

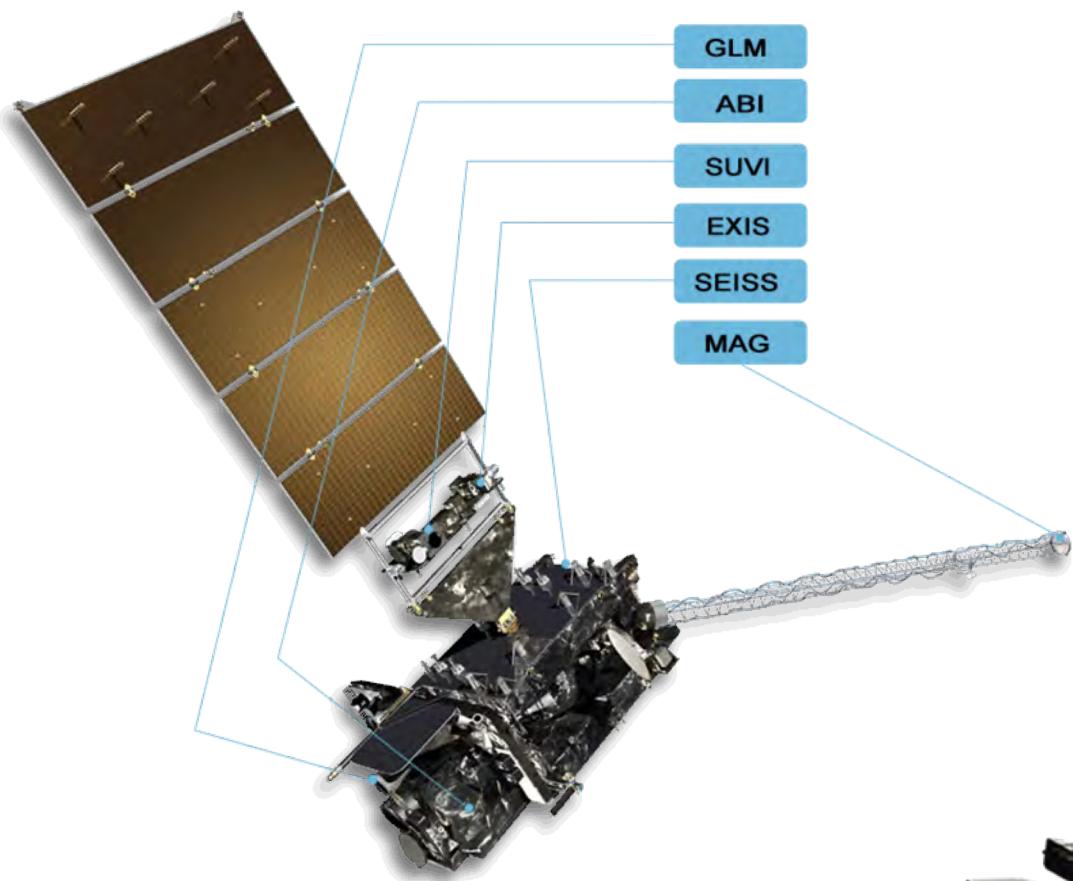
# MR3522: Remote Sensing of the Atmosphere and Ocean

## Geostationary Satellites: GOES and Himawari

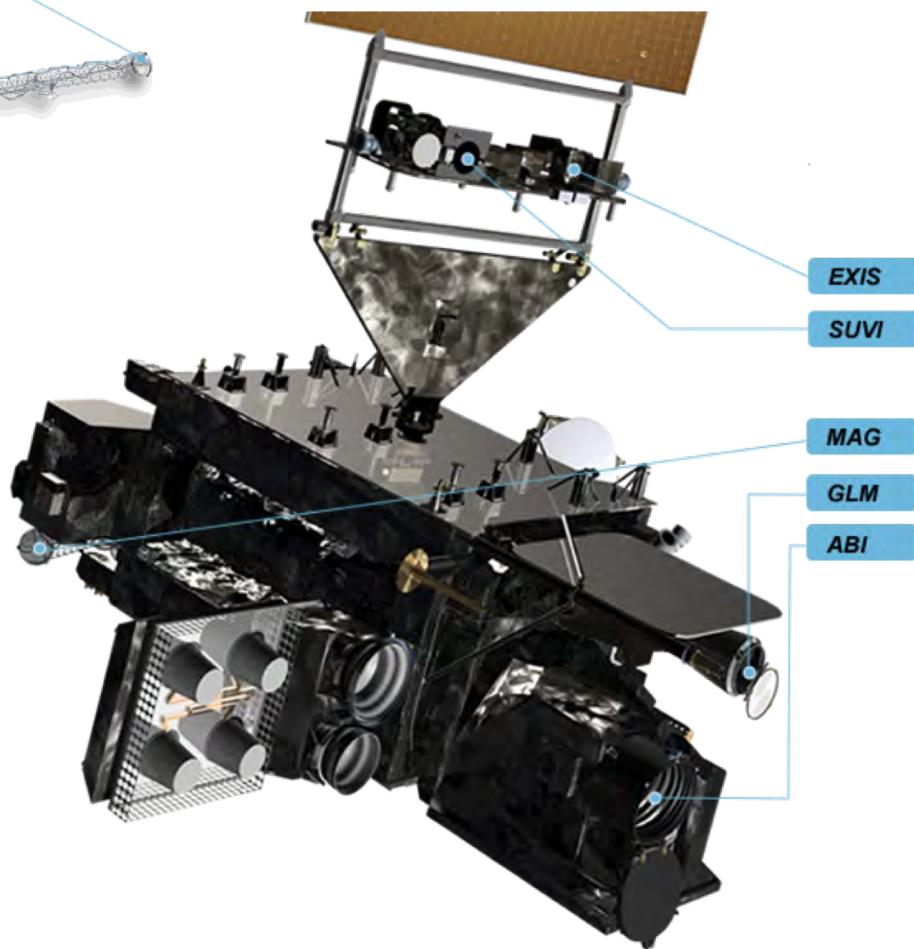
### Main Topics

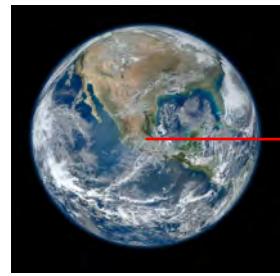
- Instrumentation
- Bands or channels used
- Absorption and scattering of radiation in GOES bands

NASA/NOAA

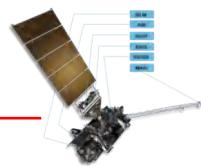


## GOES Instrumentation





42,164 km



# Features of New GOES and Himawari Satellites

- 16 channels or bands; no radiometer or “sounder” needed
- Can take full disk image every 5 minutes, or concurrently, disk images every 15 minutes, U.S. every 5 minutes, and two mesoscale regions every 30–60 seconds.
- Compared to 3<sup>rd</sup> generation, GOES-16 and GOES-17 have 3x more channels (spectral resolution), 4x the spatial resolution (up to 0.5 km spacing in the red band), and ~5x the temporal resolution.

