- Ecosystem carbon balance in the Hawaiian Islands under
- different scenarios of future climate and land use change

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18 Abstract

The State of Hawai'i passed legislation to be carbon neutral by 2045, a goal that will partly depend on carbon sequestration by terrestrial ecosystems. However, there is considerable uncertainty surrounding the future direction and magnitude of the land carbon sink in the Hawaiian Islands. We used the Land Use and Carbon Scenario Simulator (LUCAS), a spatially explicit stochastic simulation model that integrates landscape change and carbon gain-loss, to assess how projected future changes in climate and land use will influence ecosystem carbon balance in the Hawaiian Islands under all combinations of two radiative forcing scenarios (RCPs 4.5 and 8.5) and two land use scenarios (low and high) over a 90-year timespan from 2010-2100. We used simulation modeling to assess how projected future changes in climate and land use will influence ecosystem carbon balance 27 in the Hawaiian Islands under all combinations of two radiative forcing scenarios (RCP 4.5 and RCP 8.5) and two land use scenarios (low and high) over a 90-year timespan from 2010-2100. Collectively, terrestrial ecosystems of the Hawaiian Islands acted as a net carbon sink under low radiative forcing (RCP 4.5) for the entire 90-year simulation period, with low land use change further 31 enhancing carbon sink strength. In contrast, Hawaiian terrestrial ecosystems transitioned from a net 32 sink to a net source of CO₂ to the atmosphere under high radiative forcing (RCP 8.5), with high land 33 use accelerating this transition and exacerbating net carbon loss. A sensitivity test of the CO₂ fertilization effect on plant productivity revealed it to be a major source of uncertainty in projections of ecosystem carbon balance. Greater mechanistic understanding of plant productivity responses to rising atmospheric CO₂ will be essential to realistically constrain simulation models used to evaluate the effectiveness of ecosystem-based climate mitigation strategies.

39 1. Introduction

- Terrestrial ecosystems are a major sink for atmospheric carbon dioxide (CO_2), removing $\sim 30\%$ of
- human emissions on an annual basis and reducing the rate of increase in atmospheric CO₂ (Keenan
- and Williams 2018, Friedlingstein *et al* 2019). There is increasing recognition among policymakers

that natural and agricultural ecosystems can contribute to climate mitigation, which has given rise to the popularity of "natural climate solutions" (Cameron et al 2017). Defined as conservation and land management efforts aimed at enhancing ecosystem carbon storage (Griscom et al 2017), natural climate solutions are appealing because they are seen as cost-efficient and readily available (Galarraga et al 2017, Cameron et al 2017, Fargione et al 2018). However, effective implementation is complicated by the uncertainty surrounding the future direction and magnitude of the land carbon sink, especially at the regional scale. Despite this uncertainty, evidence indicates that both interannual and long-term variability in carbon uptake by land ecosystems is driven primarily by fluctuations in climate, land use, and land cover change (Ahlström et al 2015, Prestele et al 2017, Friedlingstein et al 2019). Incorporating the interactive effects of land use and climate into spatially explicit future projections of ecosystem carbon balance could therefore provide a reference point to evaluate the 53 effectiveness of land-based mitigation. Although a complex challenge, a growing number of sub-national jurisdictions plan to incorporate land-based mitigation strategies into their emissions reduction efforts. These jurisdictions would benefit from understanding how future land use and climate-biosphere feedbacks will affect ecosystem carbon balance in their respective regions (Sleeter et al 2019). The State of Hawai'i exemplifies the challenges associated with projecting the interactive effects of future climate and land use change on ecosystem carbon balance at a regional scale. Hawai'i was the first U.S. state to enact legislation committing to full carbon neutrality, requiring the state to account for and offset all of its greenhouse gas emissions by 2045 (State of Hawai'i Acts 15 and 16). This legislation emphasizes the mitigation potential of natural ecosystems as a key component to emissions reduction, necessitating baseline estimates and future projections of land carbon sink strength. However, Hawai'i's challenging terrain complicates these assessment efforts. The main Hawaiian Islands are a complex mosaic of natural and human-dominated landscapes overlain by steep climate gradients across relatively short distances, with mean annual temperature ranging from ~4-24° C (Giambelluca et al 2014) and mean annual rainfall ranging from ~200-10,200 mm (Giambelluca et al 2013). Temperatures have risen rapidly in the Hawaiian Islands since the mid

1970s (Giambelluca et al 2008, McKenzie et al 2019) and a long-term drying trend has persisted since the early 1920s (Frazier and Giambelluca 2017), resulting in reduced forest biomass and productivity (Barbosa and Asner 2017). These same drying and warming trends have increased the frequency and intensity of wildland fire (Trauernicht et al 2015, Trauernicht 2019) with predictable negative effects on ecosystem carbon balance (Selmants et al 2017). Ecosystem carbon stocks across the main Hawaiian Islands have also been strongly influenced by the legacy of past land use change (Osher et al 2003, Asner et al 2011). Thousands of hectares were deforested beginning in the late 19th century to clear land for sugar plantations and cattle pasture (Cuddihy and Stone 1990). Since 77 the mid-20th century, much of this agricultural land has been steadily converted to urban areas, commercial forestry plantations, or simply abandoned and colonized by non-native grass species (Suryanata 2009, Perroy et al 2016). Although these past trends surely inform the future impact of 80 climate and land use change on ecosystem carbon balance, high spatial and temporal heterogeneity 81 complicates realistic projection efforts. To date only one study has attempted to integrate land use, climate, and natural disturbances into future projections of Hawaiian ecosystem carbon balance, with projections limited to the mid-21st century under a single land use change scenario and moderate radiative forcing (SRES A1B, equivalent to RCP 6; Selmants et al 2017). We used a stochastic, spatially explicit simulation model to estimate ecosystem carbon balance for Hawai'i's natural and agricultural lands on an annual basis for the period 2010–2100 under a range of assumptions about future climate, land use, land cover, disturbance, and global CO2 emissions (Daniel et al 2016, 2018, Sleeter et al 2019). We developed four unique scenarios that explored different pathways, or future trajectories, of land use change and climate change. These four scenarios represent all combinations of two land use change pathways (low and high) and two radiative forcing pathways (representative concentration pathway [RCP] 4.5 and RCP 8.5). In addition to these four scenarios, we conducted a separate series of simulations to examine how ecosystem carbon balance estimates vary in response to different levels of a CO₂ fertilization effect (CFE) on net primary productivity (NPP; Sleeter et al 2019). Our goals were to estimate changes in Hawaiian ecosystem carbon balance and their uncertainties under a range of plausible future scenarios, quantify the

