



# Optical Mineralogy

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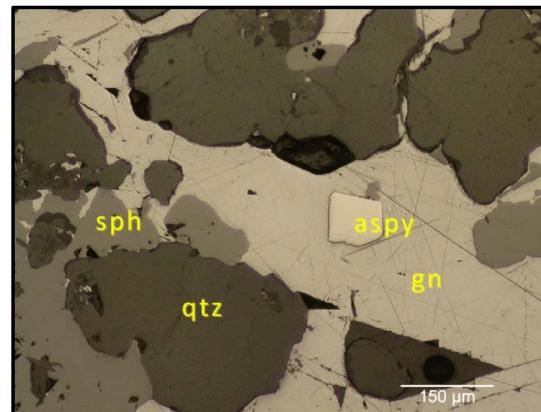
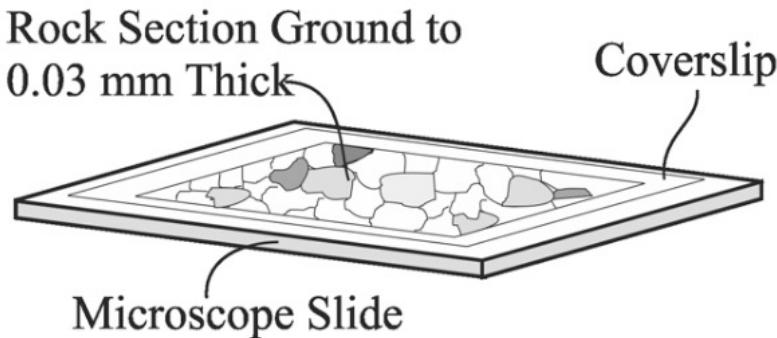
*Dr. Tapabrato Sarkar*

# Introduction

## □ What is optical Mineralogy?

Analysing thin slices of minerals or rocks under transmitted or reflected light.

- **Thin sections** are thin slices of rock or mineral mounted on a microscope slide.
- They are prepared by cementing a piece of rock to a microscope slide and then grinding it to its final thickness, usually **0.03 mm**.



Reflected light



Transmitted light

# Petrographic Microscope

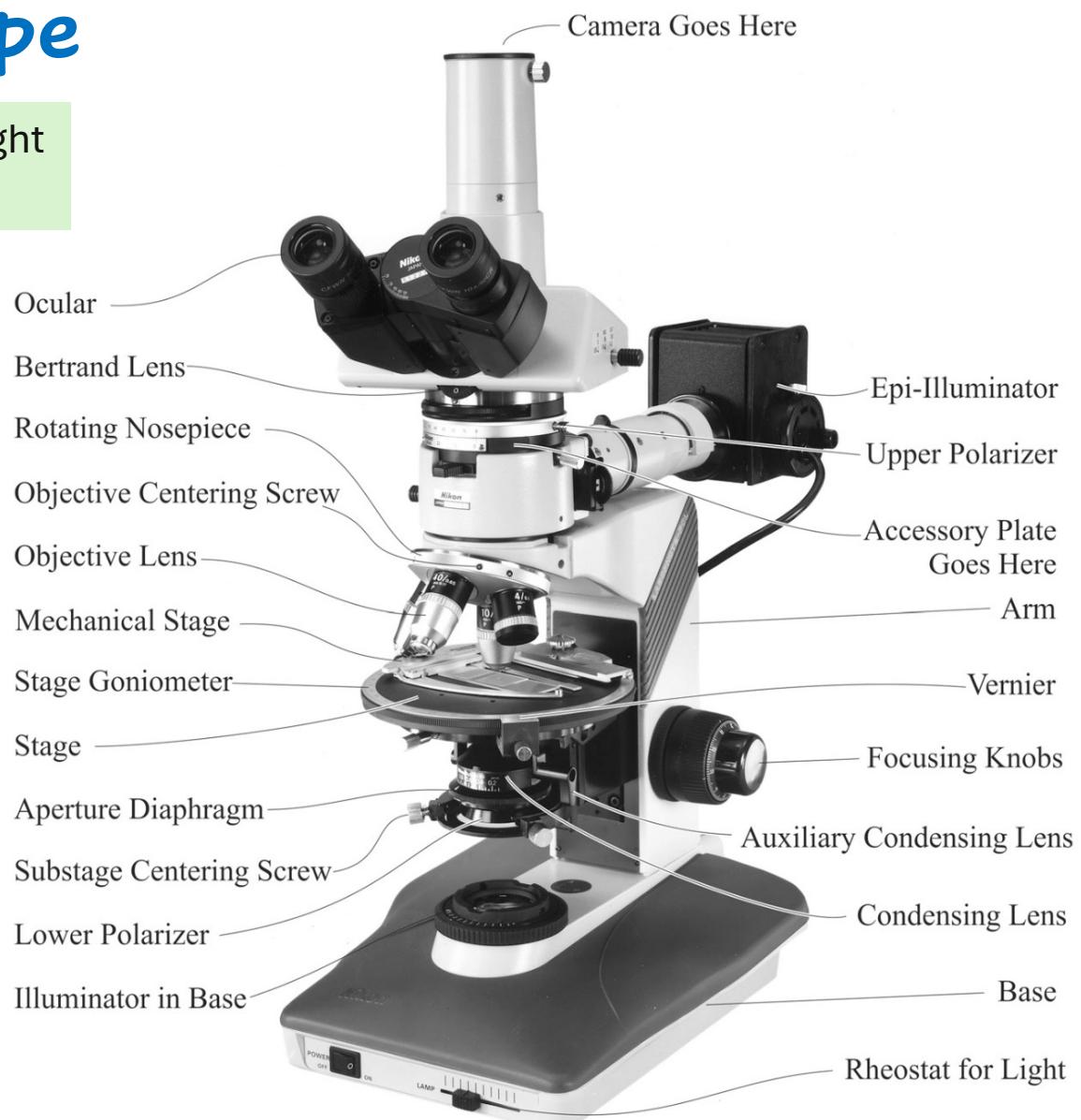
A specialized instrument that utilizes polarized light to allow measurement of a variety of properties

These properties are called **optical properties**.

Petrographic microscope provides one of the primary means of studying minerals and the rocks they comprise.

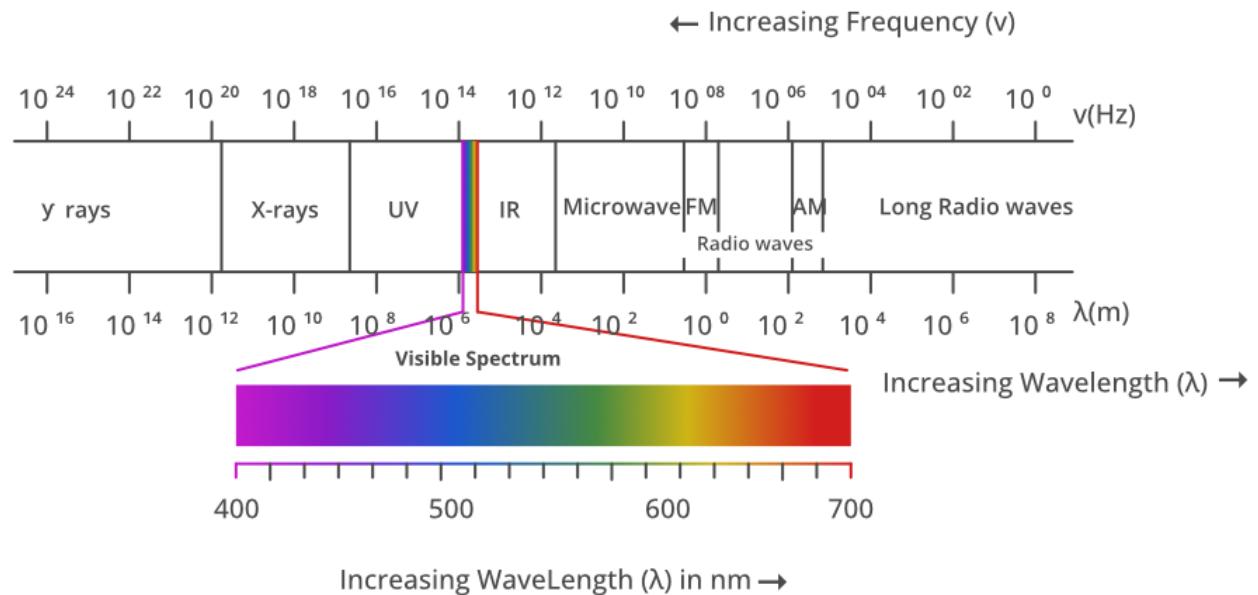
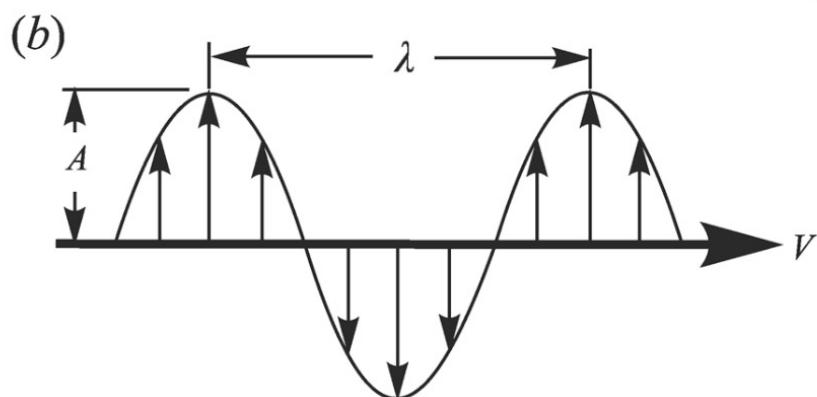
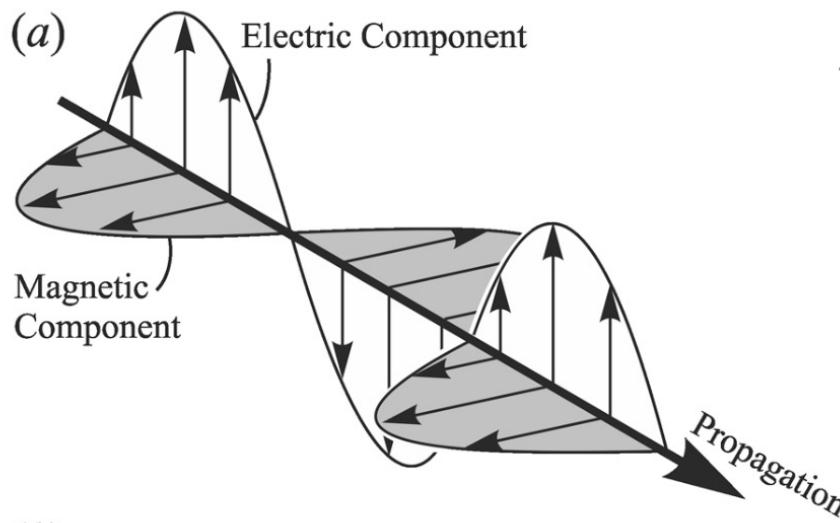
## ■ Why do we require a microscope?

- ✓ Grain size too small to observe the properties in hand specimen
- ✓ Microscopic properties which cannot be observed in a hand specimen
- ✓ Relation between the different mineral grains



# Basics of optics

For optical mineralogy, we shall treat light as a wave phenomenon

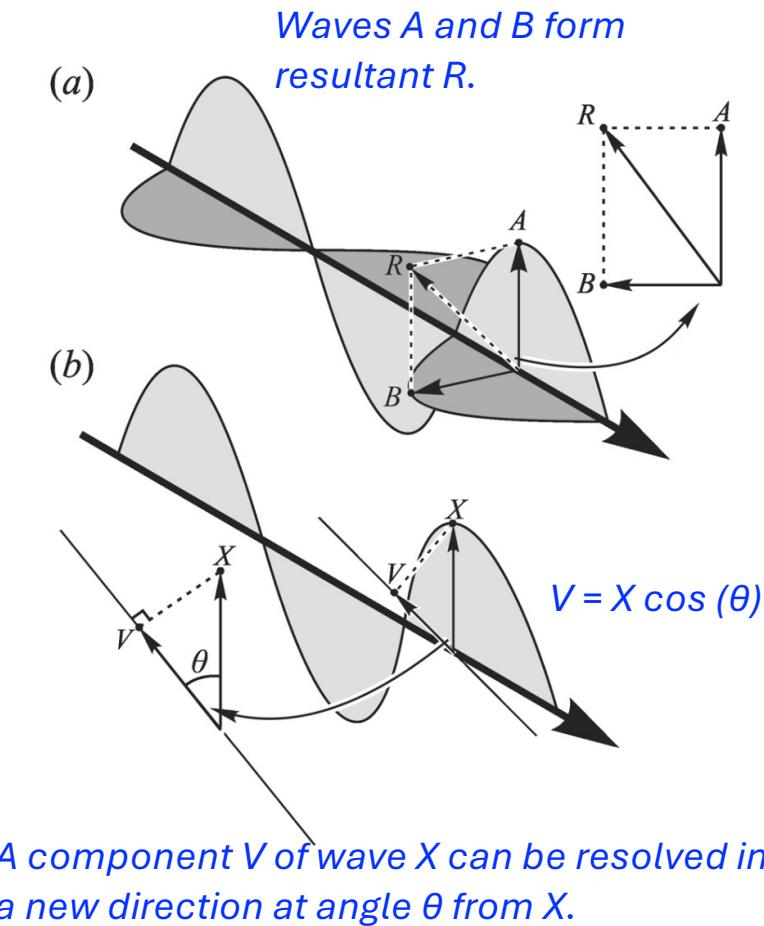


$$V = f\lambda$$

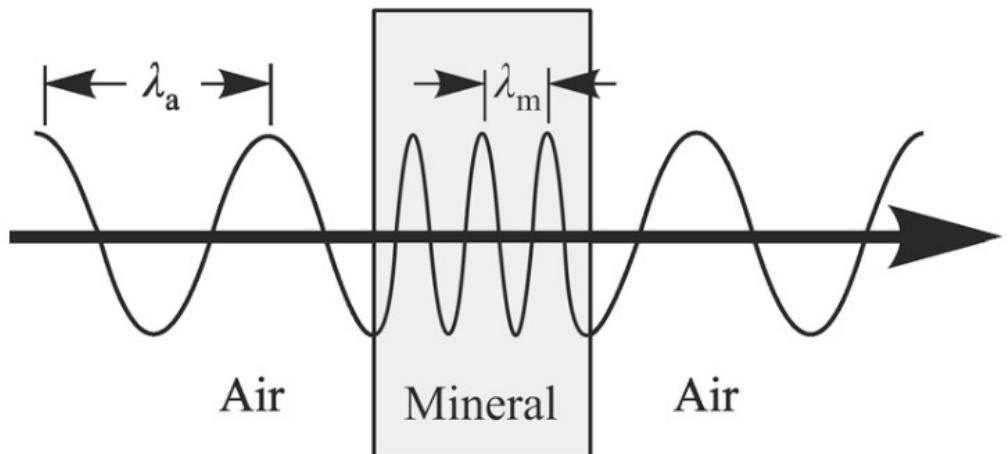
$V$ : velocity (m/s)  
 $f$ : frequency (cycles/s)  
 $\lambda$ : wavelength (m)

