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# Spacecraft Attitude Dynamics

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## Disturbance Torques – SRP and air drag

↑  
WRT THE ATMOSPHERE OF  
THE EARTH

↓  
SOLAR RADIATION PRESSURE  
DEPENDS ON ORIENTATION AND POSITION WRT THE  
SUN

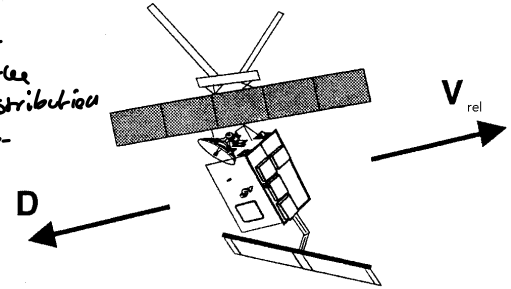
# Force due to atmospheric drag – inertial frame

The drag force on an Earth orbiting spacecraft is

It is not the same order as the atmospheric drag → the role of continuum aerodynamics are not precise anymore → we need to study the input dynamics of the distribution of particles in a volume

But we use approximate similarly to the continuum dynamics rule

$$\bar{F}_i = -\frac{1}{2} \rho C_D v_{rel}^2 \frac{\mathbf{v}_{rel}}{\|\mathbf{v}_{rel}\|} A_{cross}$$



$\rho(h, t)$  atmospheric density

$\mathbf{v}_{rel}$  relative speed

$A_{cross}$  cross sectional area perpendicular to  $\mathbf{v}_{rel}$

$C_D \approx 2.2, 1.5 < C_D < 2.6$  drag coefficient (LEO)

↳ for a flat plate and it take in account for some terms omitted in the equation.

flat plate drag coefficient

# Air density example data set

$$\bar{F}_i = -\frac{1}{2} \rho C_D v_{rel}^2 \frac{v_{rel}}{\|v_{rel}\|} A_{cross}$$

( $\rho$  need to be computed or cent for 3 surface of a cuboid satellite.)

altit. (km)	10	20	30	40	50	60	70	80	90	100
density (g/cm <sup>3</sup> )	4,02e-04	8,34e-05	1,57e-05	3,18e-06	8,37e-07	2,33e-07	5,86e-08	1,40e-08	2,99e-09	5,17e-10
altit. (km)	110	120	130	140	150	160	170	180	190	200
density (g/cm <sup>3</sup> )	8,42e-11	1,84e-11	7,36e-12	3,78e-12	2,19e-12	1,37e-12	9,00e-13	6,15e-13	4,32e-13	3,10e-13
altit. (km)	210	220	230	240	250	260	270	280	290	300
density (g/cm <sup>3</sup> )	2,27e-13	1,68e-13	1,26e-13	9,58e-14	7,35e-14	5,68e-14	4,43e-14	3,48e-14	2,75e-14	2,18e-14
altit. (km)	310	320	330	340	350	360	370	380	390	400
density (g/cm <sup>3</sup> )	1,74e-14	1,40e-14	1,13e-14	9,10e-15	7,39e-15	6,02e-15	4,92e-15	4,03e-15	3,31e-15	2,72e-15
altit. (km)	410	420	430	440	450	460	470	480	490	500
density (g/cm <sup>3</sup> )	2,25e-15	1,86e-15	1,54e-15	1,28e-15	1,07e-15	8,89e-16	7,43e-16	6,22e-16	5,22e-16	4,39e-16
altit. (km)	510	520	530	540	550	560	570	580	590	600
density (g/cm <sup>3</sup> )	3,71e-16	3,13e-16	2,66e-16	2,26e-16	1,93e-16	1,65e-16	1,41e-16	1,22e-16	1,05e-16	9,14e-17
altit. (km)	610	620	630	640	650	660	670	680	690	700
density (g/cm <sup>3</sup> )	7,96e-17	6,97e-17	6,12e-17	5,41e-17	4,79e-17	4,27e-17	3,82e-17	3,44e-17	3,10e-17	2,82e-17
altit. (km)	710	720	730	740	750	760	770	780	790	800
density (g/cm <sup>3</sup> )	2,57e-17	2,35e-17	2,16e-17	1,99e-17	1,84e-17	1,71e-17	1,59e-17	1,49e-17	1,39e-17	1,31e-17
altit. (km)	810	820	830	840	850	860	870	880	890	900
density (g/cm <sup>3</sup> )	1,23e-17	1,16e-17	1,10e-17	1,04e-17	9,86e-18	9,37e-18	8,91e-18	8,48e-18	8,09e-18	7,72e-18
altit. (km)	910	920	930	940	950	960	970	980	990	1000
density (g/cm <sup>3</sup> )	7,37e-18	7,04e-18	6,74e-18	6,45e-18	6,18e-18	5,92e-18	5,68e-18	5,44e-18	5,22e-18	5,02e-18



# Atmospheric density model

$$\rho(h, t) = \rho_0 \exp \left[ -\frac{h - h_0}{H} \right]$$

$\rho_0$  reference density

$h, h_0$  actual and reference altitude

$H$  scale height

**TABLE 7-4. Exponential Atmospheric Model.** Although a very simple approach, this method yields moderate results for general studies. Source: Wertz, 1978, 820, which uses the *U.S. Standard Atmosphere* (1976) for 0 km, CIRA-72 for 25–500 km, and CIRA-72 with  $T_\infty = 1000$  K for 500–1000 km. The scale heights have been adjusted to maintain a piecewise-continuous formulation of the density.