relative impact of major controlling processes such as land use change, disturbance, and climate change, and assess the sensitivity of model estimates to the introduction of a CFE on NPP.

2. Methods

We used the Land Use and Carbon Scenario Simulator (LUCAS), an integrated landscape change and carbon gain-loss model, to project changes in ecosystem carbon balance for the seven main Hawaiian 101 Islands. The landscape change portion of LUCAS is a state-and-transition model, where 'state' is the 102 pre-defined information specific to each simulation cell (e.g., land cover class, climate zone) and 103 'transition' is the change in state of each simulation cell over time (Daniel et al 2016). Landscape 104 change occurs by applying a Monte Carlo approach to track the state type and age of each simulation 105 cell in response to a pre-determined set of transition pathways (Daniel et al 2016). Transition 106 pathways define which state types can be converted to any other state type. The carbon gain-loss 107 portion tracks carbon stocks within each simulation cell over time as continuous state variables, along 108 with a pre-defined set of continuous flows specifying rates of change in stock levels over time (Daniel 109 et al 2018, Sleeter et al 2019). We parameterized the Hawai'i LUCAS model to estimate annual 110 changes in carbon stocks and fluxes in response to land use, land use change, wildland fire, and 111 long-term climate variability for the time period 2010-2100. 112

113 2.1 Study area

The spatial extent of this study was the terrestrial portion of the seven main Hawaiian Islands (figure 1), a total land area of 16,554 km². We subdivided this landscape into a grid of 264,870 simulation cells, each of which was 250 x 250 m in size. Each simulation cell was assigned to one of 210 possible state types based on the unique combination of three moisture zones (dry, mesic, and wet; supplemental figure 1), seven islands, and ten discrete land cover classes (figure 1).



Figure 1: Land cover classification of the seven main Hawaiian Islands, adapted from Jacobi *et al.* (2017). Agriculture in this map combines herbaceous and woody crops, but these two crop types are treated as separate land cover classes in the simulation model. Water and Wetland land cover classes are not shown.

119 2.2 States and transitions

We developed two land use scenarios (low and high) with transition pathways modified from Daniel 120 et al (2016). Transition pathways were pre-defined to represent urbanization, agricultural contraction, 121 agricultural expansion, harvesting of tree plantations, and wildfire. Agriculture, forest, grassland, 122 shrubland, and tree plantation land cover classes each had multiple transition pathways, i.e., they could each be converted into multiple other land cover classes. The barren land cover class could only transition to developed, so its sole transition pathway was urbanization. Although most land 125 cover classes had an urbanization transition pathway, there was no transition pathway out of an urbanized (developed) state. Water and wetland land cover classes remained static throughout the 127 simulation period. 128 Transition targets set the area to be transitioned over time. These targets were based on historical trends of land use change in the Hawaiian Islands from 1992-2011 (NOAA 2020) and on population projections for the State of Hawai'i (Kim and Bai 2018). For the high land use scenario, transition rates for each timestep and Monte Carlo realization were sampled from uniform distributions bounded 132 by the median and maximum historical rates of agricultural contraction, agricultural expansion, and 133 urbanization for each island. For the low land use scenario, rates of agricultural contraction and 134 expansion were sampled from uniform distributions bounded by zero and the minimum historical 135 rates for each island. Urbanization rates in the low land use scenario were based on island-level 136 population estimates and projections at five year intervals from 2010-2045 (Kim and Bai 2018). We 137 converted population projections into urbanization transition targets following Sleeter et al (2017) by 138 calculating population density for each island and then projecting future developed area based on the 139 five-year incremental change in island population. The spatial extent of agricultural contraction, 140 agricultural expansion, and urbanization was constrained in both land use scenarios based on existing 141 zoning maps (Daniel et al 2016). Tree plantation forestry in the State of Hawai'i is primarily 142 short-rotation (5-7 years) Eucalyptus spp., harvested at a rate of approximately 2 km² v⁻¹ over the 143 past decade. Transition targets for tree plantation harvest were set at ~75% of recent historical rates in