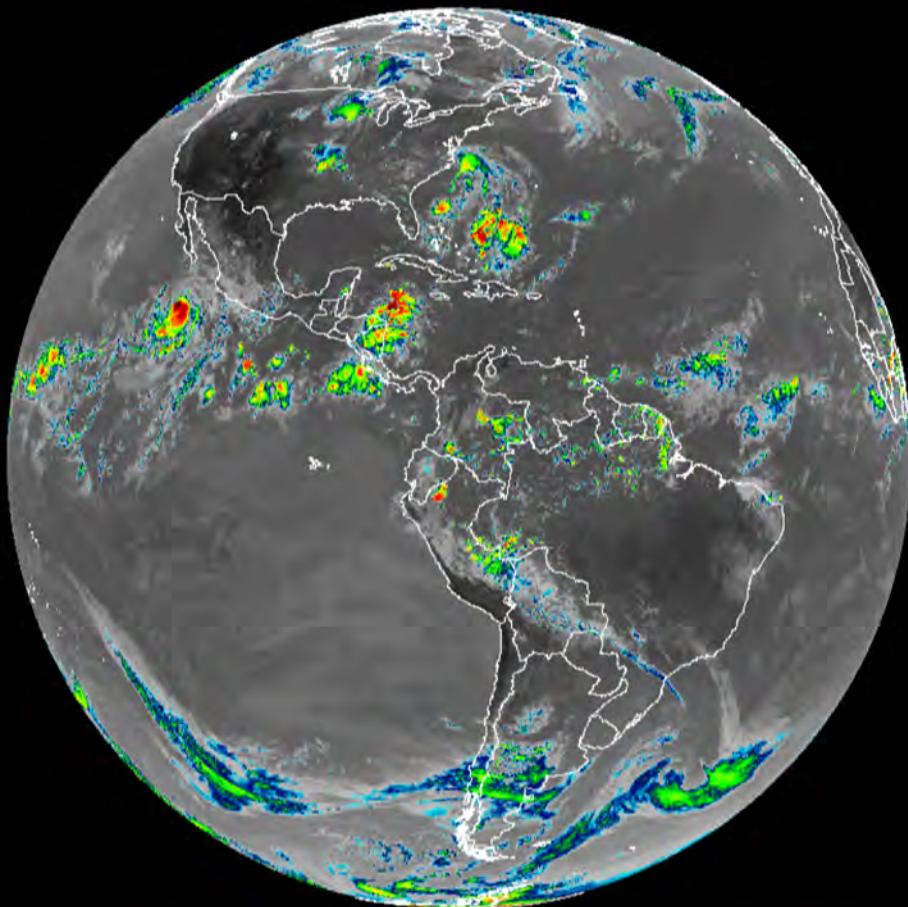
**TABLE 1.** Summary of the wavelengths, resolution, and sample use and heritage instrument(s) of the ABI bands. The minimum and maximum wavelength range represent the full width at half maximum (FWHM or 50%) points. [The Instantaneous Geometric Field Of View (IGFOV).]

Future GOES imager (ABI) band	Wavelength range ( $\mu\text{m}$ )	Central wavelength ( $\mu\text{m}$ )	Nominal subsatellite IGFOV (km)	Sample use	Heritage instrument(s)
1	0.45–0.49	0.47	1	Daytime aerosol over land, coastal water mapping	MODIS
2	0.59–0.69	0.64	0.5	Daytime clouds fog, insulation, winds	Current GOES imager/sounder
3	0.846–0.885	0.865	1	Daytime vegetation/burn scar and aerosol over water, winds	VIIRS, spectrally modified AVHRR
4	1.371–1.386	1.378	2	Daytime cirrus cloud	VIIRS, MODIS
5	1.58–1.64	1.61	1	Daytime cloud-top phase and particle size, snow	VIIRS, spectrally modified AVHRR
6	2.225–2.275	2.25	2	Daytime land/cloud properties, particle size, vegetation, snow	VIIRS, similar to MODIS
7	3.80–4.00	3.90	2	Surface and cloud, fog at night, fire, winds	Current GOES imager
8	5.77–6.6	6.19	2	High-level atmospheric water vapor, winds, rainfall	Current GOES imager
9	6.75–7.15	6.95	2	Midlevel atmospheric water vapor, winds, rainfall	Current GOES sounder
10	7.24–7.44	7.34	2	Lower-level water vapor, winds, and $\text{SO}_2$	Spectrally modified current GOES sounder
11	8.3–8.7	8.5	2	Total water for stability, cloud phase, dust, $\text{SO}_2$ , rainfall	MAS
12	9.42–9.8	9.61	2	Total ozone, turbulence, and winds	Spectrally modified current sounder
13	10.1–10.6	10.35	2	Surface and cloud	MAS
14	10.8–11.6	11.2	2	Imagery, SST, clouds, rainfall	Current GOES sounder
15	11.8–12.8	12.3	2	Total water, ash, and SST	Current GOES sounder
16	13.0–13.6	13.3	2	Air temperature, cloud heights and amounts	Current GOES sounder/GOES-12+ imager

Source: Schmit, T.J., Gunshor, M.M., Menzel, W.P., Gurka, J.J., Li, J., Bachmeier, A.S., 2005, Introducing the Next-Generation Advanced Baseline Imager on GOES-R, Bulletin of the American Meteorological Society, v. 86, p. 1079-1096.

Himawari has no 1.378  $\mu\text{m}$  band, but has a “green” band at 0.51  $\mu\text{m}$ .

