An experiment on the appearance of the complete evaporation of a land planet with a general circulation model

陸惑星における完全蒸発状態の 発生に関する大気大循環モデル実験

Planetary and Space Group

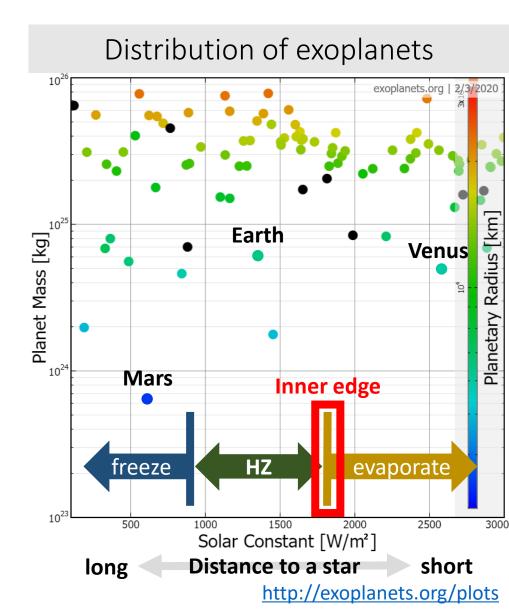
Yoshida, Tetsuji 吉田 哲治

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Introduction

Exoplanets in habitable zone

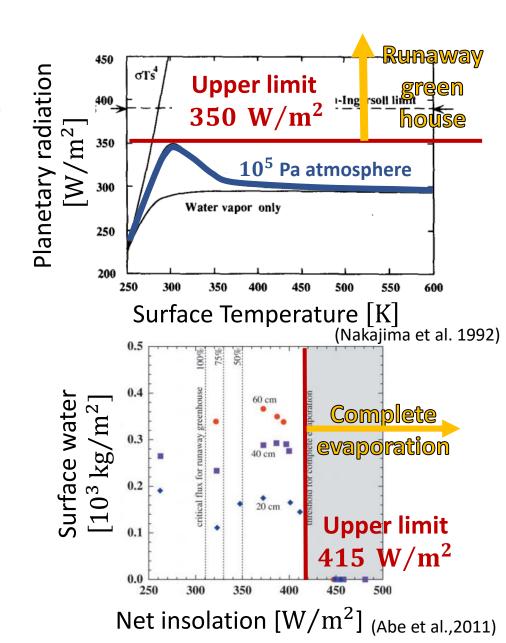
- Exoplanet is a planet out of the solar system
 - Rocky planet may exist (http://exoplanets.org/)
- In Habitable zone (HZ), liquid water exist on the planetary surface (Kasting et al. 1993)
 - Liquid water is necessary for life like that of the Earth
 - Inner edge is focused



Previous study

Inner edge of HZ

- An aqua planet (Nakajima et al., 1992; Ishiwatari et al., 2002)
 - na et al., 1992; iri et al., 2002)
 - Covered with ocean
 - Appearance of runaway green house state
 - Runaway green house state: Planetary radiation < Insolation
- A land planet (Abe et al., 2011)
 - Covered with soil and has small amount of water
 - Appearance of complete evaporation state
 - Complete evaporation state:
 All surface water evaporate



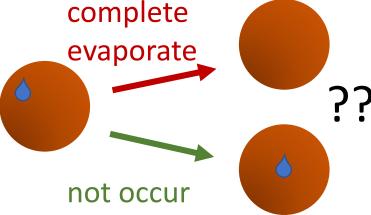
<u>Purpose</u>

- The first purpose was the investigation of dependence on obliquity and planetary rotation rate for the appearance of complete evaporation in a land planet
 - In a preliminary experiment, the complete evaporation did not occur

The purpose of this study

The examination of appearance of complete evaporation in a land planet

• re-experiment of Abe et al. (2011)



Methods

Model and Equations

Atmospheric general circulation model DCPAM (Dennou-Club Planetary Atmospheric Model)

Zonal wind
$$\frac{du}{dt} = \frac{uv\tan\varphi}{a} + fv - \frac{1}{\rho a\cos\varphi} \frac{\partial p}{\partial \lambda} + F_{\lambda}$$
 • Hydrostatic equilibrium

Hydrostatic equilibrium
$$0 = -\frac{1}{\rho} \frac{\partial p}{\partial z} - g$$

$$\text{Mass} \quad \frac{d\rho}{dt} = -\rho \left\{ \frac{1}{a\cos\varphi} \frac{\partial u}{\partial\lambda} + \frac{1}{a\cos\varphi} \frac{\partial}{\partial\varphi} (v\cos\phi) + \frac{\partial w}{\partial z} \right\} \bullet \quad \text{Vertical coordinate } \sigma = p/p_{\mathcal{S}}$$

