

Churning the Ocean: Time variable gravity for ocean applications



[Samudra Manthan, CC BY 3.0](#)

Roelof Rietbroek, 28 Sept. 2022



Analogies with today's talk



Find this presentation on <https://github.com/strawpants/grace-hackweek-ocean>



Analogy's with today's talk

- Mount Mandara: Adding mass to the ocean (e.g. melt water)



Find this presentation on <https://github.com/strawpants/grace-hackweek-ocean>



Analogyies with today's talk

- Mount Mandara: Adding mass to the ocean (e.g. melt water)
- Snake Vasuki: set the ocean in motion (e.g. forcing from wind stress)



Find this presentation on <https://github.com/strawpants/grace-hackweek-ocean>

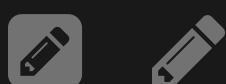


Analogy with today's talk

- Mount Mandara: Adding mass to the ocean (e.g. melt water)
- Snake Vasuki: set the ocean in motion (e.g. forcing from wind stress)
- Central Question: How can we use time variable gravity to observe ocean signals?



Find this presentation on <https://github.com/strawpants/grace-hackweek-ocean>



Adding mass to the oceans



Adding mass to the oceans

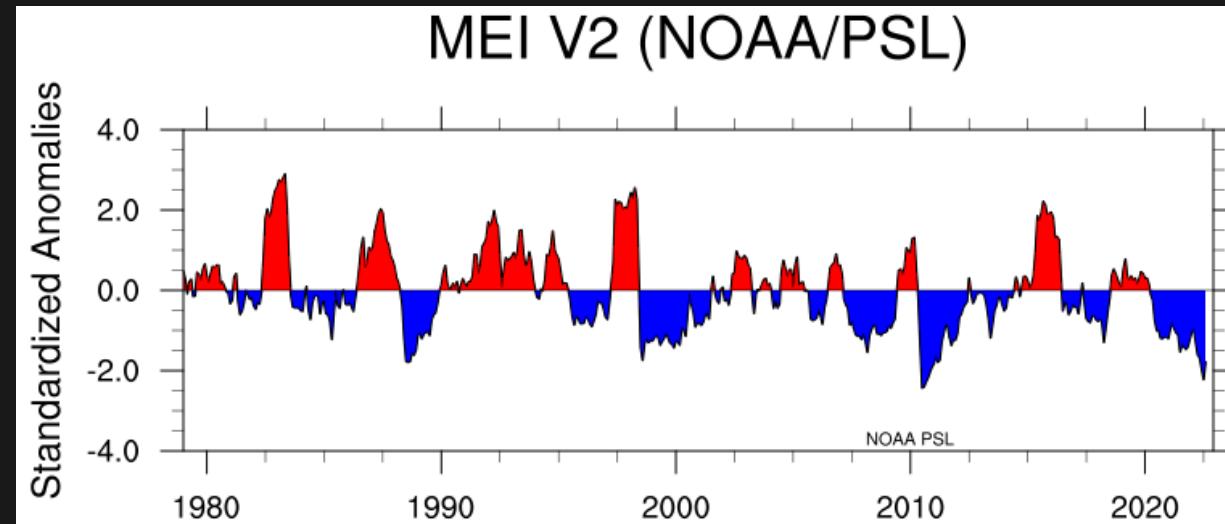
(or removing it)



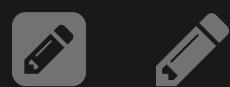
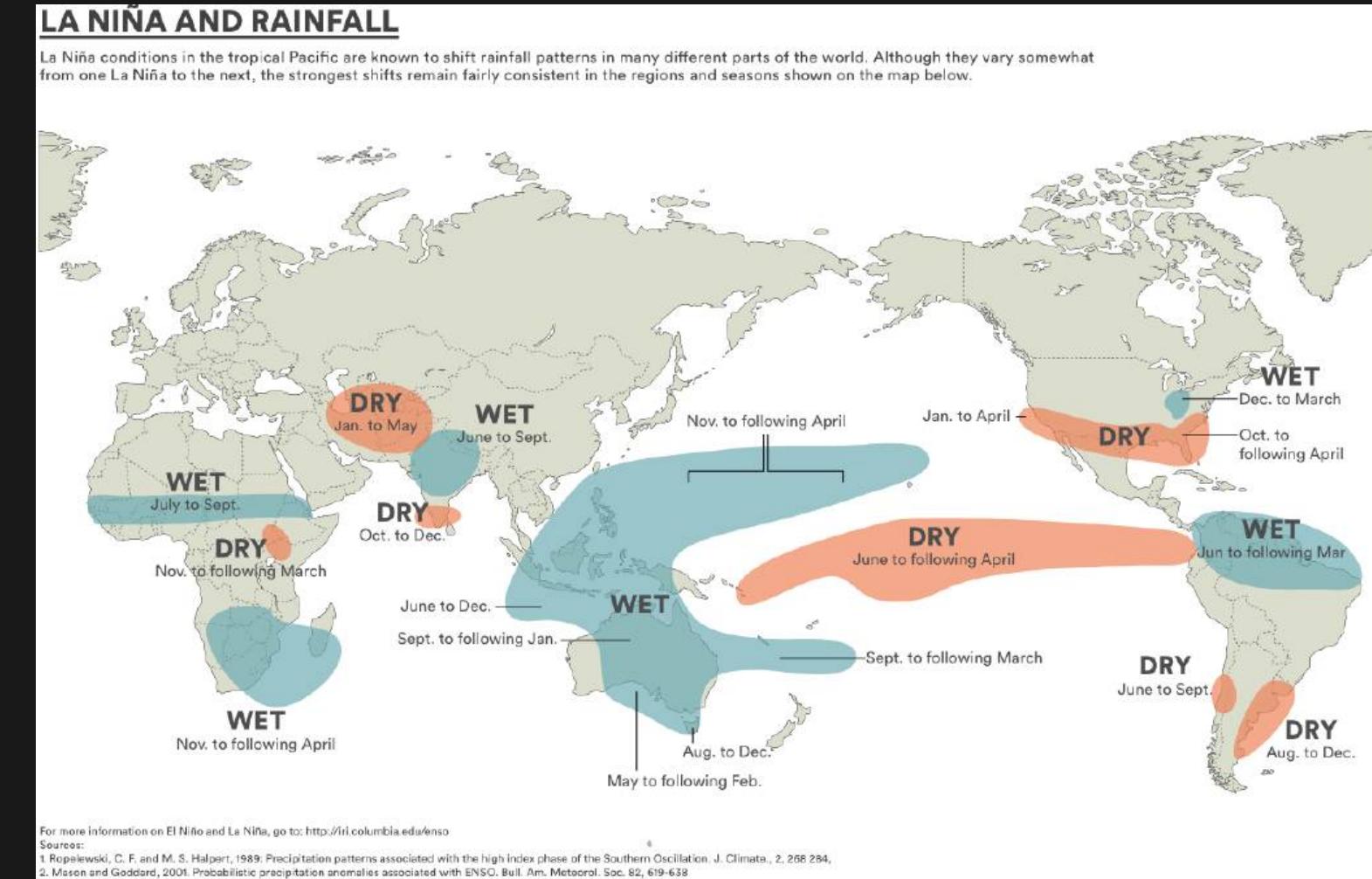
2022 Floods in Pakistan, image: AFP



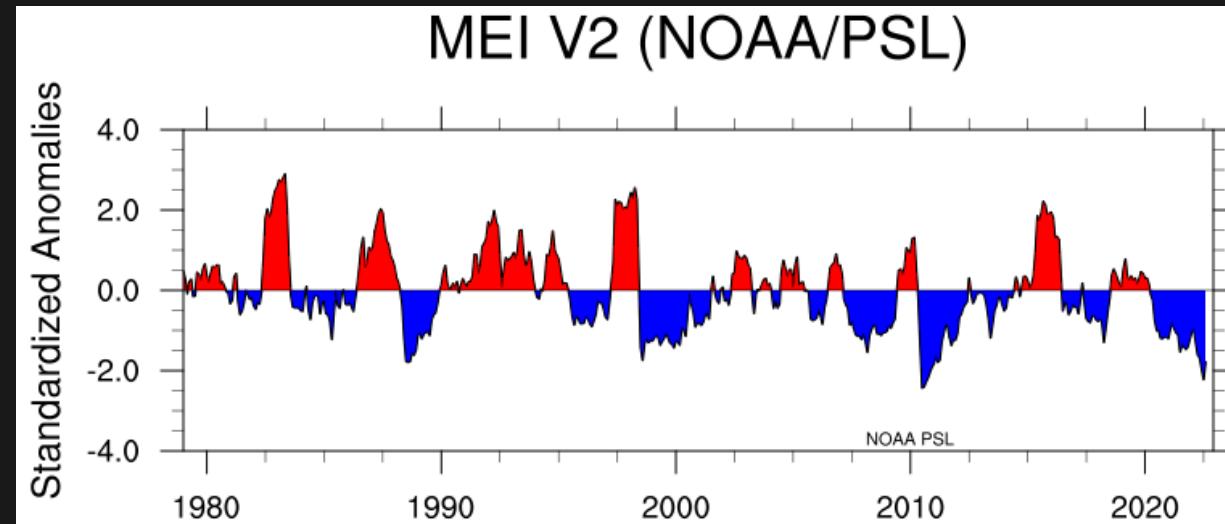
Does La Niña affect sea level?



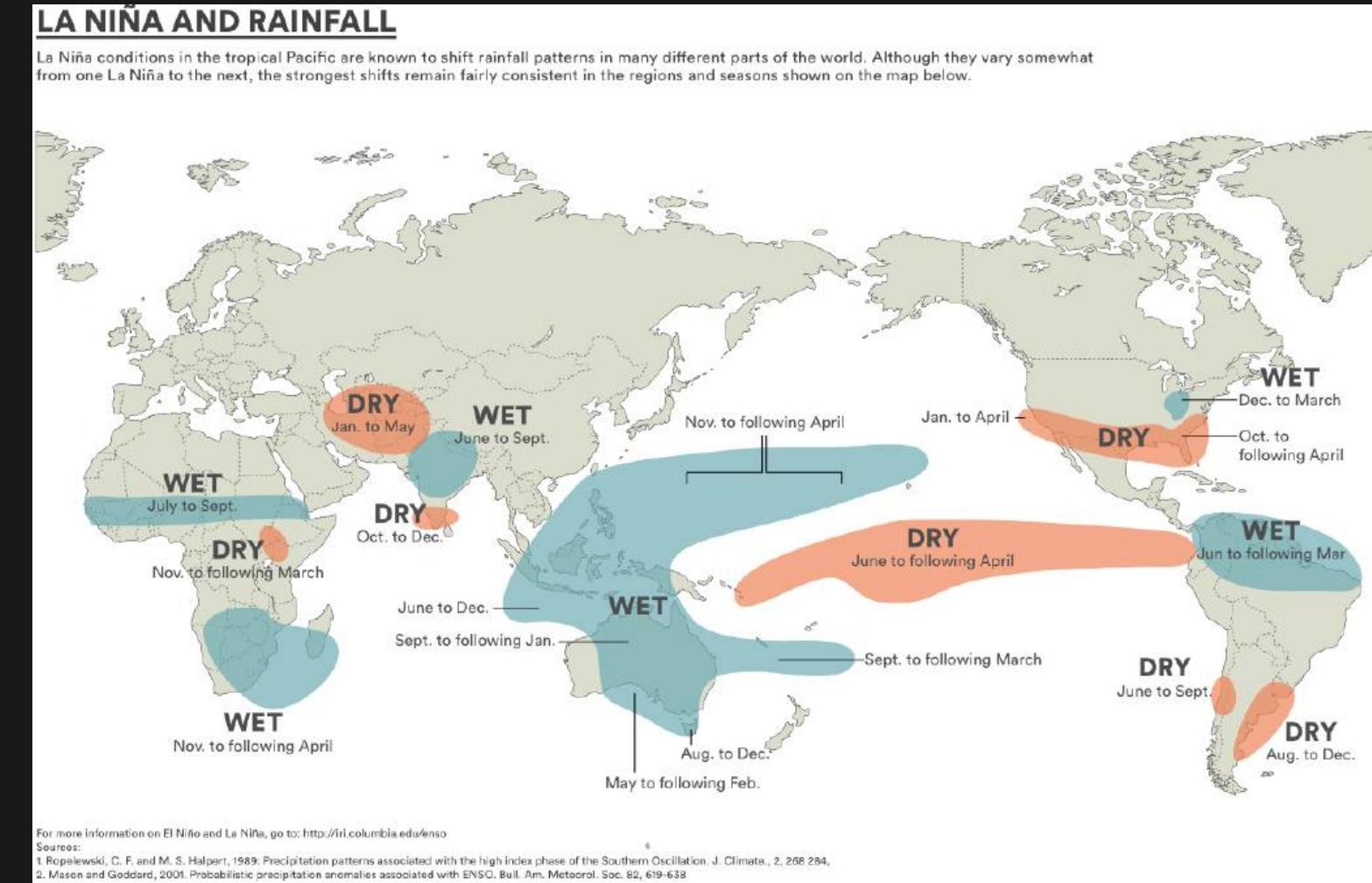
Multivariate ENSO index, <https://psl.noaa.gov/enso>



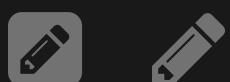
Does La Niña affect sea level?



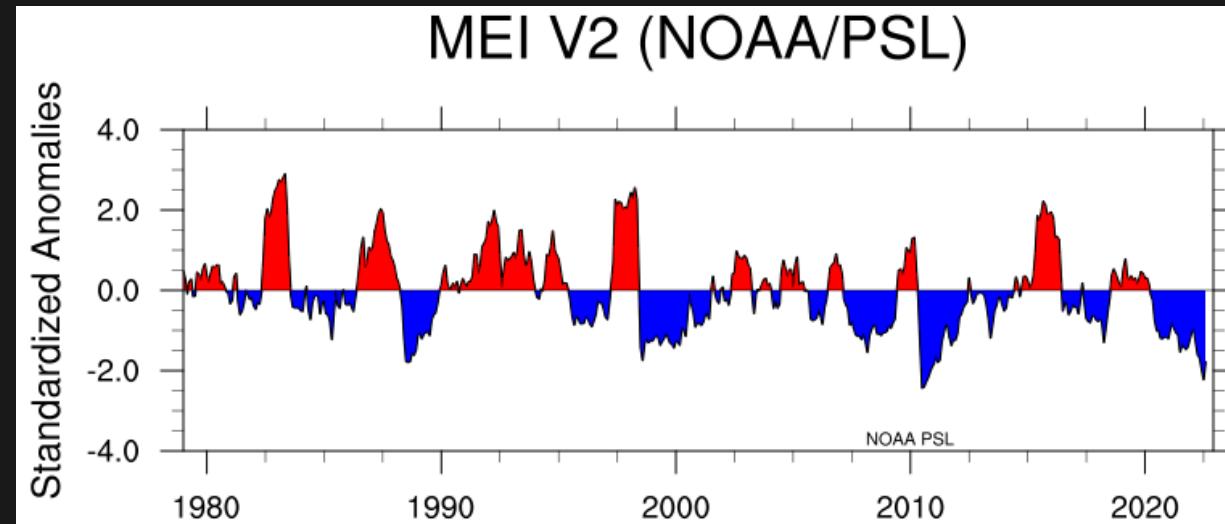
Multivariate ENSO index, <https://psl.noaa.gov/enso>



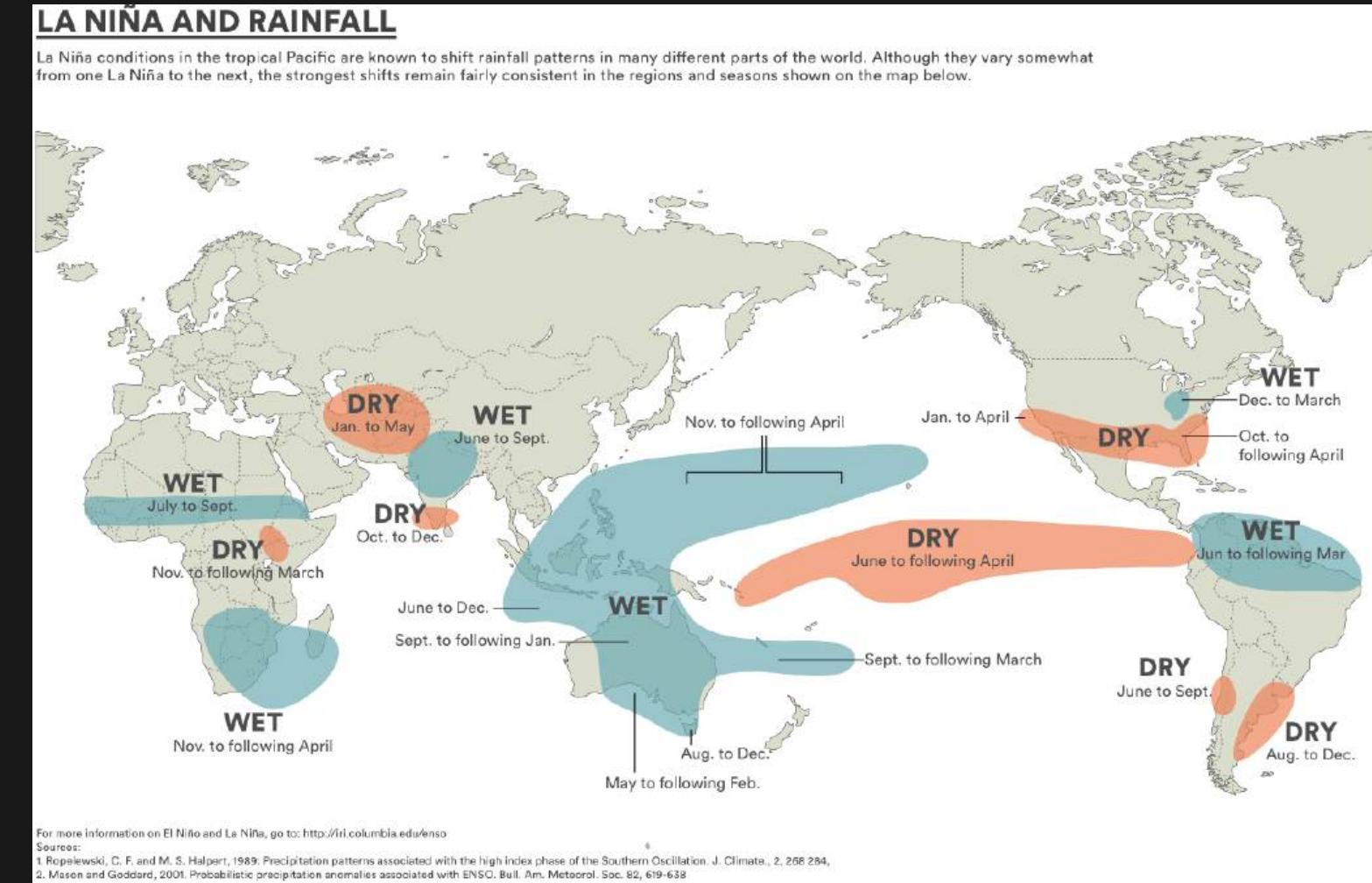
- Stronger trade winds -> warm/wet ocean blob in the western Pacific



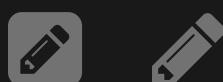
Does La Niña affect sea level?



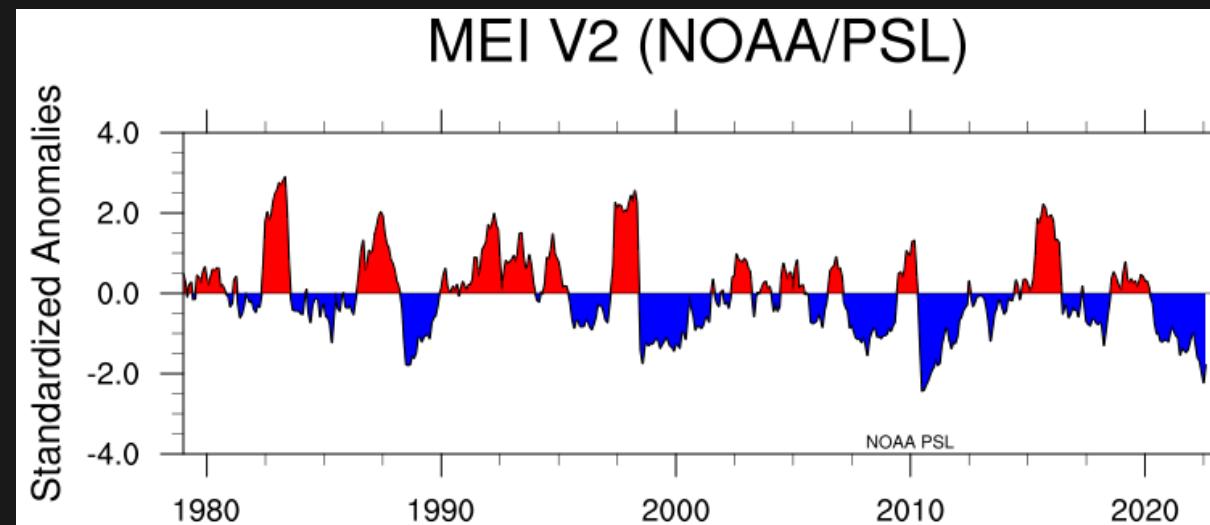
Multivariate ENSO index, <https://psl.noaa.gov/enso>



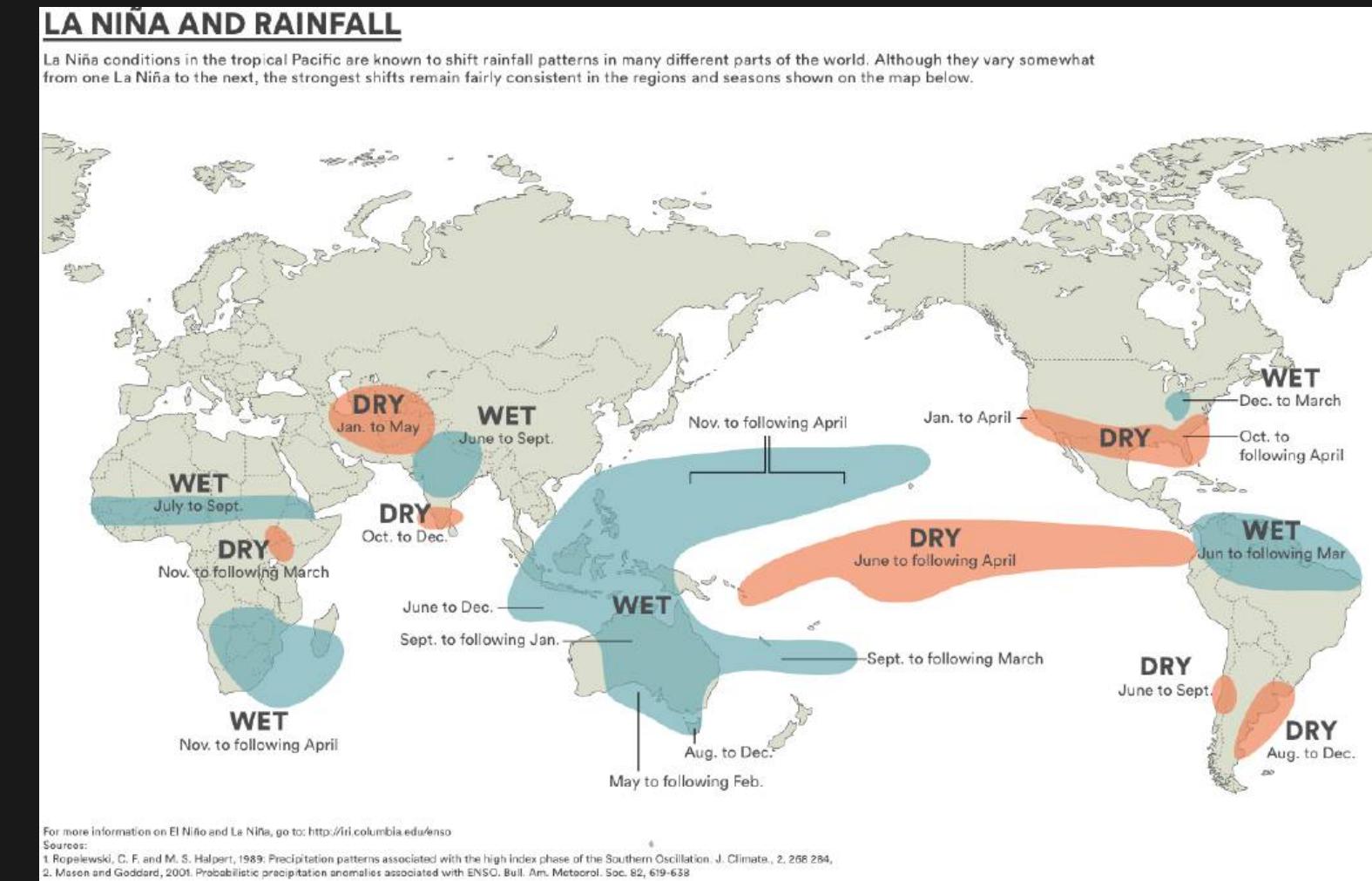
- Stronger trade winds -> warm/wet ocean blob in the western Pacific
- Wet conditions in the West Pacific cause extreme precipitation



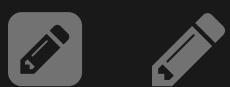
Does La Niña affect sea level?



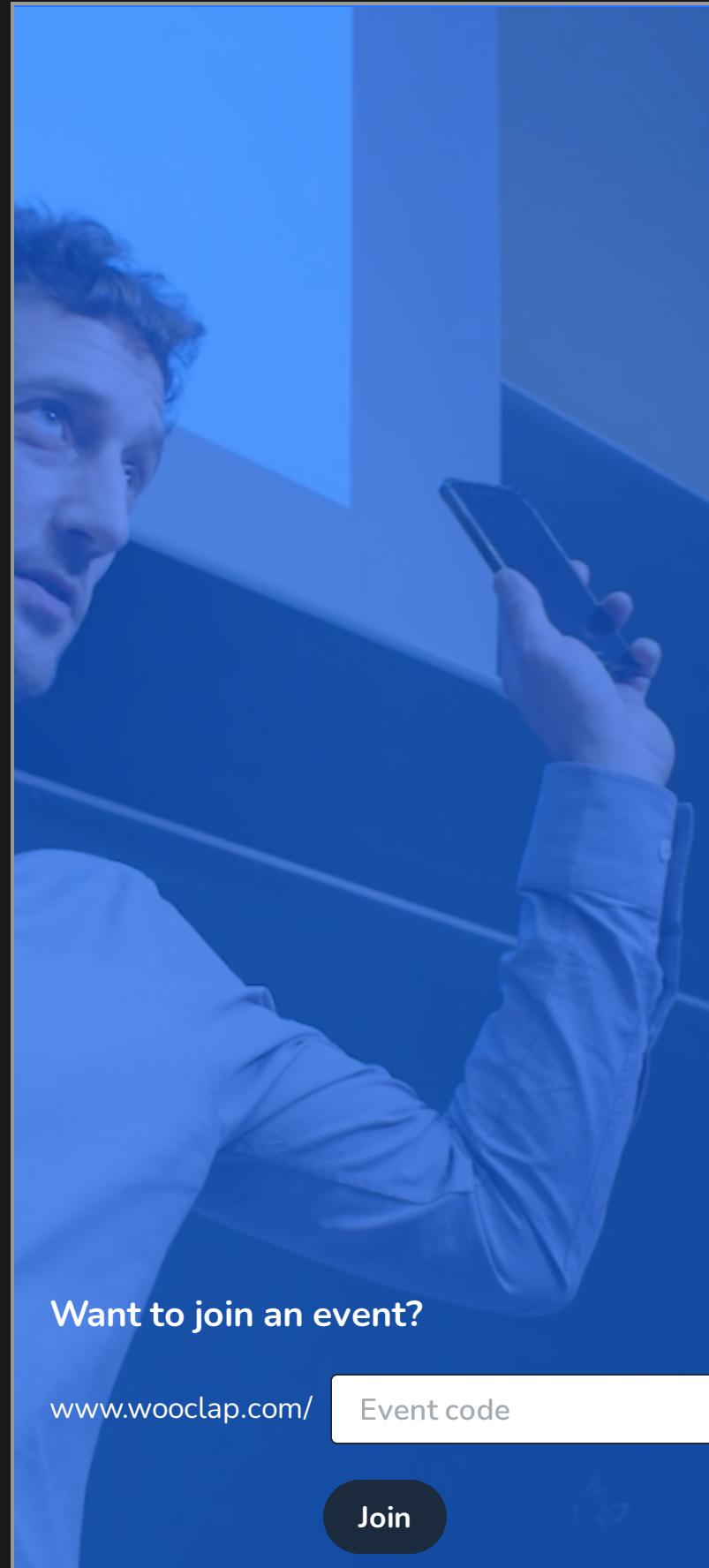
Multivariate ENSO index, <https://psl.noaa.gov/enso>



- Stronger trade winds -> warm/wet ocean blob in the western Pacific
- Wet conditions in the West Pacific cause extreme precipitation
- Teleconnections with Monsoon, Africa..



Quiz time!



Want to join an event?

www.wooclap.com/

Join

Choose a method to log in
[or sign up](#)

Your email address 

OR

 Sign up with Facebook

 Sign up with Google

 Sign up with LinkedIn

 Sign up with Microsoft

OR

Log in with your institution

Your institution 



We use [cookies](#) to improve the general experience on the platform, provide users chat support and deliver targeted ads on other websites.

