

Bayesian estimation of global glacial isostatic adjustment for sea level rise re-evaluation

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26 July 2017

GlobalMass

- combine satellite and in-situ data related to different aspects of the sea level budget,
- attribute global sea level rise to its component parts.



European Research Council
Established by the European Commission

Funded by the European Research Council (ERC) under the European Union's Horizon 2020 research and innovation programme under grant agreement No 69418.

The sea level budget *enigma*

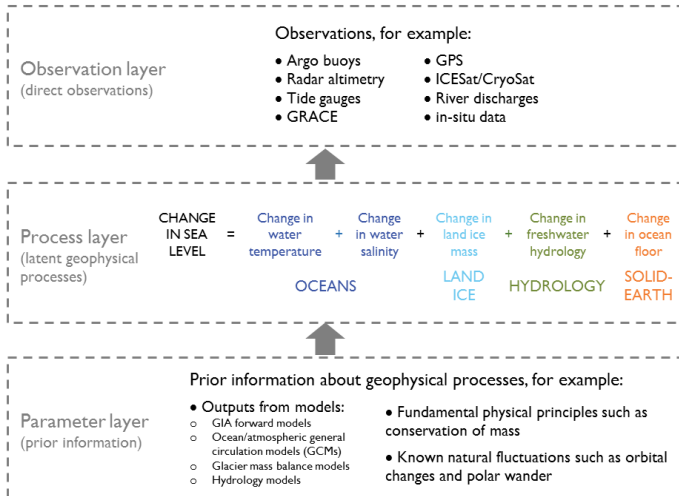
$$\Delta \text{sea level}(t) = \underbrace{\Delta \text{barystatic}(t)}_{\text{mass}} + \underbrace{\Delta \text{steric}(t)}_{\text{density}} + \underbrace{\text{GIA}}_{\text{ocean basins}}$$

- GIA: glacial isostatic adjustment
- inconsistencies between the discipline-specific estimates

GlobalMass Aims

- simultaneous global estimates of all the components
- close the sea level budget

The Bayesian hierarchical model



First step – modelling GIA

What is GIA?

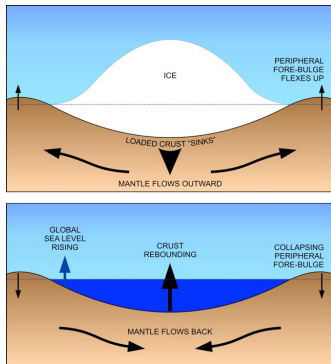
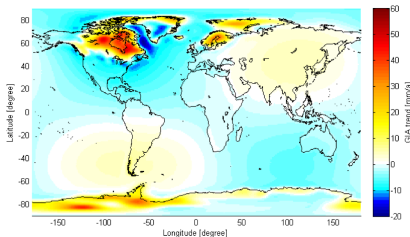


Figure courtesy of the Canadian Geodetic Survey, Natural Resources Canada.

GIA - glacial isostatic adjustment

- vertical displacement of the Earth's crust once burdened by ice-age glaciers ($\sim 20,000$ yrs ago)
- constant over the time scale of this study (1981-2020)
- Regional discrepancies in estimates from physical models, e.g. ICE-6G (Peltier et al., 2015)

Denote the true GIA process by $\mathbf{Y} : \mathbb{S}^2 \mapsto \mathbb{R}$, and the prior mean trend by $\boldsymbol{\mu} : \mathbb{S}^2 \mapsto \mathbb{R}$.



GIA estimates from a ICE-6G model.