Energy
$$\frac{dT}{dt} = \frac{1}{C_p^d \rho} \frac{dp}{dt} + \frac{Q^*}{C_p^d}$$

Water mass
$$\frac{dq}{dt} = S_q$$

Equation of state
$$p = \rho R^d T_v$$

- Primitive equations

 T_n : Virtual temperature

- The horizontal coordinate system is latitude and longitude

u:Zonal wind, v: Meridional wind, ρ : Density, T: Temperature, q: Specific humidity, a: Planetary radius, F: Forcing terms of motion, *f* : Coriolis parameter, *p* : Pressure, C_p^d : Constant pressure specific heat of dry air, Q^* : Source term of energy, S_a : Source term of vapor, R^d : Gas constant of dry air,

Methods

Numerical schemes for physical processes

- Dynamical process
 - Pseudo-spectral method is used
 - Semi-Lagrange method is used for advection of vapor
- Radiation process
 - Radiation scheme for the Earth is used (Chou and Lee, 1996; Chou et al., 1998; Chou et al., 2001)
- Vertical eddy mixing process
 - Mellor and Yamada level 2.5 scheme (Mellor and Yamada, 1982)
 - Vertical mixing of momentum, energy and water is considered

- Cumulus Convection process
 - Relaxed Arakawa-Schubert with ice is used (Arakawa and Schubert, 1974; Moorthi and Suarez, 1992)
 - Cumulus convectional advection of energy and water is considered
 - Large scale condensation process
 - Non-convective condensation with ice (Li Treut and Li, 1991)
 - Condensation of vapor in supersaturated air is considered
- Planetary surface setting
 - Thermal diffusion equation is solved
 - Bucket model is used (Manabe, 1969)

Methods

Experimental settings

Experiment Name	Surface condition	Solar const. $\left[W/m^2\right]$	Initial state
L_S1365_IniWet	bucket	1365	Isothermal, Wind static
L_S2400_IniWet	bucket	2400	A_S1365_IniWet
L_S2400_IniRun	bucket	2400	A_S2000_IniWet
L_S3600_IniWet	bucket	3600	L_S2400_IniWet
A_S1365_IniWet	swamp	1365	Isothermal, Wind static
A_S2000_IniWet	swamp	2000	Isothermal, Wind static

bucket: A land planet assumed; total water is unchanged

swamp: An aqua planet assumed; water is supplied continuously

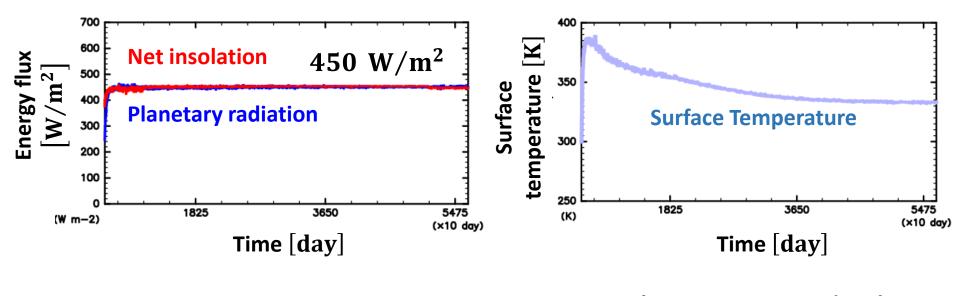
- Resolution
 - T21L26

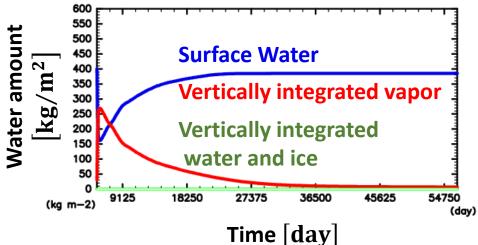
- Other parameter
 - Eccentricity, obliquity are 0
 - $\hbox{\bf Atmospheric components are} \\ \hbox{\bf same the Earth not including } O_3 \\ \\$

Results

A statistical equilibrium state is obtained

Exp. L_S2400_IniWet





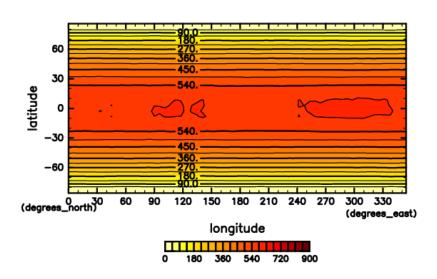
- Net insolation exceeds that of Abe et al. (2011)
- Complete evaporation does not occur

