

Heterogeneity and Anisotropy

Natural geologic processes often cause hydraulic properties to vary with location in the subsurface.



Fluvial and Alluvial



Beach



Fractured Rock



Glacial

Homogeneous: a property (say K) is independent of location in the geologic formation

- i.e., $K(x,y,z) = \text{constant}$

Heterogeneous: K varies with location in the geologic formation

- i.e., $K(x,y,z) \neq \text{constant}$

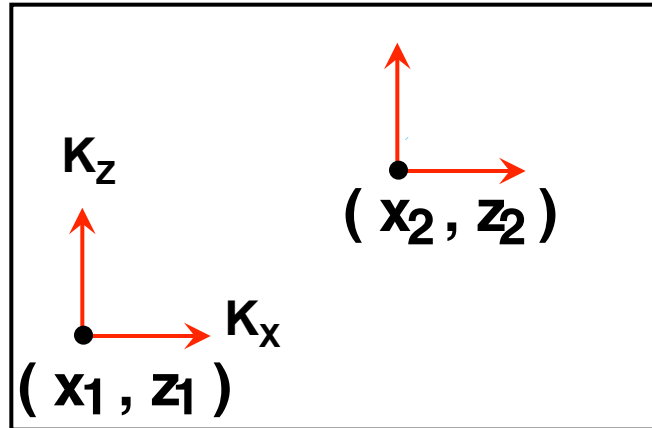
Isotropic: K is independent of the direction of measurement at a point in the geologic formation

- i.e., $K_x = K_y = K_z$

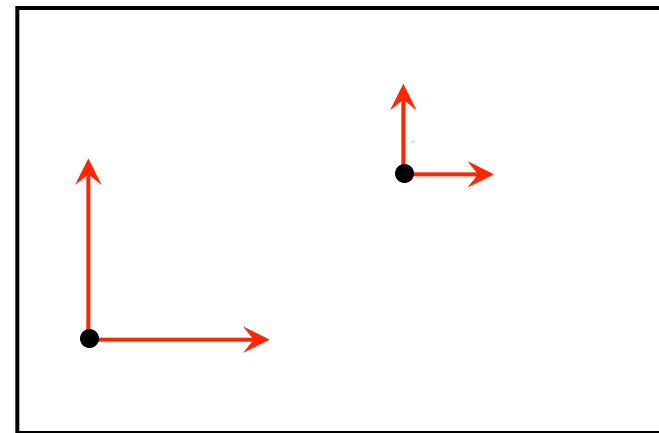
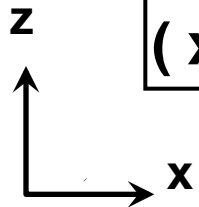
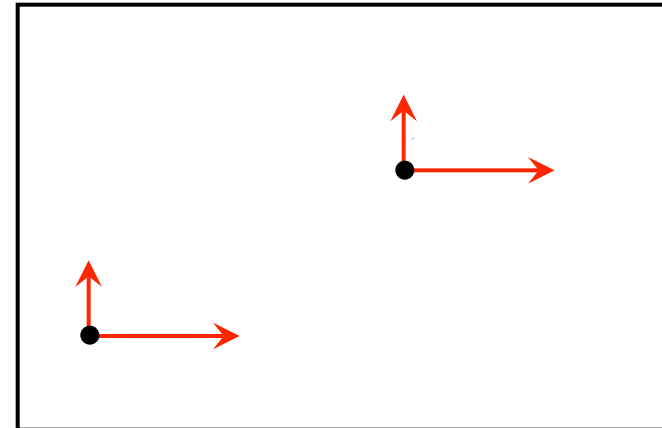
Anisotropic: K varies with the direction of measurement

- i.e., $K_x \neq K_y \neq K_z$
- K is generally larger in the horizontal direction than the vertical direction