# GOES-East (16)

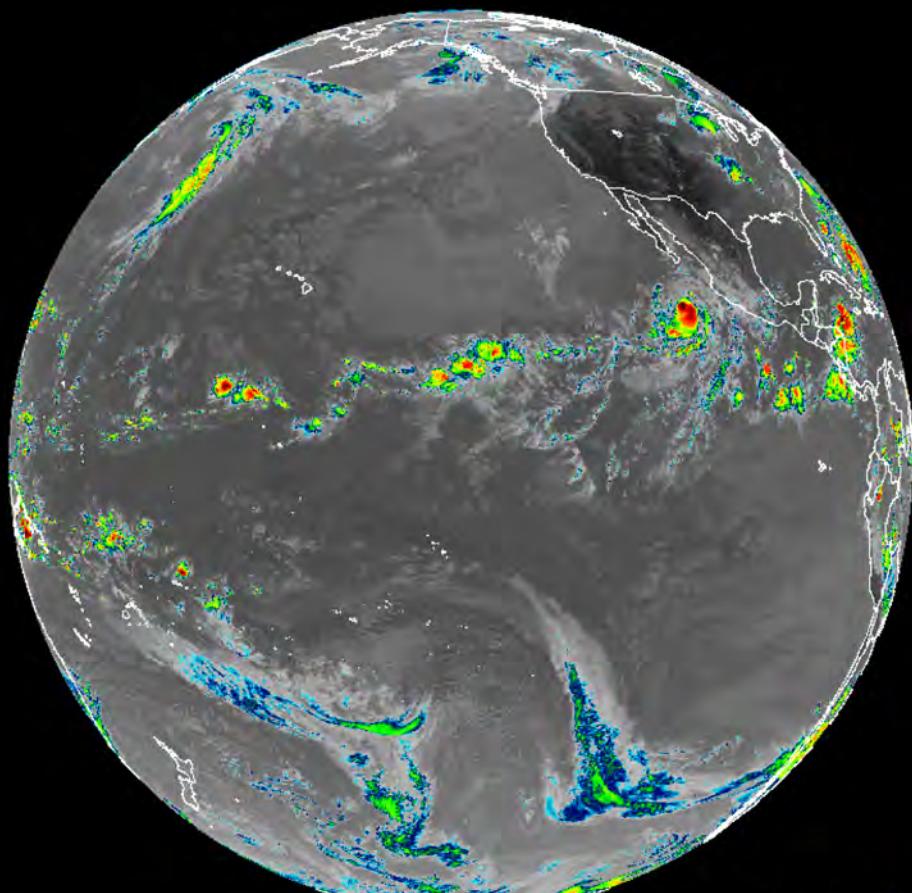


2020-07-09 17:50:24 UTC

Temperature

40 20 0 -20 -40 -60 -80 -90

# GOES-West (17)

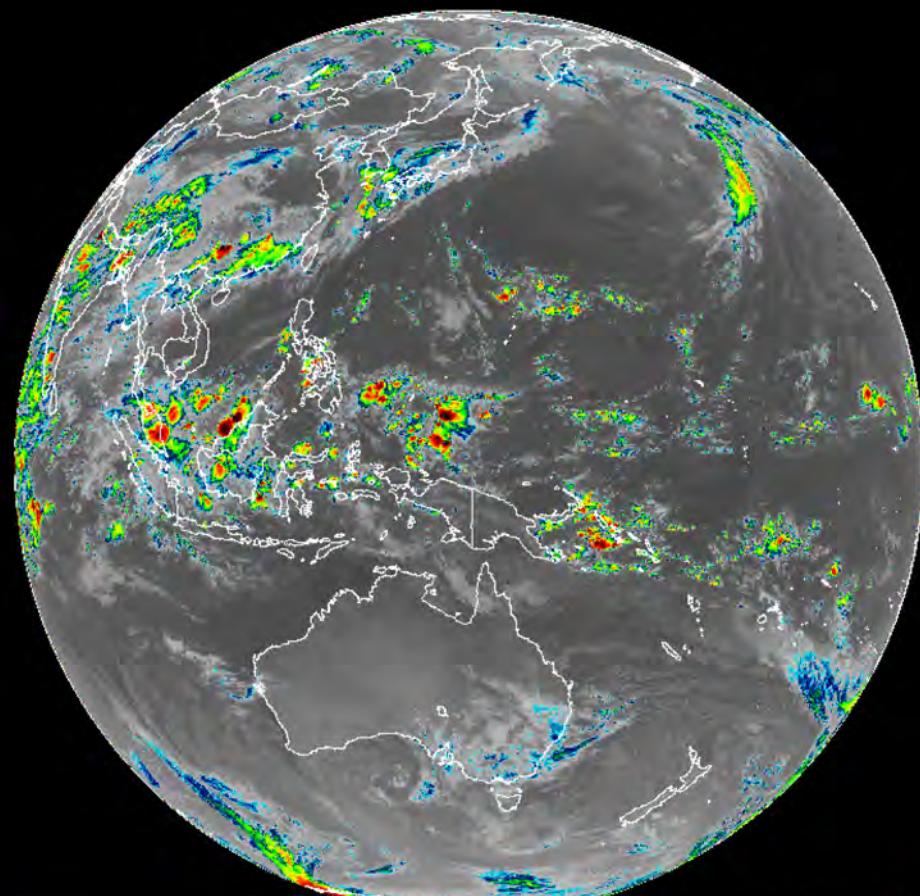


2020-07-09 17:40:32 UTC

Temperature 40 30 20 10 0 -10 -20 -30 -40 -50 -60 -70 -80 -90

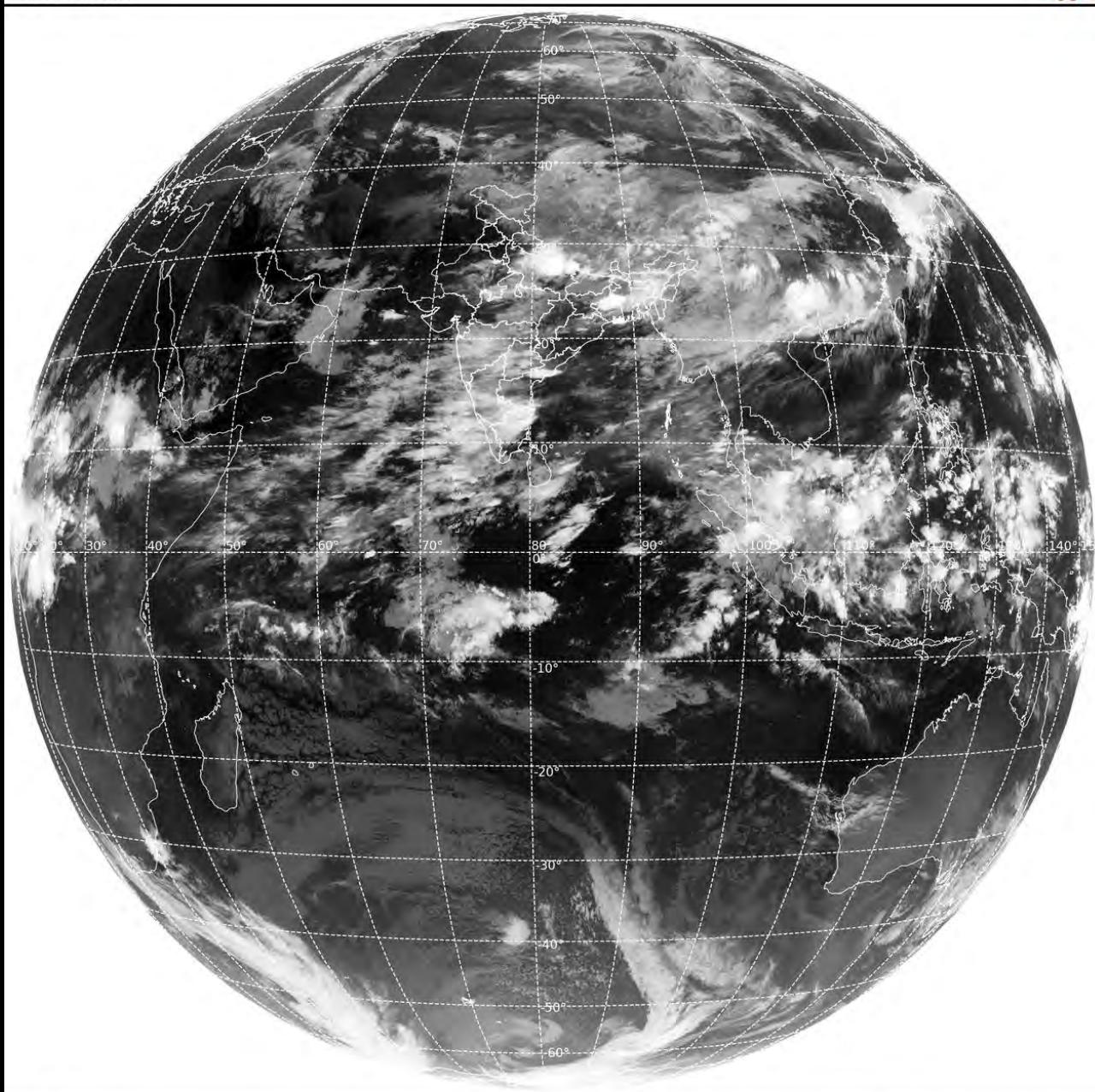


Himawari-8

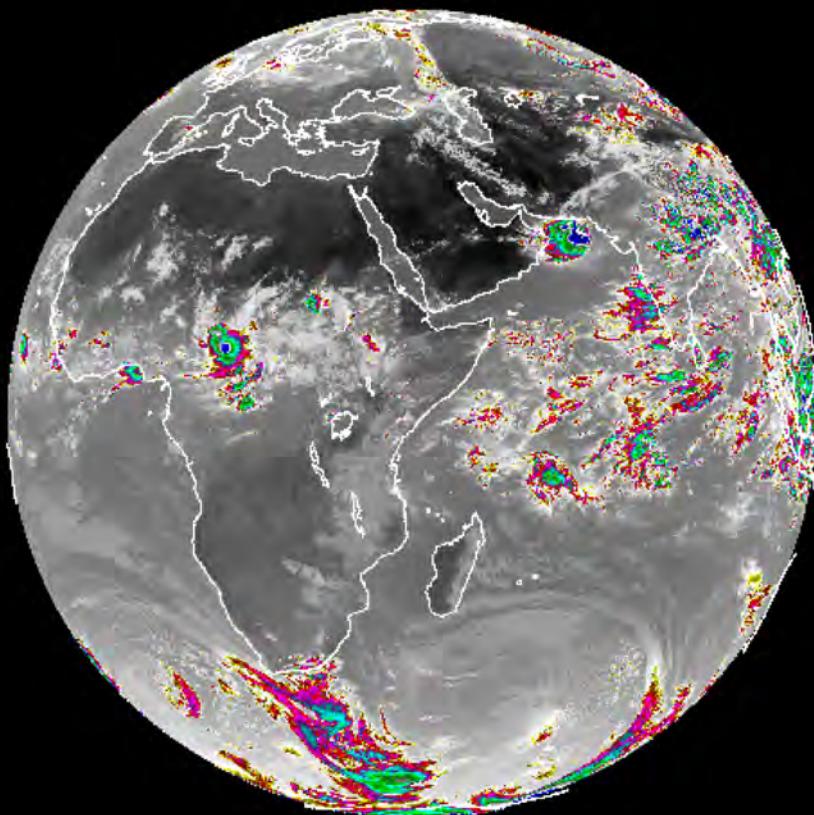




