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Massachusetts Institute of Technology

DSpace@MIT

# Extrapolated full-waveform inversion with deep learning

Hongyu Sun<sup>1</sup> and Laurent Demanet<sup>1</sup>

## ABSTRACT

The lack of low-frequency information and a good initial model can seriously affect the success of full-waveform inversion (FWI), due to the inherent cycle skipping problem. Computational low-frequency extrapolation is in principle the most direct way to address this issue. By considering bandwidth extension as a regression problem in machine learning, we have adopted an architecture of convolutional neural network (CNN) to automatically extrapolate the missing low frequencies. The band-limited recordings are the inputs of the CNN, and, in our numerical experiments, a neural network trained from enough samples can predict a reasonable approximation to the seismograms in the unobserved low-frequency band, in phase and in amplitude. The numerical experiments considered are set up on simulated P-wave data. In extrapolated FWI

(EFWI), the low-wavenumber components of the model are determined from the extrapolated low frequencies, before proceeding with a frequency sweep of the band-limited data. The introduced deep-learning method of low-frequency extrapolation shows adequate generalizability for the initialization step of EFWI. Numerical examples show that the neural network trained on several submodels of the Marmousi model is able to predict the low frequencies for the BP 2004 benchmark model. Additionally, the neural network can robustly process seismic data with uncertainties due to the existence of random noise, a poorly known source wavelet, and a different finite-difference scheme in the forward modeling operator. Finally, this approach is not subject to strong assumptions on signals or velocity models of other methods for bandwidth extension and seems to offer a tantalizing solution to the problem of properly initializing FWI.

## INTRODUCTION

Full-waveform inversion (FWI) requires low-frequency data to avoid convergence to a local minimum in cases in which the initial models miss a reasonable representation of the complex structure. However, because of the acquisition limitation in seismic processing, the input data for seismic inversion are typically limited to a band above 3 Hz. With assumptions and approximations to make inferences from tractable but simplified models, geophysicists have started reconstructing the reflectivity spectrum from the band-limited records by signal processing methods. The L<sub>1</sub>-norm minimization (Levy and Fullagar, 1981; Oldenburg et al., 1983), autoregressive modeling (Walker and Ulrych, 1983), and minimum entropy reconstruction (Sacchi et al., 1994) methods have been developed to recover the isolated spikes of seismic recordings. Recently, bandwidth extension to the low-frequency band has attracted the attention of many people in terms of FWI. For exam-

ple, they recover the low frequencies by the envelope of the signal (Wu et al., 2014; Hu et al., 2017) or the inversion of the reflectivity series and convolution with the broadband source wavelet (Wang and Herrmann, 2016; Zhang et al., 2017). However, the low frequencies recovered by these methods are still far from the true low-frequency data. Li and Demanet (2016) attempt to extrapolate the true low-frequency data based on the phase-tracking method (Li and Demanet, 2015). Unlike the explicit parameterization of phases and amplitudes of atomic events, here we propose an approach that can automatically process the raw band-limited records. The deep neural network (DNN) is trained to automatically recover the missing low frequencies from the input band-limited data.

Because of the state-of-the-art performance of machine learning in many fields, geophysicists have begun adapting such ideas in seismic processing and interpretation (Chen et al., 2017; Guitton et al., 2017; Xiong et al., 2018). By learning the probability of salt geobodies

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<sup>1</sup>Massachusetts Institute of Technology, Earth Resources Laboratory, 77 Massachusetts Ave, Cambridge, Massachusetts 02139, USA. E-mail: hongyus@mit.edu (corresponding author); laurent@math.mit.edu.

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being present at any location in a seismic image. Lewis and Vigh (2017) investigate convolutional neural network (CNN) to incorporate the long-wavelength features of the model in the regularization term. Richardson (2018) constructs FWI as recurrent neural networks. Araya-Polo et al. (2018), Wu et al. (2018), and Li et al. (2019) produce layered velocity models from shot gathers with DNN.

Like these authors and many others, we have selected DNN for low-frequency extrapolation due to the increasing community agreement in favor of this method as a reasonable surrogate for a physics-based process (Grzeszczuk et al., 1998; De et al., 2011; Araya-Polo et al., 2017). The universal approximation theorem also indicates that the neural networks can be used to replicate any function up to our desired accuracy if the DNN has sufficient hidden layers and nodes (Hornik et al., 1989). Although training is therefore expected to succeed arbitrarily well, only empirical evidence currently exists for the often-favorable performance of testing a network out of sample. Furthermore, we choose to focus on DNN with a convolutional structure, that is, CNN. The idea behind CNN is to mine the hidden correlations among different frequency components.

In the case of bandwidth extension, the relevant data are the amplitudes and phases of seismic waves, which are dictated by the physics of wave propagation. For training, large volumes of synthetic shot gathers are generated from different models, in a wide band that includes the low frequencies, and the network's parameters are fit to regress the low frequencies of those data from the high frequencies. The window to split the spectrum to low- and high-frequency band should be smooth in the frequency domain. For testing, band-limited (and not otherwise processed) data from a new geophysical scenario are used as input of the network, and the network generates a prediction of the low frequencies. In the synthetic case, validation of the testing step is possible by computing those low frequencies directly from the wave solver.

By now, neural networks have shown their ability to fulfill the task of low-frequency extrapolation. Ovcharenko et al. (2017, 2018, 2019a, 2019b) train neural networks on data generated for random velocity models (Kazei et al., 2019) to predict a single low-frequency from multiple high-frequency data. They treat each shot gather in the frequency domain as a digital image for feature detection and thus require a large number of numerical simulations to synthesize the training data. Jin et al. (2018) and Hu et al. (2019) use a deep inception-based convolutional network to synthesize data at multiple low frequencies. The input of their neural network contains the phase information of the true low frequencies by leveraging the beat tone data (Hu, 2014). In contrast, we design an architecture of CNN to directly handle the band-limited data in the time domain. The proposed architecture can flexibly import one trace or multiple traces of the band-limited shot gather to predict the data in a low-frequency band with sufficient accuracy that it can be used for FWI.

The limitations of neural networks for such signal processing tasks, however, are (1) the unreliability of the prediction when the training set is insufficient and (2) the absence of a physical interpretation for the operations performed by the network. In addition, no theory can currently explain the generalizability of a deep network, that is, the ability to perform nearly as well on testing as on training in a broad range of cases. Even so, the numerical examples indicate that the proposed architecture of CNN enjoys sufficient generalizability to extrapolate the low frequencies of unknown subsurface structures, in a range of numerical experiments.

We demonstrate the reliability of the extrapolated low frequencies to seed frequency-sweep FWI on the Marmousi model and the BP 2004 benchmark model. Two precautions are taken to ensure that trivial deconvolution of a noiseless record (by division by the high-frequency wavelet in the frequency domain) is not an option: (1) add noise to the testing records and (2) for testing, choose a hard band-pass wavelet taken as zero in the low-frequency band. In one numerical experiment involving band-limited data of greater than 0.6 Hz from the BP 2004 model, the inversion results indicate that the predicted low frequencies are adequate to initialize conventional FWI from an uninformative initial model, so that it does not suffer from the otherwise-inherent cycle-skipping at 0.6 Hz. Additionally, the proposed neural network has acceptable robustness to uncertainties due to the existence of noise, a poorly known source wavelet, and different finite-difference (FD) schemes in the forward modeling operator.

This paper is organized as follows. We start by formulating bandwidth extension as a regression problem in machine learning. Next, we introduce the general workflow to predict the low-frequency recordings with CNN. We then study the generalizability and the stability of the proposed architecture in more complex situations. Last, we illustrate the reliability of the extrapolated low frequencies to initialize FWI and analyze the limitations of this method.

## DEEP LEARNING

A neural network defines a mapping  $\mathbf{y} = f(\mathbf{x}, \mathbf{w})$  and learns the value of the parameters  $\mathbf{w}$  that result in a good fit between  $\mathbf{x}$  and  $\mathbf{y}$ . DNNs are typically represented by composing together many different functions to find complex nonlinear relationships. The chain structures are the most common structures in DNNs (Goodfellow et al., 2016):

$$\mathbf{y} = f(\mathbf{x}, \mathbf{w}) = f_L(\dots f_2(f_1(\mathbf{x}))), \quad (1)$$

where  $f_1, f_2$ , and  $f_L$  are the first, the second, and the  $L$ th layer of the network (with their own parameters omitted in this notation). Each  $f_j$  consists of three operations taken in succession: an affine (linear plus constant) transformation, a batch normalization (multiplication by a scalar chosen adaptively), and the component-wise application of a nonlinear activation function. It is the nonlinearity of the activation function that enables the neural network to be a universal function approximator. The overall length  $L$  of the chain gives the depth of the deep learning model. The final layer is the output layer, which defines the size and type of the output data. The training sets specify directly what the output layer must do at each point  $\mathbf{x}$  and constrain but do not specify the behavior of the other hidden layers. Rectified activation units are essential for the recent success of DNNs because they can accelerate convergence of the training procedure. Our numerical experiments show that, for bandwidth extension, the parametric rectified linear unit (PReLU) (He et al., 2015) works better than the rectified linear unit (ReLU). The formula for PReLU is

$$g(\alpha, \mathbf{y}) = \begin{cases} \alpha \mathbf{y}, & \text{if } \mathbf{y} < 0 \\ \mathbf{y}, & \text{if } \mathbf{y} \geq 0 \end{cases}, \quad (2)$$

where  $\alpha$  is also a learnt parameter and would be adaptively updated for each rectifier during training.

Unlike the classification problem that trains the DNNs to produce discrete labels, the regression problem trains the DNNs for the prediction of continuous-valued outputs. It evaluates the performance of the model by means of the mean-squared error (MSE) of the predicted outputs  $f(\mathbf{x}_i; \mathbf{w})$  versus the actual outputs  $\mathbf{y}_i$ :

$$J(\mathbf{w}) = \frac{1}{m} \sum_{i=1}^m L(\mathbf{y}_i, f(\mathbf{x}_i, \mathbf{w})), \quad (3)$$

where the loss  $L$  is the squared error between the true low frequencies and the estimated outputs of the neural networks. The cost function  $J$  is here minimized over  $\mathbf{w}$  by a stochastic gradient descent (SGD) algorithm, where each gradient is computed from a mini-batch, that is, a subset in a disjoint randomized partition of the training set. Each gradient evaluation is called an iteration, whereas the full pass of the training algorithm over the entire training set using minibatches is an epoch. The learning rate  $\eta$  (step size) is a key parameter for deep learning and must be fine-tuned. The gradients  $\partial J(\mathbf{w}^t)/\partial \mathbf{w}$  of the neural networks are calculated by the backpropagation method (Goodfellow et al., 2016).

CNN is an overwhelmingly popular architecture of DNN to extract spatial features in image processing, and it is the choice that we make in this paper. In this case, the matrix-vector multiplication in each  $f_j$  is a convolution. In addition, imposing local connections and weight sharing can exploit the local correlation and global features of the input image. CNNs are normally designed to deal with the image classification problem. For bandwidth extension, the data to be learned are the time-domain seismic signals, so we directly consider the amplitude at each sampling point as the pixel value image to be used as input of the CNN.

