Earth

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Earth

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"The Blue Marble" photograph of Earth,
taken from Apollo 17
Designations
Alternative names Terra, Gaia
Orbital characteristics
Epoch J2000.0[note 1]
Aphelion 152,098,232 km
1.01671388 AU[note 2]
Perihelion 147,098,290 km
0.98329134 AU[note 2]
Semi-major axis 149,598,261 km
1.00000261 AU[1]
Eccentricity 0.01671123[1]
Orbital period 365.256363004 days[2]
1.000017421 yr
Average orbital speed 29.78 km/s[3]
107,200 km/h
Mean anomaly 357.51716°[3]
Inclination 7.155° to Sun's equator
1.57869°[4] to invariable plane
Longitude of ascending node 348.73936°[3][note 3]
Argument of perihelion 114.20783°[3][note 4]
Satellites 1 natural (the Moon),
8,300+ artificial (as of 1 March 2001)[5]
Physical characteristics
Mean radius 6,371.0 km[6]
Equatorial radius 6,378.1 km[7][8]
Polar radius 6,356.8 km[9]
Flattening 0.0033528[10]
Circumference 40,075.017 km (equatorial)[8]
40,007.86 km (meridional)[11][12]
Surface area
510,072,000 km2[13][14][note 5] 148,940,000 km2 lan
d (29.2 %)
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361,132,000 km2 water (70.8 %)
Volume 1.08321×1012 km3[3]
Mass
5.9736×1024 kg[3]
3.0×10-6 Suns
Mean density 5.515 g/cm3[3]
Equatorial surface gravity 9.780327 m/s2[15]
0.99732 q
Escape velocity 11.186 km/s[3]
Sidereal rotation period 0.99726968 d[16]
23h 56m 4.100s
Equatorial rotation velocity 1,674.4 km/h (465.1
 m/s)[17]
Axial tilt 23°26'21".4119[2]
Albedo
0.367 (geometric)[3]
0.306 (Bond)[3]
Surface temp. min mean max
Kelvin 184 K[18] 288 K[19] 330 K[20]
Celsius -89.2 °C 15 °C 56.7 °C
Atmosphere
Surface pressure 101.325 kPa (MSL)
Composition 78.08% nitrogen (N2)[3] (dry air)
20.95% oxygen (O2)
0.93% argon
0.039% carbon dioxide[21]
About 1% water vapor (varies with climate)
Earth is the third planet from the Sun, and the de
nsest and fifth-largest of the eight planets in the
e Solar System. It is also the largest of the Sola
r System's four terrestrial planets. It is sometim
es referred to as the world, the Blue Planet, [22]
or by its Latin name, Terra.[note 6]
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Earth formed approximately 4.54 billion years ago, and life appeared on its surface within one billi on years.[23] Earth's biosphere then significantly altered the atmospheric and other basic physical conditions, which enabled the proliferation of organisms as well as the formation of the ozone layer

, which together with Earth's magnetic field block ed harmful solar radiation, and permitted formerly ocean-confined life to move safely to land.[24] The physical properties of the Earth, as well as it sepological history and orbit, have allowed life to persist. Estimates on how much longer the plane twill be able to continue to support life range from 500 million years (myr), to as long as 2.3 bil lion years (byr).[25][26][27]

Earth's lithosphere is divided into several rigid segments, or tectonic plates, that migrate across the surface over periods of many millions of years. About 71% of the surface is covered by salt wate roceans, with the remainder consisting of contine nts and islands which together have many lakes and other sources of water that contribute to the hyd rosphere. Earth's poles are mostly covered with ice that is the solid ice of the Antarctic ice sheet and the sea ice that is the polar ice packs. The planet's interior remains active, with a solid iron inner core, a liquid outer core that generates the magnetic field, and a thick layer of relatively solid mantle.

Earth gravitationally interacts with other objects in space, especially the Sun and the Moon. During one orbit around the sun, the Earth rotates about its own axis 366.26 times, creating 365.26 solar days, or one sidereal year.[note 7] The Earth's ax is of rotation is tilted 23.4° away from the perpendicular of its orbital plane, producing seasonal variations on the planet's surface with a period of one tropical year (365.24 solar days).[28] The Moon is Earth's only natural satellite. It began or biting the Earth about 4.53 billion years ago (bya). The Moon's gravitational interaction with Earth stimulates ocean tides, stabilizes the axial tilt, and gradually slows the planet's rotation.

The planet is home to millions of species, including humans.[29] Both the mineral resources of the p

lanet and the products of the biosphere contribute resources that are used to support a global human population.[30] These inhabitants are grouped int o about 200 independent sovereign states, which in teract through diplomacy, travel, trade, and milit ary action. Human cultures have developed many vie ws of the planet, including its personification as a planetary deity, its shape as flat, its position as the center of the universe, and in the modern Gaia Principle, as a single, self-regulating organism in its own right.

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Name and etymology

The modern English noun earth developed from Middle English erthe (recorded in 1137), itself from Old English eorthe (dating from before 725), deriving from Proto-Germanic *ertho. Earth has cognates in all other Germanic languages, including Dutch aarde, German Erde, and Swedish, Norwegian, and Danish jord.[31] The Earth is personified as a goddess in Germanic paganism (appearing as Jörð in Norse mythology, mother of the god Thor).[32]

In general English usage, the name earth can be ca pitalized or spelled in lowercase interchangeably, either when used absolutely or prefixed with "the " (i.e. "Earth", "the Earth", "earth", or "the ear th"). Many deliberately spell the name of the plan et with a capital, both as "Earth" or "the Earth". This is to distinguish it as a proper noun, disti nct from the senses of the term as a count noun or verb (e.g. referring to soil, the ground, earthin g in the electrical sense, etc.). Oxford spelling recognizes the lowercase form as the most common, with the capitalized form as a variant of it. Anot her common convention is to spell the name with a capital when occurring absolutely (e.g. Earth's at mosphere) and lowercase when preceded by "the" (e. g. the atmosphere of the earth). The term almost e xclusively exists in lowercase when appearing in c ommon phrases, even without "the" preceding it (e. q. "It does not cost the earth.", "What on earth a re you doing?").[33]

Chronology

Earth compared with the other planets (second to bo ttom row, left)