INSAT-3D

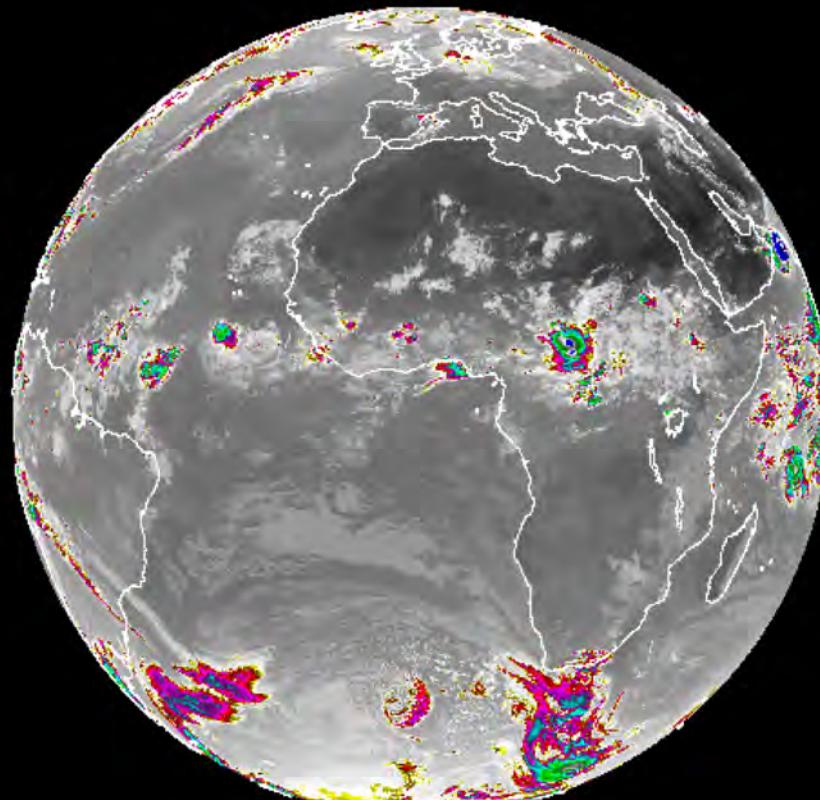


# Meteosat-8



-40 -30 -20 -10 0 10 20 30

# Meteosat-11



-40 20 20 10 0 -10 -20 40

# MR3522: Remote Sensing of the Atmosphere and Ocean

## GOES Advanced Baseline Imager Shortwave Bands

### Main Topics

- Utility of GOES shortwave bands
- Location of spectral response functions
- Reflection off the surface