# Basics of optics

## Vector resolution of light waves.



## Interaction of light with mineral



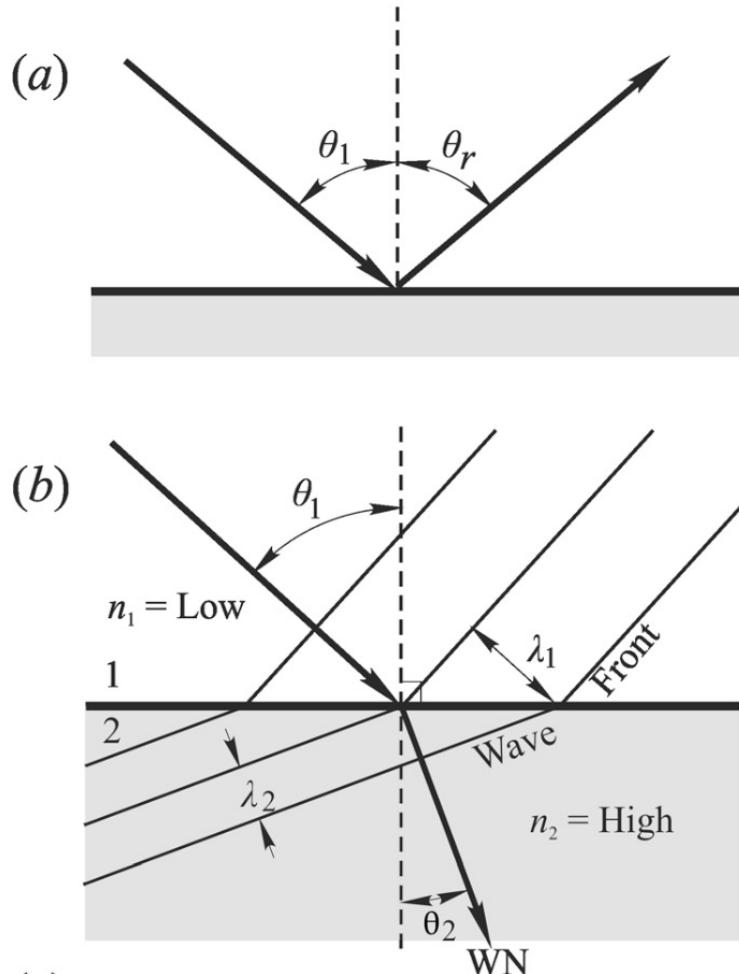
- ✓ Light slows down and wavelength decreases
- ✓ Frequency remains the same

## Refractive index

$$n = \frac{V_a}{V_m} = \frac{f\lambda_a}{f\lambda_m} = \frac{\lambda_a}{\lambda_m}$$

# Basics of optics

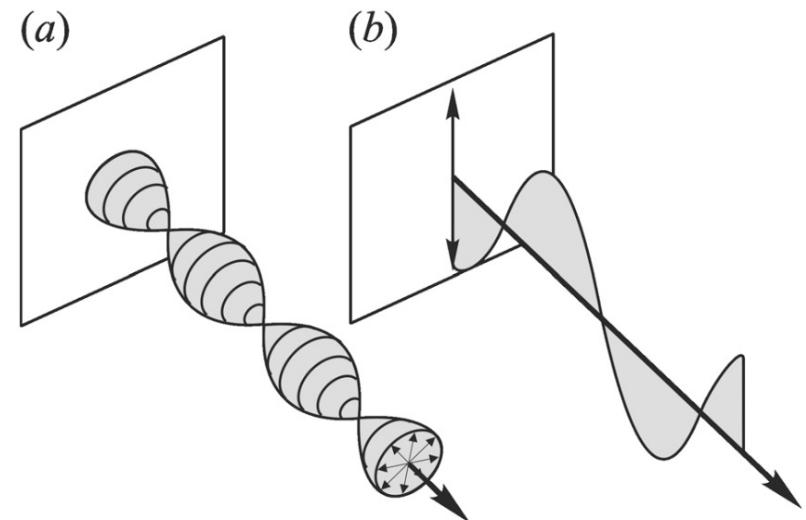
## Reflection and refraction



Snell's Law

$$\frac{\sin \theta_1}{\sin \theta_2} = \frac{n_2}{n_1}$$

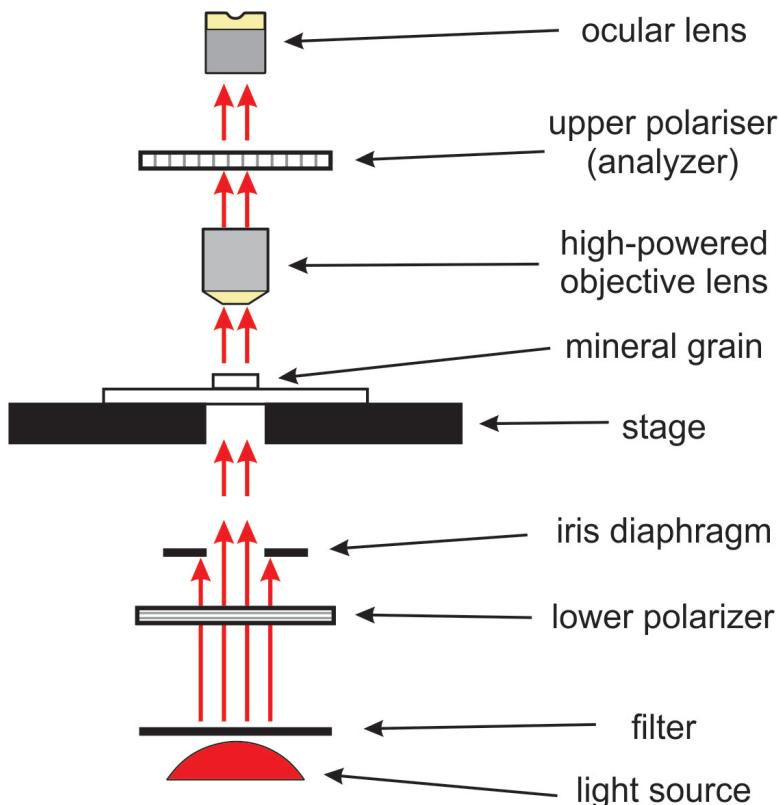
## Polarized light



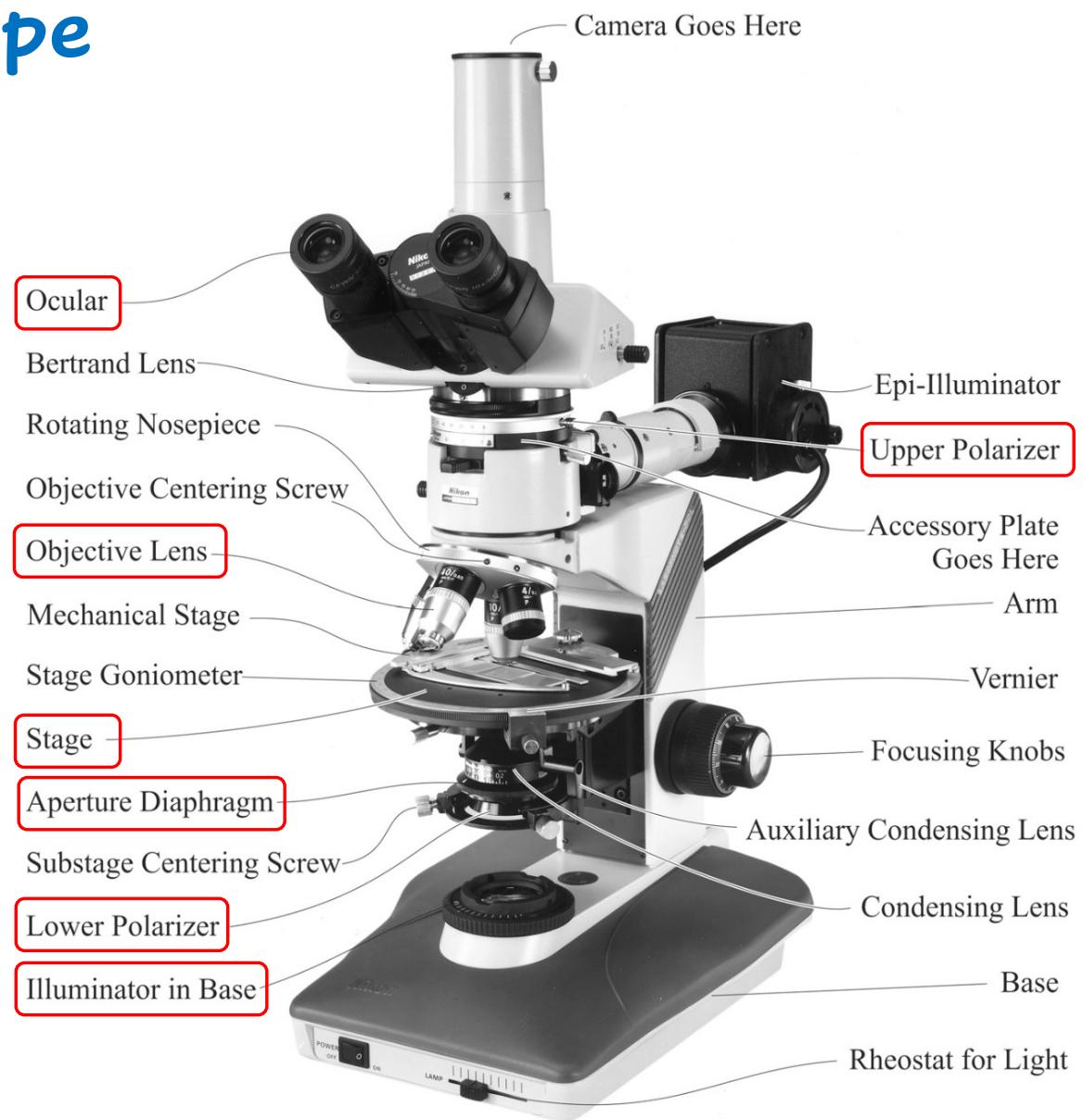
- ✓ Ordinary light coming directly vibrates in all directions at right angles to the direction of propagation – **Unpolarized light**.
- ✓ If the vibration of the light is constrained to lie in a single plane – **Plane polarized**.
- ✓ Vibration can be represented by a simple **sine wave**.

# Petrographic Microscope

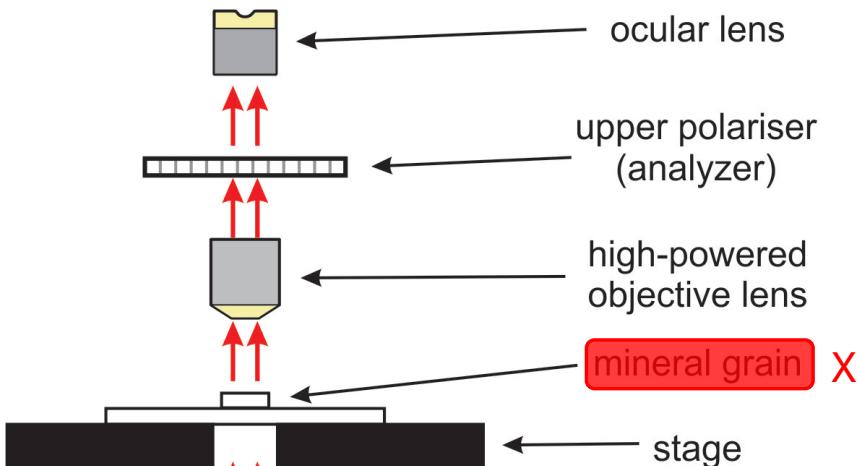
## Fundamentals of a petrographic microscope



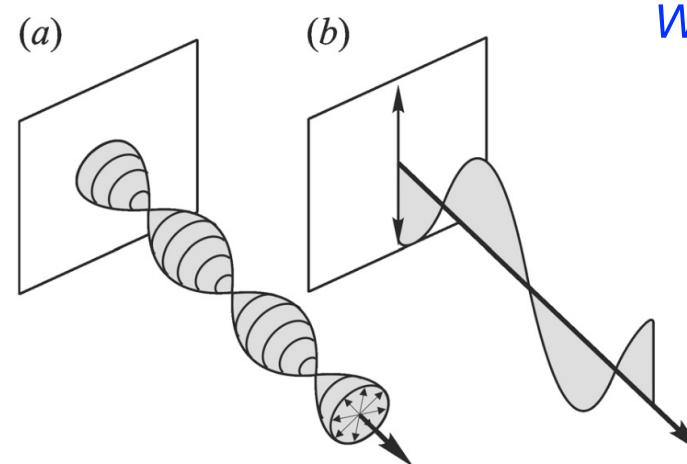
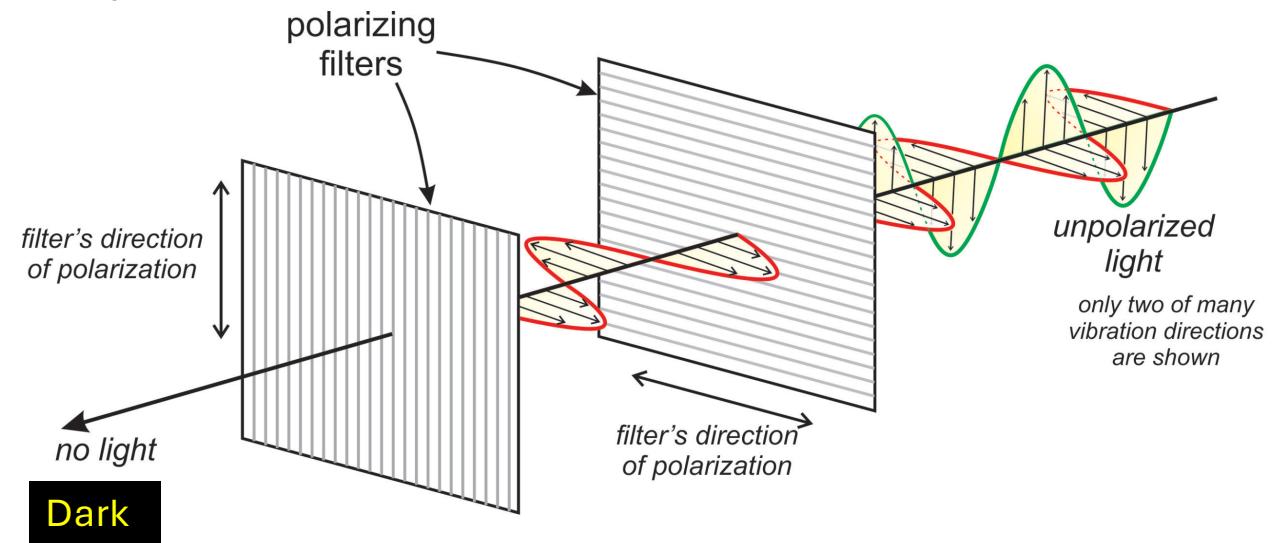
standard  
orthoscopic  
illumination