Formation

Main article: History of the Earth

The earliest material found in the Solar System is dated to 4.5672±0.0006 bya;[34] therefore, it is inferred that the Earth must have been formed by a ccretion around this time. By 4.54±0.04 bya[23] th e primordial Earth had formed. The formation and e volution of the Solar System bodies occurred in ta ndem with the Sun. In theory a solar nebula partit ions a volume out of a molecular cloud by gravitat ional collapse, which begins to spin and flatten i nto a circumstellar disk, and then the planets gro w out of that in tandem with the star. A nebula co ntains gas, ice grains and dust (including primord ial nuclides). In nebular theory planetesimals com mence forming as particulate accrues by cohesive c lumping and then by gravity. The assembly of the p rimordial Earth proceeded for 10-20 myr.[35] The M oon formed shortly thereafter, about 4.53 bya.[36]

The Moon's formation remains a mystery. The workin g hypothesis is that it formed by accretion from m aterial loosed from the Earth after a Mars-sized o bject, dubbed Theia, had a giant impact with Earth, [37] but the model is not self-consistent. In this scenario the mass of Theia is 10% of the Earth's mass, [38] it impacts with the Earth in a glancing blow, [39] and some of its mass merges with the Earth. Between approximately 3.8 and 4.1 bya, numero us asteroid impacts during the Late Heavy Bombardm ent caused significant changes to the greater surf ace environment of the Moon, and by inference, to the Earth.

Earth's atmosphere and oceans formed by volcanic a ctivity and outgassing that included water vapor. The origin of the world's oceans was condensation augmented by water and ice delivered by asteroids, proto-planets, and comets.[40] In this model, atm ospheric "greenhouse gases" kept the oceans from f

reezing while the newly forming Sun was only at 70 % luminosity.[41] By 3.5 bya, the Earth's magnetic field was established, which helped prevent the a tmosphere from being stripped away by the solar wind.[42]

A crust formed when the molten outer layer of the planet Earth cooled to form a solid as the accumul ated water vapor began to act in the atmosphere. T he two models[43] that explain land mass propose e ither a steady growth to the present-day forms[44] or, more likely, a rapid growth[45] early in Eart h history[46] followed by a long-term steady conti nental area.[47][48][49] Continents formed by plat e tectonics, a process ultimately driven by the co ntinuous loss of heat from the earth's interior. O n time scales lasting hundreds of millions of year s, the supercontinents have formed and broken up t hree times. Roughly 750 mya (million years ago), o ne of the earliest known supercontinents, Rodinia, began to break apart. The continents later recomb ined to form Pannotia, 600-540 mya, then finally P angaea, which also broke apart 180 mya.[50]

Evolution of life
Main article: Evolutionary history of life
For more details on the current eon, see Geological
history of Earth.

Stratocumulus clouds over the Pacific, viewed from orbit. Over 70% percent of Earth's surface is covered with water, which contains about half of the planet's species.[51]

Highly energetic chemistry is thought to have produced a self-replicating molecule around 4 by a and half a billion years later the last common ancestor of all life existed. [52] The development of phot osynthesis allowed the Sun's energy to be harvested directly by life forms; the resultant oxygen accumulated in the atmosphere and formed a layer of ozone (a form of molecular oxygen [03]) in the upper

r atmosphere. The incorporation of smaller cells w ithin larger ones resulted in the development of c omplex cells called eukaryotes.[53] True multicell ular organisms formed as cells within colonies bec ame increasingly specialized. Aided by the absorpt ion of harmful ultraviolet radiation by the ozone layer, life colonized the surface of Earth.[54]

Since the 1960s, it has been hypothesized that sev ere glacial action between 750 and 580 mya, during the Neoproterozoic, covered much of the planet in a sheet of ice. This hypothesis has been termed "Snowball Earth", and is of particular interest because it preceded the Cambrian explosion, when mult icellular life forms began to proliferate.[55]

Following the Cambrian explosion, about 535 mya, t here have been five major mass extinctions.[56] Th e most recent such event was 66 mya, when an aster oid impact triggered the extinction of the (non-av ian) dinosaurs and other large reptiles, but spare d some small animals such as mammals, which then r esembled shrews. Over the past 66 myr, mammalian 1 ife has diversified, and several million years ago an African ape-like animal such as Orrorin tugene nsis gained the ability to stand upright.[57] This enabled tool use and encouraged communication tha t provided the nutrition and stimulation needed fo r a larger brain, which allowed the evolution of t he human race. The development of agriculture, and then civilization, allowed humans to influence th e Earth in a short time span as no other life form had, [58] affecting both the nature and quantity o f other life forms.

The present pattern of ice ages began about 40 mya and then intensified during the Pleistocene about 3 mya. High-latitude regions have since undergone repeated cycles of glaciation and thaw, repeating every 40-100,000 years. The last continental glaciation ended 10,000 years ago.[59]

Future

Main article: Future of the Earth

See also: Risks to civilization, humans and planet

Earth

The life cycle of the Sun
The future of the planet is closely tied to that o
f the Sun. As a result of the steady accumulation
of helium at the Sun's core, the star's total lumi
nosity will slowly increase. The luminosity of the
Sun will grow by 10% over the next 1.1 byr and by
40% over the next 3.5 byr.[60] Climate models ind
icate that the rise in radiation reaching the Eart
h is likely to have dire consequences, including t
he loss of the planet's oceans.[61]

The Earth's increasing surface temperature will ac celerate the inorganic CO2 cycle, reducing its con centration to levels lethally low for plants (10 p pm for C4 photosynthesis) in approximately 500-900 myr.[25] The lack of vegetation will result in th e loss of oxygen in the atmosphere, so animal life will become extinct within several million more y ears.[62] After another billion years all surface water will have disappeared[26] and the mean globa l temperature will reach 70 °C[62] (158 °F). The E arth is expected to be effectively habitable for a bout another 500 myr from that point, [25] although this may be extended up to 2.3 byr if the nitroge n is removed from the atmosphere.[27] Even if the Sun were eternal and stable, 27% of the water in t he modern oceans will descend to the mantle in one billion years due to reduced steam venting from m id-ocean ridges.[63]

The Sun, as part of its evolution, will become a r ed giant in about 5 byr. Models predict that the S un will expand out to about 250 times its present radius, roughly 1 AU (150,000,000 km).[60][64] Ear th's fate is less clear. As a red giant, the Sun w ill lose roughly 30% of its mass, so, without tida

l effects, the Earth will move to an orbit 1.7 AU (250,000,000 km) from the Sun, when the star reach es its maximum radius. The planet was therefore in itially expected to escape envelopment by the expa nded Sun's sparse outer atmosphere, though most, i f not all, remaining life would have been destroye d by the Sun's increased luminosity (peaking at ab out 5000 times its present level).[60] A 2008 simu lation indicates that Earth's orbit will decay due to tidal effects and drag, causing it to enter th e red giant Sun's atmosphere and be vaporized.[64] After that, the Sun's core will collapse into a w hite dwarf, as its outer layers are ejected into s pace as a planetary nebula. The matter that once m ade up the Earth will be released into interstella r space, where it may one day become incorporated into a new generation of planets and other celesti al bodies.

