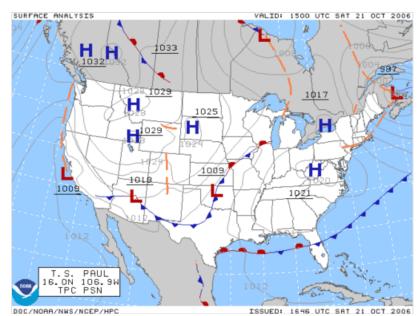


# **Surface weather analysis**

**Surface weather analysis** is a special type of <u>weather map</u> that provides a view of <u>weather</u> elements over a geographical area at a specified time based on information from ground-based weather stations. [1]

Weather maps are created by plotting or tracing the values of relevant quantities such as sea level pressure, temperature, and cloud cover onto a geographical map to help find synoptic scale features such as weather fronts.

The first weather maps in the 19th century were drawn well after the fact to help devise a theory on storm systems. [2] After the advent of the telegraph, simultaneous surface



A surface weather analysis for the United States on October 21, 2006. By that time, <u>Tropical Storm Paul</u> was active (Paul later became a hurricane).

<u>weather observations</u> became possible for the first time, and beginning in the late 1840s, the <u>Smithsonian Institution</u> became the first organization to draw real-time surface analyses. Use of surface analyses began first in the United States, spreading worldwide during the 1870s. Use of the <u>Norwegian cyclone model</u> for frontal analysis began in the late 1910s across Europe, with its use finally spreading to the United States during World War II.

Surface weather analyses have special symbols that show frontal systems, cloud cover, <u>precipitation</u>, or other important information. For example, an *H* may represent <u>high pressure</u>, implying clear skies and relatively warm weather. An *L*, on the other hand, may represent <u>low pressure</u>, which frequently accompanies precipitation. Various symbols are used not just for frontal zones and other surface boundaries on weather maps, but also to depict the present weather at various locations on the weather map. Areas of precipitation help determine the frontal type and location.

# **History of surface analysis**

The use of weather charts in a modern sense began in the middle portion of the 19th century in order to devise a theory on storm systems. [3] The development of a <u>telegraph</u> network by 1845 made it possible to gather weather information from multiple distant locations quickly enough to preserve its value for real-time applications. The Smithsonian Institution developed its network of observers over much of the

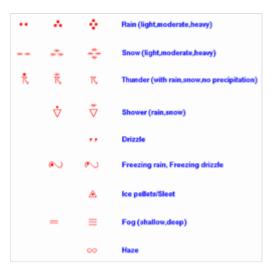
central and eastern United States between the 1840s and 1860s. The <u>U.S. Army Signal Corps</u> inherited this network between 1870 and 1874 by an act of Congress, and expanded it to the west coast soon afterwards. [4]

The weather data was at first less useful as a result of the different times at which weather observations were made. The first attempts at time standardization took hold in Great Britain by 1855. The entire United States did not finally come under the influence of time zones until 1905, when <u>Detroit</u> finally established standard time. Other countries followed the lead of the United States in taking simultaneous weather observations, starting in 1873. Other countries then began preparing surface analyses. The use of frontal zones on weather maps did not appear until the introduction of the <u>Norwegian cyclone model</u> in the late 1910s, despite Loomis' earlier attempt at a similar notion in 1841. Since the leading edge of air mass changes bore resemblance to the <u>military fronts</u> of <u>World War I</u>, the term "front" came into use to represent these lines.



Surface analysis of <u>Great</u>
<u>Blizzard of 1888</u> on March 12
at 10 pm

Despite the introduction of the Norwegian cyclone model just after World War I, the United States did not formally analyze fronts on surface analyses until late 1942, when the WBAN Analysis Center opened in downtown Washington, D.C. $^{[9]}$ The effort to automate map plotting began in the United States in 1969, [10] with the process complete in the 1970s. Hong Kong completed their process of automated surface plotting by 1987. [11] By 1999, computer systems and software had finally become sophisticated enough to allow for the ability to underlay on the same workstation satellite imagery, radar imagery, and model-derived fields such as atmospheric thickness and frontogenesis in combination with surface observations to make for the best possible surface analysis. In the United States, this development was achieved when workstations were replaced by n-AWIPS Intergraph workstations.  $\frac{[12]}{}$  By 2001, the various surface analyses done

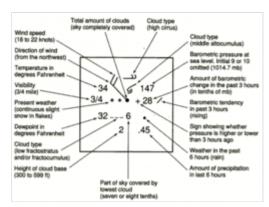


Present weather symbols used on weather maps

within the National Weather Service were combined into the Unified Surface Analysis, which is issued every six hours and combines the analyses of four different centers. [13] Recent advances in both the fields of meteorology and geographic information systems have made it possible to devise finely tailored weather maps. Weather information can quickly be matched to relevant geographical detail. For instance, icing conditions can be mapped onto the road network. This will likely continue to lead to changes in the way surface analyses are created and displayed over the next several years. [14]

