The influence of the Southland Current on the circulation within Pegasus Bay

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Objectives

Investigate the effect of the Southland Current (SC) on the coastal circulation at Pegasus Bay (PB) using a combination of long term realistic modelling and *in situ* measurements.

1. What is the contribution of the SC to the total currents on the inner shelf of PB?
2. What is the impact of remote forcing in modelling efforts at the inner shelf of PB?
3. What are the implications of underestimating the SC flow in the prediction of arbitrary passive tracers transport and dispersion in PB?

Introduction

The East coast of New Zealand is a notably complex area in terms of interaction between oceanic and coastal circulation, due to the existence of important western boundary current systems flowing along the continental slopes and outer shelves (Jannelle & Flemming, 2005). The SC is a northward/northeastward oceanic current that extends from Stewart Island in the south to the Chatham Rise in the north, formed at the subtropical convergence west of New Zealand (Chiswell, 1996). Most of it bends offshore towards the Chatham Rise (Figure 1), but a small portion bends towards PB, in the form of an anticyclonic eddy around BP (Carter & Herzer, 1979). A net southerly alongshore depth-averaged transport is reported in previous work, potentially associated with an eddy in PB.

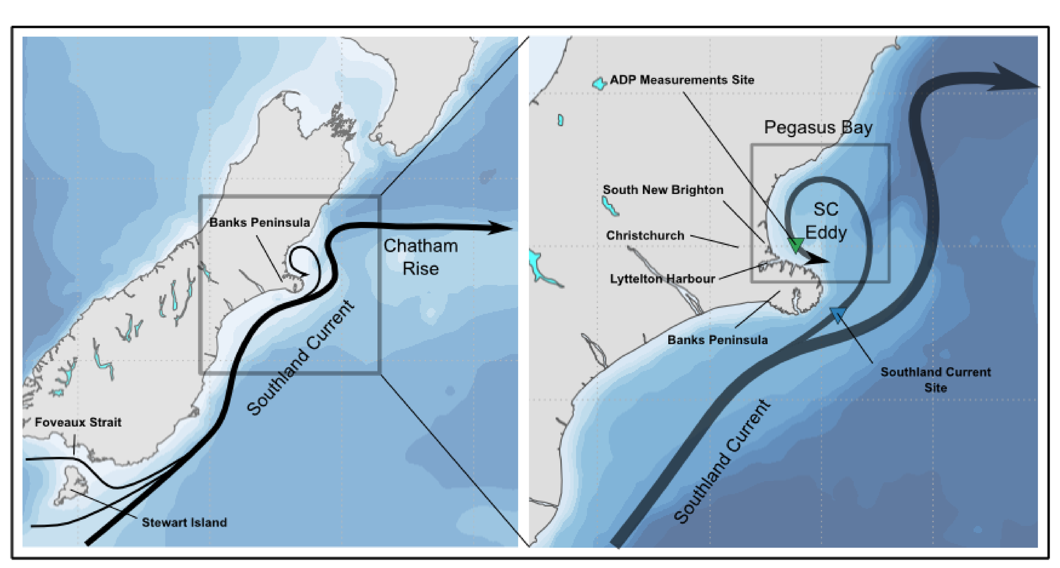

Figure 1: Figure caption

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Methods

A calibrated 10-year ROMS hindcast was carried out. Two different runs were undertaken. A CONTROL run consisted of full realistic hydrodynamical forcing (Figure 2 and Table 2) and a TEST run was deprived of most of the SC incoming flow (Table 1).

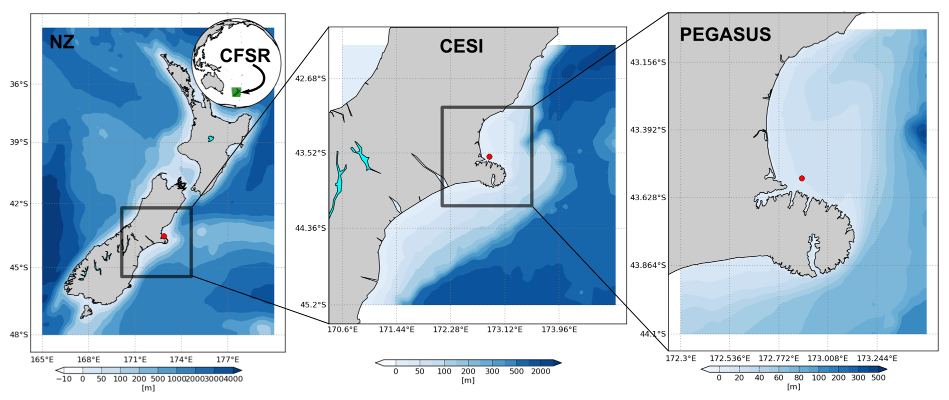

Figure 2: Figure caption

Experiment	Nests
CONTROL	CFSR → NZ → CESI → PEGASUS
TEST	CFSR → PEGASUS

Table 1: Table caption

	NZ	CESI	PEGASUS
Resolution	0.08°	0.03°	0.004°
Boundary	CFSR	NZ	CESI
Tides	No	No	Yes
Meteo	WRF12km	WRF12km	WRF12km