Composition and structure

Size comparison of inner planets (left to right): Mercury, Venus, Earth and Mars in true-color.

Main article: Earth science

Further information: Earth physical characteristics tables

Earth is a terrestrial planet, meaning that it is a rocky body, rather than a gas giant like Jupiter. It is the largest of the four solar terrestrial planets in size and mass. Of these four planets, E arth also has the highest density, the highest sur face gravity, the strongest magnetic field, and fa stest rotation, [65] and is probably the only one w ith active plate tectonics. [66]

Shape

Main article: Figure of the Earth

Chimborazo, Ecuador. The furthermost point on the E

arth's surface from its center.[67]
The shape of the Earth approximates an oblate sphe roid, a sphere flattened along the axis from pole to pole such that there is a bulge around the equa tor.[68] This bulge results from the rotation of the Earth, and causes the diameter at the equator to be 43 km (kilometer) larger than the pole-to-pole diameter.[69] For this reason the furthest point on the surface from the Earth's center of mass is the Chimborazo volcano in Ecuador.[70] The average diameter of the reference spheroid is about 12,742 km, which is approximately 40,000 km/p, as the meter was originally defined as 1/10,000,000 of the distance from the equator to the North Pole through Paris, France.[71]

Local topography deviates from this idealized sphe roid, although on a global scale, these deviations are small: Earth has a tolerance of about one part in about 584, or 0.17%, from the reference spheroid, which is less than the 0.22% tolerance allowed in billiard balls.[72] The largest local deviations in the rocky surface of the Earth are Mount Everest (8848 m above local sea level) and the Maria na Trench (10,911 m below local sea level). Due to the equatorial bulge, the surface locations farth est from the center of the Earth are the summits of Mount Chimborazo in Ecuador and Huascarán in Peru.[73][74][75]

Chemical composition of the crust[76] Compound Formula Composition Continental Oceanic silica SiO2 60.2% 48.6% 15.2% alumina Al2O3 CaO 5.5% 12.3% lime MqO 3.1% 6.8% magnesia iron(II) oxide FeO 3.8% sodium oxide Na20 3.0% potassium oxide K20 2.8% iron(III) oxide Fe2O3 2.5% water H2O 1.4% 1.1%

carbon dioxide CO2 1.2% 1.4% titanium dioxide TiO2 0.7% 1.4% phosphorus pentoxide P205 0.2% 0.3% 99.6% 99.9% Chemical composition See also: Abundance of elements on Earth The mass of the Earth is approximately 5.98×1024 k g. It is composed mostly of iron (32.1%), oxygen (30.1%), silicon (15.1%), magnesium (13.9%), sulfur (2.9%), nickel (1.8%), calcium (1.5%), and alumin ium (1.4%); with the remaining 1.2% consisting of trace amounts of other elements. Due to mass segre gation, the core region is believed to be primaril y composed of iron (88.8%), with smaller amounts o f nickel (5.8%), sulfur (4.5%), and less than 1% t race elements.[77]

The geochemist F. W. Clarke calculated that a litt le more than 47% of the Earth's crust consists of oxygen. The more common rock constituents of the Earth's crust are nearly all oxides; chlorine, sulf ur and fluorine are the only important exceptions to this and their total amount in any rock is usua lly much less than 1%. The principal oxides are si lica, alumina, iron oxides, lime, magnesia, potash and soda. The silica functions principally as an acid, forming silicates, and all the commonest min erals of igneous rocks are of this nature. From a computation based on 1,672 analyses of all kinds of rocks, Clarke deduced that 99.22% were composed of 11 oxides (see the table at right), with the other constituents occurring in minute quantities.[78]

Internal structure

Main article: Structure of the Earth
The interior of the Earth, like that of the other
terrestrial planets, is divided into layers by the
ir chemical or physical (rheological) properties,
but unlike the other terrestrial planets, it has a
distinct outer and inner core. The outer layer of
the Earth is a chemically distinct silicate solid
crust, which is underlain by a highly viscous sol

id mantle. The crust is separated from the mantle by the Mohorovicic discontinuity, and the thicknes s of the crust varies: averaging 6 km (kilometers) under the oceans and 30-50 km on the continents. The crust and the cold, rigid, top of the upper ma ntle are collectively known as the lithosphere, an d it is of the lithosphere that the tectonic plate s are comprised. Beneath the lithosphere is the as thenosphere, a relatively low-viscosity layer on w hich the lithosphere rides. Important changes in c rystal structure within the mantle occur at 410 an d 660 km below the surface, spanning a transition zone that separates the upper and lower mantle. Be neath the mantle, an extremely low viscosity liqui d outer core lies above a solid inner core.[79] Th e inner core may rotate at a slightly higher angul ar velocity than the remainder of the planet, adva ncing by $0.1-0.5^{\circ}$ per year.[80]

