**Exploring the Timescales that Storms Do Geomorphic Work in Last Chance Canyon, New Mexico**

**ABSTRACT**

Here, we investigate how discrete storm events influence sediment flux, grain size distributions, and alluvial residence times in two first and second order streams in Last Chance canyon, New Mexico, USA. Many studies assume that sediment flux out of erosional landscapes is constant, however, in ephemeral desert stream channels this assumption is likely invalid. In this study, we identify the storm properties necessary to overcome grain motion thresholds and transport differently sized sediment. To accomplish this, we will use a updated OverlandFlow model (Adams [et al., 2001](#_ENREF_30)) , from the LandLab library, which routs water from storms across a landscape. We will couple OverlandFlow to RiverBedDynamics, which uses water velocity and discharge to determine if sediment of certain sizes can be mobilized. We used precipitation data (NOAA) to reconstruct realistic storm durations and intensities for Last Chance canyon. We used drone photos to construct high resolution DEMs and orthomosaics to determine grain size distributions in different channel sections. With these methods, we plan to determine the 1) hydrograph characteristics necessary to remove the differently sized sediment armor, 2) the residence times of the differently sized sediment within channel sections of varying steepness, and 3) if the larger sediment has armored channels at timescales necessary for landscape morphology to be reflective of their presence in the channel. We hypothesize that only very intense storms will ‘do geomorphic work’ by removing larger grains from the steep dolomitic channel section, while smaller storms will be able to carry off smaller sandstone sediment from the shallow channel section. We further hypothesize that the large sediment armors bedrock at timescales necessary for stream channel morphology to reflect their presence. This study will help to constrain the imprint storms have on landscape morphology and elucidate geomorphic processes between shorter time scales and timescales necessary for the landscape to adjust.

**INTRODUCTION**

The topographic signal from variance in bedrock properties can be blurred, removed, or exaggerated by the presence of sediment armor (Duval et al., 2004; Johnson et al., 2009; Finnegan et al., 2017). Rock properties, specifically bed thickness and discontinuity spacing, has been shown to influence sediment production and grain size (Spotila et al., 2015; Keen-Zeebert et al., 2017; Sklar et al., 2017). The presence of larger boulders, sourced from thicker units and bedrock with larger discontinuity spacing, causes stream channels to steepen (Thaler and Covington, 2016; DiBiase et al. 2018). But, to imprint their signal into topography, alluvium must effectively armor stream channels for time periods which stretch over many storm events. Because intense storms are the primary drivers of sediment flux out of stream channels (Schuerch et al, 2006), residence time of boulders in stream channels depends primarily on the storm hydrograph and on relevant sediment properties like size and competency. How large must sediment be to armor channels, withstand mobilization by storms, and remain in streams at timescales necessary for the landscape to reflect their presence? How intense must storms be to do remove large boulders and incise into bedrock? Without an answer to these questions, attempts to link geomorphic processes across temporal scales will be confounded.

In this study, we explore the relationship between the storm hydrograph and the size of sediment armoring the channel on landscape morphology. More specifically, we ask: (1) how immobile is larger sediment than smaller sediment in last chance canyon? (2) is it possible for storm events to remove these large boulders? And (3) What is the relationship between channels steepness and the residence times of large alluvial armor? To address these questions, we will couple a landscape evolution model which calculates movement of sediment based on shear stress with the overland flow component developed by Jordan Adams and inform our modeled experiments from storm and sediment size data from Last Chance canyon, New Mexico. This study will rectify assumptions about the effect of differently sized sediment on topography from the storm hydrograph to time scales relevant for the landscape to reflect geomorphic processes. Furthermore, I seek to demonstrate that baselevel is effectively pinned at the transition from large to smaller sediment and knickpoint celerity is slowed in stream channels with larger sediment armor.

**Hypothesis**

Intense storms are required to overcome the threshold shear stress necessary to move large sediment. Storm duration is an important hydrograph characteristic in removing smaller sediment out of the system but will not influence sediment flux of larger sediment. Because of this, storms will more frequently incise into bedrock which is armored by smaller grains than channels with larger alluvial armor and channels will shallow.

The larger sediment grains, which are sourced from thicker bedrock, are mostly immobile on timescales necessary for landscape morphology to reflect their presence. Because these grains are so immobile, and because they armor the channel and prevent bedrock erosion, knickpoint celerity is slowed and baselevel is pinned at points above where they exist.

