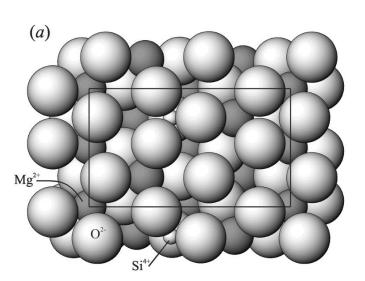
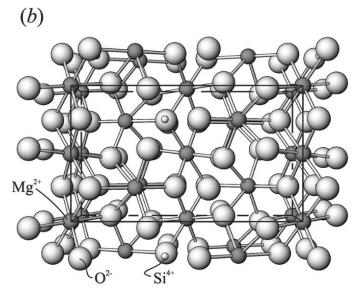


# Mineral structures

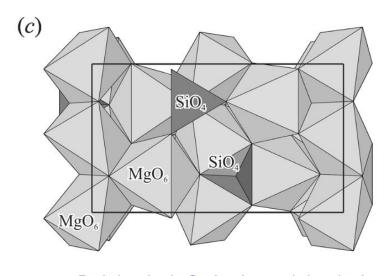
#### How are mineral structures illustrated?



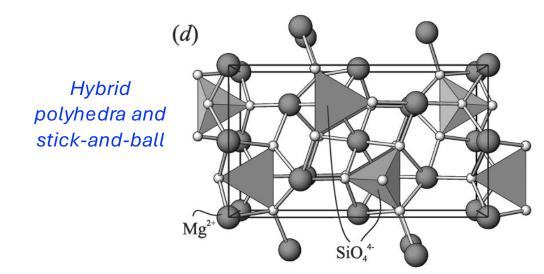
Spheres to scale

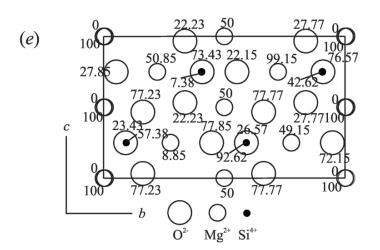


Stick-and-ball: Reduced sphere size



Polyhedral. Only the polyhedral framework is shown





"Mapped" structure.

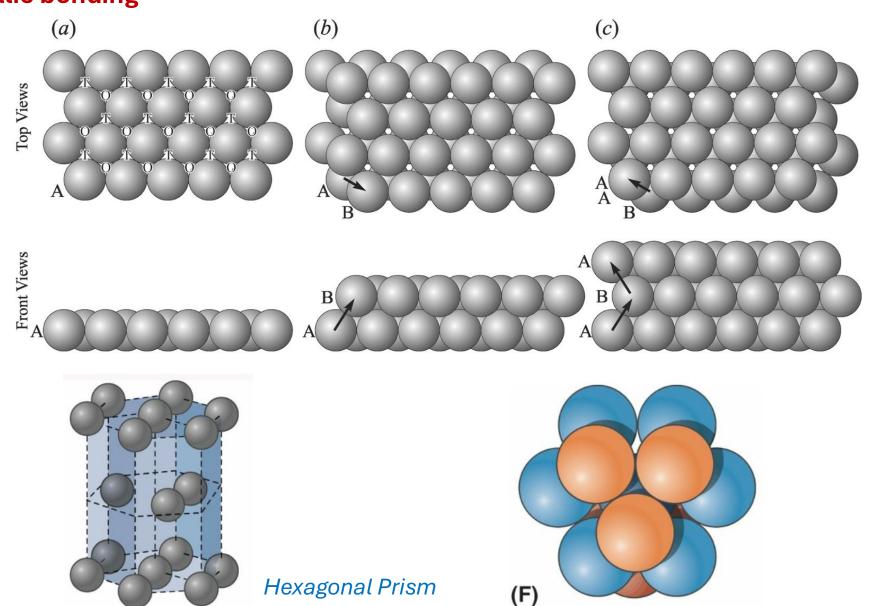
### ☐ Structure controls with metallic bonding

The metal atoms tend to pack together in highly ordered arrangements that minimize void space.