Geologic layers of the Earth[81]

```
Earth cutaway from core to exosphere. Not to scale.
Depth[82]
km Component Layer Density
a/cm3
        Lithosphere[note 8] -
0-60
        Crust[note 9]
0 - 35
                        2.2 - 2.9
35-60
        Upper mantle
                        3.4 - 4.4
            Mantle 3.4-5.6
  35-2890
100-700 Asthenosphere
                        9.9 - 12.2
2890-5100
            Outer core
                        12.8-13.1
5100-6378
            Inner core
Heat
Earth's internal heat comes from a combination of
residual heat from planetary accretion (about 20%)
 and heat produced through radioactive decay (80%)
.[83] The major heat-producing isotopes in the Ear
th are potassium-40, uranium-238, uranium-235, and
 thorium-232.[84] At the center of the planet, the
 temperature may be up to 7,000 K and the pressure
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could reach 360 GPa.[85] Because much of the heat

is provided by radioactive decay, scientists believe that early in Earth history, before isotopes with short half-lives had been depleted, Earth's heat production would have been much higher. This extra heat production, twice present-day at approximately 3 byr,[83] would have increased temperature gradients within the Earth, increasing the rates of mantle convection and plate tectonics, and allowing the production of igneous rocks such as komatites that are not formed today.[86]

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Present-day major heat-producing isotopes[87]
Isotope Heat release
W
kg isotope
Half-life
          Mean mantle concentration
years
kg isotope
kg mantle
Heat release
kg mantle
238U
          9.46 \times 10-5 \ 4.47 \times 109 \ \ 30.8 \times 10-9 \ \ 2.91 \times 10^{-2}
10 - 12
          5.69 \times 10-4 \ 7.04 \times 108 \ 0.22 \times 10-9 \ 1.25 \times
235U
10 - 13
          2.64 \times 10-5 \ 1.40 \times 1010 \ 124 \times 10-9 \ \ 3.27 \times
232Th
10-12
40K \ 2.92 \times 10-5 \ 1.25 \times 109 \ 36.9 \times 10-9 \ 1.08 \times 10-1
```

The mean heat loss from the Earth is 87 mW m-2, for a global heat loss of 4.42 × 1013 W.[88] A portion of the core's thermal energy is transported toward the crust by mantle plumes; a form of convection consisting of upwellings of higher-temperature rock. These plumes can produce hotspots and flood basalts.[89] More of the heat in the Earth is lost through plate tectonics, by mantle upwelling associated with mid-ocean ridges. The final major mode of heat loss is through conduction through the lithosphere, the majority of which occurs in the oce

ans because the crust there is much thinner than that of the continents.[90]

Tectonic plates
Earth's main plates[91]

Plate name Area
106 km2
Pacific Plate
103.3

African Plate[note 10]

78.0

North American Plate

75.9

Eurasian Plate

67.8

Antarctic Plate

60.9

Indo-Australian Plate

47.2

South American Plate

43.6

Main article: Plate tectonics

The mechanically rigid outer layer of the Earth, t he lithosphere, is broken into pieces called tecto nic plates. These plates are rigid segments that m ove in relation to one another at one of three typ es of plate boundaries: Convergent boundaries, at which two plates come together, Divergent boundari es, at which two plates are pulled apart, and Tran sform boundaries, in which two plates slide past o ne another laterally. Earthquakes, volcanic activi ty, mountain-building, and oceanic trench formatio n can occur along these plate boundaries.[92] The tectonic plates ride on top of the asthenosphere, the solid but less-viscous part of the upper mantl e that can flow and move along with the plates, [93]] and their motion is strongly coupled with convec tion patterns inside the Earth's mantle.

As the tectonic plates migrate across the planet, the ocean floor is subducted under the leading edg es of the plates at convergent boundaries. At the same time, the upwelling of mantle material at div ergent boundaries creates mid-ocean ridges. The combination of these processes continually recycles the oceanic crust back into the mantle. Due to this recycling, most of the ocean floor is less than 100 myr old in age. The oldest oceanic crust is located in the Western Pacific, and has an estimated age of about 200 myr.[94][95] By comparison, the oldest dated continental crust is 4,030 myr.[96]

The seven major plates are the Pacific, North American, Eurasian, African, Antarctic, Indo-Australian, and South American. Other notable plates include the Arabian Plate, the Caribbean Plate, the Nazca Plate off the west coast of South America and the Scotia Plate in the southern Atlantic Ocean. The Australian Plate fused with the Indian Plate between 50 and 55 mya. The fastest-moving plates are the oceanic plates, with the Cocos Plate advancing at a rate of 75 mm/year[97] and the Pacific Plate moving 52-69 mm/year. At the other extreme, the slowest-moving plate is the Eurasian Plate, progressing at a typical rate of about 21 mm/year.[98]

Surface

Main articles: Landform and Extreme points of Earth

The Earth's terrain varies greatly from place to p lace. About 70.8%[13] of the surface is covered by water, with much of the continental shelf below s ea level. This equates to 361.132 million km2 (139.43 million sq mi).[99] The submerged surface has mountainous features, including a globe-spanning m id-ocean ridge system, as well as undersea volcano es,[69] oceanic trenches, submarine canyons, ocean ic plateaus and abyssal plains. The remaining 29.2% (148.94 million km2, or 57.51 million sq mi) not covered by water consists of mountains, deserts, plains, plateaus, and other geomorphologies.

The planetary surface undergoes reshaping over geo

logical time periods due to tectonics and erosion. The surface features built up or deformed through plate tectonics are subject to steady weathering from precipitation, thermal cycles, and chemical effects. Glaciation, coastal erosion, the build-up of coral reefs, and large meteorite impacts[100] a lso act to reshape the landscape.

Present-day Earth altimetry and bathymetry. Data f rom the National Geophysical Data Center's Terrain Base Digital Terrain Model.

The continental crust consists of lower density ma terial such as the igneous rocks granite and andes ite. Less common is basalt, a denser volcanic rock that is the primary constituent of the ocean floo rs.[101] Sedimentary rock is formed from the accum ulation of sediment that becomes compacted togethe r. Nearly 75% of the continental surfaces are cove red by sedimentary rocks, although they form only about 5% of the crust.[102] The third form of rock material found on Earth is metamorphic rock, whic h is created from the transformation of pre-existi ng rock types through high pressures, high tempera tures, or both. The most abundant silicate mineral s on the Earth's surface include quartz, the felds pars, amphibole, mica, pyroxene and olivine.[103] Common carbonate minerals include calcite (found i n limestone) and dolomite.[104]

The pedosphere is the outermost layer of the Earth that is composed of soil and subject to soil form ation processes. It exists at the interface of the lithosphere, atmosphere, hydrosphere and biospher e. Currently the total arable land is 13.31% of the land surface, with only 4.71% supporting permane nt crops.[14] Close to 40% of the Earth's land surface is presently used for cropland and pasture, or an estimated 1.3×107 km2 of cropland and 3.4×107 km2 of pastureland.[105]

The elevation of the land surface of the Earth varies from the low point of -418 m at the Dead Sea, to a 2005-estimated maximum altitude of 8,848 m at the top of Mount Everest. The mean height of land above sea level is 840 m.[106]