# Petrographic Microscope

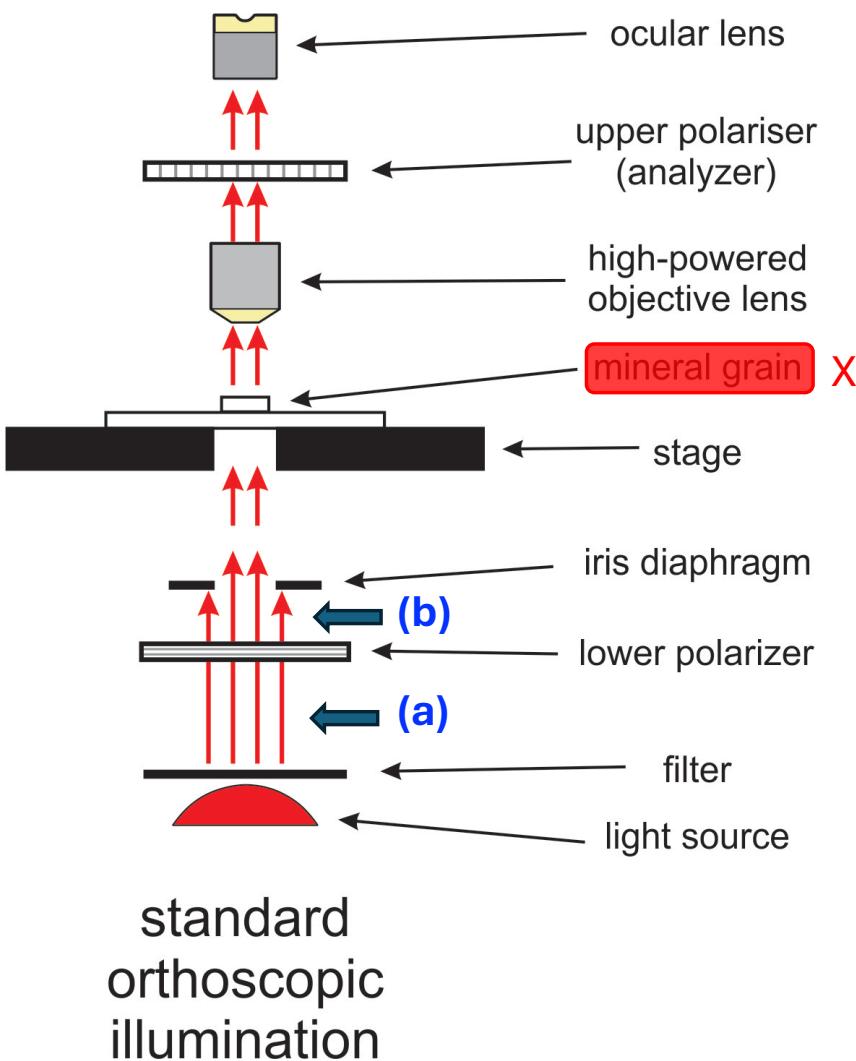


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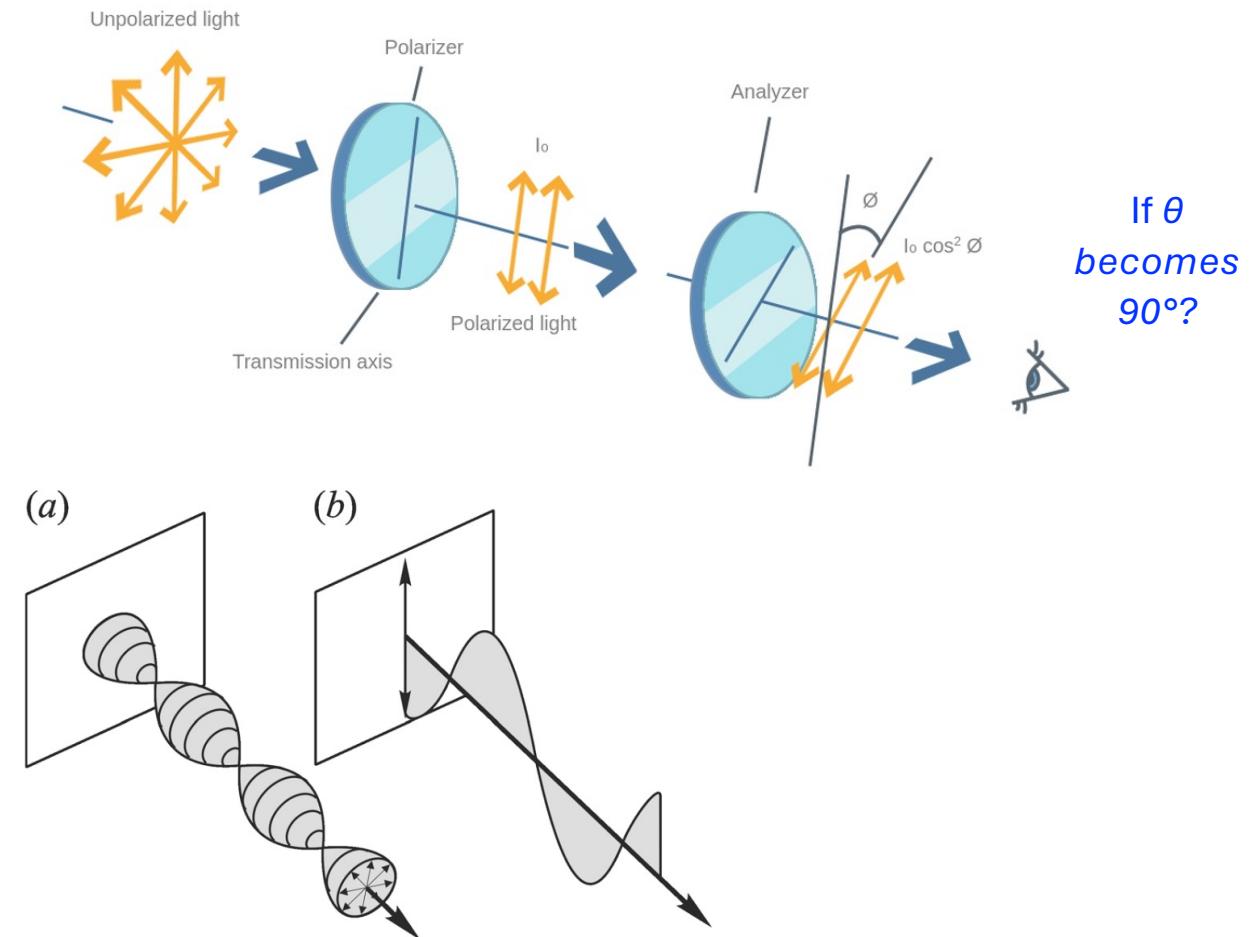


*What happens when light passes through a microscope without a mineral grain?*

# Petrographic Microscope



**What happens if the two polarizers are not perpendicular?**



# Isotropic and anisotropic mineral

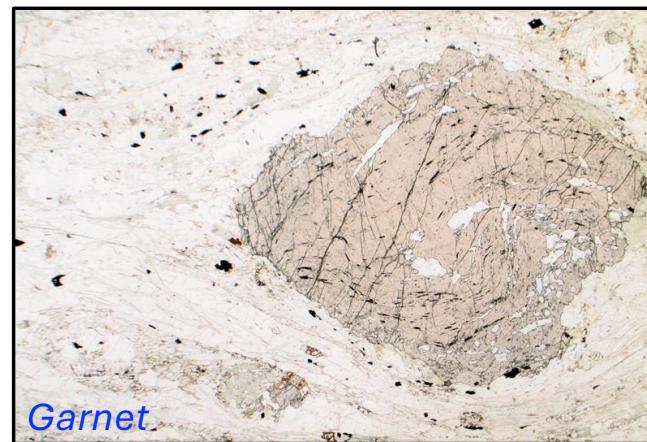
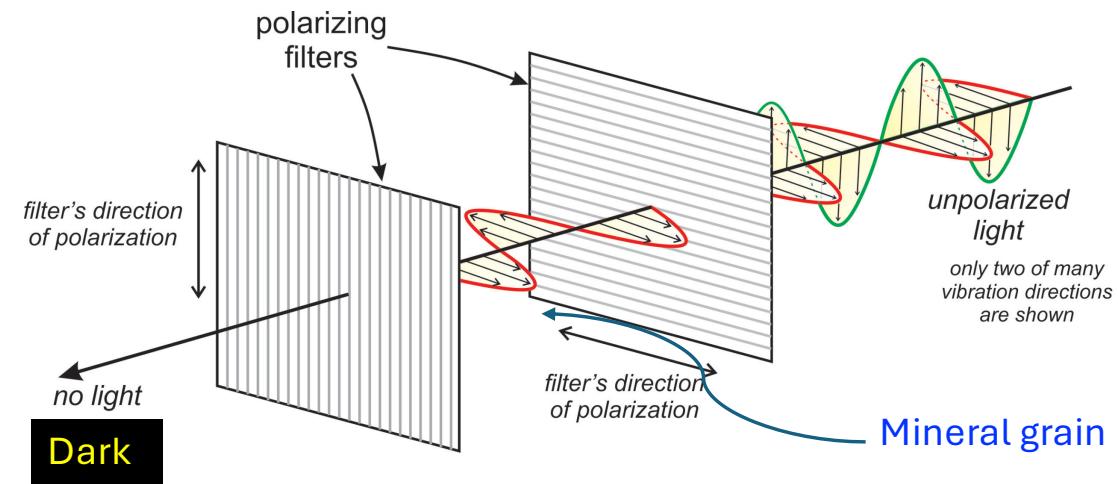
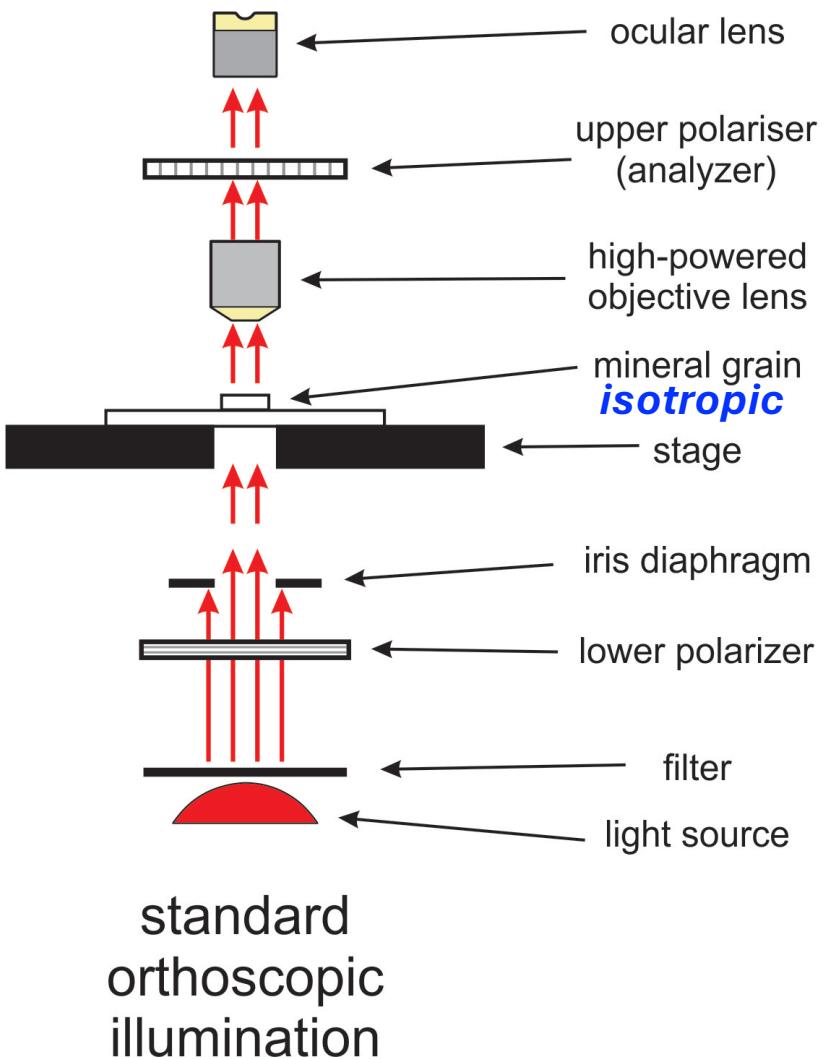
## □ Isotropic mineral

- An **optically isotropic** material is one in which the velocity of light is the same in all directions.
- The isotropic rock-forming materials include volcanic glass and minerals belonging to a specific crystal system.
- In these materials, electron density is the same in all directions, at least on average.
- The strength of the electric field with which the electric vector of light interacts also is the same regardless of direction.
- Light velocity is same in all directions.

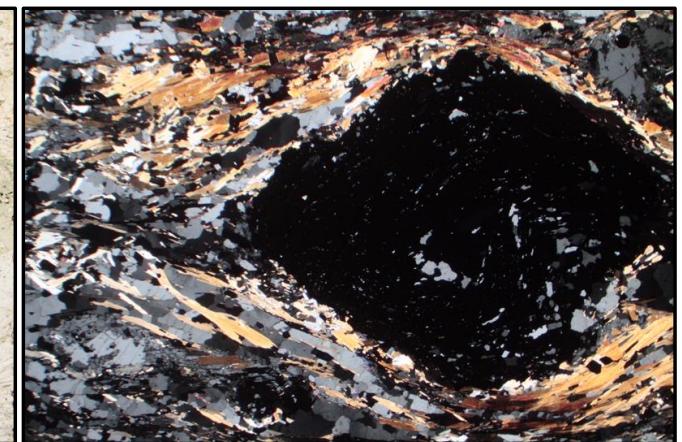
## □ Anisotropic mineral

- An **optically anisotropic** material is one in which the velocity of light is different in different directions.
- The electron density varies with direction.
- The electrons of the atoms/ions in these minerals are not able to interact with light in the same way in all directions.
- The velocity and absorption characteristics (color) of light vary with direction.