#### Packing of spheres:

### **Hexagonal Closest Packing**

- ✓ Close-packed (A) layer of spheres. Each sphere is in contact with six neighbors.
- ✓ The T site is surrounded by four spheres and the O sites are surrounded by six spheres.
- ✓ Third layer is stacked directly above the first



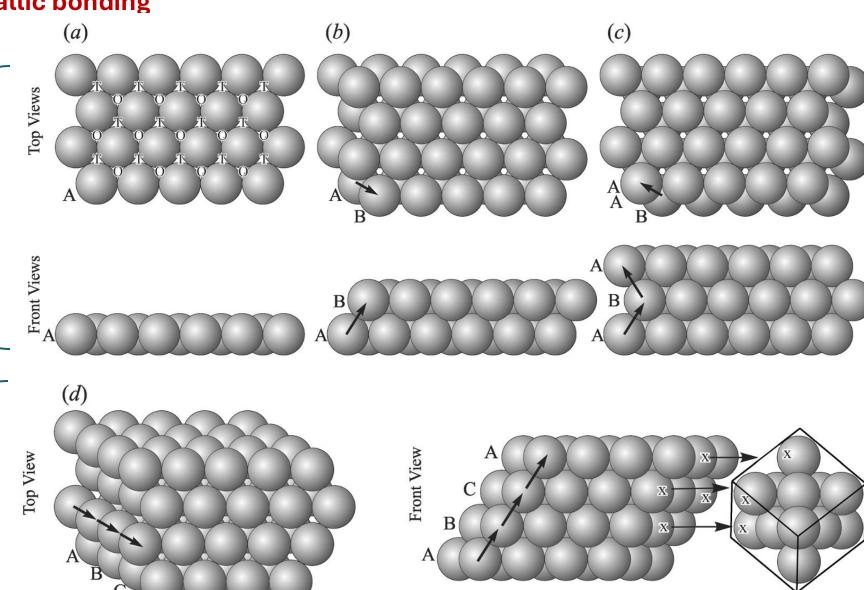
#### ☐ Structure controls with metallic bonding

The metal atoms tend to pack together in highly ordered arrangements that minimize void space.

Packing of spheres:

#### **Cubic Closest Packing**

✓ The C layer is shifted in the same direction as B, so the C spheres are directly above the O sites in the A layer



#### □ Structure controls with ionic bonding

(a) 12-fold coordination (based on cubic closest packing). The coordination polyhedron is not a regular shape.

 $R_c$  = Radius of cation

 $R_a$  = Radius of anion

$$R_c/R_a = 1$$

(b) 8-fold, or **cubic**, coordination.

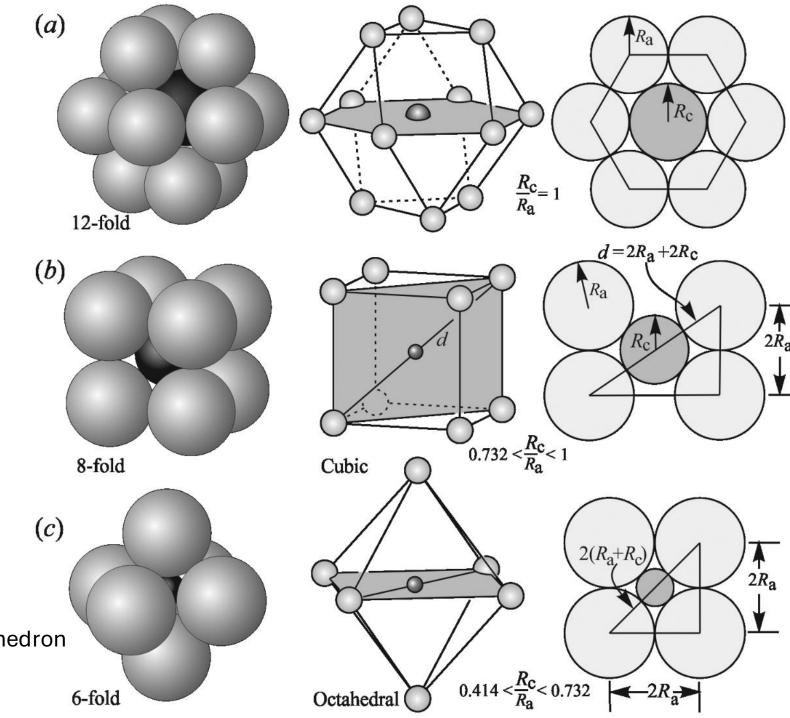
8 anions around a cation, making a cube

$$R_c/R_a = 0.732 - 1$$

(c) 6-fold, or **octahedral**, coordination.