the high land use scenario and ~40% of recent historical rates in the low land use scenario (Daniel et al 2016). In both land use scenarios, approximately 60% of tree plantation harvests were replacement harvests resulting in conversion to agriculture. The remaining 40% were rotation harvests replanted to Eucalyptus spp. No land cover class could be converted to tree plantation in either land use 148 scenario, consistent with recent historical trends of stable or declining area of tree plantations. 149 The wildfire transition sub-model was modified from Daniel et al (2016) by incorporating a new 21-year historical wildfire spatial database of the Hawaiian Islands (supplemental figure 2). We used 151 this new spatial database to calculate historical wildfire size distribution and ignition probabilities for 152 each unique combination of moisture zone (supplemental figure 1), island, and land cover class 153 (figure 1) for the years 1999-2019. Starting in 2020, the number and size of fires was randomly 154 drawn from one of these historical year-sets for each timestep and Monte Carlo realization, using 155 burn severity probabilities from Selmants et al (2017). Wildfire in the low land use scenario was 156 sampled from the subset of historical fire years at or below the median area burned statewide from 157 1999-2019. The high land use scenario sampled from historical fire years above the median area 158 burned over the same 21-year period (supplemental figure 2a). The vast majority of wildland fire in 159 Hawai'i is human-caused and initiates in non-native grasslands and shrublands (Trauernicht et al 160 2015), but can spread into nearby forest areas. Moderate to high severity fires that spread into mesic 161 and wet forests can convert these areas into grasslands about half the time. However, moderate to 162 high severity fires in mesic and wet forests account for <5% of the total annual area burned, on 163 average (Selmants et al 2017). 164

2.3 Carbon stocks and flows

The fate of carbon stocks was tracked for each simulation cell based on a suite of carbon flows (i.e., carbon fluxes) specifying the rates of change in these carbon stocks over time (Daniel *et al* 2018, Sleeter *et al* 2019). We defined carbon stocks as continuous state variables for each simulation cell, including live biomass, standing dead wood, down dead wood, litter, and soil organic carbon. We

also included and tracked carbon in atmospheric, aquatic, and harvest product pools to enforce carbon mass balance (Daniel et al 2018). To transfer carbon between stocks, we defined baseline carbon flows as continuous variables resulting from growth, mortality, deadfall, woody decay, litter decomposition, and leaching (which includes runoff). We also defined carbon flows resulting from 173 land use, land use change, and wildfire (Selmants et al 2017, Daniel et al 2018). 174 Initial carbon stocks and baseline carbon flows were estimated based on the state type and age of each simulation cell using a lookup table derived from the Integrated Biosphere Simulator (IBIS; Foley et 176 al 1996, Liu et al 2020), a process-based dynamic global vegetation model. We initiated IBIS with minimal vegetation and simulated forward for 110 years using 30-year climate normals for the 178 Hawaiian Islands (Giambelluca et al 2013, 2014). We calibrated IBIS carbon stocks with statewide 179 gridded datasets of soil organic carbon (Soil Survey Staff 2016) and forest aboveground live biomass 180 (Asner et al 2016). We also calibrated gross photosynthesis in IBIS using a Hawai'i-specific gridded 181 dataset derived from MODIS satellite imagery (Kimball et al 2017). 182 Net primary production for each simulation cell was calculated as the mean IBIS-derived value for 183 each combination of moisture zone and land cover class (supplemental Table 1) adjusted with a 184 spatially explicit stationary growth multiplier to reflect local variation driven by microclimate 185 (supplemental Figure 4; Sleeter *et al* 2019). We calculated this spatial growth multiplier as the NPP 186 anomaly for each simulation cell relative to mean NPP values for each combination of moisture zone 187 and land cover class based on empirical relationships between total annual NPP and mean annual 188 rainfall or temperature (Schuur 2003, Del Grosso et al 2008) using Hawai i-specific climate data 180 (Giambelluca et al 2013, 2014). Soil carbon flux to the atmosphere (R_h) and aquatic soil carbon 190 losses (leaching and overland flow) were estimated as the ratio of the IBIS-derived flux for each 191 combination of moisture zone and state type to the microclimate-adjusted NPP value for each 192 simulation cell. All other carbon flow rates were estimated as the ratio of the mean IBIS-derived flux 193 for each combination of moisture zone and state type to the size of the originating carbon stock at each age (Sleeter et al 2018, Daniel et al 2018).

Climate change impacts on carbon flows were represented by temporal growth and decay multipliers applied to each simulation cell based on mid-century (2049-2069) and end-of-century (2070-2099) 197 statistically downscaled CMIP5 temperature and rainfall projections for the Hawaiian Islands under 198 each of two radiative forcing scenarios (RCP 4.5 and RCP 8.5; Elison Timm et al 2015, Elison Timm 199 2017). The impact of future changes in rainfall and temperature on NPP were represented by annual 200 growth multipliers calculated using empirical NPP models (Schuur 2003, Del Grosso et al 2008) and 201 climate model projections of temperature and rainfall for each radiative forcing scenario. The effect 202 of future warming on turnover rates of dead organic matter were represented by temporal decay 203 multipliers calculated using \mathbf{Q}_{10} temperature coefficients based on climate model temperature 204 projections for each radiative forcing scenario. The Q₁₀ temperature coefficients represented the 205 proportional change in detrital carbon turnover due to a 10° C change in mean annual temperature. 206 We applied a Q₁₀ of 2.0 for wood and soil organic matter decay flows (Kurz et al 2009, Sleeter et al 207 2019) and a Q₁₀ of 2.17 for litter decay flows (Bothwell et al 2014). Transition-triggered carbon 208 flows resulting from disturbances associated with land use change, timber harvesting, and wildfire 209 were based on values from Don et al (2011), Selmants et al (2017), and Daniel et al (2018). 210

2.4 CO₂ fertilization effect

Increasing atmospheric CO₂ concentrations stimulate leaf-level photosynthesis, potentially increasing 212 NPP as well (Franks et al 2013). However, the magnitude and persistence of this effect is highly 213 uncertain, particularly across a range of climatic conditions and over long time spans (Walker et al 214 2020). Following Sleeter et al (2019), we developed a separate set of scenarios designed to test the 215 sensitivity of LUCAS model projections of ecosystem carbon balance to different rates of a CO₂ 216 fertilization effect (CFE). We incorporated a CFE multiplier for NPP that represented the percent 217 increase in NPP for every 100 ppm increase in atmospheric CO₂ concentration under the high land 218 use and high radiative forcing (RCP 8.5) scenario. We tested five CFE levels ranging from 5% to 219 15%, which is within the range of CFEs observed in free air CO₂ enrichment (FACE) experiments. For all levels, we assumed CFEs reached saturation at an atmospheric CO₂ concentration of 600 ppm, with no further stimulation of NPP despite a continued increase in CO₂ concentration to 930 ppm by 2100. This 600 ppm threshold generally coincides with the upper limit from FACE experiments and is reached by the year 2060 under RCP 8.5.