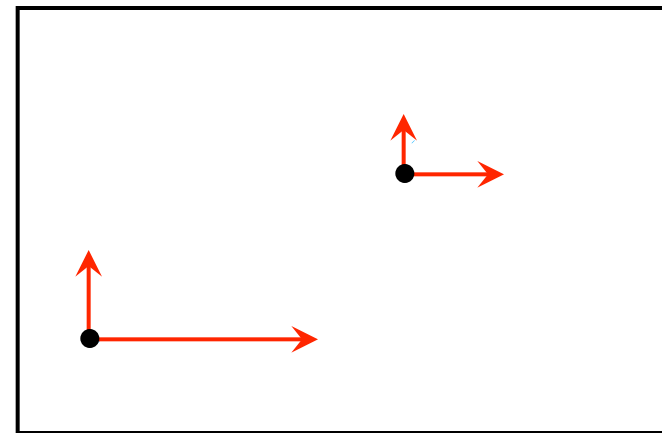
Homogeneous, Isotropic



Homogeneous, Anisotropic

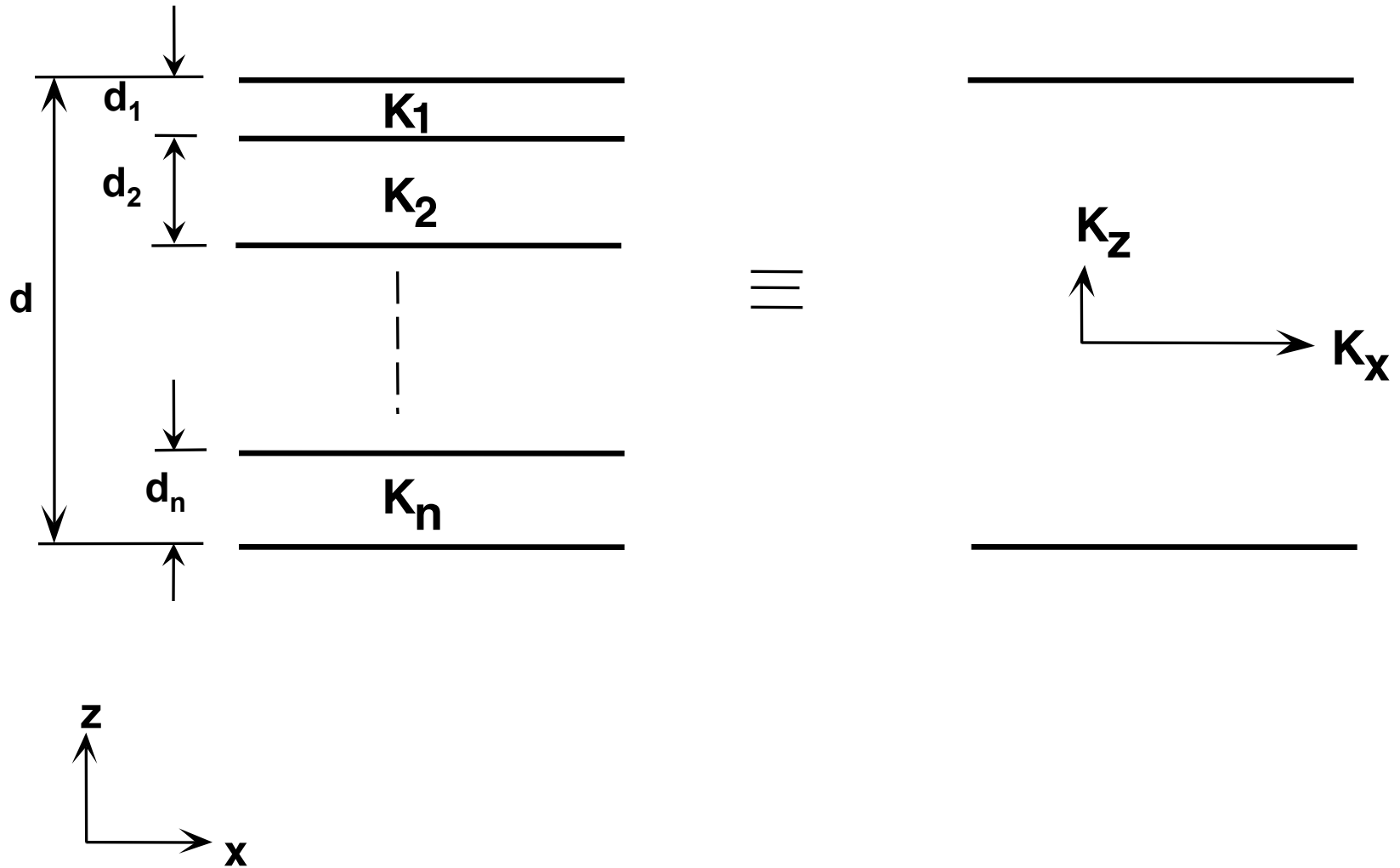