Recall that CNN involves stacks of a convolutional layer, followed by a PReLU layer, and a batch normalization layer. The filter number in each convolutional layer determines the dimensionality of the feature map or the channel of its output. Each output channel of the convolutional layer is obtained by convolving the channel of the previous layer with one filter, summing and adding a bias term. The batch normalization layer can speed up training of CNNs and reduce the sensitivity to network initialization by normalizing each input channel across a minibatch (Ioffe and Szegedy, 2015). We also leave the pooling layer out, although it is typically used in the conventional architecture of CNNs.

An essential hyperparameter for low-frequency extrapolation with deep learning is the receptive field of a neuron. It is the local region of the input volume that affects the response of this neuron — otherwise known as the domain of dependence. The spatial extent of this connectivity is related to the filter size. Unlike the small filter size commonly used in the image classification problem, we directly use a large filter in the convolutional layer to increase the receptive field of the CNN quickly with depth. The large filter size gives the neural network enough freedom to reconstruct the long-wavelength information.

The architecture of our neural network (Figure 1) is a feed-forward stack of five sequential combinations of the convolution, PReLU, and batch normalization layers, followed by one fully connected layer that outputs continuous-valued

amplitude of the time-domain signal in the low-frequency band. The first convolutional layer filters the  $nt \times 1$  input time series with 128 kernels of size  $200 \times 1 \times 1$  where  $nt$  is the number of time steps. The second convolutional layer has 64 kernels of size  $200 \times 1 \times 128$  connected to the normalized outputs of the first convolutional layer. The third convolutional layer has 128 kernels of size  $200 \times 1 \times 64$ . The fourth convolutional layer has 64 kernels of size  $200 \times 1 \times 128$ , and the fifth convolutional layer has 32 kernels of size  $200 \times 1 \times 64$ . The last layer is fully connected, taking features from the last convolutional layer as input in a vector form of length  $nt \times 32$ . The stride of the convolution is one, and zero-padding is used to make the output length of each convolution layer the same as its input. Additionally, a dropout layer (Srivastava et al., 2014) with a probability of 50% is added after the first convolution layer to reduce the generalization error.

We use CNN in the context of supervised learning, that is, inference of  $\mathbf{y}_i$  from  $\mathbf{x}_i$ . We need to first train the CNN from a large number of samples  $(\mathbf{x}_i, \mathbf{y}_i)$  to determine the coefficients of the network, and then we use the network for testing on new  $\mathbf{x}_i$ . In statistical learning theory, the generalization error is the difference between the expected and empirical error, where the expectation runs over a continuous probability distribution on the  $\mathbf{x}_i$ . This generalization error can be approximated by the difference between the errors on the training and test sets.

The object of this paper is that  $\mathbf{x}_i$  can be taken to be seismograms band limited to the high frequencies and  $\mathbf{y}_i$  can be the same seismograms in the low-frequency band. Generating training samples means collecting or synthesizing seismogram data from a variety of geophysical models, which enter as space-varying elastic coefficients in a wave equation. For the purpose of good generalization (a small generalization error), the models used to create the large training sets should be able to represent many subsurface structures, including different types of reflectors and diffractors, so we can find a representative set of parameters to handle data from different scenarios or regions. The performance of the neural network is sensitive to the architecture and the hyperparameters, so we must design them carefully. Next, we illustrate the specific choice of hyperparameters for bandwidth extension, along with numerical examples involving synthetic data from community models.

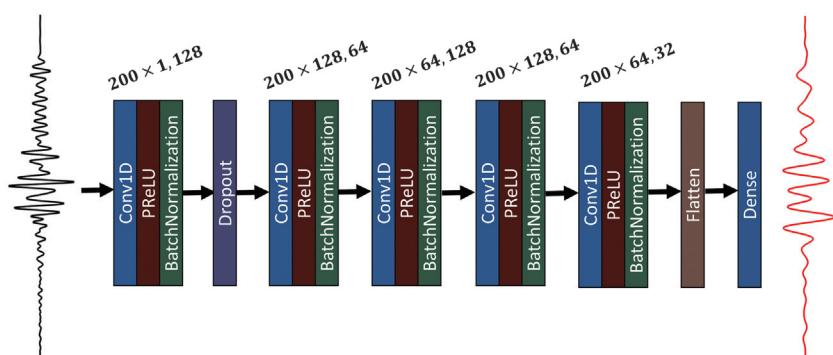


Figure 1. An illustration of the architecture of our CNN to extrapolate the low-frequency data from the band-limited data in the time domain trace by trace. The architecture is a feed-forward stack of five sequential combinations of the convolution, PReLU, and batch normalization layers, followed by one fully connected layer that outputs continuous-valued amplitude of the time-domain signal in the low-frequency band. The network's input is a 1D band-limited recording of length  $nt$  where  $nt$  is the number of time steps. The size and number of filters are labeled on the top of each convolutional layer.

## NUMERICAL EXAMPLES

In this section, we demonstrate the reliability of extrapolated FWI with CNN (EFWI-CNN) in three parts. In the first part, we show CNN's ability to extrapolate low-frequency (0.1–5.0 Hz) from band-limited data (5.0–20.0 Hz) on the Marmousi model (Figure 2). In the second part, we verify the robustness of the method with uncertainties in the seismic data due to the existence of noise, a different finite difference scheme, and a poorly known source wavelet. In the last part, we perform EFWI-CNN on the Marmousi model and the BP 2004 benchmark model (Billette and Brandsberg-Dahl, 2005), by first using the extrapolated low frequencies to synthesize the low-wavenumber background velocity model. Then, we compare the inversion results with the band-limited data in three cases that respectively start FWI from an uninformed initial model, the low-wavenumber background model created from the extrapolated low frequencies, and the low-wavenumber background model created from the true low frequencies.

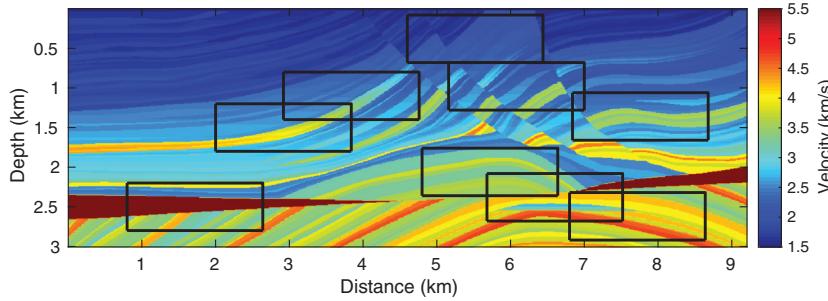


Figure 2. The nine training models randomly extracted from the Marmousi velocity model to collect the training data set. The test models are the Marmousi model and the BP 2004 benchmark model. A water layer with 300 m depth is added to the top of these training models and Marmousi model. We use the same training models to extrapolate the low frequencies on both test models.

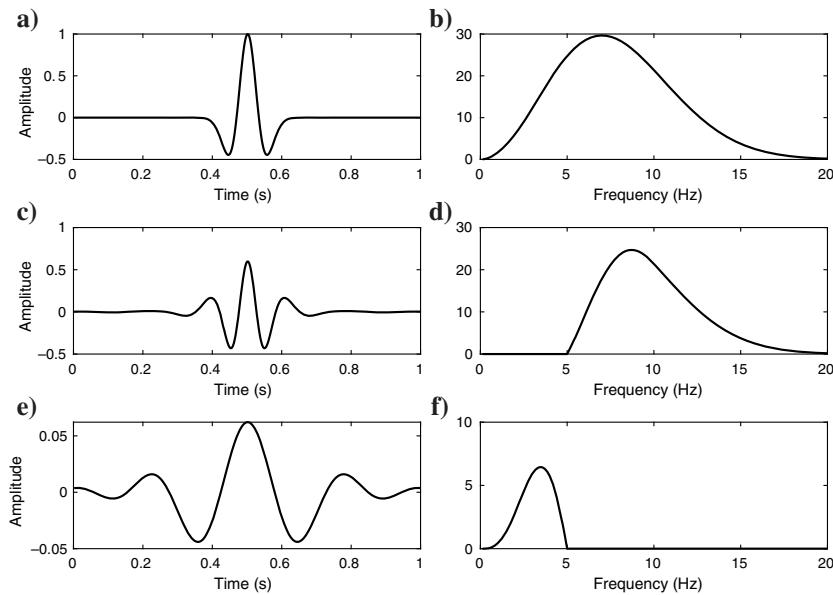


Figure 3. (a) The Ricker wavelet with 7 Hz dominant and (b) its amplitude spectrum. (c) The high-frequency wavelet band-passed from (a) and (d) its amplitude spectrum. (e) The low-frequency wavelet band-passed from (a) and (f) its amplitude spectrum.

## Low-frequency extrapolation

Following our previous work (Sun and Demanet, 2018, 2019), the true unknown velocity model for FWI is referred to as the test model because it is used to collect the test data set in deep learning. To collect the training data set, we create training models by randomly selecting nine parts of the Marmousi model (Figure 2) with a different structure but the same number of grid points  $166 \times 461$ . We also downsample the original model to  $166 \times 461$  pixels as the test model. We find that the randomized models produced in this manner are realistic enough to demonstrate the generalization of the neural network if the structures of the submodels are diversified enough.

In this example, we have the following processing steps to collect each sample (i.e., a shot record) in the training and test data sets.

- The acquisition geometry of forward modeling on each model is the same. It consists of 30 sources and 461 receivers evenly spaced at the surface. We consider each time series or trace as one sample in the data set, so we have 124,470 training samples and 13,830 test samples for the test model in total.
- We use a fourth order in space and second order in time FD modeling method with PML to solve the 2D acoustic wave equation in the time domain, to generate the synthetic shot gathers of the training and test data sets. The sampling interval and the total recording time are 2 ms and 5 s, respectively.
- We use a Ricker wavelet with dominant frequency of 7 Hz to synthesize the full-band seismic recordings. Then, the data below 5 Hz and above 5 Hz are split to synthesize the output and input of the neural networks, respectively (Figure 3). The low- and high-frequency data are obtained by a sharp windowing of the same trace.

In this example, we train the network with the Adam optimizer and use a minibatch of 64 samples at each iteration (implemented with Tensorflow [Abadi et al., 2015] and Keras [Chollet, 2015]). The initial learning rate and forgetting rate of the Adam are the same as the original paper (Kingma and Ba, 2014). The initial value of the bias is zero. The weight initialization is via the Glorot uniform initializer (Glorot and Bengio, 2010). It randomly initializes the weights from a truncated normal distribution centered on zero with the standard deviation  $\sqrt{2/n_1 + n_2}$  where  $n_1$  and  $n_2$  are the numbers of input and output units in the weight tensor, respectively.

The training process of the 20 epochs is shown in Figure 4. The training and test losses decay with the training steps, which indicates that our neural network is not overfitting. We test the performance of the neural networks by feeding the

band-limited data in the test set into the pretrained neural networks and obtaining the extrapolated low frequencies on the Marmousi model. Figure 5 compares the shot gather between the band-limited data (5.0–20.0 Hz), extrapolated, and true low frequencies (0.1–5.0 Hz) where the source is located at the horizontal distance  $x = 2.94$  km on the Marmousi model. The extrapolated results in Figure 5b show that the proposed neural network can accurately predict the recordings in the low-frequency band, which are totally missing before the test. Figure 6 compares two individual seismograms in Figure 5b where the receivers are located at the horizontal distances of  $x = 2.82$  km and  $x = 2.92$  km, respectively. The extrapolated low-frequency data match the true recordings well. Then, we combine the extrapolated low frequencies with the band-limited data and compare the amplitude spectrum of the full-band data with the extrapolated and true low frequencies. The amplitude and phase spectrum comparison of the single trace where the receiver is located at  $x = 2.92$  km (Figure 7) clearly shows that the neural networks can capture the relationship between low- and high-frequency components constrained by the wave equation.