Altitude $h_{ellp}$ (km)	Base Altitude $h_0$ (km)	Nominal Density $\rho_0$ (kg/m <sup>3</sup> )	Scale Height $H$ (km)	Altitude $h_{ellp}$ (km)	Base Altitude $h_0$ (km)	Nominal Density $\rho_0$ (kg/m <sup>3</sup> )	Scale Height $H$ (km)
0–25	0	1.225	7.249	150–180	150	$2.070 \times 10^{-9}$	22.523
25–30	25	$3.899 \times 10^{-2}$	6.349	180–200	180	$5.464 \times 10^{-10}$	29.740
30–40	30	$1.774 \times 10^{-2}$	6.682	200–250	200	$2.789 \times 10^{-10}$	37.105
40–50	40	$3.972 \times 10^{-3}$	7.554	250–300	250	$7.248 \times 10^{-11}$	45.546
50–60	50	$1.057 \times 10^{-3}$	8.382	300–350	300	$2.418 \times 10^{-11}$	53.628
60–70	60	$3.206 \times 10^{-4}$	7.714	350–400	350	$9.158 \times 10^{-12}$	53.298
70–80	70	$8.770 \times 10^{-5}$	6.549	400–450	400	$3.725 \times 10^{-12}$	58.515
80–90	80	$1.905 \times 10^{-5}$	5.799	450–500	450	$1.585 \times 10^{-12}$	60.828
90–100	90	$3.396 \times 10^{-6}$	5.382	500–600	500	$6.967 \times 10^{-13}$	63.822
100–110	100	$5.297 \times 10^{-7}$	5.877	600–700	600	$1.454 \times 10^{-13}$	71.835
110–120	110	$9.661 \times 10^{-8}$	7.263	700–800	700	$3.614 \times 10^{-14}$	88.667
120–130	120	$2.438 \times 10^{-8}$	9.473	800–900	800	$1.170 \times 10^{-14}$	124.64
130–140	130	$8.484 \times 10^{-9}$	12.636	900–1000	900	$5.245 \times 10^{-15}$	181.05
140–150	140	$3.845 \times 10^{-9}$	16.149	1000–	1000	$3.019 \times 10^{-15}$	268.00

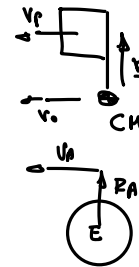
Eq. (7-31) requires knowledge of the actual altitude, found by subtracting the Earth's radius (6378.137 km) from the satellite's given radius ( $h_{ellp} = 747.2119$  km). Now, if we use values from Table 7-4, Eq. (7-31) becomes

$$\rho = 3.614 \times 10^{-14} \exp \left[ -\frac{747.2119 - 700}{88.667} \right] = 2.1219854 \times 10^{-14} \frac{\text{kg}}{\text{m}^3}$$

▲ Note that the units in the exponential cancel (all are km), and the result is less than the base value at 700 km, as we would expect.



# Relative orbital velocity in the inertial frame



$$\underline{v} = \underline{v}_s + \underline{\omega} \wedge \underline{r}$$

$$\underline{r}_p = \underline{\omega} \wedge \underline{r}_A$$

$$\underline{v}_{rel} = \underline{v}_p - \underline{v}_A$$



$$\underline{v}_{orbit} = \frac{d\vec{R}}{dt}$$

$$\underline{v}_{atmosphere} = \omega_{Earth} \times \vec{R}$$

Dependent on the orbit that we define

The atmosphere is rotating

$$\underline{v}_{rel} = \begin{bmatrix} \dot{x} + \omega_{\oplus} y \\ \dot{y} - \omega_{\oplus} x \\ \dot{z} \end{bmatrix}$$

$$\omega_{Earth} = [0 \quad 0 \quad \omega_{\oplus}]$$

point of intersection

$$\begin{bmatrix} x \\ y \\ z \end{bmatrix} = R \begin{bmatrix} \cos \theta \\ \sin \theta \cos i \\ \sin \theta \sin i \end{bmatrix}$$