# Station model used on weather maps

When analyzing a weather map, a station model is plotted at each point of observation. Within the station model, the temperature, dewpoint, wind speed and direction, atmospheric pressure, pressure tendency, and ongoing weather are plotted. The circle in the middle represents cloud cover; fraction it is filled in represents the degree of overcast. Outside the United States, temperature and dewpoint are plotted in degrees Celsius. The wind barb points in the direction from which the wind is coming. Each full flag on the wind barb represents 10 knots (19 km/h) of wind, each half flag represents 5 knots (9 km/h). When winds reach 50 knots (93 km/h), a filled in triangle is used for each 50 knots (93 km/h) of wind. In the United States, rainfall plotted in



Station model plotted on surface weather analyses

the corner of the station model are in <u>inches</u>. The international standard rainfall measurement unit is the <u>millimeter</u>. Once a map has a field of station models plotted, the analyzing <u>isobars</u> (lines of equal pressure), isallobars (lines of equal pressure change), isotherms (lines of equal temperature), and <u>isotachs</u> (lines of equal wind speed) are drawn. The abstract weather symbols were devised to take up the least room possible on weather maps.

# Synoptic scale features

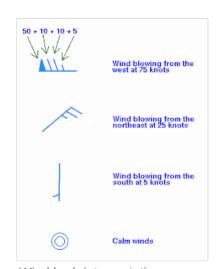
A synoptic scale feature is one whose dimensions are large in scale, more than several hundred kilometers in length. [19] Migratory pressure systems and frontal zones exist on this scale.

#### **Pressure centers**

Centers of surface high- and low-pressure areas that are found within closed isobars on a surface weather analysis are the absolute maxima and minima in the pressure field, and can tell a user in a glance what the general weather is in their vicinity. Weather maps in English-speaking countries will depict their highs as Hs and lows as Ls, [20] while Spanish-speaking countries will depict their highs as As and lows as Bs. [21]

#### Low pressure

Low-pressure systems, also known as <u>cyclones</u>, are located in minima in the pressure field. Rotation is inward at the surface and counterclockwise in the <u>northern hemisphere</u> as opposed to inward and clockwise in the <u>southern hemisphere</u> due to the <u>Coriolis force</u>. Weather is normally unsettled in the vicinity of a cyclone, with



Wind barb interpretation

increased cloudiness, increased winds, increased temperatures, and upward motion in the atmosphere, which leads to an increased chance of precipitation. Polar lows can form over relatively mild ocean

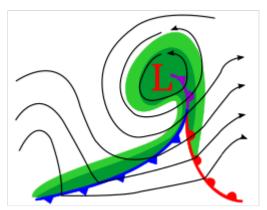
waters when cold air sweeps in from the ice cap. The relatively warmer water leads to upward convection, causing a low to form, and precipitation usually in the form of snow. Tropical cyclones and winter storms are intense varieties of low pressure. Over land, thermal lows are indicative of hot weather during the summer. [22]

#### High pressure

High-pressure systems, also known as <u>anticyclones</u>, rotate outward at the surface and clockwise in the northern hemisphere as opposed to outward and counterclockwise in the southern hemisphere. Under surface highs, sinking of the atmosphere slightly warms the air by compression, leading to clearer skies, winds that are lighter, and a reduced chance of precipitation. [23] The descending air is dry, hence less energy is required to raise its temperature. If high pressure persists, air pollution will build up due to pollutants trapped near the surface caused by the subsiding motion associated with the high. [24]

#### **Fronts**

Fronts in meteorology are boundaries between <u>air masses</u> that have different density, air temperature, and <u>humidity</u>. Strictly speaking, the front is marked at the warmer edge of a *frontal zone* where the <u>gradient</u> is very large. When a front passes over a point, it is marked by changes in temperature, moisture, wind speed and direction, a minimum of atmospheric pressure, and a change in the cloud pattern, sometimes with precipitation. <u>Cold fronts</u> develop where the cold air mass is advancing, <u>warm fronts</u> where the warm air is advancing, and a <u>stationary front</u> is not moving. Fronts classically wrap around low pressure centers as indicated in the <u>image</u> here depicted for the northern hemisphere. On a larger scale, the Earth's <u>polar front</u> is a sharpening of the general equator-to-pole temperature gradient, underlying a



Occluded cyclone example. The triple point is the intersection of the cold, warm, and occluded fronts.

high-altitude jet stream for reasons of thermal wind balance. Fronts usually travel from west to east, although they can move in a north-south direction or even east to west (a "backdoor" front) as airflow wraps around a low pressure center. Frontal zones can be distorted by such geographic features as mountains and large bodies of water. [13]

