

# Technical Talk - PhD Project

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## Introduction

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We are solving stochastic PDEs which model ocean-atmosphere interactions!

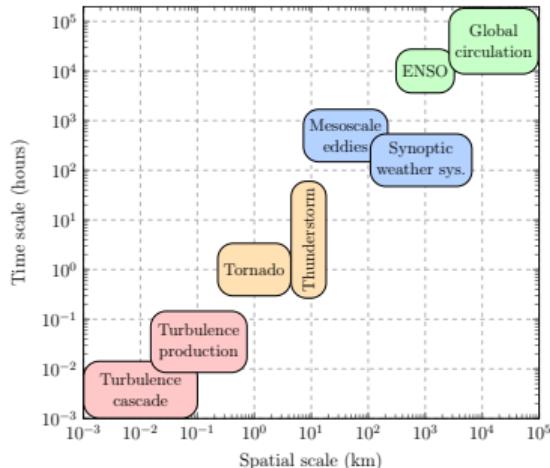
## Introduction

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Why add stochastic terms?

# Multiscale phenomena

- ▶ Earth's oceans and atmosphere are characterized by a rich hierarchy of interacting processes

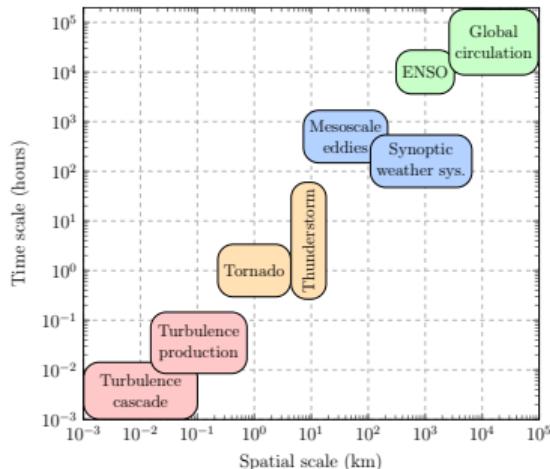


Spatial and temporal scales of various meteorological phenomena<sup>1</sup>

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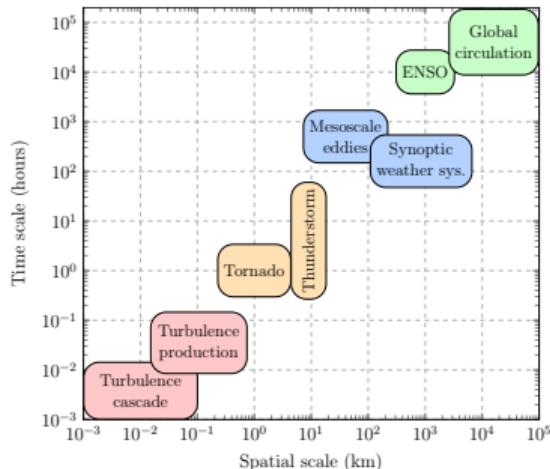
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- ▶ Numerical models have fixed resolution
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Spatial and temporal scales of various meteorological phenomena<sup>1</sup>

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- ▶ **Solution:** Parameterization or subgrid-scale modeling

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# Parameterization

- ▶ Additional terms are added to account for the missing effect of unresolved/small scales
- ▶ Parameterization techniques can be divided into two categories:
  1. Deterministic
  2. Stochastic
- ▶ We explored **stochastic parameterization** : stochastic advection by Lie transport (SALT)
- ▶ Why SALT?
  - ▶ preserves structure → conservation properties still hold
  - ▶ better forecast skills → uncertainty due to unresolved transport processes can be quantified
  - ▶ conducive to data-driven approaches → can be estimated from observations/satellite data

## Research objectives

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# Objectives

1. Explore the efficacy of stochastic parameterization (in terms of UQ skills) for a climate model
  - ▶ Previous studies focused on Euler, quasi-geostrophic, and shallow water equations
2. Numerically solve the idealized stochastic climate model
3. How to model the noise/stochastic term effectively?
4. Compare the stochastic model results with other existing models/approaches

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## The climate model

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# Model equations

2D climate model equations:

$$\text{Atmosphere : } d\mathbf{u}^a + ((\mathbf{u}^a dt + \sum_i \boldsymbol{\xi}_i \circ dW^i) \cdot \nabla) \mathbf{u}^a + \frac{1}{Ro^a} \hat{\mathbf{z}} \times (\mathbf{u}^a dt + \sum_i \boldsymbol{\xi}_i \circ dW^i)$$

$$+ \sum_i (u_1^a \nabla \xi_{i,1} + u_2^a \nabla \xi_{i,2}) \circ dW^i = (-\frac{1}{C^a} \nabla \theta + \frac{1}{Re^a} \Delta \mathbf{u}^a) dt,$$