After detrend, the residuals become stationary and assume

$$\mathbf{X} := \mathbf{Y} - \boldsymbol{\mu} \sim \mathcal{GP}(\mathbf{0}, \kappa(\rho, \sigma^2)) \quad (1)$$

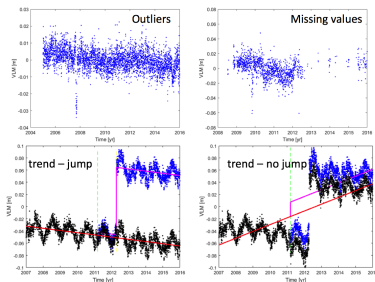
$\kappa(\rho, \sigma^2)$ - Matérn covariance function with length-scale and variance parameters ρ and σ^2 .

The GPS observations

The GPS observation at the i^{th} selected station Z_i can be decomposed as yearly trend of vertical movement in the Earth's surface and measurement error $\varepsilon_i \sim \mathcal{N}(0, e_i^2)$.



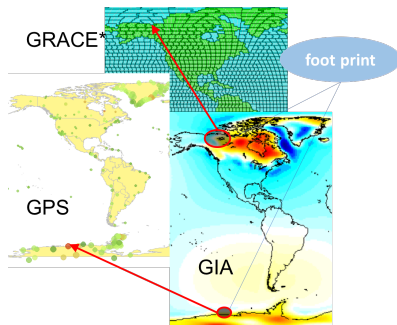
A GPS station in Antarctic.



e_i^2 is the estimated from raw GPS time series and fixed.

Forward modelling

From process to observations



*GRACE: the Gravity Recovery And Climate Experiment satellites

Define \mathcal{A}_i as the map from the latent process over the appropriate spatial **foot print** to the i^{th} observations,

$$Z_i = \mathcal{A}_i \mathbf{Y} + \varepsilon_i, \quad (2)$$

\mathcal{A}_i is usually a linear operator

- point \rightarrow point
- area \rightarrow point
- area \rightarrow area

Denote by $\mathcal{A}^T = [\mathcal{A}_1^T, \dots, \mathcal{A}_N^T]$ Then we can have the vector form for equation 2

$$\mathbf{Z} = \mathcal{A}\mathbf{Y} + \varepsilon \quad (3)$$

Bayesian hierarchical model for GIA

Combine the observation equation 3 and the process equation 1

$$\begin{cases} \tilde{\mathbf{Z}} = \mathcal{A}\mathbf{X} + \varepsilon, \quad \varepsilon \sim \mathcal{N}(\mathbf{0}, \text{diag}(e_1^2, e_2^2, \dots, e_N^2)) \\ \mathbf{X} \sim \mathcal{GP}(\mathbf{0}, \kappa(\rho, \sigma^2)) \\ \rho \sim \pi(\rho), \quad \sigma^2 \sim \pi(\sigma^2) \end{cases} \quad (4)$$

Predicted mean and uncertainty of GIA

Pointwise Bayesian update of the GIA process on a fine grid. For grid point X_i , the predictive distribution is

$$\pi(X_i|\tilde{\mathbf{Z}}) = \int_{\mathcal{X}_{-i}} \pi(\mathbf{X}|\tilde{\mathbf{Z}}) d\mathbf{X}_{-i} \quad (5)$$

predicted mean $X_i^* = \mathbb{E}(X_i|\tilde{\mathbf{Z}})$

predicted uncertainty $u_i^* = \text{sd}(X_i|\tilde{\mathbf{Z}})$

The GMRF approximation

- Bayesian update of the GP on a grid of m points scales as $\mathcal{O}(m^3)$, $m \sim 10^5$ for a 1° global grid.
- Gaussian Markov random field (GMRF) with sparse precision matrix has much better scaling property.