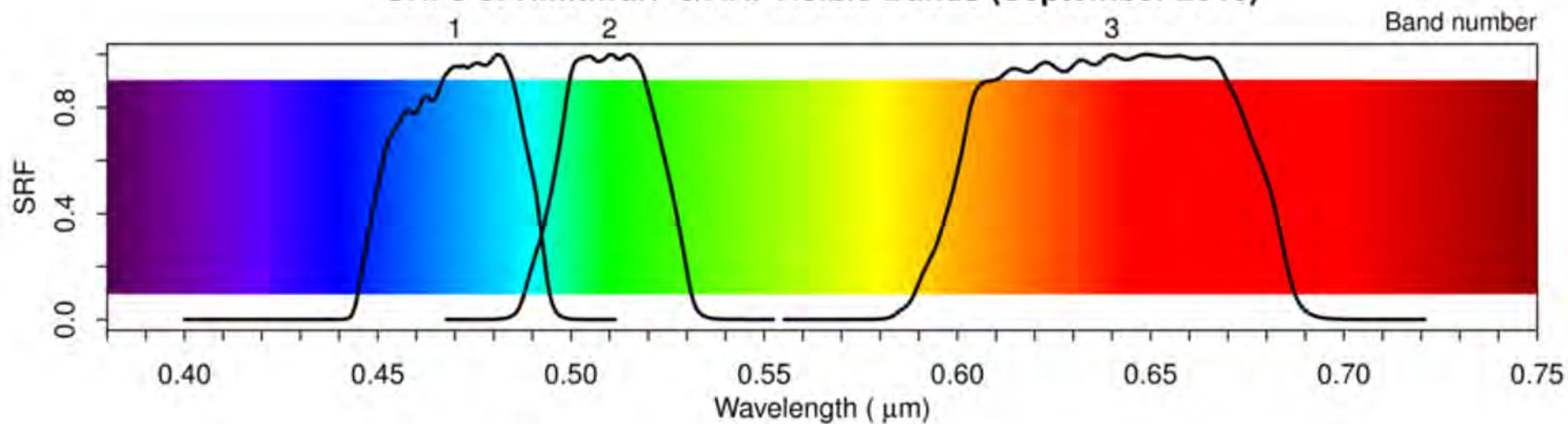
**TABLE 1.** Summary of the wavelengths, resolution, and sample use and heritage instrument(s) of the ABI bands. The minimum and maximum wavelength range represent the full width at half maximum (FWHM or 50%) points. [The Instantaneous Geometric Field Of View (IGFOV).]

Future GOES imager (ABI) band	Wavelength range ( $\mu\text{m}$ )	Central wavelength ( $\mu\text{m}$ )	Nominal subsatellite IGFOV (km)	Sample use	Heritage instrument(s)
1	0.45–0.49	0.47	1	Daytime aerosol over land, coastal water mapping	MODIS
2	0.59–0.69	0.64	0.5	Daytime clouds fog, insulation, winds	Current GOES imager/sounder
3	0.846–0.885	0.865	1	Daytime vegetation/burn scar and aerosol over water, winds	VIIRS, spectrally modified AVHRR
4	1.371–1.386	1.378	2	Daytime cirrus cloud	VIIRS, MODIS
5	1.58–1.64	1.61	1	Daytime cloud-top phase and particle size, snow	VIIRS, spectrally modified AVHRR
6	2.225–2.275	2.25	2	Daytime land/cloud properties, particle size, vegetation, snow	VIIRS, similar to MODIS
7	3.80–4.00	3.90	2	Surface and cloud, fog at night, fire, winds	Current GOES imager
8	5.77–6.6	6.19	2	High-level atmospheric water vapor, winds, rainfall	Current GOES imager
9	6.75–7.15	6.95	2	Midlevel atmospheric water vapor, winds, rainfall	Current GOES sounder
10	7.24–7.44	7.34	2	Lower-level water vapor, winds, and $\text{SO}_2$	Spectrally modified current GOES sounder
11	8.3–8.7	8.5	2	Total water for stability, cloud phase, dust, $\text{SO}_2$ , rainfall	MAS
12	9.42–9.8	9.61	2	Total ozone, turbulence, and winds	Spectrally modified current sounder
13	10.1–10.6	10.35	2	Surface and cloud	MAS
14	10.8–11.6	11.2	2	Imagery, SST, clouds, rainfall	Current GOES sounder
15	11.8–12.8	12.3	2	Total water, ash, and SST	Current GOES sounder
16	13.0–13.6	13.3	2	Air temperature, cloud heights and amounts	Current GOES sounder/GOES-12+ imager

Source: Schmit, T.J., Gunshor, M.M., Menzel, W.P., Gurka, J.J., Li, J., Bachmeier, A.S., 2005, Introducing the Next-Generation Advanced Baseline Imager on GOES-R, Bulletin of the American Meteorological Society, v. 86, p. 1079-1096.

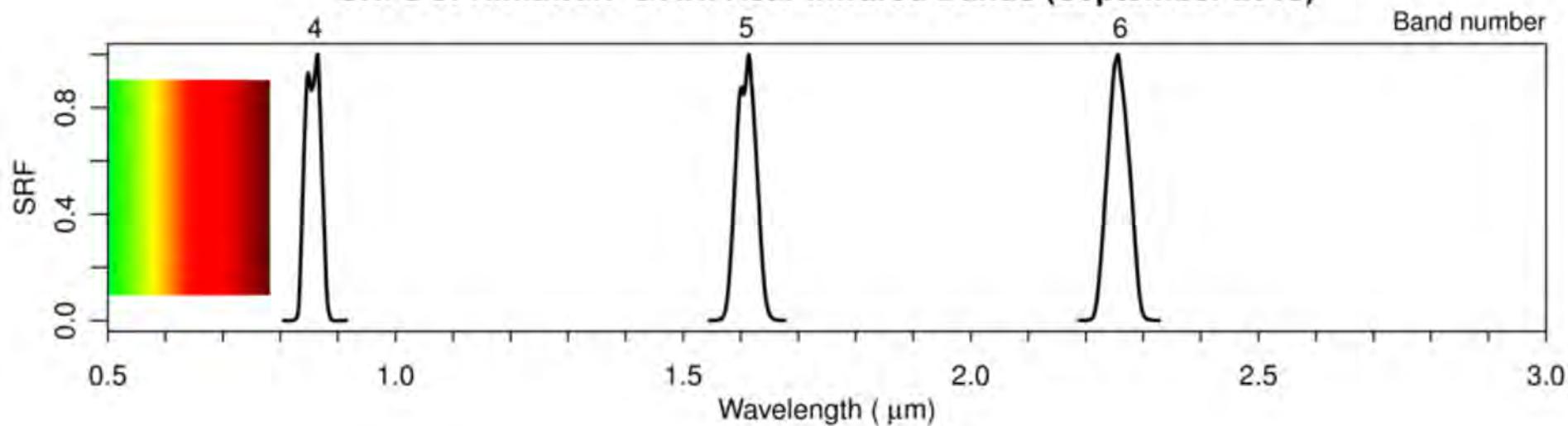
Himawari has no 1.378  $\mu\text{m}$  band, but has a “green” band at 0.51  $\mu\text{m}$ .