Hydrosphere

Main article: Hydrosphere

Elevation histogram of the surface of the Earth The abundance of water on Earth's surface is a unique feature that distinguishes the "Blue Planet" from others in the Solar System. The Earth's hydros phere consists chiefly of the oceans, but technically includes all water surfaces in the world, including inland seas, lakes, rivers, and underground waters down to a depth of 2,000 m. The deepest underwater location is Challenger Deep of the Mariana Trench in the Pacific Ocean with a depth of -10,9 11.4 m.[note 11][107]

The mass of the oceans is approximately 1.35×1018 metric tons, or about 1/4400 of the total mass of the Earth. The oceans cover an area of 3.618×108 k m2 with a mean depth of 3,682 m, resulting in an e stimated volume of 1.332×109 km3.[108] If all the land on Earth were spread evenly, water would rise to an altitude of more than 2.7 km.[note 12] About 97.5% of the water is saline, while the remaining 2.5% is fresh water. Most fresh water, about 68.7%, is currently ice.[109]

The average salinity of the Earth's oceans is about 35 grams of salt per kilogram of sea water (35 % salt).[110] Most of this salt was released from volcanic activity or extracted from cool, igneous rocks.[111] The oceans are also a reservoir of dissolved atmospheric gases, which are essential for the survival of many aquatic life forms.[112] Sea water has an important influence on the world's climate, with the oceans acting as a large heat reser

voir.[113] Shifts in the oceanic temperature distribution can cause significant weather shifts, such as the El Niño-Southern Oscillation.[114]

Atmosphere

Main article: Atmosphere of Earth

This is a picture of Earth in ultraviolet light, taken from the surface of the Moon. The day-side reflects a lot of UV light from the Sun, but the night-side shows bands of UV emission from the aurora caused by charged particles.[115]

The atmospheric pressure on the surface of the Ear

The atmospheric pressure on the surface of the Ear th averages 101.325 kPa, with a scale height of ab out 8.5 km.[3] It is 78% nitrogen and 21% oxygen, with trace amounts of water vapor, carbon dioxide and other gaseous molecules. The height of the tro posphere varies with latitude, ranging between 8 km at the poles to 17 km at the equator, with some variation resulting from weather and seasonal fact ors.[116]

Earth's biosphere has significantly altered its at mosphere. Oxygenic photosynthesis evolved 2.7 bya, forming the primarily nitrogen-oxygen atmosphere of today. This change enabled the proliferation of aerobic organisms as well as the formation of the ozone layer which blocks ultraviolet solar radiat ion, permitting life on land. Other atmospheric fu nctions important to life on Earth include transpo rting water vapor, providing useful gases, causing small meteors to burn up before they strike the s urface, and moderating temperature.[117] This last phenomenon is known as the greenhouse effect: tra ce molecules within the atmosphere serve to captur e thermal energy emitted from the ground, thereby raising the average temperature. Water vapor, carb on dioxide, methane and ozone are the primary gree nhouse gases in the Earth's atmosphere. Without th is heat-retention effect, the average surface woul d be -18 °C, in contrast to the current +15 °C, an d life would likely not exist.[118]

Weather and climate
Main articles: Weather and Climate

Satellite cloud cover image of Earth using NASA's Moderate-Resolution Imaging Spectroradiometer The Earth's atmosphere has no definite boundary, s lowly becoming thinner and fading into outer space. Three-quarters of the atmosphere's mass is contained within the first 11 km of the planet's surface. This lowest layer is called the troposphere. En ergy from the Sun heats this layer, and the surface below, causing expansion of the air. This lower density air then rises, and is replaced by cooler, higher density air. The result is atmospheric circulation that drives the weather and climate through redistribution of heat energy.[119]

The primary atmospheric circulation bands consist of the trade winds in the equatorial region below 30° latitude and the westerlies in the mid-latitud es between 30° and 60°.[120] Ocean currents are al so important factors in determining climate, particularly the thermohaline circulation that distributes heat energy from the equatorial oceans to the polar regions.[121]

Water vapor generated through surface evaporation is transported by circulatory patterns in the atmo sphere. When atmospheric conditions permit an upli ft of warm, humid air, this water condenses and se ttles to the surface as precipitation.[119] Most of the water is then transported to lower elevation s by river systems and usually returned to the oce ans or deposited into lakes. This water cycle is a vital mechanism for supporting life on land, and is a primary factor in the erosion of surface feat ures over geological periods. Precipitation patter ns vary widely, ranging from several meters of wat er per year to less than a millimeter. Atmospheric

circulation, topological features and temperature differences determine the average precipitation that falls in each region.[122]

The amount of solar energy reaching the Earth's de creases with increasing latitude. At higher latitu des the sunlight reaches the surface at lower angl es and it must pass through thicker columns of the atmosphere. As a result, the mean annual air temp erature at sea level decreases by about 0.4 °C per degree of latitude away from the equator.[123] Th e Earth can be sub-divided into specific latitudin al belts of approximately homogeneous climate. Ran ging from the equator to the polar regions, these are the tropical (or equatorial), subtropical, tem perate and polar climates.[124] Climate can also b e classified based on the temperature and precipit ation, with the climate regions characterized by f airly uniform air masses. The commonly used Köppen climate classification system (as modified by Wla dimir Köppen's student Rudolph Geiger) has five br oad groups (humid tropics, arid, humid middle lati tudes, continental and cold polar), which are furt her divided into more specific subtypes.[120]

Upper atmosphere

This view from orbit shows the full Moon partially obscured and deformed by the Earth's atmosphere. NASA image

See also: Outer space

Above the troposphere, the atmosphere is usually divided into the stratosphere, mesosphere, and ther mosphere.[117] Each layer has a different lapse rate, defining the rate of change in temperature with height. Beyond these, the exosphere thins out in to the magnetosphere, where the Earth's magnetic fields interact with the solar wind.[125] Within the stratosphere is the ozone layer, a component that partially shields the surface from ultraviolet light and thus is important for life on Earth. The

Kármán line, defined as 100 km above the Earth's s urface, is a working definition for the boundary b etween atmosphere and space.[126]