2.5 Scenario simulations and analysis

Each of the four unique scenarios were run for 90 years at an annual timestep and repeated for 30 Monte Carlo realizations, using initial conditions corresponding to the year 2010. All simulations were performed within the SyncroSim (version 2.2.4) software framework with ST-Sim (version 3.2.13) and SF (version 3.2.10) add-on modules (Daniel et al 2016, 2018). Model inputs and outputs 229 were prepared with the R statistical computing platform (R Core Team 2019) using the tidyverse 230 (Wickham et al 2019), raster (Hijmans 2020), and rsyncrosim (Daniel et al 2020) packages. Carbon 231 stocks and fluxes for the seven main Hawaiian Islands were calculated for each scenario by summing 232 within each Monte Carlo realization on an annual basis and then calculating annual means as well as 233 the annual upper and lower limits of the 30 Monte Carlo realizations. Carbon balance for the seven 234 main Hawaiian Islands was calculated on annual basis for each scenario and Monte Carlo realization 235 as net biome productivity (NBP), which was equal to annual carbon input in the form of NPP minus 236 the annual sum of all carbon losses from terrestrial ecosystems, including heterotrophic respiration 237 (R_h) from litter and soil, carbon fluxes to the atmosphere triggered by land use and land use change, 238 wildfire emissions, and aquatic carbon losses through leaching and overland flow. Positive NBP 239 values indicated ecosystems of the seven main Hawaiian Islands were acting as a net sink for 240 atmospheric CO2, while negative NBP values indicated that these ecosystems were acting as a net 241 carbon source to the atmosphere (Chapin et al 2006).

243 3. Results

3.1 Carbon stocks and fluxes

Terrestrial ecosystems of the seven main Hawaiian Islands stored an estimated 316 Tg of carbon at 245 the beginning of the simulation period in 2010 (figure 2a), with 58% in soil organic matter, 22% in 246 living biomass, and 20% in surface dead organic matter (litter and dead wood; figure 2b). Ecosystems 247 accumulated carbon in all scenarios but at different rates, with trajectories shaped primarily by 248 climate change and to a lesser extent by land use change. The highest and most consistent projected 249 accumulation of ecosystem carbon occured under the combination of low radiative forcing and low 250 land use change, yielding a ~15% increase in ecosystem carbon to an average of 363 Tg by 2100 251 (figure 2a). In contrast, high radiative forcing and high land use change resulted in the lowest 252 ecosystem carbon gain, reaching a peak of ~332 Tg in 2063 and a decline to 327 Tg in 2100, resulting 253 in a net increase of only 3% by the end of the simulation period (figure 2a). Ecosystem carbon 254 accumulation was driven exclusively by increasing soil organic carbon across all four scenarios, all 255 other stocks declined over time (figure 2b).

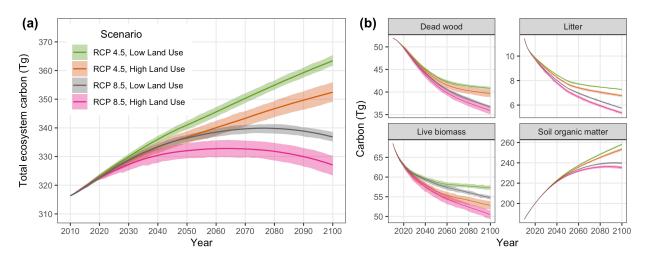


Figure 2: Projected changes in total ecosystem carbon storage (a) and individual carbon stocks (b) for the seven main Hawaiian Islands. Solid lines indicate the mean of 30 Monte Carlo realizations for each scenario, with shaded areas indicating the minimum and maximum range of Monte Carlo realizations.

Net primary production (NPP) for the seven main Hawaiian Islands declined across all four scenarios,

driven primarily by climate change and to a lesser extent by land use change (Fig. 3). The combination of high radiative forcing (RCP 8.5) and high land use change led to the steepest decline 259 in NPP over time, driven by intense long-term drying on the leeward sides of islands under RCP 8.5 (supplemental figure 4) and sustained losses of forest and shrubland land area in the high land use 261 scenario (supplemental figure 5). In contrast, climate change led to increased heterotrophic 262 respiration (R_h) over time, such that more intense warming under RCP 8.5 (supplemental figure 4) 263 resulted in R_h being $\sim 3\%$ higher by 2100 than under RCP 4.5 (figure 3). Heterotrophic respiration 264 declined substantially over time in the high land use scenario (figure 3) because of long-term 265 reductions in forest and shrubland land area (supplemental figure 5), similar to trends in NPP. 266 Transition-triggered carbon fluxes to the atmosphere from land use, land use change, and wildfire 267 were largely independent of changes in climate, stabilizing by mid-century at an average of ~0.4 Tg 268 y^{-1} in the high land use scenario and ~ 0.2 Tg y^{-1} in the low land use scenario (figure 3). Uncertainty 269 around transition-triggered carbon fluxes were higher in the high land use scenario, driven primarily 270 by greater variability in wildland fire probabilities. 271