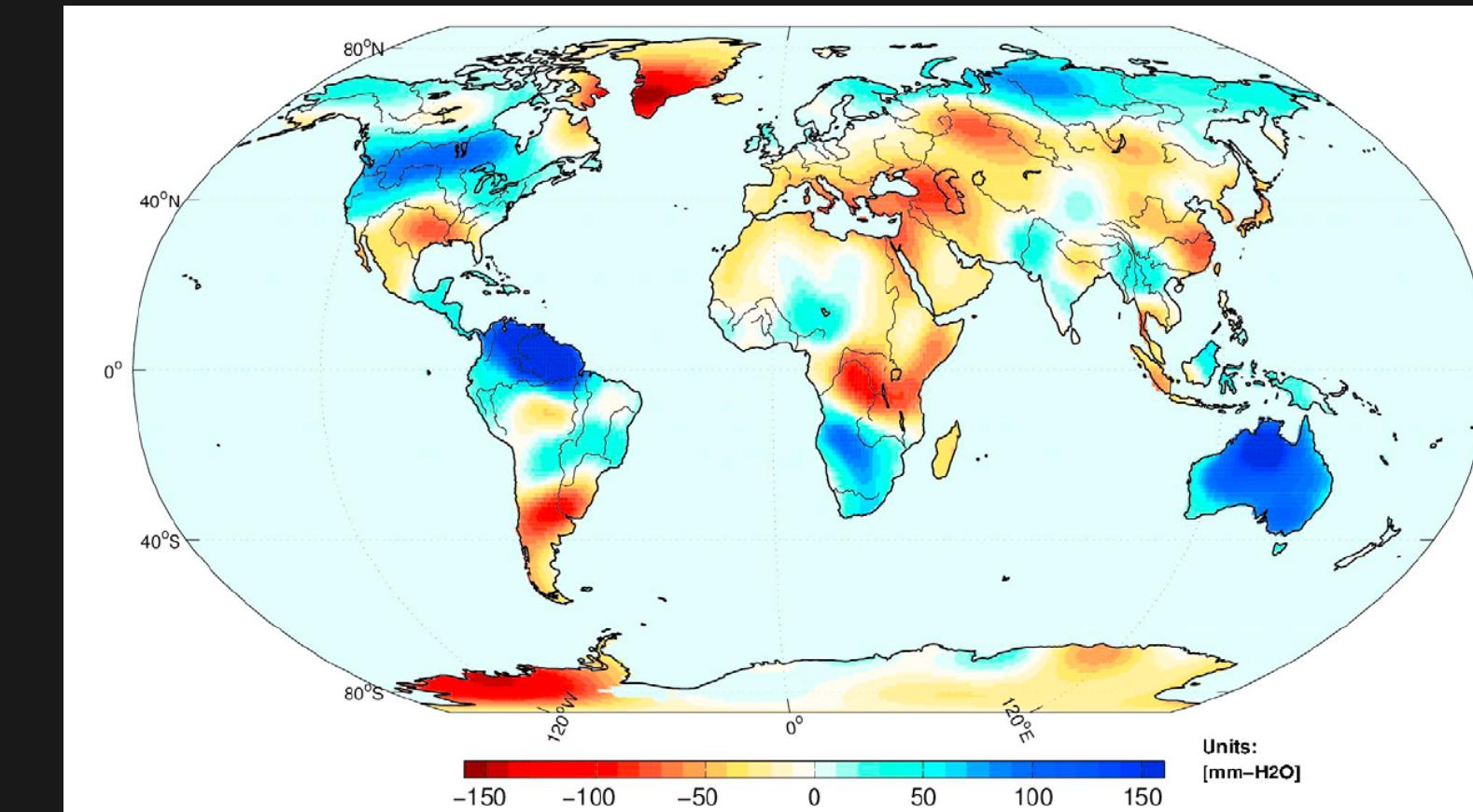
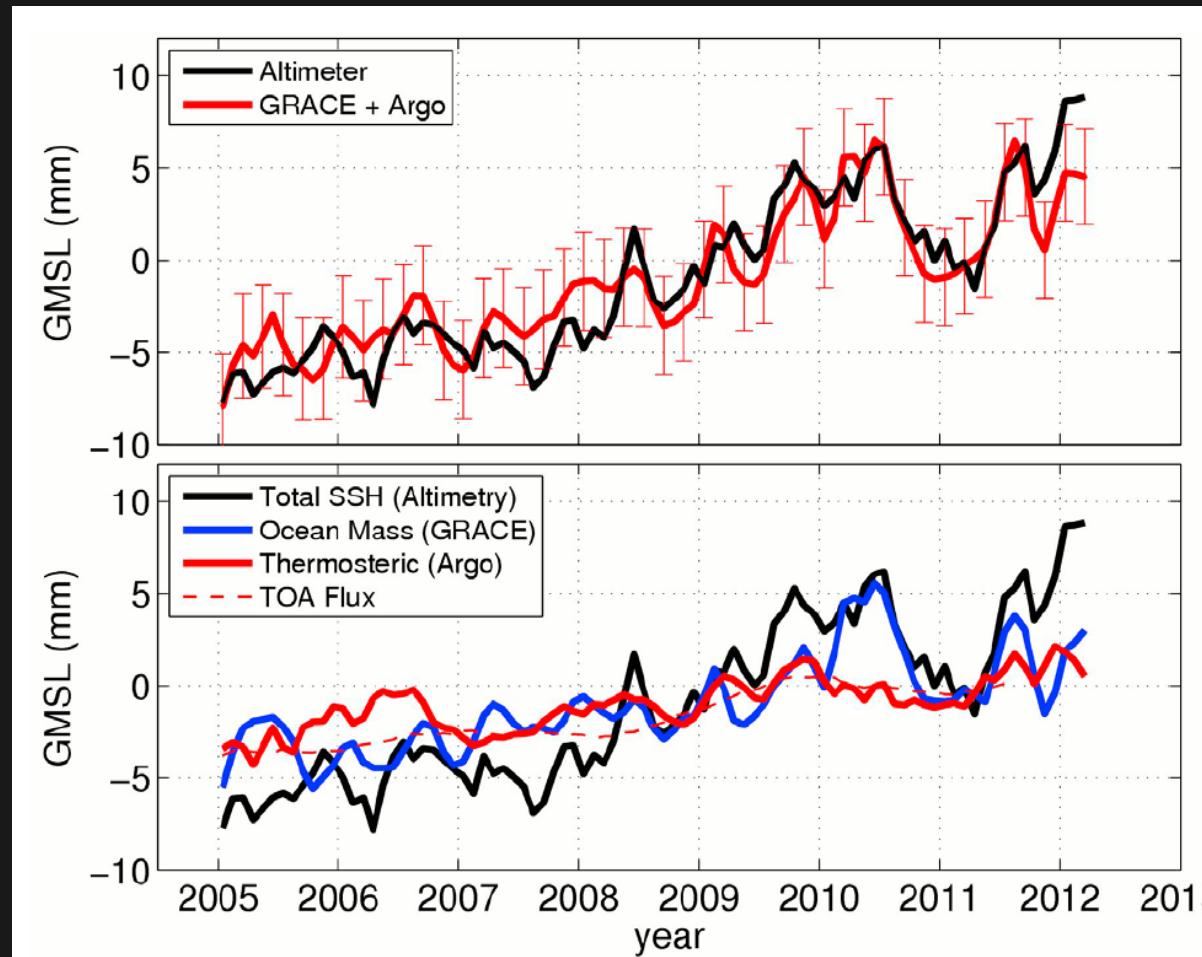
Accept all

Reject All

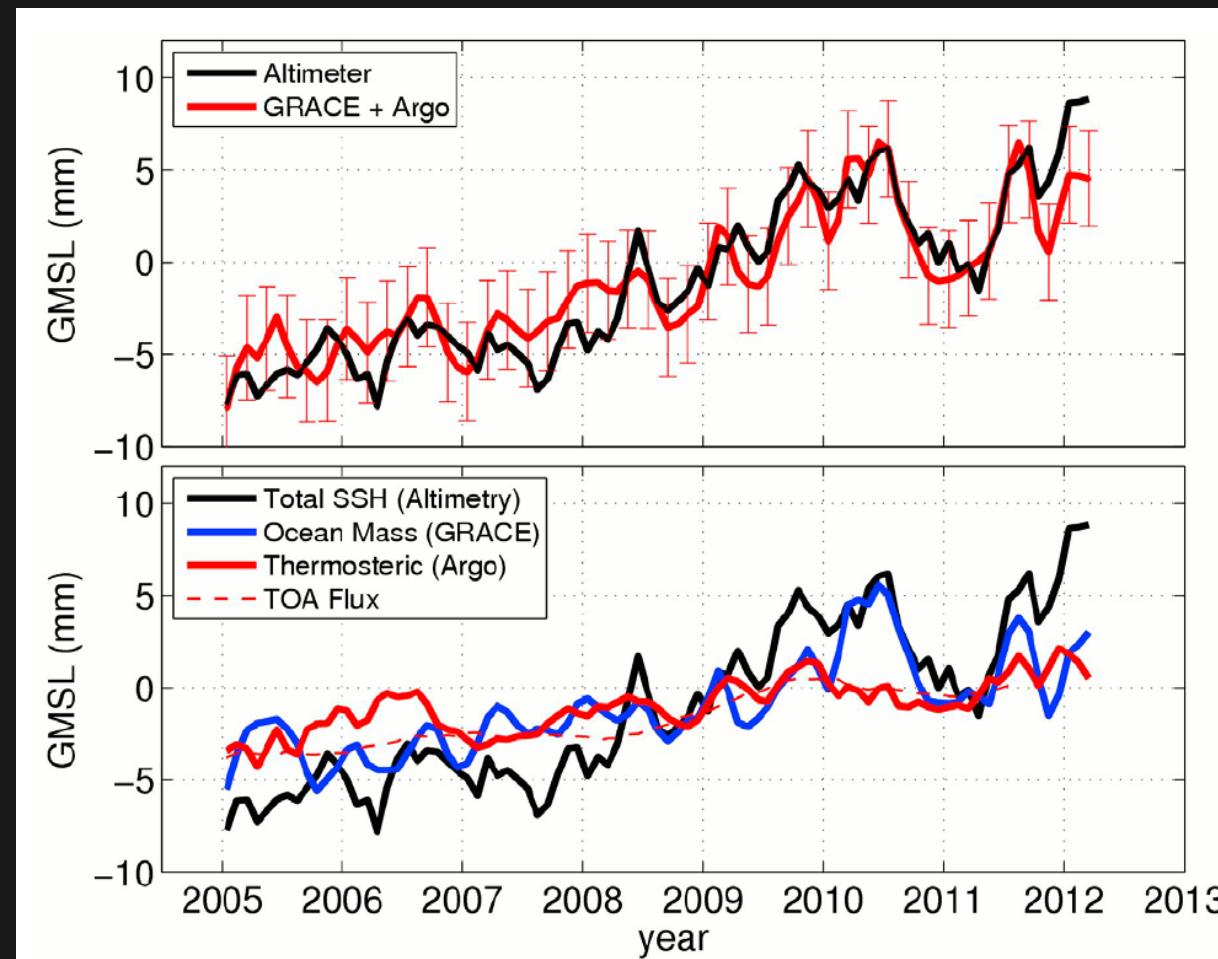
Customise



2010-2011 La Nina event

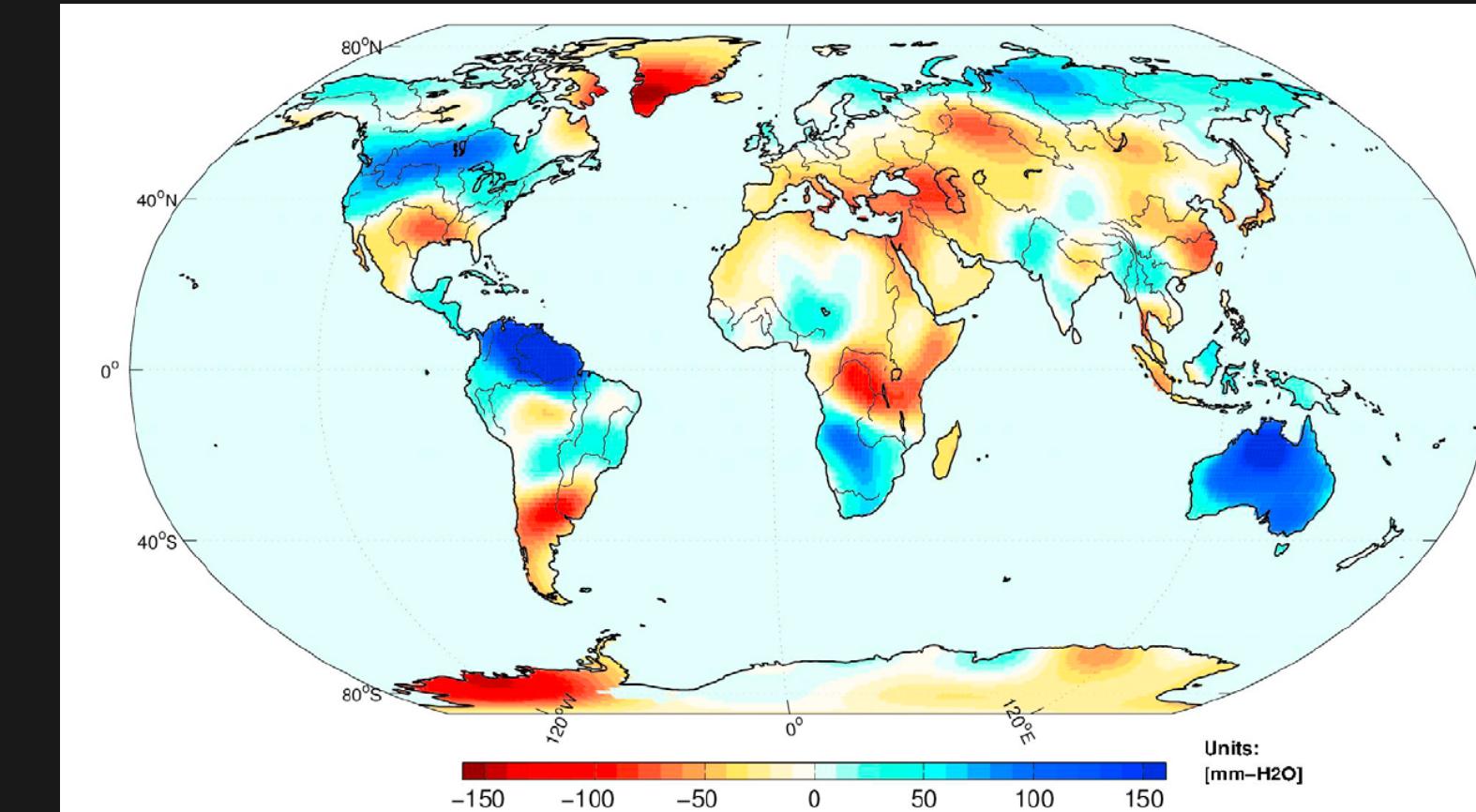


2010-2011 La Nina event

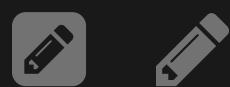


Global mean sea level variations (total, mass driven and 'steric'),
From Böning et al. 2012

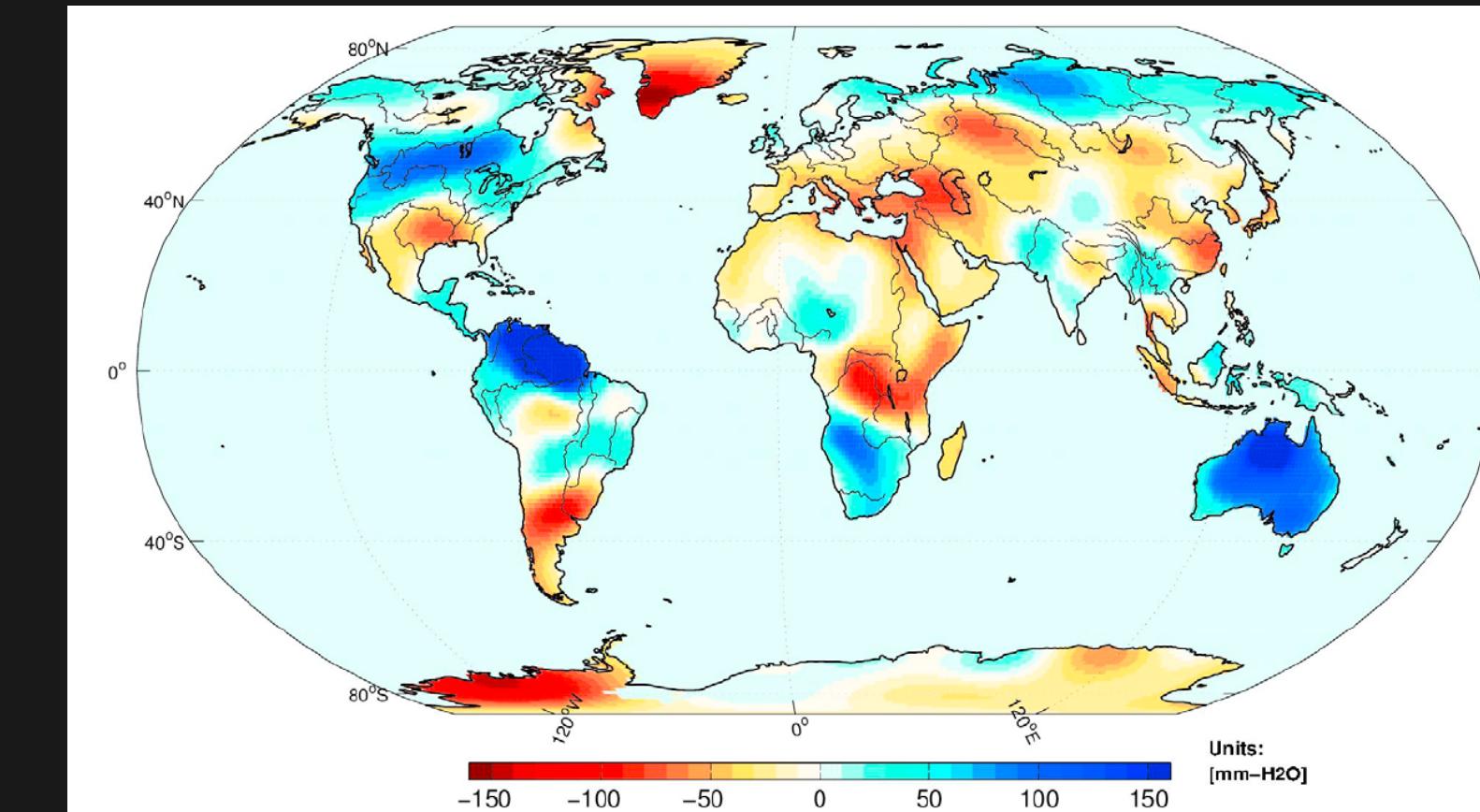
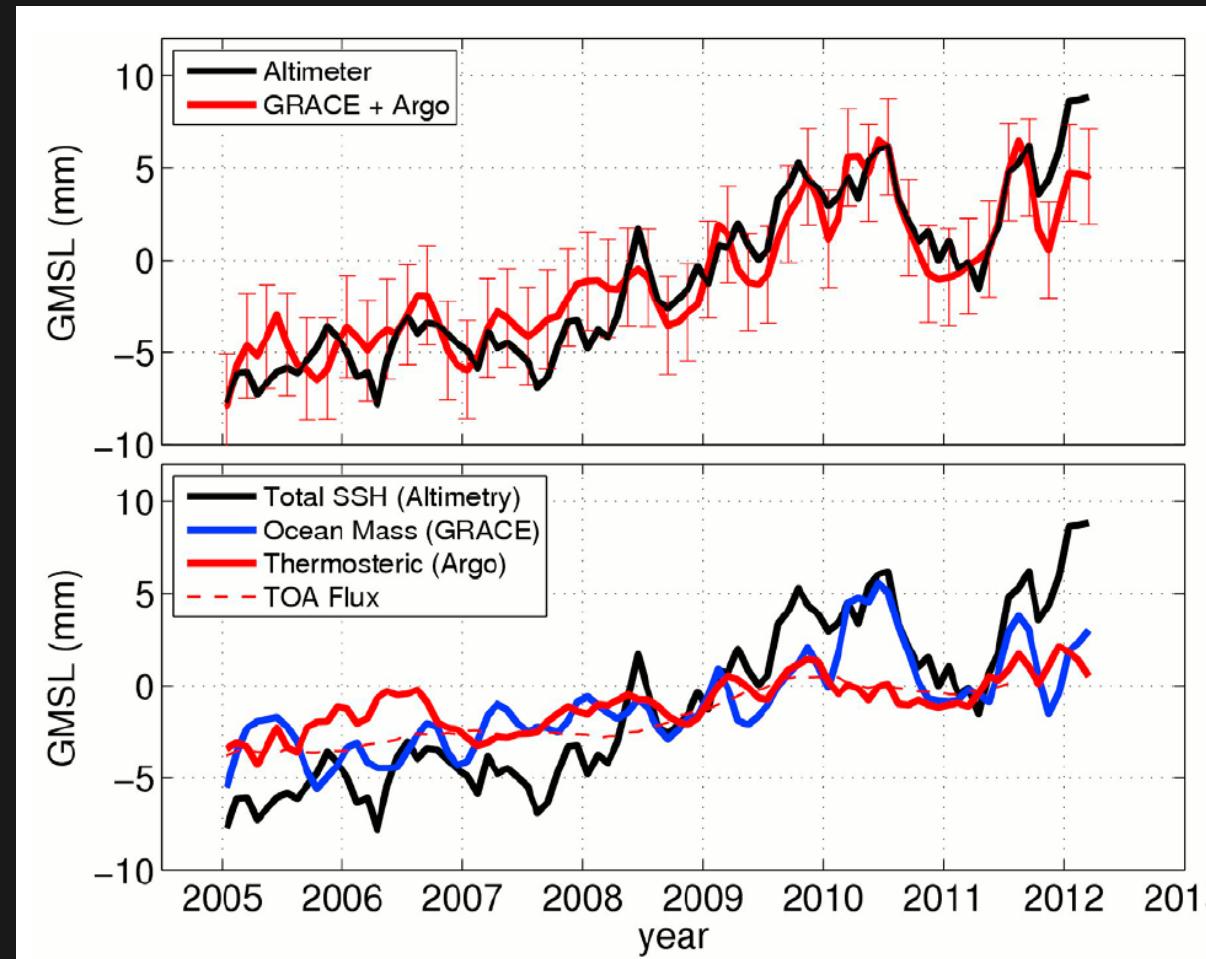
- Detectable with radar altimetry



Terrestrial water storage change (2010-2011) from GRACE, From Böning et al. 2012



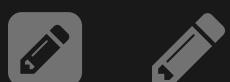
2010-2011 La Nina event



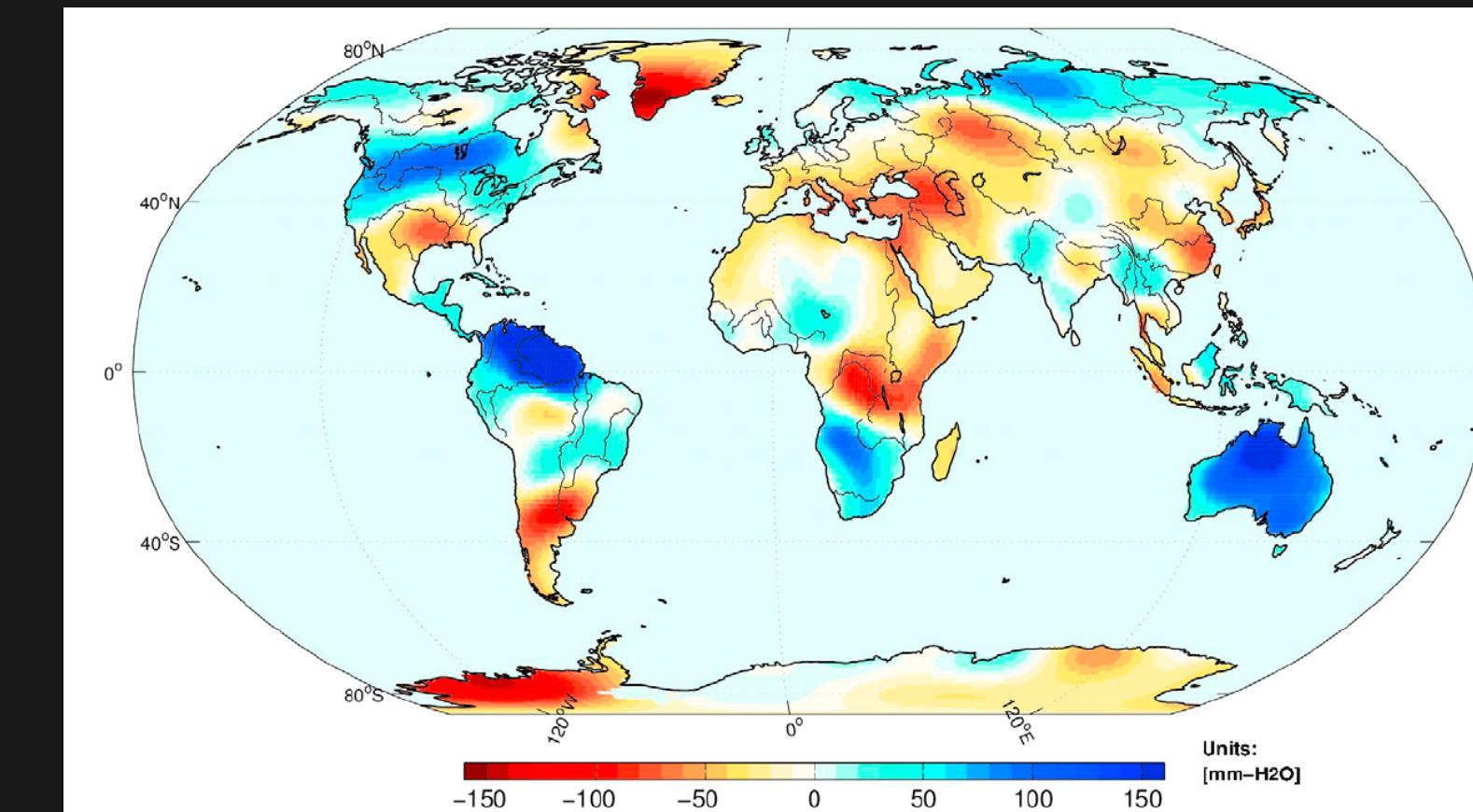
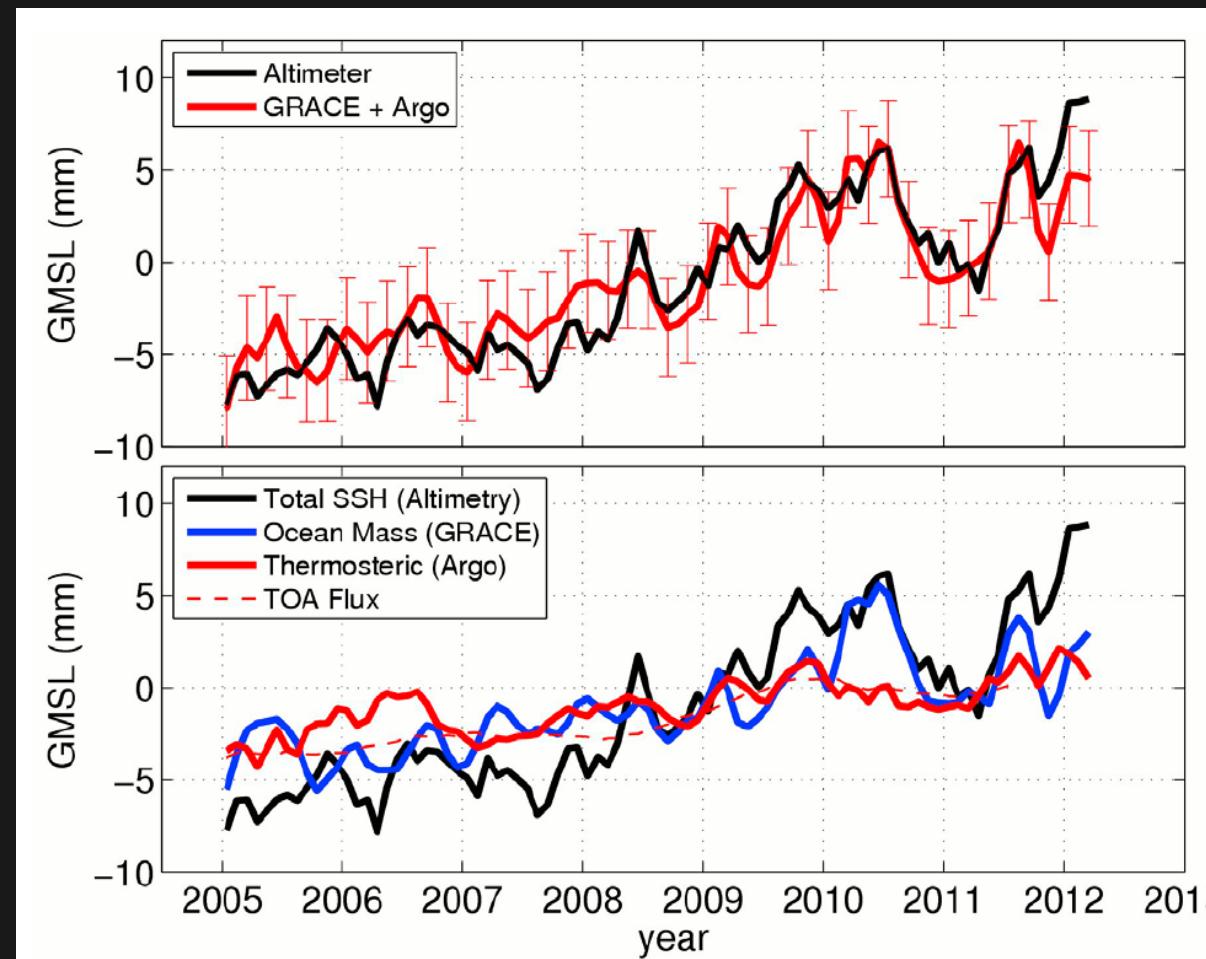
Terrestrial water storage change (2010-2011) from GRACE, From Böning et al. 2012

Global mean sea level variations (total, mass driven and 'steric'),
From Böning et al. 2012

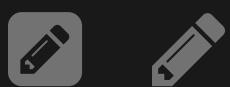
- Detectable with radar altimetry
- Can be explained by mass-driven sea level (GRACE)



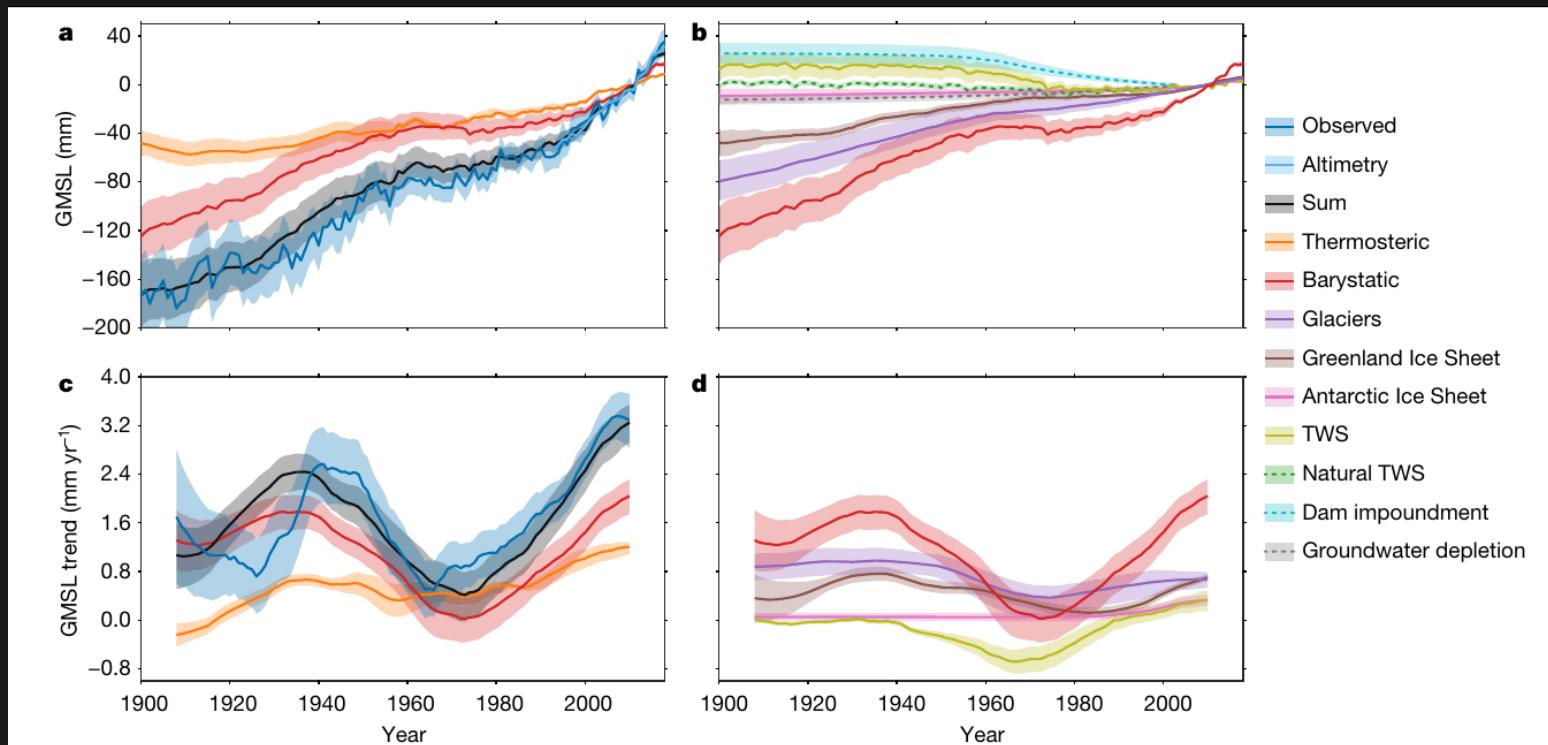
2010-2011 La Nina event



- Detectable with radar altimetry
- Can be explained by mass-driven sea level (GRACE)
- Hotspot: Australia



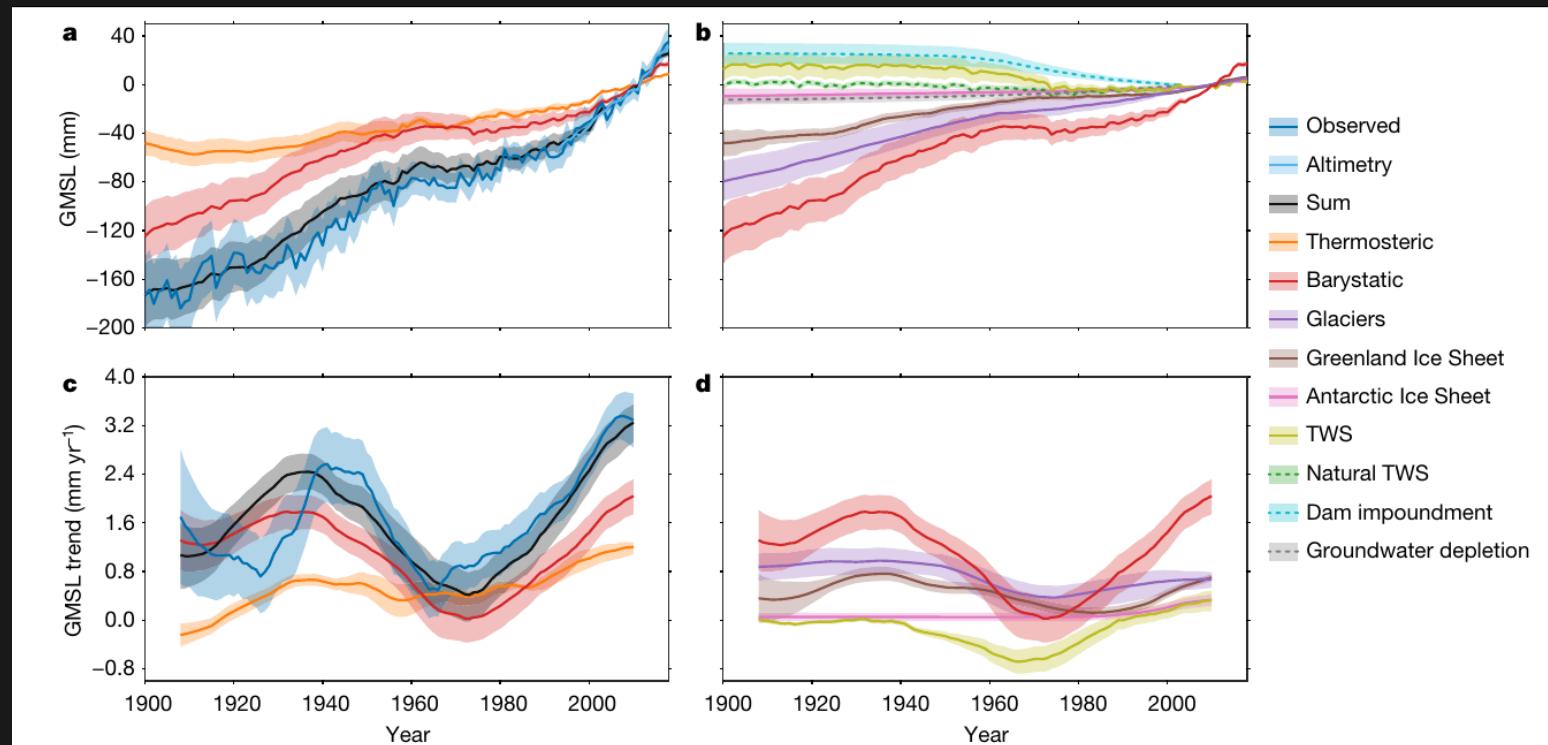
Sea level contributions since 1900



Historic sea level can be largely explained with different (modelled) sea level contributions. From Frederikse et al. 2020