Thermal energy causes some of the molecules at the outer edge of the Earth's atmosphere to increase their velocity to the point where they can escape from the planet's gravity. This causes a slow but steady leakage of the atmosphere into space. Becau se unfixed hydrogen has a low molecular weight, it can achieve escape velocity more readily and it 1 eaks into outer space at a greater rate than other gasses.[127] The leakage of hydrogen into space c ontributes to the pushing of the Earth from an ini tially reducing state to its current oxidizing one . Photosynthesis provided a source of free oxygen, but the loss of reducing agents such as hydrogen is believed to have been a necessary precondition for the widespread accumulation of oxygen in the a tmosphere.[128] Hence the ability of hydrogen to e scape from the Earth's atmosphere may have influen ced the nature of life that developed on the plane t.[129] In the current, oxygen-rich atmosphere mos t hydrogen is converted into water before it has a n opportunity to escape. Instead, most of the hydr ogen loss comes from the destruction of methane in the upper atmosphere.[130]

Magnetic field

Schematic of Earth's magnetosphere. The solar wind flows from left to right
Main article: Earth's magnetic field
The Earth's magnetic field is shaped roughly as a magnetic dipole, with the poles currently located proximate to the planet's geographic poles. At the equator of the magnetic field, the magnetic field strength at the planet's surface is 3.05 × 10-5 T, with global magnetic dipole moment of 7.91 × 101 5 T m3.[131] According to dynamo theory, the field is generated within the molten outer core region

where heat creates convection motions of conducting materials, generating electric currents. These in turn produce the Earth's magnetic field. The convection movements in the core are chaotic; the magnetic poles drift and periodically change alignment. This causes field reversals at irregular intervals averaging a few times every million years. The most recent reversal occurred approximately 700,00 years ago.[132][133]

The field forms the magnetosphere, which deflects particles in the solar wind. The sunward edge of the bow shock is located at about 13 times the radius of the Earth. The collision between the magnetic field and the solar wind forms the Van Allen radiation belts, a pair of concentric, torus-shaped regions of energetic charged particles. When the plasma enters the Earth's atmosphere at the magnetic poles, it forms the aurora.[134]

Orbit and rotation

Rotation

Main article: Earth's rotation

Earth's axial tilt (or obliquity) and its relation to the rotation axis and plane of orbit
Earth's rotation period relative to the Sun—its me an solar day—is 86,400 seconds of mean solar time (86,400.0025 SI seconds).[135] As the Earth's solar day is now slightly longer than it was during the 19th century due to tidal acceleration, each day varies between 0 and 2 SI ms longer.[136][137]

Earth's rotation period relative to the fixed star s, called its stellar day by the International Ear th Rotation and Reference Systems Service (IERS), is 86164.098903691 seconds of mean solar time (UT1), or 23h 56m 4.098903691s.[2][note 13] Earth's rotation period relative to the precessing or moving mean vernal equinox, misnamed its sidereal day, i

s 86164.09053083288 seconds of mean solar time (UT 1) (23h 56m 4.09053083288s).[2] Thus the sidereal day is shorter than the stellar day by about 8.4 m s.[138] The length of the mean solar day in SI sec onds is available from the IERS for the periods 16 23-2005[139] and 1962-2005.[140]

Apart from meteors within the atmosphere and low-orbiting satellites, the main apparent motion of celestial bodies in the Earth's sky is to the west at a rate of 15°/h = 15'/min. For bodies near the celestial equator, this is equivalent to an apparent diameter of the Sun or Moon every two minutes; from the planet's surface, the apparent sizes of the Sun and the Moon are approximately the same.[141][142]

Orbit

Main article: Earth's orbit

Earth orbits the Sun at an average distance of about 150 million kilometers every 365.2564 mean solar days, or one sidereal year. From Earth, this gives an apparent movement of the Sun eastward with respect to the stars at a rate of about 1°/day, which is one apparent Sun or Moon diameter every 12 hours. Due to this motion, on average it takes 24 hours—a solar day—for Earth to complete a full rotation about its axis so that the Sun returns to the meridian. The orbital speed of the Earth averages about 29.8 km/s (107,000 km/h), which is fast enough to travel a distance equal to the planet's diameter, about 12,742 km, in seven minutes, and the distance to the Moon, 384,000 km, in about 3.5 hours.[3]

The Moon revolves with the Earth around a common be arycenter every 27.32 days relative to the background stars. When combined with the Earth-Moon system's common revolution around the Sun, the period of the synodic month, from new moon to new moon, is 29.53 days. Viewed from the celestial north pole, the motion of Earth, the Moon and their axial rot

ations are all counterclockwise. Viewed from a van tage point above the north poles of both the Sun a nd the Earth, the Earth revolves in a counterclock wise direction about the Sun. The orbital and axia l planes are not precisely aligned: Earth's axis is tilted some 23.4 degrees from the perpendicular to the Earth-Sun plane (the ecliptic), and the Earth-Moon plane is tilted up to ±5.1 degrees against the Earth-Sun plane. Without this tilt, there would be an eclipse every two weeks, alternating between lunar eclipses and solar eclipses.[3][143]

The Hill sphere, or gravitational sphere of influe nce, of the Earth is about 1.5 Gm or 1,500,000 km in radius.[144][note 14] This is the maximum distance at which the Earth's gravitational influence is stronger than the more distant Sun and planets. Objects must orbit the Earth within this radius, or they can become unbound by the gravitational per turbation of the Sun.

Earth, along with the Solar System, is situated in the Milky Way galaxy and orbits about 28,000 ligh t years from the center of the galaxy. It is curre ntly about 20 light years above the galactic plane in the Orion spiral arm.[145]

Axial tilt and seasons

Earth and Moon from Mars, imaged by Mars Reconnais sance Orbiter. From space, the Earth can be seen to go through phases similar to the phases of the Moon.

Main article: Axial tilt

Due to the axial tilt of the Earth, the amount of sunlight reaching any given point on the surface v aries over the course of the year. This causes sea sonal change in climate, with summer in the northern hemisphere occurring when the North Pole is pointing toward the Sun, and winter taking place when the pole is pointed away. During the summer, the

day lasts longer and the Sun climbs higher in the sky. In winter, the climate becomes generally cool er and the days shorter. Above the Arctic Circle, an extreme case is reached where there is no dayli ght at all for part of the year—a polar night. In the southern hemisphere the situation is exactly r eversed, with the South Pole oriented opposite the direction of the North Pole.

