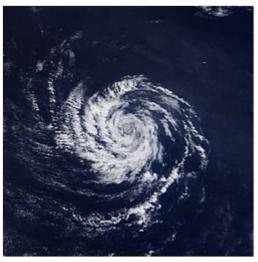


Mesoscale meteorology

Mesoscale meteorology is the study of <u>weather</u> systems and processes at horizontal scales of approximately 5 kilometres (3 mi) to several hundred kilometres. It is smaller than <u>synoptic-scale</u> systems (1,000 km or larger) but larger than <u>microscale</u> (less than 1 km). At the small end, it includes **storm-scale** phenomena (the size of an individual thunderstorm[1]). Examples of mesoscale weather systems are <u>sea breezes</u>, <u>squall lines</u>, and <u>mesoscale convective</u> complexes.

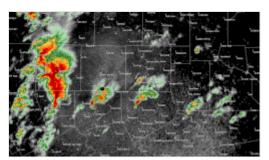
Vertical velocity often equals or exceeds horizontal velocities in mesoscale meteorological systems due to nonhydrostatic processes such as buoyant acceleration of a rising thermal or acceleration through a narrow mountain pass.



A meso-beta scale vortex

Classification

The earliest networks of weather observations in the late 1800s and early 1900s could detect the movement and evolution of larger, synoptic-scale systems like high and low-pressure areas. However, smaller and potentially hazardous meteorological phenomena were not well-captured by the sparse observation networks. The emergence of weather radar in the mid-1900s and an improved understanding of thunderstorm behavior led to a increased recognition of a need to study phenomena between the scales studied in the extant disciplines of microscale and synoptic-scale meteorology. The term "mesoscale" originated from M. G. H. Ligda at the Massachusetts Institute of Technology, who suggested a need to study phenomena at such scales in 1951: [2][3]



Mesoscale meteorology studies weather systems like thunderstorm clusters too small to be resolved by the earliest weather observation networks.

It is anticipated that radar will provide useful information concerning the structure and behavior of that portion of the atmosphere which is not covered by either micro- or synoptic-meteorological studies. We have already observed with radar that precipitation formulations which are undoubtedly of significance occur on a scale too gross to be observed from a single station, yet too small to appear even on sectional synoptic charts. Phenomena of this size might well be designated as mesometeorological.

—M. G. H. Ligda, "Radar Storm Observation", *Compendium of Meteorology* (1951)^[4]

Subclasses

Mesoscale meteorology broadly concerns meteorological phenomena larger than a few kilometres across but smaller than could be resolved by the observation networks used in the earliest <u>standardized weather maps</u>. The mesoscale regime is often divided into these subclasses based on the size of associated weather systems: 6

- **Meso-alpha** (meso-α) 200–2000 km scale of phenomena like <u>fronts</u>, <u>squall lines</u>, <u>mesoscale convective systems</u> (MCS), <u>tropical cyclones</u> at the smaller edge of <u>synoptic</u> scale. [6]
- **Meso-beta** (meso- β) 20–200 km scale of phenomena like <u>mesocyclones</u>, <u>sea breezes</u>, and lake effect snow storms. [6] Mesocale often refers to meso- β scale specifically. [7]
- **Meso-gamma** (meso-y) 2–20 km scale of phenomena like <u>thunderstorm convection</u>, complex terrain flows (at the larger edge of microscale)

As a note, tropical and subtropical cyclones are classified by <u>National Hurricane Center</u> as synoptic scale rather than mesoscale. [8]

Features the size of an individual thunderstorm are also known somewhat informally as "storm-scale", typically meso-gamma but sometimes meso-beta^[9] or microscale.

Dynamics

Mesoscale processes are characterized by having a relatively large Rossby number compared to synoptic scale processes. Thus, over shorter distances as implicated in mesoscale phenomena, the importance of geostrophic balance and the Earth's rotation in shaping atmospheric processes is small relative to synoptic-scale phenomena. This is particularly true towards the smaller end of the mesoscale range. Because the curvature of Earth is small at mesoscales, the physical models used to diagnose mesoscale phenomena often assume a constant Coriolis frequency. Nonetheless, the Coriolis force is non-negligible and comparable to the influence of atmospheric buoyancy.



Vertical motion is prominent in many mesoscale processes.

Large-scale <u>turbulence</u> and <u>eddies</u> also play a large role in mesoscale meteorology. The vertical movement of air (often expressed as <u>omega</u>) is larger at mesoscale than at synoptic scales, and the distribution of <u>air pressure</u> tends to be influenced by the behavior of winds at the mesoscale (as opposed to the converse at synoptic scales). For many mesoscale phenomena, the vertical acceleration of air is sufficiently large enough that calculations cannot assume <u>hydrostatic balance</u>. This is often true of phenomena with a vertical dimension roughly equal to their horizontal dimensions. [11]:9

Mesoscale boundaries

As in <u>synoptic frontal analysis</u>, mesoscale analysis uses cold, warm, and occluded fronts on the mesoscale to help describe phenomena. On weather maps mesoscale fronts are depicted as smaller and with twice as many bumps or spikes as the synoptic variety. In the <u>United States</u>, opposition to the use of the mesoscale versions of fronts on weather analyses, has led to the use of an overarching symbol (a trough symbol) with a label of outflow boundary as the frontal notation. [12]

See also



- Microscale meteorology
- Misoscale meteorology
- POLYGON experiment
- Surface weather analysis
- Synoptic scale meteorology

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