72 3.2 Ecosystem carbon balance

Net biome productivity (NBP) averaged approximately 0.6 Tg C y⁻¹ at the start of the simulation 273 period and declined over time in all four scenarios (figure 4). On average, terrestrial ecosystems of the 274 seven main Hawaiian Islands collectively acted as a net carbon sink throughout the simulation period 275 under the RCP 4.5 radiative forcing scenario, but carbon sink strength was ~40% lower in the high 276 land use scenario compared to the low land use scenario by the end of the simulation period (figure 4). 277 In contrast, ecosystems of the Hawaiian Islands acted as a net carbon source to the atmosphere toward 278 the latter half of the simulation period under RCP 8.5, with the transition from sink to source occurring 279 15 years earlier on average in the high land use scenario than in the low land use scenario (figure 4). 280 The high land use scenario under RCP 8.5 represented a ~40% larger net source of carbon to the 281 atmosphere by the year 2100 than the low-land use scenario under the same radiative forcing. Over 282 the entire simulation period, both global emissions reductions and local avoided land conversion

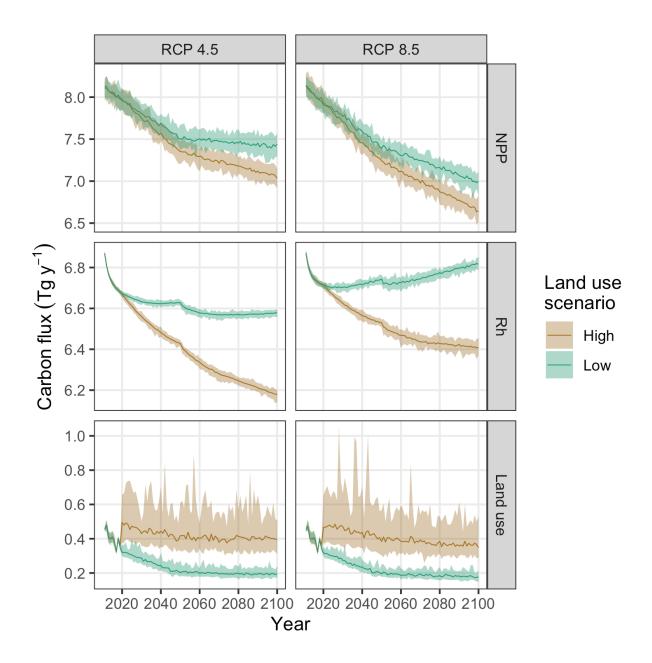


Figure 3: Projected changes in net primary production (NPP), heterotrophic respiration (Rh) and carbon fluxes induced by land use and land use change for the seven main Hawaiian Islands. Solid lines indicate the mean of 30 Monte Carlo realizations for each scenario, with shaded areas indicating the minimum and maximum range of Monte Carlo realizations.

resulted in substantial increases in cumulative NBP (figure 5). However, switching from RCP 8.5 to RCP 4.5 increased cumulative NBP in the Hawaiian Islands more than twice as much as reducing emissions from local land use change and wildfire disturbance (figure 5). Switching from RCP 8.5 to RCP 4.5 under the low land use scenario yielded the greatest cumulative increase in NBP, resulting in a median gain of 26.5 Tg of carbon over the entire 90-year simulation period.

289 3.3 CO₂ fertilization effect

Projected estimates of both total ecosystem carbon storage and ecosystem carbon balance were highly sensitive to differing rates of a CFE on plant productivity. Under the high radiative forcing (RCP 8.5) 291 and high land use scenario, the inclusion of a CFE ranging from 5-15% led to ~33-98 Tg of additional 292 carbon storage in ecosystems by the end of the century, a ~10-30% increase (figure 6a). Compared to 293 the reference scenario (0% CFE), a 5% CFE was sufficient to transform Hawaiian Island ecosystems 294 from a net carbon source to the atmosphere during the latter half of the 21st century (figure 4b) to a 295 net carbon sink for the entire simulation period (figure 6b), completely offsetting all other carbon 296 losses induced by high radiative forcing and high land use. Net carbon sink strength was further 297 enhanced at higher CFE rates, with NBP increasing by an average of 0.07 Tg C y⁻¹ for each 1% 298 increase in CFE (figure 6b). When compared to other scenarios, applying a 5% CFE to the high 290 radiative forcing and high land use scenario resulted in a mean annual NBP of 0.46 ± 0.3 Tg C y^{-1} , 300 roughly equivalent to mean annual NBP in the low radiative forcing and low land use scenario with 301 no CFE (0.52 \pm 0.12). A 15% CFE applied to the high radiative forcing and high land use scenario 302 resulted in a mean annual NBP of 1.18 ± 0.29 Tg C y^{-1} , more than double that of the low radiative forcing and low land use scenario with no CFE.