Sea level contributions since 1900

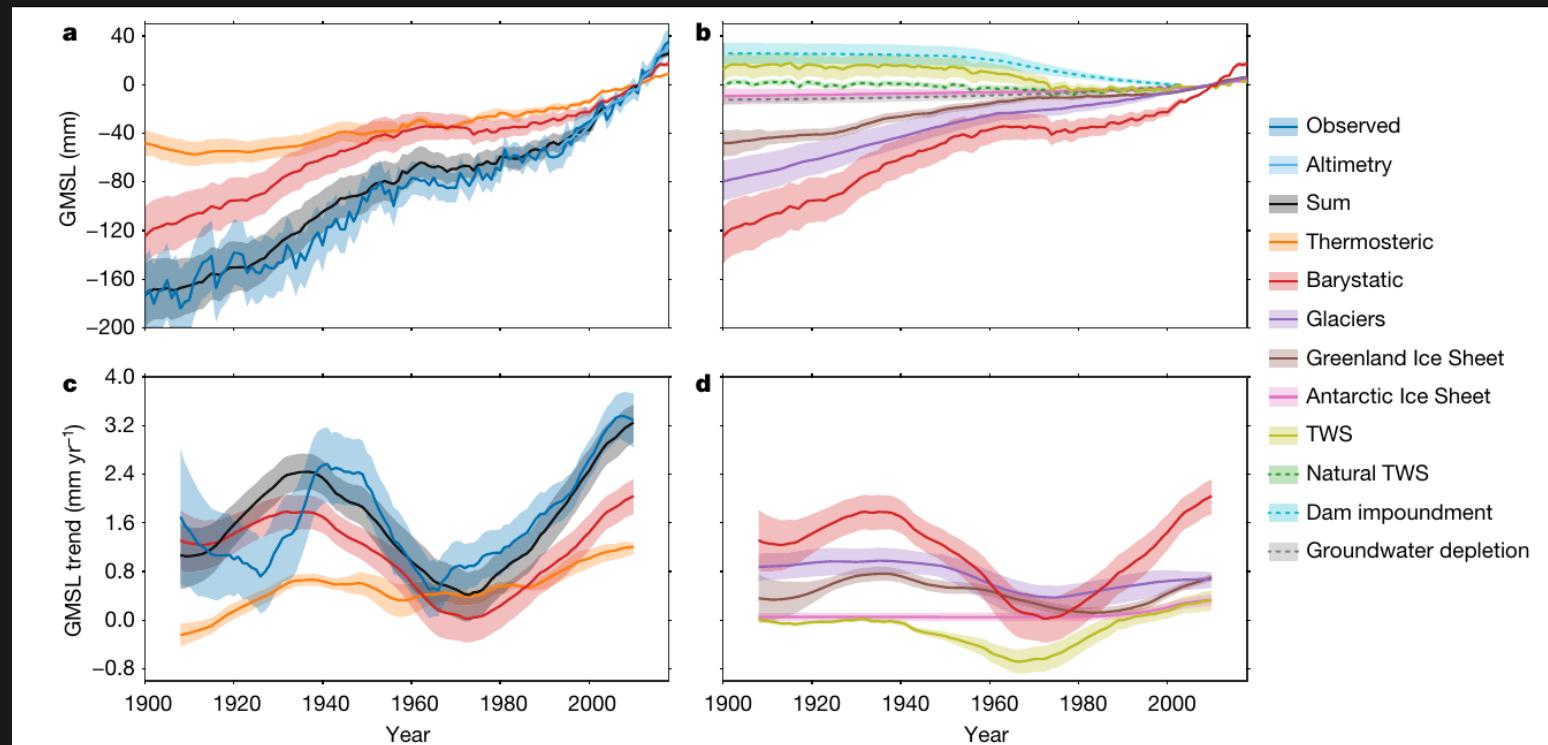


- Historic sea level curve can be inferred from tide gauges

Historic sea level can be largely explained with different (modelled) sea level contributions. From Frederikse et al. 2020



Sea level contributions since 1900

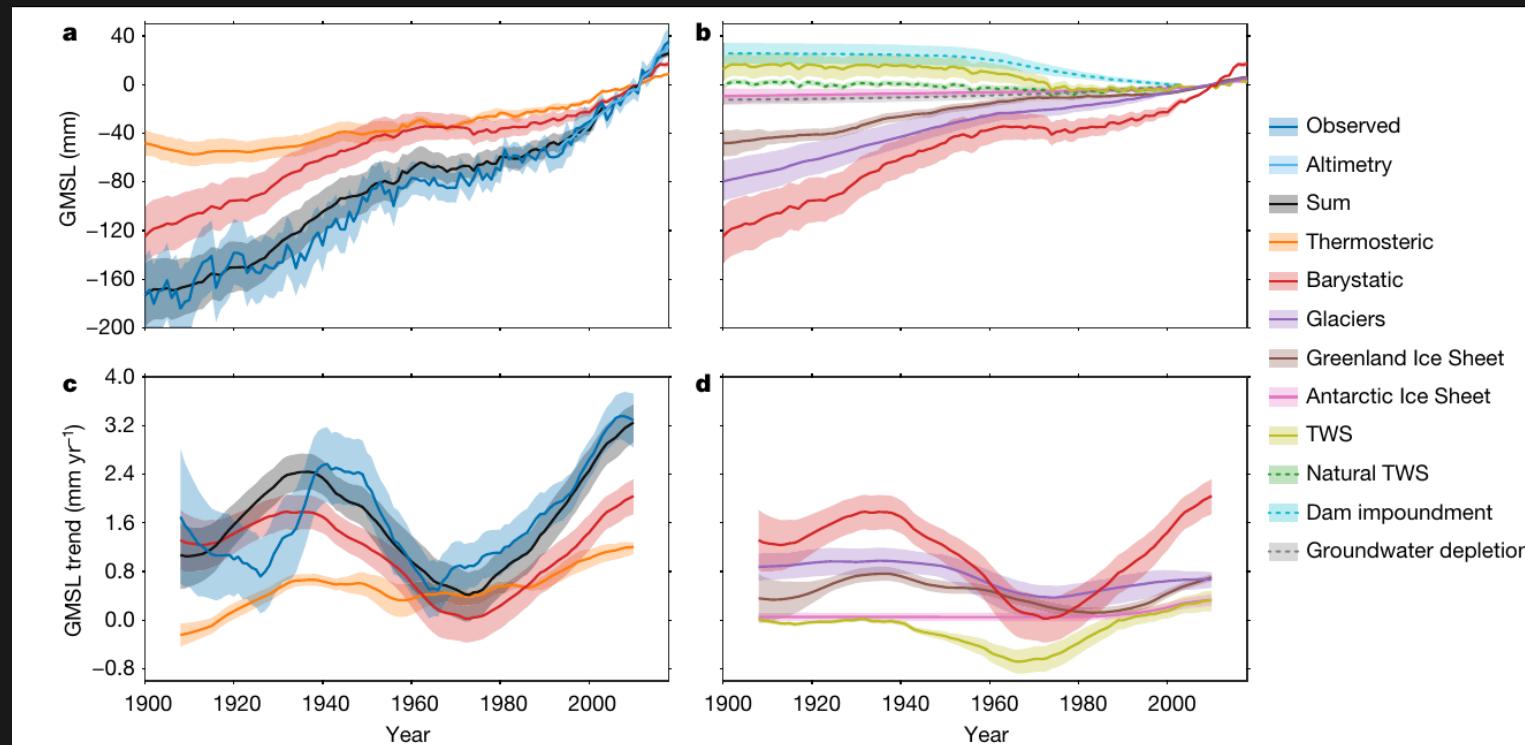


- Historic sea level curve can be inferred from tide gauges
- Uncertainty is larger for early period (sparse obs.)

Historic sea level can be largely explained with different (modelled) sea level contributions. From Frederikse et al. 2020



Sea level contributions since 1900

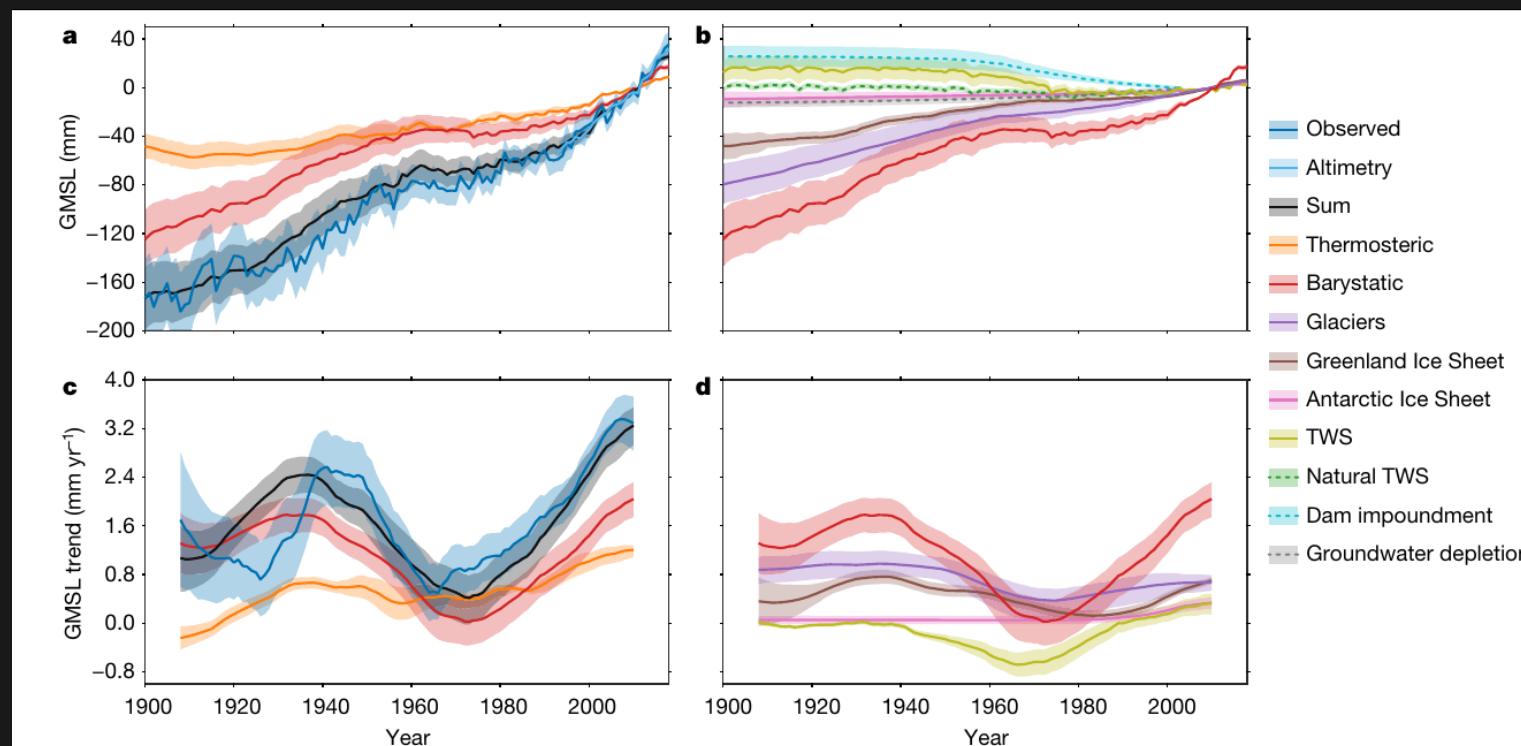


- Historic sea level curve can be inferred from tide gauges
- Uncertainty is larger for early period (sparse obs.)
- Strong interannual fluctuations in mass (barystatic)

Historic sea level can be largely explained with different (modelled) sea level contributions. From Frederikse et al. 2020



Sea level contributions since 1900

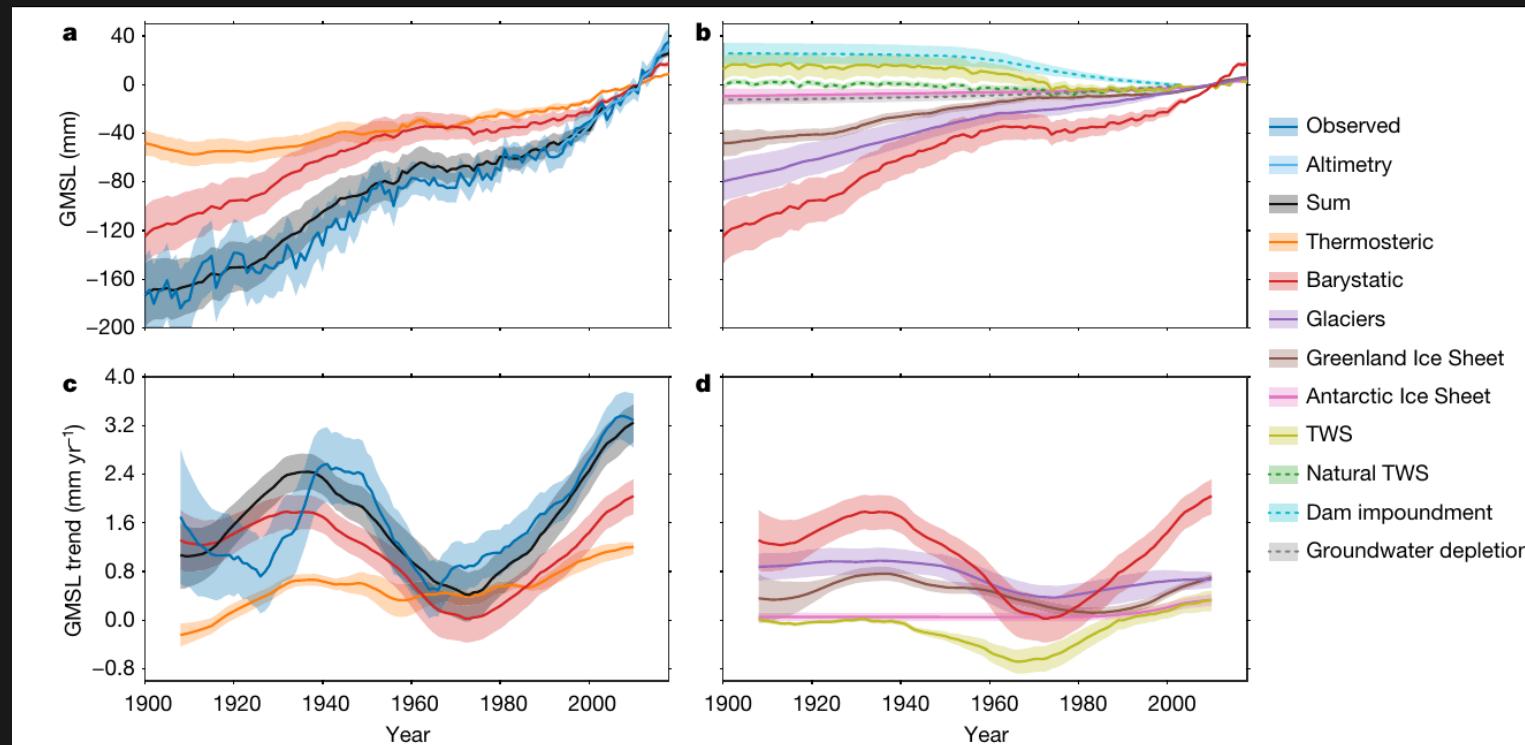


- Historic sea level curve can be inferred from tide gauges
- Uncertainty is larger for early period (sparse obs.)
- Strong interannual fluctuations in mass (barystatic)
- Altimetry only since ~1990 (~3.3 mm/yr)

Historic sea level can be largely explained with different (modelled) sea level contributions. From Frederikse et al. 2020



Sea level contributions since 1900

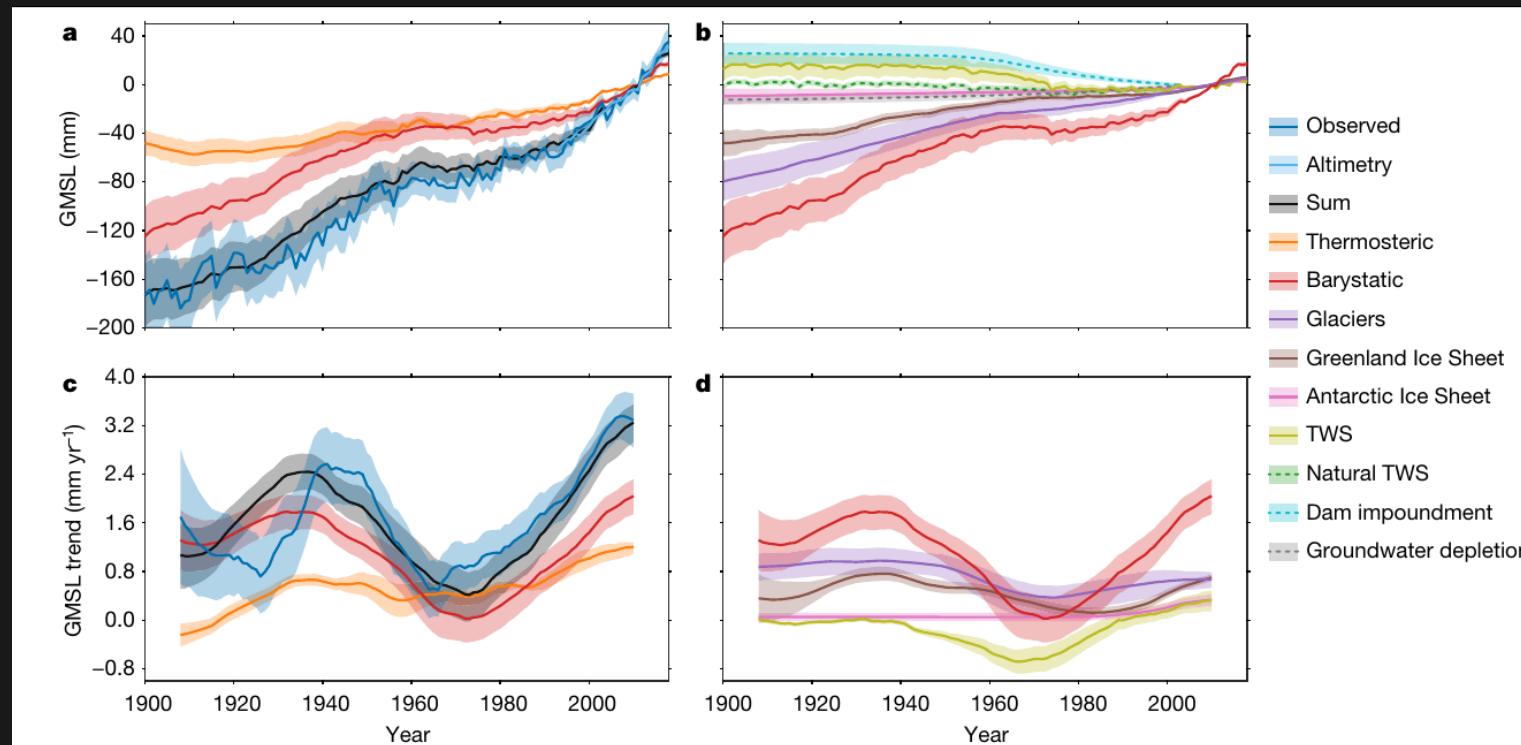


Historic sea level can be largely explained with different (modelled) sea level contributions. From Frederikse et al. 2020

- Historic sea level curve can be inferred from tide gauges
- Uncertainty is larger for early period (sparse obs.)
- Strong interannual fluctuations in mass (barystatic)
- Altimetry only since ~1990 (~3.3 mm/yr)
- Budget can be closed within uncertainties



Sea level contributions since 1900



Historic sea level can be largely explained with different (modelled) sea level contributions. From Frederikse et al. 2020

- Historic sea level curve can be inferred from tide gauges
- Uncertainty is larger for early period (sparse obs.)
- Strong interannual fluctuations in mass (barystatic)
- Altimetry only since ~1990 (~3.3 mm/yr)
- Budget can be closed within uncertainties
- Recent accelerations from icesheets and glaciers -> (observable by GRACE since 2002)



Ocean mass from GRACE (Theory)

JOURNAL OF GEOPHYSICAL RESEARCH, VOL. 107, NO. B9, 2193, doi:10.1029/2001JB000576, 2002

Methods for inferring regional surface-mass anomalies from Gravity Recovery and Climate Experiment (GRACE) measurements of time-variable gravity

Sean Swenson and John Wahr

Department of Physics and Cooperative Institute for Research in Environmental Sciences, University of Colorado, Boulder, USA

Received 8 May 2001; revised 9 March 2002; accepted 14 March 2002; published 19 September 2002.

the approximate basin average can be expressed in terms of Stokes coefficients as

$$\widetilde{\Delta\sigma}_{\text{region}} = \sum_{l,m} \frac{K_l}{\Omega_{\text{region}}} (W_{lm}^c \Delta C_{lm} + W_{lm}^s \Delta S_{lm}). \quad (27)$$



Ocean mass from GRACE (Theory)

JOURNAL OF GEOPHYSICAL RESEARCH, VOL. 107, NO. B9, 2193, doi:10.1029/2001JB000576, 2002

Methods for inferring regional surface-mass anomalies from Gravity Recovery and Climate Experiment (GRACE) measurements of time-variable gravity

Sean Swenson and John Wahr

Department of Physics and Cooperative Institute for Research in Environmental Sciences, University of Colorado, Boulder, USA

Received 8 May 2001; revised 9 March 2002; accepted 14 March 2002; published 19 September 2002.

the approximate basin average can be expressed in terms of Stokes coefficients as

$$\widetilde{\Delta\sigma}_{\text{region}} = \sum_{l,m} \frac{K_l}{\Omega_{\text{region}}} (W_{lm}^c \Delta C_{lm} + W_{lm}^s \Delta S_{lm}). \quad (27)$$

- Approach 1: Average over glacier and icesheet regions and sum up



Ocean mass from GRACE (Theory)

JOURNAL OF GEOPHYSICAL RESEARCH, VOL. 107, NO. B9, 2193, doi:10.1029/2001JB000576, 2002

Methods for inferring regional surface-mass anomalies from Gravity Recovery and Climate Experiment (GRACE) measurements of time-variable gravity

Sean Swenson and John Wahr

Department of Physics and Cooperative Institute for Research in Environmental Sciences, University of Colorado, Boulder, USA

Received 8 May 2001; revised 9 March 2002; accepted 14 March 2002; published 19 September 2002.

the approximate basin average can be expressed in terms of Stokes coefficients as

$$\widetilde{\Delta\sigma}_{\text{region}} = \sum_{l,m} \frac{K_l}{\Omega_{\text{region}}} (W_{lm}^c \Delta C_{lm} + W_{lm}^s \Delta S_{lm}). \quad (27)$$

- Approach 1: Average over glacier and icesheet regions and sum up
- Or approach 2: directly average over the ocean



Ocean mass from GRACE (Theory)

JOURNAL OF GEOPHYSICAL RESEARCH, VOL. 107, NO. B9, 2193, doi:10.1029/2001JB000576, 2002

Methods for inferring regional surface-mass anomalies from Gravity Recovery and Climate Experiment (GRACE) measurements of time-variable gravity

Sean Swenson and John Wahr

Department of Physics and Cooperative Institute for Research in Environmental Sciences, University of Colorado, Boulder, USA

Received 8 May 2001; revised 9 March 2002; accepted 14 March 2002; published 19 September 2002.

the approximate basin average can be expressed in terms of Stokes coefficients as

$$\widetilde{\Delta\sigma}_{\text{region}} = \sum_{l,m} \frac{K_l}{\Omega_{\text{region}}} (W_{lm}^c \Delta C_{lm} + W_{lm}^s \Delta S_{lm}). \quad (27)$$

- Approach 1: Average over glacier and icesheet regions and sum up
- Or approach 2: directly average over the ocean
- Other approaches (e.g. inverse problems)



Ocean mass from GRACE (Theory)

JOURNAL OF GEOPHYSICAL RESEARCH, VOL. 107, NO. B9, 2193, doi:10.1029/2001JB000576, 2002

Methods for inferring regional surface-mass anomalies from Gravity Recovery and Climate Experiment (GRACE) measurements of time-variable gravity

Sean Swenson and John Wahr

Department of Physics and Cooperative Institute for Research in Environmental Sciences, University of Colorado, Boulder, USA

Received 8 May 2001; revised 9 March 2002; accepted 14 March 2002; published 19 September 2002.

the approximate basin average can be expressed in terms of Stokes coefficients as

$$\widetilde{\Delta\sigma}_{\text{region}} = \sum_{l,m} \frac{K_l}{\Omega_{\text{region}}} (W_{lm}^c \Delta C_{lm} + W_{lm}^s \Delta S_{lm}). \quad (27)$$

- Approach 1: Average over glacier and icesheet regions and sum up
- Or approach 2: directly average over the ocean
- Other approaches (e.g. inverse problems)
- Averaging can be done purely in the spectral domain



Ocean mass from GRACE (Theory)

JOURNAL OF GEOPHYSICAL RESEARCH, VOL. 107, NO. B9, 2193, doi:10.1029/2001JB000576, 2002

Methods for inferring regional surface-mass anomalies from Gravity Recovery and Climate Experiment (GRACE) measurements of time-variable gravity

Sean Swenson and John Wahr

Department of Physics and Cooperative Institute for Research in Environmental Sciences, University of Colorado, Boulder, USA

Received 8 May 2001; revised 9 March 2002; accepted 14 March 2002; published 19 September 2002.

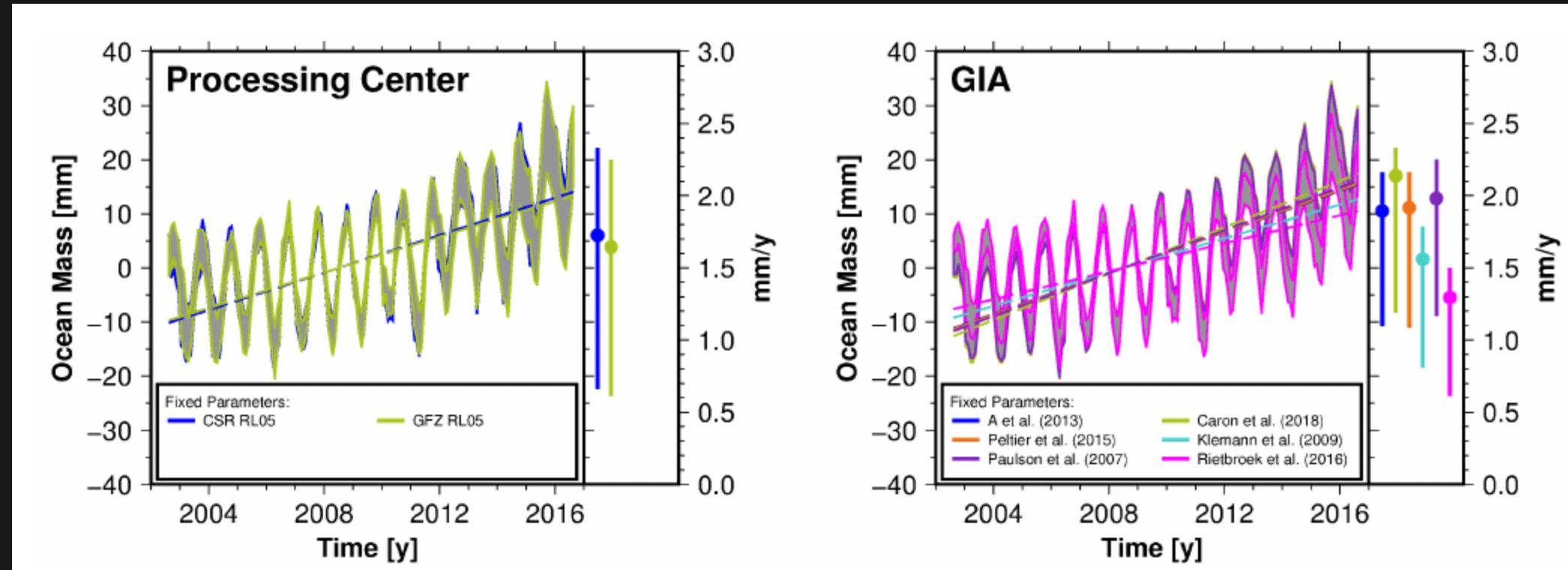
the approximate basin average can be expressed in terms of Stokes coefficients as

$$\widetilde{\Delta\sigma}_{\text{region}} = \sum_{l,m} \frac{K_l}{\Omega_{\text{region}}} (W_{lm}^c \Delta C_{lm} + W_{lm}^s \Delta S_{lm}). \quad (27)$$

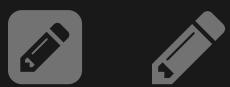
- Approach 1: Average over glacier and icesheet regions and sum up
- Or approach 2: directly average over the ocean
- Other approaches (e.g. inverse problems)
- Averaging can be done purely in the spectral domain
 - dot product of smoothed basin coefficients with GRACE-TWS coefficients



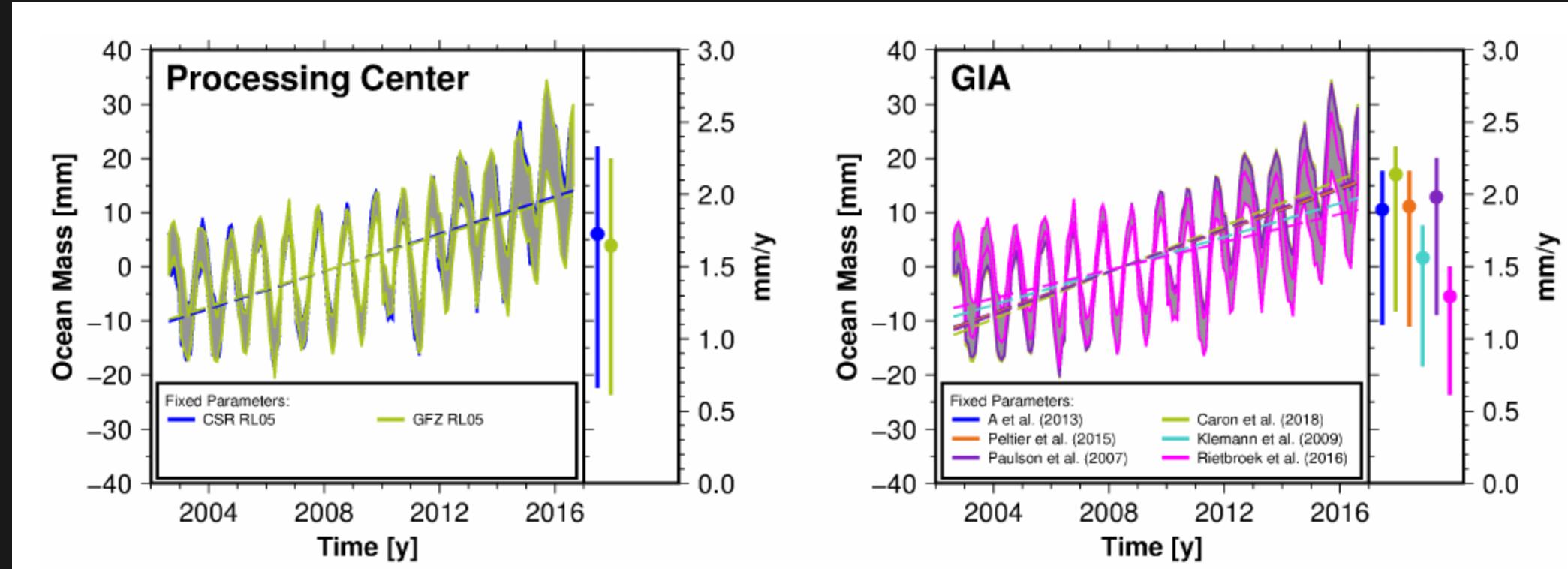
Ocean mass from GRACE (Praxis)



Effect of different processing choices on ocean mass estimates, from Uebbing et al. 2019

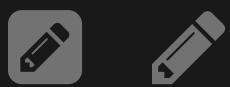


Ocean mass from GRACE (Praxis)

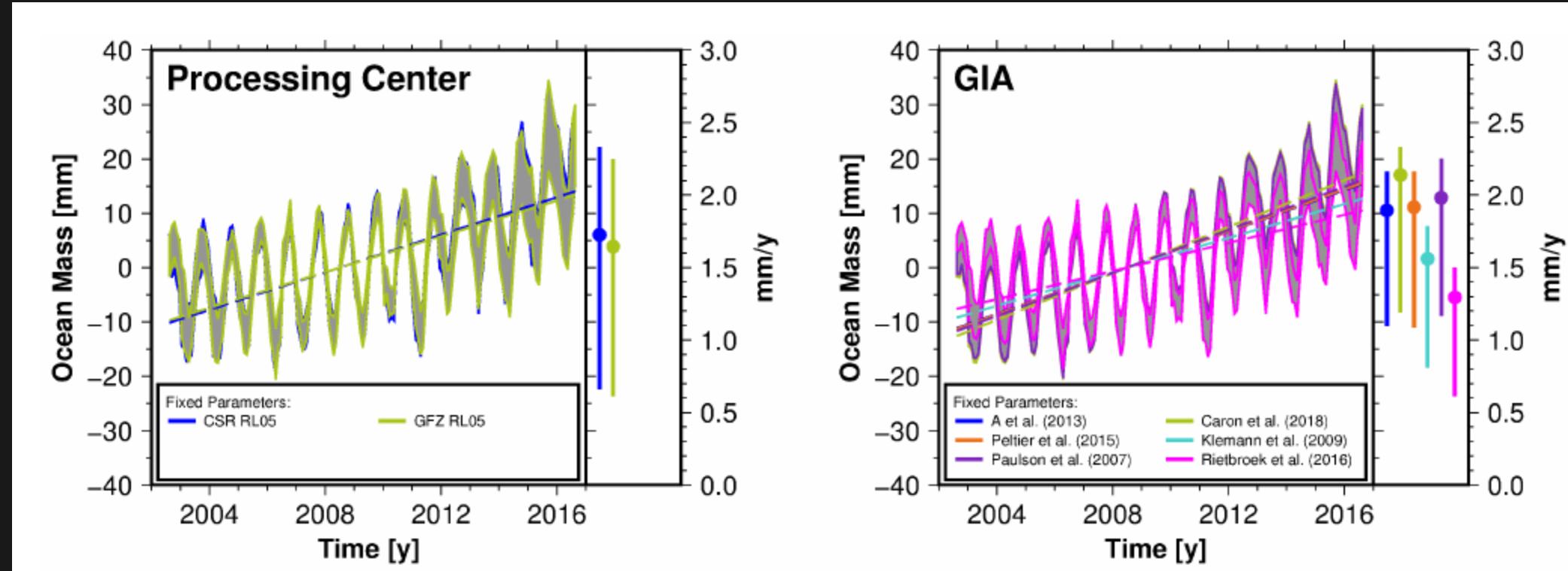


Effect of different processing choices on ocean mass estimates, from Uebbing et al. 2019

