

ARTICLE



1

https://doi.org/10.1038/s41467-022-28255-2

OPFN

Inner core composition paradox revealed by sound velocities of Fe and Fe-Si alloy

Haijun Huang ¹, Lili Fan¹, Xun Liu¹, Feng Xu ¹, Ye Wu ¹, Gang Yang ¹, Chunwei Leng ¹, Qingsong Wang ², Jidong Weng ², Xiang Wang ², Lingcang Cai ² & Yingwei Fei ^{3 ⋈}

Knowledge of the sound velocity of core materials is essential to explain the observed anomalously low shear wave velocity ($V_{\rm S}$) and high Poisson's ratio (σ) in the solid inner core. To date, neither $V_{\rm S}$ nor σ of Fe and Fe-Si alloy have been measured under core conditions. Here, we present $V_{\rm S}$ and σ derived from direct measurements of the compressional wave velocity, bulk sound velocity, and density of Fe and Fe-8.6 wt%Si up to ~230 GPa and ~5400 K. The new data show that neither the effect of temperature nor incorporation of Si would be sufficient to explain the observed low $V_{\rm S}$ and high σ of the inner core. A possible solution would add carbon (C) into the solid inner core that could further decrease $V_{\rm S}$ and increase σ . However, the physical property-based Fe-Si-C core models seemingly conflict with the partitioning behavior of Si and C between liquid and solid Fe.

¹ School of Science, Wuhan University of Technology, Wuhan, Hubei 430070, China. ² National Key Laboratory of Shock Wave and Detonation Physics, Institute of Fluid Physics, China Academy of Engineering Physics, Mianyang, Sichuan 621900, China. ³ Earth and Planets Laboratory, Carnegie Institution for Science, Washington, DC 20015, USA. [⊠]email: yfei@carnegiescience.edu

arth's core, containing a liquid outer core and a solid inner core, constitutes 32% of the mass of the planet, and its composition strongly influences the heat flow in Earth's deep interior, the crystallization of the inner core, and the evolution of the magnetic field¹. High-pressure experiments and geophysical observations have revealed that the iron-nickel outer core and inner core might contain ~10 wt% and ~4 wt% light elements, respectively²⁻⁴. The proposed light elements include carbon (C), hydrogen (H), oxygen (O), sulfur (S), and silicon (Si), but the identity and relative amounts of the light elements are still debated. Evidence from geochemistry⁵, metal-silicate partitioning experiments^{6,7}, and isotope fractionation experiments^{8,9} supports Si as the dominant light element. C might also be present in the Earth's core due to its abundance in primitive chondritic meteorites¹⁰, and its siderophile nature during metal-silicate differentiation¹¹⁻¹³. Any proposed composition models of the core must simultaneously satisfy the core density (ρ) , the bulk sound velocity $(V_{\rm B})$ of the liquid outer core, and the compressional wave velocity (V_P) and shear wave velocity (V_S) of the solid inner core, defined by seismic observations 14,15. However, it has been challenging to measure V_P at pressure (P) and temperature (T) relevant to the Earth's core conditions, and even more challenging for the measurements of V_S. Hence, measurements on core materials need to be extrapolated to core conditions to compare them with the seismic observations.

The extrapolation of $V_{\rm P}$ is usually made according to Birch's law², which assumes a linear relationship between ρ and $V_{\rm P}$. It is still debated whether Birch's law holds under high P-T conditions. For example, static data by inelastic X-ray scattering (IXS) showed that the $V_{\rm P}$ - ρ relation for iron¹6 follows Birch's law up to 93 GPa and 1100 K. On the other hand, measurements by IXS¹7,18 and nuclear-resonant inelastic X-ray scattering (NRIXS)¹9 showed that $V_{\rm P}$ decreases with temperature (<3000 K). The static data combined with the early shockwave data²0 indicate that Birch's law does not hold at moderate and high temperature. The disagreement on Birch's law would significantly affect the estimation of the amounts of Si in the inner core depending on models of the temperature effect on $V_{\rm P}$, ranging from 1–2 wt%²¹, and 3–6 wt%²², to 8 wt%¹¹.

For V_S of iron alloys, most of the data were derived from NRIXS measurements at high pressure and temperature^{23–26}. Estimate of the temperature effect at moderate and high temperature may be doubtful because the extraction of the phonen density of state is from the NRIXS spectra based on a quasi-harmonic model¹⁶. Therefore, no reliable V_S data of iron alloys are available under the Earth's core conditions. Comparison of the experimental V_S measurements with the seismic observations indicates that the extrapolated $V_{\rm S}$ of iron at 300 K is approximately two times of the inner core value, and the extrapolated Poisson's ratio (σ) of iron is approximately half of that observed in the inner core²³. High temperature and the presence of certain light elements such as C^{27,28} in the Earth's inner core are expected to decrease the V_S value and increase the σ value to match the observations. To model the V_S profile of the inner core, we need to determine the temperature and compositional dependence of V_S of iron alloys under high P-T conditions relevant to the Earth's inner core.

Computer simulations are capable of calculating ρ , V_P , and V_S under P-T conditions in the core, but the simulated results have not reached a consensus on the temperature dependence of Birch's law for V_P - ρ and V_S of iron. For example, the simulated results considering the effect of anharmonicity²⁹ revealed that both V_P and V_S of iron decreased with temperature at constant density and could match the seismic data of the inner core. On the other hand, more recent calculations^{30,31} showed that both V_P and V_S of iron increased almost linearly with density between

 $0{\sim}5500$ K, and proposed that the ρ , $V_{\rm P}$ and $V_{\rm S}$ of Fe₆₀Si₂C₂ (with 1.6 wt% Si and 0.7 wt% C) at 360 GPa and 6500 K could match the seismic data when considering the pre-melting effect³¹. However, all simulated $V_{\rm P}$ and $V_{\rm S}$ of iron at 0 K were higher than the extrapolation of experimental data at 300 K.

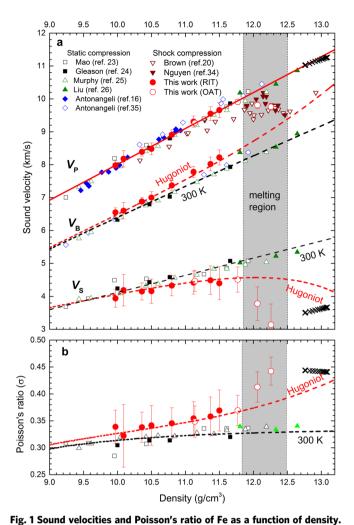
From the experimental measurements and theoretical calculations of $V_{\rm P}$ and $V_{\rm S}$ reported so far, there is no consensus on the temperature dependence of $V_{\rm P}$ and $V_{\rm S}$ of iron and iron alloys at high pressure and temperature (Fig. 1). Here, we report new direct measurements of $V_{\rm P}$ and $V_{\rm B}$ of Fe and Fe–Si alloy (8.6 wt% Si) by shock compression to investigate the temperature dependence on $V_{\rm P}$, $V_{\rm S}$ and σ under conditions relevant to the Earth's core. The results provide a test of Birch's law at simultaneous high pressure and temperature and determine if Si is a viable light element in the inner core.

Results

Sound velocity determination. Using the reverse-impact technique (RIT)^{32,33} and optical analyzer technique (OAT)^{20,33}, the sound velocities of Fe and Fe-8.6 wt% Si (hereafter Fe-8.6Si) were measured with a two-stage light gas gun. The experimental setup and the Lagrangian wave propagation diagram for the RIT are depicted in Supplementary Fig. 1. The data processing method is described in the Methods section. The RIT signals for Fe and Fe-8.6Si are shown in Supplementary Fig. 2. The RIT experiments produce both $V_{\rm P}$ and $V_{\rm B}$ measurements. To improve the accuracy, we take the derivative of the particle velocity with respect to time to pinpoint the arrival time of the rarefaction wave and the elastic-plastic transition point (Supplementary Fig. 3). Supplementary Table 1 lists the direct measurements of V_P and V_B . Then we obtained the shear wave velocity via $V_S = [3(V_P^2 (V_B^2)/4$] and the Poisson's ratio via $\sigma = 0.5(V_P^2 - 2V_S^2)/4$ $(V_{\rm P}^2 - V_{\rm S}^2)$.