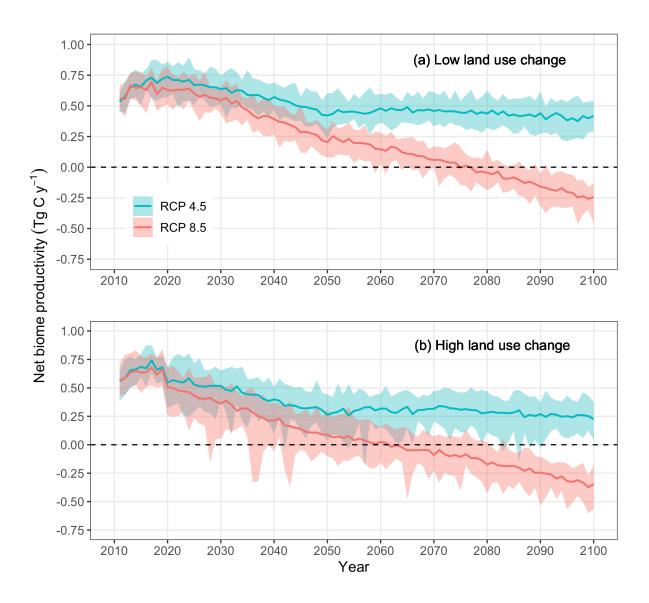


Figure 4: Projected changes in net biome productivity (NBP) for the seven main Hawaiian Islands. Values above zero indicate terrestrial ecosystems are acting as a net carbon sink for atmospheric carbon and values below zero indicate ecosystems are acting as a net carbon source to the atmosphere. Solid lines indicate the mean of 30 Monte Carlo realizations for each scenario, with shaded areas indicating the minimum and maximum range of Monte Carlo realizations. The dashed horizontal line in each panel represents the boundary between ecosystems acting as a net carbon sink (positive NBP values) and a net carbon source (negative NBP values).

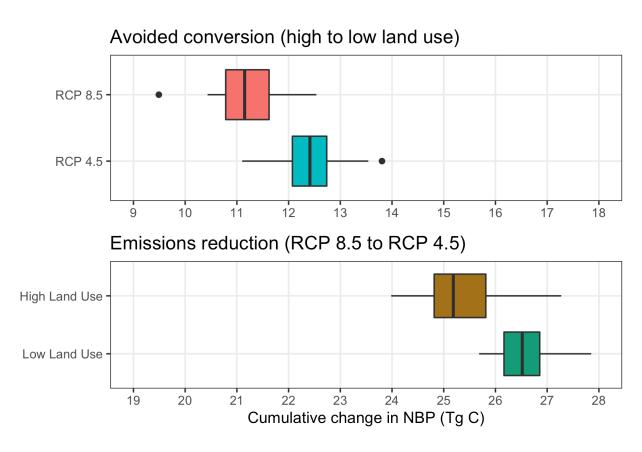


Figure 5: Projected changes in cumulative net biome productivity (NBP) for the seven main Hawaiian Islands when switching from the high to low land use change scenario under each radiative forcing scenario (top panel) and when switching from the high (RCP 8.5) to low (RCP 4.5) radiative forcing scenario under each land use scenario (bottom panel). Box plots indicate the median (vertical black line), 25th and 75th percentiles (colored boxes), 10th and 90th percentiles (thin horizontal lines), and values outside of this range (black circles). Note the different x-axis scales in each panel.

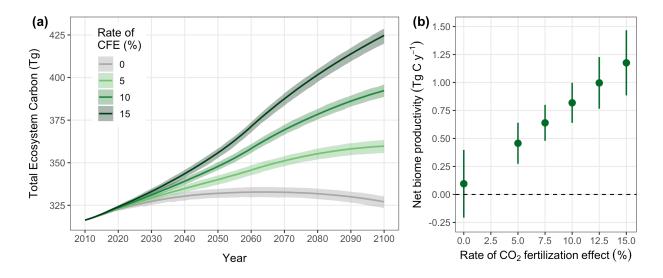


Figure 6: Sensitivity of projected changes in total ecosystem carbon storage (a) and mean annual net biome productivity (b) to different rates of carbon dioxide fertilization in the seven main Hawaiian Islands under the RCP 8.5 radiative forcing and high land use scenario. The carbon dioxide fertilization effect (CFE) is the percent change in net primary productivity (NPP) for every 100 ppm increase in atmospheric carbon dioxide. The CFE for all rates is capped at 600 ppm, which is achieved around the year 2060. Solid lines in (a) indicate the mean total ecosystem carbon storage across 30 Monte Carlo realizations for each CFE rate, with shaded areas indicating the minimum and maximum range of Monte Carlo realizations. Solid circles in (b) represent mean annual net biome productivity averaged across all years and Monte Carlo realizations for each CFE rate, with vertical lines indicating the standard deviation of the mean. The dashed horizontal line in (b) represents the boundary between ecosystems acting as a net carbon sink (positive NBP values) and a net carbon source (negative NBP values).

5 4. Discussion

We estimated that terrestrial ecosystems of the Hawaiian Islands have been a consistent net sink for 306 atmospheric carbon over the last decade (figure 4). For the time period 2011-2019, net biome 307 productivity (NBP) averaged 0.64 TgC y⁻¹ and ranged from 0.46 to 0.88 TgC y⁻¹ across all scenarios. 308 Based on this mean annual NBP estimate, Hawaiian terrestrial ecosystems offset approximately 13% 309 of 2015 statewide CO₂ emissions from energy production and transportation (5.04 TgC), the State of 310 Hawai'i's largest source of greenhouse gas emissions (State of Hawai'i 2019). Future projections 311 indicate Hawaiian terrestrial ecosystems will continue to be a net sink for atmospheric carbon if 312 global CO_2 emissions peak around 2040 and then decline (RCP 4.5), and that carbon sink strength 313 can be further enhanced by reducing the intensity and extent of future land use change. If, however,

global CO₂ emissions continue to rise throughout the 21st century (RCP 8.5), our projections indicate Hawaiian ecosystems will transition from a net sink to a net source of CO₂ to the atmosphere, with 316 high levels of land use change accelerating this transition and exacerbating net carbon loss. Our 317 model results also indicate that projections of ecosystem carbon balance are highly sensitive to the 318 introduction of a CFE. Even a 5% increase in NPP for every 100ppm increase in atmospheric CO₂ 319 was sufficient to completely offset all other carbon losses induced by the high radiative forcing and 320 high land use scenario, maintaining Hawaiian Island ecosystems as a net carbon sink for the entire 321 simulation period instead of transitioning to a net carbon source by mid-century. Reconciling the high 322 uncertainty surrounding the response of net photosynthesis to rising atmospheric CO_2 is essential to 323 more realistically constrain model projections of ecosystem carbon balance. 324