# Isotropic mineral

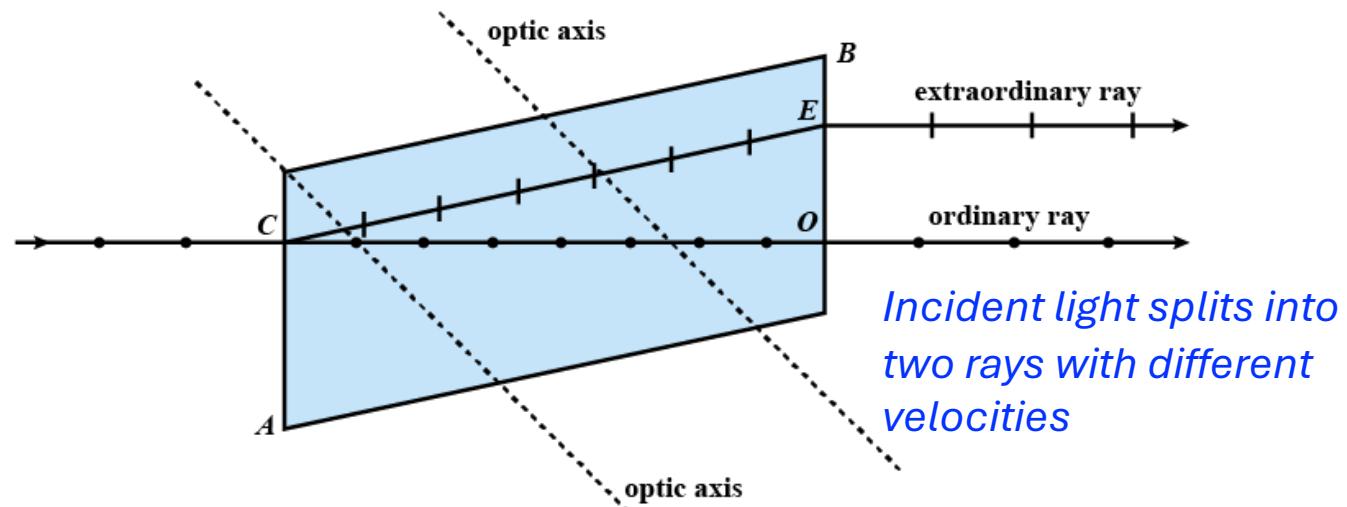
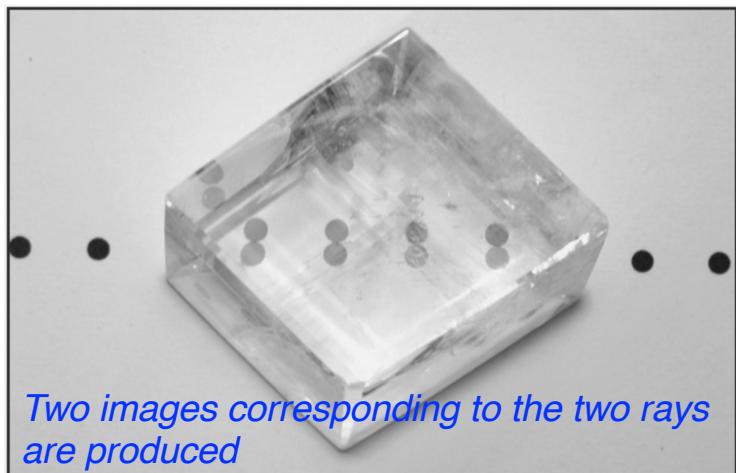


Plane Polarized Light (PPL)



# Light through an anisotropic mineral

Anisotropic minerals show double refraction

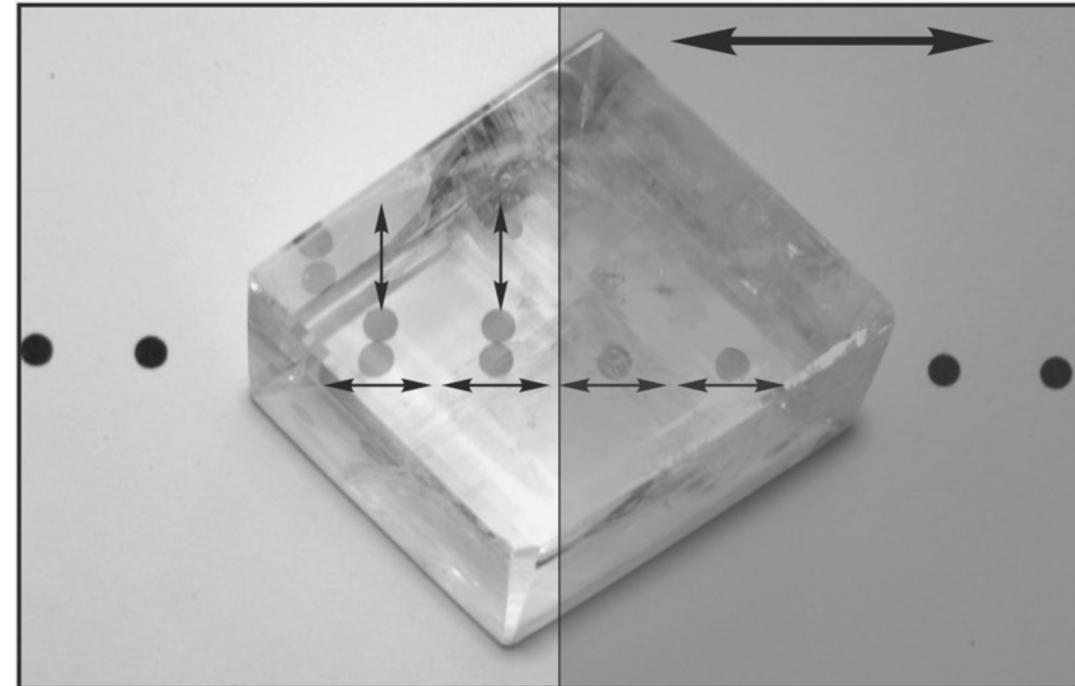
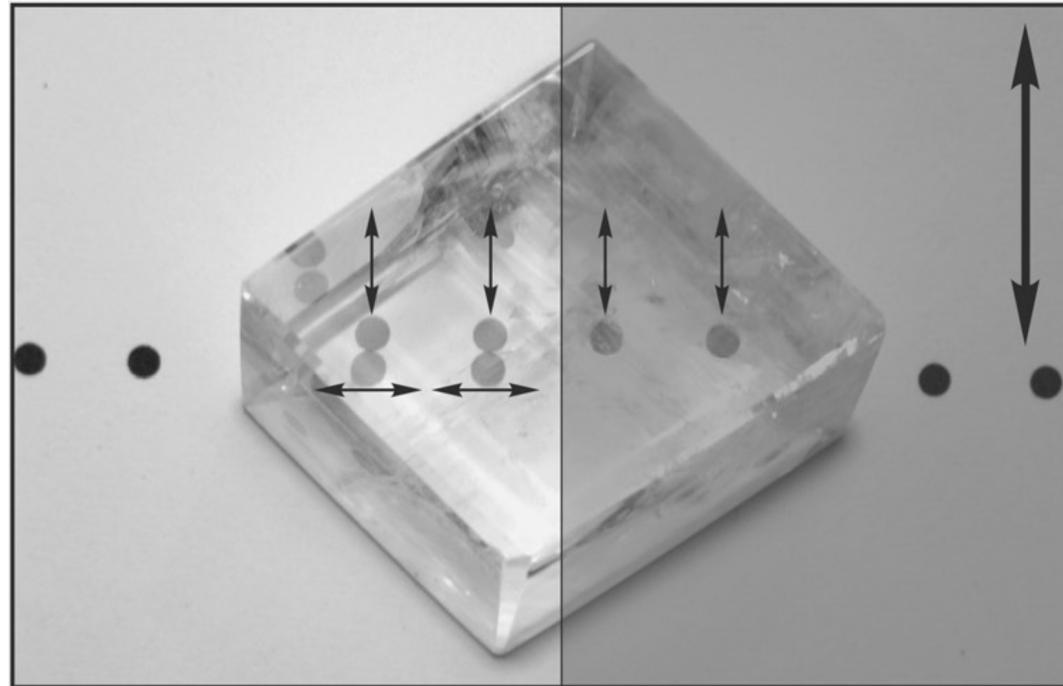


Incident light splits into two rays with different velocities



- **Ordinary ray (o-ray)** follows Snell's law. Its velocity is same in every direction
- **Extraordinary (e-ray)** does not obey Snell's law. Its velocity is different in different directions.
- These e- and o-rays are vibrating at **right angles** to each other.
- Every anisotropic mineral has one or two directions, called **optic axes**, along which the light is **not** split into two rays.

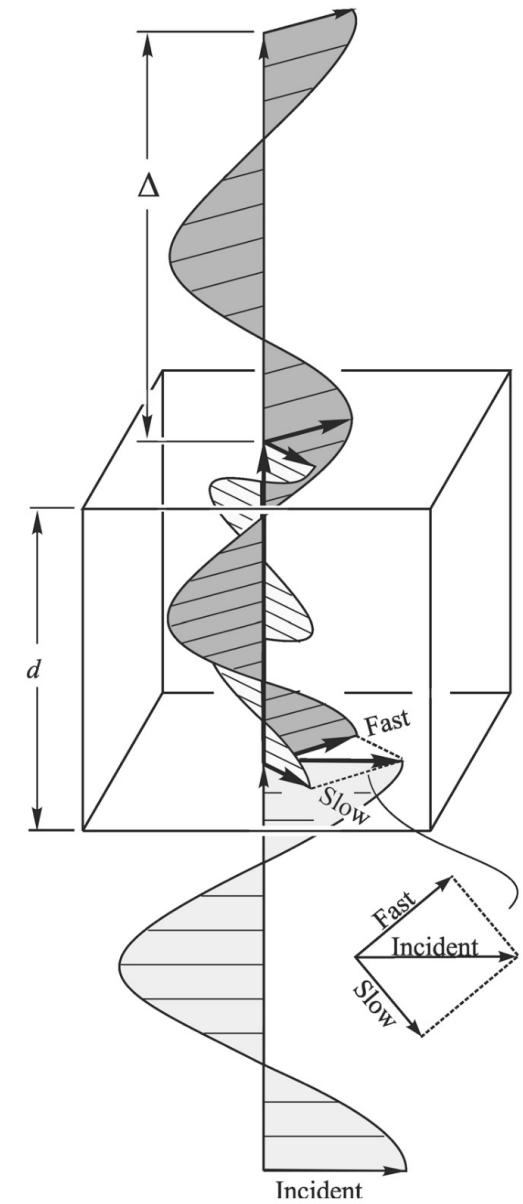
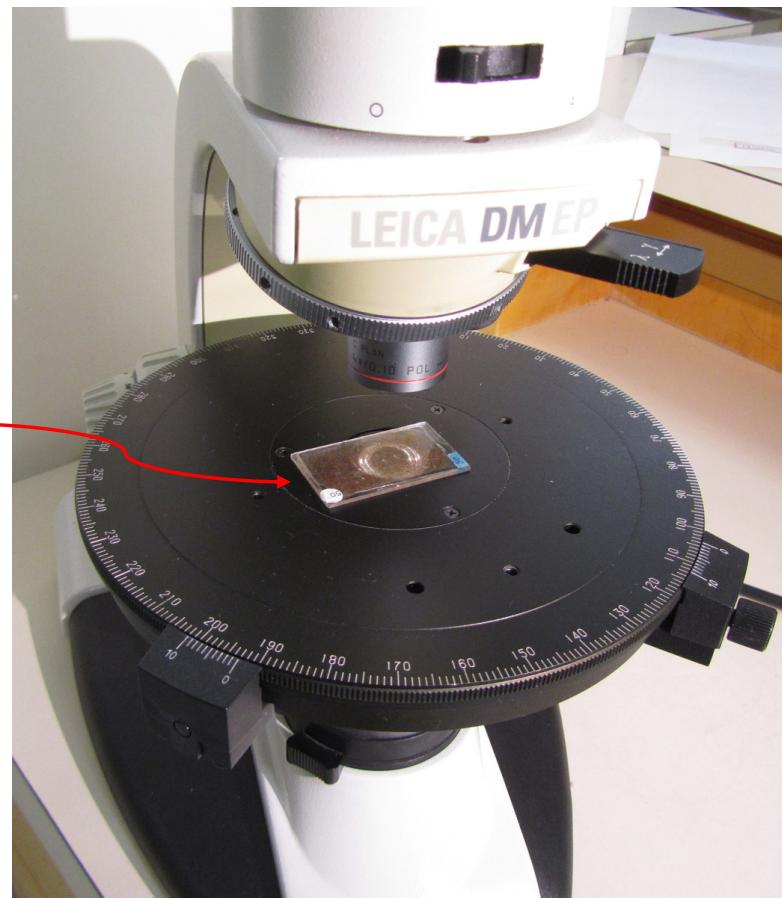
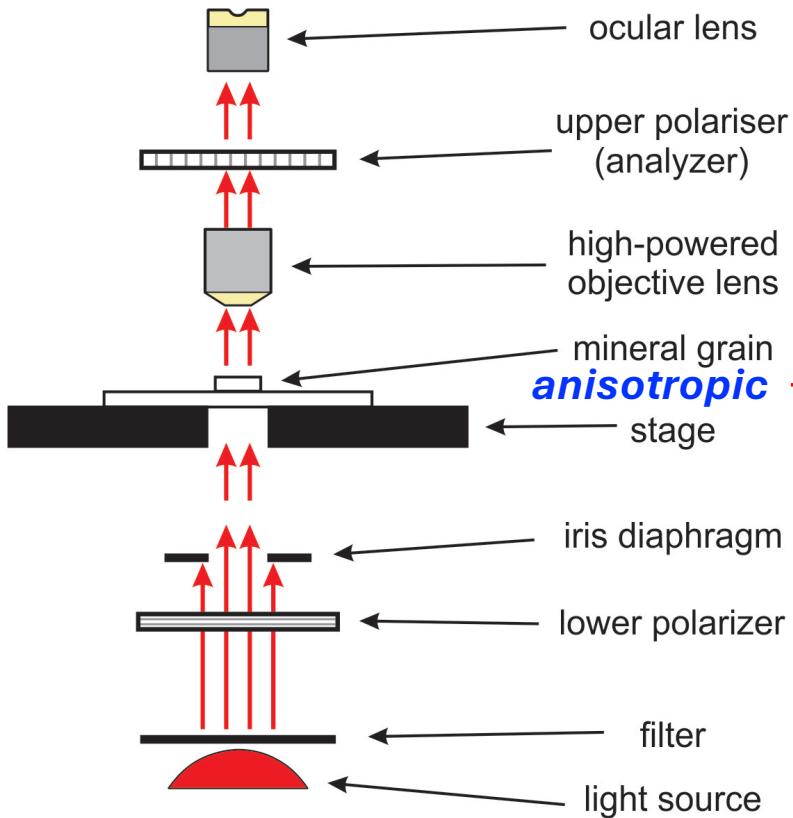
# Double refraction in calcite



- A polarizing film whose vibration direction is parallel to the short diagonal of the rhomb passes one set of dots and absorbs the other.
- The polarizing film is rotated 90°.
- The first set of dots is absorbed and the other passes.
- In intermediate orientations, both sets of dots are visible with subdued brightness.

***The two rays must therefore be plane polarized and vibrating at right angles to each other.***

# Double refraction



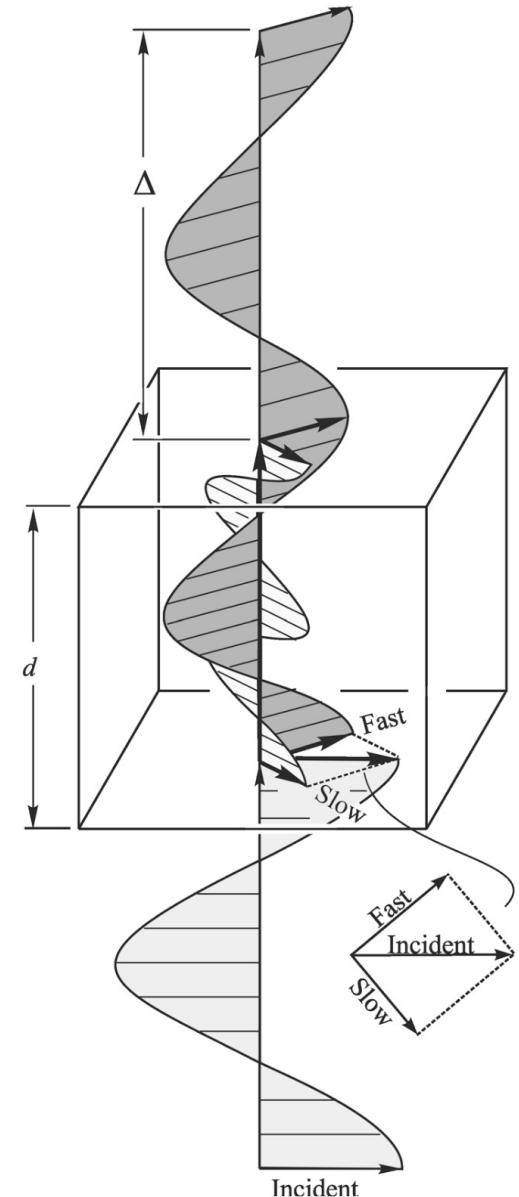
# Double refraction

- If the light velocity of the two rays is determined by measuring indices of refraction, it will be found that one ray is **faster** than the other.
- The ray with the **greater velocity** ( $V_f$ ) and **lower refractive index** ( $n_f$ ) is called the **fast ray**
- The ray with the **lower velocity** ( $V_s$ ) and **higher refractive index** ( $n_s$ ) is called the **slow ray**.
- The two rays vibrate **perpendicular** to each other.

$$n_f = \frac{V_0}{V_f}$$

$$n_s = \frac{V_0}{V_s}$$

- Due to the difference in index of refraction, the slow ray lags behind the fast ray.