**FIELD AREA**

This study focuses on 2 different, first to second order channels with intermittent flow in Last Chance canyon (Figure 1). During Permian time, a shallow lagoon existed behind a reef complex to the south and deposited what would become interbedded carbonate and siliciclastic bedrock of various thicknesses (Hill, 2000; Phelps et al., 2008; Kerans et al., 2017). The Guadalupe mountains were uplifted during basin and range extension beginning 27 million years ago, exposing the previously buried bedrock (Chapin and Cather, 1994; Ricketts et al.., 2014, Hoffman, 2014; Decker et al., 2018).

Because of their differing morphology and grain size distributions, we use data gathered from different sections of two channels, called LC1 and LC3 (grain size figure, chi, and steepness vs grain size), to identify the potential control differently sized alluvial armor has on stream channel morphology. In both channels, grain size distributions increase with channel steepness. This simple variation in grain size distribution makes LC1 and 3 ideal locations to explore the effect of varying sediment sizes on stream channel morphology.

**METHODS**

We used Landlab, a Python-language, open-source, flexible library of different components that can be easily coupled together (Adams et al., 2014; Tucker et al., 2016; Hobley et al., 2017) to model sediment transport during storm events. The Landlab OverlandFlow component (Adams citation), which is based on a simplified calculation of the shallow water equations (de Almeida et al., 2012), routs rainfall during storm events across a DEM and determines flow depth, velocity, and stream discharge at each cell. We used an updated OverlandFlow component which allowed us to include a spatially variable roughness attribute which we altered for each grain size distribution.

We coupled OverlandFlow to RiverBedDynamics (Monsalve et al., 2023), which allowed us to calculate bed load transport and elevation change due to erosion or deposition. RiverBedDynamics takes surface flow variables from OverlandFlow and grain sizes, both of which are stored in a raster grid. Sediment fluxes for each cell are then calculated, and bed surface elevation and bed properties like bed grain size distributions are updated.

We used depth-duration plots (climate figure) for Last Chance canyon from NOAA (citation). These data are used to inform the OverLandFlow component and generate storms which mimic real storms from our study area.

We used a DJI mavik 2 pro to take photos of three segments of each of the two channels- a downstream, intermediate, and upstream- from elevations of approximately 20 meters above LC1 and LC3. We then used Agisoft Metashape software to process these images and to produce orthomosaics and DEM’s with a spatial resolution that varied from # to #. We used the PebbleCounts image analysis package (Purinton and Bookhagen, 2019), to determine the b axis diameter of alluvium.

**RESULTS**

**Analysis**

To link the hydrograph and sediment size, we examine sediment flux in different channel sections for different storms. Storms which strip sediment armor of different sizes will have difference recurrence intervals. We expect that sediment residences times will vary with storm intensity and storms with higher recurrence intervals will be required to move larger alluvium. Storms will be able to remove sediment armor and do geomorphic work less frequently and for higher recurrence intervals in channel sections with larger sediment armor. However, the more frequent, less intense storms should be able to remove sediment and erode into the shallower channel sections which are armored by smaller sediments more often. We expect our modeled storms to remove smaller sediments, leaving larger alluvium behind for storms under a certain recurrence interval. We will compare modeled sediment size distributions with distributions we measured in Last Chance canyon by photosieving orthophotos generated with drone surveys. This methodology will help elucidate how variance in storm characteristics affects sediment size distributions and alluvial residence times.

To quantify the influence of storms on the time for landscape morphology to respond to the changes in sediment size, we measure the size of sediment in the channels and the recurrence interval required to remove sediments of differing sizes. We will determine the residence time of sediment necessary for the channel morphology to steepen to reflect the presence of alluvial armor. We seek to quantify a link between channel steepness and the residence time of large sediment armor. Variance in sediment size affects channel and landscape morphology (figure 4), and we will relate residence time in channel with degree of steepening. To validate our modeled results, we plan on comparing the landscape relief of bedrock armored with sediment of varying sizes (figure 5). We expect that sediment sourced from dolomite is large and competent enough (figure 6) to pin baselevel at channel sections above it. Furthermore, we will quantify the degree to which celerity in the steep knickzone has been slowed by the presence of large dolomitic armor.