

Initial report of 1976 U.S. Atmospheric Model methodology

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1 Background

“The USSA mathematical model divides the atmosphere into layers with an assumed linear distribution of absolute temperature T against geopotential altitude h . The other two values (pressure P and density ρ) are computed by simultaneously solving the equations resulting from:

$$\frac{dP}{dh} = -\rho g \quad (1)$$

$$P = \rho R_{\text{specific}} T \quad (2)$$

where R_{specific} is the specific gas constant for dry air and g is the standard gravitational acceleration.” Equation (1) evaluates the vertical pressure variation with the mean sea level as a boundary condition ($P_0 = 101,325$ Pa) and Equation (2) is the ideal gas law in molar form [1].

Combining equations (1) and (2) to find an expression for the pressure:

$$\ln P \Big|_{P_0}^{P_1} = -\frac{g}{R_{\text{specific}}} \int_{h_0}^{h_1} \frac{dh}{T(h)} \quad (3)$$

where the minimum geopotential altitude above sea level is set to $h_0 = -1,524$ m. From [The Engineer ToolBox](#), various temperatures have been given for their corresponding geopotential altitudes in Table 1 [2].

Table 1: USSA1976 Temperature for Geopotential Altitude - Imperial (BG) Units

| h [ft] | T [°F] |
|----------|----------|
| −5000 | 76.84 |
| 0 | 59 |
| 5000 | 41.17 |
| 10000 | 23.36 |
| 15000 | 5.55 |
| 20000 | −12.26 |
| 25000 | −30.05 |
| 30000 | −47.83 |
| 35000 | −65.61 |
| 40000 | −69.70 |
| 45000 | −69.70 |
| 50000 | −69.70 |
| 60000 | −69.70 |
| 70000 | −67.42 |
| 80000 | −61.98 |
| 90000 | −56.54 |
| 100000 | −51.10 |
| 150000 | 19.40 |
| 200000 | −19.78 |
| 250000 | −88.77 |

²It is interesting to note that if the maximum value of h is fed into (4) then the numerical value for the integral in the RHS becomes $323.7972921077598 \pm 1.173024627261422 \times 10^{-10}$.

where C_d is the drag coefficient and A is the cross-sectional area. The method which computes drag takes in ECEF position,³ velocity, and ballistic parameters (C_d , A , m) to output an acceleration due to force vector. The acceleration due to drag can be found using Newton's second law; $\vec{a}_d = \vec{F}_d/m$.

3 An Example

An object of mass $m = 100\text{ kg}$ is traveling with a velocity of $\vec{v} = (v_x, v_y, v_h) = (20, 100, 3000)\text{ m/s}$ at an ECEF position of $(0, 0, 10000)\text{ m}$ where the latter component is the altitude. The drag coefficient is 1 and the cross-sectional area of the object is 100 m^2 . From the applied model for atmosphere here and the fourth degree polynomial for $T(h)$ its temperature is 226.956 K and it follows that its pressure is 21846.097 Pa and the density of air at that altitude is 0.335 kg/m^3 . The applied drag force at that altitude would be $\vec{F}_d = (1.677 \times 10^5, 1.509 \times 10^8, 6.707 \times 10^3)\text{ N}$; ergo, the length of the acceleration due to drag would be $|\vec{a}_d| = 1508978.973\text{ m/s}^2$.⁴

Bibliography

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- [5] National Oceanic and Atmospheric Administration, National Aeronautics and Space Administration, United States Air Force, *U.S Standard Atmosphere, 1976*, US Department of Commerce/NOAA/NESDIS, October 1976.
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³The altitude is the only relevant coordinate when it comes to this specific calculation as Equation (5)'s only parameter dependent on position is p .

⁴Although it does change with altitude, $g = 9.81\text{ m/s}^2$ is used for all altitudes in this implementation.