$$\underline{v} = \underline{\omega} \wedge \begin{Bmatrix} x \\ y \\ z \end{Bmatrix} = \begin{bmatrix} \hat{i} & \hat{j} & \hat{k} \\ 0 & 0 & \omega_E \\ x & y & z \end{bmatrix} = \begin{Bmatrix} -\omega_E y \\ +\omega_E x \\ 0 \end{Bmatrix}$$

$$R = \frac{a(1 - e^2)}{1 + e \cos \theta}$$

$$\dot{\theta} = \frac{n(1 + e \cos \theta)^2}{(1 - e^2)^{3/2}}$$

$$\omega_{\oplus} = 0.000072921 \text{ rad/sec}$$

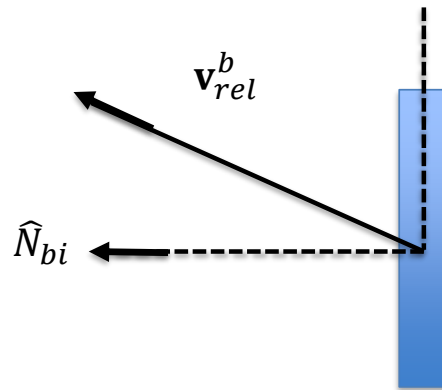


# Relative velocity in body fixed coordinates

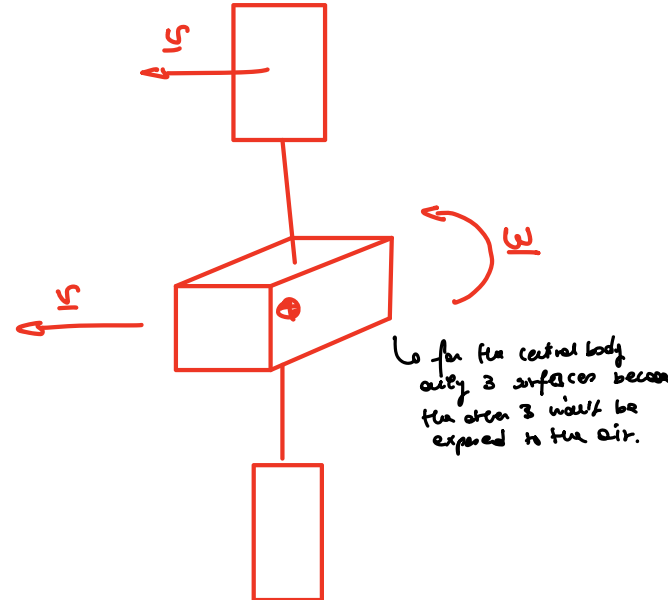
$$\mathbf{v}_{rel}^b = A_{B/N} \mathbf{v}_{rel}$$

body ↙

$$\mathbf{v}_{rel} = \begin{bmatrix} \dot{x} + \omega \oplus y \\ \dot{y} - \omega \oplus x \\ \dot{z} \end{bmatrix}$$



$$A_{cross} = A_i (\hat{N}_{bi} \cdot \frac{\mathbf{v}_{rel}^b}{\|\mathbf{v}_{rel}^b\|})$$



# Torque due to atmospheric drag

The aerodynamic force acting on a flat surface is defined by:

$$\bar{F}_i = -\frac{1}{2}\rho C_D v_{rel}^2 \frac{\mathbf{v}_{rel}^b}{\|\mathbf{v}_{rel}^b\|} (\hat{N}_{bi} \cdot \frac{\mathbf{v}_{rel}^b}{\|\mathbf{v}_{rel}^b\|}) A_i$$