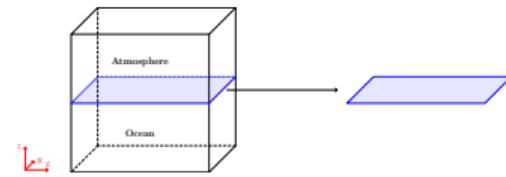
$$d\theta^a + \nabla \cdot (\theta^a (\mathbf{u}^a dt + \sum_i \boldsymbol{\xi}_i \circ dW^i)) = (\gamma(\theta^a - \theta^o) + \frac{1}{Pe^a} \Delta \theta^a) dt,$$

$$\text{Ocean : } \frac{\partial \mathbf{u}^o}{\partial t} + (\mathbf{u}^o \cdot \nabla) \mathbf{u}^o + \frac{1}{Ro^o} \hat{\mathbf{z}} \times \mathbf{u}^o + \frac{1}{Ro^o} \nabla p^a = \sigma(\mathbf{u}^o - \mathbb{E}\mathbf{u}_{sol}^a) + \frac{1}{Re^o} \Delta \mathbf{u}^o,$$

$$\nabla \cdot \mathbf{u}^o = 0,$$

$$\frac{\partial \theta^o}{\partial t} + \mathbf{u}^o \cdot \nabla \theta^o = \frac{1}{Pe^o} \Delta \theta^o.$$

Incompressible/compressible Navier-Stokes + Advection-diffusion eqs.



Model domain

- $\mathbf{u}$  : velocity,
- $\theta$  : temperature,
- $p$  : pressure

# How to get the stochastic terms?

- ▶ Difference between Lagrangian trajectories at different resolutions:

$$\sum_i \xi_i dW^i \approx \mathbf{u}_{true} dt - \mathbf{u} dt$$

$\mathbf{u}_{true}$  : true velocity

$\mathbf{u}$  : mean flow velocity i.e. the large-scale component of  $\mathbf{u}_{true}$

- ▶ Ways to estimate the true velocity  $\mathbf{u}_{true}$ :

1. Observation data (for example, from satellites)
2. Synthetic data (from high-resolution numerical simulations)

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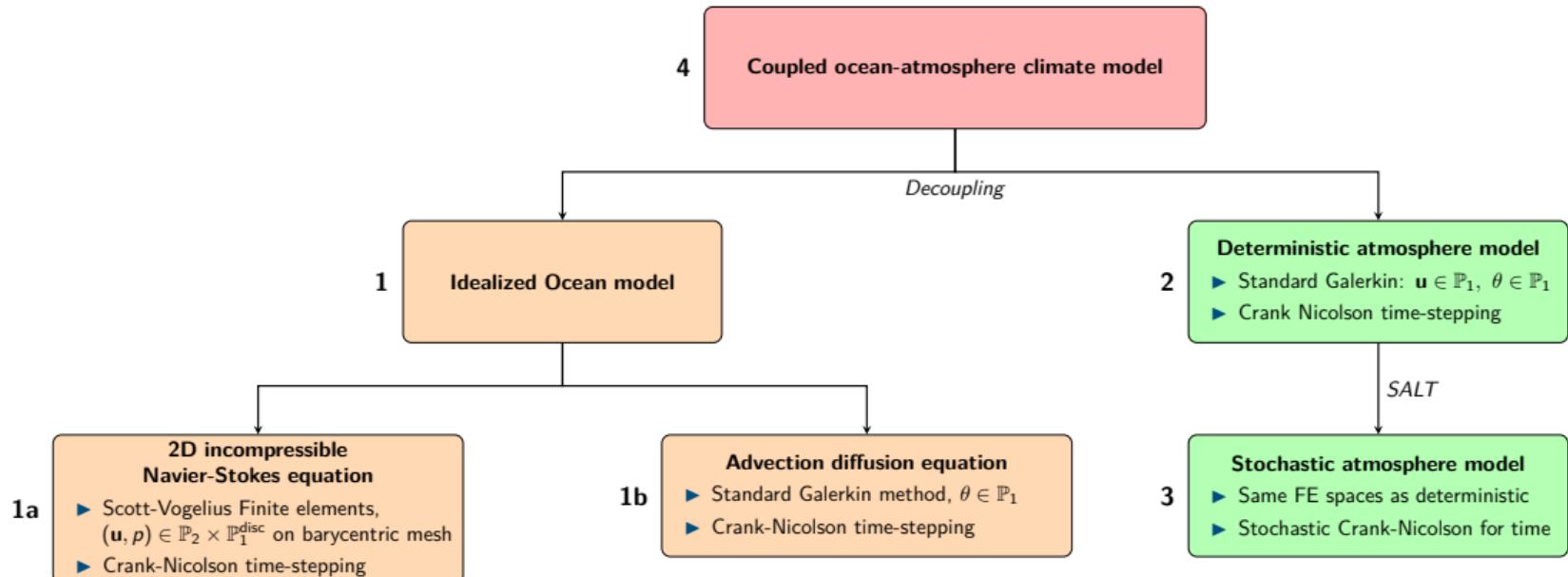
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- ▶ **We use synthetic data!**