6 anions around a cation, making an octahedron

$$R_c/R_a = 0.414 - 0.732$$

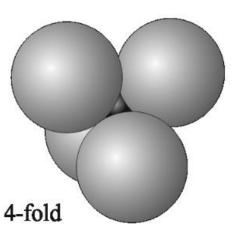


#### □ Structure controls with ionic bonding

(d) 4-fold, or tetrahedral, coordination.

4 anions around a cation, making a tetrahedron

$$R_c/R_a = 0.225 - 0.214$$



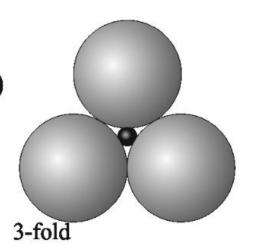


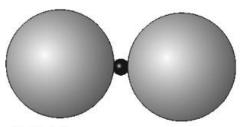
3 anions around a cation, making a (e) triangle

$$R_c/R_a = 0.155 - 0.225$$

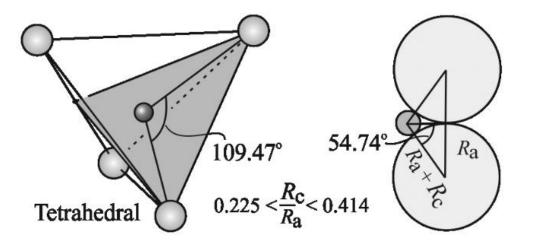
(f) 2-fold, or **linear**, coordination.

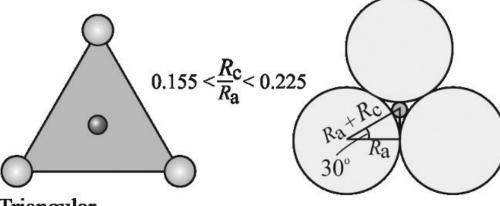
2 anions around a cation, making a **(f)** line



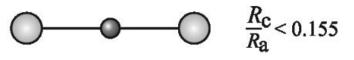


2-fold





Triangular



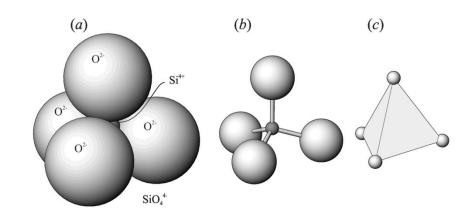
Linear

 $R_c/R_a < 0.155$ 

☐ Structure controls with ionic bonding

**Table 4.2** Cation Sizes Appropriate for Regular Coordination Polyhedra When Coordinating with O<sup>2-</sup> (~1.26 Å)

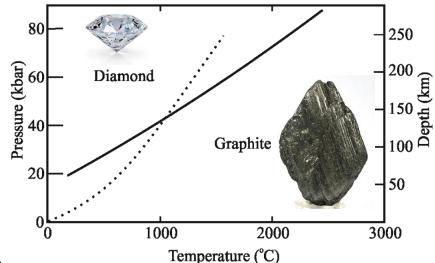
Coordination	Radius Ratio	-	~Maximum Radius (Å)	Common Cations
12	~1.00	1.26	N/A	K <sup>+</sup> , Ca <sup>2+</sup> , Na <sup>+</sup> Fe <sup>2+</sup> , Ca <sup>2+</sup> ,
8	0.732 - 1.00	0.92	1.26	Na+, Mg <sup>2+</sup>
6	0.414-0.732	0.52	0.92	Al <sup>3+</sup> , Fe <sup>2+</sup> , Fe <sup>3+</sup> , Mg <sup>2+</sup>
4	0.225-0.414	0.28	0.52	Si <sup>4+</sup> , Al <sup>3+</sup> , S <sup>6+</sup> , P <sup>5+</sup>
3	0.155-0.225	0.20	0.28	C <sup>4+ a</sup>
2	<0.155	N/A	0.20	None