number of surfaces exposed to the Air

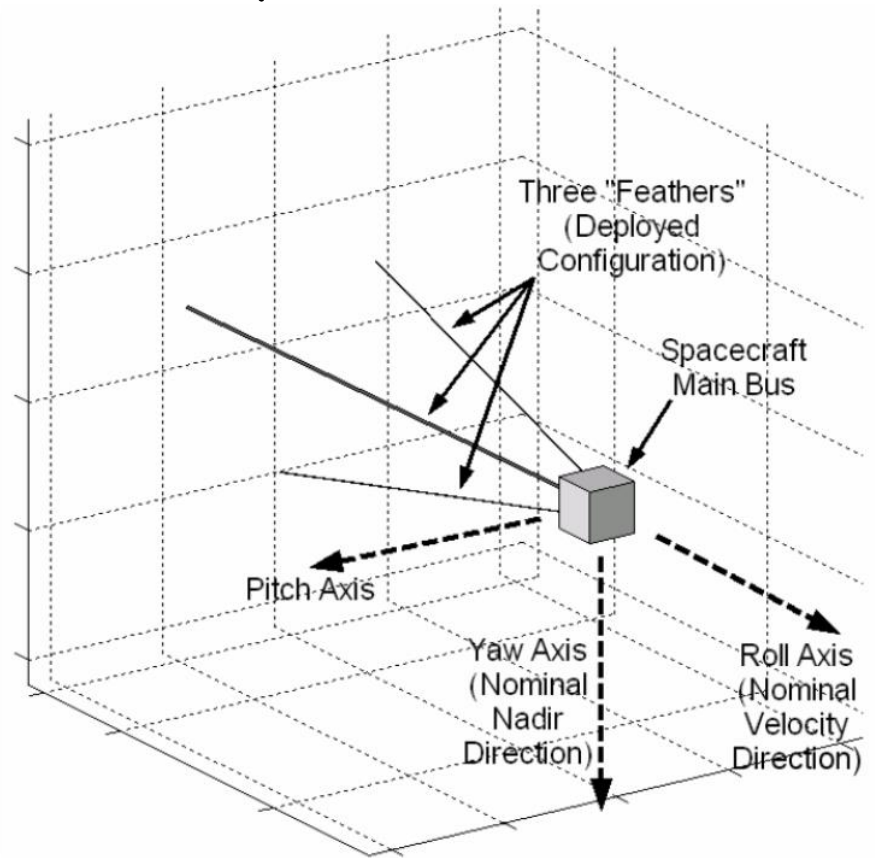
$$T_{aero} = - \sum_{i=1}^{\overset{n}{\circlearrowleft} \tilde{r}_i} \tilde{r}_i \times \frac{1}{2} \rho C_D v_{rel}^2 \frac{\mathbf{v}_{rel}^b}{\|\mathbf{v}_{rel}^b\|} (\hat{N}_{bi} \cdot \frac{\mathbf{v}_{rel}^b}{\|\mathbf{v}_{rel}^b\|}) A_i$$

Total torque on a rigid body

$$T_{aero} = \begin{cases} -\frac{1}{2}\rho C_D v_{rel}^2 \sum_{i=1}^n \tilde{r}_i \times \frac{\mathbf{v}_{rel}^b}{\|\mathbf{v}_{rel}^b\|} \sum_{i=1}^n (\hat{N}_{bi} \cdot \frac{\mathbf{v}_{rel}^b}{\|\mathbf{v}_{rel}^b\|}) A_i & \hat{N}_{bi} \cdot \frac{\mathbf{v}_{rel}^b}{\|\mathbf{v}_{rel}^b\|} > 0 \\ 0 & \hat{N}_{bi} \cdot \frac{\mathbf{v}_{rel}^b}{\|\mathbf{v}_{rel}^b\|} < 0 \end{cases}$$

