

# Literature Review Template

Cole Lord-May

## Contents

Mahrt and Gamage 1987	2
Ohata 1989	4

# Mahrt and Gamage 1987

## Observations of Turbulence in Stratified Flow

### Summary

This study also explores means of analyzing aircraft-measured turbulence data. As such, it is also somewhat less directly pertinent, but makes interesting comments on examining periodic, shear-driven turbulence. This paper is worth re-reading somewhat often, as it provides good physical insight into the mechanisms responsible for turbulence and what it “looks like”.

### Important Figures

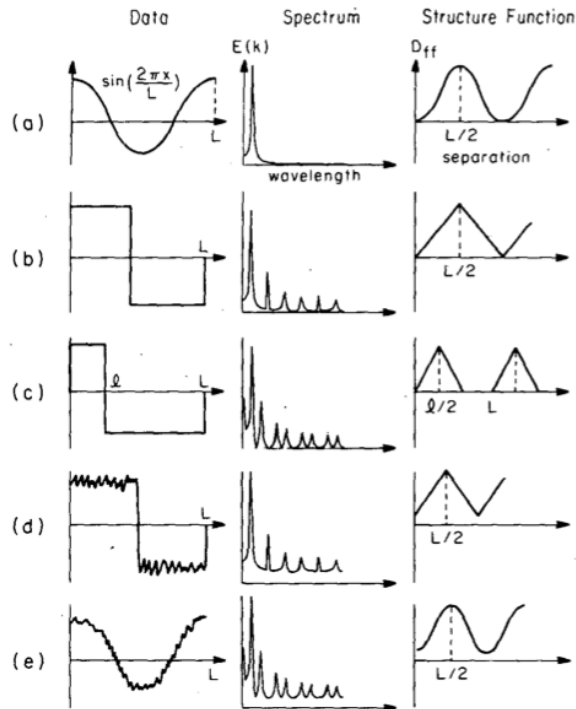


Figure 1: Illustrations of structure functions given synthetic input data

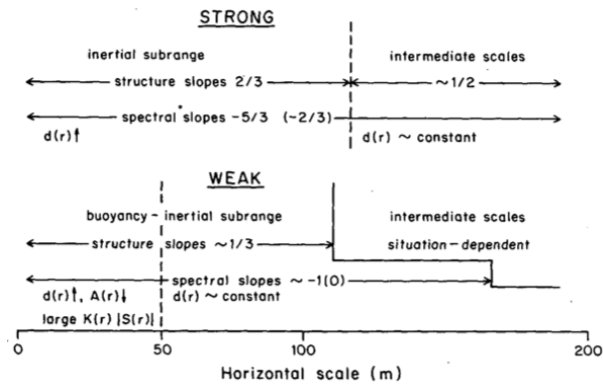


Figure 2: Scale regimes for weak and strong turbulence.  $d(r)$  is buoyancy length,  $S(r)$  is skewness, and  $S(r)$  is kurtosis.

## Questions and Comments

- The structure functions of various orders seem highly applicable to studying turbulence (as has been done many times before) fig. 1 shows the value of using a structure function instead of data or spectra. The structure function extracts the general shape of the data (the integral of the data) but smooths fluctuations and avoids the many frequencies in spectra. These might be useful inputs into PCA!
- They define the buoyancy length (as a measure of stratification strength) as  $D_{ww} / \left( g / \Theta_0 D_{\theta\theta}^{1/2} \right)$ . This might be a more general measure than  $z/L$ . I haven't seen this explored much as a classification of inversion, so it would be worth investigating. It doesn't require computation of vertical gradients, and is not tied to horizontal scales.
- fig. 2 is a very visually clear separation of turbulence. They first classify turbulence into weak and strong, and then draw out the horizontal scales over which different spectral slopes exist. Being able to distinguish turbulence in such a clear way would be very illuminating to our turbulence on the glacier slope.
- The cospectructure functions that they introduce allow them to distinguish between bora turbulence and nocturnal boundary layer turbulence. This would be fascinating if we could distinguish between origins of turbulence origins. However, I suspect that below the wind speed max, most of the turbulence would have the same origin.
- They find that at intermediate scales of turbulence, spectral slopes are influence by turbulence strength and not by mechanics. This suggests that a fairly significant key of the turbulence-in-stratified-fluids puzzles isn't being properly accounted for (as this should never be the case if you truly have critical phenomena).

