

Observational constraints on climate sensitivity derived from the 1971-2017 global energy budget

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Earth Energy Budget

On a global scale, the Earth Energy budget can be written with the **forcing/feedback framework**, assuming a linear radiative response with surface temperature:

$$N = F + \lambda T$$

The **equilibrium climate sensitivity (ECS)** is the equilibrium surface temperature increase induced by an abrupt doubling in atmospheric CO₂:

$$ECS = -\frac{F_{2\times}}{\lambda}$$

ECS is reached only at equilibrium: it cannot be directly measured in the real world and requires extremely long runs to be simulated.

Effective Climate sensitivity

In practice the climate sensitivity is derived from the transient regime and is called effective climate sensitivity (effCS):

$$\text{effCS} = -F_2 \times \frac{T}{N - F}$$

| | |
|-----------------------------------|--|
| histeffCS | CO ₂ effCS |
| from the historical energy budget | from 150yr of abrupt 2xCO ₂ |

CO₂effCS is a precise estimate of the ECS (although biased by ~17%).

The aim of this study is to derive an observational constraint on CO₂effCS

The Pattern Effect

The radiative response of the Earth depends on the geographic pattern in surface air temperature. (Sherwood et al. 2020; Gregory et al. 2020)

This effect is called **the pattern effect**. It arises from:

- ▶ Mix of radiative forcings
- ▶ Lag-dependent responses to forcings
- ▶ Unforced variability.

The pattern effect leads to apparent time variations in λ and thus:

$$\text{histeffCS} \neq \text{CO}_2\text{effCS}$$

Observational constraints on CO₂effCS

- Compute histeffCS using a regression on observations of the energy budget from 1971-2017:
 - ▶ F non aerosol (F_{NA}) from Sherwood et al. 2020
 - ▶ F aerosol (F_{AER}) from Bellouin et al. 2020
 - ▶ T from the Cowtan and Way 2014 corrected for the surface bias due to satellite data (Richardson et al. 2016)
 - ▶ N from Ocean heat content estimates derived from optimal interpolation of ocean in situ data (Ishii et al. 2017, Cheng et al. 2017)

- Include all sources of observational uncertainty

- Quantify the pattern effect from CMIP6:
 - ▶ Internal variability i.v. in λ
 - ▶ A distance between histeffCS and CO₂effCS

Summary

Approach

- 1 Regression on **observational data** from 1971 to 2017
- 2 Include all sources of **observational uncertainty**
- 3 Quantify and include the **pattern effect**

Results

- 1 We find an CO₂effCS of **5.5 K** with a 5-95% interval of 2.4-35.6 K
- 2 **CO₂effCS below 2.4 K is inconsistent with the observed energy budget**
- 3 The upper bound is not constrained by observations
- 4 The main observational sources of uncertainty are
 1. the **aerosol forcing** and 2. the **earth energy balance** estimate from the ocean heat content

Comparison with previous studies

The two most recent studies estimating climate sensitivity from observational data are **Lewis and Curry (2018)** and **Sherwood et al. (2020)**

| CO ₂ effCS estimates: | 5% | Med. | 95% |
|----------------------------------|-----|------|------|
| Lewis & Curry (2018) | 1.2 | 1.7 | 2.7 |
| Sherwood et al. (2020) | 2.0 | 4.3 | 16.1 |
| Chenal et al. (2022) | 2.4 | 5.5 | 35.6 |

L&C18 use the IPCC AR5 gaussian aerosol forcing and do not correct for the **pattern effect**

Both L&C18 and S20 use a state difference method with $N(\sim 1860) = +0.2 \text{ W/m}^2$ but it is an uncertain value (probably < 0 ?). The regression approach removes this dependence

The higher bound difference between our estimate and S20 is unsignificant

Comparison with previous estimates

- 1 Our lower bound estimate is **1.2 K** above Lewis and Curry (2018) and **0.4 K** above Sherwood et al. (2020)
- 2 The difference with Sherwood et al. is explained by the high ocean heat uptake of $+0.2 \text{ W/m}^2$ they use as the reference state in 1860
- 3 The further difference with Lewis and Curry is explained by the fact that, in addition, they ignore the pattern effect and use gaussian aerosol forcing from AR5

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