The SPDE approach (Lindgren et al., 2011)

Denote by $\tilde{\mathbf{X}}$ the GMRF approximation of \mathbf{X} on a given triangulation with basis functions $\{\phi_i\}_{i \in \mathbb{N}}$, then $s \in \mathbb{S}^2$

$$\mathbf{X}(s) \approx \phi_i(s)^T \tilde{\mathbf{X}} \quad (6)$$

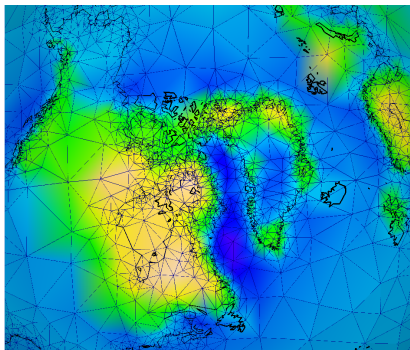
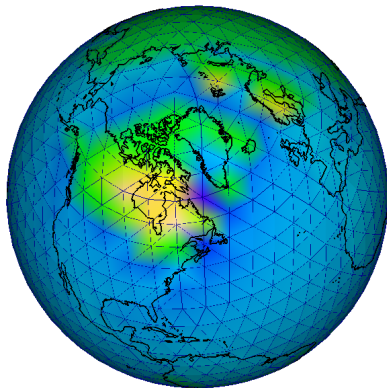
For a grid $\mathbf{S} \subset \mathbb{S}$ of locations, we find the weight matrix \mathbf{C} and write $\mathbf{X}(\mathbf{S}) \approx \mathbf{C}\tilde{\mathbf{X}}$.

Now the BHM 4 becomes

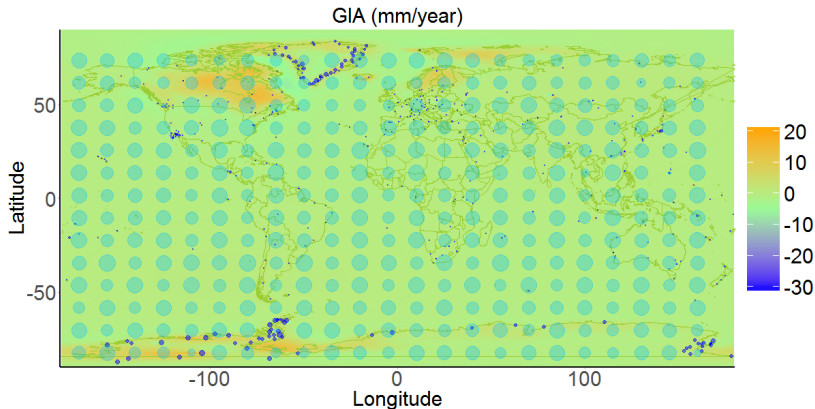
$$\begin{cases} \tilde{\mathbf{Z}} = \mathbf{A}\mathbf{C}\tilde{\mathbf{X}} + \boldsymbol{\varepsilon}, \quad \boldsymbol{\varepsilon} \sim \mathcal{N}(\mathbf{0}, \text{diag}(e_1^2, e_2^2, \dots, e_N^2)) \\ \tilde{\mathbf{X}} \sim \mathcal{N}(\mathbf{0}, \mathbf{Q}^{-1}(\rho, \sigma^2)) \\ \rho \sim \pi(\rho), \quad \sigma^2 \sim \pi(\sigma^2) \end{cases} \quad (7)$$

where $\tilde{\mathbf{X}}$ is defined by the precision matrix \mathbf{Q}

Choosing the mesh on sphere



First solution for GIA



- Test the our statistical framework on the GIA.
- Find a better mesh.
 - Adaptive – fast computation
 - regularity – stable approximation
- Need better quality GPS data.
- Extend to the full system.

- Lindgren, F., Rue, H., and Lindström, J. (2011). An explicit link between Gaussian fields and Gaussian Markov random fields: the stochastic partial differential equation approach. *Journal of the Royal Statistical Society: Series B (Statistical Methodology)*, 73(4):423–498.
- Peltier, W. R., Argus, D. F., and Drummond, R. (2015). Space geodesy constrains ice age terminal deglaciation: The global ICE-6G_C (VM5a) model. *Journal of Geophysical Research: Solid Earth*, 120(1):450–487.

Thank you!