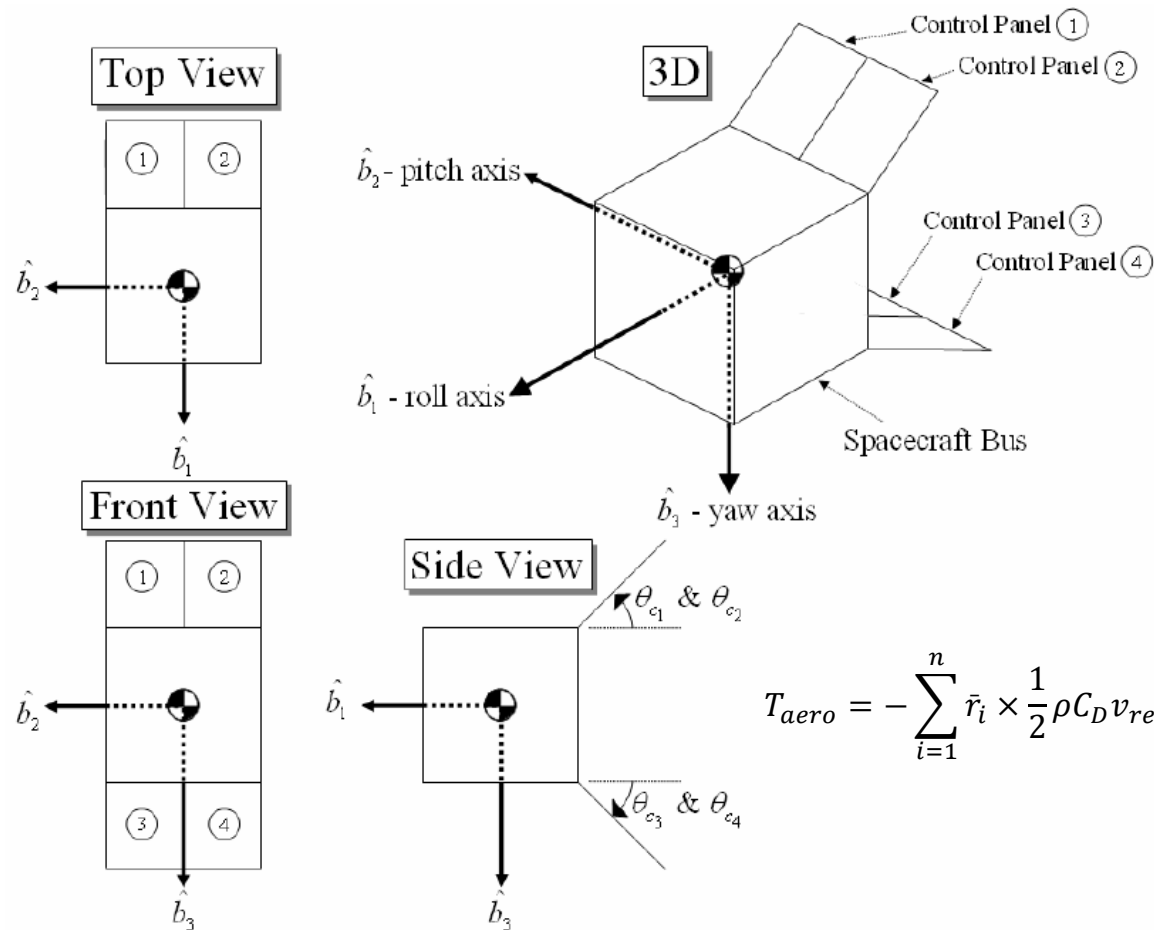
# Shuttlecock concept

*It is generating a passive stability but introduce some drag so we will have some orbit decay.*





# CubeSat – air drag control



$$T_{aero} = - \sum_{i=1}^n \bar{r}_i \times \frac{1}{2} \rho C_D v_{rel}^2 \frac{\mathbf{v}_{rel}^b}{\|\mathbf{v}_{rel}^b\|} \sum_{i=1}^n (\hat{N}_{bi} \cdot \frac{\mathbf{v}_{rel}^b}{\|\mathbf{v}_{rel}^b\|}) A_i$$

# Force due to solar radiation pressure

for orbit around center the distance from the Sun is constant

Depends on the relative position of the center body with to the Sun!

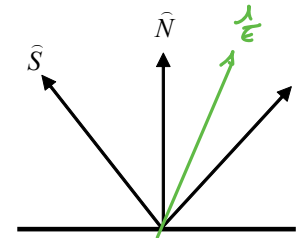
center body = our orbit  
→ But it is general we can compute always the radiation emitted from the center body. It depends on the temperature

Altitude from the surface of the Earth (Km)	Direct solar radiation (W/m <sup>2</sup> )	Radiation reflected by the Earth (W/m <sup>2</sup> )	Earth radiation (W/m <sup>2</sup> )
500	1358	600	150
1000	1358	500	117
2000	1358	300	89
4000	1358	180	62
8000	1358	75	38
15000	1358	30	14
30000	1358	12	3
60000	1358	7	2

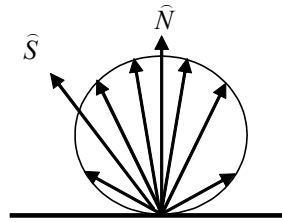
The average pressure due to radiation can be evaluated as

$$P = \frac{F_e}{c}$$

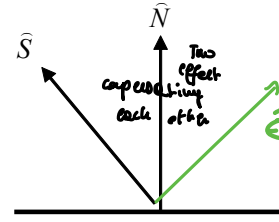
where  $c$  is the speed of light and  $F_e$  is the power per unit surface.



Specular reflection  
perfect mirror  
direction of the force is  $\hat{N}$



Diffuse reflection  
we decompose the vector



Absorption

More  
it is important about the direction of the radiation from the Earth and not only its module. for at least 200 km it should be considered

$$\rho_s + \rho_d + \rho_a = 1$$

→ The equivalent force has the same direction of  $\hat{N}$ .

We can use a combination of these effects and the total of the radiation that is reflected diffuse and absorb should be 100% of the total incoming radiation.



# Force due to solar radiation pressure (from the Sun)

Force on a flat panel

$$\bar{F}_i = -PA \left[ \rho_a (\hat{S} \cdot \hat{N}) \hat{S} + 2\rho_s (\hat{S} \cdot \hat{N})^2 \hat{N} + \rho_d (\hat{S} \cdot \hat{N}) \left( \hat{S} + \frac{2}{3} \hat{N} \right) \right]$$

$$\bar{F}_i = -PA (\hat{S} \cdot \hat{N}) \left[ (1 - \rho_s) \hat{S} + (2\rho_s (\hat{S} \cdot \hat{N}) + \frac{2}{3} \rho_d) \hat{N} \right]$$

$$\bar{F}_i = -PA_i (\hat{S}_b \cdot \hat{N}_{bi}) \left[ (1 - \rho_s) \hat{S}_b + (2\rho_s (\hat{S}_b \cdot \hat{N}_{bi}) + \frac{2}{3} \rho_d) \hat{N}_{bi} \right]$$