# Ohata 1989

## Katabatic Wind on Melting Snow and Ice Surfaces (I) Stationary Glacier Wind on a Large Maritime Glacier

### Summary

Observations made on the 40 km-long San Rafael Glacier in the Patagonia Northern Icefield from 1983-1984. Data was collected from four observation sites on the ground and two observation sites on the glacier. In addition to higher-frequency measurements of mean temperature, mean wind speed, and temperature inversion at a fixed height, 19 wind profile measurements were made at random times throughout the day. Glacier wind height  $h$  is measured as the maximum height at which predominant wind direction is still down-slope, or when wind speed decreases below  $1 \text{ m s}^{-1}$  – whichever is lower. The measured thickness of the glacial wind varied between 30 m and 120 m. They measured that the height of the wind maximum  $h_m$  was typically half of  $h$ ,  $h_m/h \approx 0.5$ . This is substantially higher than Prandtl's model, which is attributed to a higher surface roughness. They concluded that even with a low lapse rate, ( $\Delta T < 1^\circ \text{C}$  over 100 m), glacier winds are still the predominant daytime feature, occurring 80 – 90% of the time. There appears to be a consistent fluctuation in mean wind speed, with a nighttime period of 3 hours and a daytime period of 1-2 hours.

### Important Figures

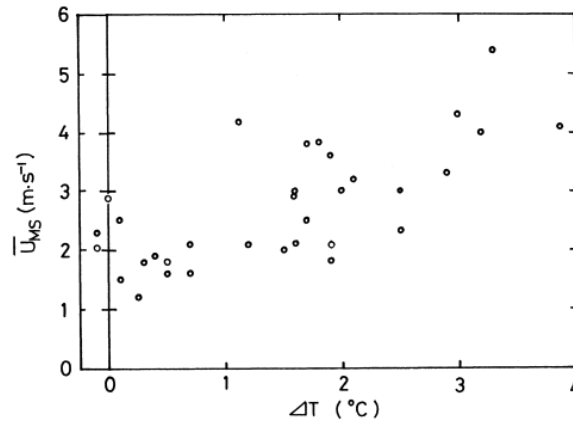


Figure 3: Comparison of observed temperature inversion strength  $\Delta T$  and wind speed – confirming that high wind speeds are driven by strong temperature inversions

### Questions and Comments

- The two glacier sites are approximately 50 m away from the edge of the glacier. Does this boundary condition affect the measured winds?
- Wind profile measurements were made at “random times” throughout the day between 8:00 and 17:00. It’s not clear how these random times were determined. Was it truly random or whenever it was most convenient? If the latter, subjectivity (When weather conditions are most favourable) might have affected the launch times, which could change the wind profiles recorded.
- On the note of  $h_m/h \approx 0.5$ : Their vertical resolution is approximately 10 m. We now know that this is not high enough resolution to capture the true wind maximum that typically resides a couple of metres above the surface.
- I think it might be interesting to further explore geostrophic forcing. It affects katabatic winds in antarctica, but is claimed not have no affect here since  $\mathcal{O}(g\Theta \sin \alpha / \Theta_0) = 10^{-2}$  and  $\mathcal{O}(fv) = 10^{-3}$

to  $10^{-4}$ . Particularly if looking at time-dependent cases, it's not immediately clear to me that this perturbation would be entirely negligible. Especially since this is an argument used to explain why katabatics are less frequent in the mornings (when geostrophic winds are stronger) and more frequent in afternoons.

## References

- [1] L. Mahrt and N. Gamage. “Observations of Turbulence in Stratified Flow”. In: *Journal of Atmospheric Sciences* 44.7 (Apr. 1987), pp. 1106–1122. DOI: 10.1175/1520-0469(1987)044<1106:00TISF>2.0.CO;2.
- [2] Tetsuo Ohata. “Katabatic Wind on Melting Snow and Ice Surfaces (I) Stationary Glacier Wind on a Large Maritime Glacier”. In: *Journal of the Meteorological Society of Japan* 67.1 (Jan. 1989), pp. 99–112. DOI: 10.2151/jmsj1965.67.1\\_99.