# Our methodology



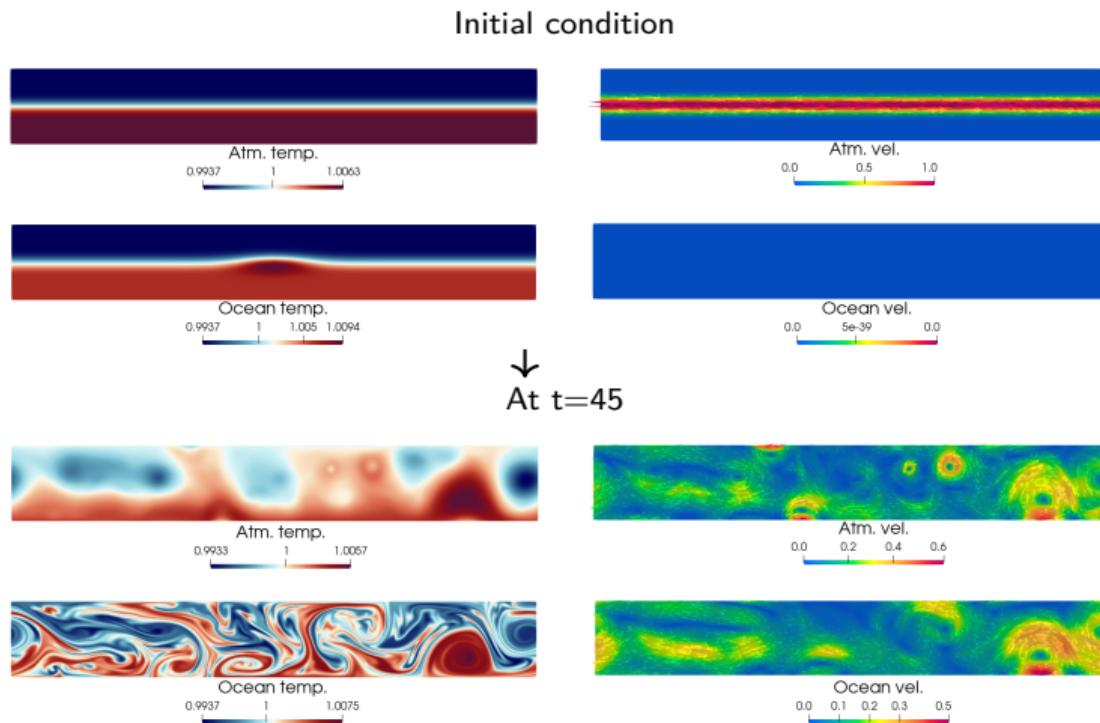
We used an open-source Python Finite Element package for numerical implementation and simulation

## Numerical simulation of the climate model

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# Deterministic model simulation and model calibration

- We run the climate model in high-res. ( $\Delta x = 1/128 \sim 30$  km) for 45 time units ( $\sim 25$  days)



- Analyze data → extract small-scale features (using statistical algorithms) → obtain  $\xi_i$

## Numerical simulation of the climate model

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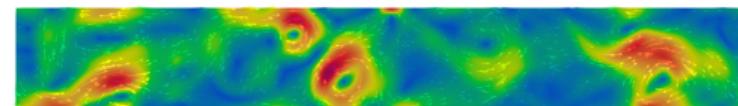
Stochastic model simulation