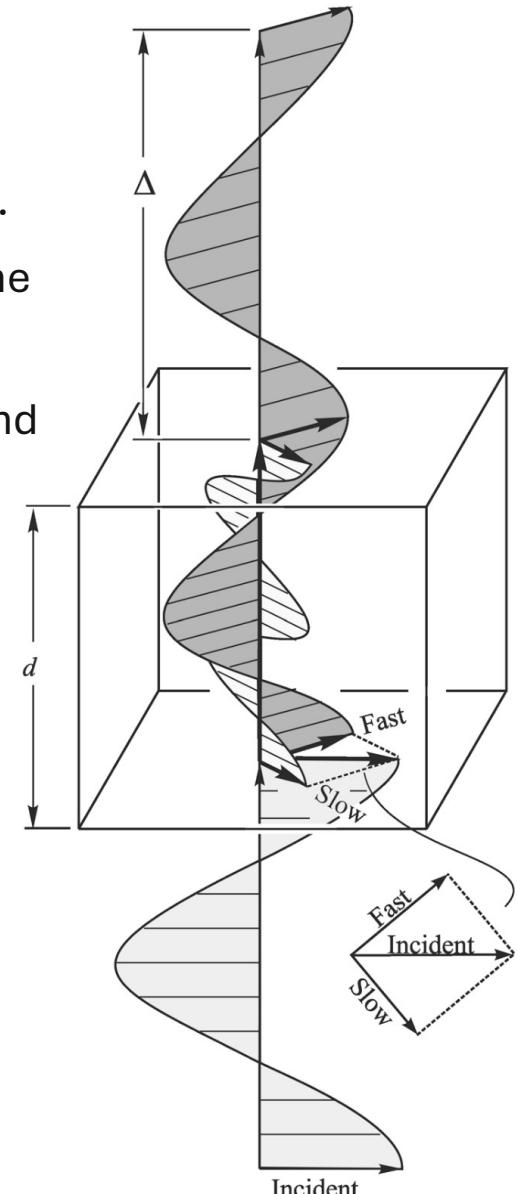


# Double refraction

- During the *time taken by slow ray* to pass through the mineral, the fast ray will pass through the mineral plus an *additional distance* called the **retardation ( $\Delta$ )**.
- The *retardation remains constant* after the slow and fast rays have exited into the air above the mineral grain because both have the same velocity there.
- The magnitude of the retardation depends on the thickness of the mineral ( $d$ ) and on the difference in index of refraction of the slow ray ( $n_s$ ) and fast ray ( $n_f$ ) in the mineral.

$$\Delta = d(n_s - n_f) = d(\delta)$$

- ( $\delta$ ) is the birefringence equal to the value ( $n_s - n_f$ ).
- The numerical value of birefringence depends on the direction followed by the light through the mineral.
- Parallel to an optic axis: zero birefringence (Mineral appears dark).
- Other directions show a maximum birefringence, and **most show an intermediate value**.
- **The maximum birefringence is a useful diagnostic property of minerals.**



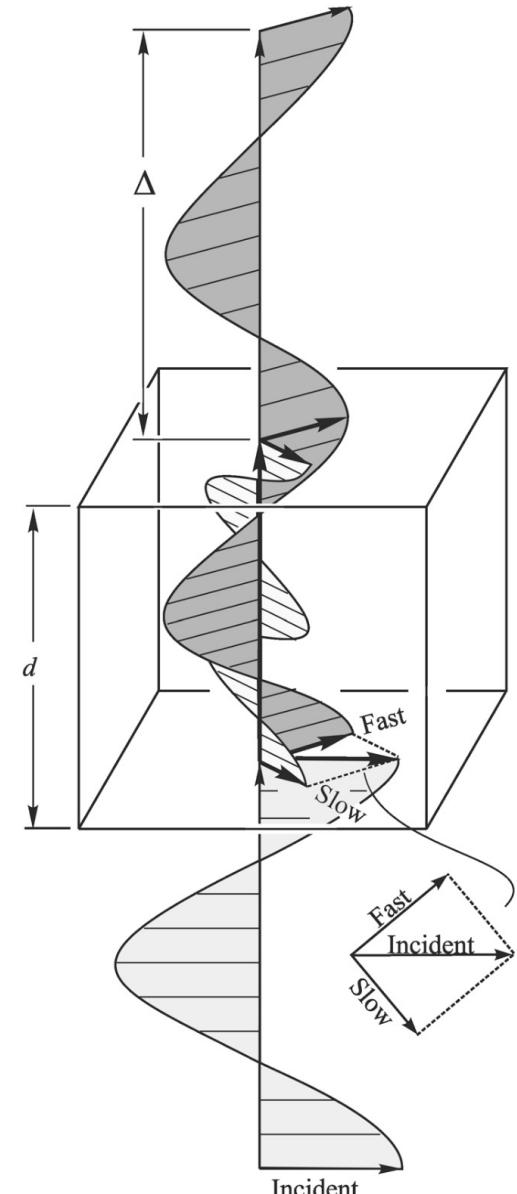
## Double refraction

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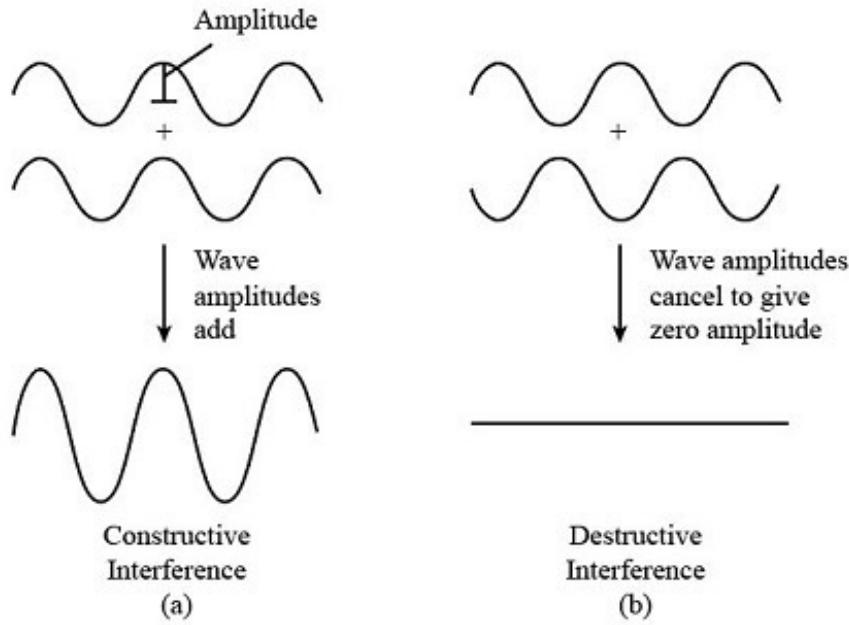
$$\Delta = d(n_s - n_f) = d(\delta)$$

$$t = \frac{d}{v_s} = \frac{d}{v_f} + \frac{\Delta}{v_0} \Rightarrow \Delta = d \left( \frac{v_0}{v_s} - \frac{v_0}{v_f} \right),$$

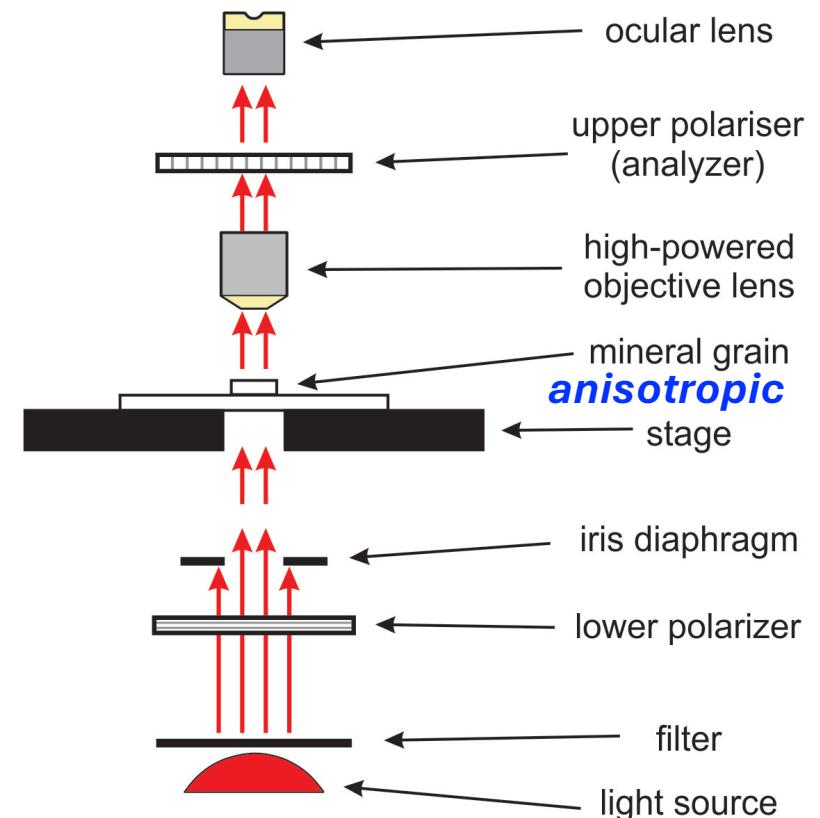
$$\Delta = d(n_s - n_f)$$



# Interference of light waves



**Interference color:** The color seen between crossed polarizers, produced as a consequence of light being split into two rays on passing through the mineral. Components of the two mutually perpendicular rays interfering along a single plane.



# Interference of light waves

- Case 1: the slow ray is retarded an integer number of wavelengths relative to the fast ray

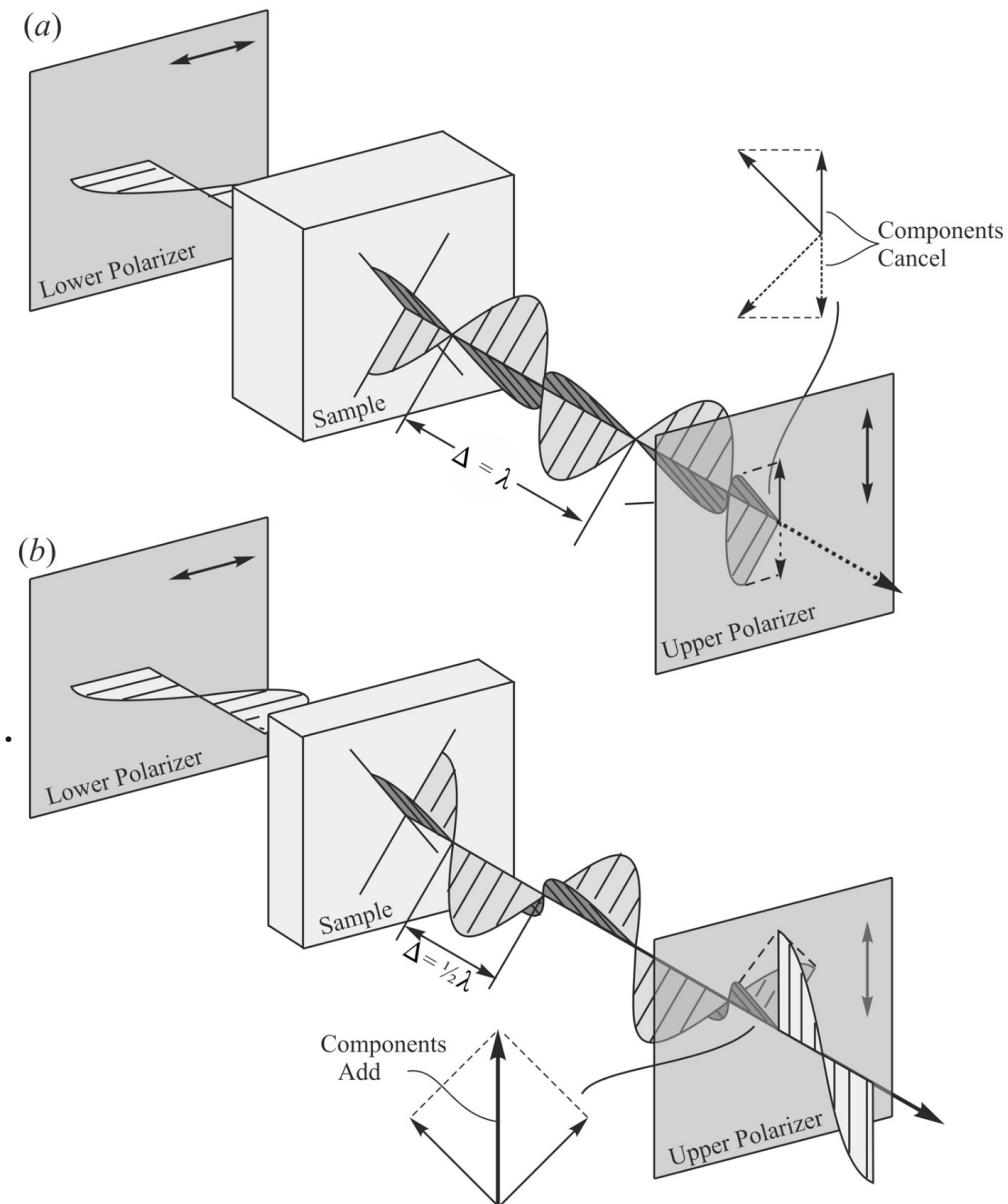
$$\Delta = i\lambda \quad i : \text{integer}; \lambda : \text{wavelength}$$

- Components of the two rays resolved into the vibration direction of the upper polarizer  $\rightarrow$  equal magnitude; opposite directions  $\rightarrow$  cancel each other
- No light passes the upper polarizer  $\rightarrow$  mineral grain appears dark.

- Case 2: the retardation is equal to one-half wavelength.