- Processing choices introduce uncertainties in trends

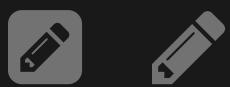


Ocean mass from GRACE (Praxis)

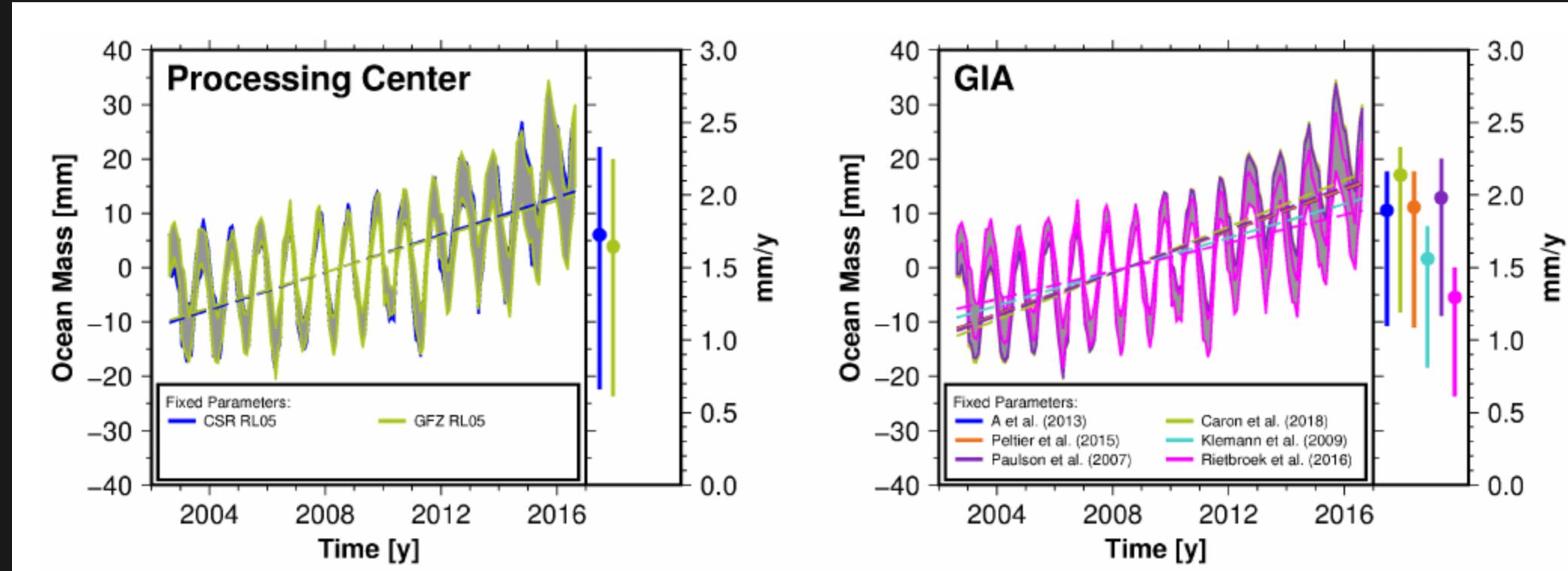


Effect of different processing choices on ocean mass estimates, from Uebbing et al. 2019

- Processing choices introduce uncertainties in trends
 - Correction for glacial isostatic adjustment

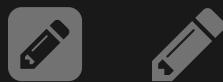


Ocean mass from GRACE (Praxis)

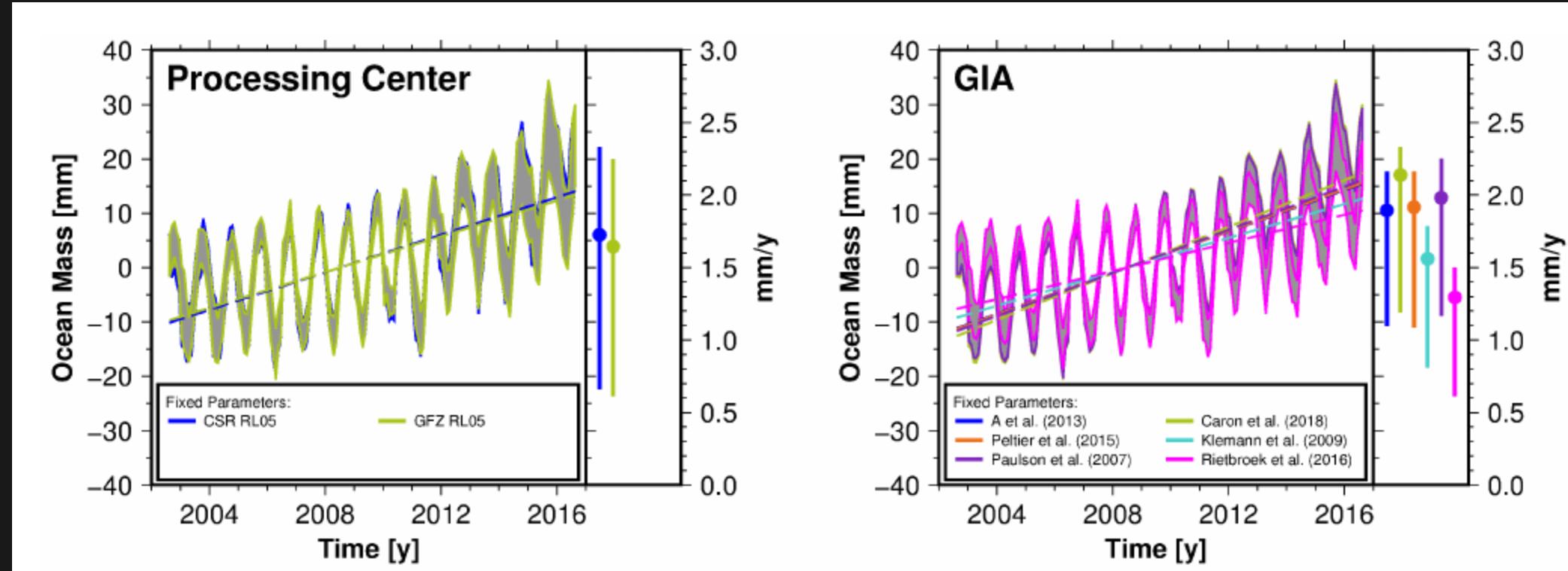


Effect of different processing choices on ocean mass estimates, from Uebbing et al. 2019

- Processing choices introduce uncertainties in trends
 - Correction for glacial isostatic adjustment
 - restoring background models (GAC, GAD)

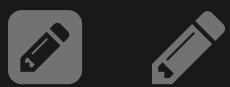


Ocean mass from GRACE (Praxis)

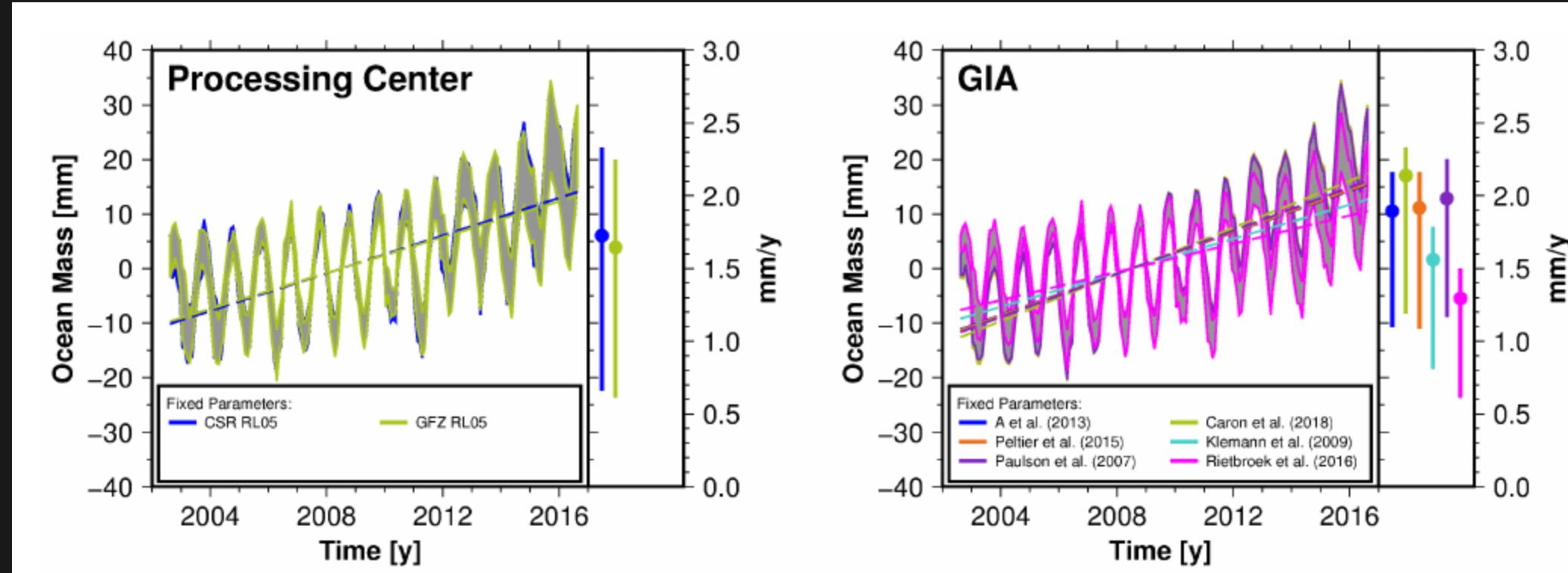


Effect of different processing choices on ocean mass estimates, from Uebbing et al. 2019

- Processing choices introduce uncertainties in trends
 - Correction for glacial isostatic adjustment
 - restoring background models (GAC, GAD)
 - but also averaging region

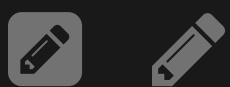


Ocean mass from GRACE (Praxis)

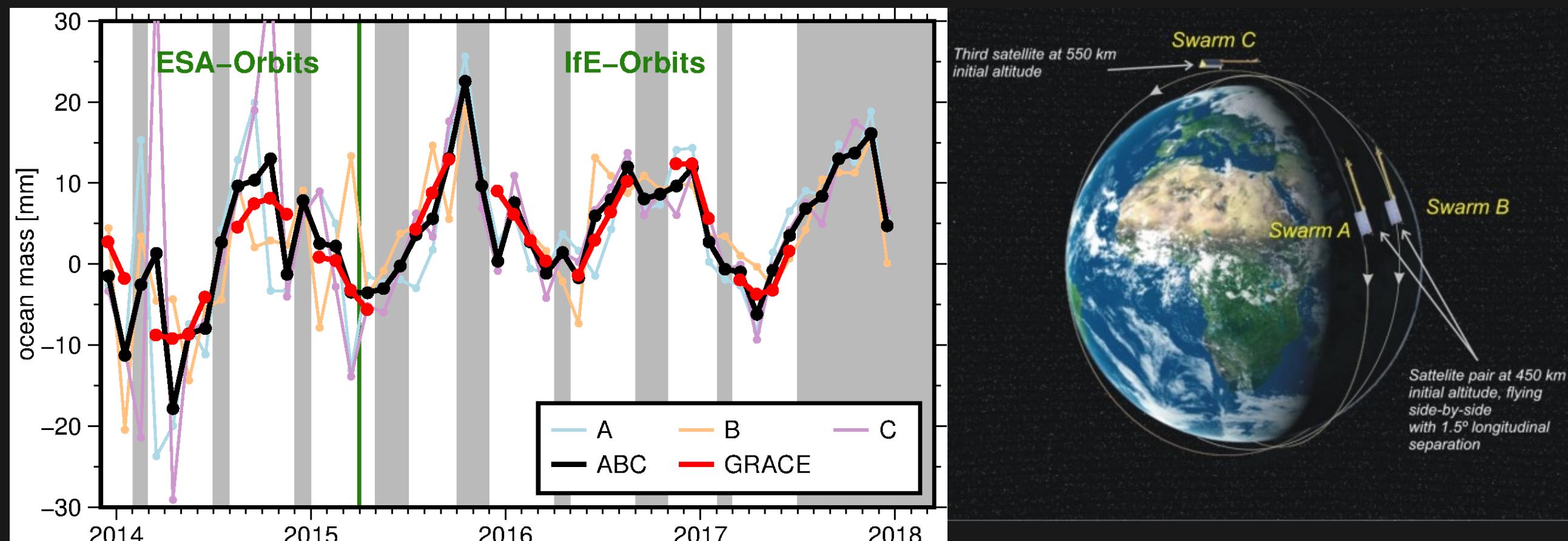


Effect of different processing choices on ocean mass estimates, from Uebbing et al. 2019

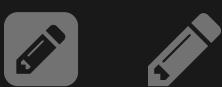
- Processing choices introduce uncertainties in trends
 - Correction for glacial isostatic adjustment
 - restoring background models (GAC, GAD)
 - but also averaging region
 - Current mass trends \sim 1.5-2mm/yr (c.f. total trend \sim 3.3 mm/yr)



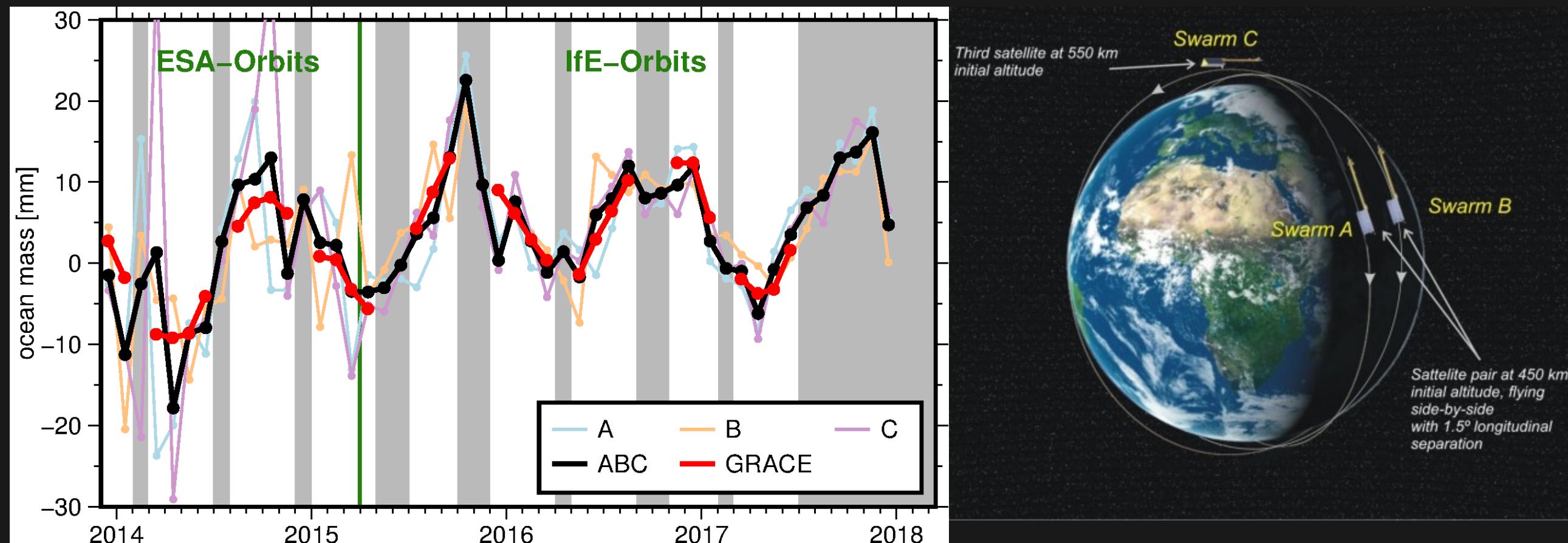
Ocean mass from the SWARM Mission



Comparison of ocean mass from GRACE and the SWARM (A,B,C satellites). Update of Luck et al. 2018

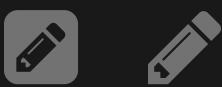


Ocean mass from the SWARM Mission

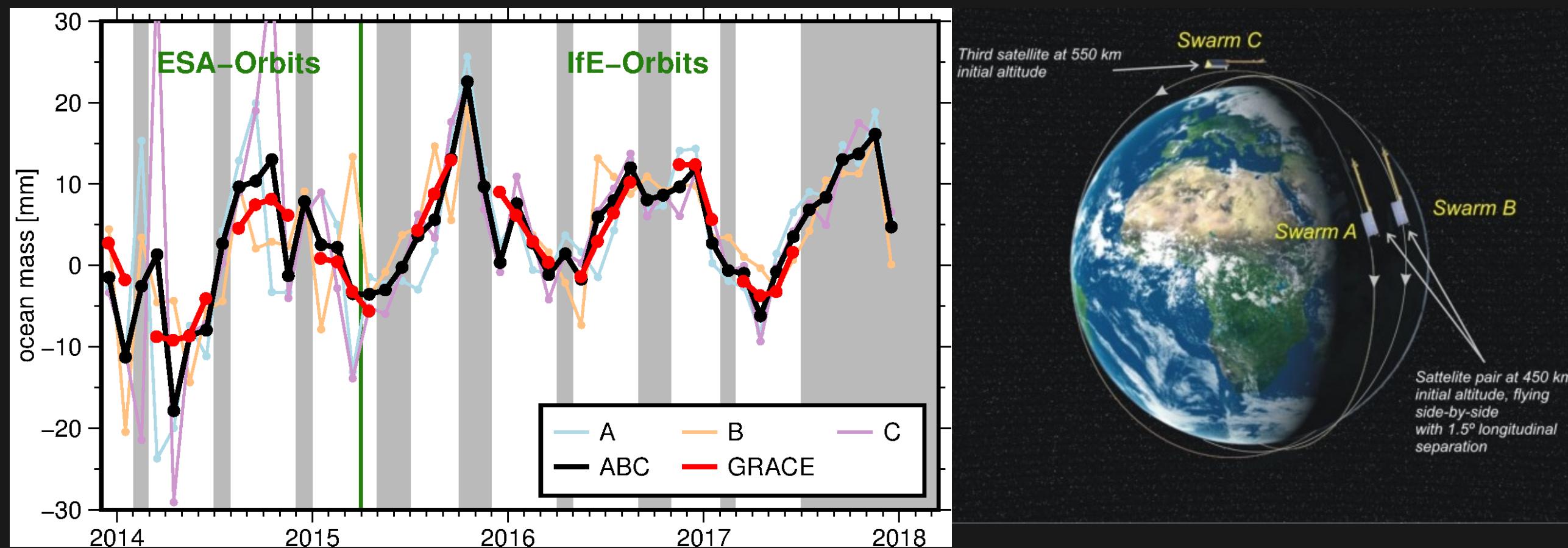


Comparison of ocean mass from GRACE and the SWARM (A,B,C satellites). Update of Luck et al. 2018

- SWARM has no inter-satellite tracking, but precise orbits (sat. A,B and C)

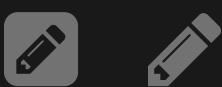


Ocean mass from the SWARM Mission

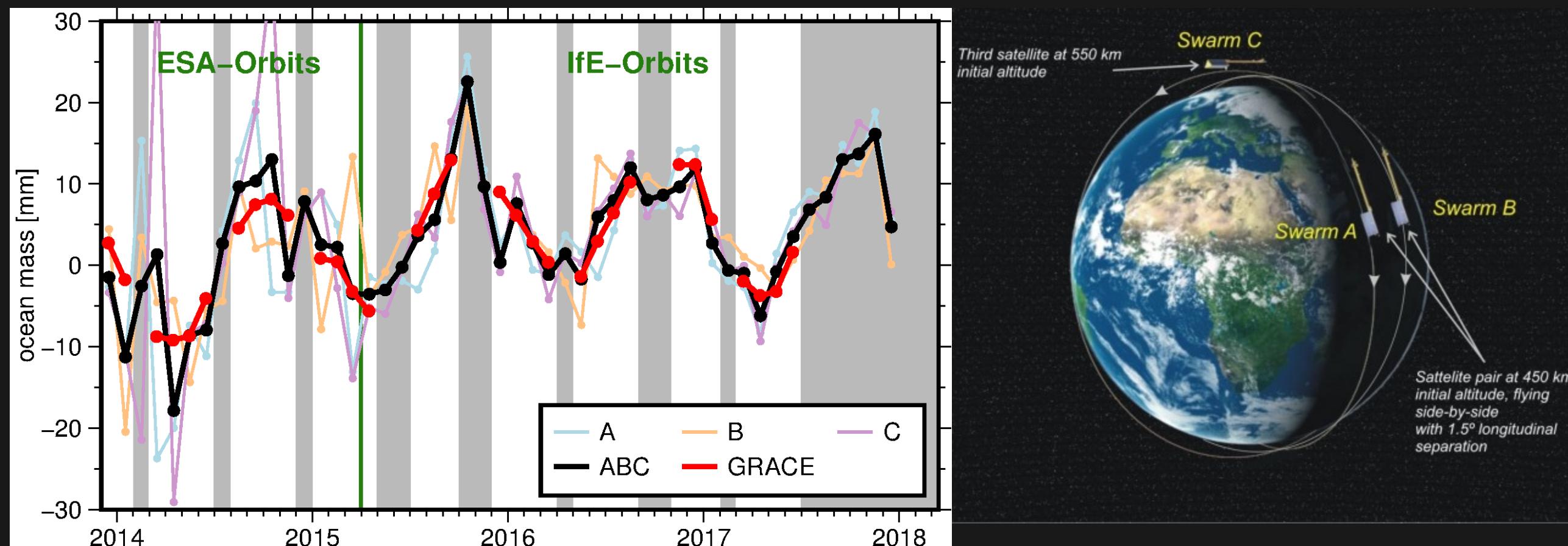


Comparison of ocean mass from GRACE and the SWARM (A,B,C satellites). Update of Luck et al. 2018

- SWARM has no inter-satellite tracking, but precise orbits (sat. A,B and C)
- Precise (kinematic) orbits contain information on low-resolution gravity changes -> ocean scales

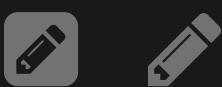


Ocean mass from the SWARM Mission

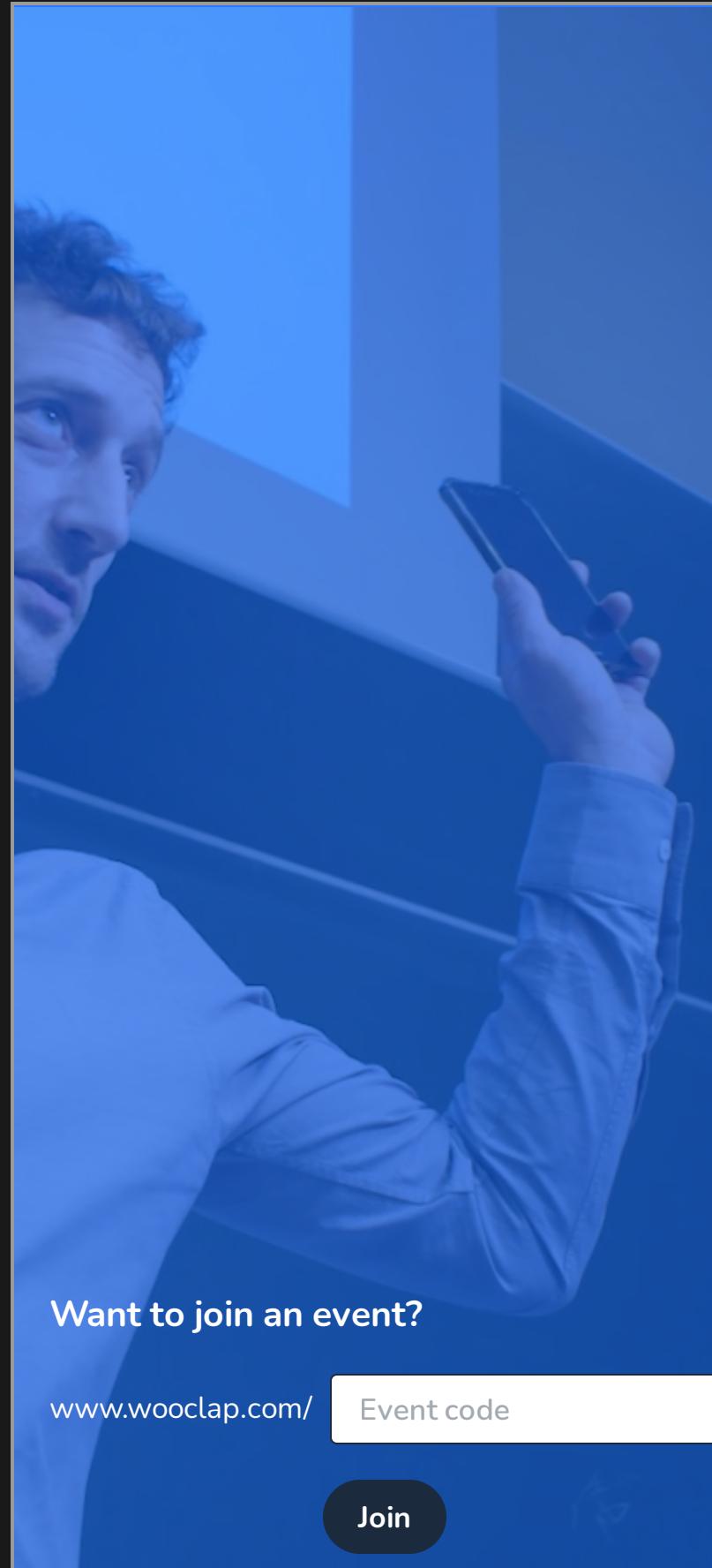


Comparison of ocean mass from GRACE and the SWARM (A,B,C satellites). Update of Luck et al. 2018

- SWARM has no inter-satellite tracking, but precise orbits (sat. A,B and C)
- Precise (kinematic) orbits contain information on low-resolution gravity changes -> ocean scales
- Principle is useful for filling gaps and potentially pre-GRACE era



Does Greenland induce regional sea level changes?



Want to join an event?

www.wooclap.com/

Event code

Join



[Sign up](#)

Choose a method to log in

[or sign up](#)

Your email address

OR

Sign up with Facebook

Sign up with Google

Sign up with LinkedIn

Sign up with Microsoft

OR

Log in with your institution

Your institution



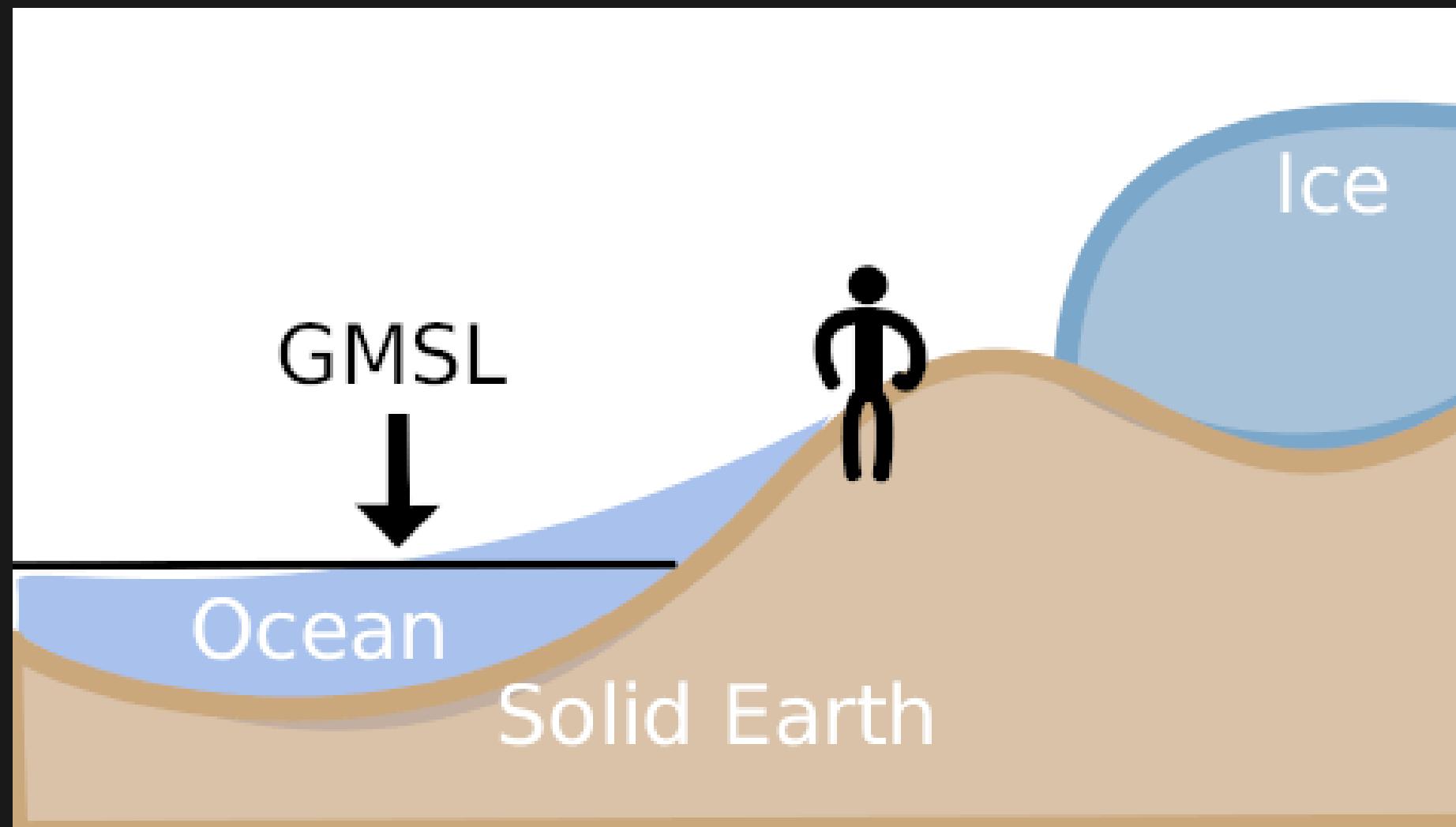
We use [cookies](#) to improve the general experience on the platform, provide users chat support and deliver targeted ads on other websites.

[Accept all](#)

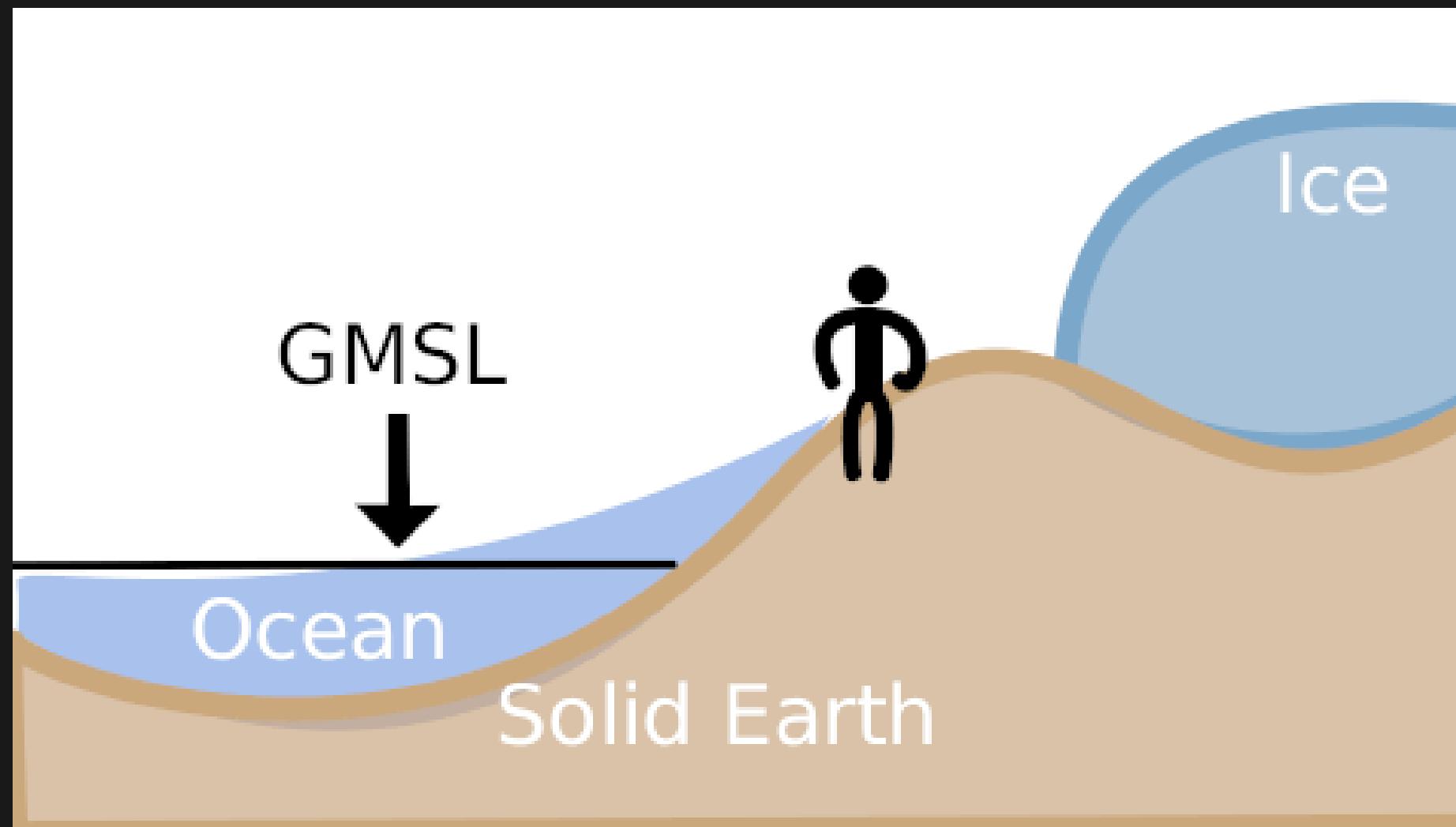
[Reject All](#)

[Customise](#)

Relative sea level is affected by self attraction and loading (SAL)..