Because of the low impedance of the lithium fluoride (LiF) window, the maximal pressure reached by the RIT is limited to ~160 GPa. For higher pressures, the $V_{\rm P}$ was obtained by the OAT, which utilizes multiple samples with different thicknesses to determine the catch-up thickness for accurate measurements of the sound velocity. The experimental setup and Lagrangian wave propagation diagram for the OAT are illustrated in Supplementary Fig. 4. Supplementary Fig. 5 shows representative signals and the determined catch-up thickness of experiment 180102 at 126 GPa. Similarly, we also take the derivative of the particle velocity with respect to time to accurately determine the arrival time of the rarefaction wave at the sample/window interface. Supplementary Table 2 lists the measured $V_{\rm P}$ along with the Hugoniot parameters. Using the thermodynamic parameters shown in Supplementary Table 3, we calculated $V_{\rm B}$, and then obtained $V_{\rm S}$ and σ from the calculated $V_{\rm B}$ and measured $V_{\rm P}$.

Hugoniot velocity and density measurements on iron. The compressional wave velocity $V_{\rm P}$ of iron along the Hugoniot measured by Brown and McQueen²⁰ is often used to assess the effect of temperature on $V_{\rm P}$ by comparing them with static data at room temperature. Figure 1a shows our experimental $V_{\rm P}$ for iron at simultaneously high pressures and temperatures in the range of 56.8~234.0 GPa and 1073~5417 K, respectively. The data show a linear velocity-density relationship, $V_{\rm P} = -2.89(\pm 0.10) + 1.09(\pm 0.02)$ ρ , up to 180 GPa (corresponding to 11.77 g/cm³) for solid iron. At the same density, our measured $V_{\rm P}$ values are 1.7~5.3% larger than previous data²⁰ by shock compression but are consistent with subsequent measurements by improved techniques³⁴. The typical uncertainties in velocity measurements are 1.7~3.1% (Supplementary Tables 1 and 2). The calculated temperatures of our Hugoniot data range from 1073 K at 56.8 GPa to 3919 K at 180.1 GPa. However, the



a The measured V_P along Hugoniot by the RIT (red solid circles) and the OAT (red open circles) are compared with previous shockwave data^{20,34}. The solid red line represents the linear fit to the V_P measurements for solid Fe. The static V_P data by NRIXS²³⁻²⁶ and IXS^{16,35} are also plotted for comparison. The measured V_B along Hugoniot by the RIT and the derived V_S from this study are shown by red circles, compared with NRIXS data with a natural ⁵⁷Fe isotope concentration at 300 K²³⁻²⁶. The black dashed line for V_B is calculated from the equation of state at 300 K, and the red dashed V_B -line is the calculated V_B from Hugoniot equation of state. The black and red dashed lines for V_S represent the calculated results from V_P and V_B , at 300 K and along Hugoniot, respectively. **b** The derived Poisson's ratios σ from this study are compared with the room-temperature NRIXS data. All measurements are compared with the observed values (black

measured V_P data along the Hugoniot are almost identical to the recent static data for hexagonal close-packed (hcp) iron at 300 K by IXS³⁵ and the corrected data (refer to the method²⁶) by NRIXS^{23–26} for iron with a natural ⁵⁷Fe isotope concentration. The measured V_P data are also consistent with the IXS data up to 1110 K obtained by stable external heating¹⁶. These results demonstrate that Birch's law still holds for the V_P of iron at high temperature.

crosses) from PREM14. The boundaries of the melting region (shaded area)

were determined according to the discontinuities of the sound velocity.

Our new Hugoniot density-pressure measurements are consistent with previous shock compression data³⁶ (Fig. 2). Using Hugoniot density-pressure data³⁶ and static 300 K isotherm⁴, we calculate V_P as a function of pressure constrained by Birch's law. Figure 2 shows consistent V_P results at room temperature and along Hugoniot between the measurements^{23–26} and calculations.

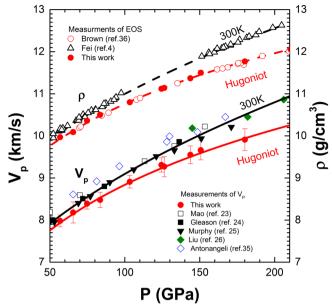


Fig. 2 The density and compressional wave velocity (V_P) of Fe as a function of pressure along the Hugoniot and 300-K isotherms. The Hugoniot densities of Fe³⁶ are compared with new Hugoniot data of this study and static data at 300 K⁴, along with the fitted results represented by the dashed lines. The red solid circles represent the measured V_P under shock compression (this work), compared with the static V_P data^{23–26,35}. The black and red curves represent the calculated V_P as a function of pressure along 300 K and shock temperature, using Birch's law and equations of state of Fe.

Due to the quasi-harmonic and anharmonic effect at high temperature, both ρ and $V_{\rm P}$ of iron under shock compression gradually deviate from the static room-temperature data at same pressure. But the linear relationship between $V_{\rm P}$ and ρ does not change, indicating the anharmonic effect does not play a significant role in Birch's law. The linear density-velocity relationship defined by Birch's law would allow a reliable extrapolation of $V_{\rm P}$ to core conditions for comparison with seismic models such as PREM¹⁴.

Figure 1a also shows the direct measurements of $V_{\rm B}$ from the RIT along with the calculated $V_{\rm B}$ from equation of state. The $V_{\rm B}$ obtained by both methods are in an excellent agreement. Static data²³⁻²⁶ showed that the $V_{\rm B}$ of pure iron increases linearly with density at 300 K. The shockwave data show a systematic deviation from the room-temperature data due to the temperature effect. From the measurements of V_P and V_B , we calculate V_S at the measured densities (Fig. 1a). At room temperature, V_S linearly increases with density, expressed as $V_S = -1.08(\pm 0.03) + 0.52(\pm 0.03)\rho$. The V_S along Hugoniot are systematically lower than the room-temperature values, showing shear softening with increasing temperature. The $V_{\rm S}$ values drop significantly at 12.06 and 12.26 g/cm³, indicating initiation of melting (Fig. 1a). Within the solid region, the effect of temperature on V_S is dependent on density and temperature (Fig. 3a). The rate of change of V_S in temperature, $|(\Delta V_S/\Delta T)_V|$, decreases with density by an exponential function and increases with temperature at a constant density. At 6000 K and 13.04 g/cm³, relevant to inner core condition, V_S of iron decreases at a rate of $(\Delta V_{\rm S}/\Delta T)_{\rm V} \approx 0.12~{\rm ms}^{-1}{\rm K}^{-1}$, much less than a calculated value of $0.48 \,\mathrm{ms}^{-1}\mathrm{K}^{-1}$ by simulations²⁹. The weak temperature dependence of V_S makes more difficult to match the observed low V_S in the inner core by the high-temperature effect alone.

Figure 1b shows the calculated Poisson's ratio σ that increases with density along the Hugoniot. The elevated σ values along the

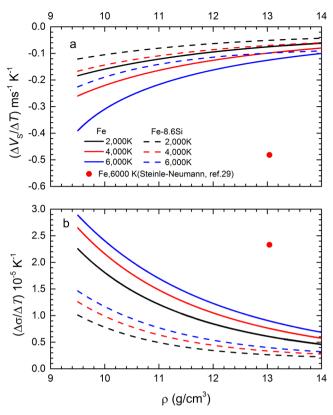


Fig. 3 The rate of change of $V_{\rm S}$ and σ in temperature as a function of density at different temperatures. a The calculated $\Delta V_{\rm S}/\Delta T$ increase with density along 2000K (black line), 4000 K (red line) and 6000 K (blue line) isotherms. b The calculated $\Delta \sigma/\Delta T$ decrease with density along 2000 K (black line), 4000 K (red line) and 6000 K (blue line) isotherms. The solid and dashed lines represent the values of Fe and Fe-8.6Si, respectively, compared with the ab initio calculations (red solid circle)²⁹.

Hugoniot relative to the 300 K isotherm indicate a positive effect by high temperature (Fig. 1b). Figure 3b shows the effect of temperature and density on σ . The $(\Delta\sigma/\Delta T)_V$ for iron can be expressed as an exponential function of density, with a positive temperature correlation. At 6000 K and 13.04 g/cm³, σ increases at a rate of $(\Delta\sigma/\Delta T)_V = 0.90 \times 10^{-5}$ K⁻¹, less than a predicted value of 2.33×10^{-5} K⁻¹ by simulations²⁹.

