Simulations of stratified rotating turbulence and comparison with oceanic wave turbulence

Rémy Guillermin - supervised by Pierre Augier

Research Training III





2024-2025

୶ୢଌଊ

Abstract

In fluid mechanics, turbulent flows occur when the Reynolds number $\text{Re} = UL/\nu \gg 1$, where U and L are the characteristic velocity and length, respectively, and ν is the kinematic viscosity. This condition is satisfied in cases such as the flow of an inviscid fluid¹, e.g., superfluids; when the flow is extremely fast, e.g., jet engine exhaust; or when the flow occurs at large scales, e.g., climate phenomena.

When we look at large scale climate flows such as the ocean or the atmosphere, it is mandatory to consider two major changes in the basic Navier-Stokes equations: We need to account for the rotation of Earth as well as the stratification in density (ocean) or temperature(atmosphere). In this project, we have been focused on the effect of a density stratification of the flow, which will induce internal wave turbulence.

1 Theoretical background

The Navier-Stokes equations we will work with are

$$\frac{\partial \rho}{\partial t} + \vec{\nabla} \left(\rho \vec{u} \right) = 0 \tag{1a}$$

$$\rho \left(\frac{\partial \vec{u}}{\partial t} + \left(\vec{u} \cdot \vec{\nabla} \right) \vec{u} \right) = -\vec{\nabla} p - \rho g \hat{z} + \mu \Delta \vec{u} \quad (1b)$$

It is useful to define the Froude number as ${\rm Fr}^2=U^2/gH$ as the ratio of the advection term to the gravitational term. One must be familiar with the Boussinesq's approximation [Bou97] in order to work with density stratified fluid. After applying this approximation on system 1, and after some rework, we find the wave propagation equation written as

$$\partial_t^2 \vec{u} = \left(N^2 \nabla_h^2 \nabla^{-2} \right) \vec{u} \tag{2}$$

where ∇_h and ∇ are the horizontal gradient operator and the spatial gradient operator, respectively, and $N^2 = -\frac{g}{\rho} \frac{\partial \rho}{\partial z}$ is the Brunt-Väisälä frequency [Ped79]. This equation allows us to find the dispersion relation of the gravity wave

$$\omega = \pm N \sin \theta \tag{3}$$

where $\theta = k_z/|\vec{k}|$ is the angle between the wavevector \vec{k} and the z-axis. This means that ω is bounded inside the range [-N, +N].

We will consider an incompressible flow $(\vec{\nabla} \cdot \vec{u} = 0)$, which means that we can use the poloidal-toroidal decomposition [SW92] in the vertical direction. This allows us to express the kinetic energy E_k^2 as the sum of the poloidal and toroidal contribution $E_K = E_p + E_t$ where

$$E_{p,t} = \frac{1}{2} \int_{\mathbb{R}^3} (\vec{u}_{p,t})^2 d^3 \vec{r}$$

We can also express the potential energy as

$$E_{\phi} = \left\langle \frac{g^2 \delta \rho^2}{2N^2 \rho_0^2} \right\rangle \tag{4}$$

where ρ_0 is the mean density of the fluid and $\delta \rho$ are the density perturbation along the density stratification $\rho_0 + \rho_b$.

¹a fluid with no viscosity

²In this paper, the specific kinetic energy E_k/ρ will be named kinetic energy

References References

References

- [Bou97] Joseph Boussinesq. Théorie de l'écoulement tourbillonnant et tumultueux des liquides dans les lits rectilignes a grande section. fre. Paris, Gauthier-Villars et fils, 1897. URL: http://archive.org/details/thbeoriedelbeco01bousrich (visited on 12/05/2024).
- [Ped79] Joseph Pedlosky. Geophysical Fluid Dynamics. en. New York, NY: Springer US, 1979. ISBN: 978-0-387-90745-1 978-1-4684-0071-7. DOI: 10.1007/978-1-4684-0071-7. URL: https://link.springer.com/10.1007/978-1-4684-0071-7 (visited on 10/17/2024).
- [SW92] Burkhard J. Schmitt and Wolf von Wahl. "Decomposition of solenoidal fields into poloidal fields, toroidal fields and the mean flow. Applications to the boussinesq-equations". In: The Navier-Stokes Equations II Theory and Numerical Methods. Ed. by John G. Heywood et al. Berlin, Heidelberg: Springer Berlin Heidelberg, 1992, pp. 291–305. ISBN: 978-3-540-47498-2.