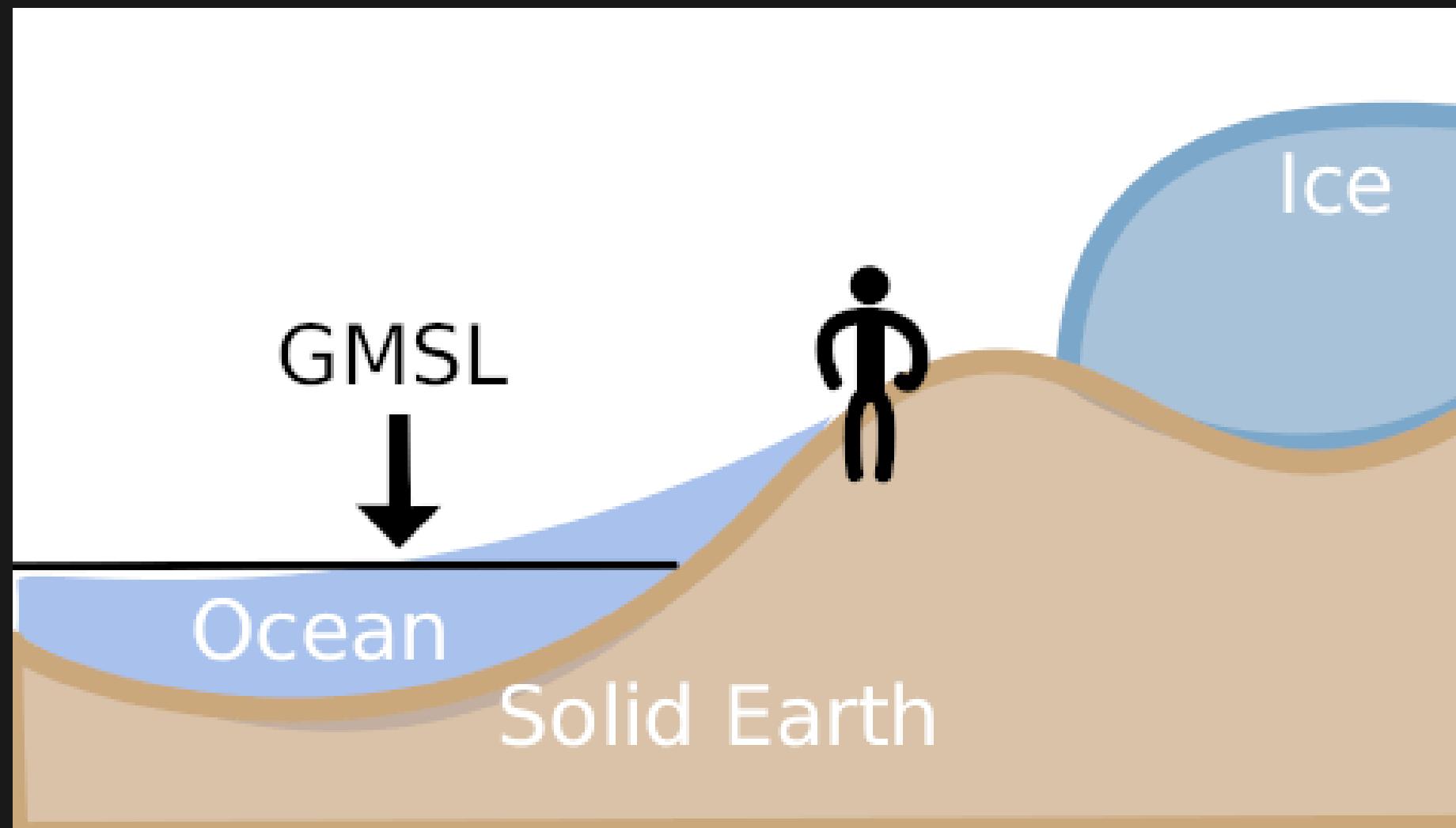


Relative sea level is affected by self attraction and loading (SAL)..



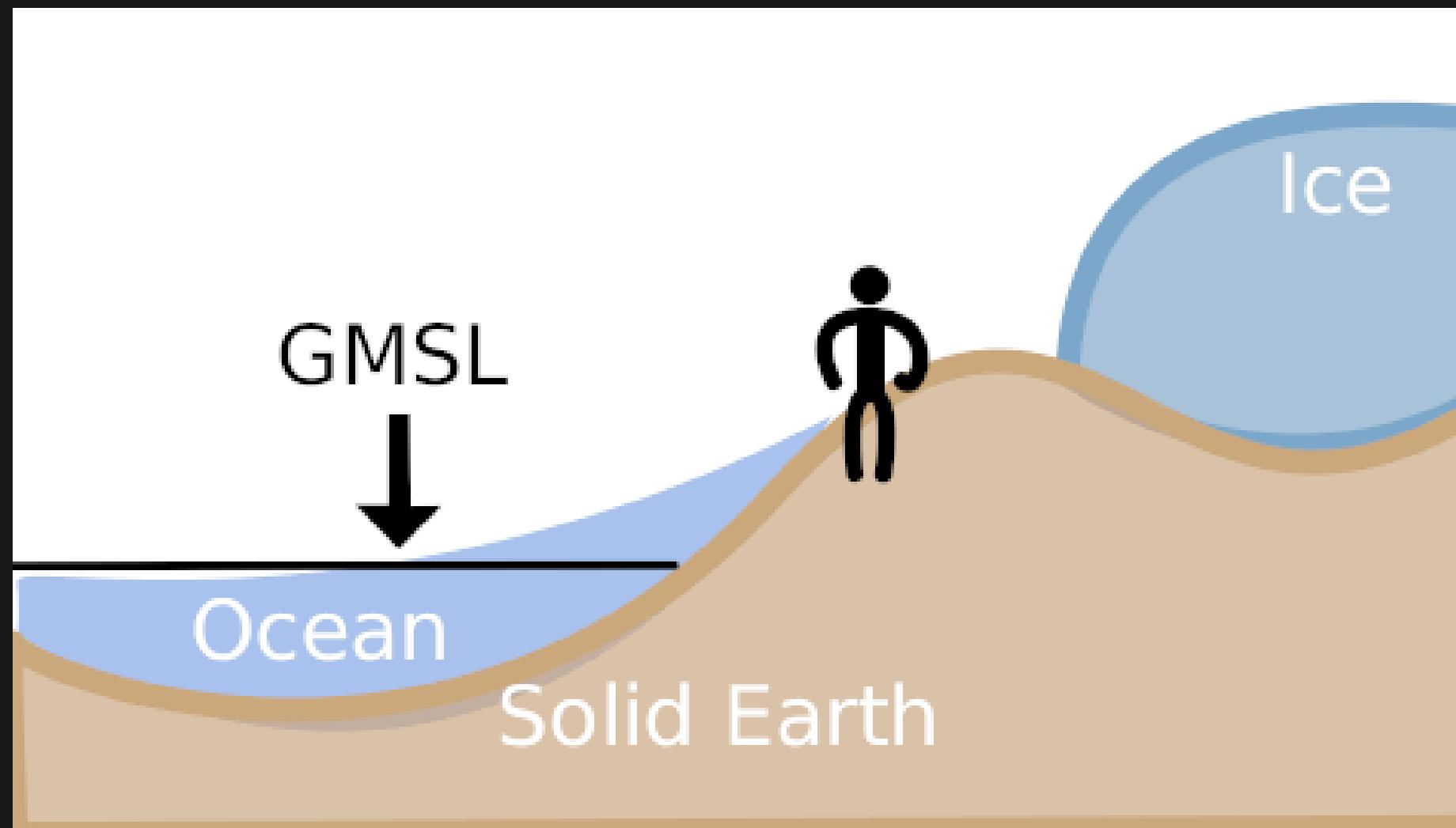
- Theory goes back to 1888
(Woodward)

Relative sea level is affected by self attraction and loading (SAL)..



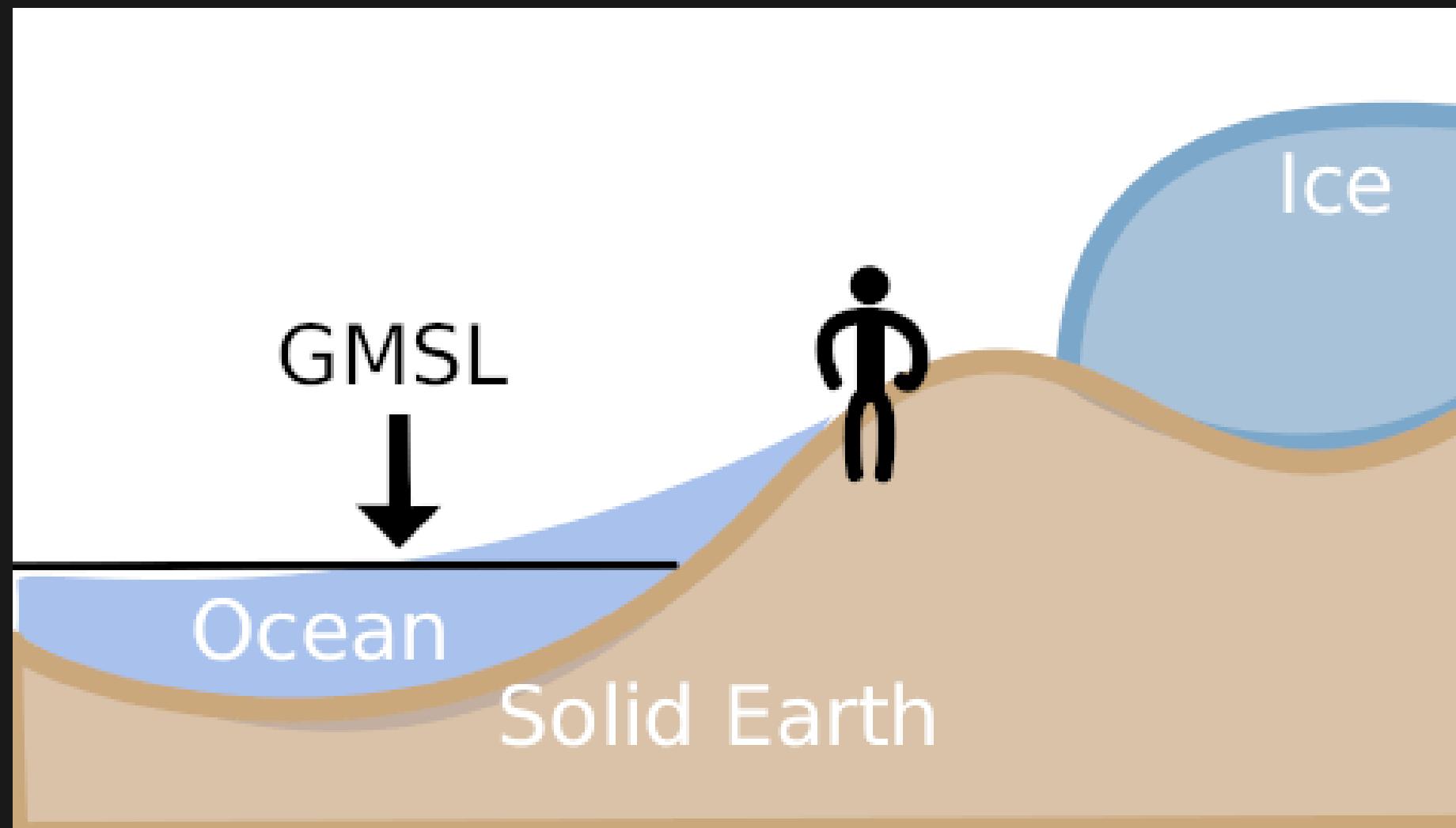
- Theory goes back to 1888 (Woodward)
- Modelled as a passive ocean response (no currents)

Relative sea level is affected by self attraction and loading (SAL)..



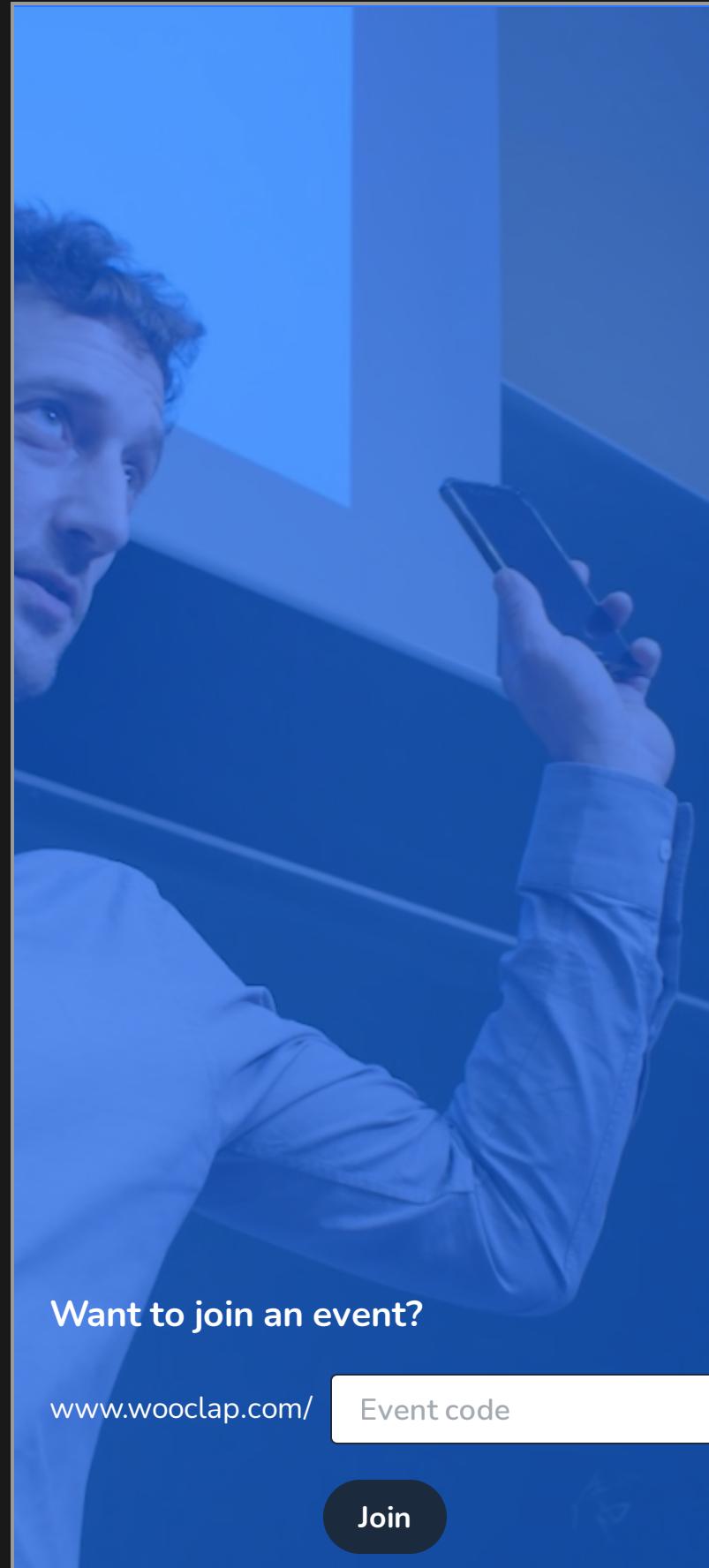
- Theory goes back to 1888 (Woodward)
- Modelled as a passive ocean response (no currents)
- mass conserving

Relative sea level is affected by self attraction and loading (SAL)..



- Theory goes back to 1888 (Woodward)
- Modelled as a passive ocean response (no currents)
- mass conserving
- Takes into account the deformation of the Earth

Other factors affecting relative sea level?



Want to join an event?

www.wooclap.com/

Event code

Join



[Sign up](#)

Choose a method to log in

[or sign up](#)

Your email address

OR

[Sign up with Facebook](#)

[Sign up with Google](#)

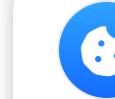
[Sign up with LinkedIn](#)

[Sign up with Microsoft](#)

OR

Log in with your institution

Your institution



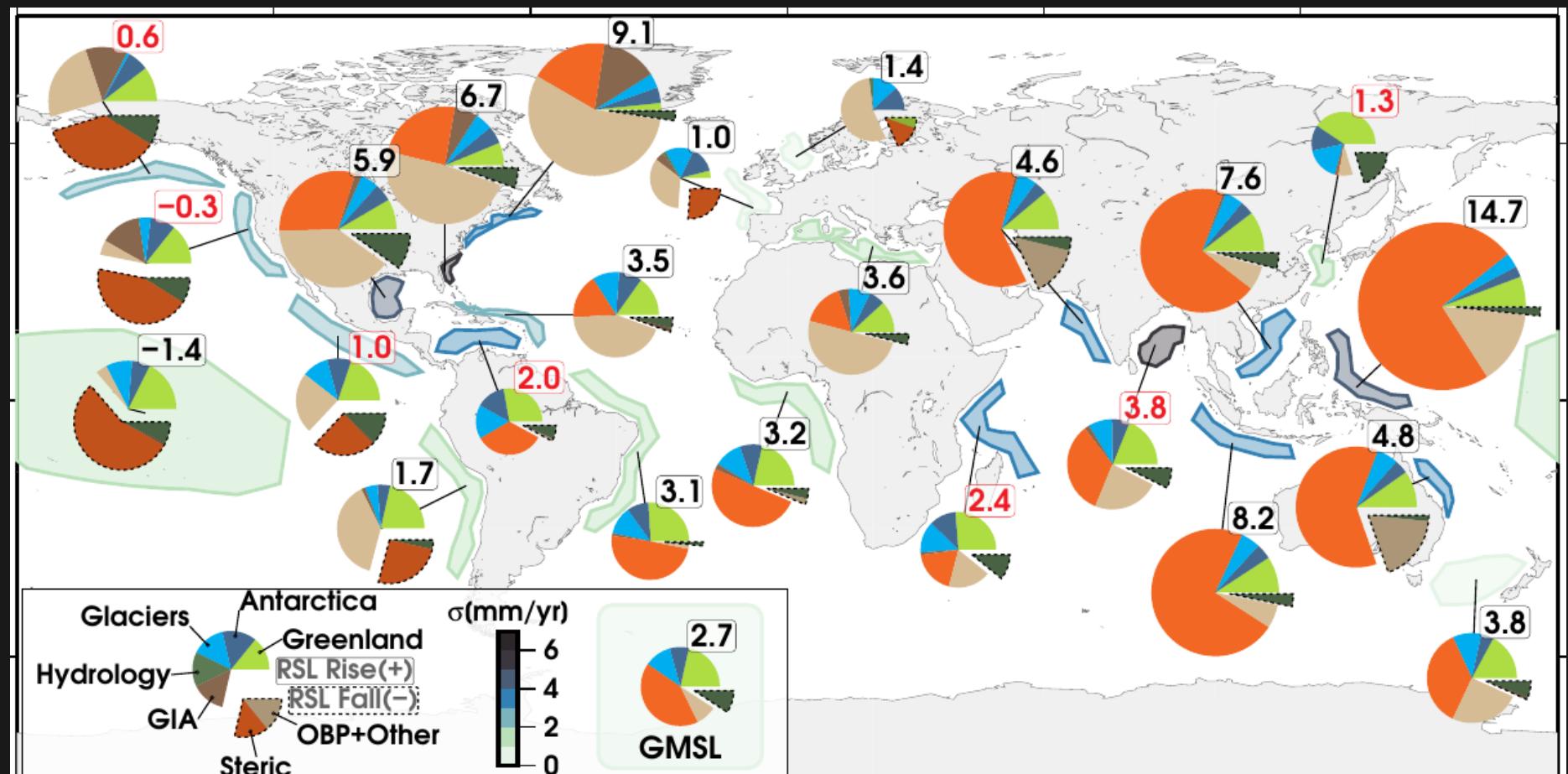
We use [cookies](#) to improve the general experience on the platform, provide users chat support and deliver targeted ads on our websites.

[Accept all](#)

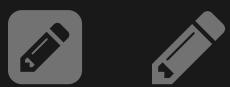
[Reject All](#)

[Customise](#)

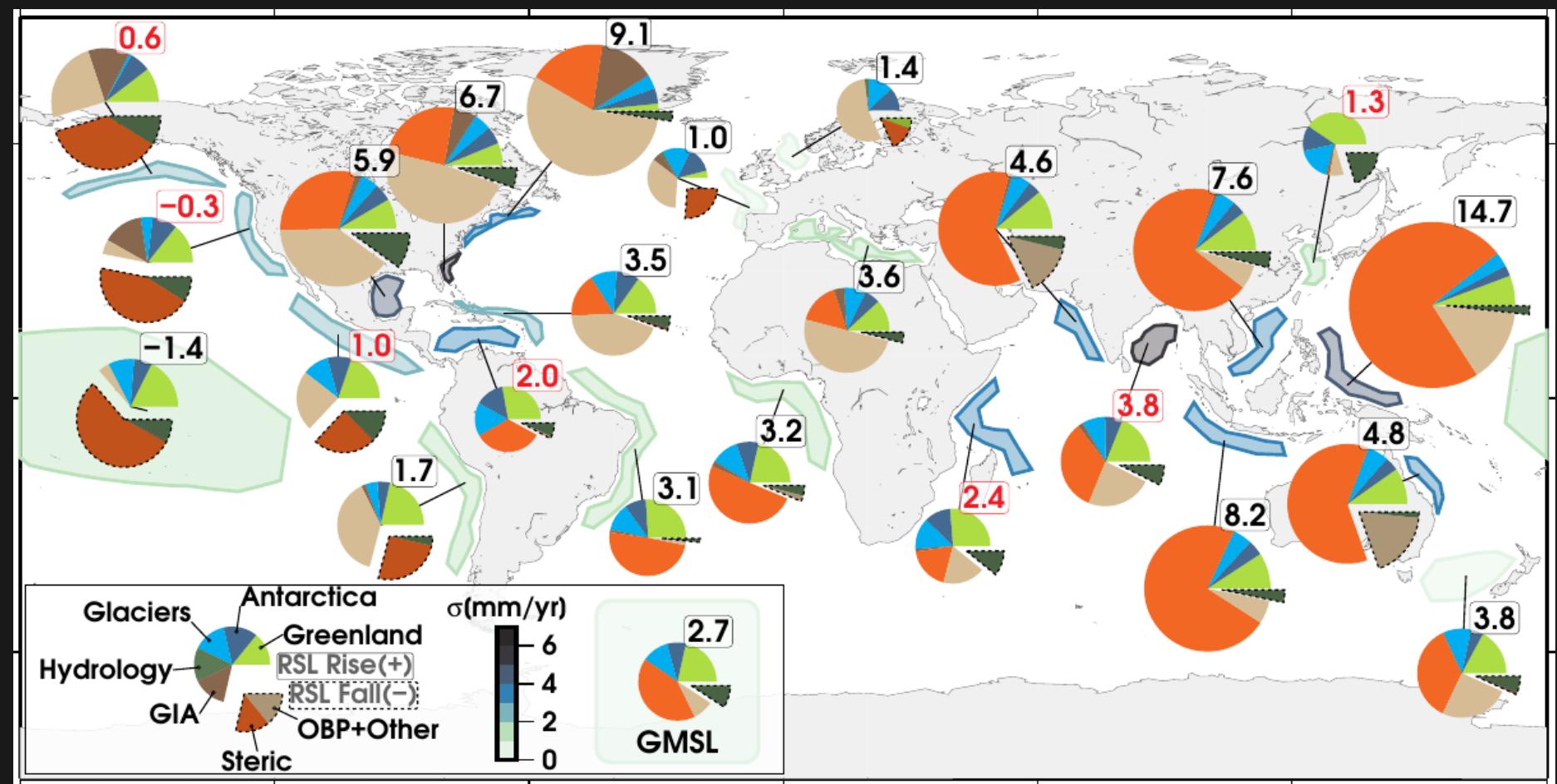
Regional sea level budgets



Regional sea level budgets from Rietbroek et al. 2016

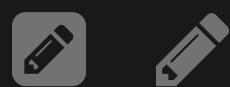


Regional sea level budgets

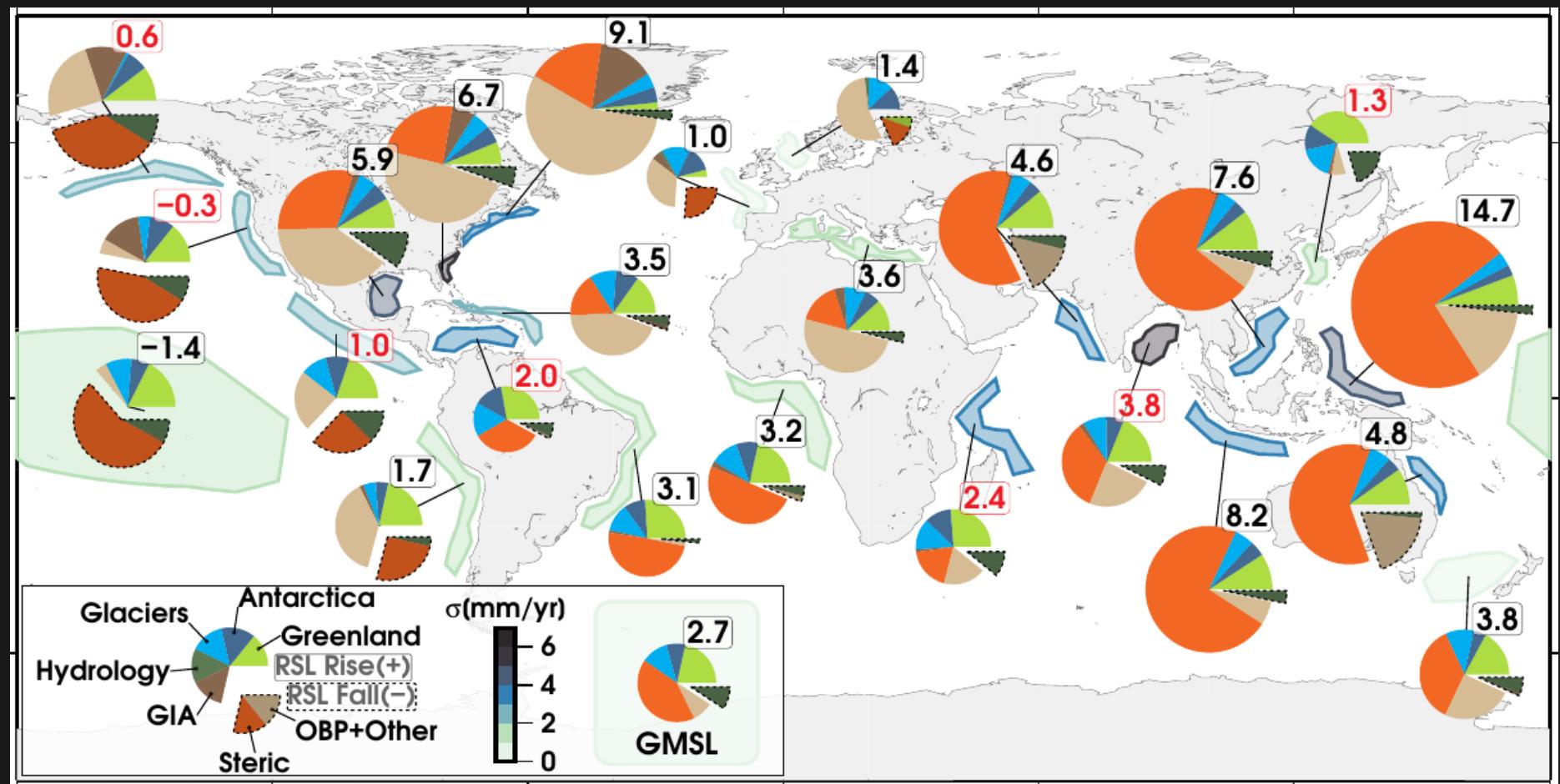


Regional sea level budgets from Rietbroek et al. 2016

- Sea level rise is not uniform



Regional sea level budgets

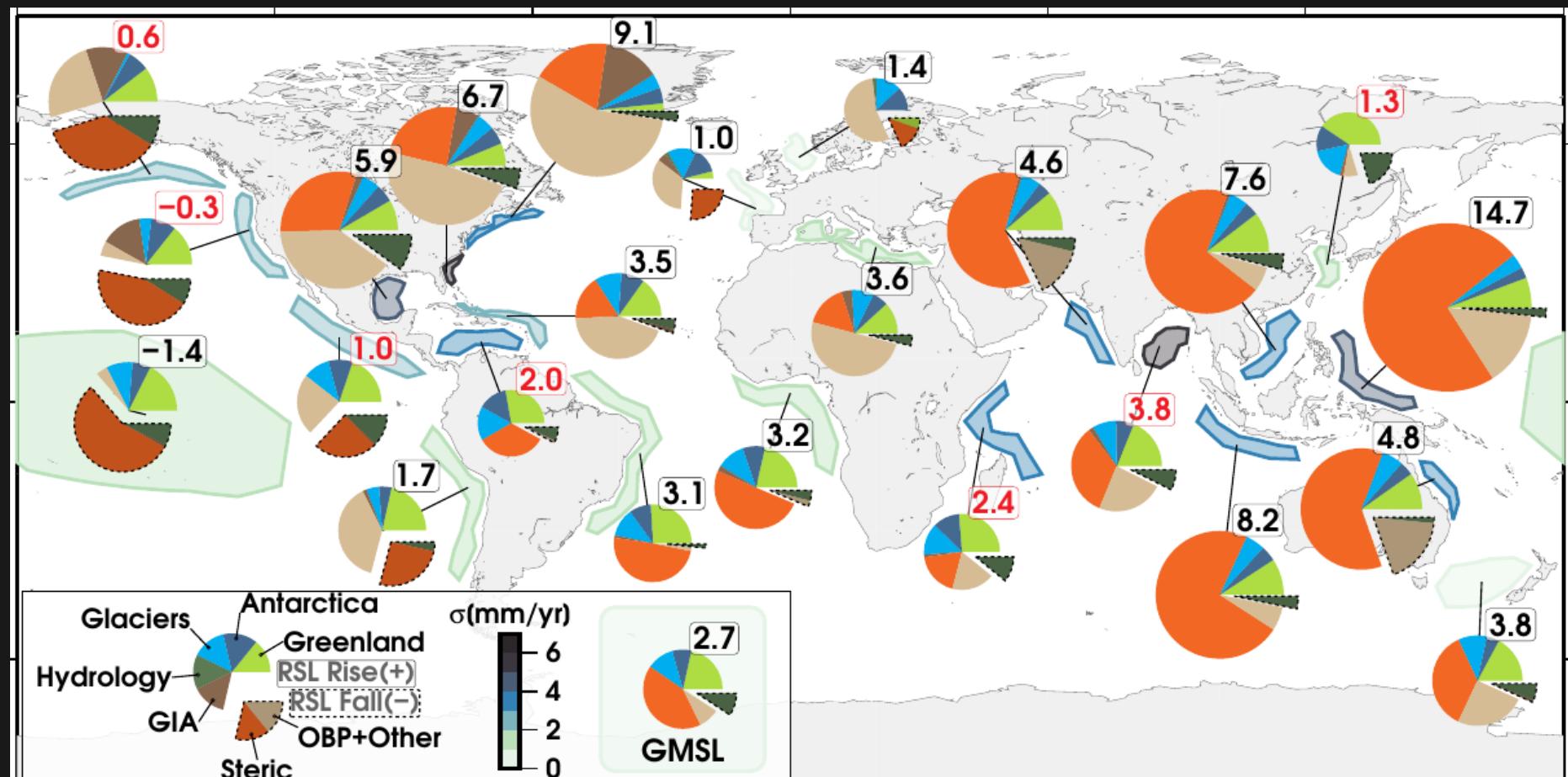


Regional sea level budgets from Rietbroek et al. 2016

- Sea level rise is not uniform
- SAL effects



Regional sea level budgets



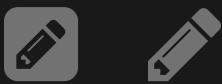
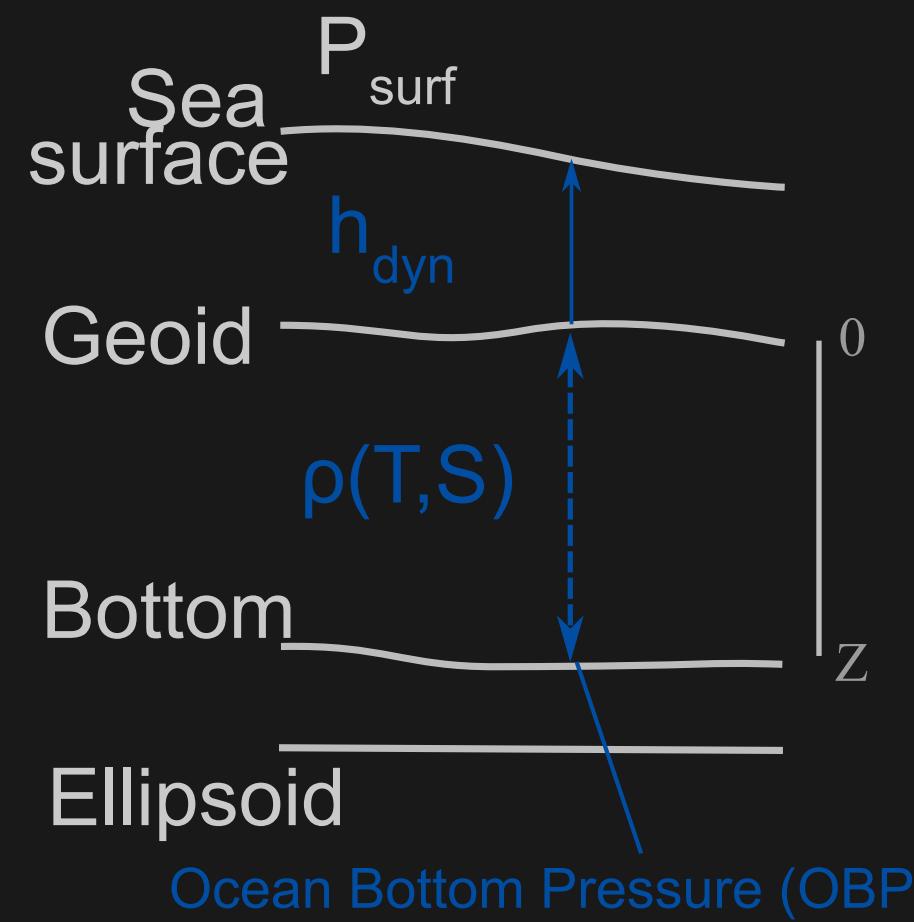
Regional sea level budgets from Rietbroek et al. 2016