Our new Hugoniot velocity measurements on pure iron below melting provide a direct evaluation of the temperature effect on the shear wave velocity and the Poisson's ratio. We also observed melting at a pressure between 180 and 210 GPa, indicated by a large drop in $V_{\rm P}$ and $V_{\rm S}$ and a fast rise in σ . The onset of melting of iron under shock compression has been an intense debate. The data by Brown and McQueen²⁰ showed two discontinuities in sound velocity at 200 GPa and 243 GPa, respectively, whose explanation has been controversial. New in-situ X-ray diffraction measurements by static compression have ruled out possible solid phase transition below melting³⁷. Additional shock experiments by Nguyen and Holmes³⁴ showed a single discontinuity in $V_{\rm P}$ of iron, indicating onset of melting at ~225 GPa. Our $V_{\rm P}$ measurements indicate that experiments at 210 GPa and 234 GPa are in the melting region.

Hugoniot velocity measurements on iron-silicon alloy. In order to evaluate the effect of Si on the shear wave velocity, we also measured the sound velocities of Fe-8.6Si alloy by shock compression, using the same methodology described for the study of pure iron. Figure 4 shows the measurements of V_P and V_B for Fe-

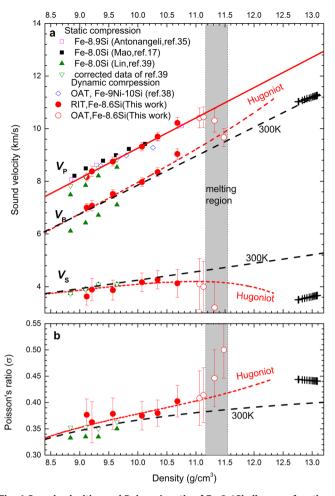


Fig. 4 Sound velocities and Poisson's ratio of Fe-8.6Si alloy as a function of density. a The measured V_P along Hugoniot by the RIT (red solid circles) and the OAT (red open circles) are compared with shockwave data for Fe-9Ni-10Si³⁸. The solid red line represents the linear fit to the V_P measurements for solid Fe-8.6Si alloy. The static V_P data of Fe-8Si by NRIXS³⁹ and of Fe-8.9Si³⁵, Fe-8Si¹⁷ by IXS are also plotted for comparison. The measures V_B along Hugoniot by the RIT and the derived V_S from this study are shown by red circles, compared with NRIXS data at 300 K³⁹. The olive open inverted triangles are the corrected values based on the equation of state for Fe-8.6 Si 40 and the measured $V_{\rm D}$ by NRIXS³⁹. The black dashed line for $V_{\rm B}$ is calculated from the equation of state at 300 K, and the red dashed V_B -line is the calculated V_B from Hugoniot equation of state. The black and red dashed lines for V_S represent the calculated results from V_P and $V_{\rm B}$, at 300 K and along Hugoniot, respectively. **b** The derived Poisson's ratios σ from this study are compared with the room-temperature NRIXS data. All measurements are compared with the observed values (black crosses) from PREM¹⁴. The boundaries of the melting region (shaded area) were determined according to the discontinuities of the sound velocity.

8.6Si up to 162 GPa (corresponding to 10.68 g/cm³) by the RIT (Supplementary Table 1). The $V_{\rm P}$ measurements were extended to higher pressures between 208 and 239 GPa, with the OAT. We observed the onset of melting at 239 GPa based on the velocity drop. The estimated melting temperature is slightly lower than that of Fe-9Ni-10Si³8, which could be caused by uncertainties in the calculated Hugoniot temperature and the estimated porosity (Supplementary Fig. 6).

Similar to the results of Fe, the measured V_P of the Fe–Si alloy also follows Birch's law, expressed by $V_P = -2.24(\pm 0.10) + 1.15(\pm 0.03)\rho$. The results are in agreements with shockwave data on solid Fe-9 wt% Ni-10 wt%Si alloy³⁸. They are also comparable to the static room-

temperature data for hcp Fe-9 wt% Si³⁵ and Fe-8 wt% Si¹⁷by IXS, showing no resolvable difference between the room-temperature and Hugoniot data. Early $V_{\rm P}$ measurements for Fe-8 wt% Si by NRIXS³⁹ showed systematically lower values compared to all other studies. The $V_{\rm P}$ and $V_{\rm S}$ by NRIXS were determined according to the calculated $V_{\rm B}$ inferred from the equation of state and measurements of the Debye sound velocity ($V_{\rm D}$) via $3V_{\rm D}^{-3}=V_{\rm P}^{-3}+2V_{\rm S}^{-3}$. Thus, $V_{\rm P}$ is sensitive to the difference in the calculated $V_{\rm B}$. At room temperature, the calculated $V_{\rm B}$ of Fe-8.0Si³⁹ is systematically smaller than our calculations based on the equation of state of Fe-8.6Si⁴⁰ (Fig. 4a). Combining the $V_{\rm D}$ measured by NRIXS³⁹ with the $V_{\rm B}$ of Fe-8.6Si obtained in this work, the corrected $V_{\rm P}$ of Fe-8.0 wt% Si is in excellent agreement with other static data by IXS and our Hugoniot data (Fig. 4a). The combined dataset shows no temperature effect on Birch's law.

Both datasets for pure Fe and Fe-8.6Si alloy demonstrated the validity of Birch's law. Using the linear V_P - ρ relationship defined by Birch's law, we extrapolated V_P of Fe-8.6Si to inner core conditions. The extrapolated value is consistent with the ab initio results for Fe-6.7Si between 0~4000 K³¹, and the calculated V_P for Fe-6.6 Si at 5500 K³⁰, but it is about 13.5% larger than the seismically observed value.

Our derived V_S values from Hugoniot V_P and V_B measurements are smaller than those at room temperature, derived from NRIXS, showing temperature-induced softening. It is noted that the $V_{\rm S}$ derived from NRIXS is insensitive to $V_{\rm B}$ data, thus, the corrected $V_{\rm S}$ values are indistinguishable from the original data of Fe-8.0 wt% Si³⁹. The room-temperature V_S is linearly related to the density, $V_S = 1.05(\pm 0.04) + 0.32(\pm 0.03)\rho$. The extrapolated data to inner-core density are much smaller than the simulated values at $0 \, \mathrm{K}^{30,31}$. Under shock compression, V_S gradually decreases (Fig. 4a) and σ gradually increases (Fig. 4b) because of the effect of temperature. Similar to the results for pure iron discussed above, the magnitudes of $|(\Delta V_S/\Delta T)_V|$ and $(\Delta \sigma/\Delta T)_V$ for Fe-8.6Si also decrease with density by an exponential function and increase with temperature (Fig. 3). At 6000 K and $\rho = 13.04 \text{ g/cm}^3$, V_S decreases at a rate of $(\Delta V_{\rm S}/\Delta T)_V = -0.10\,{\rm ms}^{-1}\,{\rm K}^{-1}$, and σ increases at a rate of $(\Delta \sigma / \Delta T)_V = 0.4 \times 10^{-5} \text{ K}^{-1}$. These rates are generally smaller than that of pure iron at the same conditions.

The above analysis of V_P , V_S and σ for Fe-8.6Si alloy is assumed no phase transition along the Hugoniot. Based on the calculated shock temperature of Fe-8.6Si⁴⁰ and the phase relations in Fe-Si alloys under static compression^{41,42}, it is possible that the shocked hcp Fe-8.6Si phase would decompose into a mixture of Si-poor hcp and Si-rich B2 (CsCl-type) phases at pressures above 145 GPa. However, there is no detectable change in the trends of V_P , V_S and σ above 145 GPa and 2400 K. This implies that either the decomposition process is kinetically prohibited because of the extremely short shock duration, or the physical properties of the decomposed mixture follow the ideal mixing rule⁴³. Neither of these scenarios would affect our analysis because the equation of state and the compressional wave velocity for Fe-Si system can be interpolated using an ideal mixing model.