### SRFs of Himawari-8/AHI Visible Bands (September 2013)

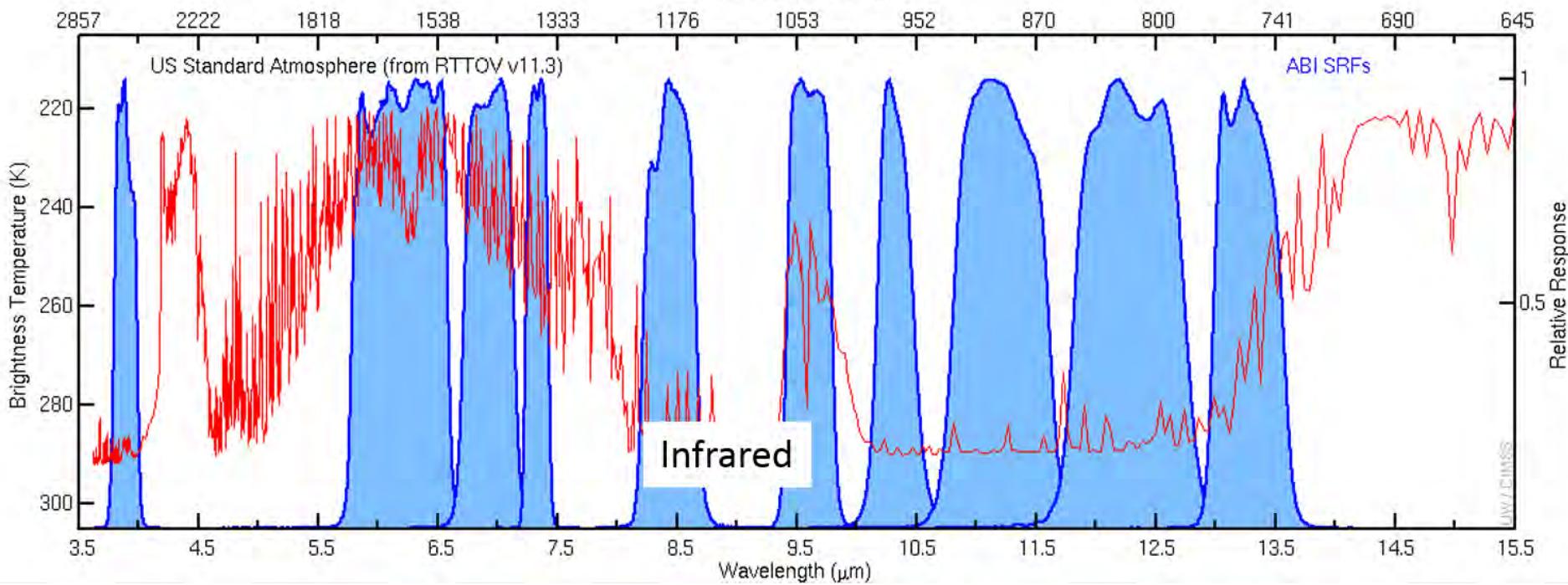
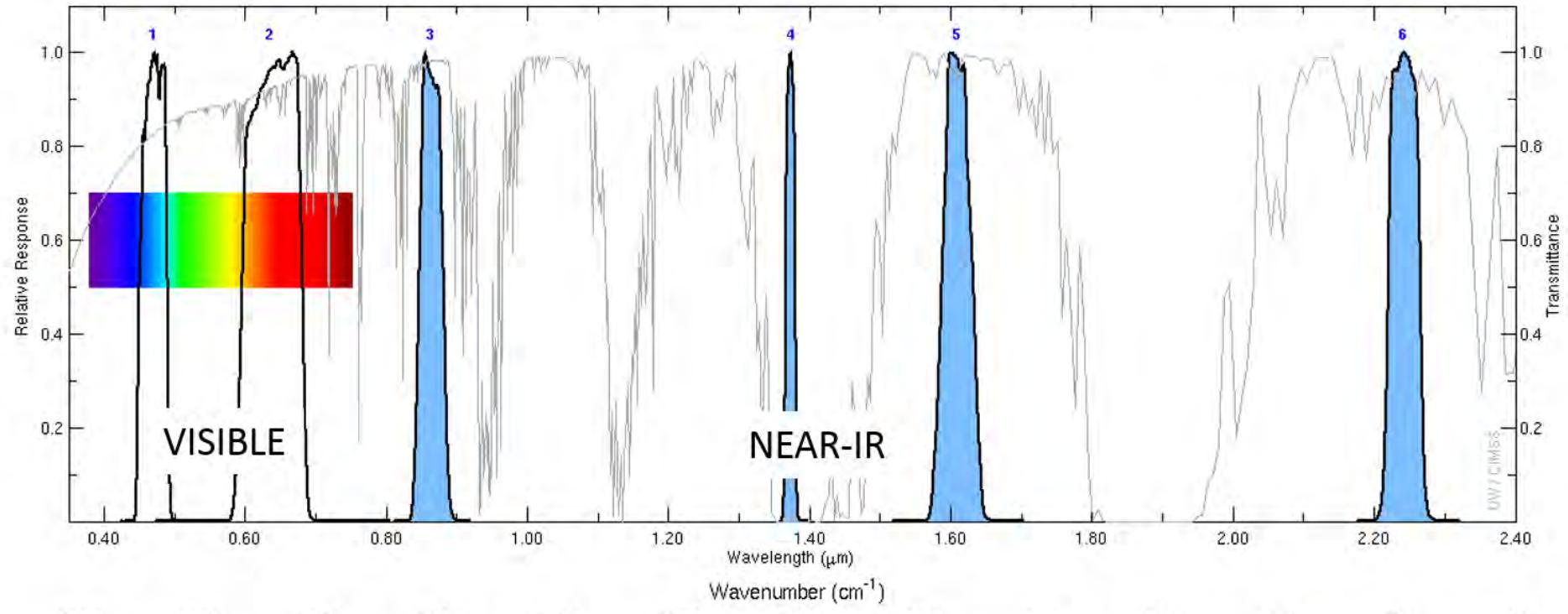


