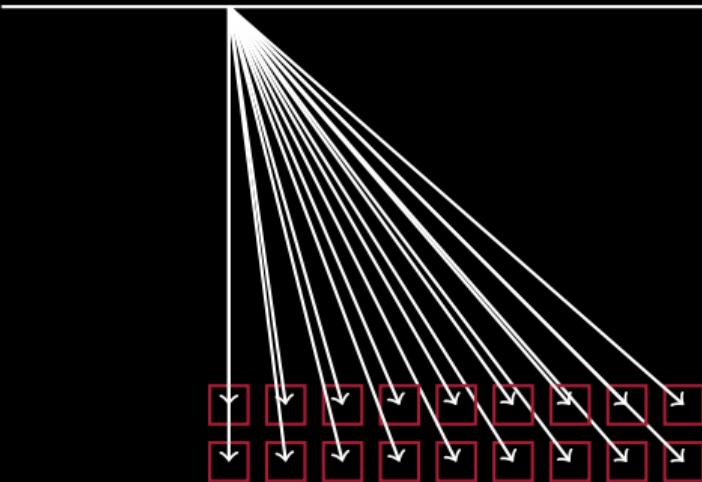
$$\vec{F}(\vec{r}) = \sum_i G \frac{dM_i}{r_i^2} \hat{r}_i$$

For  $i$  point masses the effect adds up.

# Beyond point masses

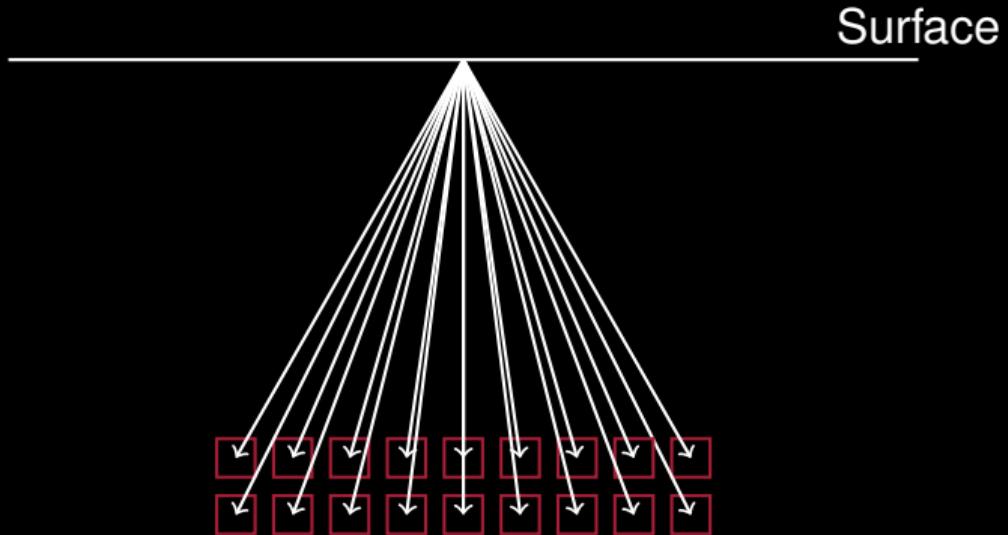


Surface



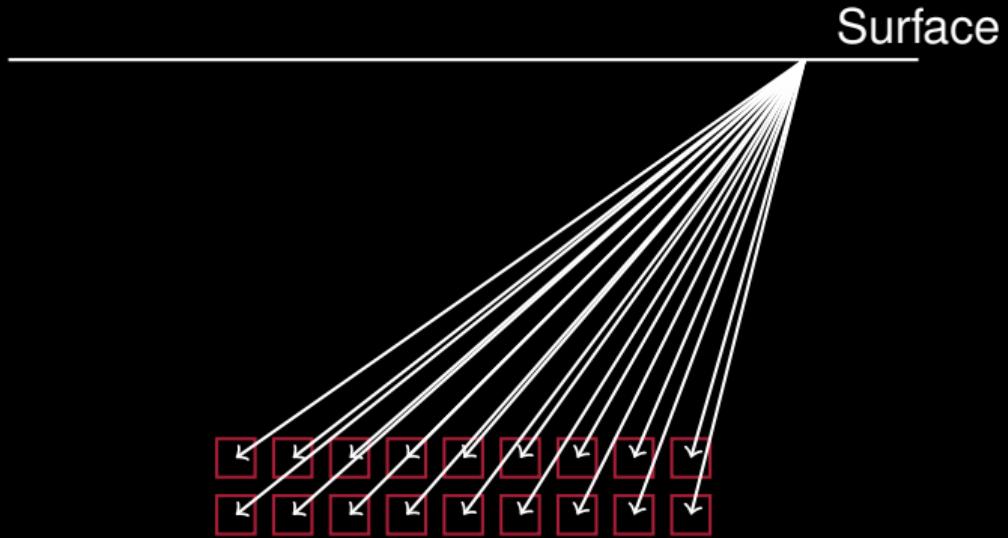
$$\vec{F}(\vec{r}) = \sum G \frac{dM_i}{r_i^2} \hat{r}_i$$

# Beyond point masses



$$\vec{F}(\vec{r}) = \sum G \frac{dM_i}{r_i^2} \hat{r}_i$$

# Beyond point masses



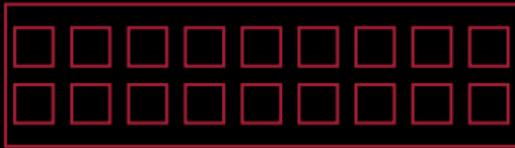
$$\vec{F}(\vec{r}) = \sum G \frac{dM_i}{r_i^2} \hat{r}_i$$



## Surface

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$$\vec{F}(\vec{r}) = G \int \rho \frac{1}{r^2} \hat{r} dV$$



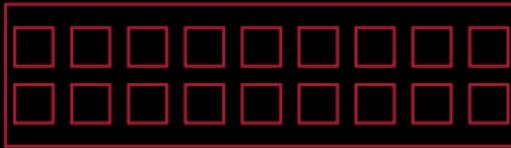
The summation can be replaced by an integration over a volume enclosing a continuous density.