Using the ideal linear mixing model⁴³, we interpolated the ρ and $V_{\rm P}$ for Fe-6 wt%Si, Fe-4.5 wt% Ni-3.7 wt%Si, and Fe-9 wt% Ni-2.3 wt%Si alloys based on the corresponding values of Fe-23 wt% Ni⁴⁴, Fe-8.6Si, and Fe. The measured densities of several Fe-Si alloys from static compression experiments are consistent with the calculated results based on our ideal mixing model (Supplementary Fig. 7). The interpolated $V_{\rm P}$ are in a good agreement with the measured $V_{\rm P}$ of Fe-9 wt%Ni-2.3 wt%Si²⁶ and Fe-6 wt%Si²², and broadly consistent with those of Fe-4.5 wt% Ni-3.7 wt%Si²¹ by IXS (Fig. 5). The measured $V_{\rm P}$ of Fe-5 wt%Si by picosecond acoustics measurements (PAM)⁴⁵ showed systematically higher values than the interpolation. Figure 5 illustrates that $V_{\rm P}$ of Fe-Si system increase with the Si content of the Fe-Si

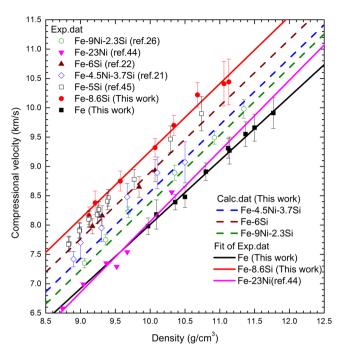


Fig. 5 The relationship between density and compressional velocity for the Fe-Ni-Si system. The black solid squares and red solid circles represent the new V_P measurements for Fe and Fe-8.6Si, respectively, compared with static data for various intermediate compositions including Fe-6Si²², Fe-4.5Ni-3.7Si²¹, Fe-9Ni-2.3Si²⁶, and Fe-5Si⁴⁵. The data for Fe-Ni alloy, Fe-23Ni⁴⁴, are also shown for comparison. The solid lines are the linear fits to the experimental data for the end members. The dashed lines represent the calculated V_P for Fe-6Si, Fe-4.5Ni-3.7Si and Fe-9Ni-2.3Si (in weight percent) according to an ideal linear mixing model⁴³.

alloys. It should be noted that the slope of the sound velocity of Fe-Ni is larger than that of pure iron which could influence the extrapolation when Ni is considered as a core component. However, the data range is limited, and data are more scattered. Additional data for Fe-Ni alloy are required to incorporate Ni into the core models.

Compared with the $V_{\rm P}$, the effect of Si on $V_{\rm S}$ of the Fe–Si alloys is more complicated because $V_{\rm S}$ is expected to decrease with temperature. Supplementary Fig. 8 shows that pure Fe has a stronger dependence of $V_{\rm S}$ on density than the Fe-8.6Si alloy, leading to a higher $V_{\rm S}$ for Fe than that of Fe-8.6Si at high density (>10.5 g/cm³). In contrast to $V_{\rm S}$, σ of Fe-8.6Si is almost parallel and ~16% larger than that of pure Fe at the same density (Supplementary Fig. 9). Therefore, the net effect is that an increase in the Si content in the alloy would decrease $V_{\rm S}$ and increase σ of the Fe–Si alloy under conditions of Earth's inner core.

Discussion

We established a thermodynamic model (see Methods) with optimized parameters (Supplementary Table 3) that best describes the experimental data, and calculated the ρ , $V_{\rm P}$, $V_{\rm S}$ and σ of Fe and Fe-8.6Si along an adiabatic temperature profile $T=T_{\rm ICB}~(\rho/\rho_{\rm ICB})^{\gamma}$, with a Grüneisen parameter γ of 1.5. The estimated temperature at the inner core boundary (ICB) $T_{\rm ICB}$ ranges from 5200 K to 5700 K, and the calculations were based on a preferred value of 5440 K³. Figure 6 shows the calculated ρ , $V_{\rm P}$, $V_{\rm S}$ and σ of Fe and Fe-8.6Si with uncertainties, compared with the PREM values in the inner core. The calculated values are listed in Supplementary Table 4. The uncertainties were assessed through propagation of errors associated with the parameters of equations

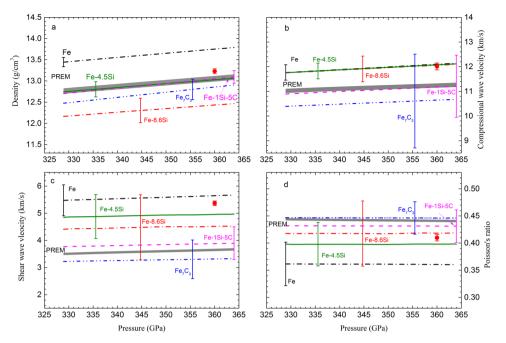


Fig. 6 Comparison of densities, compressional wave velocities, shear wave velocities and Poisson's ratios of various core compositions with PREM values. a The calculated densities of Fe (black dash-dot line), Fe-8.6Si (red dash-dot line), Fe-4.5 wt% Si (olive solid line), Fe₇C₃ (blue dash-dot-dot line) and Fe-1 wt% Si-5 wt% C mixture (magenta dash line) in the inner core assuming a temperature at inner core boundary of 5440 K, are compared to the PREM density profile. **b** The same comparison is shown for the compressional wave velocity. **c** The same comparison is shown for the shear wave velocity. **d** The same comparison is shown for the Poisson's ratio. The solid circles represent the simulated values for a model composition Fe₆₀Si₂C₂³¹ at 360 GPa and 5500 K. The thick gray lines represent the PREM¹⁴ for the inner core, with the line width representing 0.5% uncertainty in density, 0.5% in V_P and 1.1% in V_S . The PREM uncertainty were derived by comparing different Earth's model^{14,15}. Error bars are included for the calculated values. For the best-matched composition, Fe-1 wt% Si-5 wt% C, the error bars show 1.1 % errors in density, 11.4 % errors in V_S and 6.8% errors in σ .

of state, the Grüneisen parameters, and the parameters of Birch's law for $V_{\rm P}$. The uncertainties in ρ , $V_{\rm P}$, $V_{\rm S}$ and σ caused by the temperature uncertainty in the core were also assessed. A 10% uncertainty in $T_{\rm ICB}$ contributes ~0.5% uncertainty in ρ and $V_{\rm P}$, ~1% uncertainty in $V_{\rm S}$ and ~2% uncertainty in σ for iron.

Throughout the Earth's inner core, the density of Fe and Fe-8.6Si are ~5.3% higher and ~4.6% lower than the PREM values, respectively. A Si content of 4.5 wt.% could explain the density deficit in the inner core (Fig. 6a). For comparison, we calculated the sound velocity and σ for Fe-4.5Si alloy by interpolation. The V_P values of Fe, Fe-4.5Si and Fe-8.6Si are almost coincident with each other, ~6.9% higher than the PREM (Fig. 6b). It is not surprising that the V_P values of Fe–Si alloys are not sensitive to the Si content under inner core conditions because V_P of Fe-Si alloy linearly increases with the Si content at constant density (Supplementary Fig. 10), while the density of Fe-Si alloy linearly decreases with the Si content at the same rate (Supplementary Fig. 11). The analyses show that incorporation of Si alone in the inner core cannot simultaneously satisfy both the density and compressional sound velocity requirements defined by the observations.

Under the conditions of the inner core, the $V_{\rm S}$ of Fe–Si alloy decreases with the Si content. The $V_{\rm S}$ of Fe-4.5Si is ~11 % less than that of Fe, but it is still ~38% larger than the PREM value even with a temperature-induced softening (Fig. 6c). The σ of the Fe–Si system increases with the Si content under inner core's condition. Accounting for the temperature effect, the σ value of pure Fe at 5440 K increases ~9 % compared with that at room temperature. As a net effect, the σ value of Fe-4.5 Si is ~ 11% less than the PREM value (Fig. 6d).

The observed anomalously low $V_{\rm S}$ and high σ in the inner core were previously attributed to the temperature effect^{29–31}. In this work, we provide the first estimate of $V_{\rm S}$ and σ from direct

measurements of both $V_{\rm P}$ and $V_{\rm B}$ under simultaneous high temperature and pressure, relevant to the core conditions, and find that neither the temperature nor silicon is the main factor that would explain the observed low $V_{\rm S}$ and high σ in the inner core. Although we observed a considerable decrease in $V_{\rm S}$ and a considerable increase in σ during melting (Figs. 1 and 4), these drastic changes were caused by the effect in the solid-liquid mixing-phase region and not associated with pre-melting behavior. Here, we do not consider that the pre-melting behavior 30,31 is a viable mechanism to explain the low $V_{\rm S}$ and high σ in the inner core, because the observed low shear velocity is not limited to the top region of the inner core.