#### **Cold front**

A cold front is located at the leading edge of a sharp temperature gradient on an <u>isotherm</u> analysis, often marked by a sharp surface pressure <u>trough</u>. Cold fronts can move up to twice as quickly as warm fronts and produce sharper changes in <u>weather</u> since cold air is denser than warm air and rapidly lifts as well as pushes the warmer air. Cold fronts are typically accompanied by a narrow band of clouds, showers and thunderstorms. On a weather map, the surface position of the cold front is marked with a blue line of triangles (pips) pointing in the direction of travel, at the leading edge of the cooler air mass. [13]

#### Warm front

Warm fronts mark the position on the Earth's surface where a relatively warm body of air is advancing into colder air. The front is marked on the warm edge of the gradient in isotherms, and lies within a low pressure trough that tends to be broader and weaker than that of a cold front. Warm fronts move more slowly than cold fronts because cold air is denser, and is only pushed along (not lifted from) the Earth's surface. The warm air mass overrides the cold air mass, so temperature and cloud changes occur at higher altitudes before those at the surface. Clouds ahead of the warm front are mostly stratiform with precipitation that increases gradually as the front approaches. Ahead of a warm front, descending cloud bases will often begin with cirrus and cirrostratus (high-level), then altostratus (mid-level) clouds, and eventually lower in the atmosphere as the front passes through. Fog can precede a warm front when precipitation falls into areas of colder air, but increasing surface temperatures and wind tend to dissipate it after a warm front passes through. Cases with environmental instability can be conducive to thunderstorm development. On weather maps, the surface location of a warm front is marked with a red line of half circles pointing in the direction of travel.

#### Occluded front

The classical view of an <u>occluded front</u> is that they are formed when a cold front overtakes a warm front. A more modern view suggests that they form directly during the wrap-up of the <u>baroclinic zone</u> during <u>cyclogenesis</u>, and lengthen due to flow <u>deformation</u> and rotation around the cyclone.

Occluded fronts are indicated on a weather map by a purple line with alternating half-circles and triangles pointing in direction of travel: that is, with a mixture of warm and cold frontal colors and symbols. Occlusions can be divided into

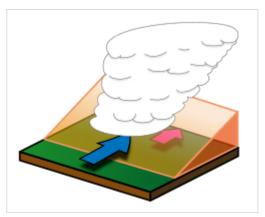


Illustration clouds overriding a warm front

warm vs. cold types.<sup>[27]</sup> In a cold occlusion, the air mass overtaking the warm front is cooler than the cool air ahead of the warm front, and plows under both air masses. In a warm occlusion, the air mass overtaking the warm front is not as cool as the cold air ahead of the warm front, and rides over the colder air mass while lifting the warm air. Occluded fronts are indicated on a weather map by a purple line with alternating half-circles and triangles pointing in direction of travel.<sup>[13]</sup>

Occluded fronts usually form around low pressure systems in the mature or late stages of their life cycle, but some continue to deepen after occlusion, and some do not form occluded fronts at all. The weather associated with an occluded front includes a variety of cloud and precipitation patterns, including dry slots and banded precipitation. Cold, warm and occluded fronts often meet at the point of occlusion or triple point. [28]

#### Stationary fronts and shearlines

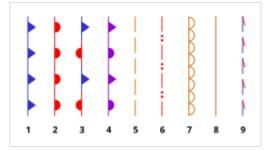
A stationary front is a non-moving boundary between two different air masses. They tend to remain in the same area for long periods of time, sometimes undulating in waves. [29] Often a less-steep temperature gradient continues behind (on the cool side of) the sharp frontal zone with more widely spaced isotherms. A wide variety of weather can be found along a stationary front, characterized more by its prolonged presence than by a specific type. Stationary fronts may dissipate after several days, but can change into a

cold or warm front if conditions aloft change, driving one air mass toward the other. Stationary fronts are marked on weather maps with alternating red half-circles and blue spikes pointing in opposite directions, indicating no significant movement.

As airmass temperatures equalize, stationary fronts may become smaller in scale, degenerating to a narrow zone where wind direction changes over a short distance, known as a shear line, [30] depicted as a blue line of single alternating dots and dashes. [13][31]

# **Mesoscale features**

<u>Mesoscale</u> features are smaller than <u>synoptic scale</u> systems like fronts, but larger than <u>storm-scale</u> systems like thunderstorms. Horizontal dimensions generally range from over ten kilometres to several hundred kilometres. [32]



A guide to the symbols for weather fronts that may be found on a weather map:

- 1. cold front
- 2. warm front
- 3. stationary front
- 4. occluded front
- 5. surface trough
- 6. squall line
- 7. dry line
- 8. tropical wave
- 9. Trowal

### **Dry line**

The <u>dry line</u> is the boundary between dry and moist air masses east of mountain ranges with similar orientation to the <u>Rockies</u>, depicted at the leading edge of the <u>dew point</u>, or moisture, gradient. Near the surface, warm moist air that is denser than warmer, dryer air wedges under the drier air in a manner similar to that of a cold front wedging under warmer air. [33] When the warm moist air wedged under the drier mass heats up, it becomes less dense and rises and sometimes forms thunderstorms. [34] At higher altitudes, the warm moist air is less dense than the cooler, drier air and the boundary slope reverses. In the vicinity of the reversal aloft, severe weather is possible, especially when a triple point is formed with a cold front.