# Stochastic model simulation

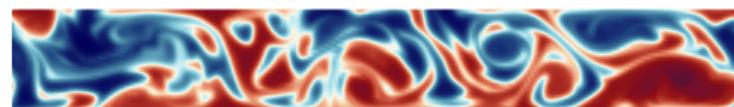
- We run the stochastic model on a coarse grid ( $\Delta x = 1/32 \sim 120$  km)
- Initial conditions correspond to coarse grained velocity and temperature fields at  $t = 25$



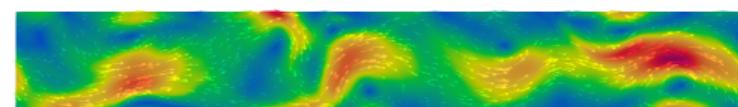
Atm. temp.  
0.9934 1 1.0054



Atm. vel.  
0.0 0.2 0.4 0.7

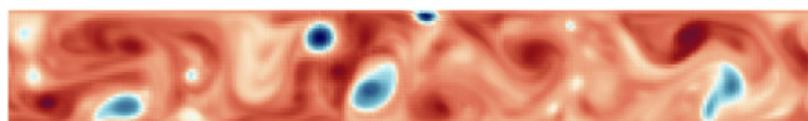


Ocean temp.  
0.9938 1 1.0072



Ocean vel.  
0.0 0.1 0.2 0.3 0.4

Atmospheric vorticity at  $t = 25$

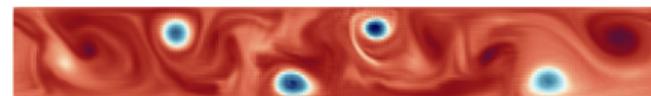


Atm. vort.  
-12.1 -5 0 6.4

# Simulation results

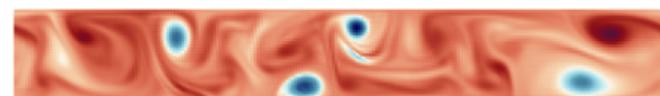
- ▶ Stochastic model is simulated for 20 time units ( $t = 25$  to  $t = 45$ )
- ▶ Stochastic vs. deterministic (without parameterization) model results at  $t = 35$

Particle 1



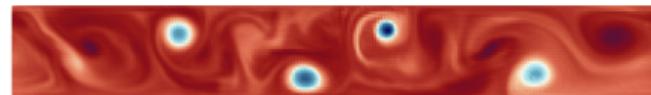
Atm. vort. particle 1  
-9.5 -5 0 3.4

Deterministic solution



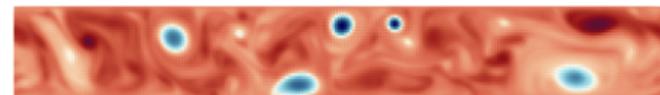
Atm. vort. deterministic  
-7.6 -5 0 4.0