The mismatch in both $V_{\rm P}$ and $V_{\rm S}$ for Fe–Si alloys would rule out Si as the dominant light element in the solid inner core. Any additional light elements incorporated in the inner core would have to further decrease $V_{\rm P}$ and $V_{\rm S}$ to match the observed inner core values. Compared with the available sound velocity measurements of FeH_x⁴⁶, Fe₃S^{47,48}, Fe₇C₃²⁷, and Fe₃C⁴⁹, iron carbide Fe₇C₃ (8.4 wt% C) is the only iron alloy that reduces the $V_{\rm P}$ and $V_{\rm S}$ comparing with pure Fe. Phase relations in the Fe–C system support that a mixture of Fe₇C₃ and Fe could be a potential composition of the inner core if the C content is high enough^{50,51}. Therefore, we explore possible composition space for an Fe–Si–C core.

Measurements of the sound velocities (V_P and V_S) of Fe₇C₃ at room temperature²⁷ yielded $V_P = 2.16(\pm 1.16) + 0.66(\pm 0.12)\rho$ and $V_{S,300K} = 0.843(\pm 0.439) + 0.242(\pm 0.045)\rho$. The V_S values at high temperature were calculated by $V_S = V_{S,300K} + (dV_S/dT)_\rho$ (T -300). Because there is no measurement of the temperature-dependence of V_S for Fe₇C₃, we assumed that Fe₇C₃ has the same rate of change of V_S at different densities for a given temperature as pure Fe (Fig. 3). Our measurements indicate the upper bound of $(dV_S/dT)_V$ for pure Fe is about -0.124 ms $^{-1}K^{-1}$ at an inner-

core temperature of 6000 K. Using the equation-of-state parameters (Supplementary Table 3) and the velocity functions described above, we calculated ρ , V_P , V_S and σ for Fe₇C₃ under core conditions and then compared them to the observed values. Throughout the Earth's inner core, ρ , V_P and V_S for Fe₇C₃ are ~2.2%, ~5.7%, and ~7.7% less than the inner core observations, respectively, whereas the σ value of Fe₇C₃ is in the range of PREM. Thus, a combination of Si and C in the inner core could match PREM in all four key parameters (ρ , V_P , V_S , and σ). Using Fe, Fe-8.6Si, and Fe₇C₃ as the endmembers, the composition of Fe-1Si-5C (1 wt.% Si and 5 wt.% C) would yield the best fit to the observations (Fig. 6). The derived composition represents a global minimum using the best available thermodynamic parameters of the endmembers, and its uncertainty shown in Fig. 6 is propagated from errors associated with the parameters.

Our model calculations were based on interpolation from data of endmembers Fe, Fe-8.6Si and Fe₇C₃, using ideal mixing model⁴³. Among the measured physical properties, the density measurement has the highest precision. To examine whether the equations of state of the Fe-C-Si alloys follow the ideal mixing model, we calculated the density of Fe₉₃C₄Si₃ (corresponding to 0.9 wt.% C and 1.6 wt.% Si) using data from the endmembers, Fe⁴, Fe-8.6Si, Fe₇C₃⁵² and Fe₃C⁵³, and compared the calculated results with the density measurements of the same alloy composition⁵⁴. The agreement between the calculated and measured densities (Supplementary Fig. 12) suggests that the ideal mixing model is a reasonable interpolation for the analysis. The effect of C on the density in the core can be effectively modeled from the measurements of iron carbides if the C substitution is not interstitial in the structure, which may have a different effect⁵⁵.

The Fe–Si–C core model has also been tested using ab initio molecular dynamics calculations. Li et al. 31 calculated the density and velocities of hcp-FeSiC alloys for various composition combinations and found that the properties of a hcp-Fe $_{60}$ Si $_{2}$ C $_{2}$ (corresponding to 1.6 wt% Si and 0.7 wt% C) could match the density and sound velocity of the inner core, but at relatively high core temperature. The simulated $V_{\rm P}$, $V_{\rm S}$ and σ for Fe $_{60}$ Si $_{2}$ C $_{2}$ at 360 GPa and 5500 K would deviate from the PREM values (Fig. 6).

Adding C into the inner core is merely based on the required match of physical properties between core alloys and the inner core. The incorporation of C into the solid inner core would have consequence for the chemistry of the core. Based on cosmochemical and geochemical consideration, the abundance of the carbon in the Earth's core ranges from ~0.2 to ~1 wt%^{5,13}. Presence of ~5 wt.% C in the inner core would require that most of carbon entered the solid inner core during core solidification, with less than 0.8 wt.% C in the liquid outer core. Such a distribution of C between the inner and outer cores requires significant modification of melting relations in the Fe-C system at the ICB pressure, namely shifting the eutectic C composition toward the Fe endmember to keep the bulk C content of the core at the iron carbide + liquid region. This apparently contradicts with recent result on the evolution of the eutectic composition in the Fe-C system as a function of pressure, which did not detect significant change of the eutectic C content up to 260 GPa⁵¹. Similarly, a silicon-poor inner core requires to reconcile with seemingly robust conclusion of high Si content in the core from accretion and differentiation models^{6,7}. The solution to this inner core paradox between the mineral physics models and geochemical and petrological constraints might lie in our understanding of melting relations in a multi-component iron alloy system at the ICB conditions. Therefore, further experiments should investigate the partitioning of C and Si between the solid and liquid iron at the inner core boundary, as well as the sound velocity of Fe-C-Si alloy under conditions relevant to the Earth's inner core.

Methods

We used two-stage light gas guns at the Institute of Fluid Physics of China Academy of Engineering Physics and at the High Pressure Physics and Novel Materials Research Center in Wuhan University of Technology to measure the sound velocities of the Fe and Fe-8.6Si alloys. The discs of iron were from Trillion Metals Co., Ltd, with a chemical purity of $99.97(\pm0.05)$ wt% and an average density of $7.859(\pm0.007)$ g/cm³. The Fe–Si alloy powder with an average grain size of $\sim\!10\,\mu m$ was from Goodfellow Co., Ltd., and was sintered in a large-volume cubic multi-anvil apparatus at 5 GPa and $1200\,^{\circ} \text{C}$. The average bulk immersion density of the sintered sample is $7.386(\pm0.021)$ g/cm³. Electron microprobe analysis of the sintered discs showed uniform distribution of Si with an average Si content of $8.61(\pm0.02)$ wt%. The preparation of the Fe-8.6Si sample was described in our previous work 40 .

The sound velocities of pure Fe and Fe-8.6Si alloy were measured using two complementary methods, the reverse-impact technique (RIT)^{32,33} and optical analyzer technique (OAT)^{20,33}. For the RIT, a 12-mm sample disc was used as a flyer to directly impact a LiF single-crystal window coated with a 3-µm aluminum film. An 8-µm aluminum foil was mounted in front of the film with epoxy (Supplementary Fig. 1a). Supplementary Fig. 1b shows the Lagrangian distance-time diagram for the reverse-impact experiments. The particle velocity history u(t) at the film/LiF window interface was measured by the displacement interferometer system for any reflector (DISAR), as shown in Supplementary Fig. 2. When the flyer with thickness h impacted the window at time t_0 , indicated by a sharp increase in particle velocity, a shock wave with velocity $D_{\rm S}$ and an elastic precursor wave with velocity $D_{\rm C}$ were produced in the flyer. If $D_{\rm S}$ is less than $D_{\rm C}$, the elastic precursor wave $D_{\rm C}$ first reaches the rear surface of flyer, then is reflected as a rarefaction wave with velocity $V_{\rm P,0}$ (compressional velocity under ambient conditions), and finally meets the oncoming shock wave $D_{\rm S}$ at position $h_{\rm I}^{32}$.

$$h_1 = h \frac{D_s}{D_e} \left(\frac{D_e + V_{P,0}}{D_s + V_{P,0}} \right) \tag{1}$$

Then, $D_{\rm S}$ is reflected as a rarefaction wave transported with velocity $V_{\rm P,L}$ in the flyer and reaches the Al foil/LiF interface at time t_1 , indicating a decrease in particle velocity (Supplementary Fig. 3). The compressional velocity in the Lagrangian coordinates is determined by

$$V_{P,L} = \frac{h_1}{(t_1 - t_0) - h_1/D_s} \tag{2}$$

If the plastic shock velocity $D_{\rm S}$ is equal to or greater than the elastic precursor wave velocity at the Hugoniot elastic limit (HEL), only a single shock wave is observed, which is known as the "overdriven" condition. The overdriven conditions for Fe and Fe-8.6Si are expected to occur at stresses of $61(\pm 2)$ GPa and $49(\pm 2)$ GPa, respectively, based on their compressional velocities⁵⁶ under ambient condition. When the pressure is above the overdriven condition, position h_1 in Eqs. (1) and (2) is replaced with the flyer thickness h. Multiplying $V_{\rm P,L}$ with ρ_0/ρ (the ratio of the initial and compressed densities), we obtain the compressional wave velocity $V_{\rm P}$ in Eulerian coordinates. The main sources of the uncertainty in the measured $V_{\rm P}$ are from the uncertainties in the thickness h of the sample, the shock wave $D_{\rm S}$, the arrival time of the rarefaction wave t_1 , and the initial and compressed densities. The errors in the measured $V_{\rm P}$ for Fe and Fe-8.6Si range from 1.5~3% (Supplementary Table 1).