Heterogeneous, Isotropic



Heterogeneous, Anisotropic

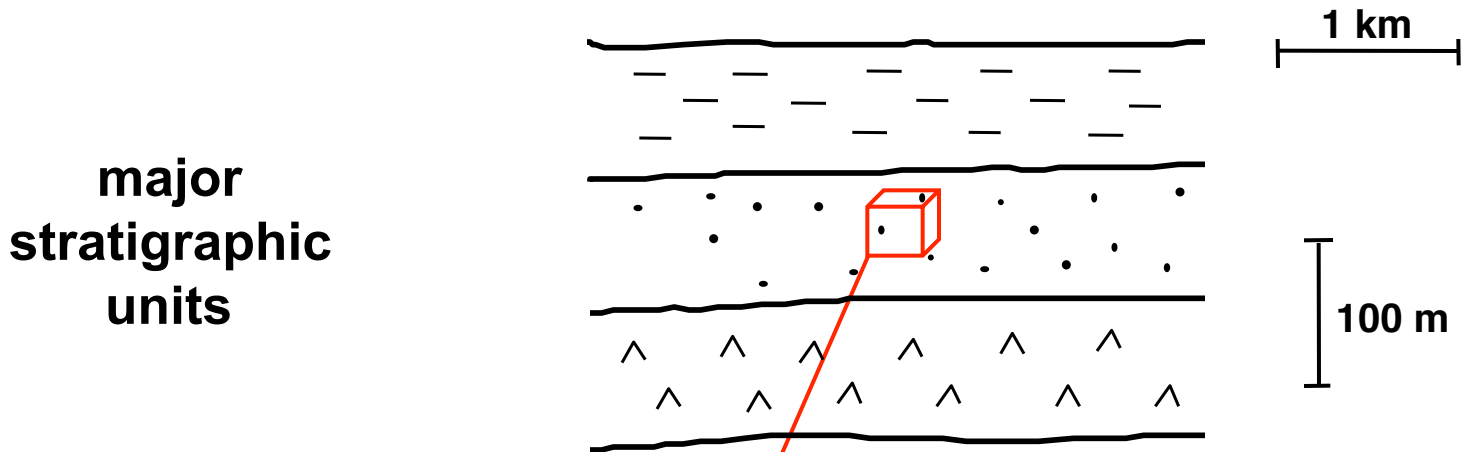
At a larger scale, heterogeneity leads to anisotropy.