- Sea level rise is not uniform
- SAL effects
- Thermosteric and ocean bottom pressure changes play a larger role in the regional budgets

Churning the ocean (forcing from wind stress, density contrasts)

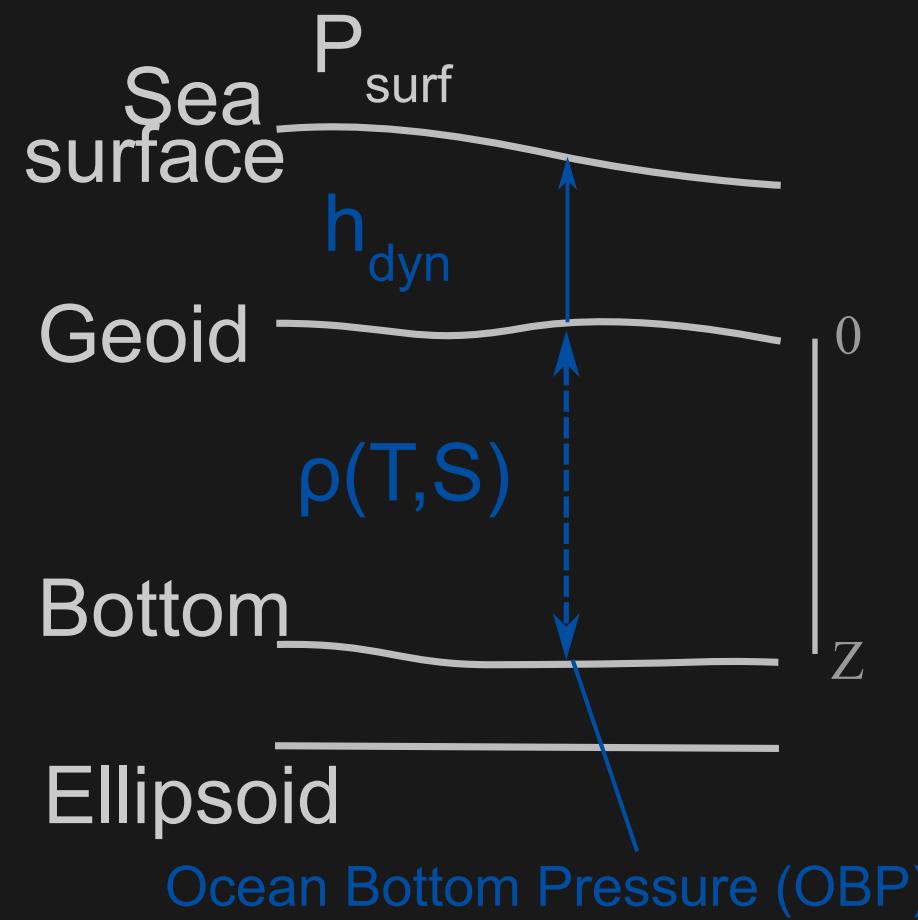


Integrating the water column to obtain ocean bottom pressure

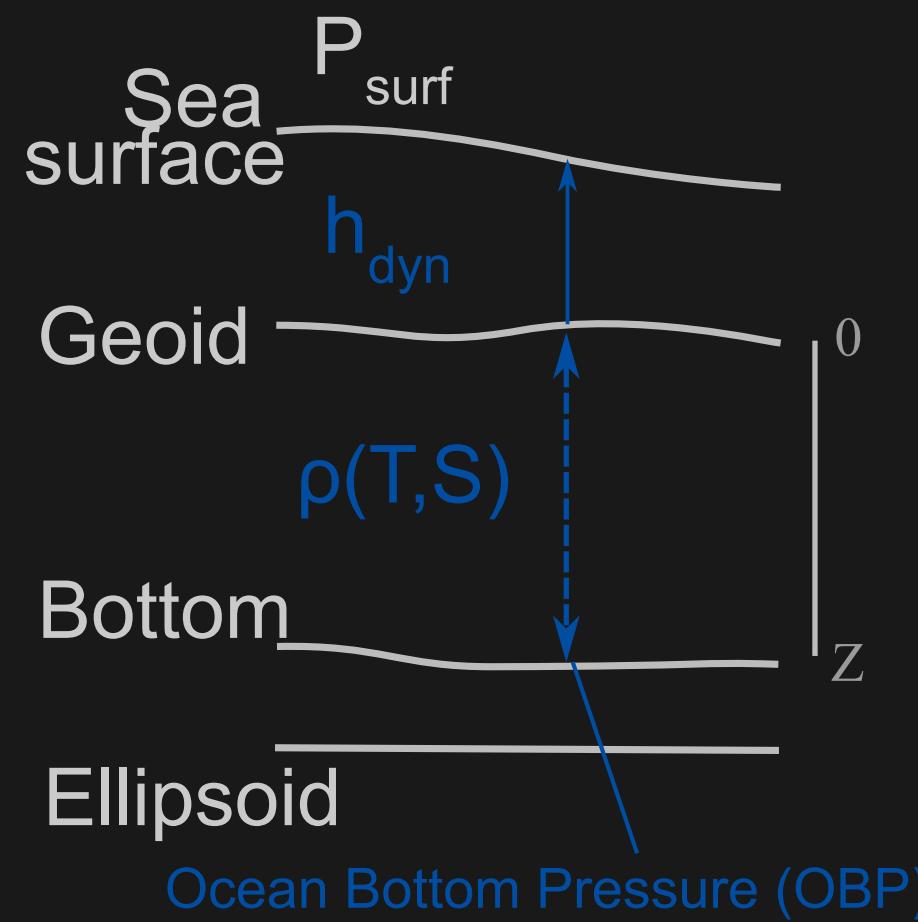


Integrating the water column to obtain ocean bottom pressure

$$\begin{aligned} P_{OBP} &= P_{surf} + g \\ &\int_{-Z}^0 \rho(z, T, S) dz \\ &+ g\rho_{sea} h_{dyn} \end{aligned}$$



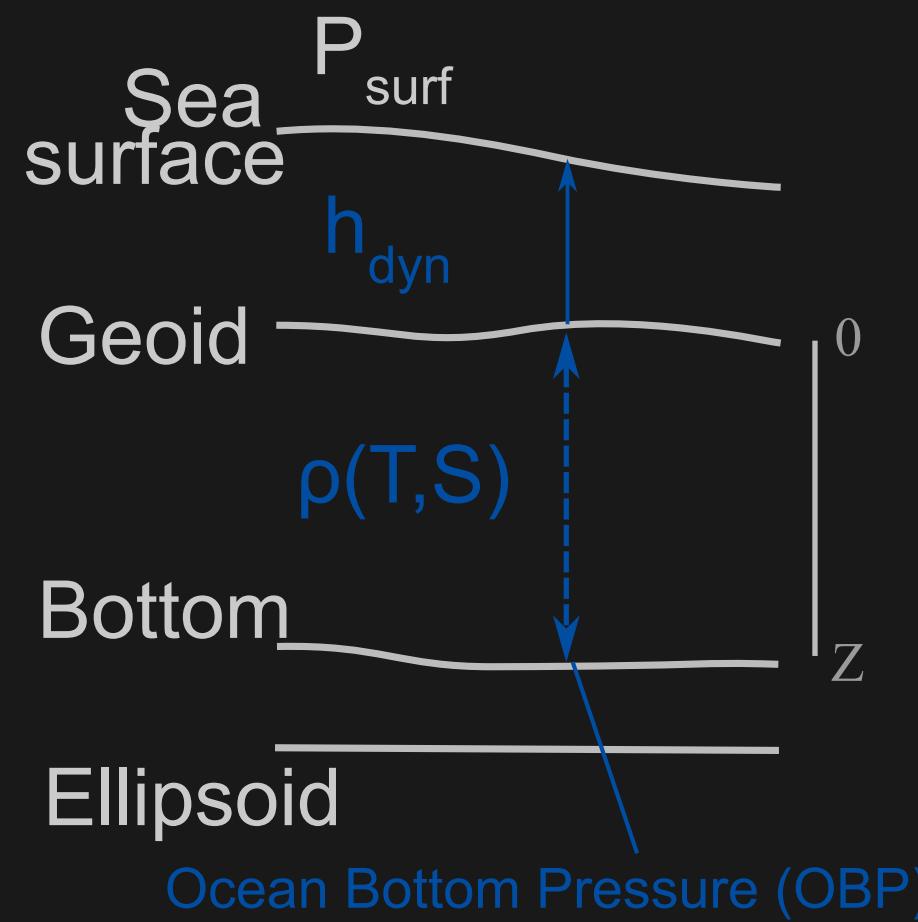
Integrating the water column to obtain ocean bottom pressure



$$\begin{aligned} P_{OBP} &= P_{surf} + g \\ &\int_{-Z}^0 \rho(z, T, S) dz \\ &+ g\rho_{sea} h_{dyn} \end{aligned}$$

No OBP change \rightarrow Steric change ($h_{dyn} = h_{ster}$)
 $0 = g \int_{-Z}^0 \delta\rho(z, T, S) dz + g\rho_{sea} h_{ster}$

Integrating the water column to obtain ocean bottom pressure



$$POBP = P_{surf} + g$$

$$\int_{-Z}^0 \rho(z, T, S) dz$$

$$+ g\rho_{sea} h_{dyn}$$

No OBP change \rightarrow Steric change ($h_{dyn} = h_{ster}$)

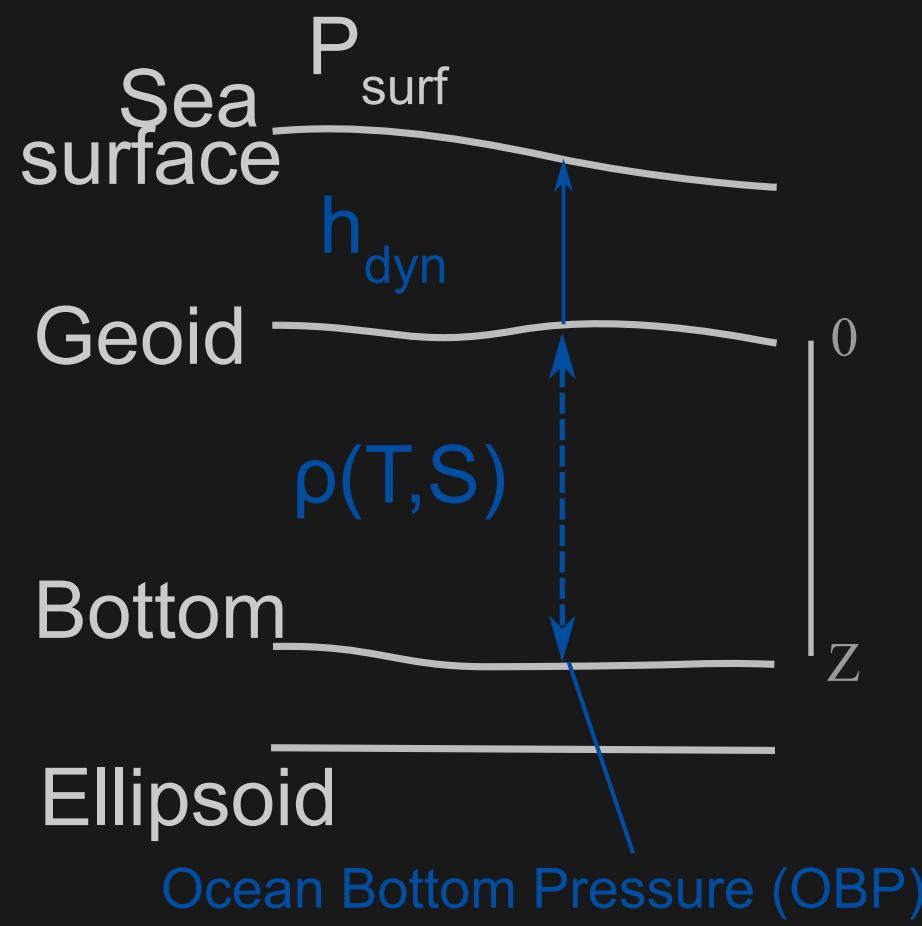
$$0 = g \int_{-Z}^0 \delta\rho(z, T,$$

$$S) dz + g\rho_{sea} h_{ster}$$

$$h_{ster} = -\frac{1}{\rho_{sea}}$$

$$\int_{-Z}^0 \delta\rho(z, T, S) dz$$

Integrating the water column to obtain ocean bottom pressure



$$P_{OBP} = P_{surf} + g$$

$$\int_{-Z}^0 \rho(z, T, S) dz$$

$$+ g\rho_{sea} h_{dyn}$$

No OBP change \rightarrow Steric change ($h_{dyn} = h_{ster}$)

$$0 = g \int_{-Z}^0 \delta\rho(z, T,$$

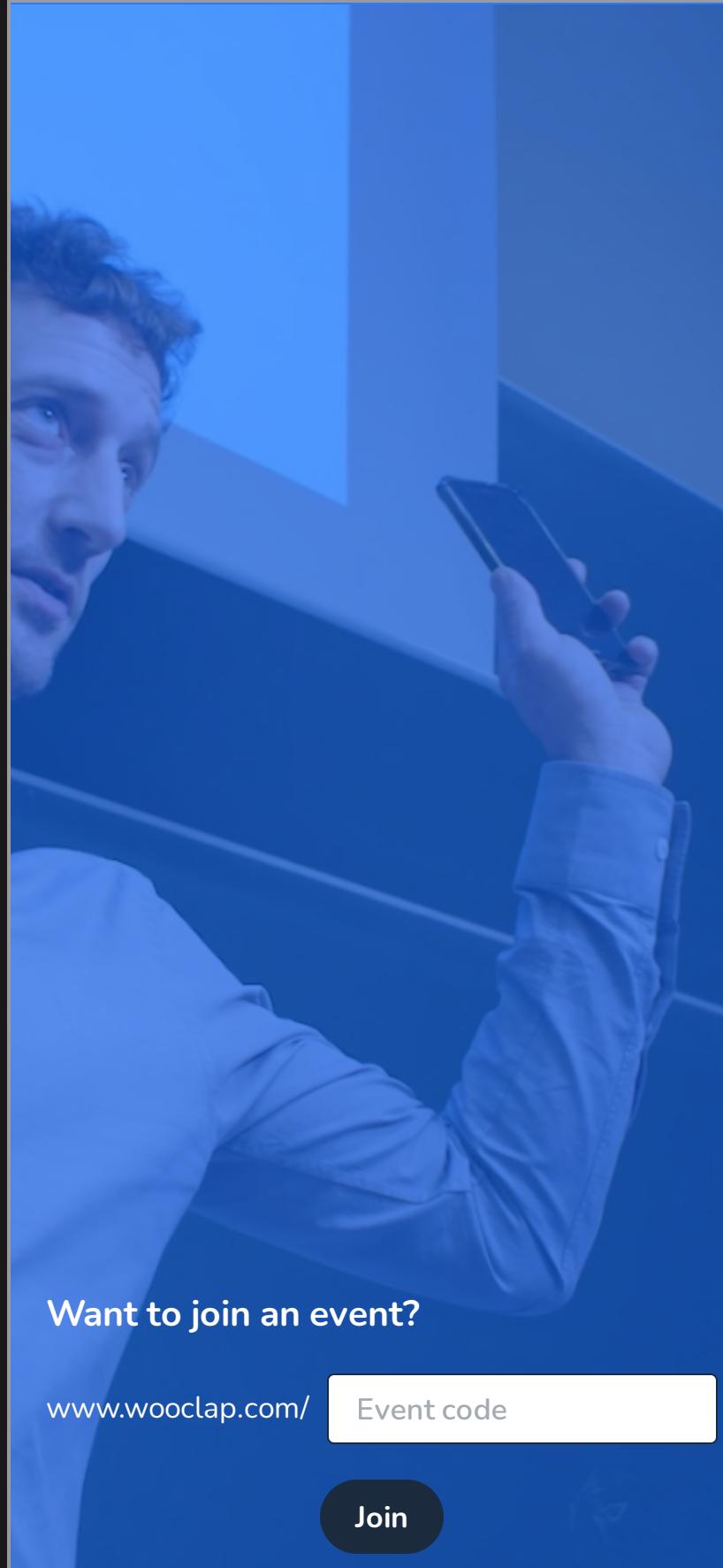
$$S) dz + g\rho_{sea} h_{ster}$$

$$h_{ster} = -\frac{1}{\rho_{sea}}$$

$$\int_{-Z}^0 \delta\rho(z, T, S) dz$$

- h_{ster} invisible to GRACE! ($P_{OBP} = 0$)
- h_{ster} visible by radar altimetry!
- Density \uparrow results in \downarrow of sea level
- Wind driven (quick) \rightarrow change in h_{dyn}
- Density driven (slow) $\rightarrow h_{dyn} \approx h_{ster}$

Quiz time! (again)



Want to join an event?

www.wooclap.com/

Join

Choose a method to log in
[or sign up](#)

Your email address 

OR

 Sign up with Facebook

 Sign up with Google

 Sign up with LinkedIn

 Sign up with Microsoft

OR

Log in with your institution

Your institution 



We use [cookies](#) to improve the general experience on the platform, provide users chat support and deliver targeted ads on other websites.

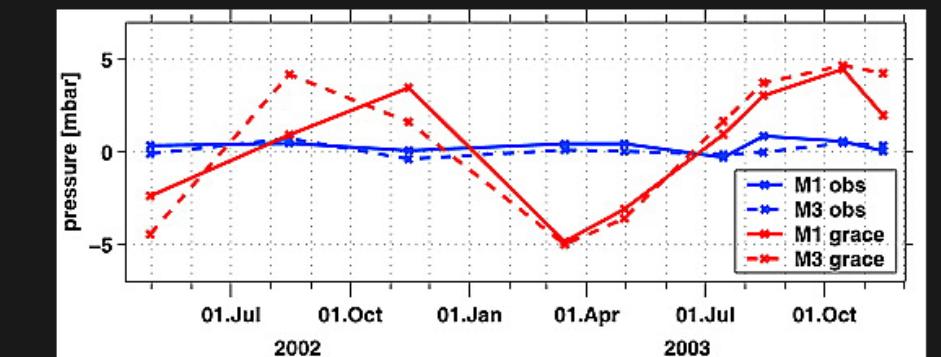
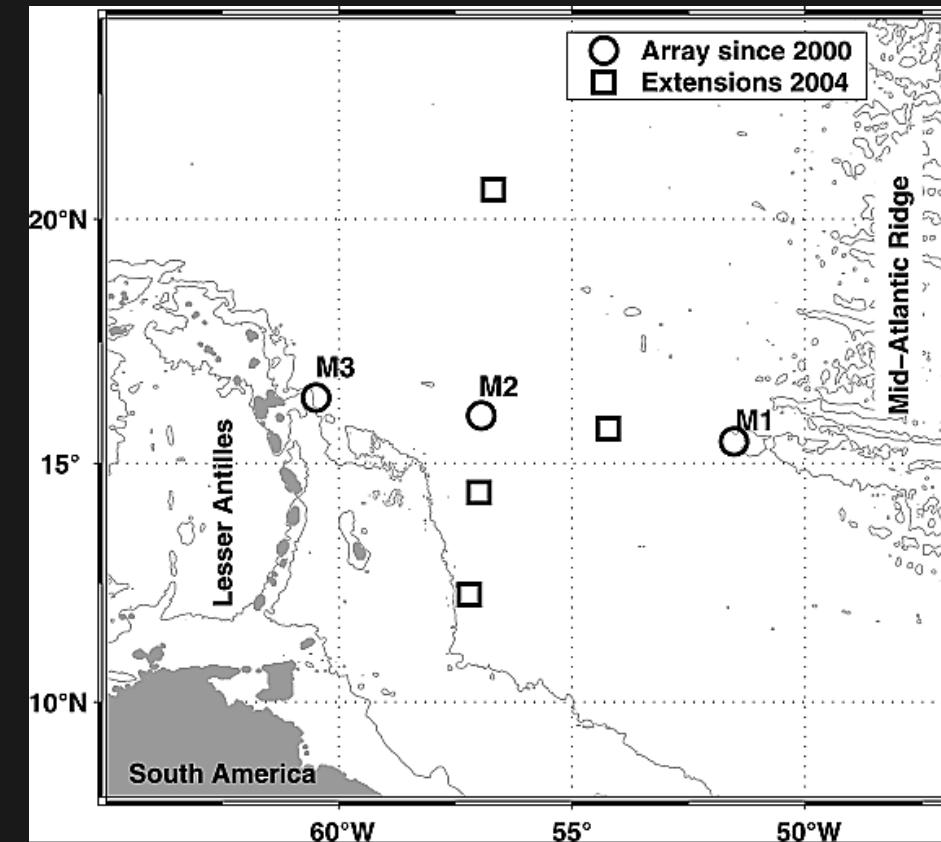
Accept all

Reject All

Customise

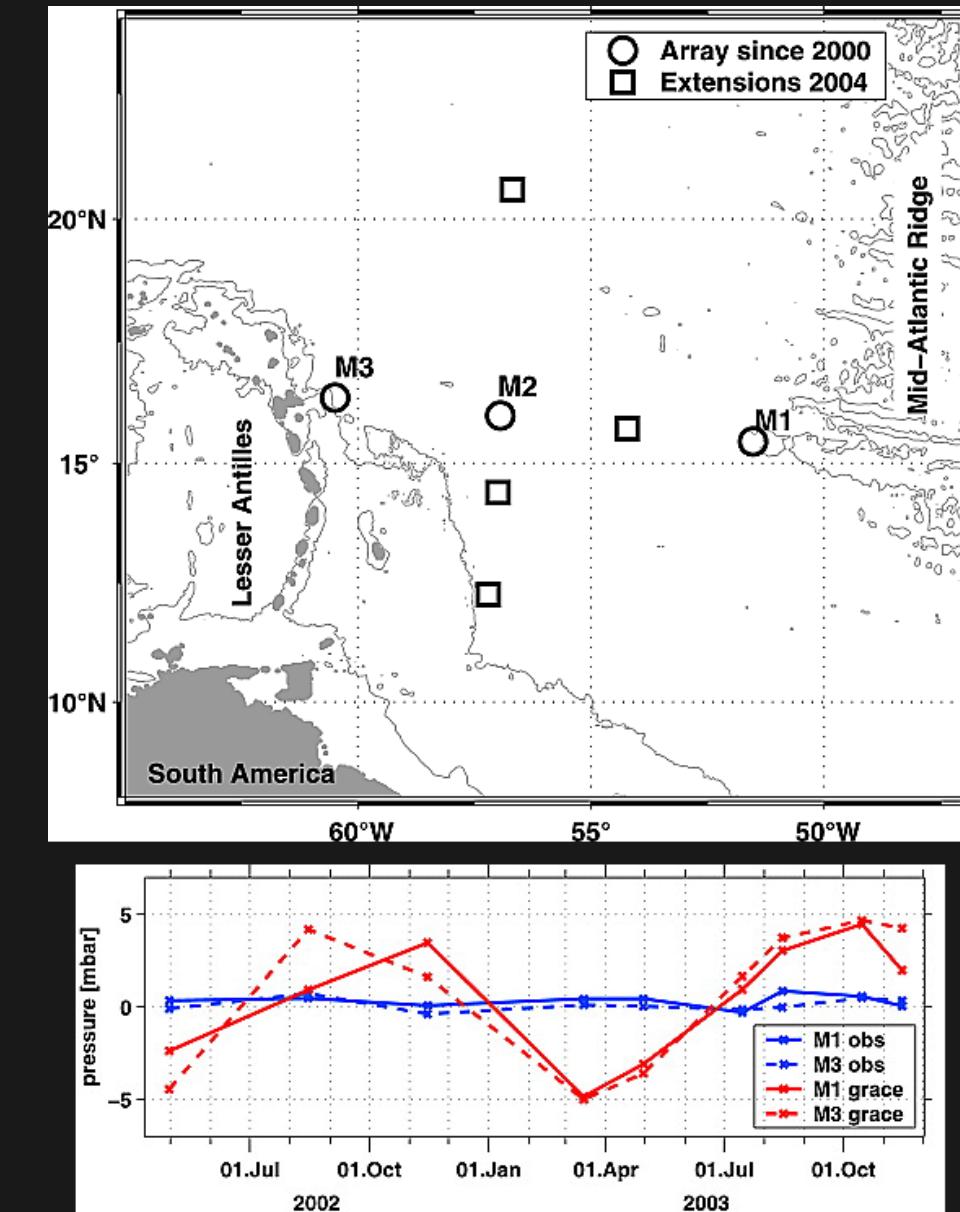


Early GRACE validation of ocean bottom pressure changes (Kanzow et al. 2005)



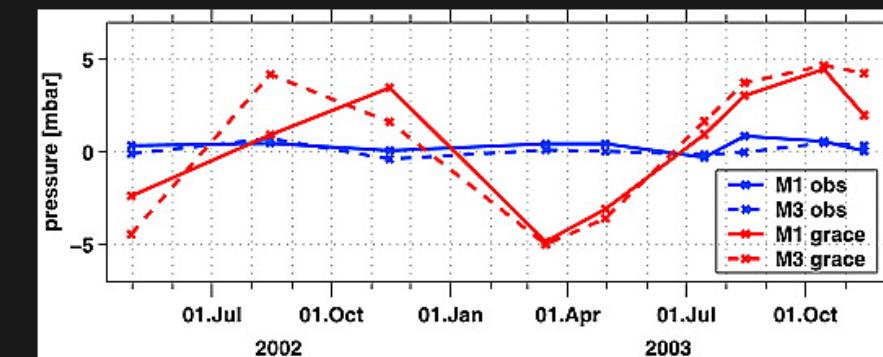
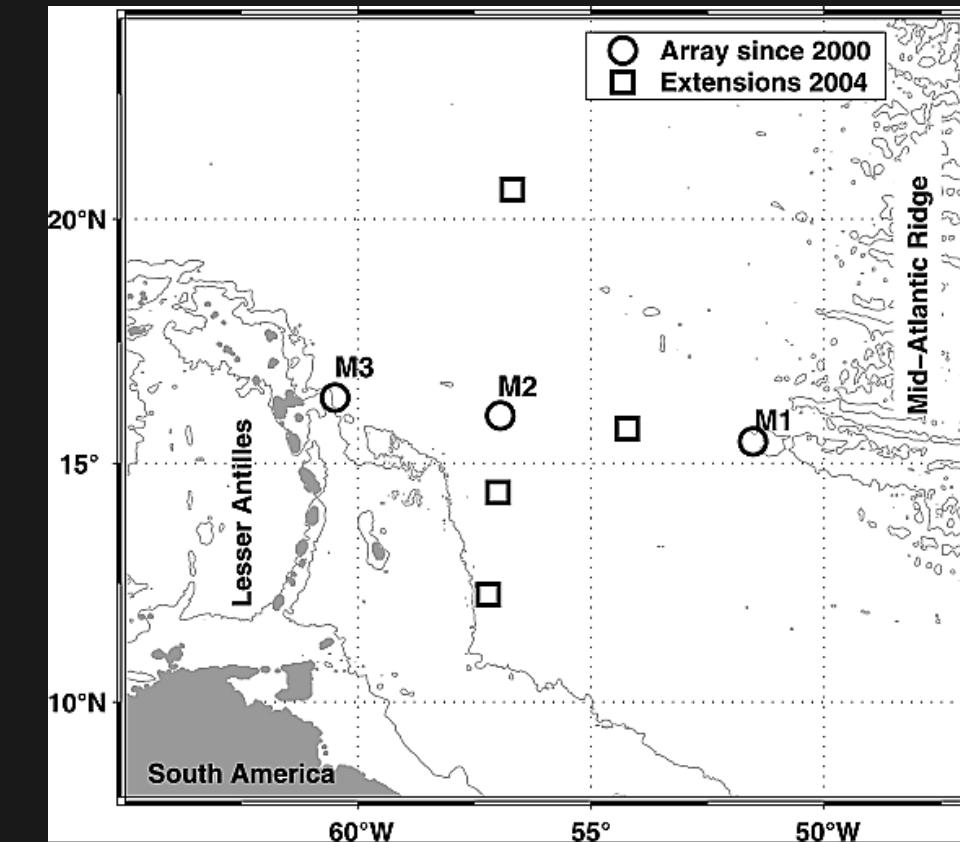
Early GRACE validation of ocean bottom pressure changes (Kanzow et al. 2005)

- In situ validation with pressure sensors



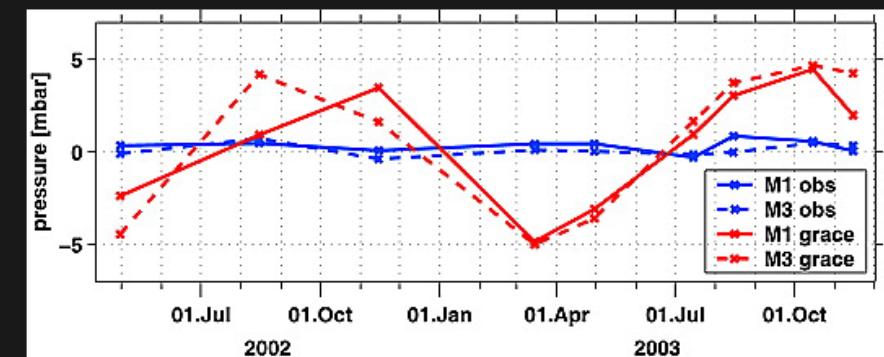
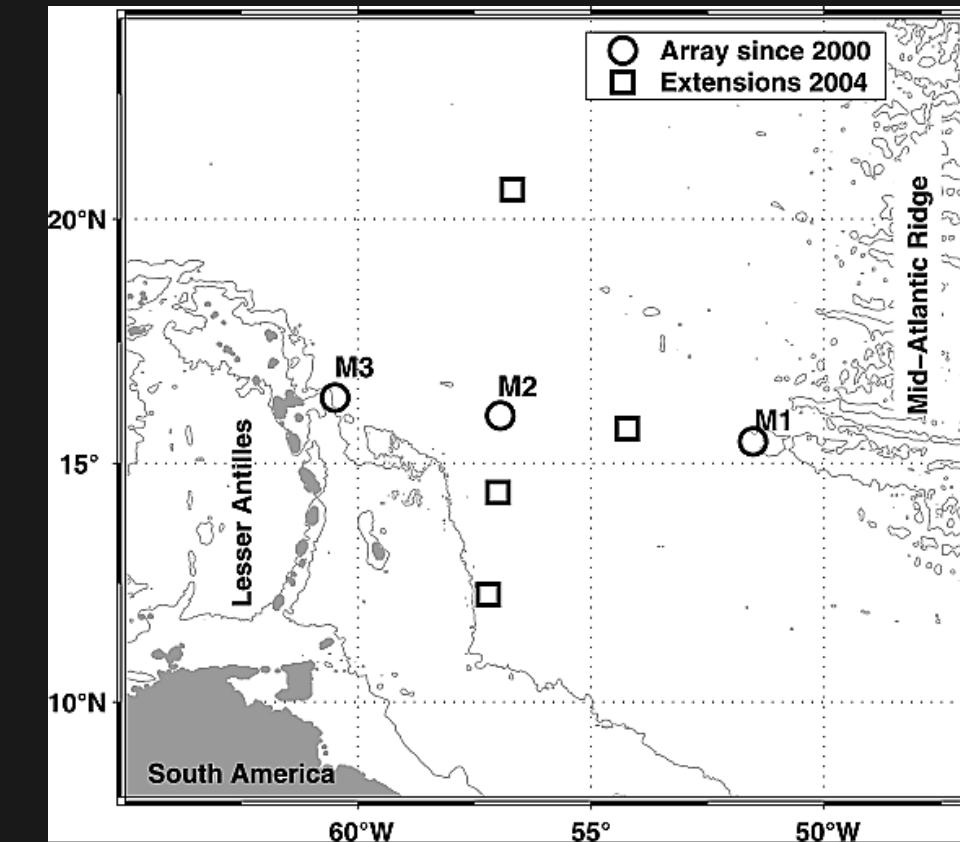
Early GRACE validation of ocean bottom pressure changes (Kanzow et al. 2005)