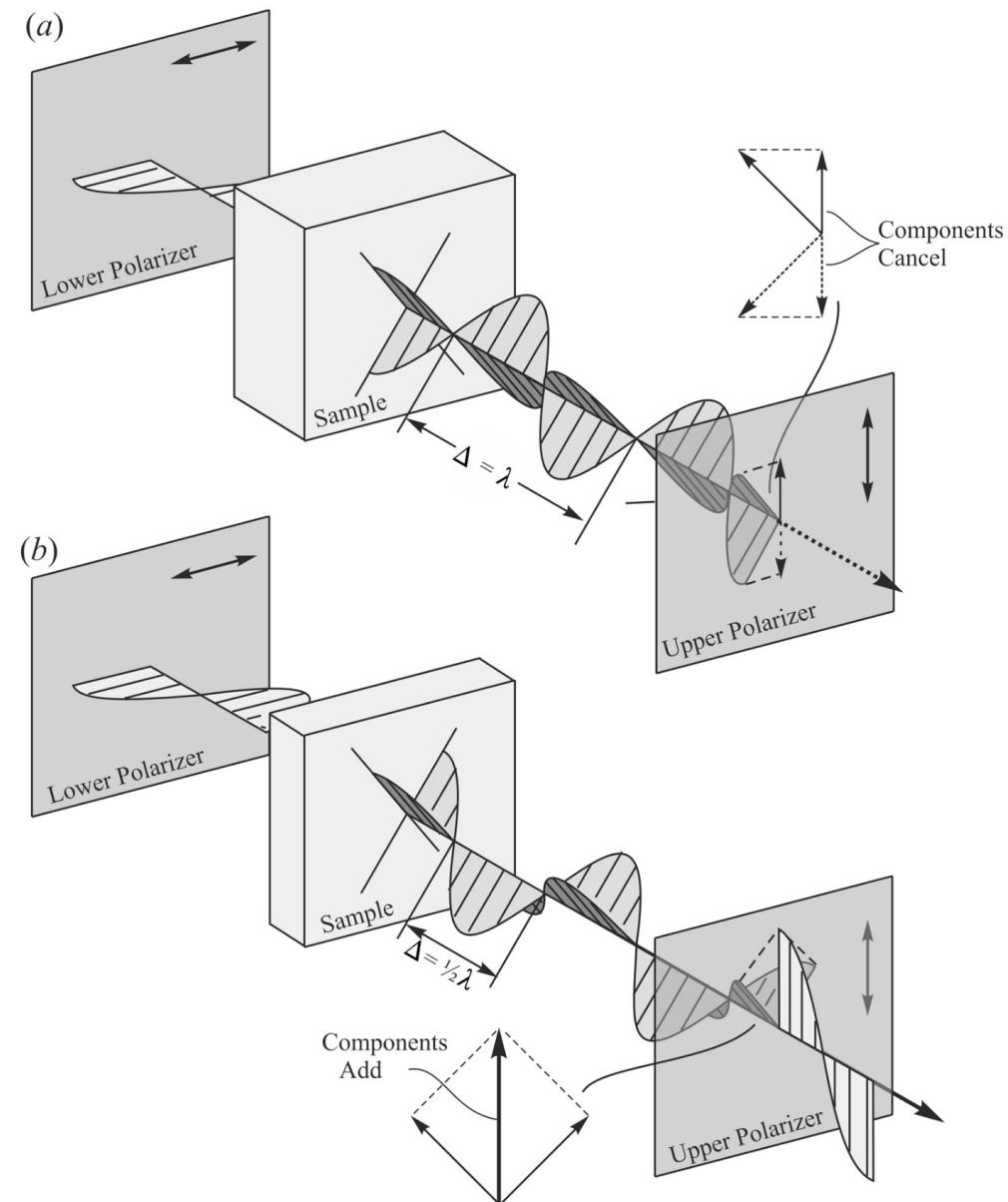
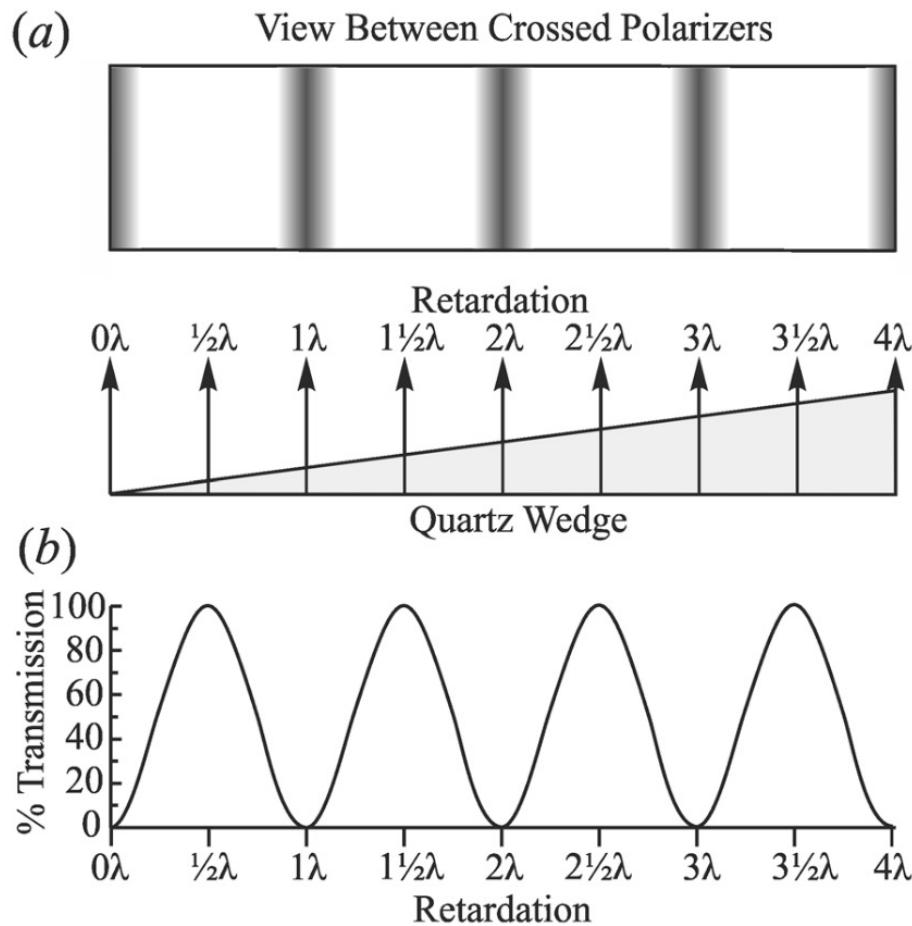
$$\Delta = \left(i + \frac{1}{2}\right)\lambda \quad i : \text{integer}; \lambda : \text{wavelength}$$

- The resolved components are both in the same direction  $\rightarrow$  light constructively interferes  $\rightarrow$  light passes the upper polarizer.
- The observed color is the interference color



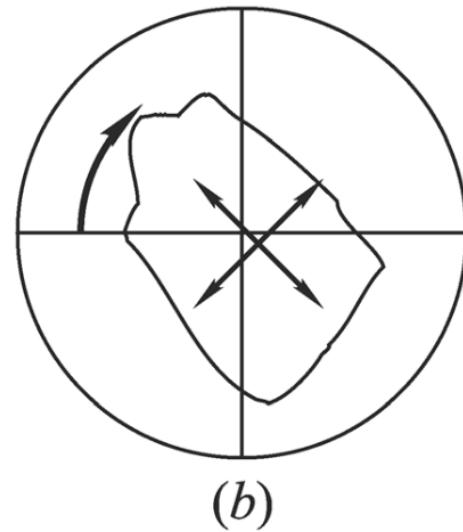
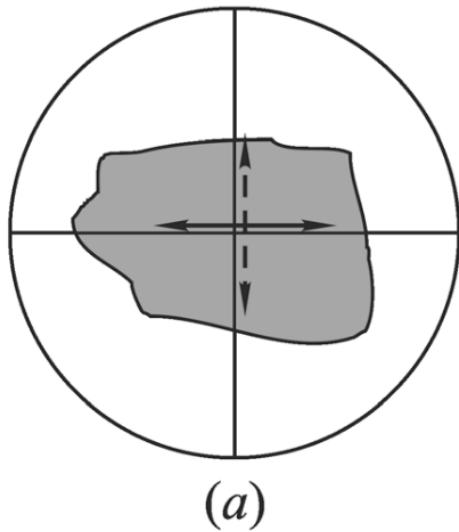
# Interference of light waves

## □ Viewing a quartz wedge under crossed polars



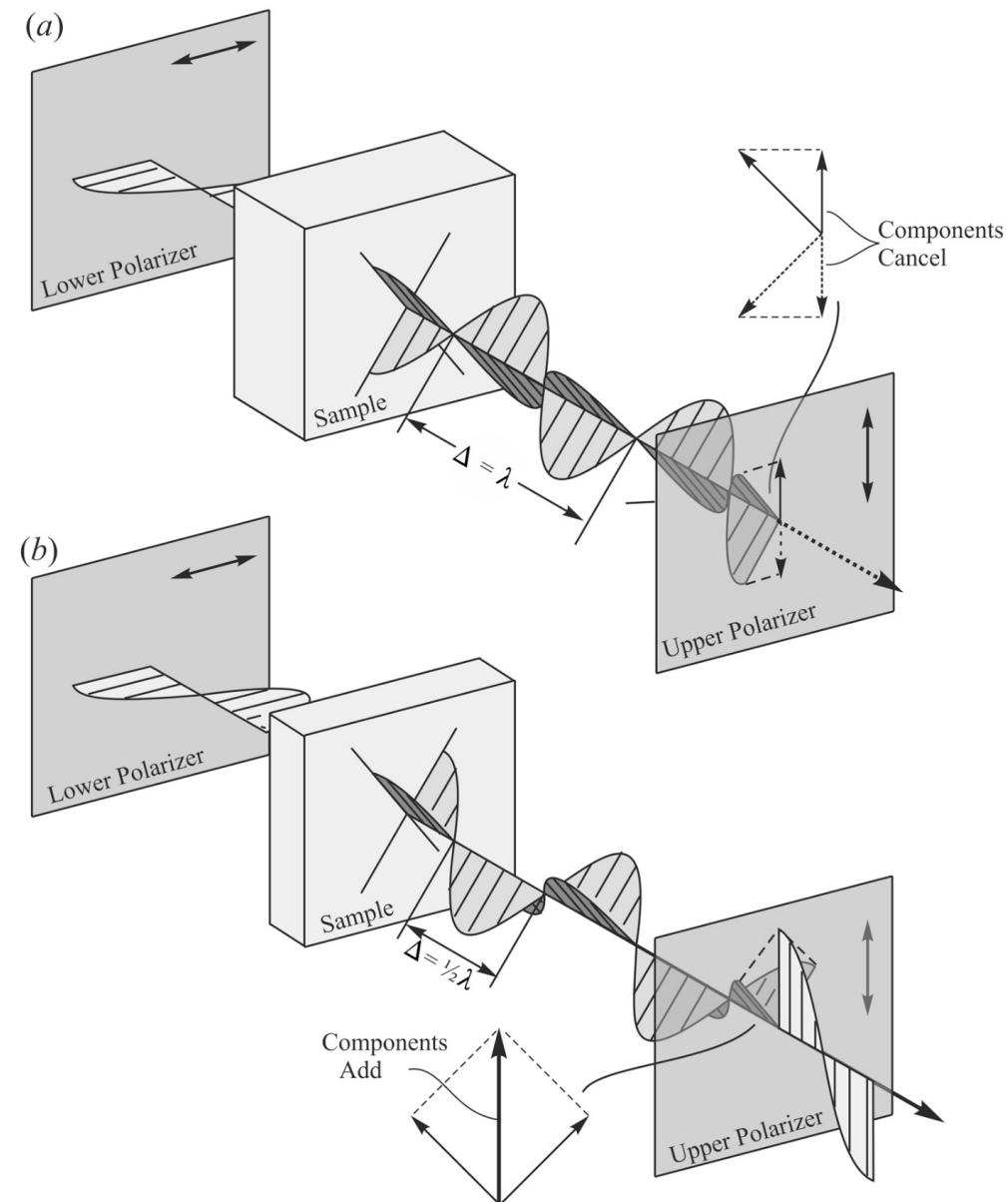
# Extinction

**What happens when we rotate the stage?**



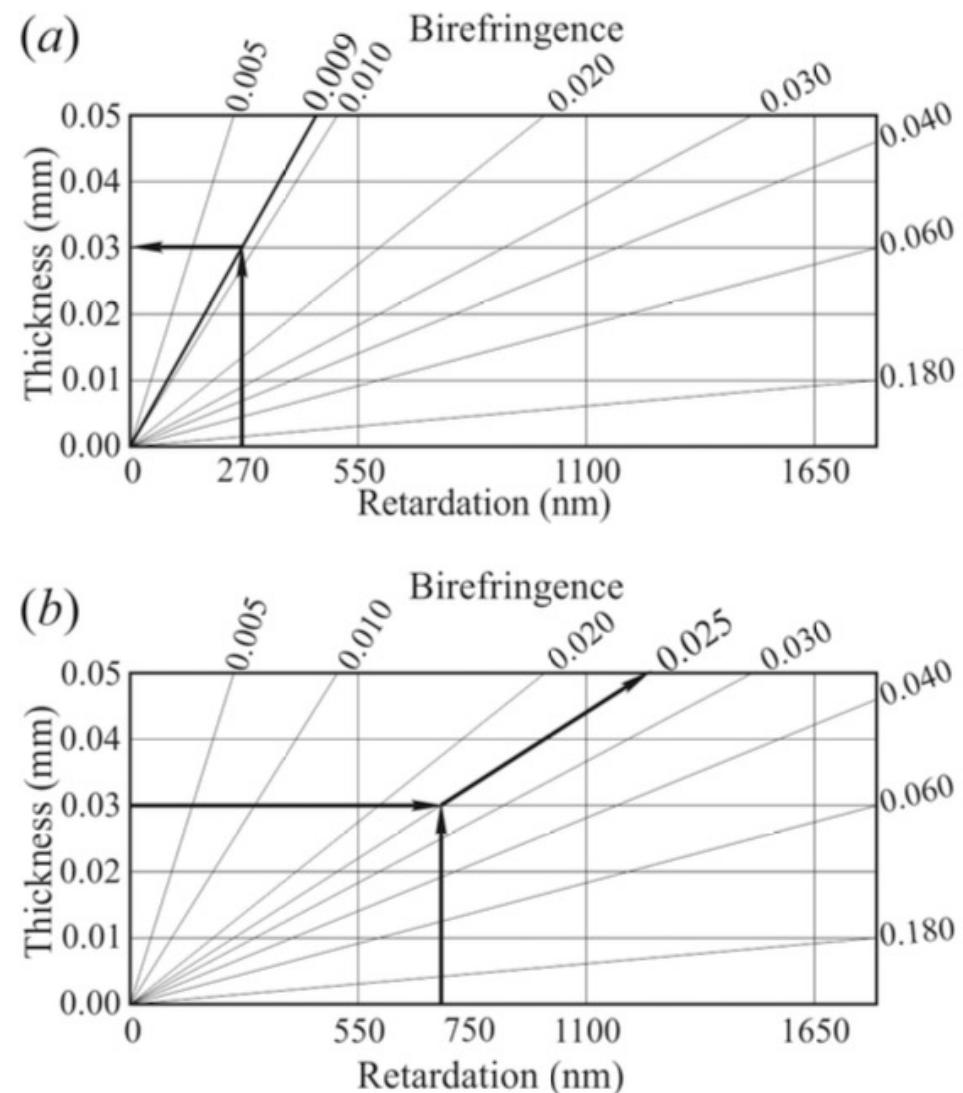
All anisotropic minerals go dark or extinct between crossed polarizers once in every  $90^\circ$  of stage rotation.

The interference color does not change with rotation.  
It gets brighter and dimmer, because the retardation  
between the two rays remains constant.



# Use of interference color chart

- The interference color chart shows the interference colors produced for **retardations (path differences) between 0 and 1800 nm**. These values are plotted along the lower edge of the diagram.
- Thickness is plotted along the y-axis
- Birefringence is plotted along the upper and right edge of the diagram.
- This color sequence is conveniently divided into orders, with the breaks between orders occurring **every 550 nm** of retardation.
- First- and second-order colors are most vivid; higher-order colors are progressively more and more washed out.
- Above the fourth order, the colors degenerate into a creamy white.