When the elastic-plastic transition (EPT) point at time t_2 during unloading is determined, the bulk velocities $V_{\rm B}$ at released pressure can also be obtained by replacing t_1 with t_2 in Eq. (2). However, the EPT point at t_2 is very subtle in the particle velocity history record and difficult to determine precisely. To obtain $V_{\rm B}$ precisely, we take the derivative of particle velocity u(t) with respect to time (du/ dt). The du/dt data have a minimum value due to EPT which is considered as the inflexion of particle velocity history u(t). Supplementary Fig. 3a, b shows the particle velocity and its derivative as a function of time for experiment 170314. Supplementary Fig. 3c shows the Lagrangian sound velocity as a function of the particle velocity during unloading. During plastic unloading, the Lagrangian sound velocity increases linearly with the particle velocity. Extrapolating the linear relation to the particle velocity just before unloading yields a Lagrangian bulk velocity $V_{\rm B,L}$ under compression. Multiplying $V_{\rm B,L}$ by $\rho_{\rm 0}/\rho$, we obtain the bulk velocity $V_{\rm B}$ in Eulerian coordinates. The main sources of the uncertainty in V_B are the uncertainties in h, D_S, t₂ and the fitted parameters for the Lagrangian sound velocity during plastic unloading. The errors in the measured V_B range from 1.9~3.8% (Supplementary Table 1).

From the direct measurements of $V_{\rm P}$ and $V_{\rm B}$, we obtained the shear wave velocity via $V_{\rm S}=[3(V_{\rm P}^2-V_{\rm B}^2)/4]^{1/2}$ and Poisson's ratio via $\sigma=0.5(V_{\rm P}^2-2V_{\rm S}^2)/(V_{\rm P}^2-V_{\rm S}^2)$, listed in Supplementary Table 1. Compared with $V_{\rm P}$ and $V_{\rm B}$, the propagated errors in $V_{\rm S}$ and σ are considerably large, ranging from 6~13% and 8~18%, respectively. Nevertheless, the derived $V_{\rm S}$ and σ along Hugoniot provide a direct evaluation of the temperature effect on shear wave velocity and Poisson's ratio when compared with static measurements at room temperature.

We also performed $V_{\rm P}$ measurements on pure Fe and Fe-8.6Si alloy at higher pressure using the optical analyzer technique (OAT). Supplementary Fig. 4 illustrates the experimental configuration of a flyer (Fe or Ta) impacting the sample with different thicknesses and the Lagrangian distance-time diagram showing the flyer and sample interaction. Supplementary Fig. 5a-c show three particle velocity history records at the interface between the Ta foil and LiF window for the

experiment 180102 with three different sample thicknesses. A linear fit of the measured time interval as a function of the sample thickness provides the catch-up thickness (Supplementary Fig. 5d). The details of the sound velocity determination from the measured catch-up thickness of the sample and the known flyer properties were discussed by Huang et al 33 . The main sources of the uncertainty are from the uncertainties in the thickness of the sample, the shock velocities in the flyer and sample, and the compressional wave velocity in the flyer if its material is different from that of the sample. The uncertainty in $V_{\rm P}$ measured by OAT ranges from $1{\sim}4\%$ (Supplementary Table 2).

We also calculated the bulk sound velocities $V_{\rm B}$ from the Hugoniot equation of state for solid Fe and Fe-8.6Si alloy based on following thermodynamic model.

$$V_{\rm B}^2 = -V^2 \left(\frac{dP}{dV}\right)_{VT} + \gamma_{\rm eff}^2 C_V T \tag{3}$$

The value $(dP/dV)_{V,T}$ was obtained from the Grüneisen equation of state.

$$P(V,T) = P_{T_0}(V) + \int_{T_0}^{T} \frac{\gamma_{\text{eff}}}{V} C_V dT$$
 (4)

where y_{eff} the effective Grüneisen parameter, includes the lattice and electrons contributions.

$$\gamma_{\text{eff}} = \frac{C_{V,l} \gamma_{l} + C_{V,e} \gamma_{e}}{C_{V,l} + C_{V,e}}$$
 (5)

The lattice Grüneisen parameter is defined as $y_1 = y_0 (\rho_0/\rho)^q$. y_e is the electrons Grüneisen parameter. $C_{V,1} = 3 \text{ R}/\mu$ and $C_{V,e} = \beta_0 (\rho_0/\rho)^k T$ are the specific heat contributed by the lattice and electrons, respectively. The parameters for Fe, Fe-8.6Si and Fe₇C₃ were listed in Supplementary Table 3. $P_{T_0}(V)$ is the equation of state at 300 K, which was calculated from Hugoniot pressure $P_H(V,T_H)$ by

$$P_{T_0}(V) = P_{H}(V, T_{H}) - \int_{T_0}^{T_{H}} \frac{\gamma_{\text{eff}}}{V} C_V dT$$
 (6)

The shock temperature $T_{\rm H}$ were calculated from the thermodynamic relation²⁰.

$$dT = \left[-T \frac{\gamma_{\text{eff}}}{V} + \frac{P_{\text{H}} + (V_0 - V)}{2C_V} \frac{dP_{\text{H}}}{dV} \right] dV \tag{7}$$

Using the third-order Birch–Murnaghan equation of state, the fitted parameters of isothermal bulk modulus K_0 and its pressure derivative K_0 ' are given in Supplementary Table 3.

The calculated bulk sound velocity from the Hugoniot equation of state agrees well with the experimental measurements (Figs. 1 and 4). The main sources of uncertainty in the model calculations of $V_{\rm B}$ are uncertainties in the Hugoniot parameters (C_0,λ) , the lattice contributions to the Grüneisen parameter, and the electronic and anharmonic contributions to the specific heat. The thermodynamic parameters and their errors for Fe and Fe-8.6Si are listed in Supplementary Table 3. The uncertainties in the Hugoniot parameters directly contribute to the uncertainties in the derived equation of state, having a large influence on the propagated errors in $V_{\rm B}$. Along Hugoniot, the propagated errors in the calculated $V_{\rm B}$ for Fe and Fe-8.6Si are approximately 2% and 5%, respectively, which are smaller than the errors in the calculated $V_{\rm S}$ in the ranges of 7~20% and 23~30%, respectively. The errors in σ for Fe and Fe-8.6Si are 6~8% and 12%, respectively.

Under the conditions of Earth's inner core, the density and sound velocity in the Fe-Si-C system are estimated based on an ideal mixing model

$$\rho = 1 / \sum_{i} \frac{W_i}{\rho_i} \tag{8}$$

and

$$V_{\text{P(or S)}} = 1 / \left(\rho \sum_{i} \frac{W_{i}}{\rho_{i} V_{\text{P(or S)}, i}} \right)$$
 (9)

where W_i is the weight percent, $V_{P(\text{or S}),i}$ and ρ_i represent the sound velocity and density of the endmember (Fe, Fe-8.6Si or Fe₇C₃) at the same pressure and temperature. For Fe and Fe-8.6Si, ρ and V_B were calculated with Eqs. (3) and (4), and V_P data were calculated using Birch's law. The V_S and σ were then calculated from ρ , V_B and V_P . Because of no direct shock wave measurement for Fe₇C₃, the thermodynamic parameters for Fe₇C₃ at high pressure and temperature have not been constrained. We therefore estimated the V_S of Fe₇C₃ using the data at 300 K with an assumed $(dV_S/dT)_V$ value. The errors of the density and sound velocity of the mixture were calculated based on the following equations.

$$\delta \rho = \left[\sum \left(\frac{\partial \rho}{\partial \rho_i} \delta \rho_i \right)^2 \right]^{1/2} \tag{10}$$

and

$$\delta V_{P(\text{or S})} = \left[\sum \left(\frac{\partial V_{P(\text{or S})}}{\partial \rho_i} \delta \rho_i \right)^2 + \sum \left(\frac{\partial V_{P(\text{or S})}}{\partial V_{P(\text{or S}),i}} \delta V_{P(\text{or S}),i} \right)^2 \right]^{1/2}$$
(11)

where the subscript P(or S) indicates the compressional or shear velocity. The subscript i represents the ith endmember (Fe, Fe-8.6Si or Fe₇C₃) at the same

pressure and temperature. $\delta \rho_i$ and $\delta V_{P(or~S),i}$ represent the errors of density and sound velocity of the *i*th endmember.