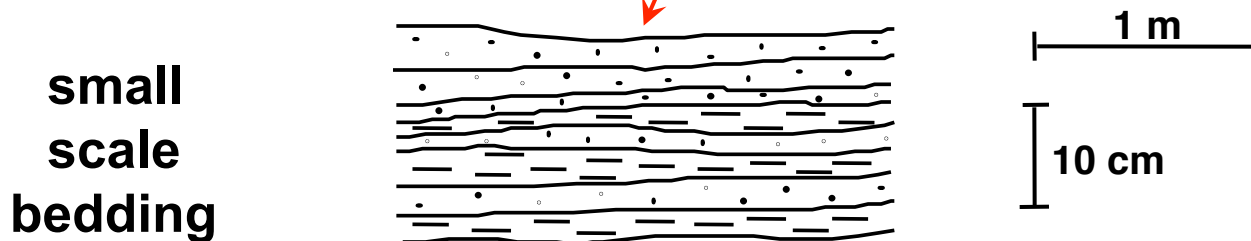
Types of Heterogeneity

1. Layering

(a) can result from large scale depositional processes

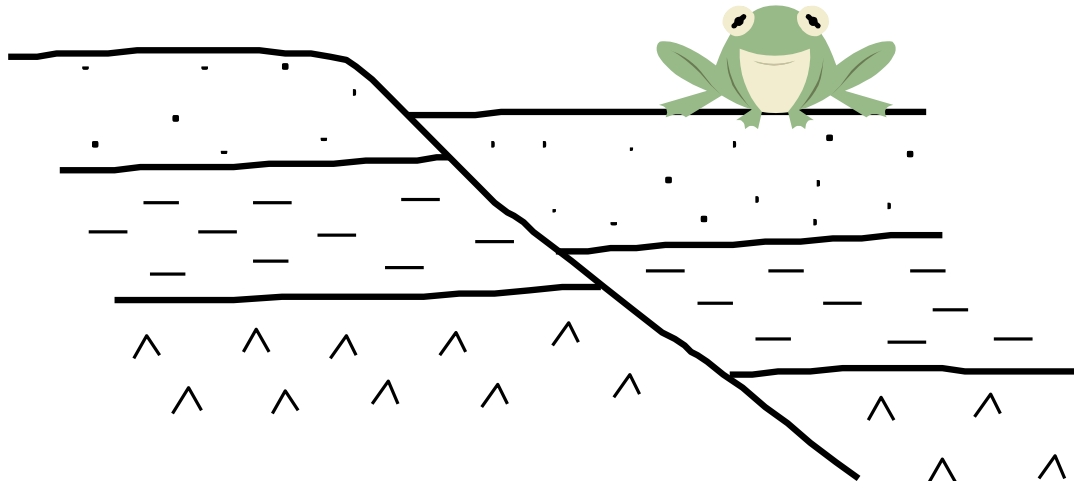


(b) or as small-scale variations within a larger system

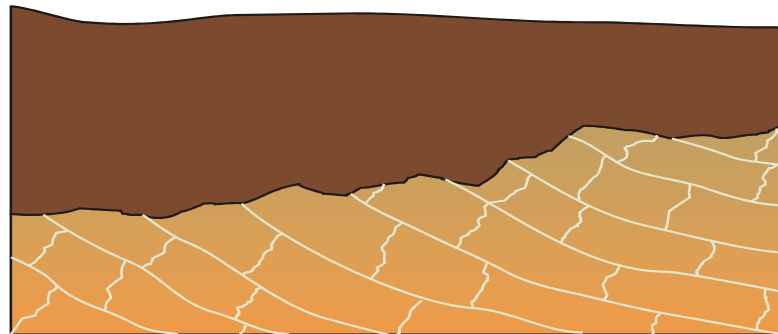


2. Discontinuities

(a) Fold, Faults, etc.