(b) (a) (c) ☐ Structure controls with ionic bonding SiO<sub>4</sub><sup>4</sup> Na⁺ +2 +5 +3 +1 +6 anions cations

Polymorphism: The ability of a chemical compound to crystallize with more than one structure

- ✓ At a given T and P, one structure may represent the lowest energy configuration, therefore stable
- ✓ At other T and P, a different structure may be more stable.
- ✓ In general, high pressures favor tightly packed structures, reflected in higher mineral density.

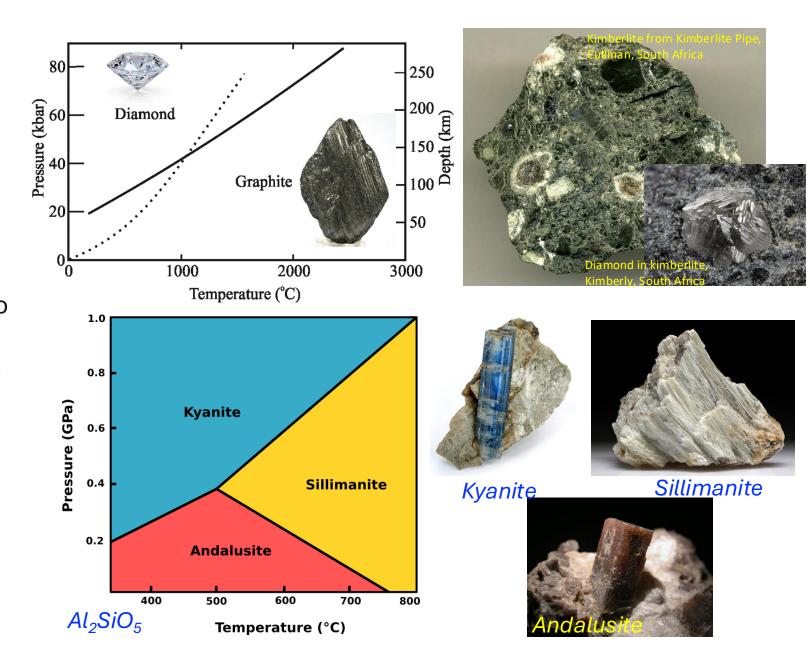




- ✓ **High temperature** tends to favor somewhat more open, lower density structures and structures that allow greater diversity in the occupancy of specific structural sites.
- ✓ Different polymorphs of the same substance are stable under different sets of conditions
- ✓ The presence of a given polymorph in a rock may indicate the conditions under which the rock was formed or to which it has been subjected.

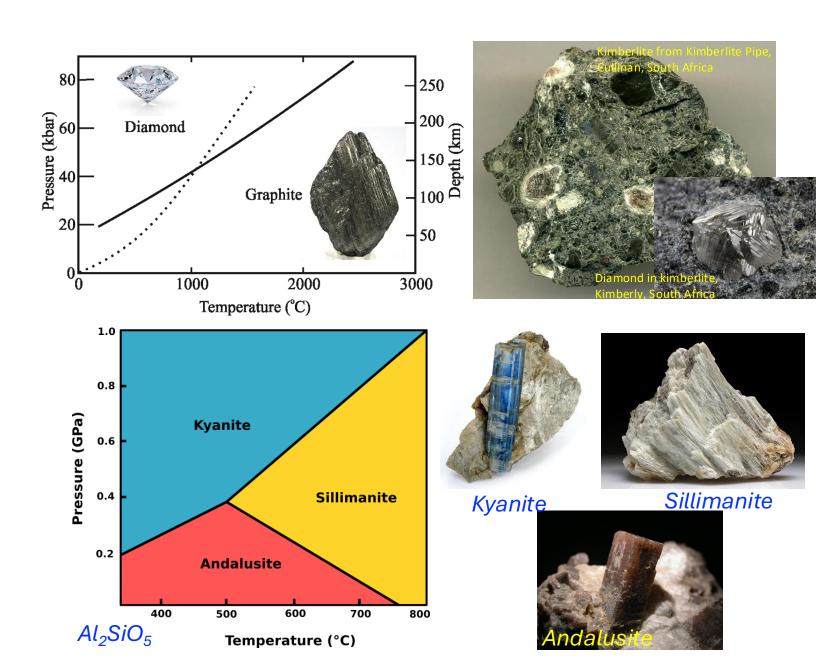
#### Reconstructive polymorphism:

- ✓ Conversion from one polymorph to another involves a major reorganization of the crystal structure.
- ✓ The chemical bonds that hold one structure together must be broken so that the atoms/ions can be rearranged and bonded into the new structure.
- ✓ Usually involves a large change in energy of the structure which must occur at the transformation temperature and/or pressure.

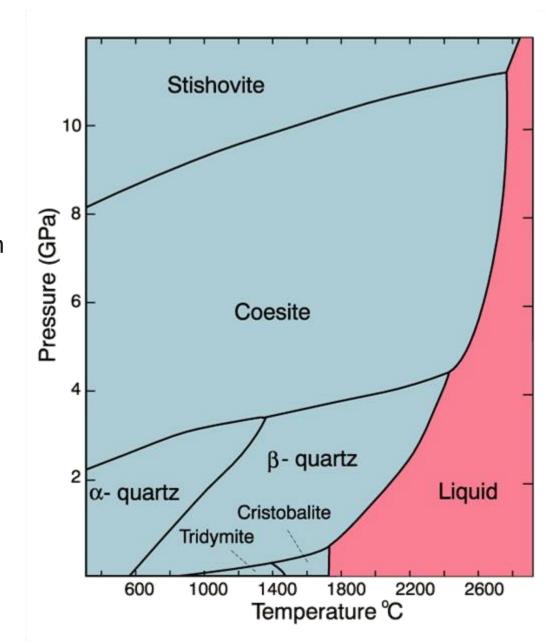


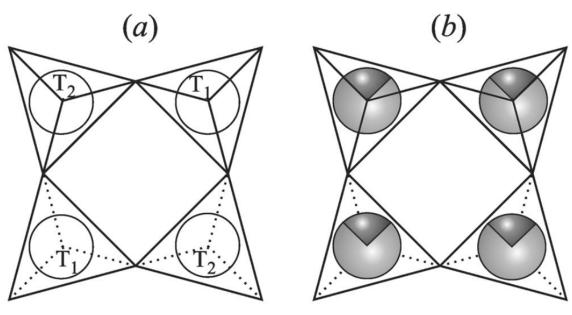
#### Reconstructive polymorphism:

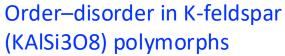
- ✓ Because of the extensive rearrangement involved, the rate at which this type of *transformation* occurs may be *very slow*.
- ✓ If the rate of the transformation is very slow, unstable polymorphs (metastable) may exist for long periods of time.



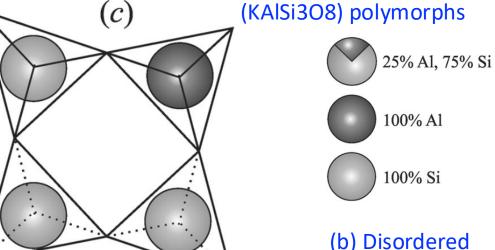
- Displacive polymorphism
- ✓ Polymorphic inversions do not involve breaking chemical bonds; the difference between polymorphs is simply a distortion or bending of the crystal structure.
- ✓ High-temperature forms typically have higher symmetry than low-temperature polymorphs.
- ✓ The crystal shape of the high-temperature polymorph will be retained on inversion to the low-temperature polymorph, though internal strains in the crystal lattice may lead to the formation of transformation twins







(c) Ordered



#### Order–disorder polymorphism

- ✓ The mineral structure remains more or less the same.
- ✓ Changes in the cation distribution within structural sites.
- ✓ If two cations X and Y can occupy two equivalent structural sites T1 and T2, the structure is considered *disordered* if there is an *equal probability* of finding X in either T1 or T2.
- ✓ If all X (e.g. Al) cations are located in one site T1 and all Y (e.g. Si) cations are in other T1 and T2 sites, the structure is considered fully *ordered*.