During daylight hours, drier air from aloft drifts down to the surface, causing an apparent movement of the dryline eastward. At night, the boundary reverts to the west as there is no longer any solar heating to help mix the lower atmosphere. If enough moisture converges upon the dryline, it can be the focus of afternoon and evening thunderstorms. A dry line is depicted on United States surface analyses as a brown line with scallops, or bumps, facing into the moist sector. Dry lines are one of the few surface fronts where the special shapes along the drawn boundary do not necessarily reflect the boundary's direction of motion. [37]

# Outflow boundaries and squall lines

Organized areas of thunderstorm activity not only reinforce pre-existing frontal zones, but they can outrun cold fronts. This outrunning occurs in a pattern where the upper level jet splits into two streams. The resultant mesoscale convective system (MCS) forms at the point of the upper level split in the wind pattern at the area of the best low-level inflow. The convection then moves east and equatorward into the warm sector, parallel to low-level thickness lines. When the convection is strong and linear or curved, the MCS is called a squall line, with the feature placed at the leading edge where the significant wind shifts

and pressure rises. [38] Even weaker and less organized areas of thunderstorms will lead to locally cooler air and higher pressures, and <u>outflow boundaries</u> exist ahead of this type of activity, "SQLN" or "SQUALL LINE", while outflow boundaries are depicted as troughs with a label of "OUTFLOW BOUNDARY" or "OUTFLOW BNDRY".

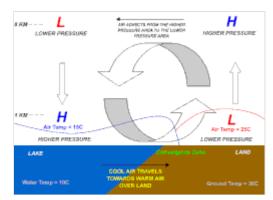
#### Sea and land breeze fronts

<u>Sea breeze</u> fronts occur on sunny days when the landmass warms the air above it to a temperature above the water temperature. Similar boundaries form downwind on lakes and rivers during the day, as well as offshore landmasses at night. Since the <u>specific heat</u> of water is so high, there is little diurnal temperature change in bodies of water, even on the sunniest days. The water temperature varies less than 1 °C (1.8 °F). By contrast, the land, with a lower specific heat, can vary several degrees in a matter of hours. [39]

During the afternoon, air pressure decreases over the land as the warmer air rises. The relatively cooler air over the sea rushes in to replace it. The result is a relatively cool onshore wind. This process usually reverses at night where the water



A <u>shelf cloud</u> such as this one can be a sign that a squall is imminent



Idealized circulation pattern associated with a sea breeze

temperature is higher relative to the landmass, leading to an offshore land breeze. However, if water temperatures are colder than the land at night, the sea breeze may continue, only somewhat abated. This is typically the case along the California coast, for example.

If enough moisture exists, thunderstorms can form along sea breeze fronts that then can send out outflow boundaries. This causes chaotic wind/pressure regimes if the steering flow is weak. Like all other surface features, sea breeze fronts lie inside troughs of low pressure.

# Microscale features

# **Descending reflectivity core**

A descending reflectivity core (DRC) is a <u>meteorological</u> phenomenon observed in <u>supercell</u> thunderstorms, characterized by a localized, small-scale area of enhanced <u>radar reflectivity</u> that <u>descends</u> from the echo overhang into the lower levels of the storm.

## See also

- Bowditch's American Practical Navigator
- Extratropical cyclone
- Frontolysis

- Outline of meteorology
- Ridge (meteorology)

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# **External links**

- "The Mid-Latitude Cyclone" (http://www.physicalgeography.net/fundamentals/7s.html)
- Norwegian Cyclone Model NWS (https://web.archive.org/web/20160104214105/http://www.srh.weather.gov/srh/jetstream/synoptic/cyclone.htm)
- Unified Surface Analysis Manual NWS (http://www.wpc.ncep.noaa.gov/sfc/UASfcManual Version1.pdf)
- Unified Surface Analysis NWS (https://web.archive.org/web/20171016122221/http://www.opc.ncep.noaa.gov/unified\_analysis.php)
- Glossary of Meteorology (http://amsglossary.allenpress.com/glossary/)
- Cold Front Page (http://ww2010.atmos.uiuc.edu/(GI)/guides/mtr/af/frnts/cfrnt/def.rxml)

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