(b) Overburden-Bedrock Contacts, Unconformities

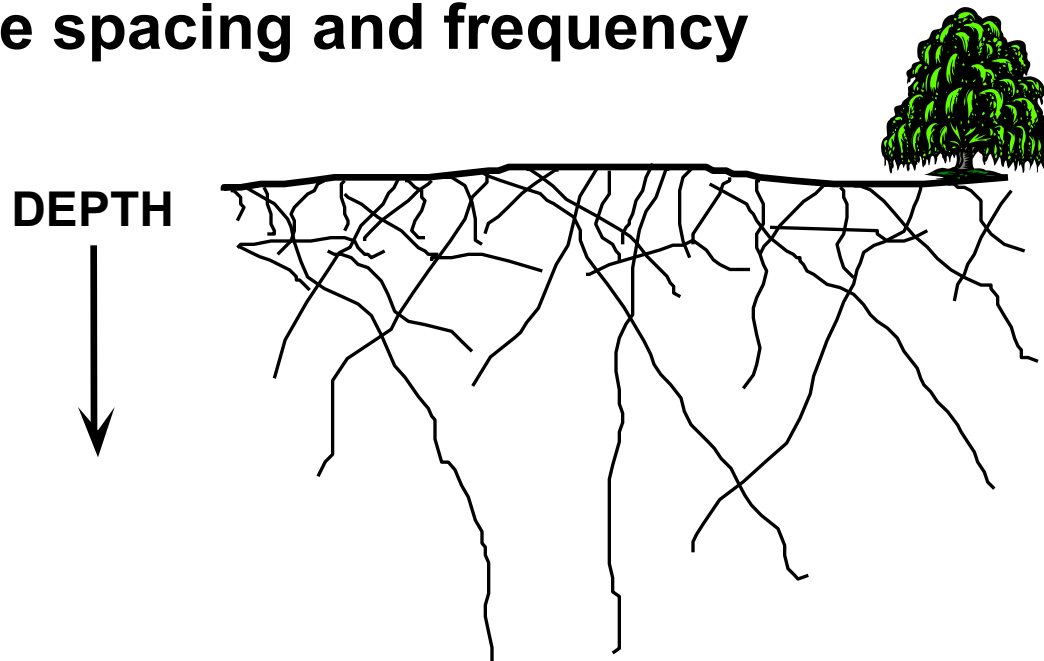


3. Spatial Trends

(a) depositional trends

- e.g., deltas, glacial outwash, and some fluvial deposits

(b) fracture spacing and frequency

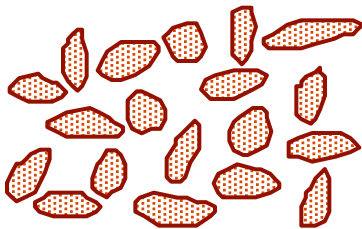




Causes of Anisotropy

1. Grain Orientation

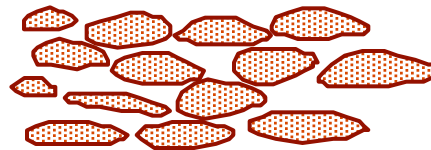
- particularly prevalent in clays



$$K_x = K_y = K_z$$

Random Orientation

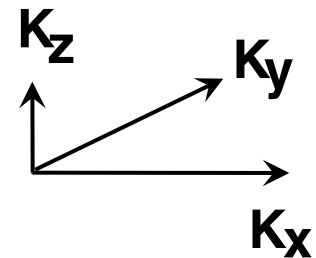
Isotropic



$$K_x = K_y > K_z$$

Directional Orientation

Anisotropic



- Anisotropy caused by grain orientation is generally no greater than 10

•i.e.,

$$1 < \frac{K_x}{K_z} < 10$$

2. Layering

- anisotropy results from layering at both the local and regional scales
- anisotropy ratios can exceed 100 (i.e., $K_x/K_z > 100$)

3. Fracture Orientation

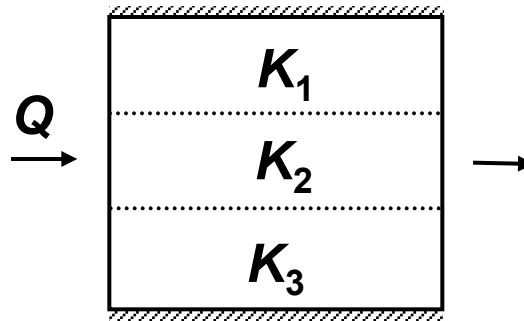
- fractures often have a preferred orientation, resulting in preferential flow in one direction



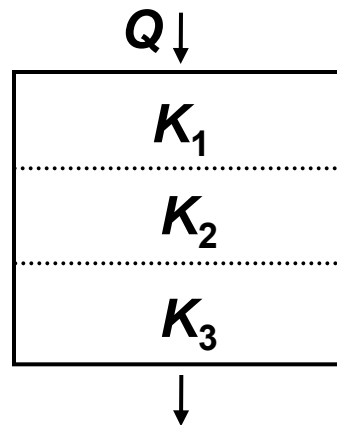
Estimating K in Layered Media

We would like to define a bulk average K value for layered media. Consider two particular cases.

1. Flow Parallel to Layering



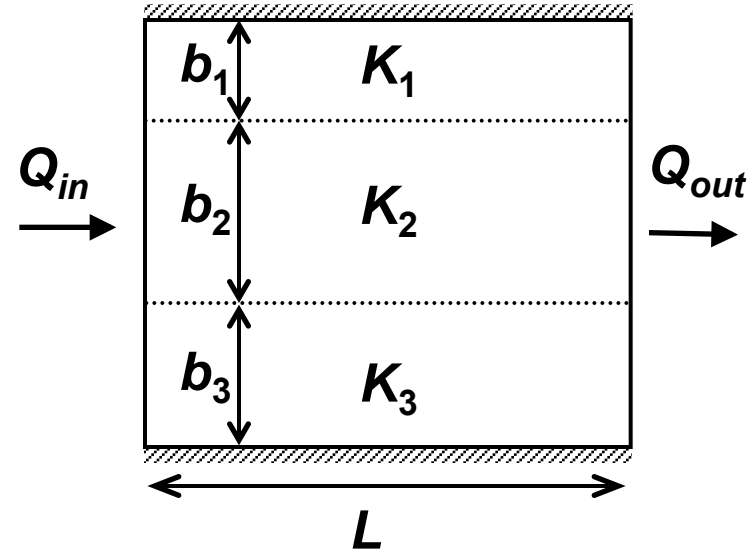
2. Flow Perpendicular to Layering



Flow Parallel to Layering

Find bulk K_x

The total volumetric flow is the sum in all layers.