## Mineral classification

Minerals other than the native elements are conventionally classified based on the identity of the major *anion* or *anionic group*. (First degree of classification)

Differences between 'ides' and '-ates'

Only Anion: '-ides'

Anionic group: '-ates'

**Table 4.5** Chemical Classification of Minerals

Mineral Group	Anion or Anionic Group	Mineral Group	Anion or Anionic Group
Native elements	N/A	Carbonates	$CO_3$
Oxides	O	Nitrates	$NO_3$
Hydroxides	ОН	Borates	$BO_3$ , $BO_4$
Halides	Cl, Br, F	Chromates	$CrO_4$
Sulfides	S	Tungstates	$WoO_4$
Arsenides	As	Molybdates	$MoO_4$
Antimonides	Sb	Phosphates	$PO_4$
Selenides	Se	Arsenates	$AsO_4$
Tellurides	Te	Vanadates	$VO_4$
Sulfates	$SO_4$	Silicates	$SiO_4$

- Minerals have a definite but not fixed chemical composition (as minerals are crystalline solids).
- Variation in composition takes place keeping the crystal structure constant.

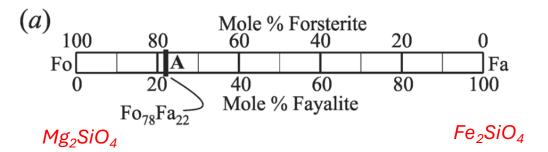
#### **☐** Solid solution

- ✓ Compositional variation is not a matter of mixing various solutes into a solid solvent.
- ✓ Compositional variation is a consequence of the ability of different elements, mostly cations, to substitute for each other within the crystal structure.
- ✓ Compositional variation is possible because different cations can interchangeably occupy the various sites among the anions.





- Minerals have a *definite* but *not fixed* chemical composition (as minerals are crystalline solids).
- Variation in composition takes place keeping the crystal structure constant.
- **Some important terminologies related to compositional variation**
- ✓ **Solid solution series:** The range of compositions produced by solid solution in a given mineral
- ✓ **End members:** The compositional extremes of a substitution series
- ✓ A continuous or complete substitution series is one in which all intermediate compositions are possible.



✓ An incomplete or discontinuous substitution series is one in which only a restricted range of compositions between the end members is found.

#### **☐** Substitution solid solution

Substitution involves interchanging one cation for another in a structural site. *Two* requirements control whether substitution can occur.

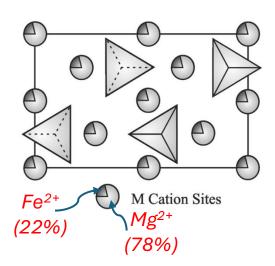
#### Size

- ✓ Ion sizes must be similar.
- ✓ If the difference in ion size is <15%, extensive substitution is usually possible; if the difference in ion size is >15%, substitution becomes limited.
- ✓ For example, Si<sup>4+</sup> and Al<sup>3+</sup> may substitute in tetrahedral sites; Mg<sup>2+</sup>, Fe<sup>2+</sup>, Fe<sup>3+</sup> and Al<sup>3+</sup> may substitute in octahedral sites; and Na<sup>+</sup> and Ca<sup>2+</sup> may substitute in 12-fold sites.
- ✓ Temperature has a substantial influence on the degree of substitution.

#### Charge

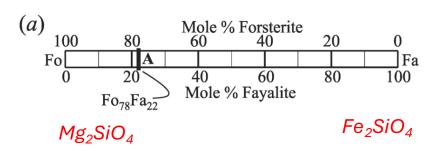
- ✓ Charge neutrality must be maintained.
- ✓ The ions that substitute for each other in a crystal structure may have the same or different charges.
- ✓ Based on charge difference between the substituting ions, four different substitutions can take place.
- ✓ Simple substitution is possible if the cations have the same charge.
- ✓ Coupled, omission, or interstitial substitution is invoked if cation charges are different.