# Torque due to solar radiation pressure

$$\bar{F}_i = -PA_i(\hat{S}_b \cdot \hat{N}_{bi}) \left[ (1 - \rho_s)\hat{S}_b + (2\rho_s(\hat{S}_b \cdot \hat{N}_{bi}) + \frac{2}{3}\rho_d)\hat{N}_{bi} \right]$$

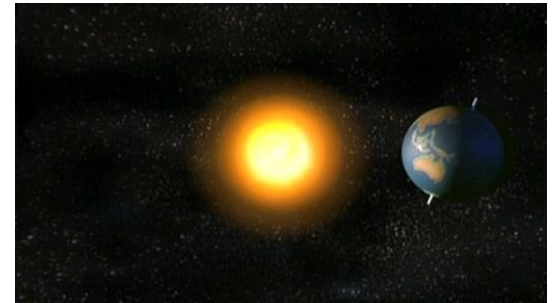
## Modelling the Sun in body coordinate

All the computation must be done in the principal axis frame because the torque in the principal axis frame will be the input of the system.

$$\hat{S}_b = A_{B/N} \hat{S}_i$$

### NOTE

there is a difference between the maximum force and the maximum torque because the geometry have a really important role



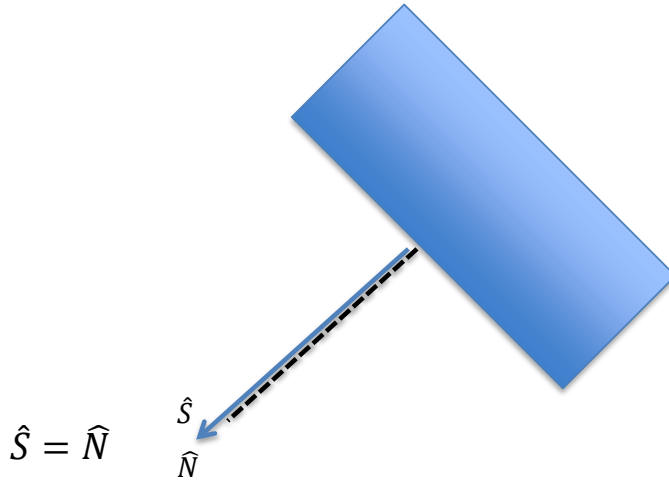
For a rigid body we have

$$T_{SRP} = \sum_{i=1}^n \bar{r}_i \times \bar{F}_i$$

$$T_{SRP} = \begin{cases} \sum_{i=1}^n \bar{r}_i \times \bar{F}_i & \hat{S}_b \cdot \hat{N}_b > 0 \\ 0 & \hat{S}_b \cdot \hat{N}_b < 0 \end{cases}$$

## Maximum SRP Force and corresponding torque

$$\bar{F}_i = -PA(\hat{S} \cdot \hat{N}) \left[ (1 - \rho_s)\hat{S} + (2\rho_s(\hat{S} \cdot \hat{N}) + \frac{2}{3}\rho_d)\hat{N} \right]$$



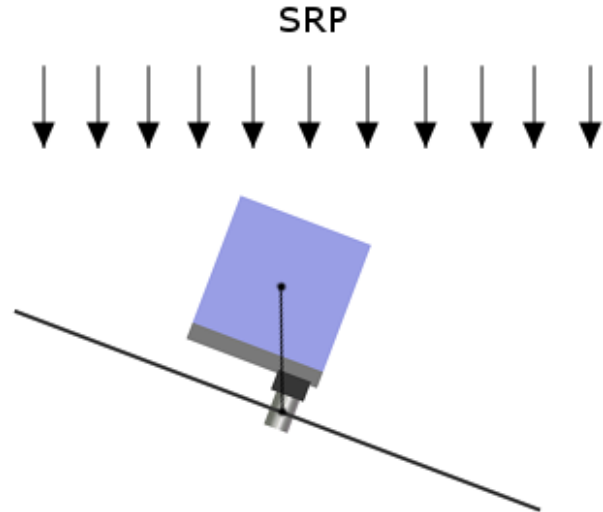
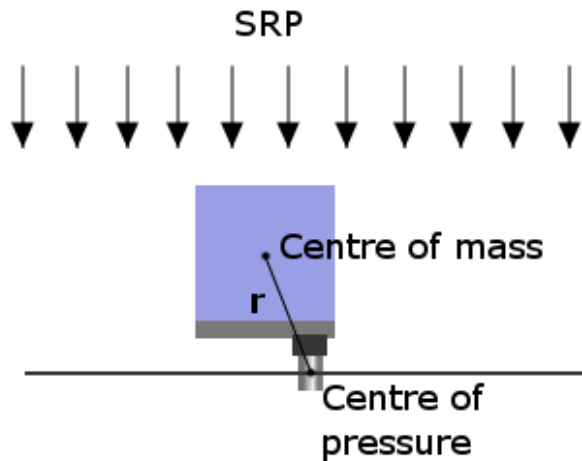
then