Figure 8 shows the low-frequency extrapolation without direct waves. The direct waves are muted from the full-band shot gathers with a smooth time window before splitting into the band-limited recordings and the low frequencies. The low frequencies of reflections are recovered without the existence of the direct waves. Therefore, the neural network is robust with the presence of muting.

### Uncertainty analysis

With a view toward dealing with complex field data, we investigate the stability of the neural network's predictive performance under three kinds of discrepancies, or uncertainties, between the training and the test: additive noise, different FD operator in the forward modeling, and different source wavelet. In each case, we compare the extrapolation accuracy with the reference in Figure 5, where training and testing are set up the same way (noiseless band-limited data, finite difference operator with second order in time and fourth order in space, and the Ricker wavelet with a 7 Hz dominant frequency). The root-mean-square error (rms error) between data with extrapolated and true low frequencies of the 30 shot gathers in Figure 5 is  $2.1304 \times 10^{-4}$ .

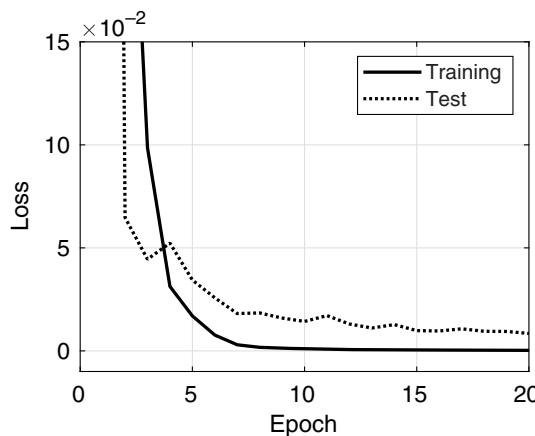


Figure 4. The learning curves. Training and test losses decay with the training steps.

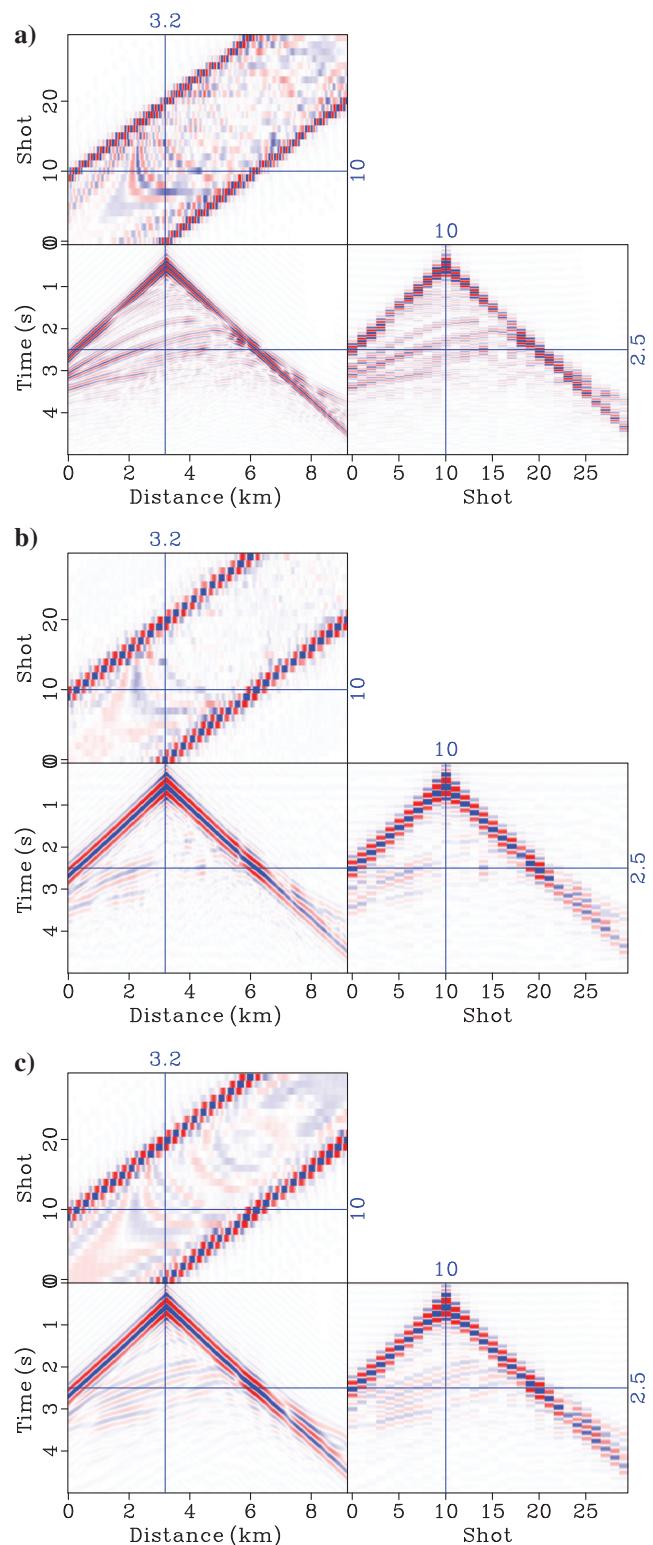


Figure 5. The extrapolation result on the Marmousi model: comparison among the (a) band-limited recordings (5.0–20.0 Hz), (b) predicted, and (c) true low-frequency recordings (0.1–5.0 Hz). The band-limited data in (a) are the inputs of CNNs to predict the low frequencies in (b).

In the first case, the neural network is expected to extrapolate the low frequencies from the noisy band-limited data. We add 20% additive Gaussian noise to the band-limited data in the test data set and 30% additive Gaussian noise to the band-limited data in the training data set. The low frequencies in the training set are noiseless as before. Even though noise will disturb the neural network to find the correct mapping between the band-limited data with their low frequencies, Figure 9 shows that the proposed neural network can still successfully extrapolate the low frequencies of the main reflections. The rms error between data with extrapolated and true low frequencies of the 30 shot gathers in Figure 9 is  $2.4156 \times 10^{-4}$ . The neural network is able to perform extrapolation as well as denoising. Incidentally, we make the (unsurprising) observation that CNN has the potential for the denoising of seismic data.

Another challenge of FWI is that the observed and calculated data can come from different wave-propagation schemes. For example, under the control of different numerical dispersion curves, the phase velocity would have different behavior if we used different finite difference operators to simulate the observed and calculated data. Therefore, it is necessary to study the influence of different discretization, or other details of the simulation, on the accuracy of low-frequency extrapolation. In our case, the shot gathers in the test data set are simulated with a sixth-order spatial FD operator, but the neural network is trained on the samples simulated with a fourth-order spatial FD operator. The extrapolation result in Figure 10b shows that the neural network trained on the fourth-order operator is able to extrapolate the low frequencies of the band-limited data collected with the sixth-order operator. In this case,

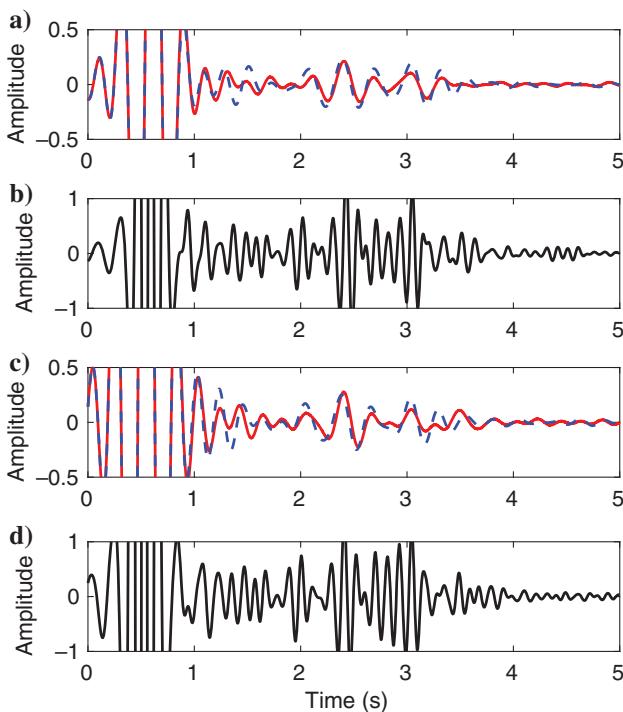


Figure 6. The extrapolation result on the Marmousi model: comparison among the predicted (the red line), the true (the blue dashed line) recording in the low-frequency band (0.1–5.0 Hz), and the band-limited recording (the black line) (5.0–20.0 Hz) at the horizontal distance (a and b)  $x = 2.82$  km and (c and d)  $x = 2.92$  km.

the rms error between data with extrapolated (Figure 10b) and true (Figure 10c) low frequencies of the 30 shot gathers is  $2.2248 \times 10^{-4}$ . The neural network appears to be stable with respect to mild modifications to the forward modeling operator, at least in the examples tried.

Another uncertainty is the unknown source wavelet. To check the extrapolation capability of the neural network in the context of data excited by an unknown source wavelet, we train the neural network with a 7 Hz Ricker wavelet but we test it with an Ormsby wavelet. The four corner frequencies of the Ormsby wavelet are 0.2, 1.5, 8, and 14 Hz, respectively. Figure 11 shows that the neural network trained on the data from the 7 Hz Ricker wavelet source wavelet is able to extrapolate the data synthesized with the Ormsby source wavelet. However, the recovery of the amplitude is much poorer than the phase. The rms error between data with extrapolated and true low frequencies of the 30 shot gathers in Figure 11 is increased to  $1.1717 \times 10^{-3}$ . The commonplace issue of the source wavelet being unknown or poorly known in FWI has seemingly little effect on the performance of the proposed neural network to extrapolate the phase of low-frequency data, at least in the examples tried.

Even though all of the uncertain factors hurt the accuracy of extrapolated low frequencies to some extent, the CNN's prediction has a degree of robustness that surprised us. All of the extrapolation results in the above numerical examples can be further improved by increasing the diversity of the training data set because subjecting the network to a broader range of scenarios can fundamentally reduce the generalization error of the deep learning predictor (e.g., we can simulate the training data set with multiple kinds of source wavelets and FD operators).

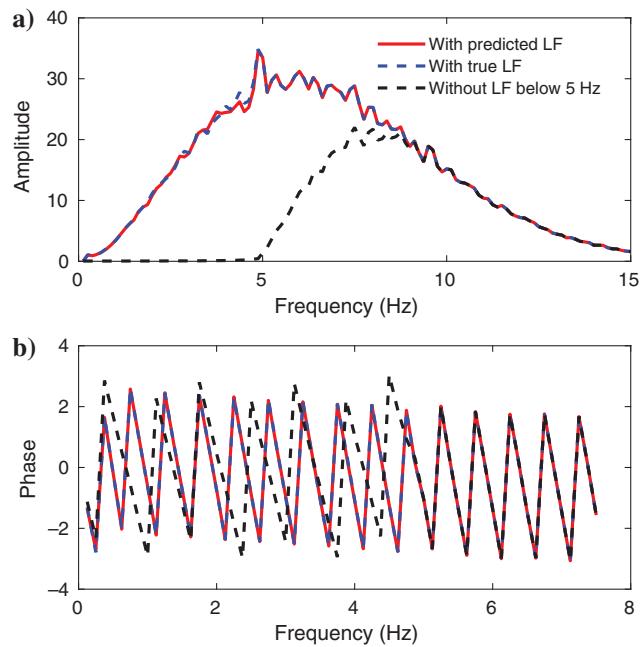


Figure 7. The extrapolation result on the Marmousi model: comparison of (a) the amplitude spectrum and (b) the phase spectrum at  $x = 2.92$  km among the band-limited recording (5.0–20.0 Hz), the recording (0.1–20.0 Hz) with the true and predicted low frequencies (0.1–5.0 Hz).