Particle 2



Atm. vort. particle 2  
-11.0 -5 0 3.3

True solution



Atm. vort. truth  
-9.1 -5 0 4.4

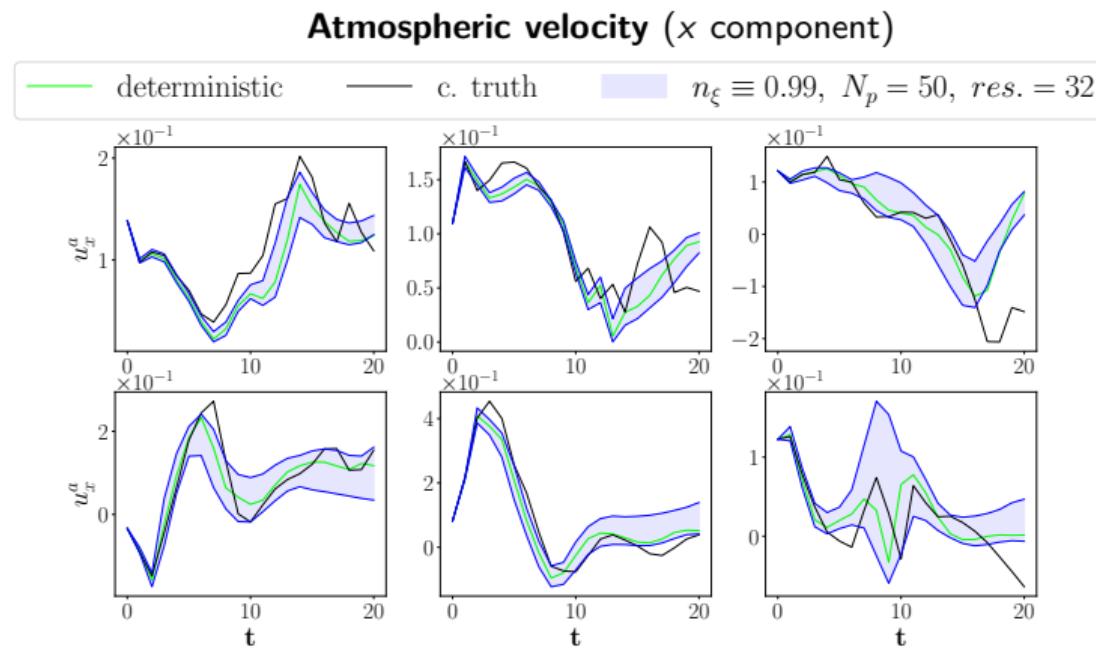
Particle 3



Atm. vort. particle 3  
-7.5 -5 0 3.9

# Simulation results

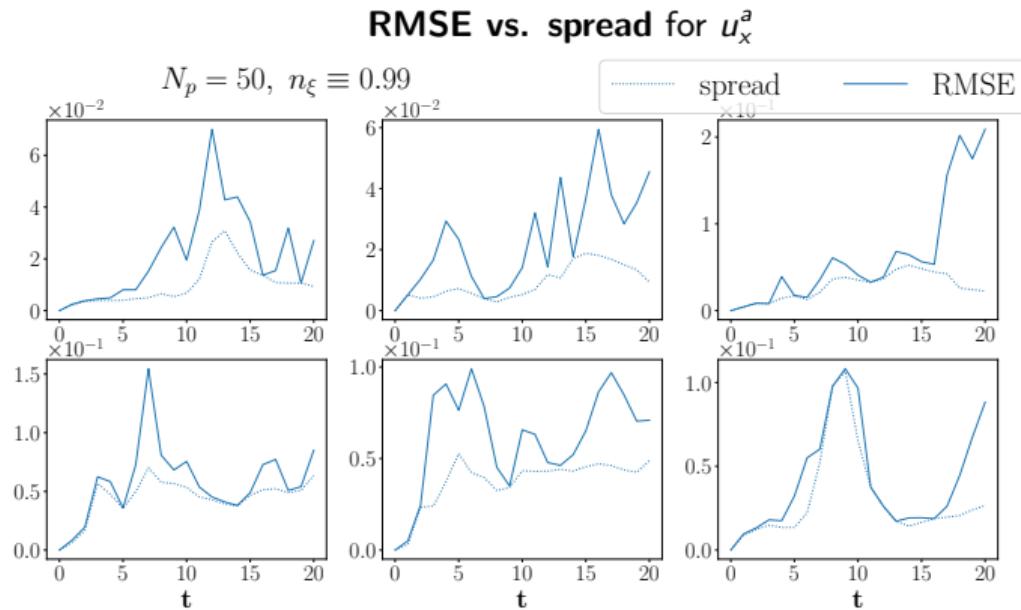
- ▶ Comparison between the stochastic ( $\mu \pm 1\sigma$ ) and deterministic sol. at 6 grid points:



1. Ensemble spread increases over time
2. Stochastic model captures the truth for 5 to 10 time units

# Uncertainty quantification

- ▶ Plots for the evolution of ensemble spread and RMSE at six different locations on the grid



- ▶ The size of ensemble spread is proportional to the RMSE (for at least 15 time units)
- ▶ Error in the stochastic solution can be estimated by its own spread
  - ▶ SPDE solution is suitable for data assimilation methods!

## Summary

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# Summary & outcomes

- ▶ **Aim:** solve a stochastic coupled ocean-atmosphere climate model
- ▶ **Novelty:** Data-driven model
  - ▶ stochastic terms (modelling unresolved processes) are obtained from observation data
  - ▶ has desired properties: structure-preservation
  - ▶ We used high-res. model sim. data to estimate  $\xi$ ; → data analysis & statistical modelling
- ▶ **Methodology** for solving the model equations:
  - ▶ Solve simpler models first
  - ▶ build the climate model step-by-step
  - ▶ verify numerical schemes at each step
- ▶ **Results:**
  - ▶ Stochastic model solution capture the truth for some time units
  - ▶ Shows promising UQ test results: ensemble spread size  $\propto$  RMSE
- ▶ **Outcome:** first step towards using SALT in a highly complex coupled system!

Thank you for your attention!