Results

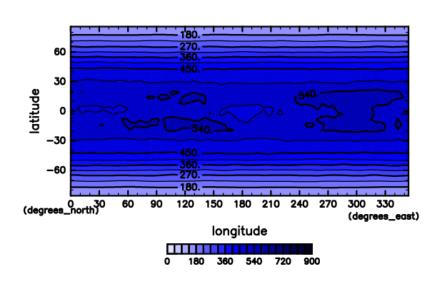
The structure of statistical equilibrium state Exp. L S2400 IniWet

- A statistical equilibrium state is maintained
 - Equatorial Planetary radiation is $520 \, \text{W/m}^2$
 - Much larger than $350 \, \text{W/m}^2$ (upper limit for aqua planet)
 - Because of dry atmosphere in the equatorial region
 - Can be balanced with large solar constant (Abe et al., 2011)

Net insolation (annual mean)



Planetary radiation (annual mean)

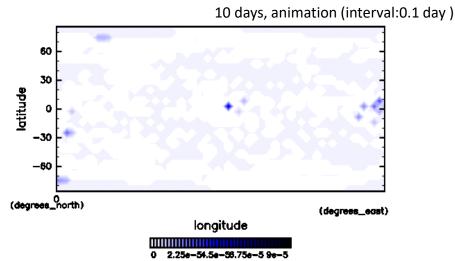


Results

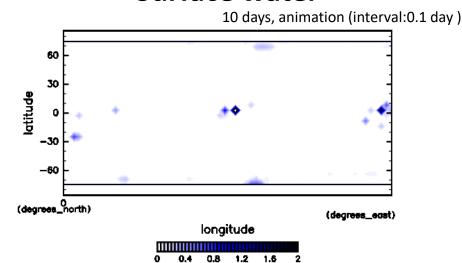
Precipitation and Evaporation

L_S2400_IniWet

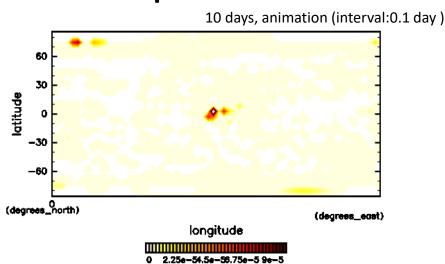




Surface water



Evaporation flux



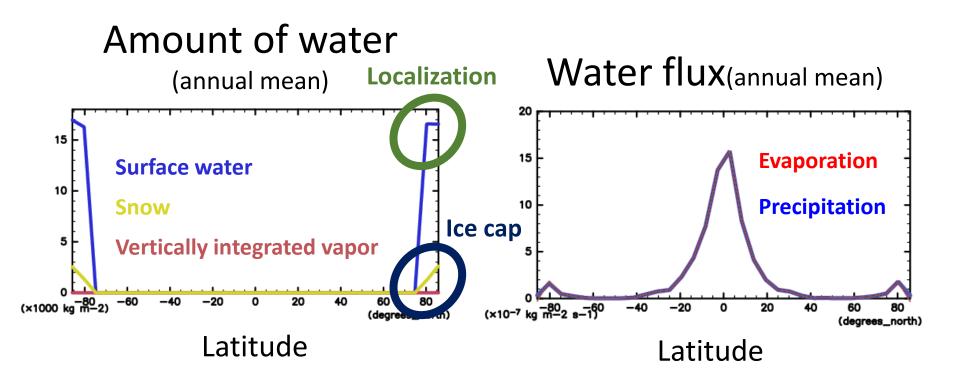
- Precipitation occurs in the region where evaporation occurs
 - Equatorial region
 - Polar region

<u>Results</u>

Water localization

L S2400 IniWet

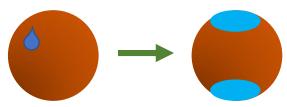
- Surface water is localized in polar region
- Polar ice cap exists
- Because temperature of polar region is low, surface water does not evaporate



Conclusions

- Complete evaporation did not occur in the experiment of this study
 - HZ for a land planet can be wider than that for previous study
 - Temperature in polar region did not increase
 - Insolation to the polar region is small
 - Atmospheric meridional heat transport is small?

snow and water



- Further experiments
 - Experiments with different settings (increasing polar region temperature)
 - Larger solar constant
 - Various obliquity and planetary rotation rate
 - Comparative experiment using other GCMs