By astronomical convention, the four seasons are d etermined by the solstices—the point in the orbit of maximum axial tilt toward or away from the Sun—and the equinoxes, when the direction of the tilt and the direction to the Sun are perpendicular. In the northern hemisphere, Winter Solstice occurs on about December 21, Summer Solstice is near June 21, Spring Equinox is around March 20 and Autumnal Equinox is about September 23. In the Southern he misphere, the situation is reversed, with the Summ er and Winter Solstices exchanged and the Spring a nd Autumnal Equinox dates switched.[146]

The angle of the Earth's tilt is relatively stable over long periods of time. The tilt does undergo nutation; a slight, irregular motion with a main p eriod of 18.6 years.[147] The orientation (rather than the angle) of the Earth's axis also changes o ver time, precessing around in a complete circle o ver each 25,800 year cycle; this precession is the reason for the difference between a sidereal year and a tropical year. Both of these motions are ca used by the varying attraction of the Sun and Moon on the Earth's equatorial bulge. From the perspec tive of the Earth, the poles also migrate a few me ters across the surface. This polar motion has mul tiple, cyclical components, which collectively are termed quasiperiodic motion. In addition to an an nual component to this motion, there is a 14-month cycle called the Chandler wobble. The rotational velocity of the Earth also varies in a phenomenon known as length of day variation.[148]

In modern times, Earth's perihelion occurs around January 3, and the aphelion around July 4. These d ates change over time due to precession and other orbital factors, which follow cyclical patterns kn own as Milankovitch cycles. The changing Earth-Sun distance causes an increase of about 6.9% [note 15]] in solar energy reaching the Earth at perihelion relative to aphelion. Since the southern hemisphe re is tilted toward the Sun at about the same time that the Earth reaches the closest approach to th e Sun, the southern hemisphere receives slightly m ore energy from the Sun than does the northern ove r the course of a year. This effect is much less s ignificant than the total energy change due to the axial tilt, and most of the excess energy is abso rbed by the higher proportion of water in the sout hern hemisphere.[149]

Moon

Characteristics
Diameter 3,474.8 km
Mass 7.349×1022 kg
Semi-major axis 384,400 km
Orbital period 27 d 7 h 43.7 m

Details of the Earth-Moon system. Besides the radius of each object, the radius to the Earth-Moon barycenter is shown. Photos from NASA. Data from NASA. The Moon's axis is located by Cassini's third law.

Main article: Moon

The Moon is a relatively large, terrestrial, plane t-like satellite, with a diameter about one-quarter of the Earth's. It is the largest moon in the Solar System relative to the size of its planet, although Charon is larger relative to the dwarf planet Pluto. The natural satellites orbiting other planets are called "moons" after Earth's Moon.

The gravitational attraction between the Earth and

Moon causes tides on Earth. The same effect on the Moon has led to its tidal locking: its rotation period is the same as the time it takes to orbit the Earth. As a result, it always presents the same face to the planet. As the Moon orbits Earth, different parts of its face are illuminated by the Sun, leading to the lunar phases; the dark part of the face is separated from the light part by the solar terminator.

Due to their tidal interaction, the Moon recedes f rom Earth at the rate of approximately 38 mm a year. Over millions of years, these tiny modification s—and the lengthening of Earth's day by about 23 μ s a year—add up to significant changes.[150] During the Devonian period, for example, (approximately 410 mya) there were 400 days in a year, with each day lasting 21.8 hours.[151]

The Moon may have dramatically affected the development of life by moderating the planet's climate. Paleontological evidence and computer simulations show that Earth's axial tilt is stabilized by tidal interactions with the Moon.[152] Some theorists believe that without this stabilization against the torques applied by the Sun and planets to the Earth's equatorial bulge, the rotational axis might be chaotically unstable, exhibiting chaotic changes over millions of years, as appears to be the case for Mars.[153]

Viewed from Earth, the Moon is just far enough awa y to have almost the same apparent-sized disk as the Sun. The angular size (or solid angle) of these two bodies match because, although the Sun's diameter is about 400 times as large as the Moon's, it is also 400 times more distant.[142] This allows total and annular solar eclipses to occur on Earth.

The most widely accepted theory of the Moon's origin, the giant impact theory, states that it formed from the collision of a Mars-size protoplanet cal

led Theia with the early Earth. This hypothesis ex plains (among other things) the Moon's relative la ck of iron and volatile elements, and the fact that its composition is nearly identical to that of the Earth's crust.[154]

A scale representation of the relative sizes of, and average distance between, Earth and Moon
Asteroids and artificial satellites

The International Space Station is an artificial sa tellite that orbits Earth.

Earth has at least five co-orbital asteroids, including 3753 Cruithne and 2002 AA29.[155][156]

On July 27, 2011, astronomers reported a trojan as teroid companion, 2010 TK7, librating around the leading Lagrange triangular point, L4, of Earth in Earth's orbit around the Sun.[157][158]

As of 2011, there are 931 operational, man-made sa tellites orbiting the Earth.[159] There are also i noperative satellites and over 300,000 pieces of s pace debris. Earth's largest artificial satellite is the International Space Station.

Habitability

See also: Planetary habitability

A planet that can sustain life is termed habitable, even if life did not originate there. The Earth provides liquid water—an environment where complex organic molecules can assemble and interact, and sufficient energy to sustain metabolism.[160] The distance of the Earth from the Sun, as well as its orbital eccentricity, rate of rotation, axial til t, geological history, sustaining atmosphere and p rotective magnetic field all contribute to the cur

rent climatic conditions at the surface.[161]

Biosphere

Main article: Biosphere

The planet's life forms are sometimes said to form a "biosphere". This biosphere is generally believ ed to have begun evolving about 3.5 bya. The biosphere is divided into a number of biomes, inhabited by broadly similar plants and animals. On land, b iomes are separated primarily by differences in latitude, height above sea level and humidity. Terrestrial biomes lying within the Arctic or Antarctic Circles, at high altitudes or in extremely arid a reas are relatively barren of plant and animal life; species diversity reaches a peak in humid lowlands at equatorial latitudes.[162]

Natural resources and land use Main article: Natural resource

The Earth provides resources that are exploitable by humans for useful purposes. Some of these are n on-renewable resources, such as mineral fuels, that are difficult to replenish on a short time scale.

Large deposits of fossil fuels are obtained from the Earth's crust, consisting of coal, petroleum, natural gas and methane clathrate. These deposits a resused by humans both for energy production and a sefection for chemical production. Mineral ore bodies have also been formed in Earth's crust through a process of Ore genesis, resulting from action sof erosion and plate tectonics.[163] These bodies form concentrated sources for many metals and other useful elements.