Data availability

All data generated or analyzed during this study are included in this published article and its Supplementary Information files.

Received: 3 May 2021; Accepted: 10 January 2022; Published online: 01 February 2022

References

- Buffett, B. A., Huppert, H. E., Lister, J. R. & Woods, A. W. On the thermal evolution of the Earth's core. J. Geophys. Res. 101, 7989–8006 (1996).
- Birch, F. Composition of the Earth's mantle. Geophys. J. R. Astron. Soc. 4, 295–311 (1961).
- Hirose, K., Labrosse, S. & Hernlund, J. Composition and state of the core. *Annu. Rev. Earth Planet. Sci.* 41, 657–691 (2013).
- Fei, Y., Murphy, C., Shibazaki, Y., Shahar, A. & Huang, H. Thermal equation of state of hcp-iron: constraint on the density deficit of earth's solid inner core. *Geophys. Res. Lett.* 43, 6837–6843 (2016).
- McDonough, W. F. Compositional model for the Earth's core. In *The mantle and core.Treatise on Geochemistry*, Vol. 2, (ed. Carlson, R. W.) chap. 15 (Elsevier, 2003), pp. 547–568.
- Rubie, D. C. et al. Heterogeneous accretion, composition and core–mantle differentiation of the earth. Earth Planet. Sci. Lett. 301, 31–42 (2011).
- Fischer, R. A. et al. High pressure metal-silicate partitioning of Ni, Co, V, Cr, Si, and O. Geochim. et Cosmochim. Acta 167, 177–194 (2015).
- Fitoussi, C., Bourdon, B., Kleine, T., Oberli, F. & Reynolds, B. C. Si isotope systematics of meteorites and terrestrial peridotites: implications for Mg/Si fractionation in the solar nebula and for Si in the Earth's core. *Earth Planet.* Sci. Lett. 287, 77–85 (2009).
- Shahar, A. et al. Experimentally determined Si isotope fractionation between silicate and Fe metal and implications for Earth's core formation. *Earth Planet.* Sci. Lett. 288, 228–234 (2009).
- Wasson, J. T. & Kallemeyn, G. W. Compositions of chondrites. *Philos. Trans. R. Soc. Lond. Ser. A Math. Phys. Sci.* 325, 535–544 (1988).
- Dasgupta, R., Chi, H., Shimizu, N., Buono, A. S. & Walker, D. Carbon solution and partitioning between metallic and silicate melts in a shallow magma ocean: Implications for the origin and distribution of terrestrial carbon. *Geochimica et. Cosmochimica Acta* 102, 191–212 (2013).
- Fischer, R. A., Cottrell, E., Hauri, E., Lee, K. K. & Le Voyer, M. The carbon content of Earth and its core. *Proc. Natl Acad. Sci. USA.* 117, 8743–8749 (2020).
- Wood, B. J., Li, J. & Shahar, A. Carbon in the core: its influence on the properties of core and mantle. Rev. Mineral. Geochem. 75, 231–250 (2013).
- Dziewonski, A. M. & Anderson, D. L. Preliminary reference earth model. *Phys. Earth Planet. Inter.* 25, 297–356 (1981).
- Tkalčić, H. & Pham, T. S. Shear properties of Earth's inner core constrained by a detection of J waves in global correlation wavefield. Science 362, 329–332 (2018).
- Antonangeli, D. et al. Simultaneous sound velocity and density measurements of hcp iron up to 93 GPa and 1100 K: An experimental test of the Birch's law at high temperature. Earth Planet. Sci. Lett. 331, 210–214 (2012).
- Mao, Z. et al. Sound velocities of Fe and Fe-Si alloy in the earth's core. Proc. Natl Acad. Sci. USA. 109, 10239–10244 (2012).
- Sakamaki, T. et al. Constraints on Earth's inner core composition inferred from measurements of the sound velocity of hcp-iron in extreme conditions. Sci. Adv. 2, e1500802 (2016).
- Lin, J. F. et al. Sound velocities of hot dense iron: Birch's law revisited. Science 308, 1892–1894 (2005).
- Brown, J. M. & McQueen, R. G. Phase transitions, Grüneisen parameter, and elasticity for shocked iron between 77 GPa and 400 GPa. J. Geophys. Res. 91, 7485–7494 (1986).
- Antonangeli, D. et al. Composition of the Earth's inner core from highpressure sound velocity measurements in Fe-Ni-Si alloys. *Earth Planet. Sci. Lett.* 295, 292–296 (2010).
- Sakairi, T. et al. Sound velocity measurements of hcp Fe-Si alloy at high pressure and high temperature by inelastic x-ray scattering. Am. Mineral. 103, 85–90 (2018).
- Mao, H. K. et al. Phonon density of states of iron up to 153 gigapascals. Science 292, 914–916 (2001).
- Gleason, A. E., Mao, W. L. & Zhao, J. Y. Sound velocities for hexagonally close packed iron compressed hydrostatically to 136 GPa from phonon density of states. *Geophys Res. Lett.* 40, 2983–2987 (2013).

- Murphy, C. A., Jackson, J. M. & Sturhahn, W. Experimental constraints on the thermodynamics and sound velocities of hcp-Fe to core pressures. *J. Geophys. Res.* 118, 1999–2016 (2013).
- 26. Liu, J. et al. Seismic parameters of hcp-Fe alloyed with Ni and Si in the Earth's inner core. *J. Geophys. Res.* **121**, 610–623 (2016).
- Chen, B. et al. Hidden carbon in Earth's inner core revealed by shear softening in dense Fe₂C₃. Proc. Natl Acad. Sci. USA 111, 17755–17758 (2014).
- 28. Prescher, C. et al. High Poisson's ratio of Earth's inner core explained by carbon alloying. *Nat. Geosci.* **8**, 220–223 (2015).
- Steinle-Neumann, G., Stixrude, L., Cohen, R. E. & Gülseren, O. Elasticity of iron at the temperature of the Earth's inner core. *Nature* 413, 57–60 (2001).
- Martorell, B., Wood, I. G., Brodholt, J. & Vočadlo, L. The elastic properties of hcp-Fe_{1-x}Si_x at Earth's inner-core conditions. *Earth Planet. Sci. Lett.* 451, 89–96 (2016).
- Li, Y., Vocadlo, L. & Brodholt, J. P. The elastic properties of hcp -Fe alloys under the conditions of the Earth's inner core. *Earth Planet. Sc. Lett.* 493, 118–127 (2018).
- Duffy, T. S. & Ahrens, T. J. Compressional sound velocity, equation of state, and constitutive response of shock-compressed magnesium oxide. *J. Geophys. Res.* 100, 529–542 (1995).
- Huang, H. et al. Evidence for an oxygen-depleted liquid outer core of the Earth. Nature 479, 513–516 (2011).
- Nguyen, J. H. & Holmes, N. C. Melting of iron at the physical conditions of the Earth's core. Nature 427, 339–342 (2004).
- Antonangeli, D. et al. Sound velocities and density measurements of solid hcp-Fe and hcp-Fe-Si (9 wt.%) alloy at high pressure: Constraints on the Si abundance in the Earth's inner core. Earth Planet. Sci. Lett. 482, 446–453 (2018).
- Brown, J. M., Fritz, J. N. & Hixson, R. S. Hugoniot data for iron. *J. Appl. Phys.* 88, 5496–5498 (2000).
- Anzellini, S., Dewaele, A., Mezouar, M., Loubeyre, P. & Morard, G. Melting of iron at Earth's inner core boundary based on fast X-ray diffraction. *Science* 340, 464–466 (2013).
- 38. Zhang, Y. et al. Experimental constraints on light elements in the earth's outer core. *Sci. Rep.* **6**, 22473 (2016).
- Lin, J. et al. Sound velocities of iron-nickel and iron-silicon alloys at high pressures. Geophys. Res. Lett. 30, 2112–2115 (2003).
- Huang, H. et al. Equation of state for shocked Fe-8.6wt% Si up to 240 GPa and 4,670 K. J. Geophys. Res. 124, 8300–8312 (2019).
- Fischer, R. A. et al. Phase relations in the Fe-FeSi system at high pressures and temperatures. Earth Planet. Sci. Lett. 373, 54-64 (2013).
- 42. Tateno, S., Kuwayama, Y., Hirose, K. & Ohishi, Y. The structure of Fe-Si alloy in Earth's inner core. *Earth Planet. Sci. Lett.* **418**, 11–19 (2015).
- Badro, J. et al. Effect of light elements on the sound velocities in solid iron: implications for the composition of earth's core. Earth Planet. Sci. Lett. 254, 233–238 (2007).
- Kantor, A. P. et al. Sound wave velocities of fcc Fe-Ni alloy at high pressure and temperature by mean of inelastic X-ray scattering. *Phys. Earth Planet. Inter.* 164, 83–89 (2007).
- Edmund, E. et al. Velocity-density systematics of Fe-5wt% Si: constraints on Si content in the Earth's inner core. J. Geophys. Res. 124, 3436–3447 (2019).
- Shibazaki, Y. et al. Sound velocity measurements in dhcp-FeH up to 70 GPa with inelastic X-ray scattering: Implications for the composition of the Earth's core. *Earth Planet. Sci. Lett.* 313, 79–85 (2012).
- Lin, J. F. et al. Magnetic transition and sound velocities of Fe₃S at high pressure: implications for Earth and planetary cores. *Earth Planet. Sci. Lett.* 226, 33–40 (2004).
- Kamada, S. et al. The sound velocity measurements of Fe₃S. Am. Mineralogist 99, 98–251 (2014).
- Gao, L. et al. Pressure-induced magnetic transition and sound velocities of Fe₃C: Implications for carbon in the Earth's inner core. *Geophys. Res. Lett.* 35, L17306 (2008).
- Fei, Y. & Brosh, E. Experimental study and thermodynamic calculations of phase relations in the Fe-C system at high pressure. *Earth Planet. Sci. Lett.* 408, 155–162 (2014).
- Mashino, I., Miozzi, F., Hirose, K., Morard, G. & Sinmyo, R. Melting experiments on the Fe-C binary system up to 255 GPa: constraints on the carbon content in the Earth's core. Earth Planet. Sci. Lett. 515, 135–144 (2019).