## Extrapolated FWI: Marmousi model

In this example, we construct the low-wavenumber velocity model for the Marmousi model, by leveraging the extrapolated low-frequency data (Figure 5b) to solve the cycle-skipping problem for FWI on the band-limited data. The objective function of the inversion is formulated as the least-squares misfit between the observed and calculated data in the time domain. Starting from an initial model in Figure 12b, we use the L-BFGS (a limited-memory variant of the quasi-Newton BFGS method named after its discoverers Broyden, Fletcher, Goldfarb, and Shanno; see Nocedal and Wright, 2006) optimization method to update the model gradually (for the implementation in this paper, see Hewett and Demanet, 2013). To help the gradient-based iterative inversion method avoid falling into local minima, we also perform the inversion from the lowest frequency that the data contained to successively higher frequencies. Additionally, an accurate wavelet is assumed to be known for FWI.

With this inversion scheme, we test the reliability of the extrapolated low frequencies (Figure 5b) on the Marmousi model (Figure 12a). The velocity structure of the initial model is far from the true model. The true model was not used in the training stage. The acquisition geometry and source wavelet are the same as in the example in the previous section. The observed data below 5.0 Hz are totally missing. Therefore, we first use the band-limited data in 5.0–20.0 Hz to recover the low frequencies in 0.1–5.0 Hz and then use the low frequencies to invert the low-wavenumber velocity model for the band-limited FWI. Figure 13 compares the inverted models from FWI using the true and extrapolated 0.5–3.0 Hz low-frequency data. Because the low-frequency extrapolation accuracy of reflections after 4.0 s is limited (as seen in Figure 5b), the low-wavenumber model constructed from the extrapolated low frequencies has lower resolution in the deeper section compared with that from the true low frequencies. However, both models capture the low-wavenumber information of the Marmousi model. These models are used as the starting models for FWI on the band-limited data.

Figure 14 compares the inverted models from FWI using the band-limited data (5–15 Hz) with different starting models. The resulting model in Figure 14b starts from the low-wavenumber model constructed from the extrapolated low frequencies (Figure 13b), which is almost the same as the one from the true low frequencies (Figure 14a). Because the highest frequency component in the low-frequency band is 3 Hz when we invert the starting model, both inversion results have a slight cycle-skipping phenomenon. However, Figure 14c performs band-limited FWI with the linear initial model and it shows a much more pronounced effect of cycle-skipping. We cannot find much meaningful information about the subsurface structure if the band-limited inversion starts at 5 Hz from a linear initial model (Figure 12b).

Figure 15 compares the velocity profile among the resulting models in Figure 14 (the initial and true velocity models) at the horizontal locations of  $x = 3$  km,  $x = 5$  km, and  $x = 7$  km. The final inversion result started from the extrapolated low frequencies gives us almost the same model as the true low frequencies, which illustrates that the extrapolated low-frequency data are reliable enough to provide an adequate low-wavenumber velocity model. However, both inversion workflows have difficulty in the recovery of velocity

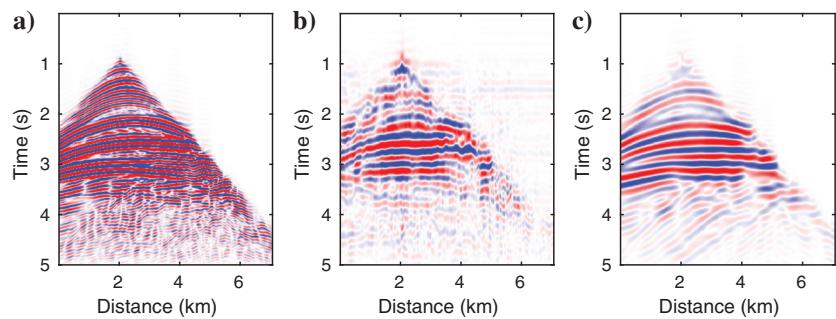


Figure 8. Low-frequency extrapolation without direct waves: comparison among the (a) band-limited recordings (5.0–20.0 Hz), (b) predicted, and (c) true low-frequency recordings (0.1–5.0 Hz) on the Marmousi model. The direct waves are muted from the full-band shot gathers with a smooth time window before splitting into the band-limited recordings and the low frequencies. The extrapolation is robust with the presence of muting.

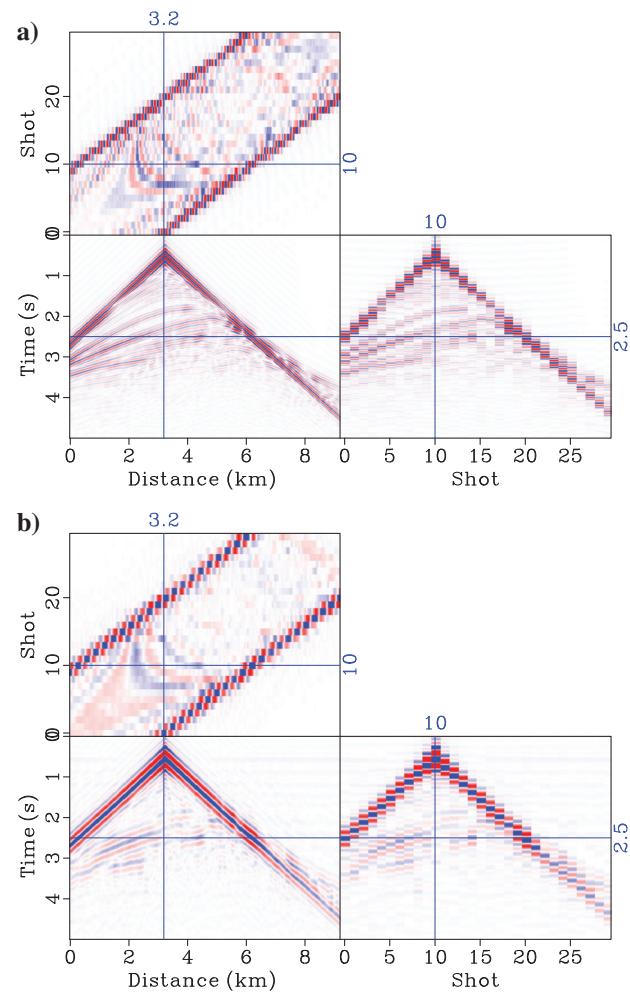


Figure 9. Noise robustness: comparison between the (a) band-limited recordings (5.0–20.0 Hz) and (b) predicted low-frequency recordings (0.1–5.0 Hz) on the Marmousi model. We add 20% additive Gaussian noise to the band-limited data in the test data set and 30% additive Gaussian noise to the band-limited data in the training data set. Even though noise will disturb the neural network find the correct mapping between the band-limited data with their low frequencies, the proposed neural network can still extrapolate the low frequencies of the main reflections. The neural network is able to perform extrapolation as well as denoising.

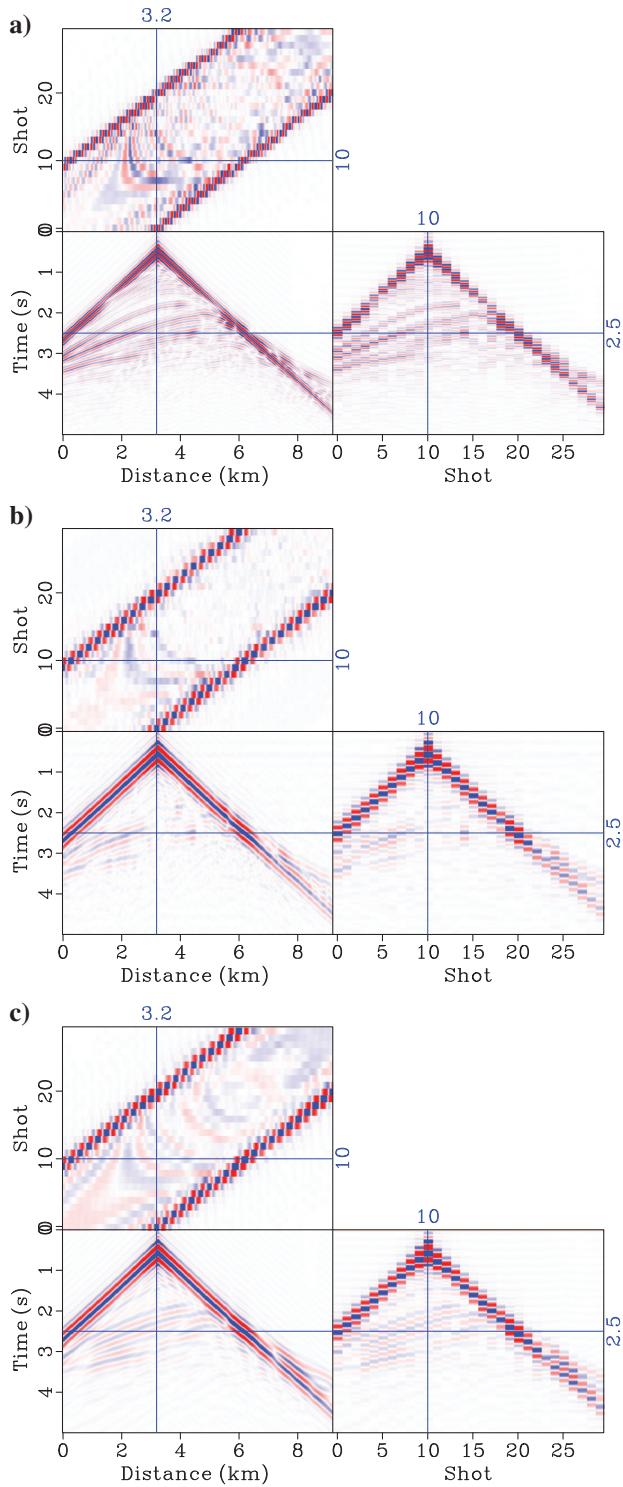


Figure 10. Forward modeling operator robustness: comparison among the (a) band-limited recordings (5.0–20.0 Hz), (b) predicted, and (c) true low-frequency recordings (0.1–5.0 Hz) on the Marmousi model. The shot gather in the test data set is simulated with the sixth-order operator, whereas the neural network is trained with the samples simulated with the fourth-order operator. The extrapolation result in (b) shows that the neural network trained on the fourth-order FD operator can extrapolate the low frequencies of the band-limited data coming from the sixth-order operator. The neural network is stable with the variance of the forward modeling operator.