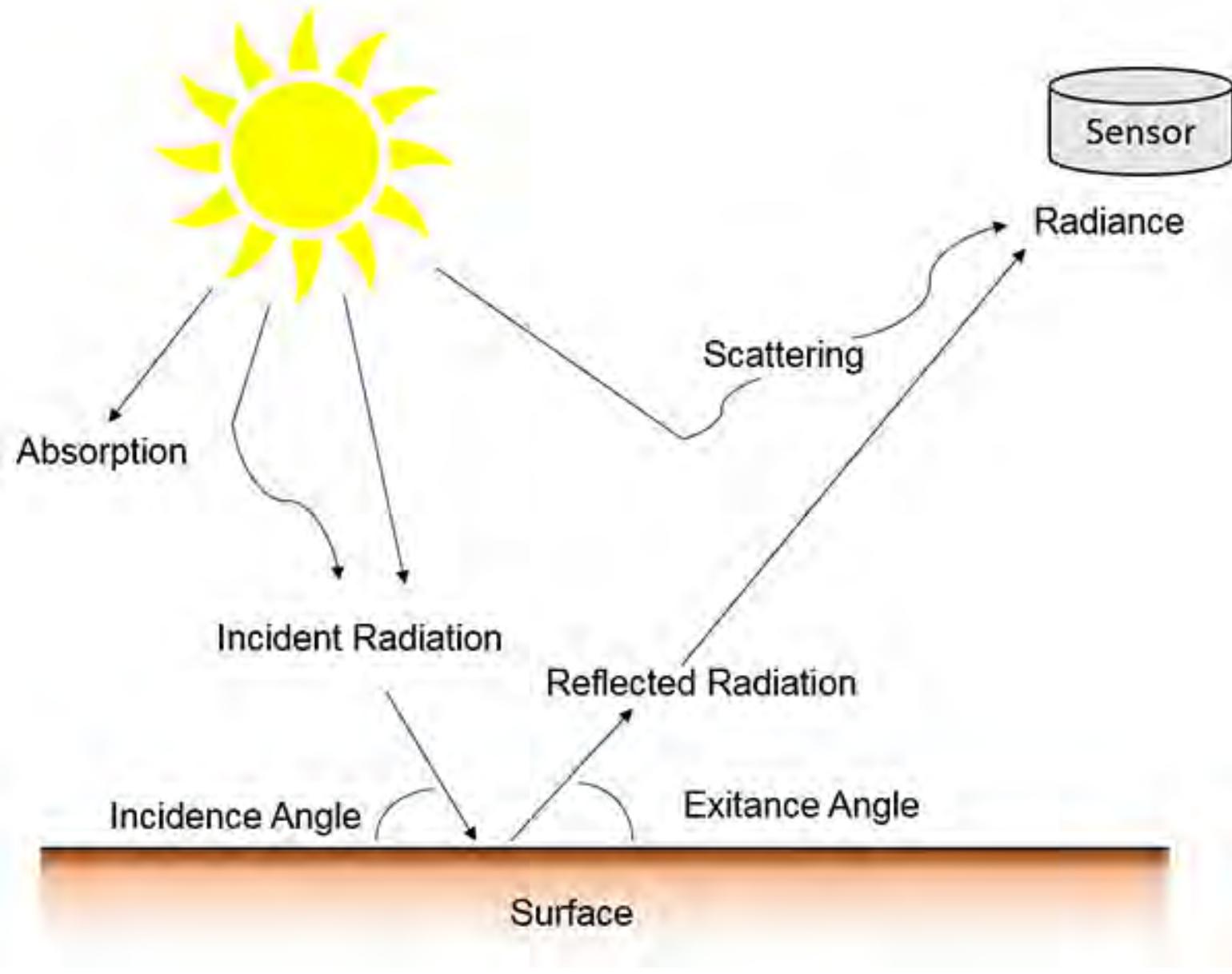
RGB VALUES FOR VISIBLE WAVELENGTHS by Dan Bruton (<http://www.physics.sfasu.edu/astro/color/spectra.html>)

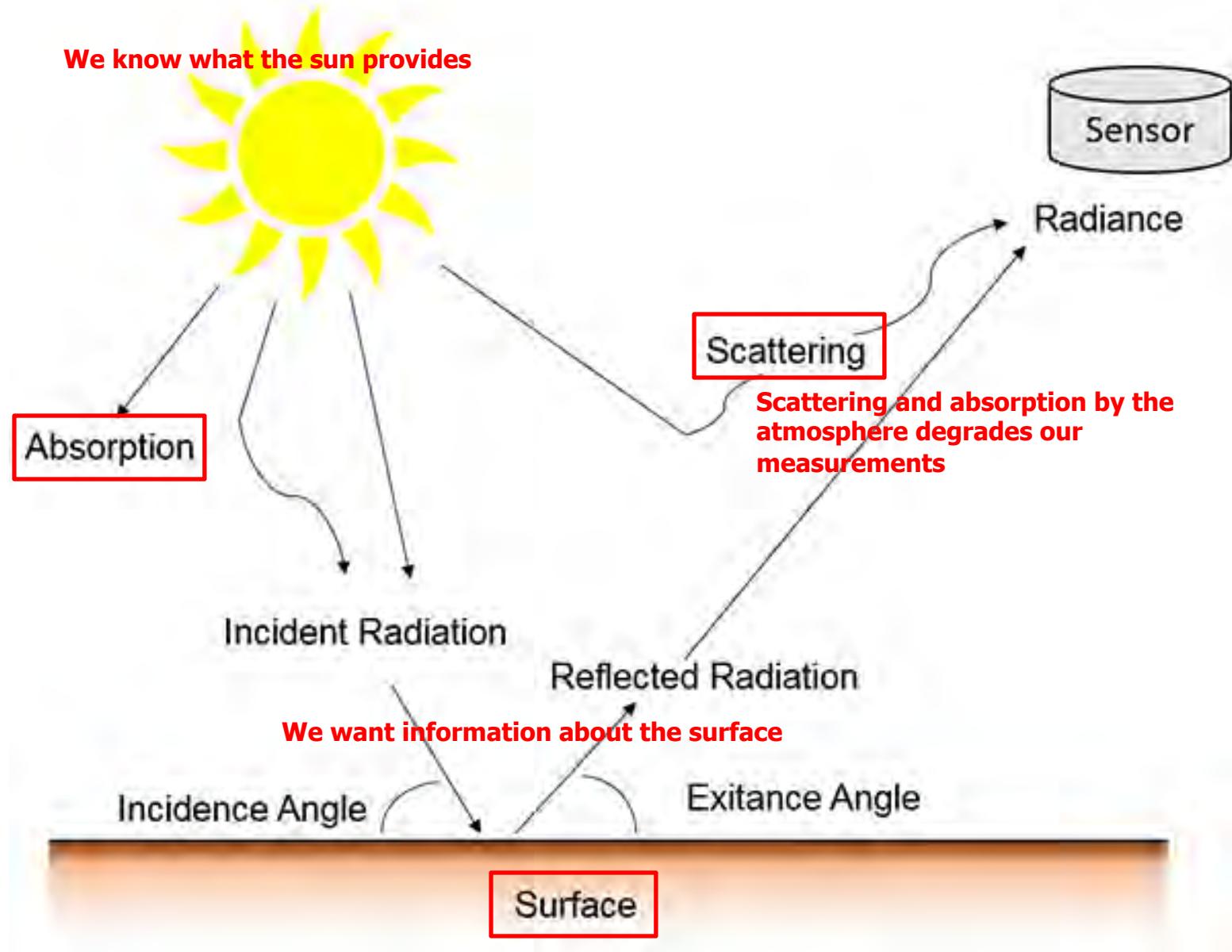
### SRFs of Himawari-8/AHI Near Infrared Bands (September 2013)