- In situ validation with pressure sensors
- GRACE fluctuations were too large



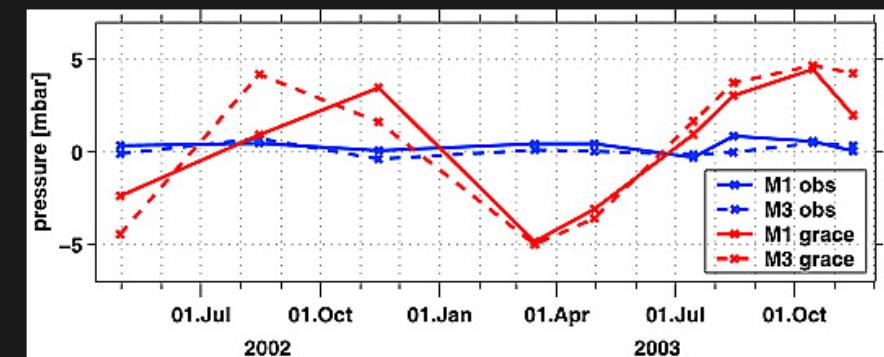
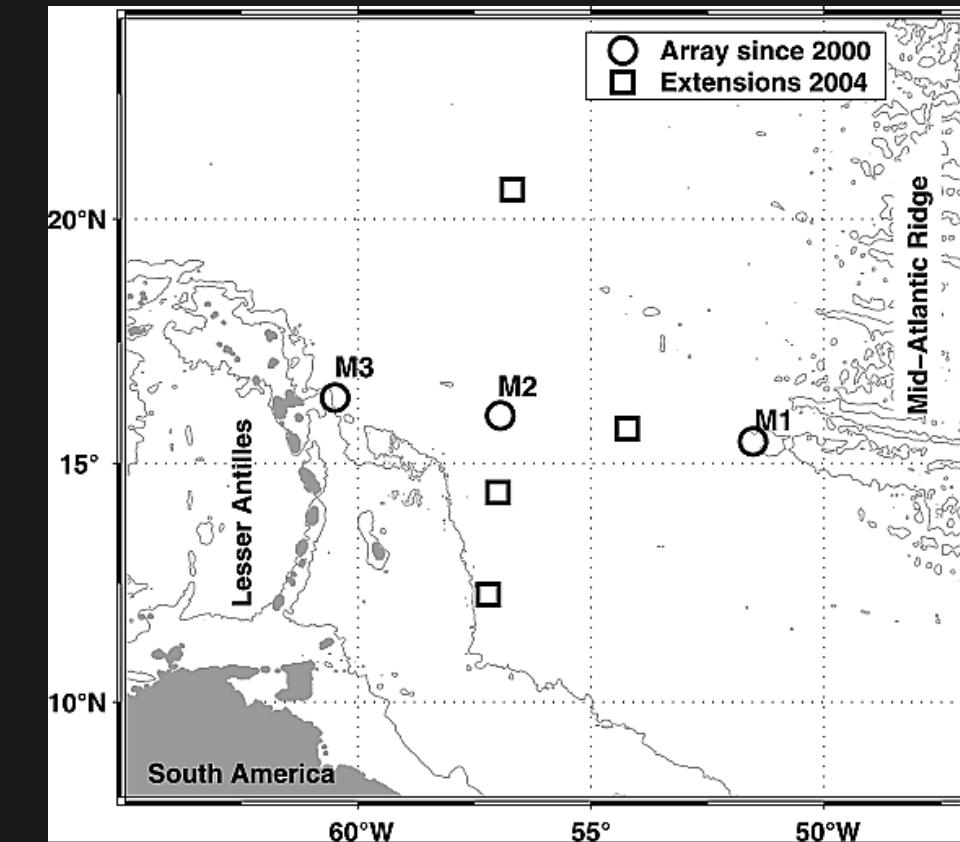
Early GRACE validation of ocean bottom pressure changes (Kanzow et al. 2005)

- In situ validation with pressure sensors
- GRACE fluctuations were too large
 - Signal leakage

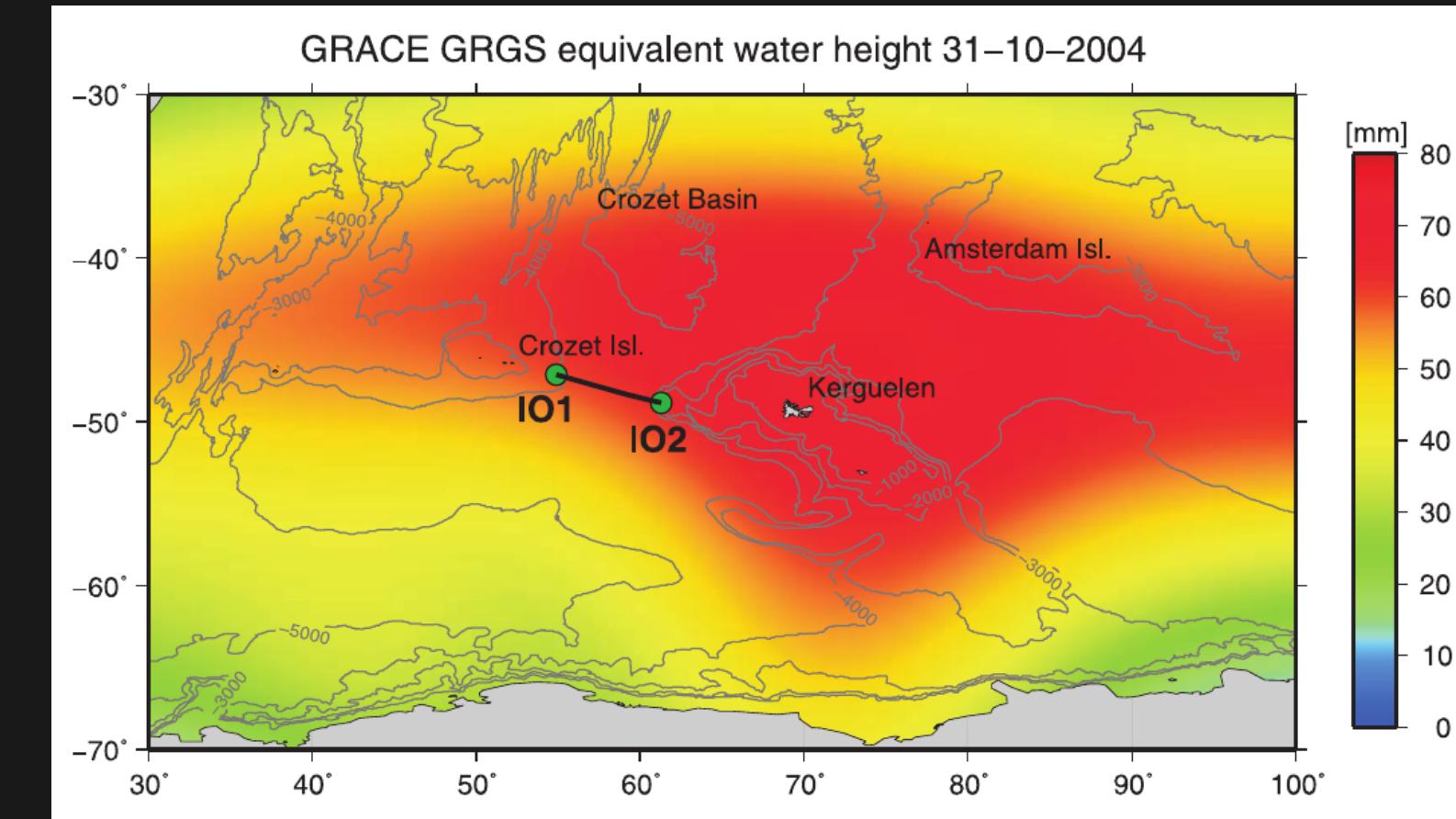
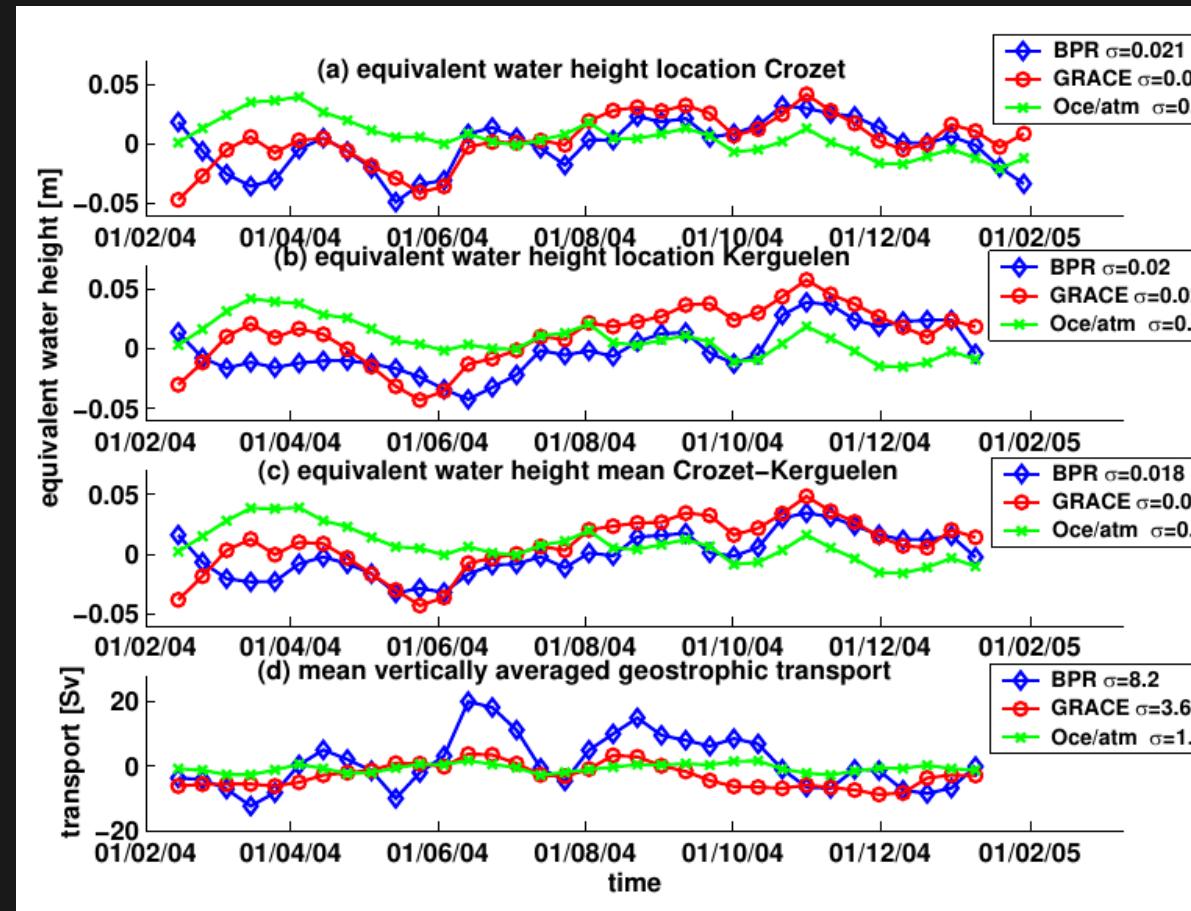


Early GRACE validation of ocean bottom pressure changes (Kanzow et al. 2005)

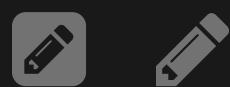
- In situ validation with pressure sensors
- GRACE fluctuations were too large
 - Signal leakage
 - Processing errors



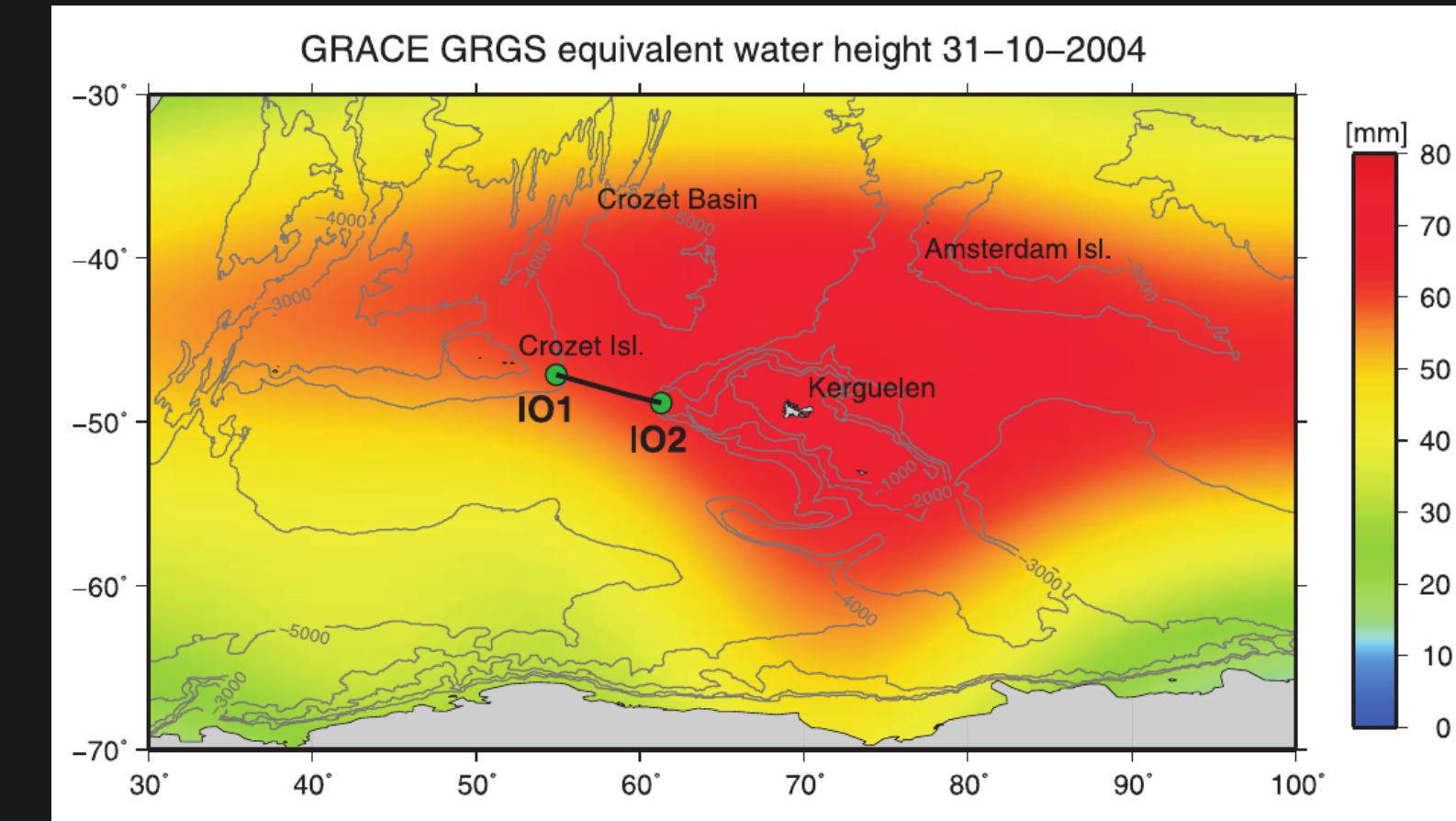
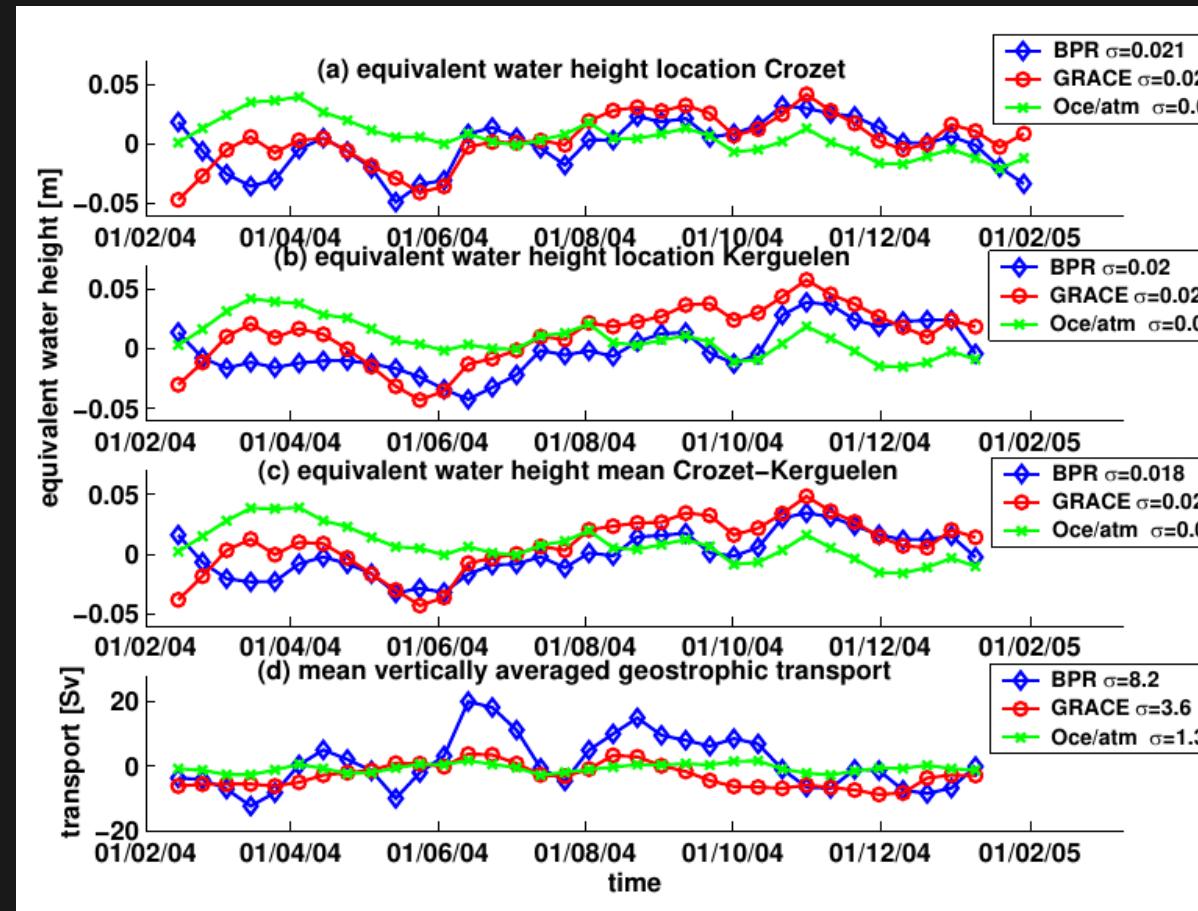
Validation in the Southern Ocean



From Rietbroek et al. 2006



Validation in the Southern Ocean

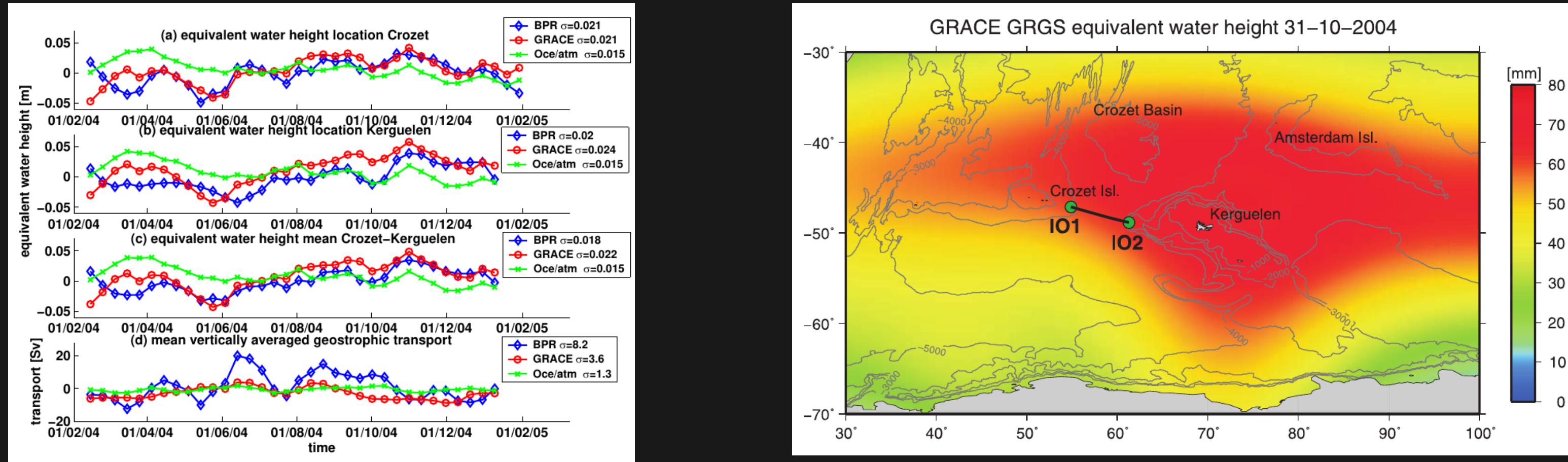


From Rietbroek et al. 2006

- In situ validation with pressure sensors

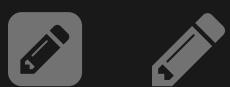


Validation in the Southern Ocean

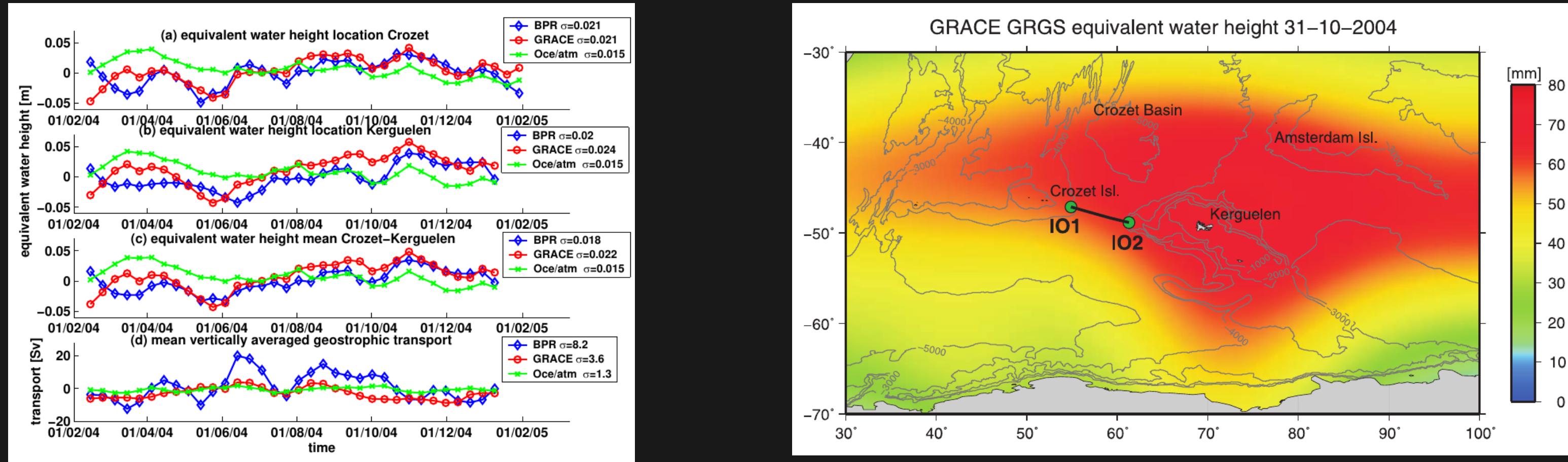


From Rietbroek et al. 2006

- In situ validation with pressure sensors
- 10-day GRACE solution was better than background model

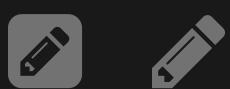


Validation in the Southern Ocean

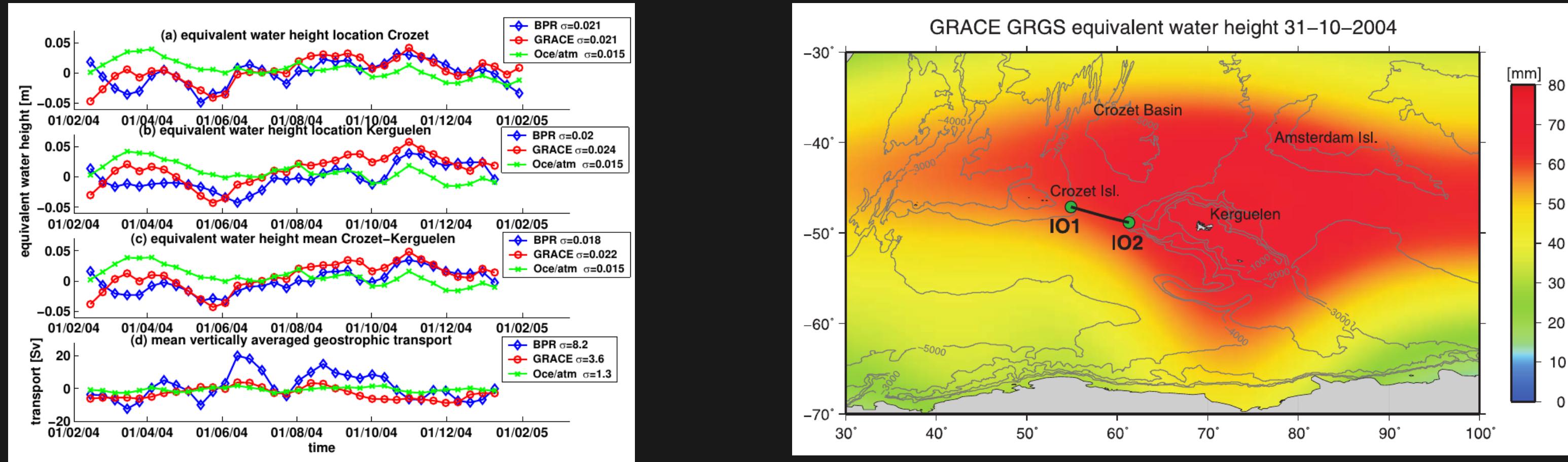


From Rietbroek et al. 2006

- In situ validation with pressure sensors
- 10-day GRACE solution was better than background model
- Large scale ocean signals picked up



Validation in the Southern Ocean

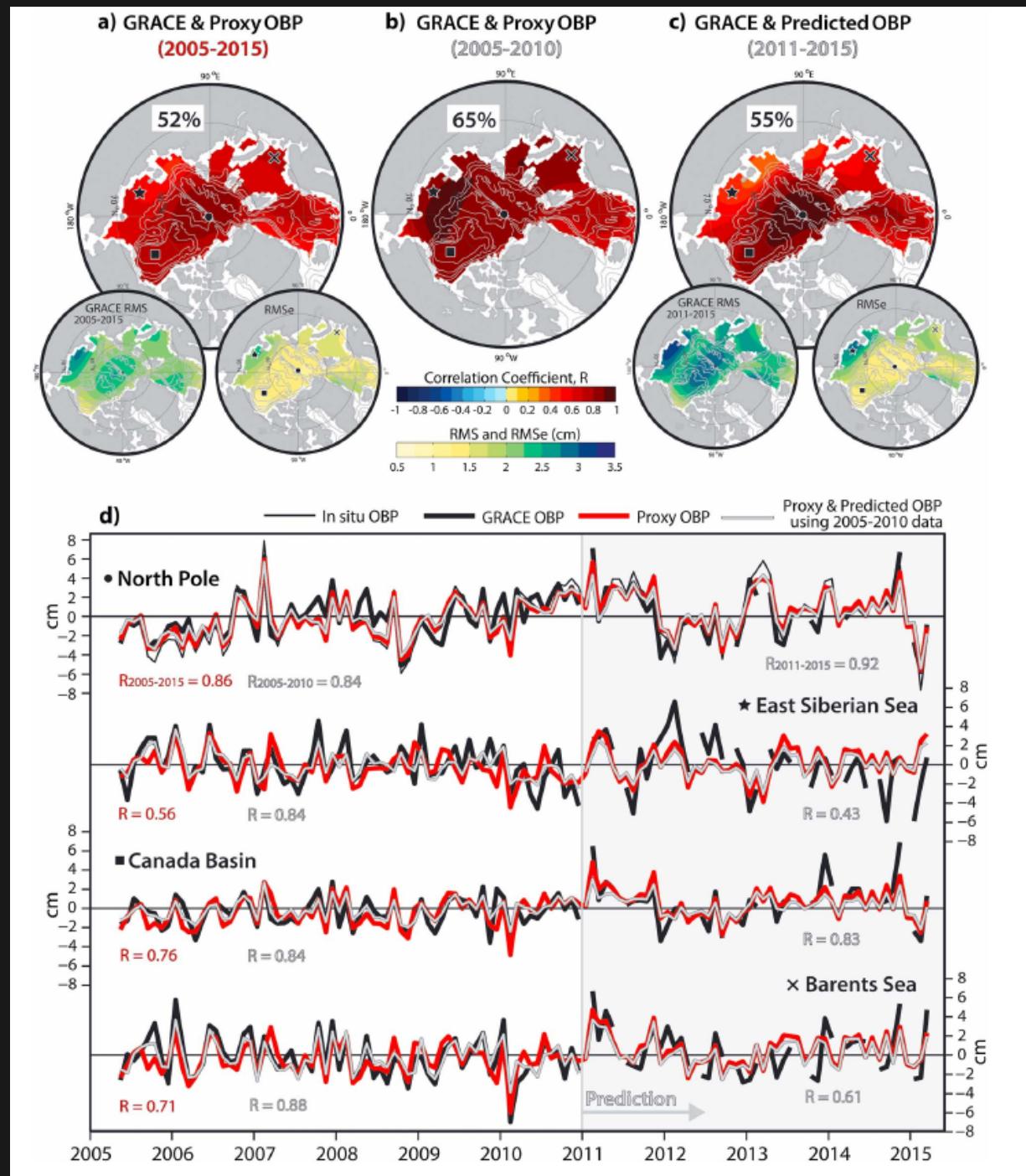


From Rietbroek et al. 2006

- In situ validation with pressure sensors
- 10-day GRACE solution was better than background model
- Large scale ocean signals picked up
- Due to smoothing geostrophic bottom currents not resolved



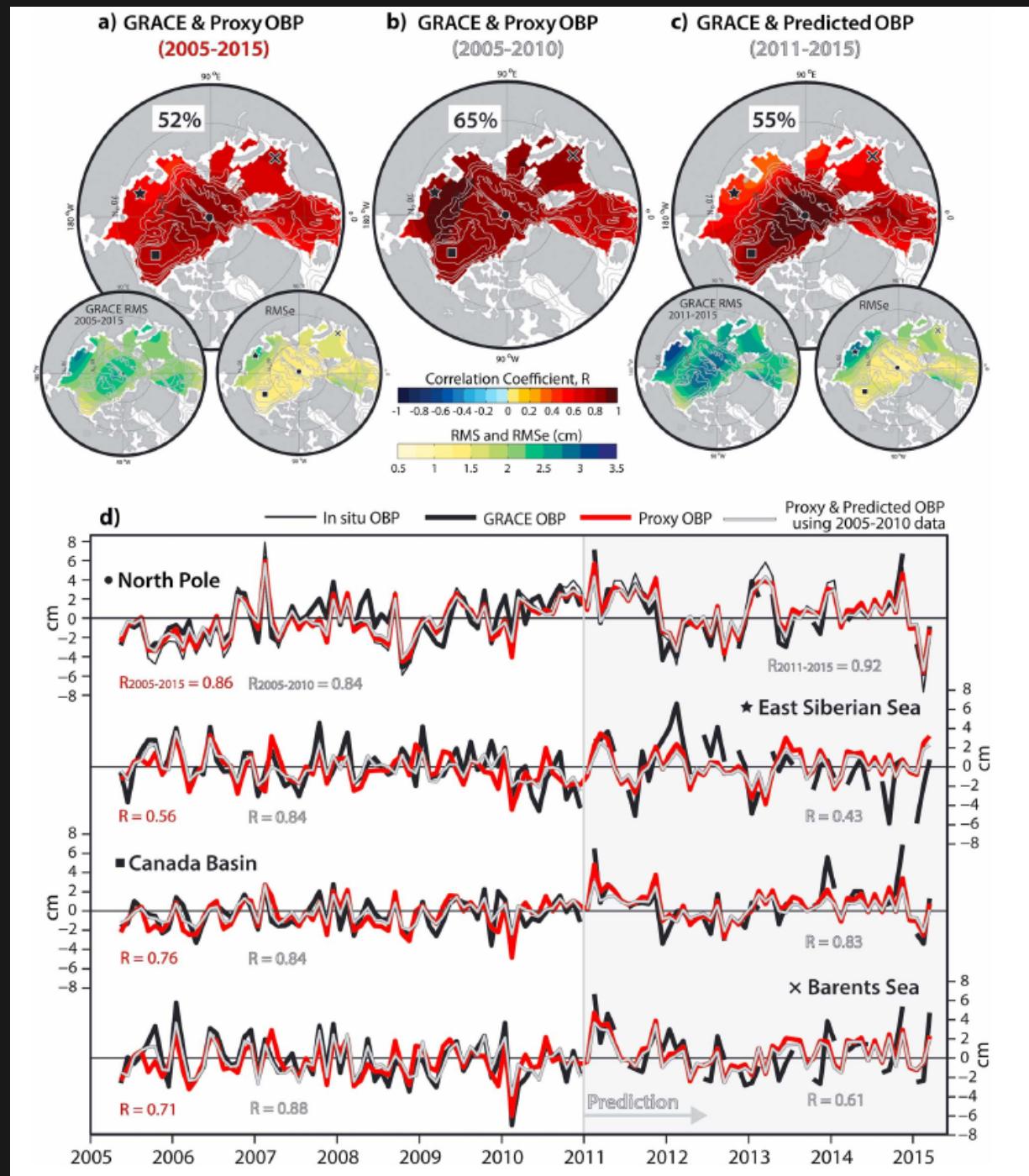
Arctic Ocean Bottom pressure variations well captured by GRACE