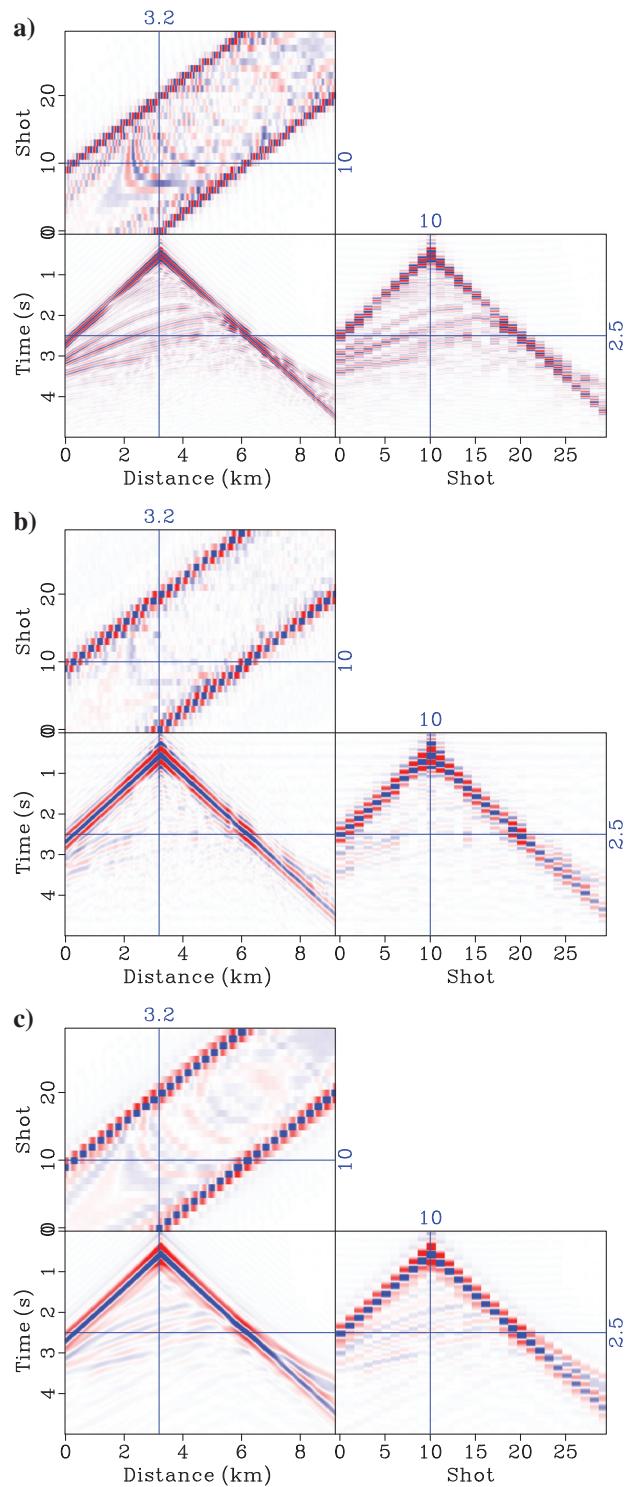


Figure 11. Unknown source wavelet robustness: comparison among the (a) band-limited recordings (5.0–20.0 Hz), (b) predicted, and (c) true low-frequency recordings (0.1–5.0 Hz) on the Marmousi model. In this case, we use an Ormsby wavelet with the four corner frequencies 0.2, 1.5, 8.0, and 14.0 Hz to synthesize the output and input of neural network for samples in the test data set. The result in (b) shows that the neural network trained on the data from 7 Hz Ricker wavelet is able to extrapolate the data synthesized with an Ormsby wavelet. However, the recovery of the amplitude is poorer than the phase.

structures below 2 km depth. Because the velocity model in Figure 14c has fallen into a local minimum, the inversion cannot converge to the true model in the subsequent iterations. The inversion results can be further improved by a multiscale FWI method (Bunks et al., 1995) with a frequency band selection method (Sirgue and Pratt, 2004).

### Extrapolated FWI: BP model

In deep learning, it is essential to estimate the generalization error of the proposed neural network for understanding its performance. Clearly, the intent is not to compute the generalization error exactly because it involves an expectation over an unspecified probability distribution. Nevertheless, we can access the test error in the framework of synthetic shot gathers; hence, we can use the test error minus the training error as a good proxy for the generalization error. For the purpose of assessing whether the network can truly generalize “out of sample” (when the training and testing geophysical models are very different), we train it with the samples collected from the submodels of Marmousi, but we test it on the BP 2004 benchmark model (Figure 16). With the extrapolated low-frequency data predicted by the neural network trained on the submodels of Marmousi, we perform the EFWI-CNN on the BP 2004 benchmark model (Figure 16).

To reduce the computation burden, we downsample the BP 2004 benchmark model to  $80 \times 450$  grid points with a grid interval of 150 m. It is challenging for FWI to use only the band-limited data to invert the shallow salt overhangs and the salt body with steeply dipping flanks in the BP model. Numerical examples show that, starting from very erroneous initial model (shown later), the highest starting frequency to avoid cycle-skipping on this model is 0.3 Hz. Therefore, we should extrapolate the band-limited data to at least 0.3 Hz to invert the BP model successfully.

In this example, we use a 7 Hz Ricker wavelet as the source to simulate the full band seismic records on the training models (submodels of Marmousi) and the test model (BP model). Here, the same training models with  $166 \times 461$  pixels from Marmousi are used but the grid spacing is increased to 150 m to be consistent with the test model (the BP model). The sampling interval and the total recording time are 5 ms and 10 s, respectively. To collect the input of the CNN, a high-pass filter where the low-frequency band (0.1–0.5 Hz) is exactly zero is applied to the full-band seismic data. The band-limited data (0.6–20.0 Hz) are fed into the proposed CNN model to extrapolate the low-frequency data in 0.1–0.5 Hz trace by trace.

Figure 17 shows the learning curves of training process across the 20 epochs. Figure 18 compares the extrapolation result of one shot gather on the

BP model where the shot is located at 31.95 km. The neural network can recover the low frequencies of reflections with some degree of accuracy. Even though the information contained in the data collected on Marmousi is physically unlike that of the salt dome model, the

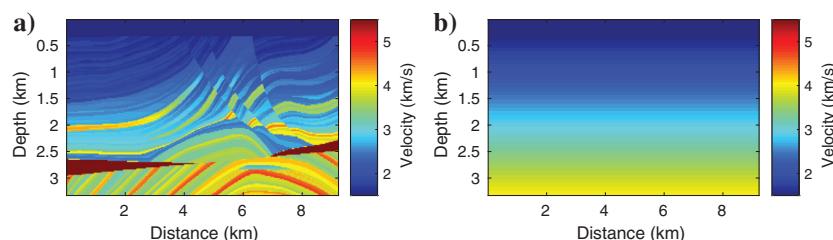


Figure 12. (a) The Marmousi model (the true model in FWI and the test model in deep learning) and (b) the initial model for FWI.

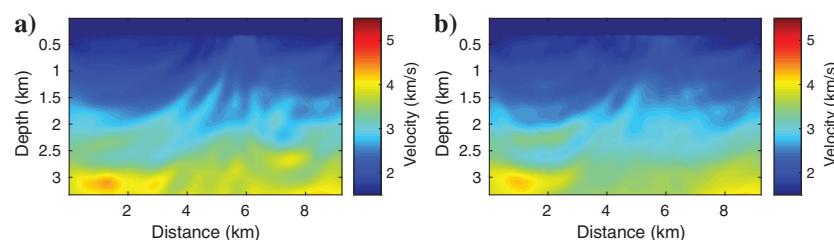


Figure 13. Comparison between the inverted low-wavenumber models using the (a) true and (b) extrapolated 0.5–3.0 Hz low frequencies. The model constructed from the extrapolated low frequencies has lower resolution in the deeper section compared to the model from the true low frequencies because the extrapolation accuracy of the deeper reflections is poor. However, both models capture the low-wavenumber information of the Marmousi model.

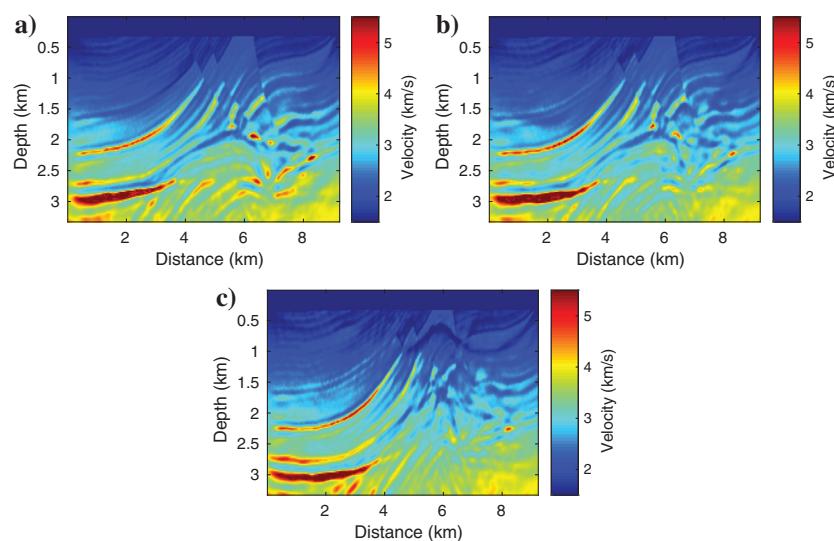


Figure 14. Comparison among the inverted models from FWI using the band-limited data (5.0–15.0 Hz). (a) The resulting model is started from the low-wavenumber velocity model constructed with the true low frequencies in Figure 13a. (b) The resulting model is started from the low-wavenumber velocity model constructed with the extrapolated low frequencies in Figure 13b. (c) The resulting model is started from the initial model in Figure 12b.

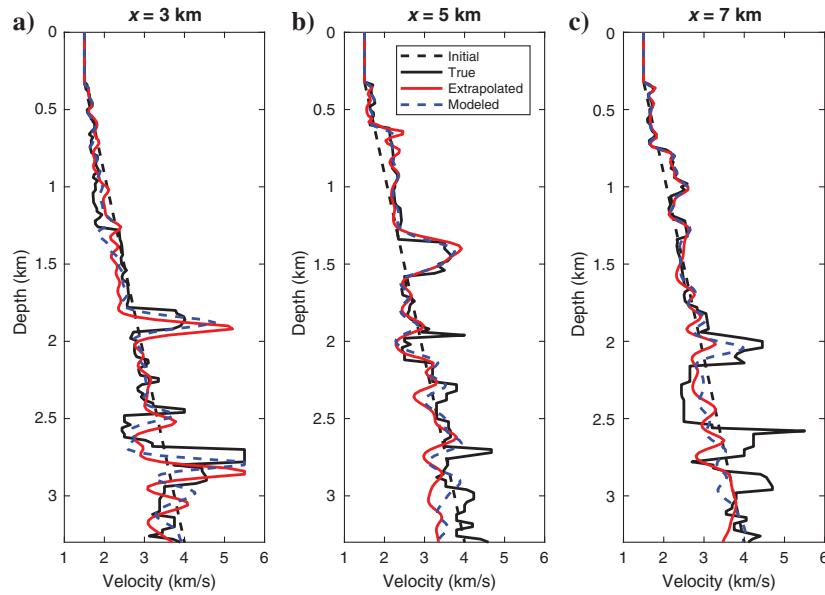


Figure 15. Comparison of velocity profiles among the initial model (the black dashed line), the true model (the black line), and the resulting models started from the low-wavenumber models constructed with extrapolated (the red line) and true (the blue dashed line) low frequencies at the horizontal locations of (a)  $x = 3$  km, (b)  $x = 5$  km, and (c)  $x = 7$  km.