4.1 Impact of different climate and land use pathways

By comparing ecosystem carbon balance estimates under different scenario combinations, we were 326 able to assess the relative impact of both global emissions reductions and regional actions to reduce 327 emissions from land use, land use change, and wildland fire (figure 5). Global adherence to a lower 328 emissions trajectory (i.e., switching from RCP 8.5 to RCP 4.5) had the largest impact, resulting in a 329 median cumulative increase of 26 Tg C sequestered by Hawaiian ecosystems over the 90-year 330 simulation period. Long-term reductions in the intensity of land use change also consistently led to an 331 increase in ecosystem carbon sequestration, but to a lesser degree than global emissions reductions. 332 Switching from the high to the low land use scenario resulted in a median cumulative retention of an 333 additional 11.6 Tg C in Hawaiian ecosystems by 2100. The combination of global climate mitigation 334 and local reductions in land use conversion had the largest potential benefit to ecosystem carbon 335 sequestration, reducing cumulative net losses by over 400% (37.7 Tg C). Notably, the relative impact 336 of reducing emissions from land use change was much greater under the high radiative forcing 337 pathway (RCP 8.5). Cumulative NBP increased by 130% when switching from the high to low land 338 use scenario under RCP 8.5, as opposed to a 37% cumulative increase in NBP when switching from high to low land use under RCP 4.5. These results demonstrate that reducing ecosystem carbon losses from land use change, harvest, and wildland fire can be an important component of greenhouse gas
reduction efforts by sub-national jurisdictions like the State of Hawai'i, regardless of the global
emissions trajectory. These results also highlight the utility of Hawai'i's multi-pronged approach of
participating in global climate mitigation efforts by reducing emissions from the energy and
transportation sectors while also reducing land use emissions to minimize positive feedbacks to the
climate system.

4.2 Comparison to other studies

There are few estimates of contemporary ecosystem carbon balance for the main Hawaiian Islands, 348 and even fewer model projections of future ecosystem carbon balance in response to climate and land use change. Our mean annual NBP estimate of 0.64 TgC y⁻¹ for the period 2011-2020 agrees well 350 with a recent State of Hawai'i Greenhouse Gas Inventory report, which estimated an annual net 351 carbon sink of 0.66 Tg C in 2015 from agriculture, forestry, and other land uses (State of Hawai'i 352 2019). In contrast, our NBP estimate for the past decade was ~ 88% higher than a previous statewide 353 LUCAS model estimate covering the same time period (0.341 Tg C y⁻¹; Selmants et al 2017). This 354 discrepancy was likely driven by modifications in how we calculated NPP, soil R_h, and soil aquatic 355 carbon loss compared to previous versions of the LUCAS model, as well our model's finer spatial 356 resolution (Selmants et al 2017, Daniel et al 2018). Previous versions of a Hawai'i LUCAS model 357 were run at 1-km spatial resolution and simulation cells within each unique combination of moisture 358 zone and state type all had the same mean IBIS-derived NPP value applied to them at the beginning 350 of the simulation period. In contrast, our NPP estimates at 250-m spatial resolution were adjusted on 360 a cell-by-cell basis using Hawai'i-specific climate data as described in section 2.3. As a result, our 361 statewide NPP estimates from 2011-2020 were 9.5% lower on average than previous LUCAS model 362 estimates for Hawai'i during the same time period (Selmants et al 2017), likely because of the greater 363 influence of more arid simulation cells. Soil carbon losses via R_h, leaching, and overland flow in 364 previous versions of the LUCAS model were calculated as the ratio of the IBIS-derived flux to the size of the originating carbon stock, in this case soil organic carbon to 1-m depth (Daniel et al 2018).

Here we calculated soil R_h and aquatic carbon losses as the ratio of the mean IBIS-derived flux to the microclimate-adjusted NPP value of each simulation cell, which is a more realistic driver than stock size (Jackson *et al* 2017). Compared to previous Hawai'i LUCAS model estimates (Selmants *et al* 2017), soil R_h and aquatic carbon losses from 2011-2020 were reduced by an average of 15% and 21%, respectively, which widened the gap between between carbon gain (NPP) and carbon losses and accounted for the overall increase in NBP estimates for this time period.