Table 2: Table caption

Results

As in Figure 3, the SC clearly bifurcates, where some of the water recirculates into PB in the form of an anticyclonic eddy. There is a clear difference in the magnitude of the mean currents predicted by the different runs. The average flow from CONTROL explained 80% of the observations, while TEST, only 33%.

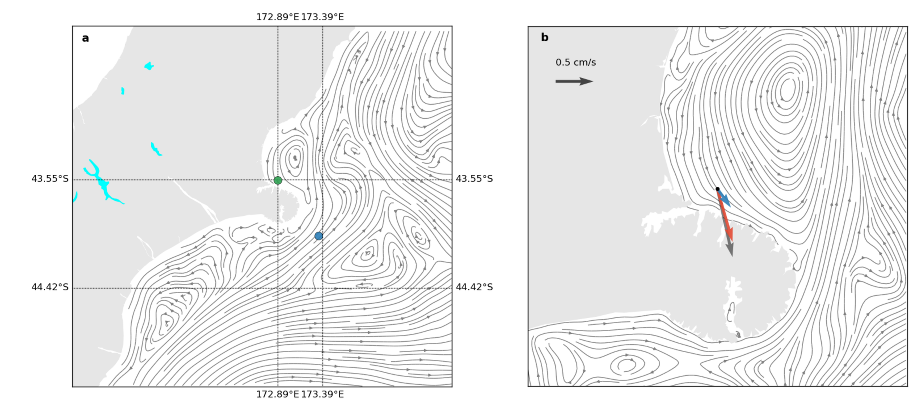

Figure 3: Figure caption

According to Figure 5, CONTROL shows good skill in reproducing the total currents. The residual currents are notably matched by the model for most events. TEST, however, produced relatively poor results.

A 20 h lag between perturbations in the SC and PB (Figure 4) seems to occur. The spectra confirms the variability match at time scales of western boundary current variability between both locations, corroborating a clear influence of the SC in the PB circulation.