#### Simple substitution



- Fe<sup>2+</sup> ↔ Mg<sup>2+</sup>

  Radii of Fe<sup>2+</sup> and Mg<sup>2+</sup> in octahedral coordination with O<sup>2-</sup> are 0.75 and 0.86 Å, respectively.
  - ✓ Same charge
  - ✓ Readily substituted.



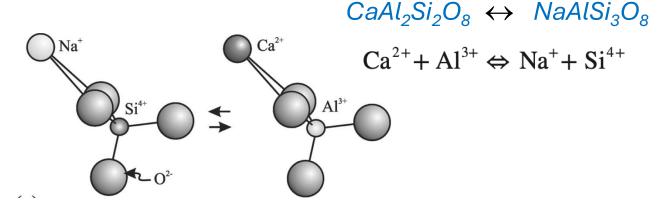


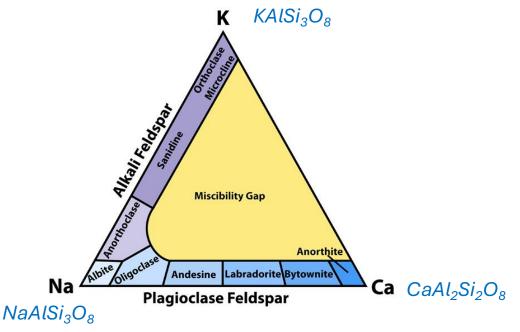


Forsterite Mg<sub>2</sub>SiO<sub>4</sub>

Fayalite Fe<sub>2</sub>SiO<sub>4</sub>

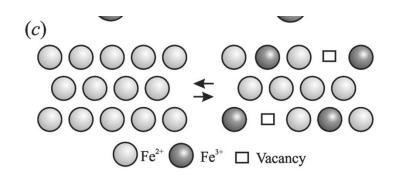
#### Coupled substitution



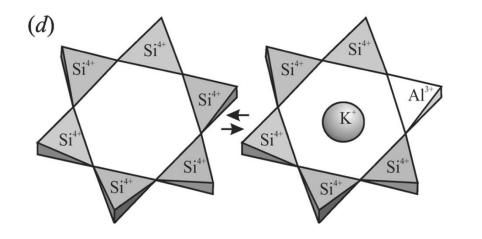


Ref: Chapter 4. Introduction to Mineralogy, William D. Nesse

#### Ommision substitution



Interstitial substitution



$$(n+1)M^{n+} \Leftrightarrow nM^{(n+1)+} + \square$$

$$3Fe^{2+} \Leftrightarrow 2Fe^{3+} + \square$$
(Pyrrhotite FeS)

Upto 13% of the sites may be vacant in Pyrrhotite

M<sup>n+</sup> and M<sup>(n+1)+</sup>: two different
cations substituting for each other,
n: the charge of lower charged
cation,

[]: a vacant site that normally would be occupied by M<sup>n+</sup>

$$\Box + \operatorname{Si}^{4+} \Leftrightarrow \operatorname{Al}^{3+} + (\operatorname{K}^+, \operatorname{Rb}^+, \operatorname{Cs}^+) \text{ or }$$

$$\Box + \operatorname{Si}^{4+} \Leftrightarrow \operatorname{Be}^{2+} + 2(\operatorname{K}^+, \operatorname{Rb}^+, \operatorname{Cs}^+)$$

## Mineral formula

- Cations are written first, followed by the anion(s) or anionic group.
- Charges must balance. The total charge of cations must equal the total charge of anions.
- Cations in the same structural site are grouped together.
- Cations in different structural sites are listed in order of decreasing coordination number.