The Earth's biosphere produces many useful biological products for humans, including (but far from limited to) food, wood, pharmaceuticals, oxygen, and the recycling of many organic wastes. The land-based ecosystem depends upon topsoil and fresh water, and the oceanic ecosystem depends upon dissolved nutrients washed down from the land.[164] Humans

also live on the land by using building materials to construct shelters. In 1993, human use of land is approximately:

Land use Arable land Permanent crops Permanent pastures Forests and woodland Urban areas Other

Percentage 13.13%[14] 4.71%[14] 26% 32% 1.5% 30%

The estimated amount of irrigated land in 1993 was 2,481,250 km2.[14]

Natural and environmental hazards
Large areas of the Earth's surface are subject to
extreme weather such as tropical cyclones, hurrica
nes, or typhoons that dominate life in those areas
. From 1980 to 2000, these events caused an averag
e of 11,800 deaths per year.[165] Many places are
subject to earthquakes, landslides, tsunamis, volc
anic eruptions, tornadoes, sinkholes, blizzards, f
loods, droughts, wildfires, and other calamities a
nd disasters.

Many localized areas are subject to human-made pollution of the air and water, acid rain and toxic substances, loss of vegetation (overgrazing, defore station, desertification), loss of wildlife, species extinction, soil degradation, soil depletion, erosion, and introduction of invasive species.

According to the United Nations, a scientific consensus exists linking human activities to global warming due to industrial carbon dioxide emissions. This is predicted to produce changes such as the melting of glaciers and ice sheets, more extreme te mperature ranges, significant changes in weather and a global rise in average sea levels.[166]

Human geography
Main articles: Human geography and World

"Nightime imagery from orbit provides an instant u nderstanding about humanity's footprint on the Ear th's surface, .. " -NASA Scientist[167] This view of Liège, Belgium at night reveals surrounding towns, roads, and agriculture

Cartography, the study and practice of map making, and vicariously geography, have historically been the disciplines devoted to depicting the Earth. Surveying, the determination of locations and distances, and to a lesser extent navigation, the determination of position and direction, have developed alongside cartography and geography, providing and suitably quantifying the requisite information.

Earth has reached approximately 7,000,000,000 human inhabitants as of October 31, 2011.[168] Project ions indicate that the world's human population will reach 9.2 billion in 2050.[169] Most of the growth is expected to take place in developing nations. Human population density varies widely around the world, but a majority live in Asia. By 2020, 60% of the world's population is expected to be living in urban, rather than rural, areas.[170]

It is estimated that only one-eighth of the surface of the Earth is suitable for humans to live on—three-quarters is covered by oceans, and half of the land area is either desert (14%),[171] high mountains (27%),[172] or other less suitable terrain. The northernmost permanent settlement in the world is Alert, on Ellesmere Island in Nunavut, Canada. [173] (82°28'N) The southernmost is the Amundsen-Scott South Pole Station, in Antarctica, almost exactly at the South Pole. (90°S)

Independent sovereign nations claim the planet's e ntire land surface, except for some parts of Antar ctica and the odd unclaimed area of Bir Tawil betw een Egypt and Sudan. As of 2013 there are 206 sove reign states, including the 193 United Nations mem ber states. In addition, there are 59 dependent te rritories, and a number of autonomous areas, terri

tories under dispute and other entities.[14] Histo rically, Earth has never had a sovereign governmen t with authority over the entire globe, although a number of nation-states have striven for world do mination and failed.[174]

The United Nations is a worldwide intergovernmental organization that was created with the goal of intervening in the disputes between nations, thereby avoiding armed conflict.[175] The U.N. serves primarily as a forum for international diplomacy and international law. When the consensus of the membership permits, it provides a mechanism for armed intervention.[176]

The first human to orbit the Earth was Yuri Gagari n on April 12, 1961.[177] In total, about 487 peop le have visited outer space and reached Earth orbit as of July 30, 2010, and, of these, twelve have walked on the Moon.[178][179][180] Normally the only humans in space are those on the International Space Station. The station's crew, currently six people, is usually replaced every six months.[181] The furthest humans have travelled from Earth is 400,171 km, achieved during the 1970 Apollo 13 miss ion.[182]

The 7 continents of Earth:[183] North America, South America, Antarctica, Africa, Europe, Asia, Australia

The Earth at night in 2000, a composite of DMSP/OL S ground illumination data on a simulated night-time image of the world. This image is not photographic and many features are brighter than they would appear to a direct observer.

ISS video beginning just south-east of Alaska. The first city that the ISS passes over (seen approximately 10 seconds into the video) is San Francisco and the surrounding areas in California. The Gold

en Gate Bridge is located by a small strip of lights just before the city of San Francisco, nearest to the clouds on the right of the image. Lightning storms can be seen on the Pacific Ocean coastline, with clouds overhead. As the video continues, the ISS passes over Central America (green lights are visible here), with the Yucatan Peninsula on the left. The pass ends as the ISS is over the capital city of Bolivia, La Paz. Cultural viewpoint

Main article: Earth in culture

The first photograph ever taken by astronauts of an "Earthrise", from Apollo 8
The standard astronomical symbol of the Earth cons ists of a cross circumscribed by a circle.[184]

Unlike the rest of the planets in the Solar System , humankind did not begin to view the Earth as a m oving object in orbit around the Sun until the 16t h century.[185] Earth has often been personified a s a deity, in particular a goddess. In many cultur es a mother goddess is also portrayed as a fertili ty deity. Creation myths in many religions recall a story involving the creation of the Earth by a s upernatural deity or deities. A variety of religio us groups, often associated with fundamentalist br anches of Protestantism[186] or Islam,[187] assert that their interpretations of these creation myth s in sacred texts are literal truth and should be considered alongside or replace conventional scien tific accounts of the formation of the Earth and t he origin and development of life.[188] Such asser tions are opposed by the scientific community[189] [190] and by other religious groups.[191][192][193] A prominent example is the creation-evolution co ntroversy.

In the past there were varying levels of belief in a flat Earth, [194] but this was displaced by the

concept of a spherical Earth due to observation and circumnavigation.[195] The human perspective regarding the Earth has changed following the advent of spaceflight, and the biosphere is now widely viewed from a globally integrated perspective.[196][197] This is reflected in a growing environmental movement that is concerned about humankind's effects on the planet.[198]