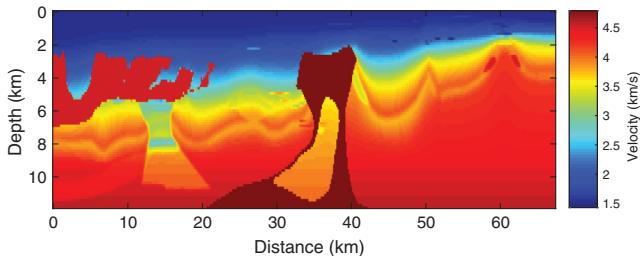


Figure 16. The 2004 BP benchmark velocity model used to collect the test data set for studying the generalizability of the proposed neural network. This model is the true model in extrapolated FWI.

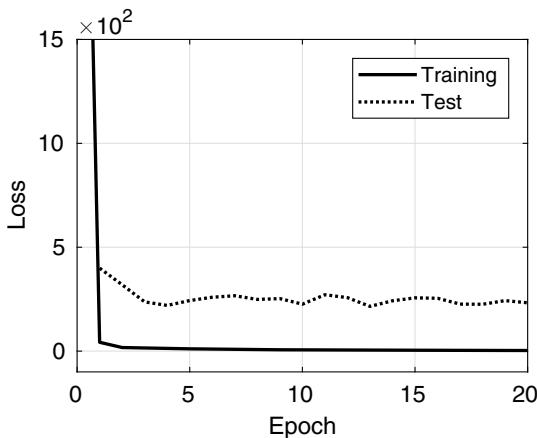


Figure 17. The learning curves.

pretrained neural network can successfully find an approximation of their low frequencies based on the band-limited inputs.

The extrapolated low-frequency data are used to invert the low-wavenumber velocity model with the conventional FWI method. We observe that the accuracy of extrapolated low frequencies decreases as the offset increases, so we limit the maximum offset to 12 km. Starting from the initial model (Figure 19a), Figure 19b and 19c shows the low-wavenumber inverted models using 0.3 Hz extrapolated data and 0.3 Hz true data, respectively. Compared with the initial model, the resulting model using the 0.3 Hz extrapolated data reveals the positions of the high- and low-velocity anomalies, which is almost the same as that of true data. The low-wavenumber background velocity models can initialize the frequency-sweep FWI in the right basin of attraction.

Figure 20 compares the inverted models from FWI using 0.6–2.4 Hz band-limited data, starting from the respective low-wavenumber models in the previous figure. In Figure 20a, the resulting model starts from the original initial model. In Figure 20b, the resulting model starts from the inverted low-wavenumber velocity model using 0.3 Hz extrapolated data. In Figure 20c, the resulting model starts from the inverted low-wavenumber velocity model using 0.3 Hz true data. With the low-wavenumber velocity model, FWI can find the accurate velocity boundaries by exploring band-limited data. However, the inversion falls in a local minimum with only the band-limited data. The low frequencies extrapolated with deep learning are reliable enough to overcome the cycle-skipping problem on the BP model, even though the training data set is ignorant of the particular subsurface structure of BP — salt bodies. Therefore, the neural network approach has the potential to deal favorably with real field data.

So far, the experiments on BP 2004 have assumed that data are available in a band starting at 0.6 Hz. We now study the performance of EFWI-CNN when this band starts at a frequency higher than 0.6 Hz. We still start the frequency-sweep FWI with 0.3 Hz extrapolated data, and the highest frequency is still fixed at 2.4 Hz. Figures 21 and 22 compare the conventional FWI and EFWI-CNN results with data band limited above 0.9 and 1.2 Hz, respectively. With the increase of the lowest frequency of band-limited data, Figure 23 compares the quality of the inverted models at each iteration for FWI using full-band data, EFWI-CNN, and FWI using only the band-limited data. The norm of the relative model error is used to evaluate the model quality, as (Brossier et al., 2009)

$$mq = \frac{1}{N} \left\| \frac{\mathbf{m}_{\text{inv}} - \mathbf{m}_{\text{true}}}{\mathbf{m}_{\text{true}}} \right\|_2, \quad (4)$$

where  $\mathbf{m}_{\text{inv}}$  and  $\mathbf{m}_{\text{true}}$  are the inverted model and the true model, respectively, and  $N$  denotes the number of grid point in the computational domain. The performance of EFWI-CNN of course decreases with the increase of the lowest frequency of the band-limited data because this leads to more extrapolated data involved in the frequency-sweep FWI. The more iterations of FWI with the extrapolated

lated data, the more errors the inverted model will have before exploring the true band-limited data. Overfitting of the unfavorable extrapolated data makes the inversion worse after several iterations with the extrapolated data. However, EFWI-CNN is still superior to using FWI with only band-limited data. We observe that EFWI-CNN with the current architecture still helps to reduce the inverted model error on the BP model when the lowest available frequency is as high as 1.2 Hz.

Finally, we encountered a puzzling numerical phenomenon: The accuracy of the extrapolated data at the single frequency 0.3 Hz depends very weakly on the band in which data are available,

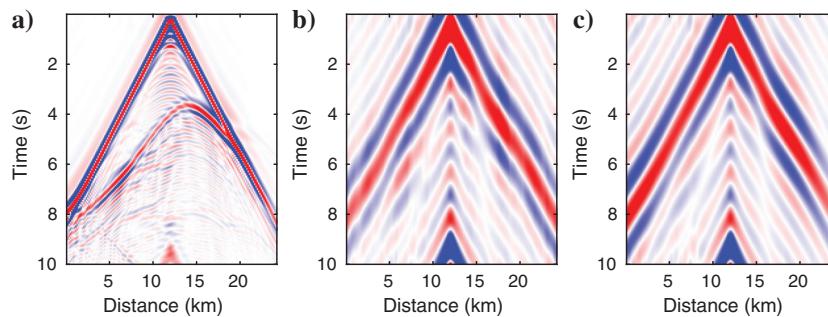


Figure 18. The extrapolation result on the BP model: comparison among the (a) band-limited recordings (0.6–20.0 Hz), (b) predicted, and (c) true low-frequency recordings (0.1–0.5 Hz). The neural network trained on the Marmousi submodels can recover the low frequencies synthesized from the BP model, which illustrates that the proposed neural network can generalize well.

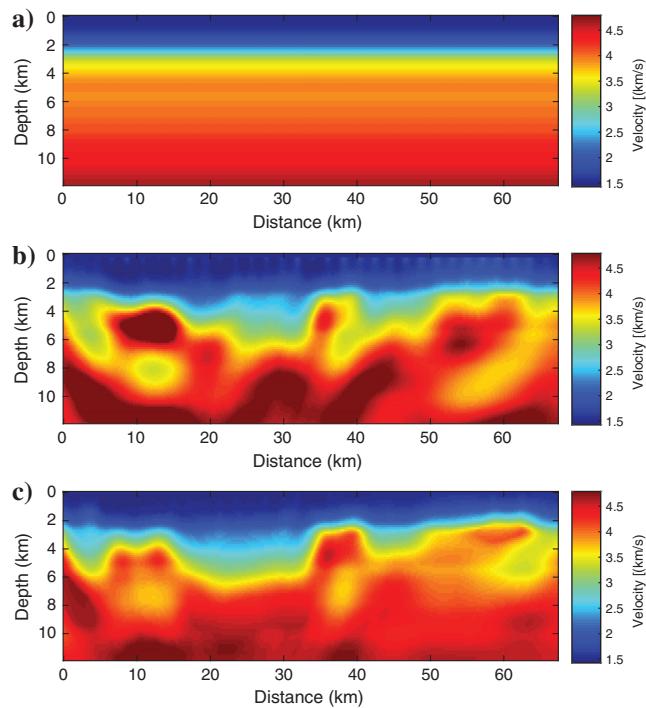


Figure 19. Comparison among (a) the initial model for FWI on the BP model, the inverted low-wavenumber velocity models using (b) 0.3 Hz extrapolated data, and (c) 0.3 Hz true data. The inversion results in (b) and (c) are started from the initial model in (a).

whether it be [0.6, 20.0] Hz or [1.2, 20.0] Hz, for instance. As mentioned earlier, extrapolating data from 0.3 to 1.2 Hz, so as to be useful for EFWI starting at 1.5 Hz, is the much tougher task.

## DISCUSSIONS AND LIMITATIONS

The most important limitation of CNN for bandwidth extension is the possibly large generalization error that can result from an incomplete training set or an architecture unable to predict well out of sample. As a data-driven statistical optimization method, deep learning requires a large number of samples (usually millions) to become an effective predictor. Because the training data set in this example is small but the model capacity (trainable parameters) is large, it is very easy for the neural network to overfit, which seriously deteriorates the extrapolation accuracy. Therefore, in practice, it is standard to use regularization or dropout, with only empirical evidence that this addresses the overfitting problem.

In addition, the training time for deep learning is highly related to the size of the data set and the model capacity and, thus, is very demanding. To speed up the training by reducing the number of weights of the neural networks, we can downsample the inputs and outputs and then use band-limited interpolation method to recover the signal after extrapolation.

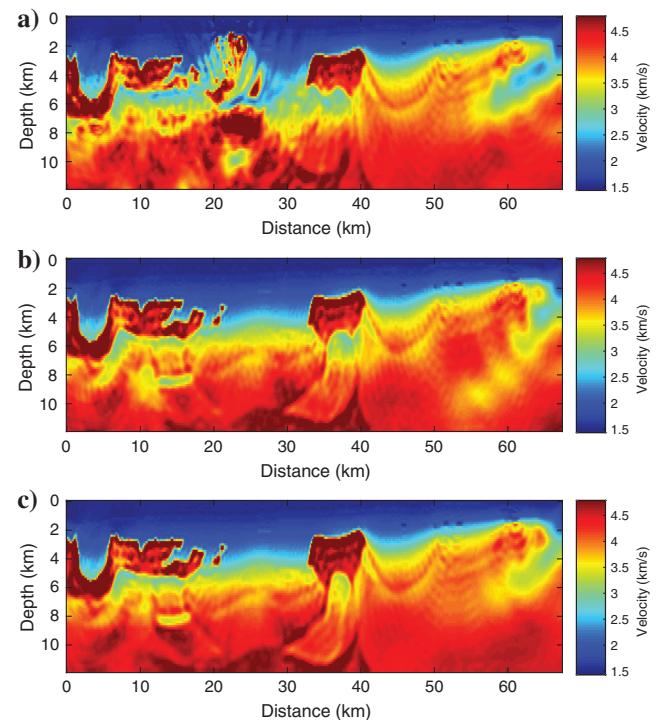


Figure 20. Comparison of the inverted models from FWI using 0.6–2.4 Hz band-limited data. (a) The resulting model starts from the original initial model. (b) The resulting model starts from the inverted low-wavenumber velocity model using 0.3 Hz extrapolated data. (c) The resulting model starts from the inverted low-wavenumber velocity model using 0.3 Hz true data.

Another limitation of deep learning is due to the unbalanced data. The energy of the direct wave is very strong compared with that of the reflected waves, which biases the neural networks toward fitting the direct wave and contributing less to the reflected waves. Therefore, the extrapolation accuracy of the reflected waves is not as good as that of the primary wave in this example.