#### Diopside (CaMgSi<sub>2</sub>O<sub>6</sub>)

Cation	Charge	Stoichiometric Coefficient	Total Charge <sup>a</sup>
Ca	2+	1	+2
Mg	2+	1	+2
Si	4+	2	+8
O	2-	6	-12
Total			0

Coordination	Common Cations	
12	K+, Ca <sup>2+</sup> , Na+	
	Fe <sup>2+</sup> , Ca <sup>2+</sup> ,	VIIICaVIMgIVSi <sub>2</sub> O <sub>6</sub>
8	Na+, Mg <sup>2+</sup>	Carling Si <sub>2</sub> O <sub>6</sub>
6	$Al^{3+}$ , $Fe^{2+}$ ,	
	$Fe^{3+}, Mg^{2+}$	Diopside
4	$Si^{4+}, Al^{3+},$	•
	$S^{6+}, P^{5+}$	(End member
3	C <sup>4+ a</sup>	composition)
2	None	Composition)

#### Representing mineral formula with solid solution

If Fe<sup>2+</sup> and Mg<sup>2+</sup> can readily interchange in the octahedral site

If the number  $Fe^{2+}$  of cations is x, then the formula can be written as:

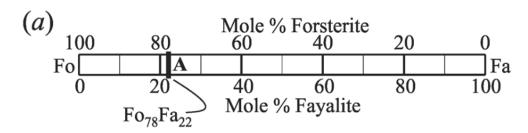
If x=22%, i.e. x=0.22, then the formula will be written as:

$$Ca(Mg_{0.78}, Fe_{0.22})Si_2O_6$$

In minerals with just two end members this is sometimes shortened to report just one end member (e.g.,Di<sub>78</sub>) because the sum of both end members (diopside plus hedenbergide in clinopyroxene) must equal 100%.

# Graphical representation of mineral formula

#### **Complete solid solution**



Trusted website for Mineral information: https://www.mindat.org/ https://webmineral.com/

> Wo,00 wollastonite CaSiO<sub>3</sub>

> > augite

pigeonite

50

CaFeSi<sub>2</sub>O<sub>6</sub>

Fe →

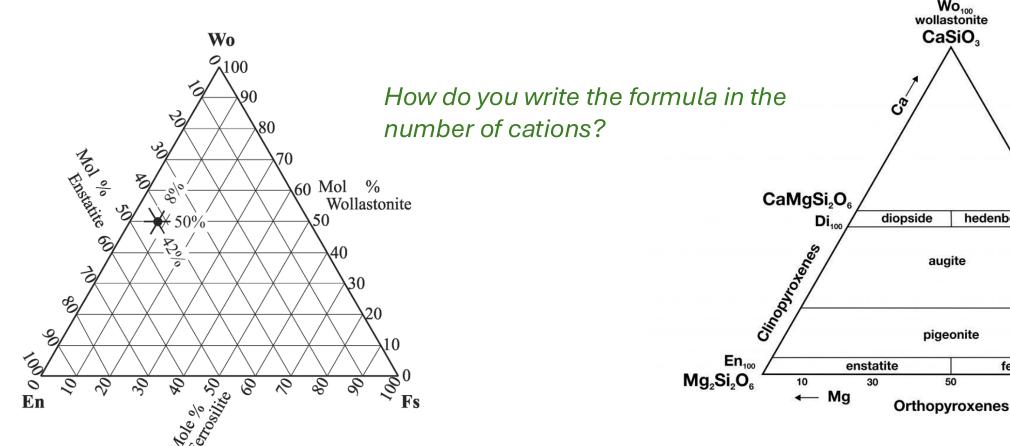
Fs<sub>100</sub>

Fe<sub>2</sub>Si<sub>2</sub>O<sub>6</sub>

Hd,00

hedenbergite

ferrosilite



# Thank

Minerals have a *definite* but *not fixed* chemical composition (as minerals are crystalline Variation in composition takes place keeping the crystal structure constant.

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- ✓ **Solid solution series:** The range of compositions produced by solid solution in a given mineral (a)
- ✓ *End members:* The compositional extremes of a substitution series
- ✓ A continuous or complete substitution series is one in which all intermediate compositions are possible.
- ✓ An incomplete or discontinuous substitution series is one in which only a restricted range of compositions between the end members is found.



Mole % Forsterite

Mole % Fayalite

40

Mg<sub>2</sub>SiO<sub>4</sub>

Fe<sub>2</sub>SiO<sub>4</sub>

100