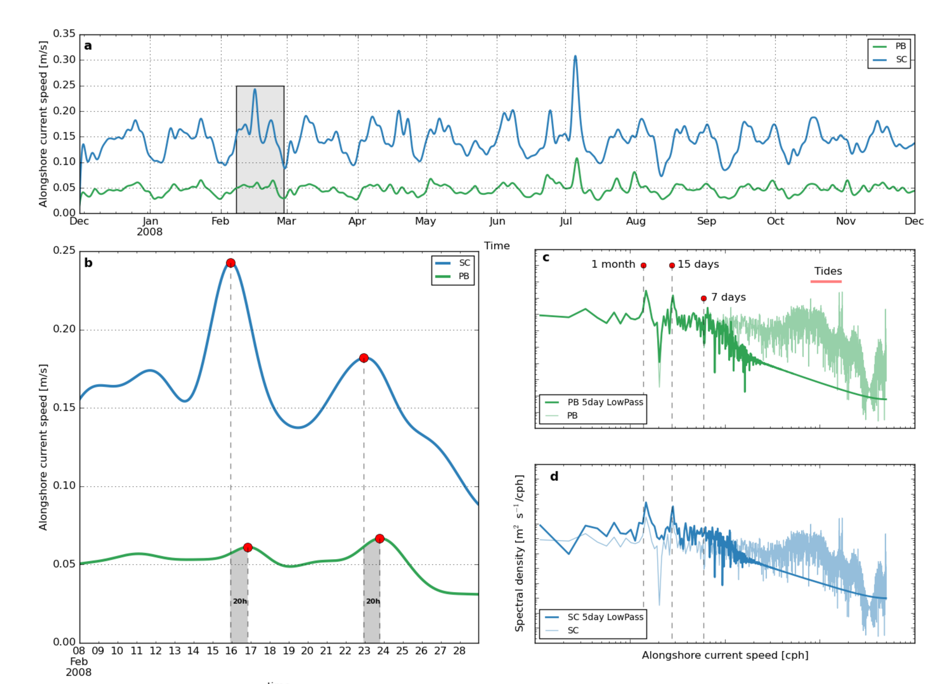

Figure 4: Figure caption

Model Validation

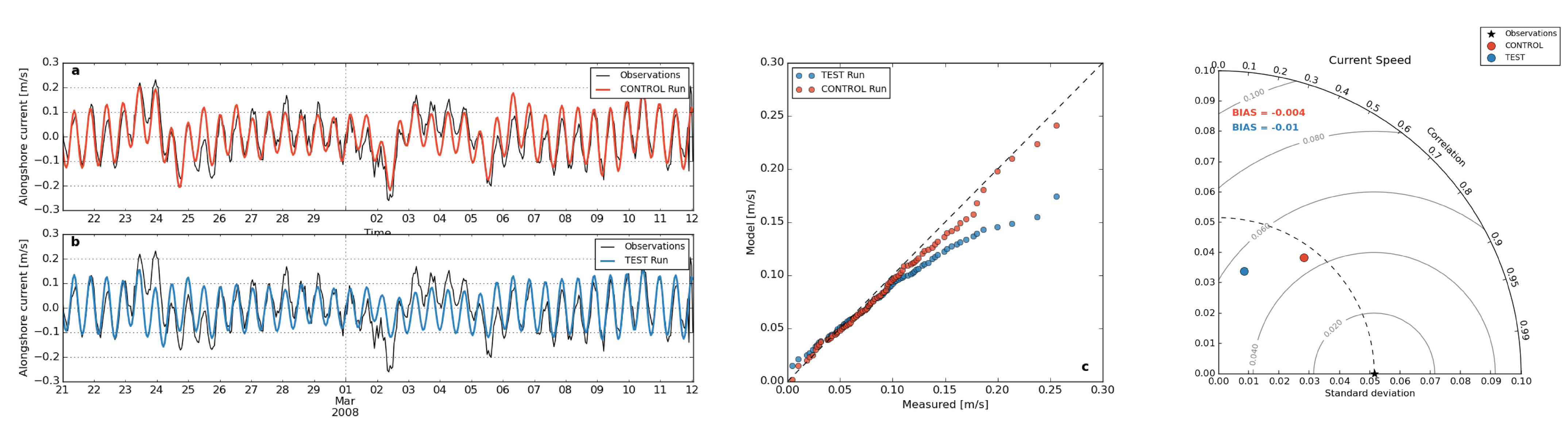

Figure 5: Figure caption

Plume propagation

A passive tracer point source was introduced in both experiments at a constant rate at the surface during 15 days, and than halted and allowed to evolve for 15 additional days. The plume extends over a greater spatial extent on CONTROL, especially in the SE direction. There is also more northward dispersion in CONTROL. The TEST run overestimates concentrations near the source and underestimates at larger distances.

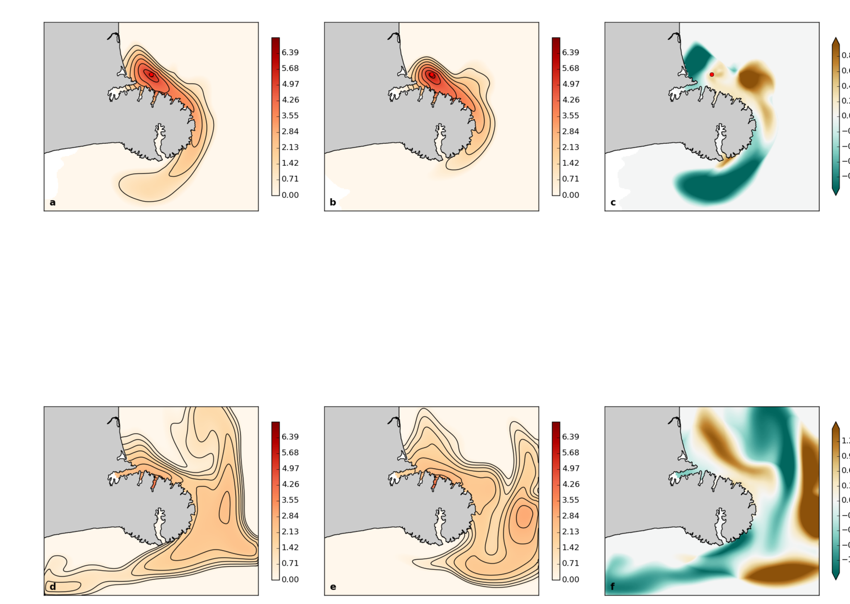

Figure 6: Figure caption

Conclusions

1. The SC has a substantial influence in the circulation at PB, accounting for more than half of the total currents at times and controling most of its variability.
2. The absence of proper open ocean forcing greatly affects model performance for this particular area.
3. Underestimating the SC implies in less efficient dispersion of passive tracer plumes at the PB inner shelf.

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