One limitation of trace-by-trace extrapolation is that the accumulation of the predicted errors reduces the coherence of the event across traces. Hence, multitrace extrapolation can alleviate this problem to a certain degree by leveraging the spatial relationship

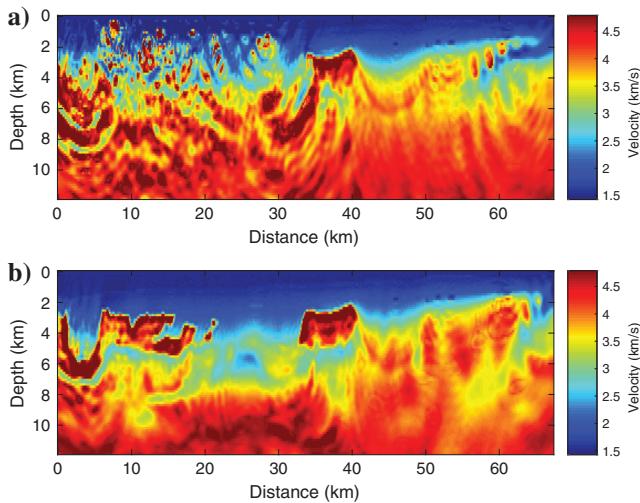


Figure 21. Comparison of the inverted models from FWI using 0.9–2.4 Hz band-limited data. (a) The resulting model starts from the original initial model. (b) The resulting model starts from the inverted low-wavenumber velocity model using 0.3 and 0.6 Hz extrapolated data. The extrapolated data below 0.9 Hz are recovered by 0.9–20.0 Hz band-limited data.

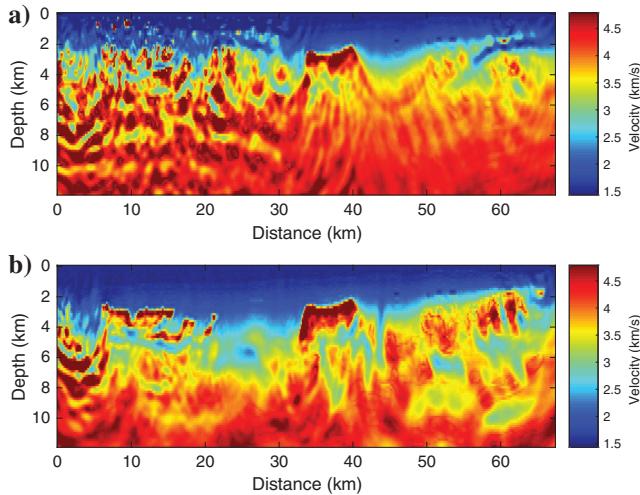


Figure 22. Comparison of the inverted models from FWI using 1.2–2.4 Hz band-limited data. (a) The resulting model starts from the original initial model. (b) The resulting model starts from the inverted low-wavenumber velocity model using 0.3, 0.6, and 0.9 Hz extrapolated data. The extrapolated data below 1.2 Hz are recovered by 1.2–20.0 Hz band-limited data.

existing in the input. The design of the architecture in Figure 1 is flexible to import multiple traces as the input of the neural network. To extrapolate the low-frequency signal of a single trace,  $ntr$  traces in the neighborhood of the single trace have been used as the input of the neural network. Then we only need to change the size of the filter on the first convolutional layer from  $200 \times 1 \times 1$  to  $200 \times ntr \times 1$  and keep the rest the same. Figure 24 compares the extrapolated low-frequency data (0.1–5.0 Hz) on the full-size Marmousi model using one trace ( $ntr = 1$ ), five traces ( $ntr = 5$ ), and seven traces ( $ntr = 7$ ) as the input of neural network. The predicted low-frequency data using multiple-trace input show better coherence along traces compared to that using trace-by-trace extrapolation. Additionally, more numerical experiments show that multiple-trace extrapolation helps to reduce the random noise but is unhelpful to correlated noise.

Even though we are encouraged by the ability of a CNN to generate [0.1, 0.5] Hz data for the BP 2004 model, much work remains to be done to be able to find the right architecture that will generate data in broader frequency bands, for instance, in the [0.1, 1.4] Hz band. Finding a suitable network architecture, hyperparameters, and training schedule for such cases remains an important open problem. Other community models, and more realistic physics such as elastic waves, are also left to be explored.

Finally, the influence of different physics between the training and test data set is left to be studied. This point is important for the application to field data. Even though we show the robustness of the proposed method in dealing with uncertainty due to random noise, a different forward modeling operator and a poorly known source wavelet, more stable neural network, and training strategies are yet to be proposed to overcome the challenges of field data, such as strong correlated and uncorrelated noise, complex and unknown wavelet shot by shot, viscoelasticity, and anisotropy.

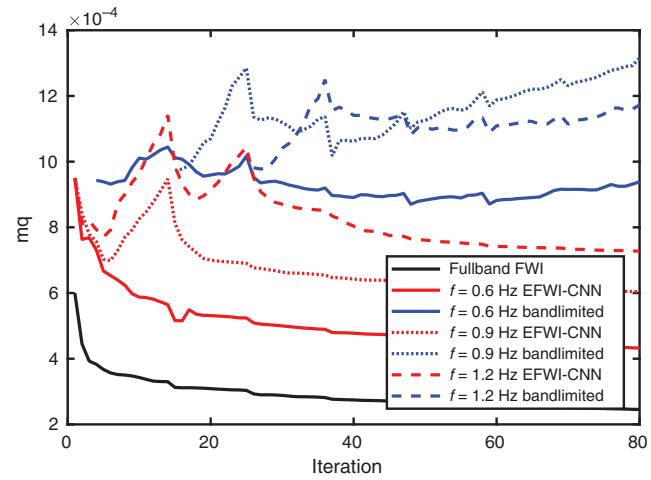


Figure 23. Comparison of the quality of the inverted models at each iteration for FWI using full band data (the black line), EFWI-CNN (the red line), and FWI using only band-limited data (the blue line).  $f$  denotes the lowest frequency of the band-limited data. The highest frequency of inversion is fixed at 2.4 Hz. The performance of EFWI-CNN decreases with the increase of the lowest frequency of the band-limited data. However, compared to FWI using only band-limited data, EFWI-CNN improves the quality of the inverted model very well.

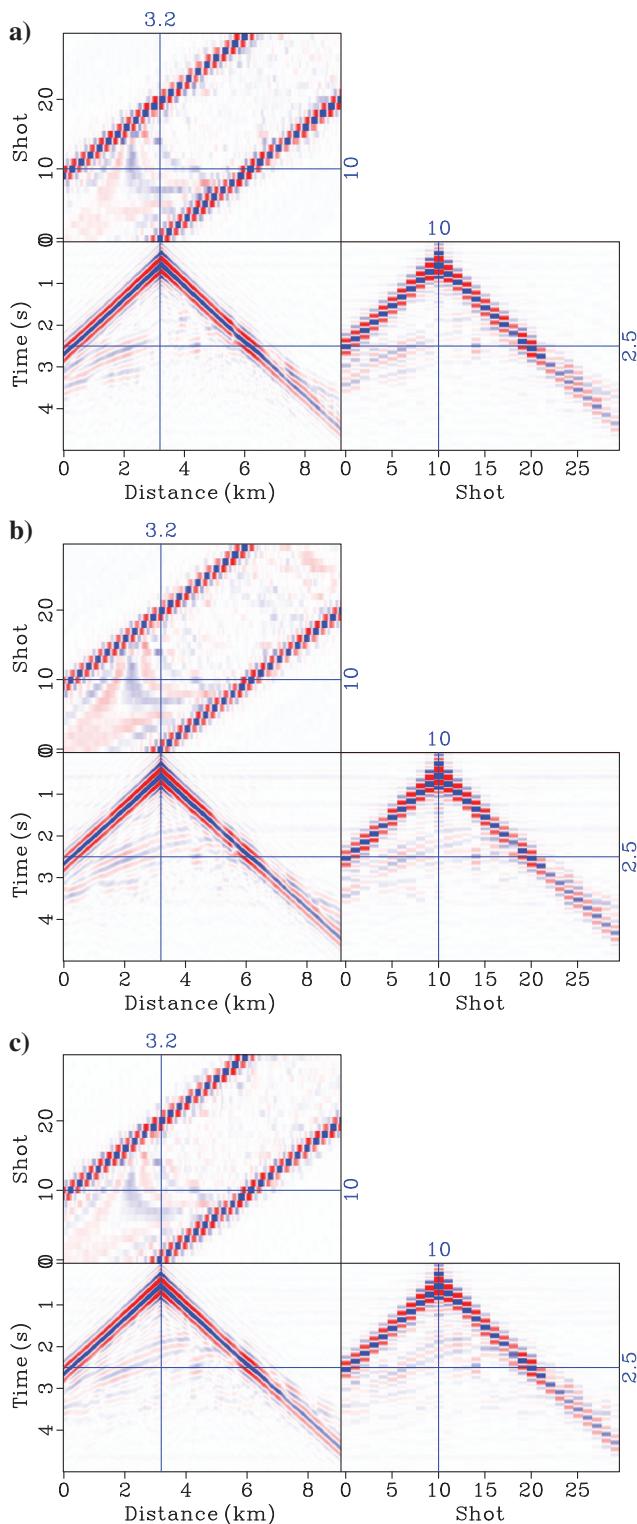


Figure 24. The extrapolation result on the Marmousi model: comparison of extrapolation using (a) one trace, (b) five traces, and (c) seven traces as the input of neural network. Multiple-trace extrapolation shows better coherence along traces by leveraging the spatial relationship existing in the input.

## CONCLUSION

In this paper, deep learning is applied to the challenging bandwidth extension problem that is essential for FWI. We formulate bandwidth extension as a regression problem in machine learning and propose an end-to-end trainable model for low-frequency extrapolation. Without preprocessing on the band-limited data and postprocessing on the extrapolated low frequencies, CNN sometimes have the ability to recover the low frequencies of unknown subsurface structure that are completely missing at the training stage. The extrapolated low-frequency data can be reliable to invert the low-wavenumber velocity model for initializing FWI on the band-limited data without cycle-skipping. Even though there is freedom in choosing the architectural parameters of the DNN, making the CNN have a large receptive field is necessary for low-frequency extrapolation. The extrapolation accuracy can be further modified by adjusting the architecture and hyperparameters of the neural networks depending on the characteristics of the band-limited data.

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## DATA AND MATERIALS AVAILABILITY

Data associated with this research are available and can be obtained by contacting the corresponding author.