## Surface

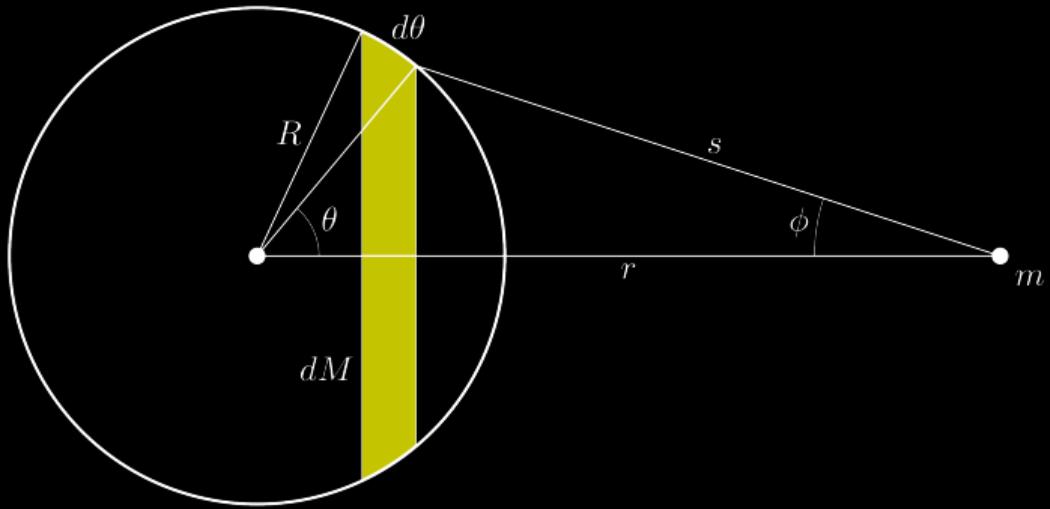
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$$\vec{F}(\vec{r}) = G \int \rho \frac{1}{r^2} \hat{r} dV$$



The integration is a triple integral. Integration limits and coordinates depend on the viewpoint. Example is a Bouger plate, in general not easy to solve ( $\rightarrow$  Exercises).

# Example: Shell



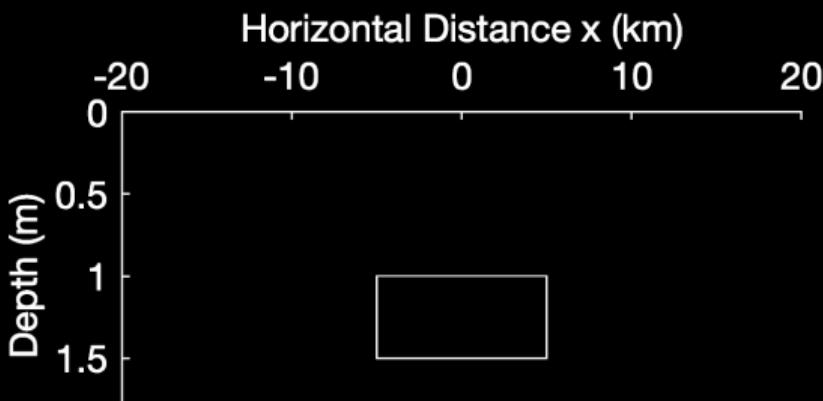
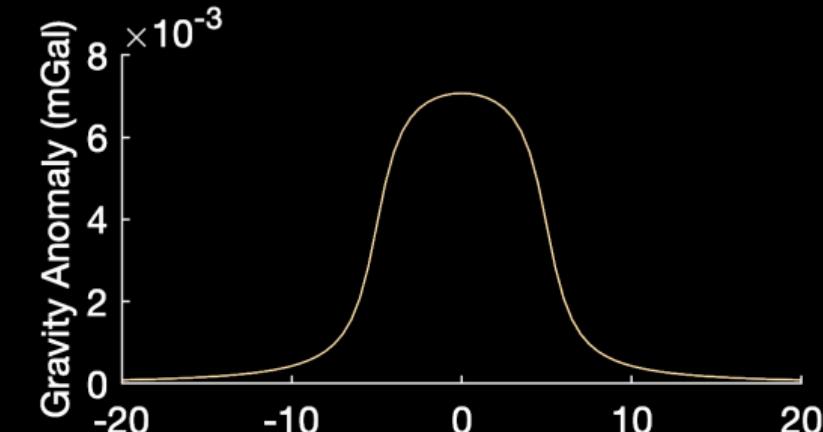
[Xaononl CC BY-SA 4.0]

Newton's shell theorem solves the volume integral inside and outside spherical objects (→ Ex.-Discussion)



- ▶ The field outside a shell is the same as the one from an equivalent point mass
- ▶ The field inside a shell is zero. Everywhere.

# Numerical forward modelling ( $\rightarrow$ Ex)



# Vector fields