$$Q = Q_1 + Q_2 + Q_3$$

$$= -A_1 K_1 \frac{\Delta h}{L} - A_2 K_2 \frac{\Delta h}{L} - A_3 K_3 \frac{\Delta h}{L}$$

For a unit thickness, $A_n = b_n \times 1$

$$Q = -\frac{\Delta h}{L} (b_1 K_1 + b_2 K_2 + b_3 K_3)$$

If $b = b_1 + b_2 + b_3$, the bulk horizontal K_x can be written as:

$$Q = -\frac{\Delta h}{L} b \cdot K_x$$

Therefore, K_x becomes:

$$K_x = \frac{b_1 K_1 + b_2 K_2 + b_3 K_3}{b}$$

In more general form:

$$K_x = \frac{\sum_{i=1}^n b_i K_i}{b}$$

Weighted Arithmetic Mean

Key points:

- Effective K is controlled by the **most** conductive layer
- Layer thickness serves as a weighting factor

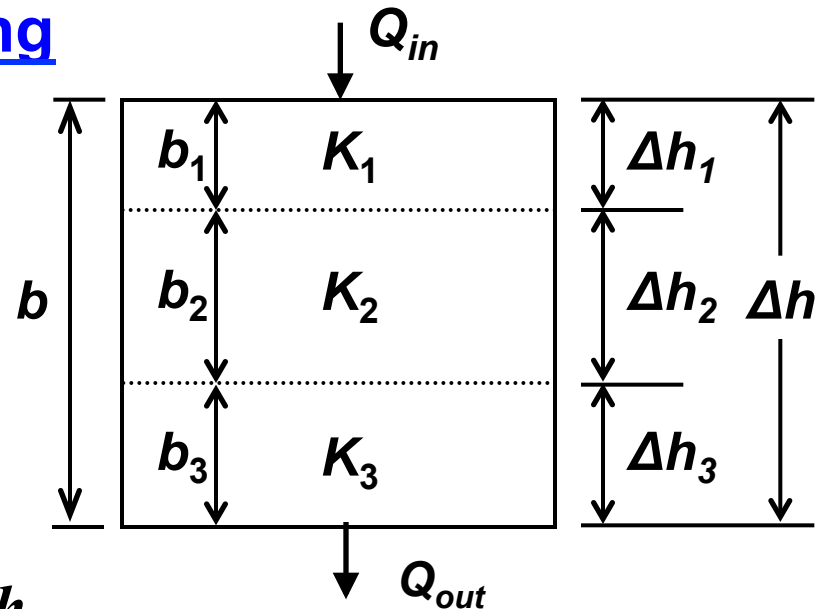
Flow Perpendicular to Layering

Find bulk K_z

In this case we have the same Q or q in all layers.

$$q = q_1 = q_2 = q_3$$

$$q = -K_1 \frac{\Delta h_1}{b_1} = -K_2 \frac{\Delta h_2}{b_2} = -K_3 \frac{\Delta h_3}{b_3}$$



The sum of head losses in each layer must equal the total head loss, Δh .

$$\begin{aligned} \Delta h &= \Delta h_1 + \Delta h_2 + \Delta h_3 \\ \Delta h_1 &= -\frac{q_1 b_1}{K_1} \\ &= -\frac{q_1 b_1}{K_1} - \frac{q_2 b_2}{K_2} - \frac{q_3 b_3}{K_3} = -q \left(\frac{b_1}{K_1} + \frac{b_2}{K_2} + \frac{b_3}{K_3} \right) \end{aligned}$$

We can transform Darcy's Law for bulk vertical K_z to get:

$$q = -K_z \frac{\Delta h}{b} \quad \text{or} \quad \Delta h = -\frac{qb}{K_z}$$

Equating the previous two expressions, K_z becomes:

$$K_z = \frac{b}{\frac{b_1}{K_1} + \frac{b_2}{K_2} + \frac{b_3}{K_3}}$$

More generally:

$$K_z = \frac{b}{\sum_{i=1}^n \frac{b_i}{K_i}}$$

Weighted Harmonic Mean

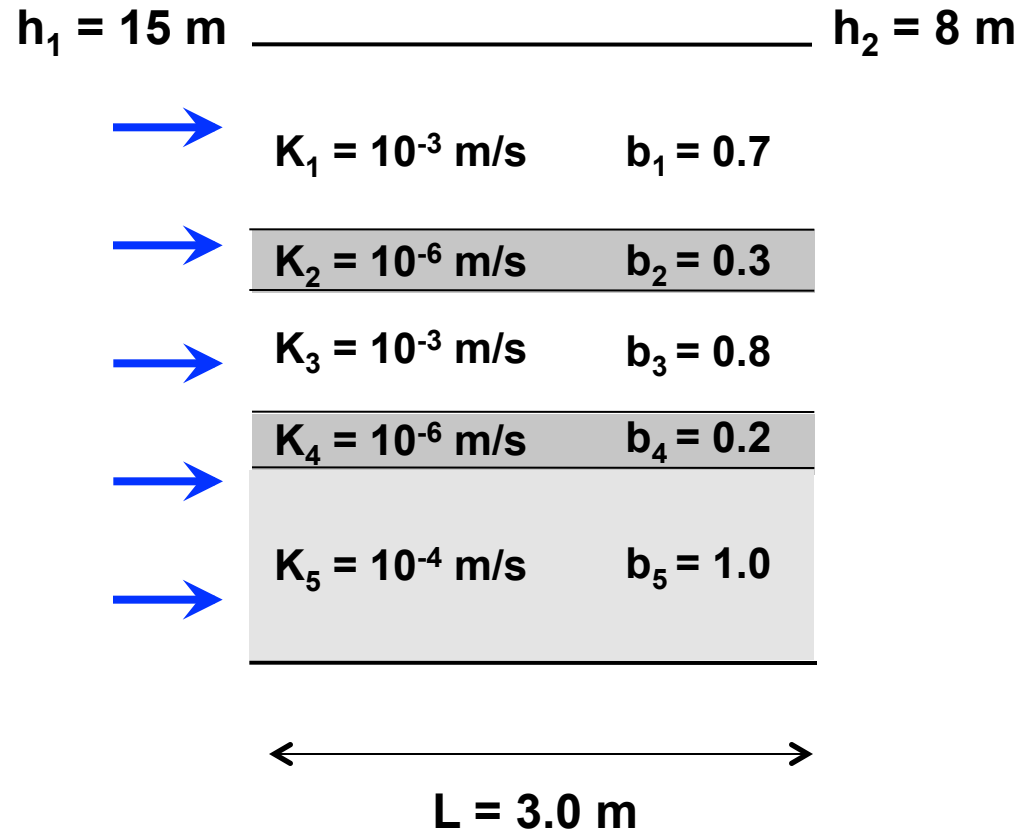
Key points:

- Effective K is controlled by the *least* conductive layer
- Layer thickness serves as a weighting factor