- 52. Chen, B. et al. Magneto-elastic coupling in compressed Fe₇C₃ supports carbon inEarth's inner core. *Geophys. Res. Lett.* **39**, L18301 (2012).
- Sata, N. et al. Compression of FeSi, Fe₃C, Fe₀. ₉₅O, and FeS under the core pressures and implication for light element in the Earth's core. *J. Geophys. Res.* 115, B09204 (2010).
- Pamato, M. G. et al. Equation of state of hcp Fe-C-Si alloys and the effect of C incorporation mechanism on the density of hcp Fe alloys at 300 K. J. Geophys. Res. 125, e2020]B020159 (2020).
- Yang, J., Fei, Y., Hu, X., Greenberg, E. & Prakapenka, V. B. Effect of carbon on the volume of solid iron at high pressure: Implications for carbon substitution in iron structures and carbon content in the Earth's inner core. *Minerals* 9, 720 (2019).
- Liu, J., Lin, J. F., Alatas, A. & Bi, W. Sound velocities of bcc-Fe and Fe_{0.85}Si_{0.15} alloy at high pressure and temperature. *Phys. Earth Planet. Inter.* 233, 24–32 (2014)

Acknowledgements

We thank Q. Shen (the Sate Key Lab of Advanced Technology for Materials Synthesis), M. Yang and C. Shen (the Materials Research and Testing Center) for sample preparation, and R. Tao for chemical analysis of the samples. This work was supported by the National Science Foundation of China (grant 41874103, 51932006), and the Fundamental Research funds for the Central Universities (grants WUT 14050). Additional support for this collaborative research was provided by the Carnegie Institution for Science, and National Science Foundation grants (EAR-1619868) to YF.

Author contributions

H.H. and Y.F. designed the project. H.H., L.F., X.L., F.X., Y.W., G.Y., C.L., Q.W., J.W., X.W., L.C., and Y.F. participated the data collection and analysis. H.H and Y.F. were responsible for data interpretation and wrote the manuscript. L.F., X.L., F.X., Y.W., G.Y., C.L., Q.W., J.W., X.W., and L.C. participated the discussion and provided comments on the manuscript.

Competing interests

The authors declare no competing interests.

Additional information

Supplementary information The online version contains supplementary material available at https://doi.org/10.1038/s41467-022-28255-2.

Correspondence and requests for materials should be addressed to Yingwei Fei.

Peer review information *Nature Communications* thanks Guillaume Morard and the other, anonymous, reviewer(s) for their contribution to the peer review of this work. Peer reviewer reports are available.

Reprints and permission information is available at http://www.nature.com/reprints

Publisher's note Springer Nature remains neutral with regard to jurisdictional claims in published maps and institutional affiliations.

CC Op

Open Access This article is licensed under a Creative Commons Attribution 4.0 International License, which permits use, sharing,

adaptation, distribution and reproduction in any medium or format, as long as you give appropriate credit to the original author(s) and the source, provide a link to the Creative Commons license, and indicate if changes were made. The images or other third party material in this article are included in the article's Creative Commons license, unless indicated otherwise in a credit line to the material. If material is not included in the article's Creative Commons license and your intended use is not permitted by statutory regulation or exceeds the permitted use, you will need to obtain permission directly from the copyright holder. To view a copy of this license, visit https://creativecommons.org/licenses/by/4.0/.

© The Author(s) 2022

Terms and Conditions

Springer Nature journal content, brought to you courtesy of Springer Nature Customer Service Center GmbH ("Springer Nature").

Springer Nature supports a reasonable amount of sharing of research papers by authors, subscribers and authorised users ("Users"), for small-scale personal, non-commercial use provided that all copyright, trade and service marks and other proprietary notices are maintained. By accessing, sharing, receiving or otherwise using the Springer Nature journal content you agree to these terms of use ("Terms"). For these purposes, Springer Nature considers academic use (by researchers and students) to be non-commercial.

These Terms are supplementary and will apply in addition to any applicable website terms and conditions, a relevant site licence or a personal subscription. These Terms will prevail over any conflict or ambiguity with regards to the relevant terms, a site licence or a personal subscription (to the extent of the conflict or ambiguity only). For Creative Commons-licensed articles, the terms of the Creative Commons license used will apply.

We collect and use personal data to provide access to the Springer Nature journal content. We may also use these personal data internally within ResearchGate and Springer Nature and as agreed share it, in an anonymised way, for purposes of tracking, analysis and reporting. We will not otherwise disclose your personal data outside the ResearchGate or the Springer Nature group of companies unless we have your permission as detailed in the Privacy Policy.

While Users may use the Springer Nature journal content for small scale, personal non-commercial use, it is important to note that Users may not:

- 1. use such content for the purpose of providing other users with access on a regular or large scale basis or as a means to circumvent access control:
- 2. use such content where to do so would be considered a criminal or statutory offence in any jurisdiction, or gives rise to civil liability, or is otherwise unlawful;
- 3. falsely or misleadingly imply or suggest endorsement, approval, sponsorship, or association unless explicitly agreed to by Springer Nature in writing:
- 4. use bots or other automated methods to access the content or redirect messages
- 5. override any security feature or exclusionary protocol; or
- 6. share the content in order to create substitute for Springer Nature products or services or a systematic database of Springer Nature journal content.

In line with the restriction against commercial use, Springer Nature does not permit the creation of a product or service that creates revenue, royalties, rent or income from our content or its inclusion as part of a paid for service or for other commercial gain. Springer Nature journal content cannot be used for inter-library loans and librarians may not upload Springer Nature journal content on a large scale into their, or any other, institutional repository.

These terms of use are reviewed regularly and may be amended at any time. Springer Nature is not obligated to publish any information or content on this website and may remove it or features or functionality at our sole discretion, at any time with or without notice. Springer Nature may revoke this licence to you at any time and remove access to any copies of the Springer Nature journal content which have been saved.

To the fullest extent permitted by law, Springer Nature makes no warranties, representations or guarantees to Users, either express or implied with respect to the Springer nature journal content and all parties disclaim and waive any implied warranties or warranties imposed by law, including merchantability or fitness for any particular purpose.

Please note that these rights do not automatically extend to content, data or other material published by Springer Nature that may be licensed from third parties.

If you would like to use or distribute our Springer Nature journal content to a wider audience or on a regular basis or in any other manner not expressly permitted by these Terms, please contact Springer Nature at

onlineservice@springernature.com