$$\bar{F}_i = -PA \left[ (1 + \rho_s) + \frac{2}{3}\rho_d \right] \hat{N}$$

$$T_{SRP} = \sum_{i=1}^n \bar{r}_i \times \bar{F}_i$$

# Exploiting SRP for attitude control - CubeSail

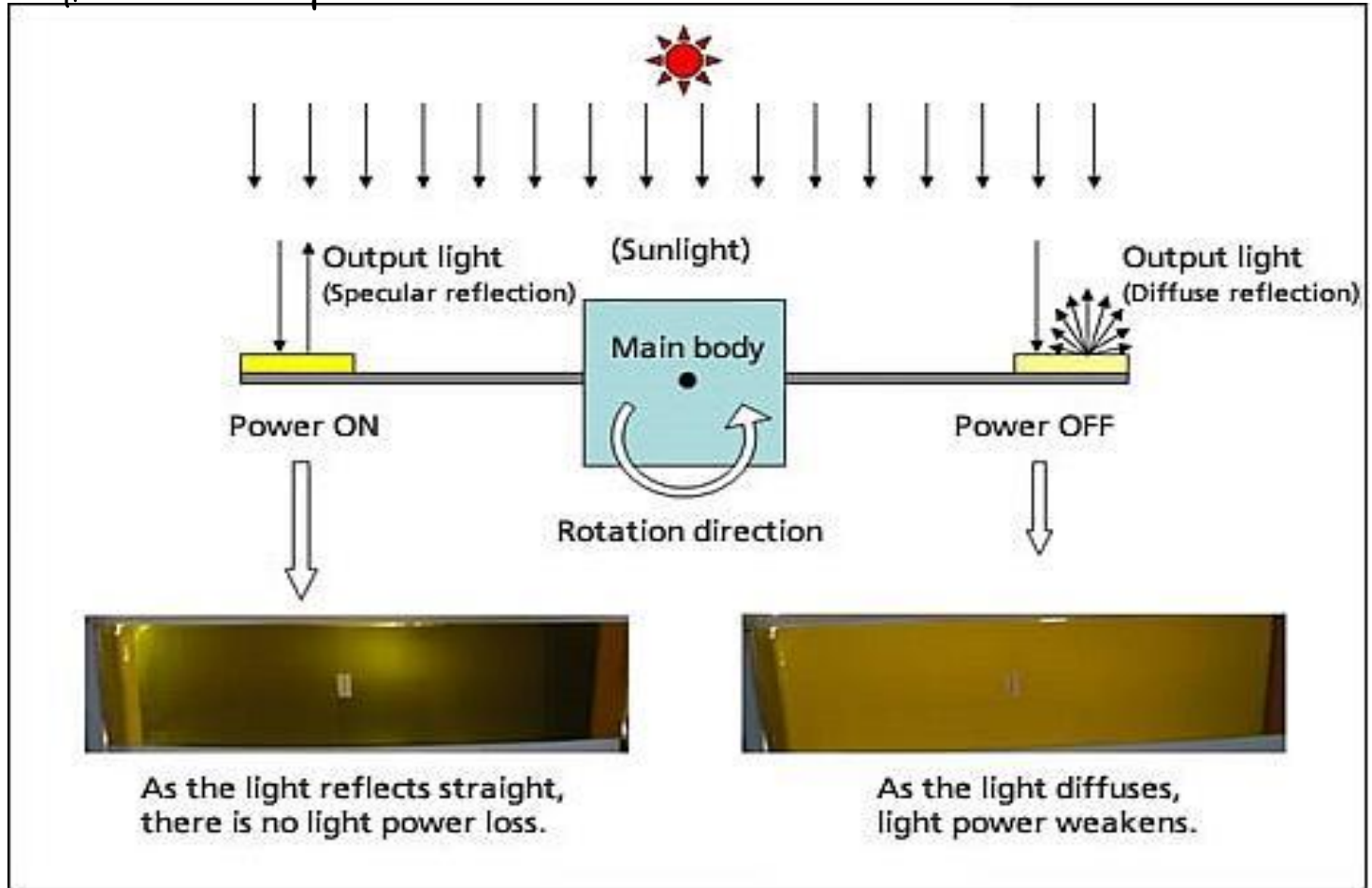
Using a sail is a way to generate some torque.