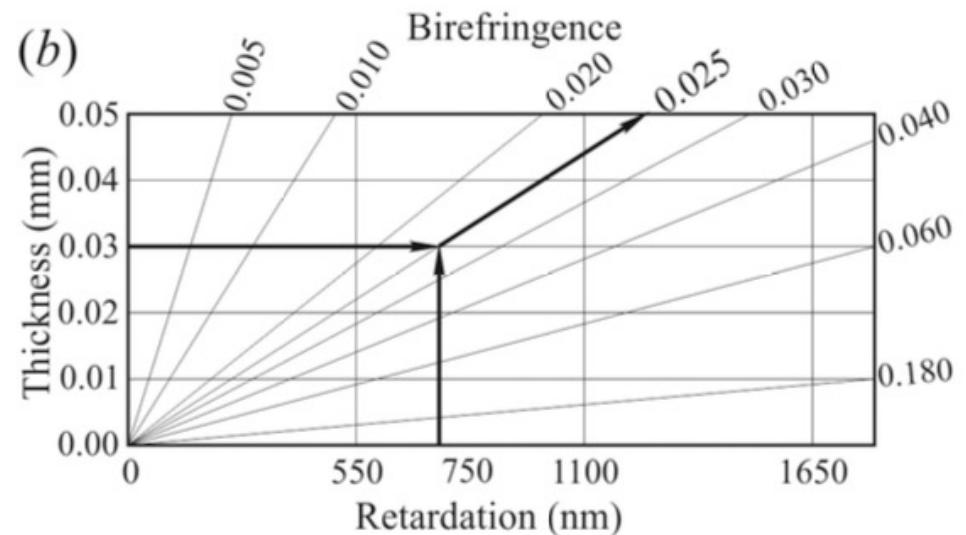
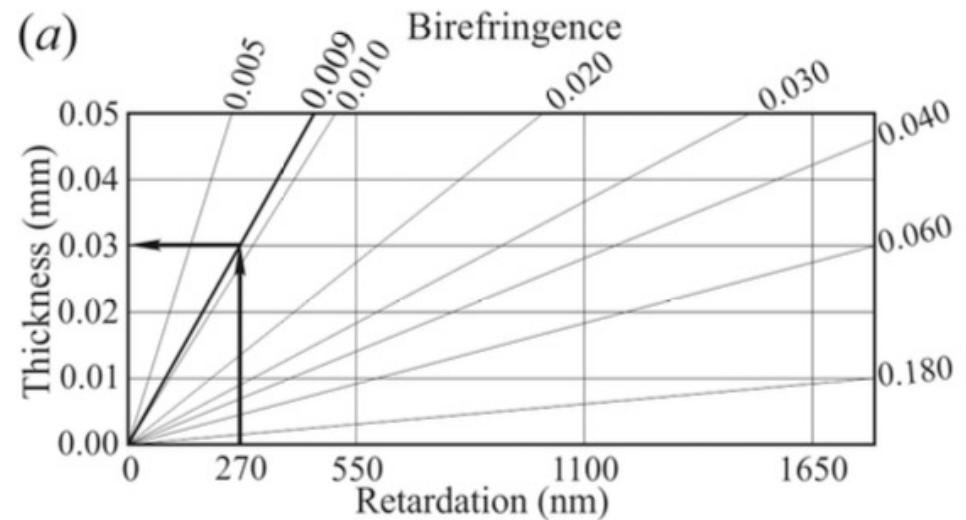


# Use of interference color chart

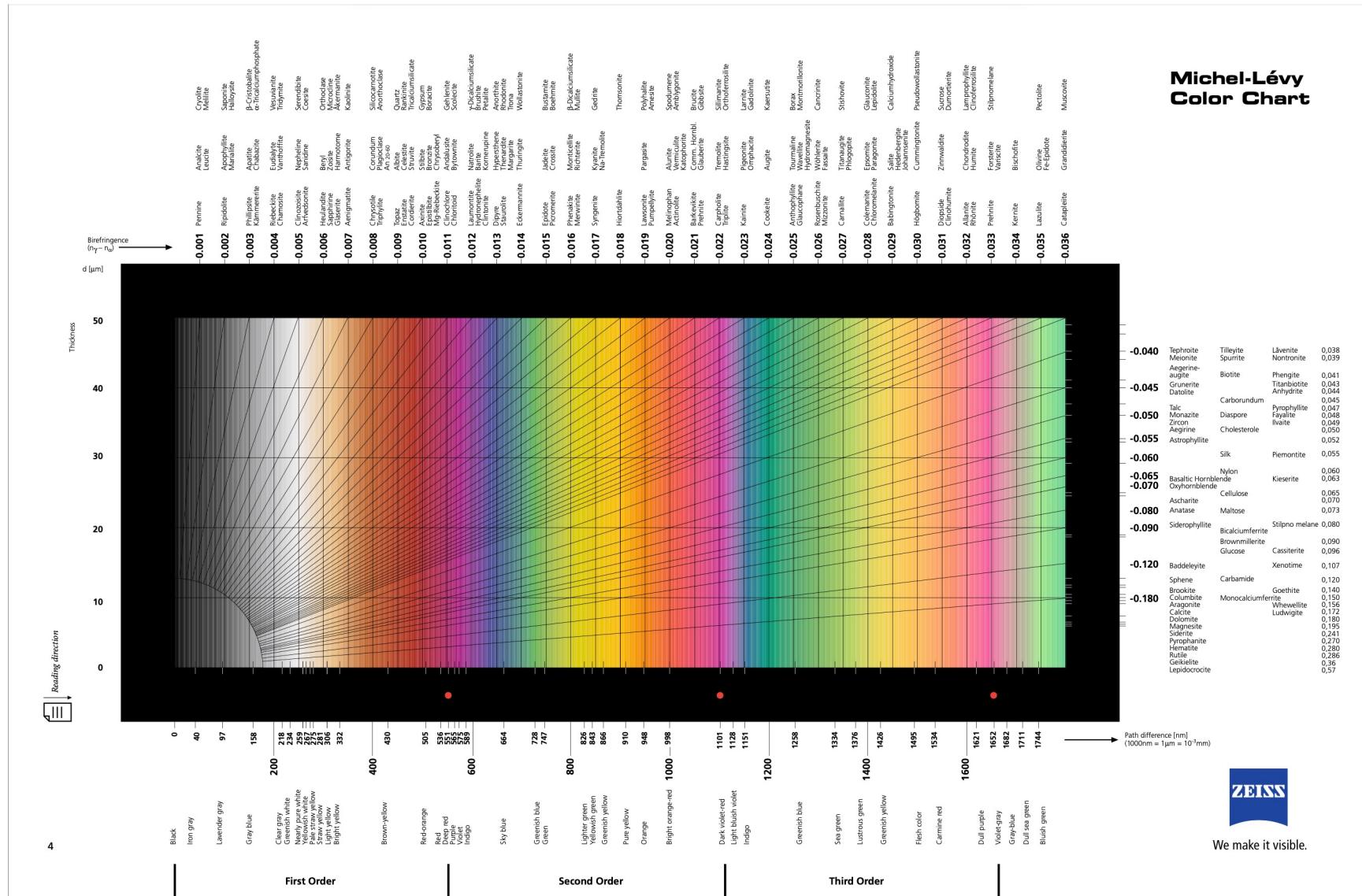
Used for:

- Determination of retardation between the slow and fast rays.
- Determination of thickness

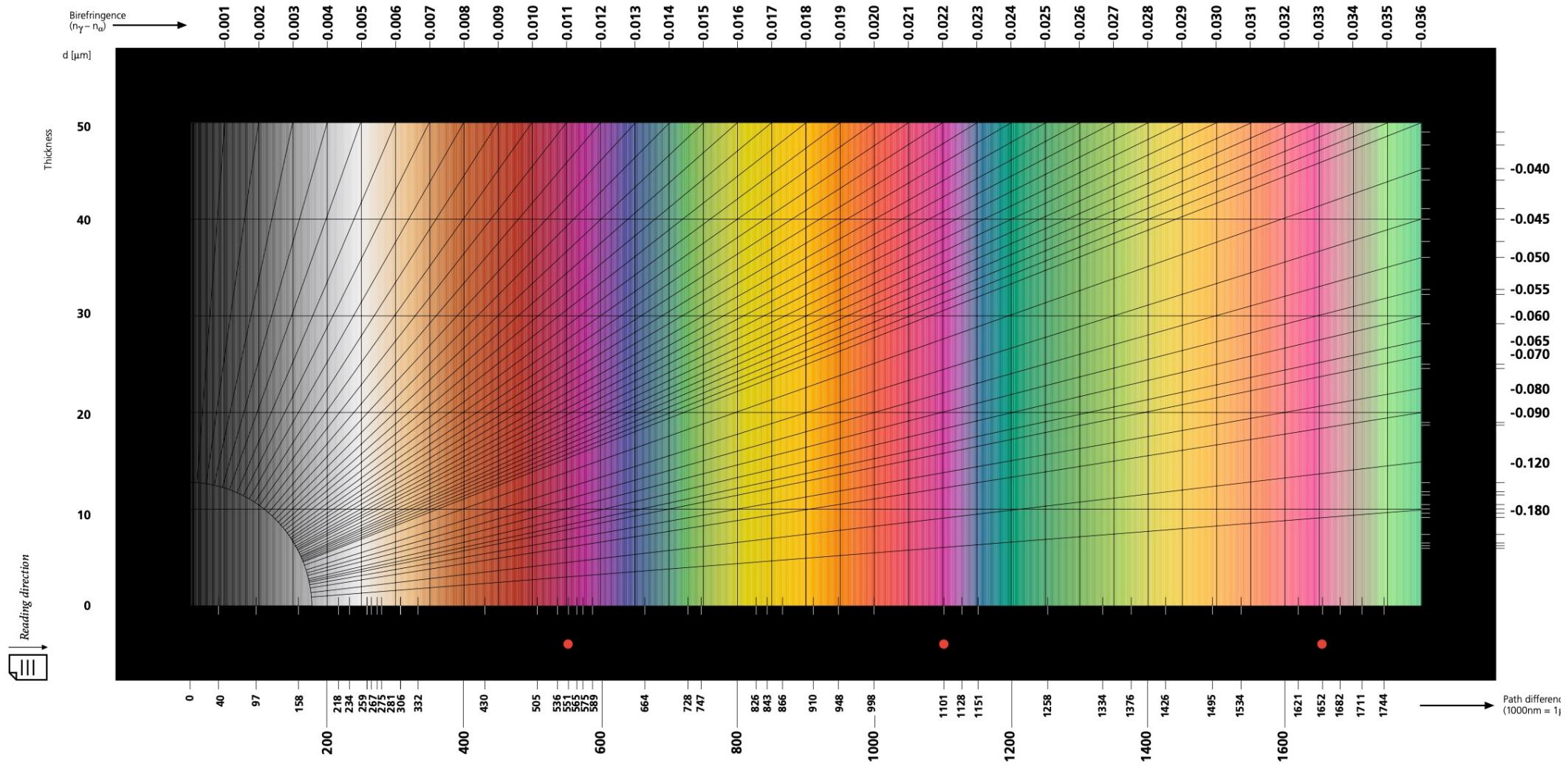
- (a) Thickness is indicated by the horizontal line that goes through the point where birefringence and interference color intersect.
- (b) Birefringence is indicated by the diagonal line that goes through the intersection of interference color (retardation) and thickness.



# Michel-Levy Color Chart



# Michel-Levy Color Chart



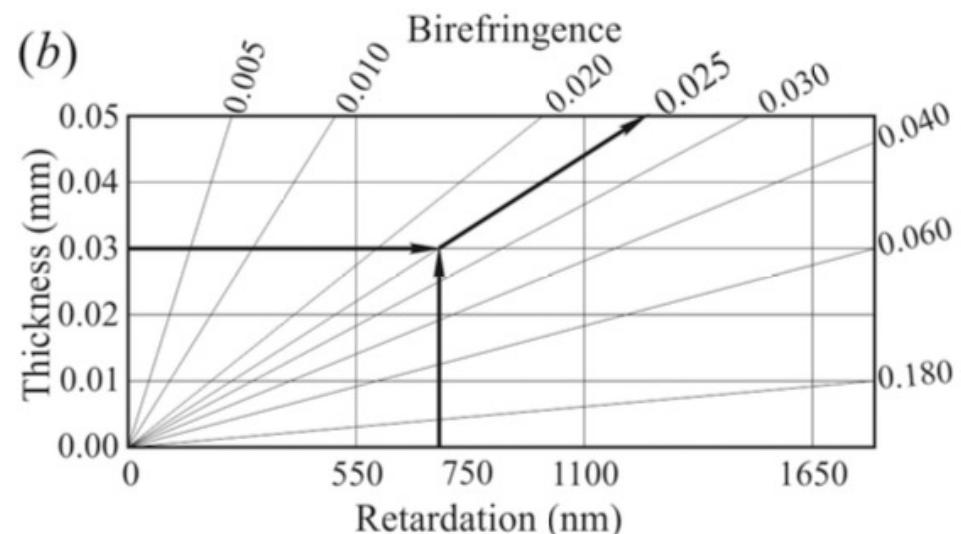
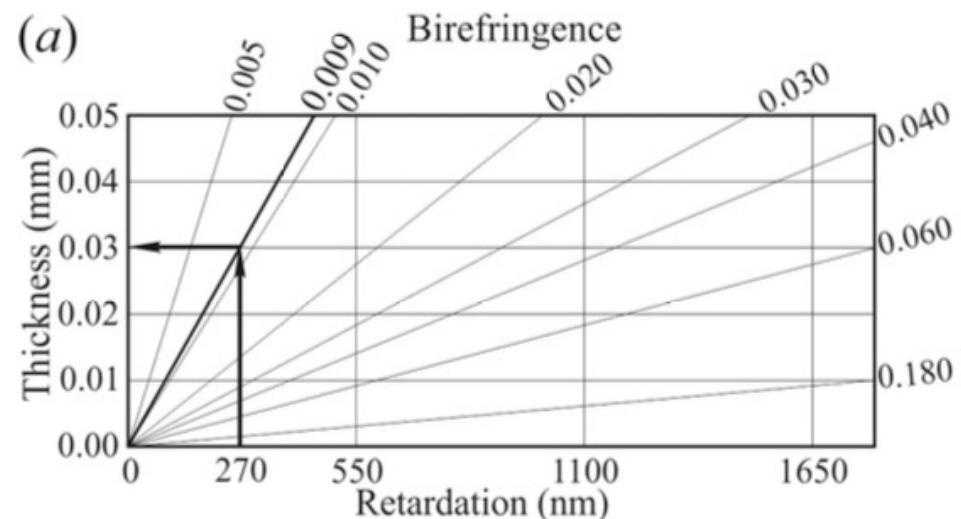
# Determining thickness

## Example: Quartz

- Scan the slide to find grains of quartz that display the highest interference color → farthest to the right on the interference color chart.
- A grain whose c axis is parallel to the microscope stage (horizontal) will display the maximum birefringence.
- Only these grains are oriented to display the known 0.009 (maximum) birefringence
- Identify the highest interference color displayed by the quartz and find it on the interference color chart.
- Read the retardation corresponding to that color from the bottom edge of the chart.
- If, for example, the color is first-order white with a tinge of yellow, the retardation is 270 nm.