Example

Calculate K

Calculate q



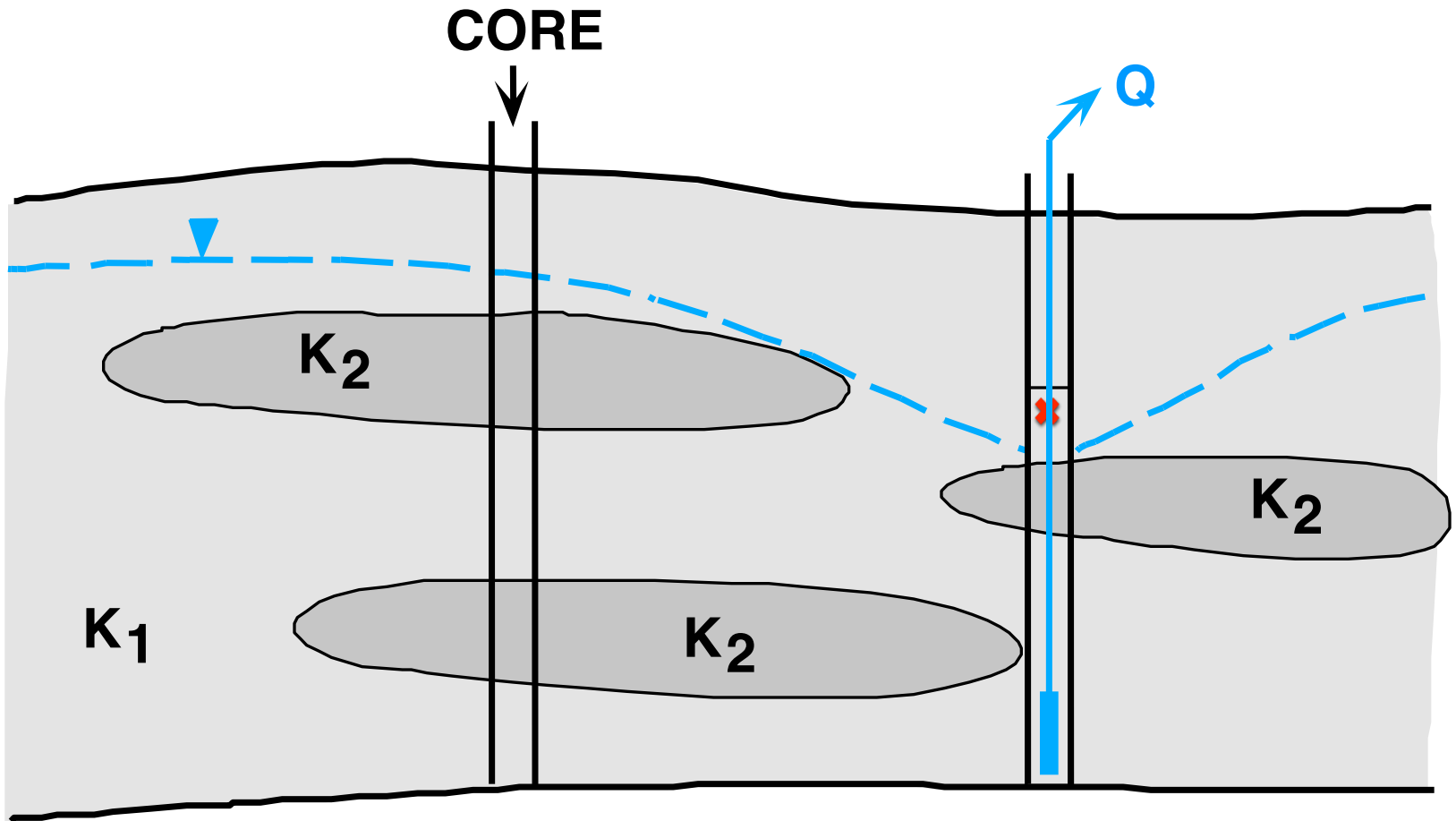
Example

Calculate K

Calculate q

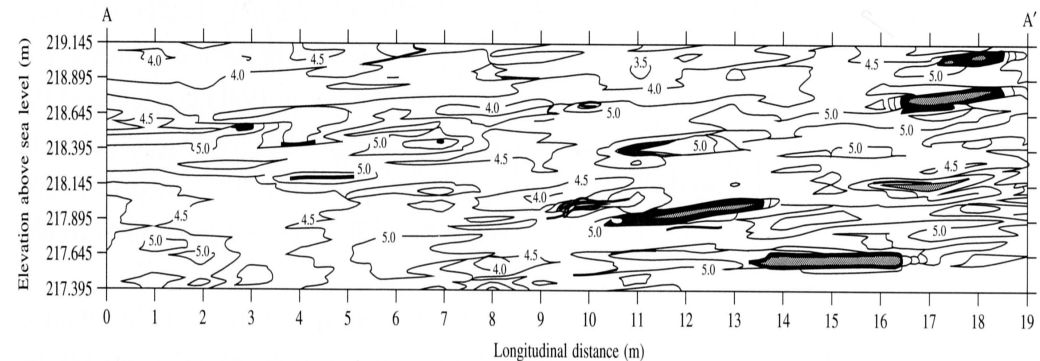
		$h_1 = 15 \text{ m}$
$K_1 = 10^{-3} \text{ m/s}$	$b_1 = 0.7$	
$K_2 = 10^{-6} \text{ m/s}$	$b_2 = 0.3$	
$K_3 = 10^{-3} \text{ m/s}$	$b_3 = 0.8$	
$K_4 = 10^{-6} \text{ m/s}$	$b_4 = 0.2$	
$K_5 = 10^{-4} \text{ m/s}$	$b_5 = 1.0$	
		$h_2 = 8 \text{ m}$

How do we define this layering or anisotropy?



We even have layering in “homogeneous” materials. Is the Borden sand really homogeneous? How would this influence flow and transport?

**Contours are given as $-\log K$
(e.g., $K=1 \times 10^{-4}$ m/s is 4.0 contour)**



**Hydraulic conductivity distribution of the Borden aquifer in cross-section.
(Fetter, Contaminant Hydrogeology, 1999)**