373 4.3 Limitations of this study

There is ample evidence that increasing atmospheric CO₂ concentrations can stimulate NPP (Norby 374 et al 2005, Zhu et al 2016), but the magnitude and persistence of this effect remains highly uncertain (Franks et al 2013, Walker et al 2020). Our results demonstrate that long-term projections of ecosystem carbon balance are highly sensitive to uncertainty in CFE strength. With no CFE, 377 Hawaiian ecosystems became a net source of CO₂ to the atmosphere beginning in the latter half of the 378 21st century under the high land use and high radiative forcing scenario. However, a CFE equivalent 379 to a 5% increase in NPP for every 100 ppm increase in atmospheric CO₂ applied to the same scenario 380 resulted in Hawaiian ecosystems remaining a net carbon sink throughout the entire simulation period. 381 A 15% CFE applied to the high land use and high radiative forcing scenario resulted in a nearly 382 5-fold increase in mean annual NBP averaged across all years and Monte Carlo realizations. Despite 383 this demonstrated sensitivity to a CFE, several potentially attenuating factors complicate the selection 384 of a realistic CFE value with any degree of confidence (Walker et al 2020). Nitrogen and phosphorus 385 limitation can reduce or eliminate a CFE (Reich et al 2006, Norby et al 2010, He et al 2017, Terrer et 386 al 2019), as can water limitation and heat stress (Obermeier et al 2017, Birami et al 2020). Forest 387 age may also be a factor, with young aggrading forests showing a strong positive growth response to 388 CO₂ fertilization (Walker et al 2019), while old-growth forests show little to no response (Jiang et al 2020, Yang et al 2020). This evidence indicates that a CFE may be highly variable across space and 390 time, suggesting it may be unrealistic to apply a single CFE rate value across an entire region over 391 several decades. Until mechanistic understanding is improved, the most conservative approach when

projecting future ecosystem balance in the context of climate mitigation planning may be to assume no CFE, with the knowledge that any realized CFE will only enhance ecosystem carbon sequestration. Our model does not currently differentiate between forests dominated by native versus non-native tree 395 species, which could influence estimates of ecosystem carbon balance. Native forests in the Hawaiian Islands are dominated by 'ōhi'a (Metrosideros polymorpha Gauditch), an endemic foundational tree species found across a broad range of climatic and edaphic conditions (Ziegler 2002). Beginning in 2010, a fungal disease termed Rapid 'Ōhi'a Death (ROD) caused by two Ceratocystus spp. has emerged and caused widespread mortality to mature 'ōhi'a trees across a range of size classes, primarily on Hawai'i Island (Mortenson et al 2016, Fortini et al 2019). The distribution and potential 401 range of this emerging threat to Hawai'i's dominant native tree species have only recently been 402 mapped (Vaughn et al 2018, Fortini et al 2019), but the pulse of newly dead organic matter and 403 reduction in photosynthetic capacity induced by widespread tree mortality could significantly alter 404 ecosystem carbon balance over the long-term (Sleeter et al 2019). ROD-affected forests could also 405 undergo a replacement of canopy dominant stress-tolerating 'ōhi'a trees by non-native tree species 406 with faster relative growth rates, lower wood density, and faster tissue turnover, potentially altering 407 the long-term trajectory of carbon cycling in Hawaiian forest ecosystems. New model projections for 408 Hawai'i that incorporate ROD spread rates and forest restoration scenarios will therefore require 400 differentiation among forest ecosystems dominated by native and non-native tree species. 410 Interannual climate variability is a primary factor influencing spatial and temporal patterns of global 411 wildland fire activity (Abatzoglou et al 2018), with climate warming expected to increase wildland 412 fire frequency and wildfire season length across a wide range of biomes (Sun et al 2019). Although 413 our model projections capture the spatial and temporal variation in ignition probability and fire extent 414 by sampling from previous fire years (1999-2019), we did not incorporate the effects of projected 415 future climate change on the frequency and extent of wildland fire. Recent fire probability modeling 416 for the northwest portion of Hawai'i Island indicated that projected drying and warming trends under 417 RCP 8.5 could increase maximum fire probability values more than three-fold and shift areas of peak 418

flammability to higher elevation by mid-century (Trauernicht 2019). Extending this probability fire
modeling approach statewide would provide a quantitative, spatially explicit assessment of wildland
fire probability for the main Hawaiian Islands as predicted by climate, land cover, and ignition
density, which is highly correlated with population density (Trauernicht *et al* 2015, Trauernicht 2019).
This approach would provide future simulation model projections of Hawaiian ecosystem carbon
balance with more realistic scenarios of expected annual area burned based on the integrated effects
of future climate and land use change.

5. Conclusion

Although terrestrial ecosystems are currently an important sink for atmospheric CO2, the future 427 direction and magnitude of the land carbon sink are highly uncertain, especially at regional scales. 428 Our simulation modeling results indicated that projected climate change, dictated by long-term 429 trajectories in global greenhouse gas emissions, was the primary factor influencing terrestrial 430 ecosystem carbon balance in the Hawaiian Islands. Long-term reductions in the intensity of land use 431 change and wildland fire also consistently led to an increase in ecosystem carbon sequestration, but to 432 a lesser degree than global emissions reductions. CO₂ fertilization of NPP was the largest source of 433 uncertainty in long-term projections of ecosystem carbon balance in the Hawaiian Islands, 434 highlighting the need for greater mechanistic understanding of the cascading effects of rising 435 atmospheric CO₂ on ecosystem carbon sequestration. By incorporating the interactive effects of land use and climate change into future projections of ecosystem carbon balance, our model results could 437 be used as a set of baseline projections for the State of Hawai'i to evaluate different ecosystem-based climate mitigation strategies. Studies like ours that incorporate stochasticity into spatially explicit simulation models could also provide a framework for the growing number of sub-national jurisdictions that plan to incorporate ecosystem carbon sequestration into their emissions reduction 441 efforts. These long-term projections will be critical to understanding how future land use and 442 climate-biosphere feedbacks could influence the achievement of climate mitigation goals.

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Data and code availability

- Tabular model output data and metadata are available in machine readable format from the USGS
- ScienceBase data repository at https://doi.org/10.5066/P9AWLFKZ. Model input data and R code
- used to format input data, summarize output data, and compile this manuscript are available from a
- GitHub repository at https://github.com/selmants/HI Model.

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