$$\Delta = d(n_s - n_f) = d(\delta)$$

$$d = \frac{270 \text{ nm}}{0.009} = 30,000 \text{ nm} = 0.03 \text{ mm}$$

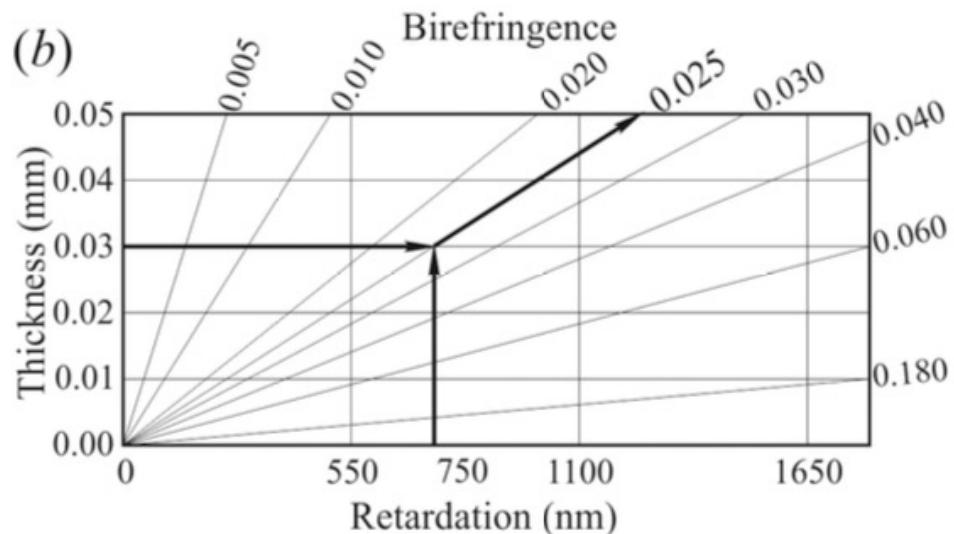
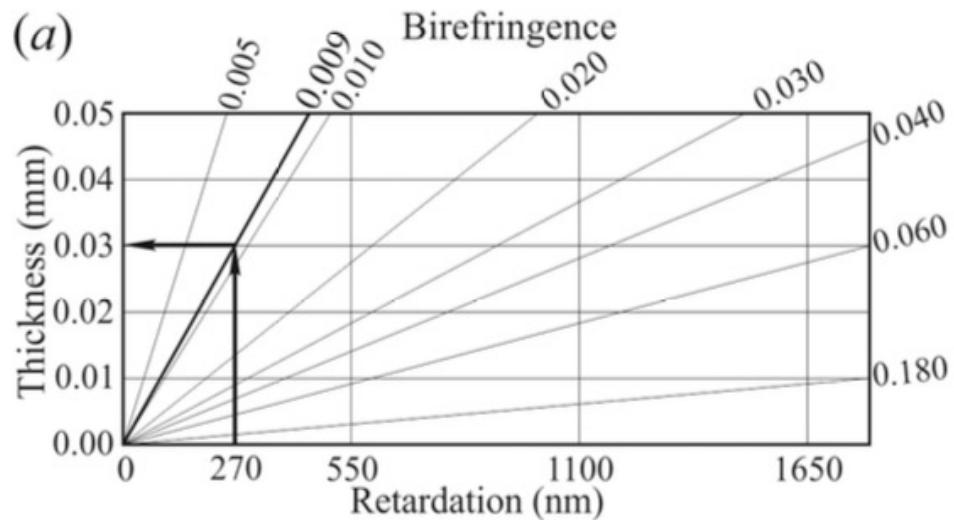


# Determining birefringence

- The maximum birefringence of a mineral is a very useful diagnostic property.
- It is easy to measure in thin sections whose thickness is known.

## For an unknown mineral:

- Scan the thin section to find a sample of the unknown mineral whose interference color is highest, i.e. farthest to the right on the interference color chart.
- Only these grains are oriented to display the maximum birefringence
- Identify that interference color, find it on the interference color chart.
- Read the retardation corresponding to that interference color from the bottom of the chart.



# Determining birefringence

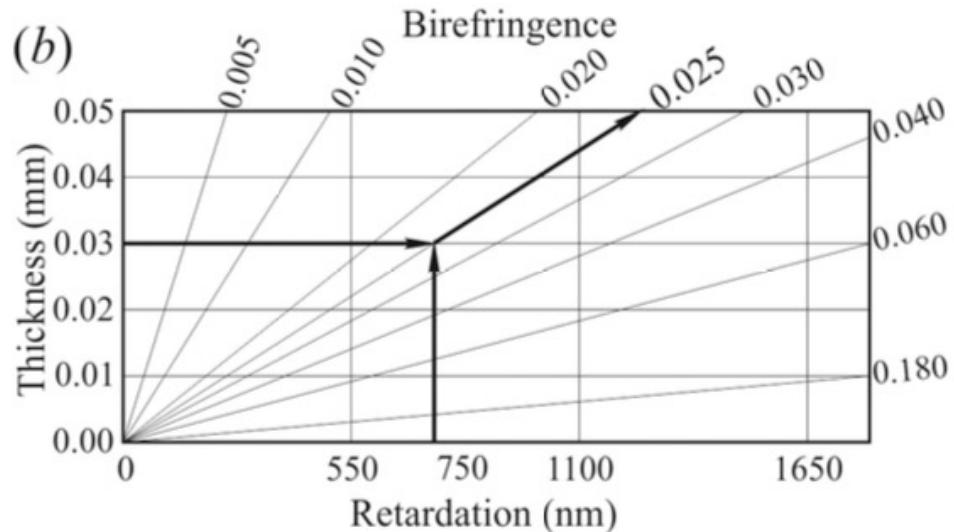
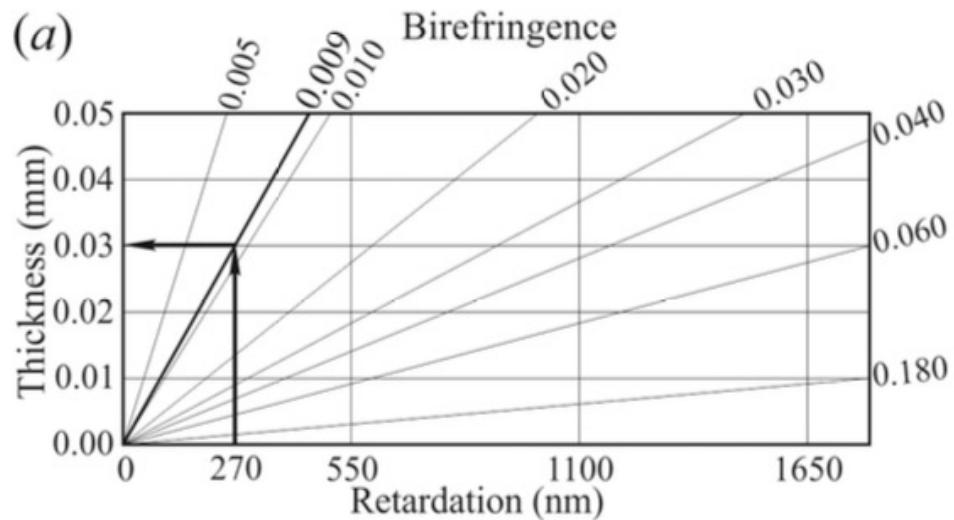
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**For an unknown mineral:**

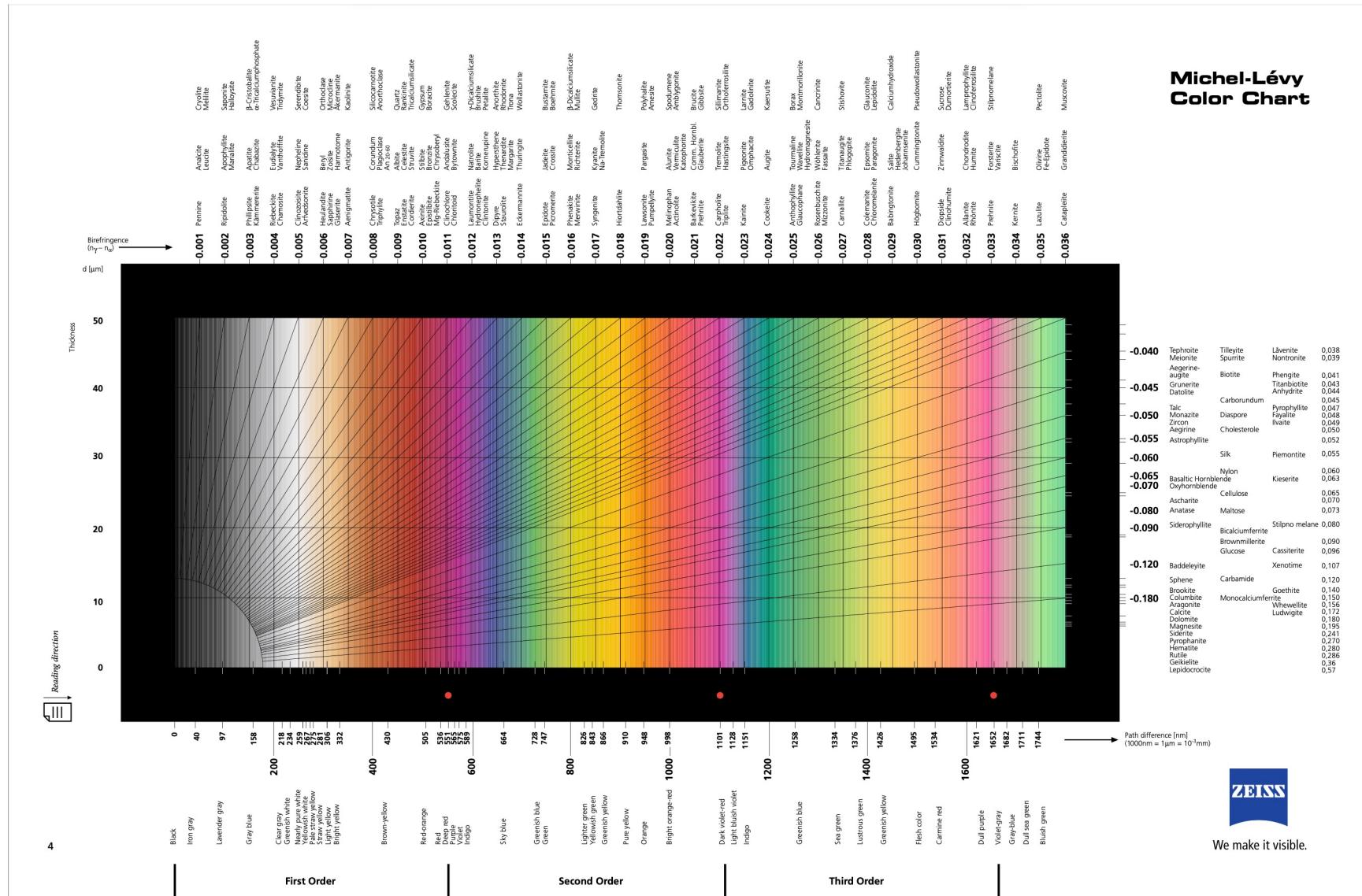
- For example, the interference color is second-order yellowish green, the retardation is about 750 nm.
- Calculate the birefringence ( $\delta$ ) by using the known thickness of the thin section (e.g., 0.03 mm =  $3 \times 10^4$  nm), and the retardation determined from interference color.

$$\delta = \frac{\Delta}{d} = \frac{750\text{nm}}{3 \times 10^4 \text{nm}} = 0.025$$

- Check the chart to look for possible minerals



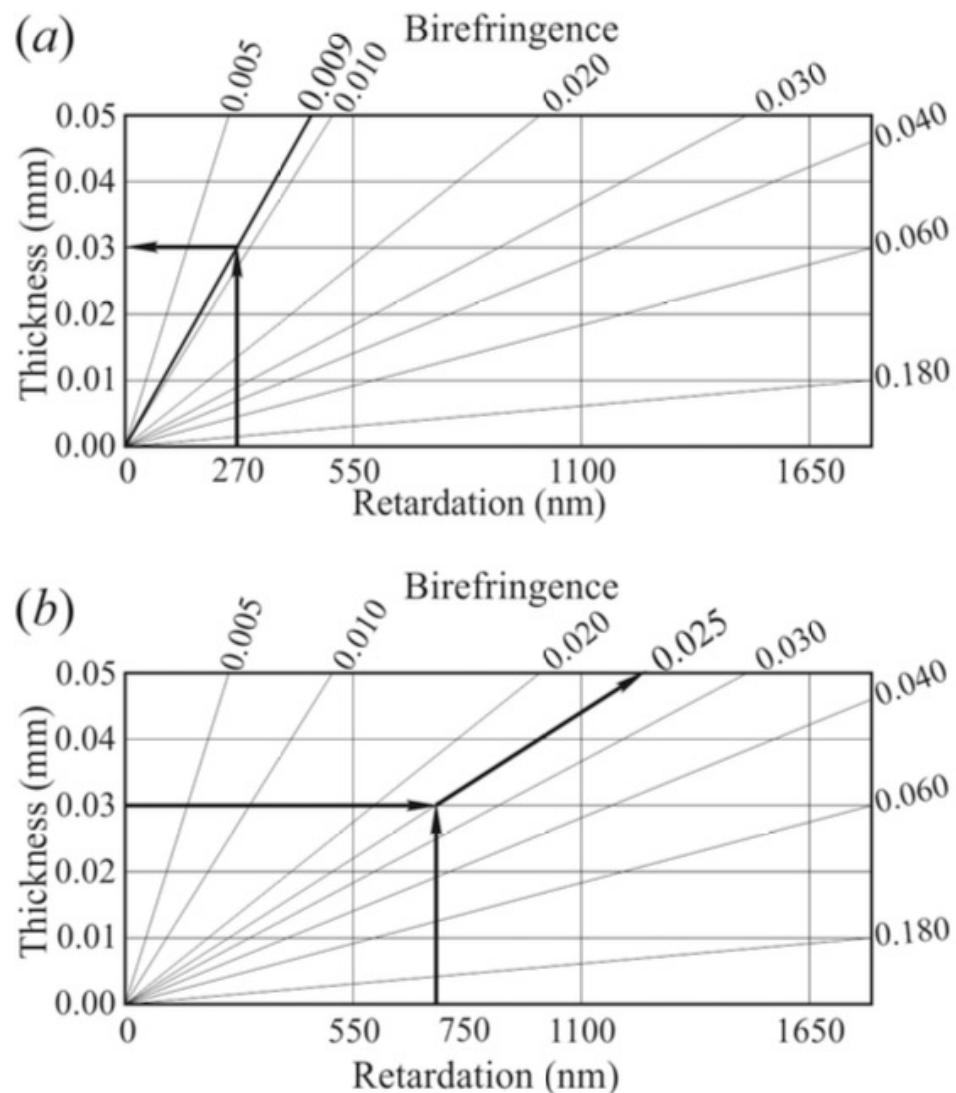
# Michel-Levy Color Chart



# Determining birefringence

## Sources of error

- Thickness of the thin section.
- The inability to accurately recognize interference colors of minerals
- Inaccurate color recognition or interpretation of the interference color chart.
- None of the grains in the sample may be cut along the right orientation so that the birefringence is the maximum value.
- ✓ Regardless of these limitations, the method provides sufficient accuracy for routine mineral identification work with the petrographic microscope.
- ✓ Recognition of unknown mineral should be aided by other optical properties as well. **E.g. habit of the grain, color in PPL, pleochroism, cleavages and extinction angles.**



Thank  
You