## REFERENCES

- Abadi, M., A. Agarwal, P. Barham, E. Brevdo, Z. Chen, C. Citro, G. S. Corrado, A. Davis, J. Dean, M. Devin, S. Ghemawat, I. Goodfellow, A. Harp, G. Irving, M. Isard, Y. Jia, R. Jozefowicz, L. Kaiser, M. Kudlur, J. Levenberg, D. Mané, R. Monga, S. Moore, D. Murray, C. Olah, M. Schuster, J. Shlens, B. Steiner, I. Sutskever, K. Talwar, P. Tucker, V. Vanhoucke, V. Vasudevan, F. Viégas, O. Vinyals, P. Warden, M. Wattenberg, M. Wicke, Y. Yu, and X. Zheng, 2015, TensorFlow: Large-scale machine learning on heterogeneous systems (Software available from tensorflow.org).
- Araya-Polo, M., T. Dahlke, C. Frogner, C. Zhang, T. Poggio, and D. Hohl, 2017, Automated fault detection without seismic processing: The Leading Edge, **36**, 208–214, doi: [10.1190/tle36030208.1](https://doi.org/10.1190/tle36030208.1).
- Araya-Polo, M., J. Jennings, A. Adler, and T. Dahlke, 2018, Deep-learning tomography: The Leading Edge, **37**, 58–66, doi: [10.1190/tle37010058.1](https://doi.org/10.1190/tle37010058.1).
- Billette, F., and S. Brandsberg-Dahl, 2005, The 2004 BP velocity benchmark: 67th Annual International Conference and Exhibition, EAGE, Extended Abstracts, B035.
- Brossier, R., S. Operto, and J. Virieux, 2009, Seismic imaging of complex onshore structures by 2D elastic frequency-domain full-waveform inversion: Geophysics, **74**, no. 6, WCC105–WCC118, doi: [10.1190/1.3215771](https://doi.org/10.1190/1.3215771).
- Bunks, C., F. M. Saleck, S. Zaleski, and G. Chavent, 1995, Multiscale seismic waveform inversion: Geophysics, **60**, 1457–1473, doi: [10.1190/1.1443880](https://doi.org/10.1190/1.1443880).
- Chen, Y., J. Hill, W. Lei, M. Lefebvre, J. Tromp, E. Bozdag, and D. Komatsitsch, 2017, Automated time-window selection based on machine learning for full-waveform inversion: 87th Annual International Meeting, SEG, Expanded Abstracts, 1604–1609, doi: [10.1190/segam2017-17734162.1](https://doi.org/10.1190/segam2017-17734162.1).
- Chollet, F., 2015, Keras, <https://github.com/fchollet/keras>, accessed 8 April 2020.
- De, S., D. Deo, G. Sankaranarayanan, and V. S. Arikatla, 2011, A physics-driven neural networks-based simulation system (phyness) for multimodal interactive virtual environments involving nonlinear deformable objects: Presence: Teleoperators and Virtual Environments, **20**, 289–308, doi: [10.1162/PRES\\_a\\_00054](https://doi.org/10.1162/PRES_a_00054).
- Glorot, X., and Y. Bengio, 2010, Understanding the difficulty of training deep feedforward neural networks: Proceedings of the Thirteenth International Conference on Artificial Intelligence and Statistics, 249–256.
- Goodfellow, I., Y. Bengio, and A. Courville, 2016, Deep learning: MIT Press, 1.

- Grzeszczuk, R., D. Terzopoulos, and G. Hinton, 1998, Neuroanimator: Fast neural network emulation and control of physics-based models: Proceedings of the 25th Annual Conference on Computer Graphics and Interactive Techniques, 9–20.
- Guitton, A., H. Wang, and W. Trainor-Guitton, 2017, Statistical imaging of faults in 3D seismic volumes using a machine learning approach: 87th Annual International Meeting, SEG, Expanded Abstracts, 2045–2049, doi: [10.1190/segam2017-17589633.1](https://doi.org/10.1190/segam2017-17589633.1).
- He, K., X. Zhang, S. Ren, and J. Sun, 2015, Delving deep into rectifiers: Surpassing human-level performance on ImageNet classification: Proceedings of the IEEE International Conference on Computer Vision, 1026–1034.
- Hewett, R., and L. Demanet, 2013, PySIT: Python Seismic Imaging Toolbox v0.5, <http://www.pysit.org>.
- Hornik, K., M. Stinchcombe, and H. White, 1989, Multilayer feedforward networks are universal approximators: Elsevier, 2.
- Hu, W., 2014, FWI without low frequency data-beat tone inversion: 84th Annual International Meeting, SEG, Expanded Abstracts, 1116–1120, doi: [10.1190/segam2014-0978.1](https://doi.org/10.1190/segam2014-0978.1).
- Hu, W., Y. Jin, X. Wu, and J. Chen, 2019, A progressive deep transfer learning approach to cycle-skipping mitigation in FWI: 89th Annual International Meeting, SEG, Expanded Abstracts, 2348–2352, doi: [10.1190/segam2019-3216030.1](https://doi.org/10.1190/segam2019-3216030.1).
- Hu, Y., L. Han, Z. Xu, F. Zhang, and J. Zeng, 2017, Adaptive multi-step full waveform inversion based on waveform mode decomposition: Elsevier.
- Ioffe, S., and C. Szegedy, 2015, Batch normalization: Accelerating deep network training by reducing internal covariate shift: arXiv preprint, arXiv:1502.03167.
- Jin, Y., W. Hu, X. Wu, and J. Chen, 2018, Learn low wavenumber information in FWI via deep inception based convolutional networks: 88th Annual International Meeting, SEG, Expanded Abstracts, 2091–2095, doi: [10.1190/segam2018-2997901.1](https://doi.org/10.1190/segam2018-2997901.1).
- Kazei, V., O. Ovcharenko, T. Alkhalfah, and F. Simons, 2019, Realistically textured random velocity models for deep learning applications: 81st Annual International Conference and Exhibition, EAGE, Extended Abstracts, doi: [10.3997/2214-4609.201901340](https://doi.org/10.3997/2214-4609.201901340).
- Kingma, D. P., and J. Ba, 2014, Adam: A method for stochastic optimization: arXiv preprint, arXiv:1412.6980.
- Levy, S., and P. K. Fullagar, 1981, Reconstruction of a sparse spike train from a portion of its spectrum and application to high-resolution deconvolution: Geophysics, **46**, 1235–1243, doi: [10.1190/1.1441261](https://doi.org/10.1190/1.1441261).
- Lewis, W., and D. Vigh, 2017, Deep learning prior models from seismic images for full-waveform inversion: 87th Annual International Meeting, SEG, Expanded Abstracts, 1512–1517, doi: [10.1190/segam2017-17627643.1](https://doi.org/10.1190/segam2017-17627643.1).
- Li, S., B. Liu, Y. Ren, Y. Chen, S. Yang, Y. Wang, and P. Jiang, 2019, Deep learning inversion of seismic data: arXiv preprint, arXiv:1901.07733.
- Li, Y. E., and L. Demanet, 2015, Phase and amplitude tracking for seismic event separation: Geophysics, **80**, no. 6, WD59–WD72, doi: [10.1190/geo2015-0075.1](https://doi.org/10.1190/geo2015-0075.1).
- Li, Y. E., and L. Demanet, 2016, Full-waveform inversion with extrapolated low-frequency data: Geophysics, **81**, no. 6, R339–R348, doi: [10.1190/geo2016-0038.1](https://doi.org/10.1190/geo2016-0038.1).
- Noceval, J., and S. J. Wright, 2006, Numerical optimization: Springer.
- Oldenburg, D., T. Scheuer, and S. Levy, 1983, Recovery of the acoustic impedance from reflection seismograms: Geophysics, **48**, 1318–1337, doi: [10.1190/1.1441413](https://doi.org/10.1190/1.1441413).
- Ovcharenko, O., V. Kazei, M. Kalita, D. Peter, and T. A. Alkhalfah, 2019a, Deep learning for low-frequency extrapolation from multi-offset seismic data: Geophysics, **84**, no. 6, R989–R1001, doi: [10.1190/geo2018-0884.1](https://doi.org/10.1190/geo2018-0884.1).
- Ovcharenko, O., V. Kazei, D. Peter, and T. Alkhalfah, 2017, Neural network based low-frequency data extrapolation: Presented at the 3rd SEG FWI workshop: What are we getting, SEG.
- Ovcharenko, O., V. Kazei, D. Peter, and T. Alkhalfah, 2019b, Transfer learning for low frequency extrapolation from shot gathers for FWI applications: 81st Annual International Conference and Exhibition, EAGE, Extended Abstracts, ThR1112.
- Ovcharenko, O., V. Kazei, D. Peter, X. Zhang, and T. Alkhalfah, 2018, Low-frequency data extrapolation using a feed-forward ANN: 80th Annual International Conference and Exhibition, EAGE, Extended Abstracts, WeP314.
- Richardson, A., 2018, Seismic full-waveform inversion using deep learning tools and techniques: arXiv preprint, arXiv:1801.07232.
- Sacchi, M. D., D. R. Velis, and A. H. Cominguez, 1994, Minimum entropy deconvolution with frequency-domain constraints: Geophysics, **59**, 938–945, doi: [10.1190/1.1443653](https://doi.org/10.1190/1.1443653).
- Sirgue, L., and R. G. Pratt, 2004, Efficient waveform inversion and imaging: A strategy for selecting temporal frequencies: Geophysics, **69**, 231–248, doi: [10.1190/1.1649391](https://doi.org/10.1190/1.1649391).
- Srivastava, N., G. Hinton, A. Krizhevsky, I. Sutskever, and R. Salakhutdinov, 2014, Dropout: A simple way to prevent neural networks from overfitting: The Journal of Machine Learning Research, **15**, 1929–1958.
- Sun, H., and L. Demanet, 2018, Low frequency extrapolation with deep learning: 88th Annual International Meeting, SEG, Expanded Abstracts, 2011–2015, doi: [10.1190/segam2018-2997928.1](https://doi.org/10.1190/segam2018-2997928.1).
- Sun, H., and L. Demanet, 2019, Extrapolated full waveform inversion with convolutional neural networks: 89th Annual International Meeting, SEG, Expanded Abstracts, 4962–4966, doi: [10.1190/segam2019-3197987.1](https://doi.org/10.1190/segam2019-3197987.1).
- Walker, C., and T. J. Ulrych, 1983, Autoregressive recovery of the acoustic impedance: Geophysics, **48**, 1338–1350, doi: [10.1190/1.1441414](https://doi.org/10.1190/1.1441414).
- Wang, R., and F. Herrmann, 2016, Frequency down extrapolation with TV norm minimization: 86th Annual International Meeting, SEG, Expanded Abstracts, 1380–1384, doi: [10.1190/segam2016-13879674.1](https://doi.org/10.1190/segam2016-13879674.1).
- Wu, R.-S., J. Luo, and B. Wu, 2014, Seismic envelope inversion and modulation signal model: Geophysics, **79**, no. 3, WA13–WA24, doi: [10.1190/geo2013-0294.1](https://doi.org/10.1190/geo2013-0294.1).
- Wu, Y., Y. Lin, and Z. Zhou, 2018, Inversionnet: Accurate and efficient seismic-waveform inversion with convolutional neural networks: 88th Annual International Meeting, SEG, Expanded Abstracts, 2096–2100, doi: [10.1190/segam2018-2998603.1](https://doi.org/10.1190/segam2018-2998603.1).
- Xiong, W., X. Ji, Y. Ma, Y. Wang, N. M. BenHassan, M. N. Ali, and Y. Luo, 2018, Seismic fault detection with convolutional neural network: Geophysics, **83**, no. 5, O97–O103, doi: [10.1190/geo2017-0666.1](https://doi.org/10.1190/geo2017-0666.1).
- Zhang, P., L. Han, Z. Xu, F. Zhang, and Y. Wei, 2017, Sparse blind deconvolution based low-frequency seismic data reconstruction for multiscale full waveform inversion: Journal of Applied Geophysics, **139**, 91–108, doi: [10.1016/j.jappgeo.2017.02.021](https://doi.org/10.1016/j.jappgeo.2017.02.021).