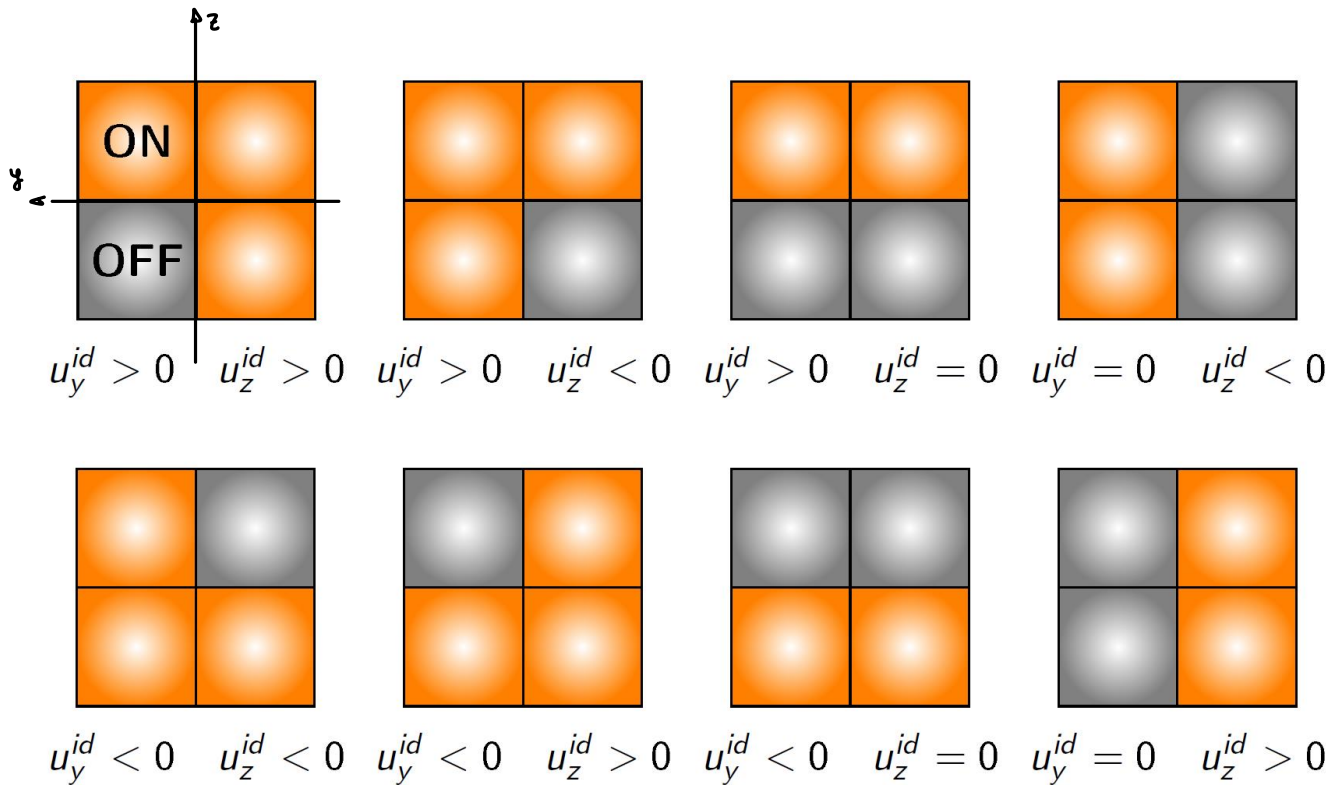
$$T_{srp} = \sum_{i=1}^n \vec{r}_i \times \vec{F}_i$$

# IKAROS – reflectivity control devices

*Different panel can give a new torque.*

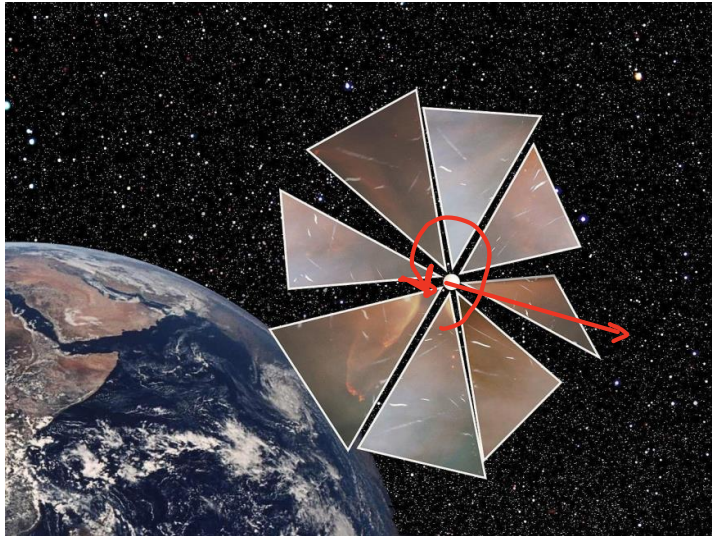


# Basic control logic *→ we are interested in seeing which loops for the Torque*





# Future concepts for attitude control with SRP



$$\bar{F}_i = -PA_i(\hat{S}_b \cdot \hat{N}_{bi}) \left[ (1 - \rho_s)\hat{S}_b + (2\rho_s(\hat{S}_b \cdot \hat{N}_{bi}) + \frac{2}{3}\rho_d)\hat{N}_{bi} \right]$$