From Peralta-ferriz et al. 2016



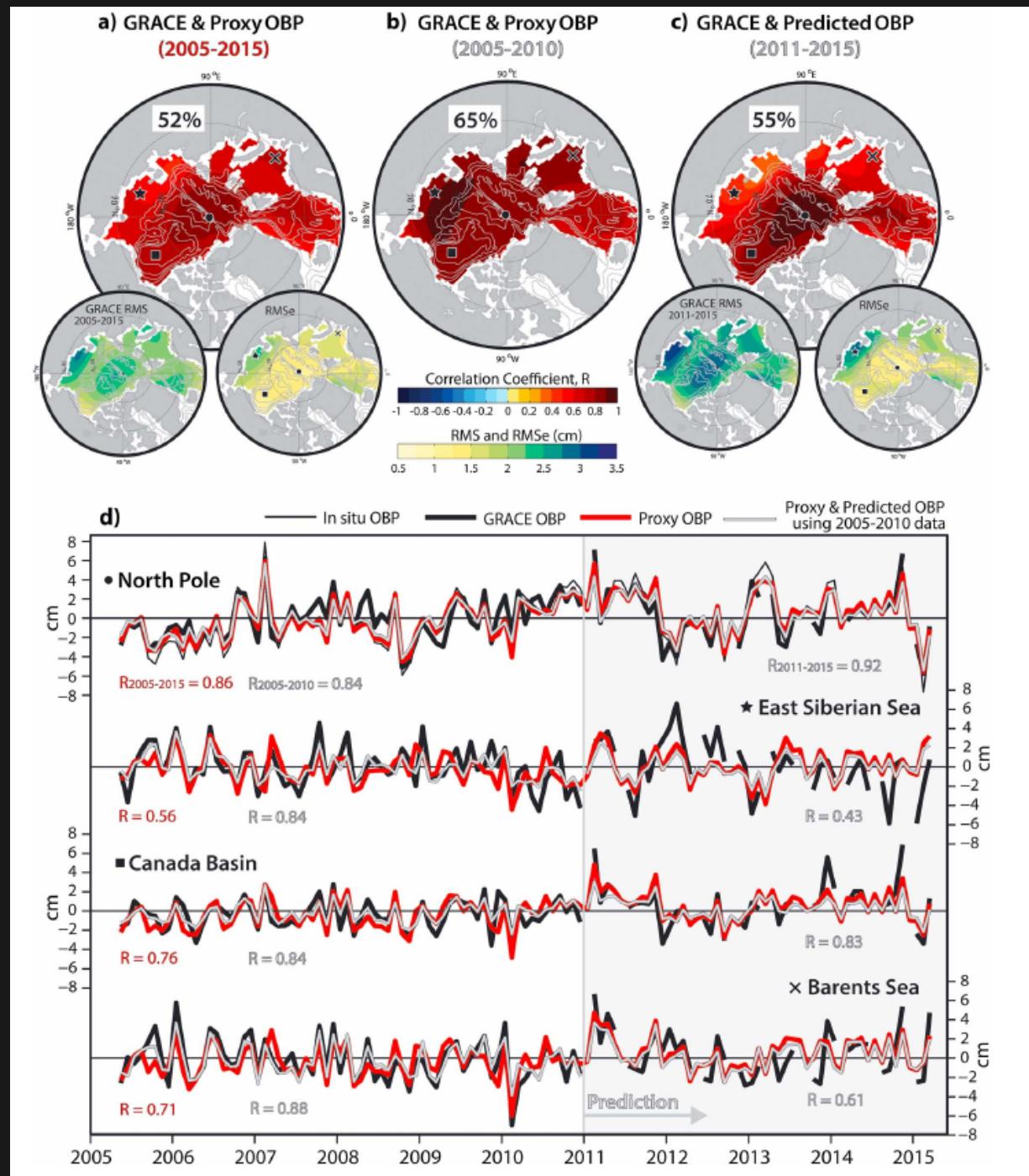
Arctic Ocean Bottom pressure variations well captured by GRACE



From Peralta-ferriz et al. 2016

- In situ OBP well captured by GRACE

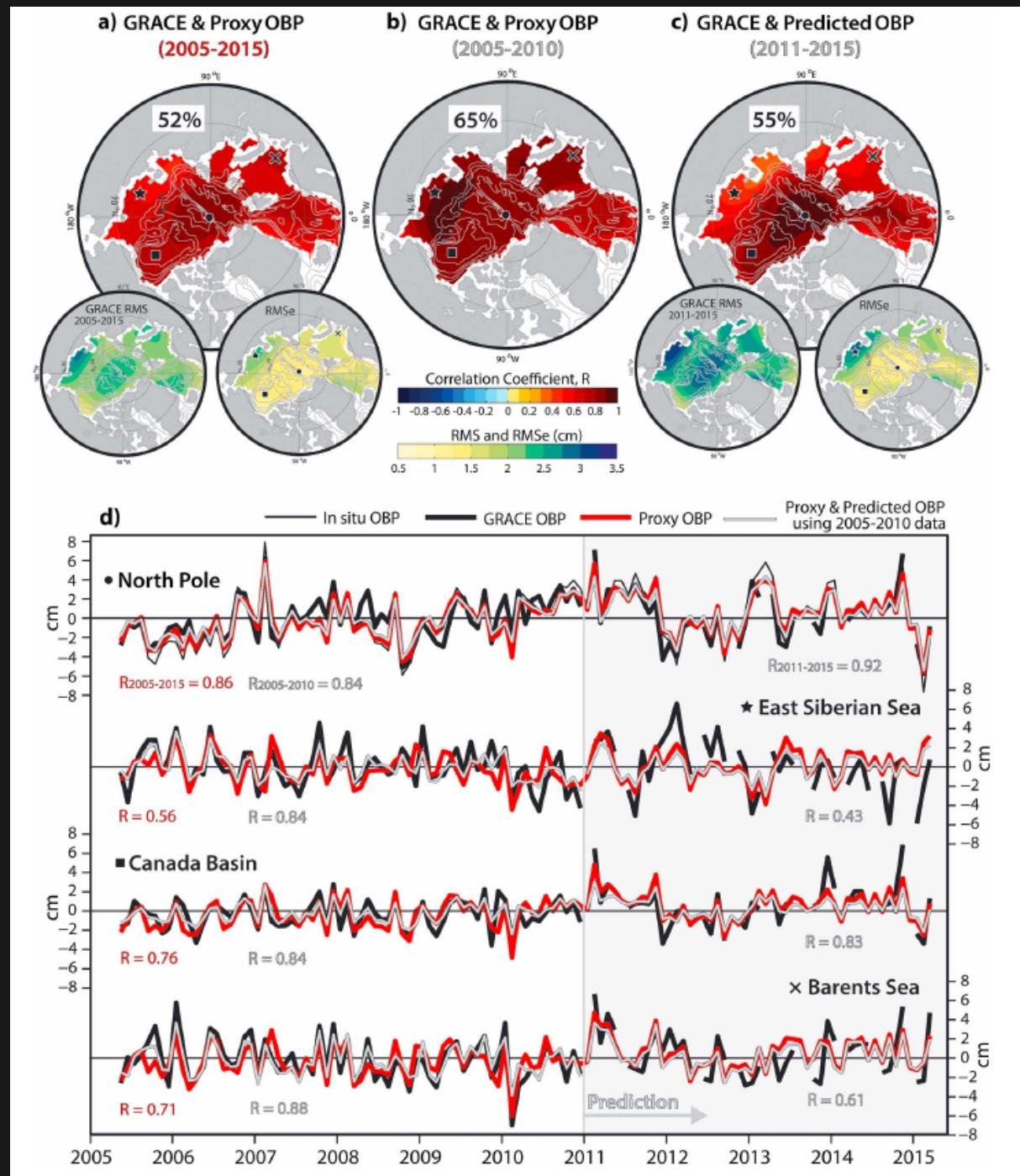
Arctic Ocean Bottom pressure variations well captured by GRACE



From Peralta-ferriz et al. 2016

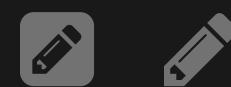
- In situ OBP well captured by GRACE
- Coherent OBP signals in the Arctic

Arctic Ocean Bottom pressure variations well captured by GRACE

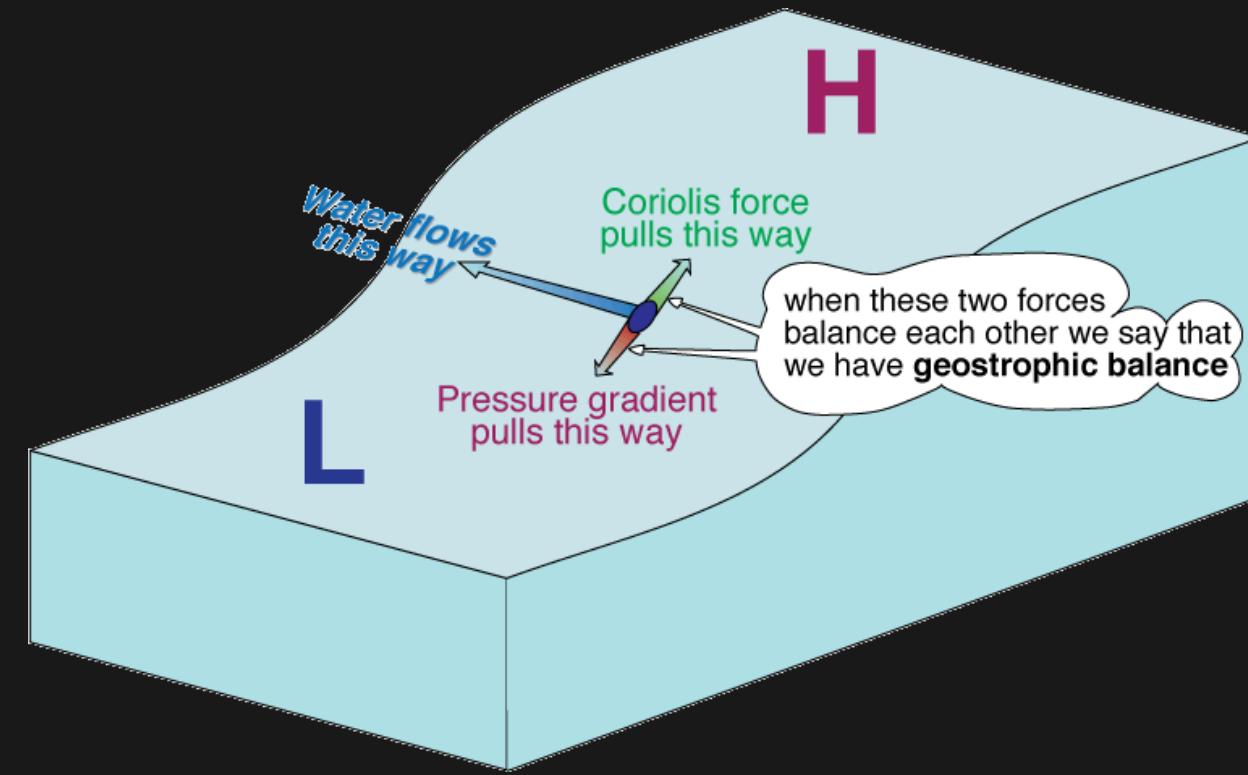
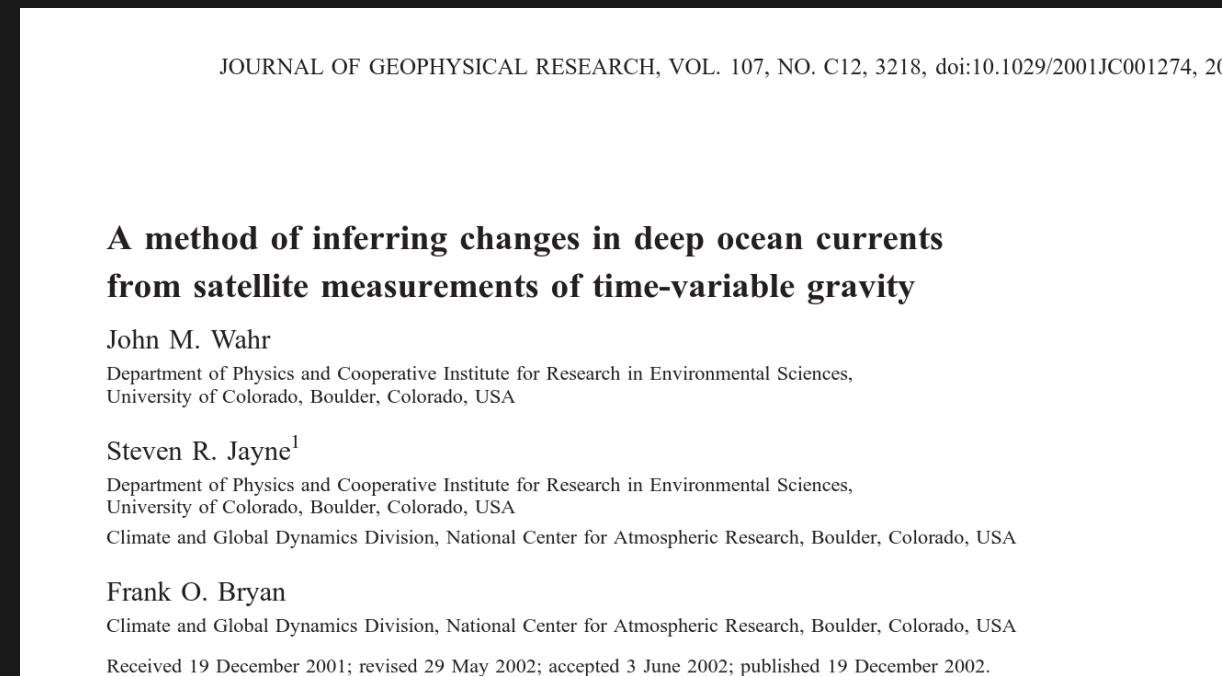


From Peralta-ferriz et al. 2016

- In situ OBP well captured by GRACE
- Coherent OBP signals in the Arctic
- Comparison showed that a single in situ OBP can serve as proxy for the entire Arctic



Geostrophic currents from time variable gravity



CC-NC-SA, <https://www.seos-project.eu>

the effects of friction and external forcing are less important,
(1) reduces to the geostrophic approximation:

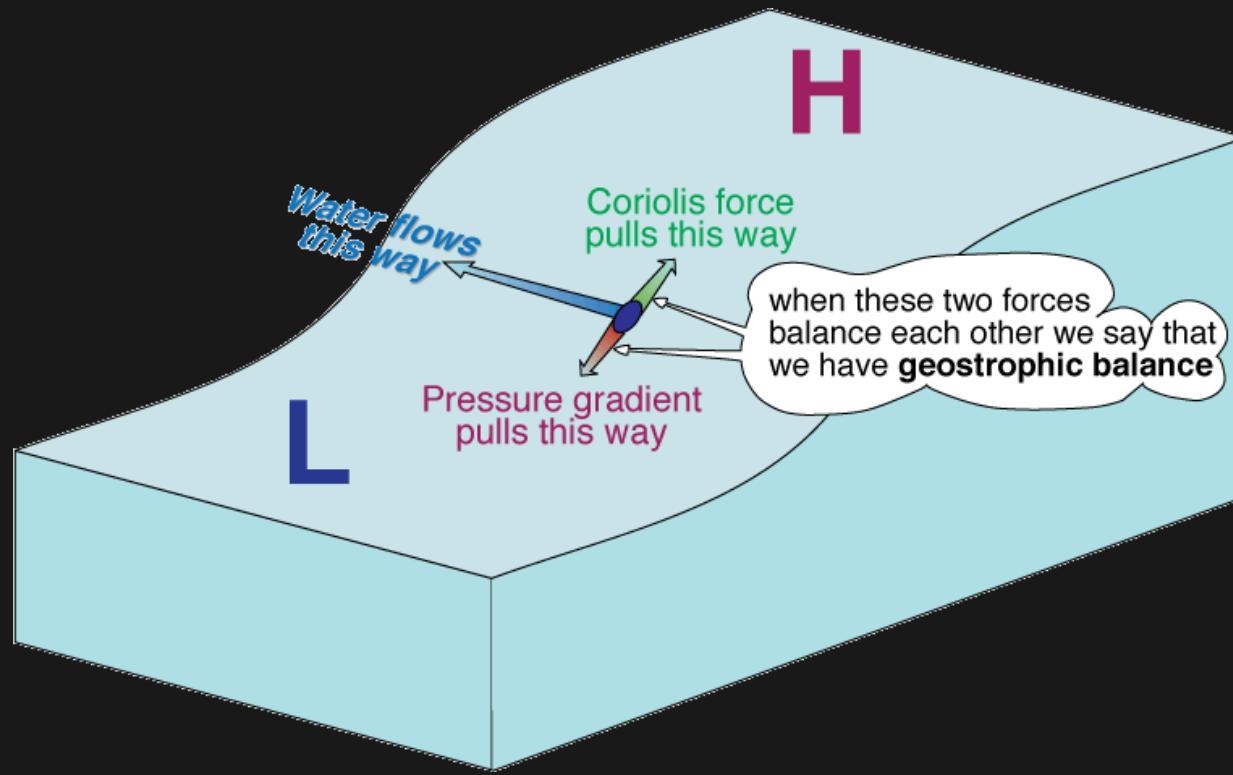
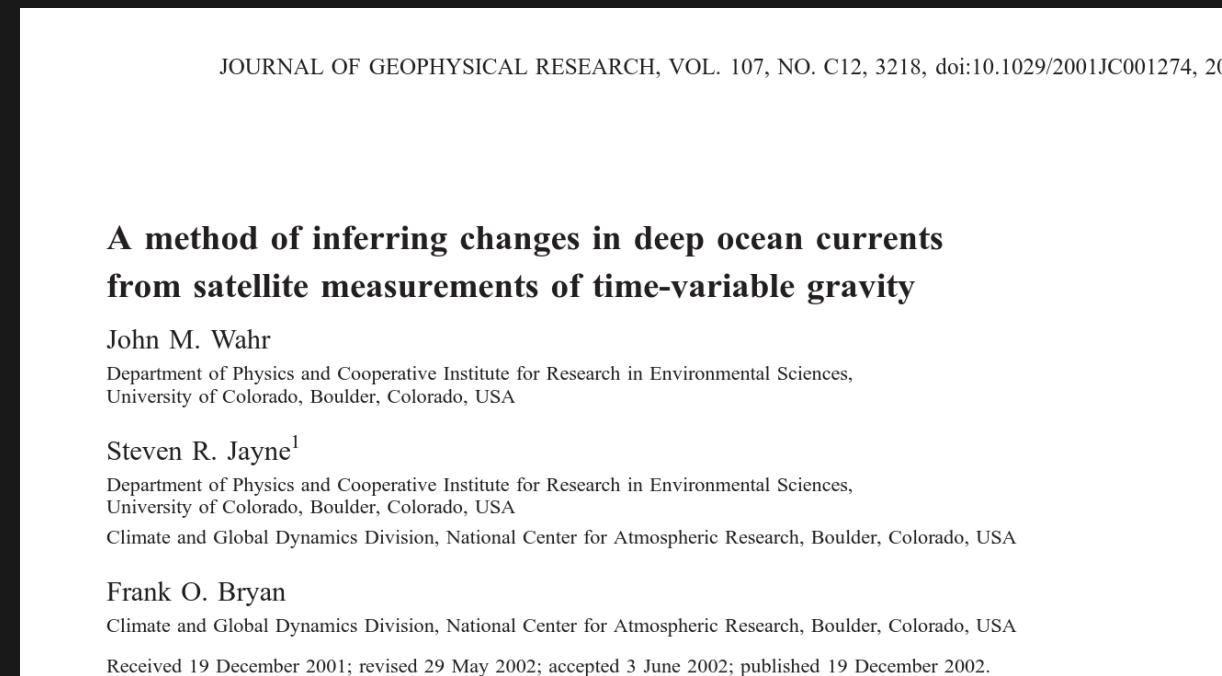
Pressure gradient

$$2\rho_0\Omega \times \mathbf{v} \approx -\nabla P \quad (2)$$

Earth rotation vector ↑ velocity



Geostrophic currents from time variable gravity



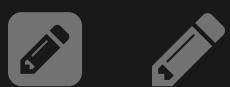
CC-NC-SA, <https://www.seos-project.eu>

- Assume no accelerations

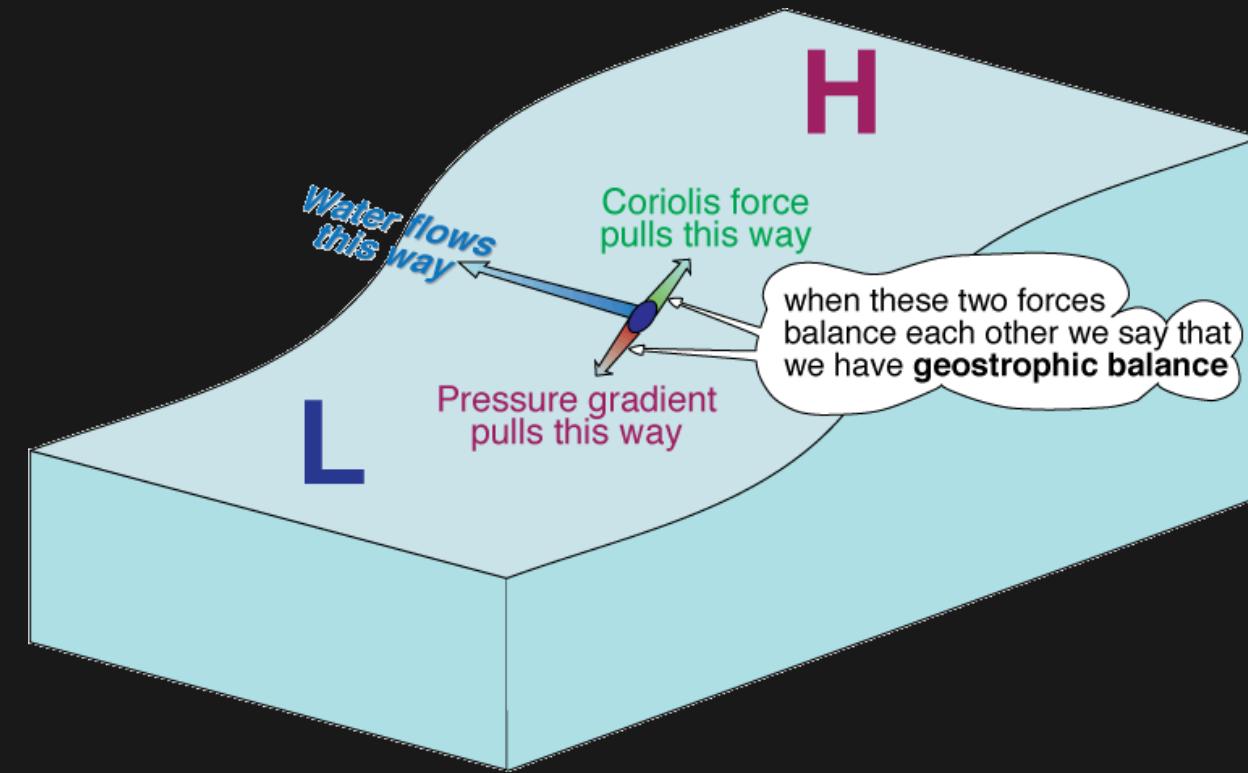
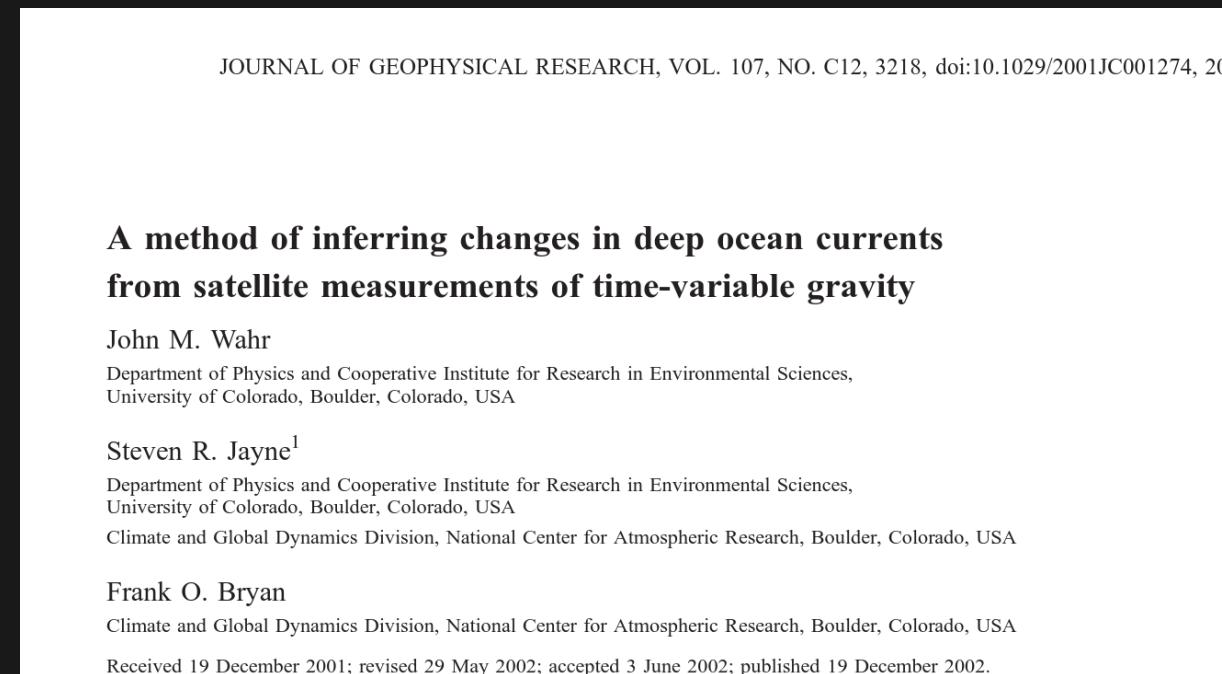
the effects of friction and external forcing are less important,
(1) reduces to the geostrophic approximation:

$$\text{Pressure gradient} \\ 2\rho_0 \Omega \times \mathbf{v} \approx -\nabla P \quad (2)$$

Earth rotation vector ↑ velocity



Geostrophic currents from time variable gravity



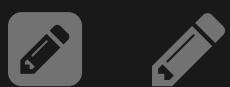
CC-NC-SA, <https://www.seos-project.eu>

the effects of friction and external forcing are less important,
(1) reduces to the geostrophic approximation:

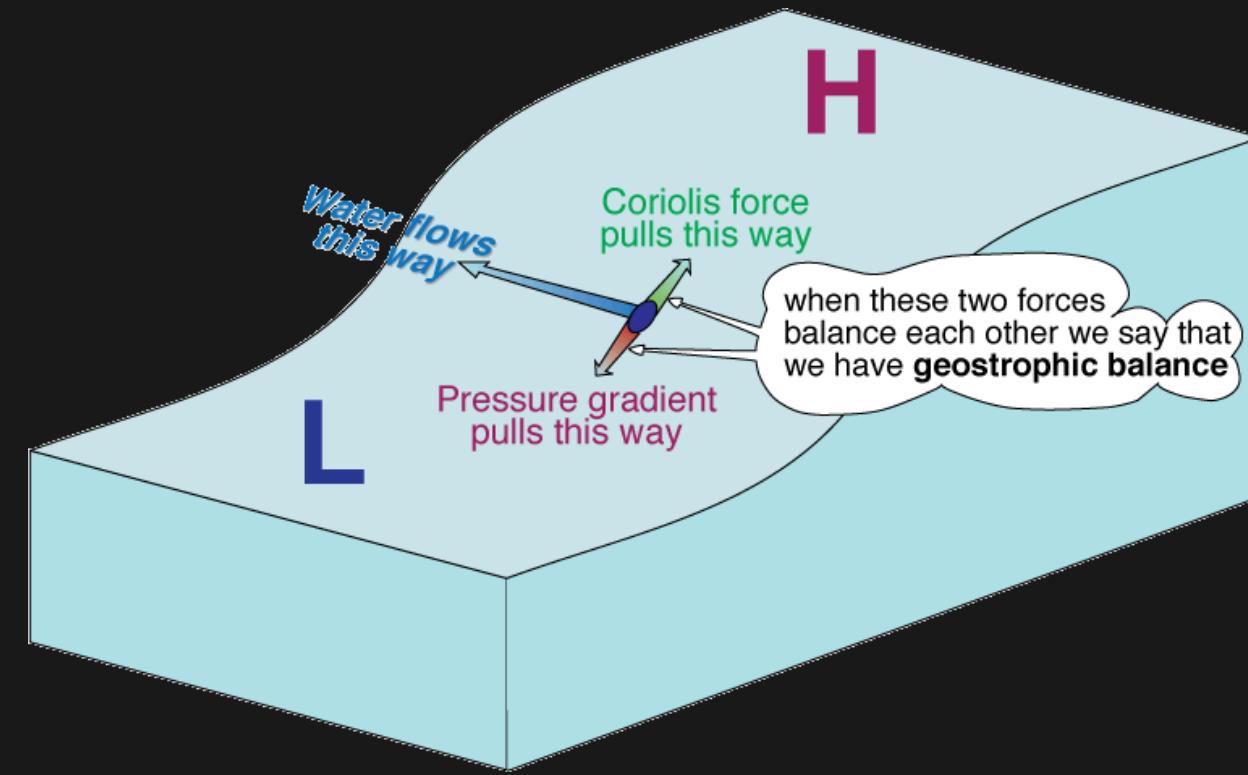
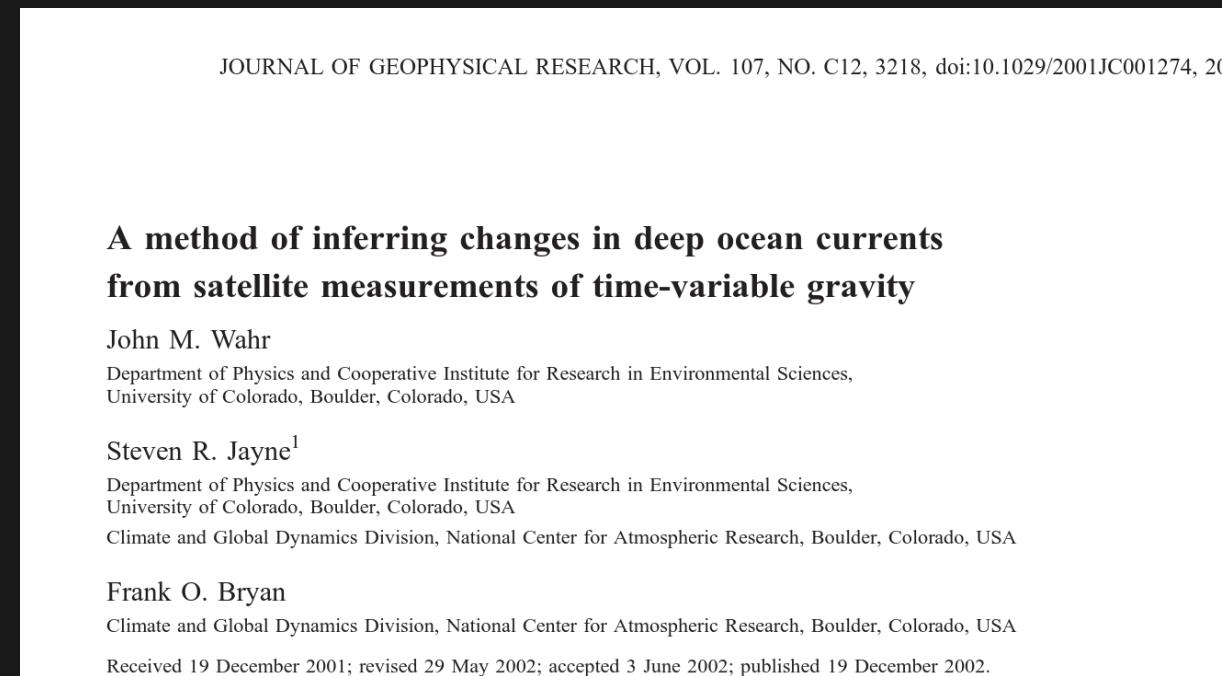
$$\text{Pressure gradient} \\ 2\rho_0 \Omega \times \mathbf{v} \approx -\nabla P \quad (2)$$

Earth rotation vector ↑ velocity

- Assume no accelerations
- cross-product -> latitude dependency



Geostrophic currents from time variable gravity



CC-NC-SA, <https://www.seos-project.eu>

the effects of friction and external forcing are less important,
(1) reduces to the geostrophic approximation:

$$\text{Pressure gradient} \\ 2\rho_0\Omega \times \mathbf{v} \approx -\nabla P \quad (2)$$

Earth rotation vector



- Assume no accelerations
- cross-product \rightarrow latitude dependency
- Velocity is orthogonal to gradient!



What have you learned?

- Adding mass to the ocean
 - Observable in time-variable gravity
 - GRACE (1.5 - 2 mm/yr global mean ocean mass change)
 - Different causes have non-uniform contributions (Ice sheets, hydrology, dam impoundment,..)
 - Regional sea level also influenced by: Self attraction and loading
- Churning the Ocean (currents and wind)
 - Integration of the water column -> OBP
 - OBP changes due to dynamic height differences and column density change
 - Steric sea level: density driven dynamic height changes under the assumption of no OBP change
 - Pressure gradients cause circulation (surface gradient may be different from bottom)



Outlook

- Still many ocean applications unexplored (geostrophic currents!)
- Added value comes from combined data e.g. altimetry, Argo, etc.