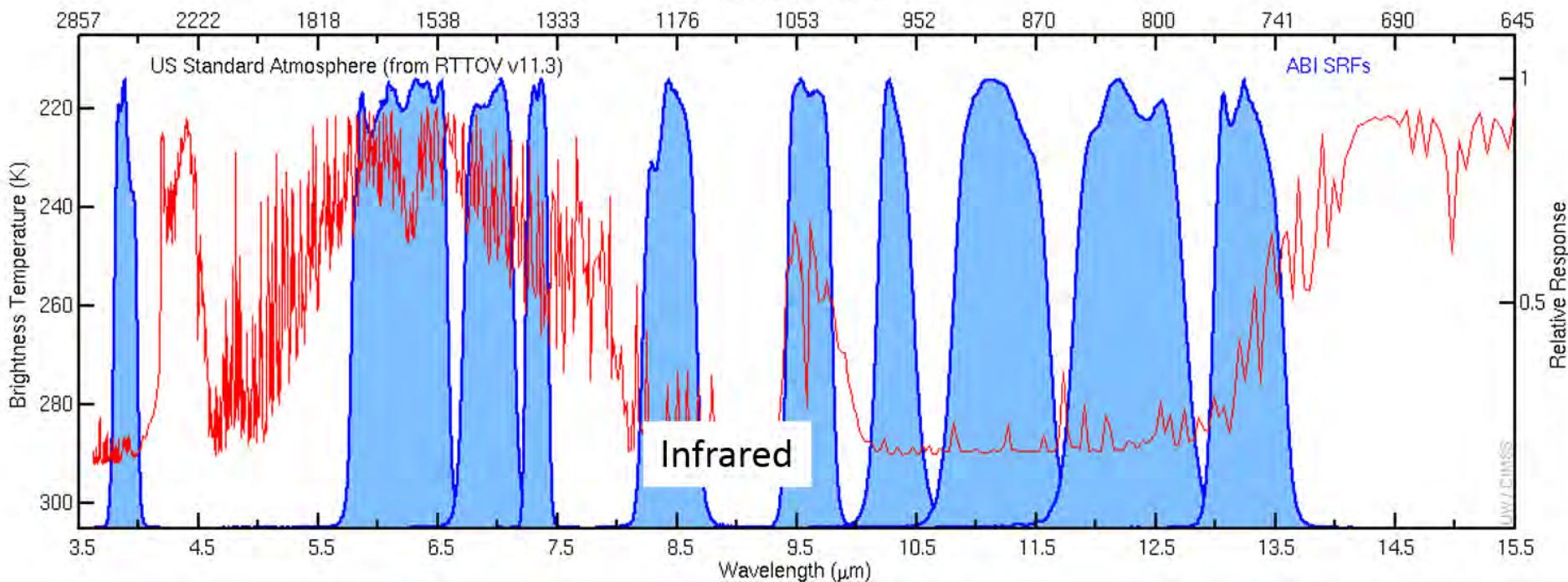
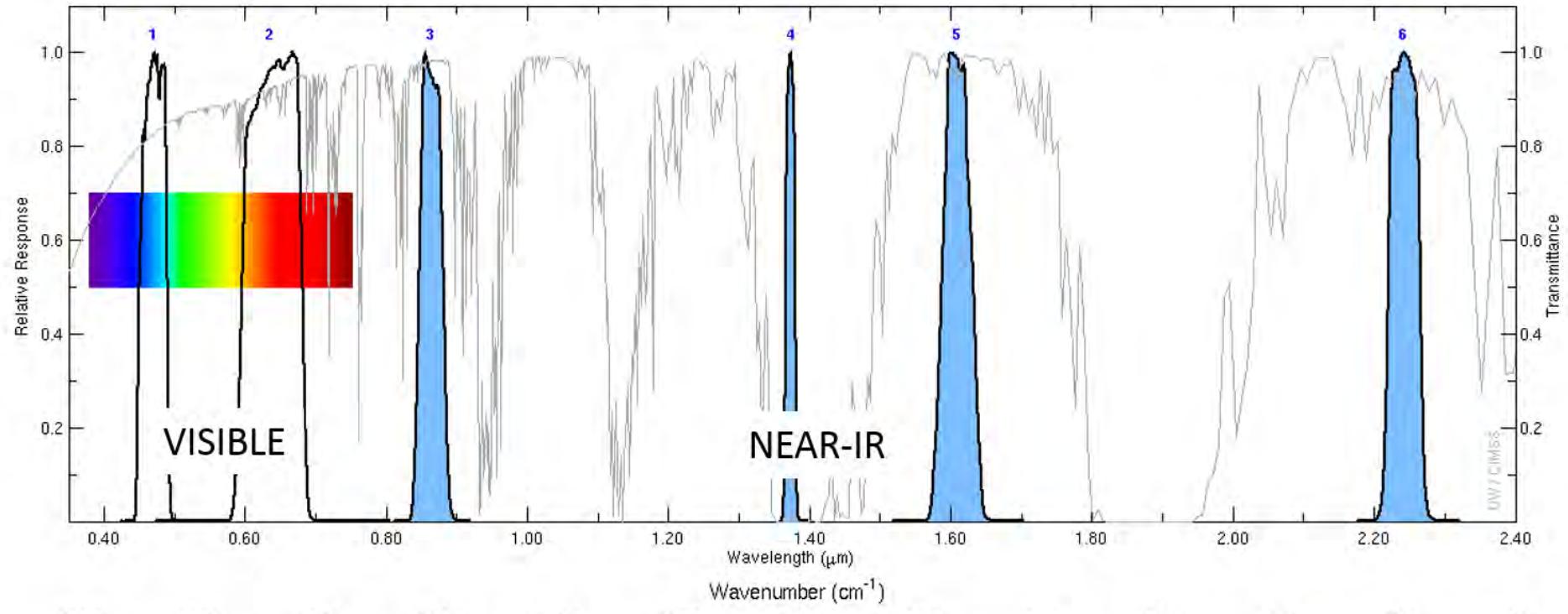


Spectral response functions (SRFs) describe the sensitivity of each spectral band to radiation as a function of wavelength.







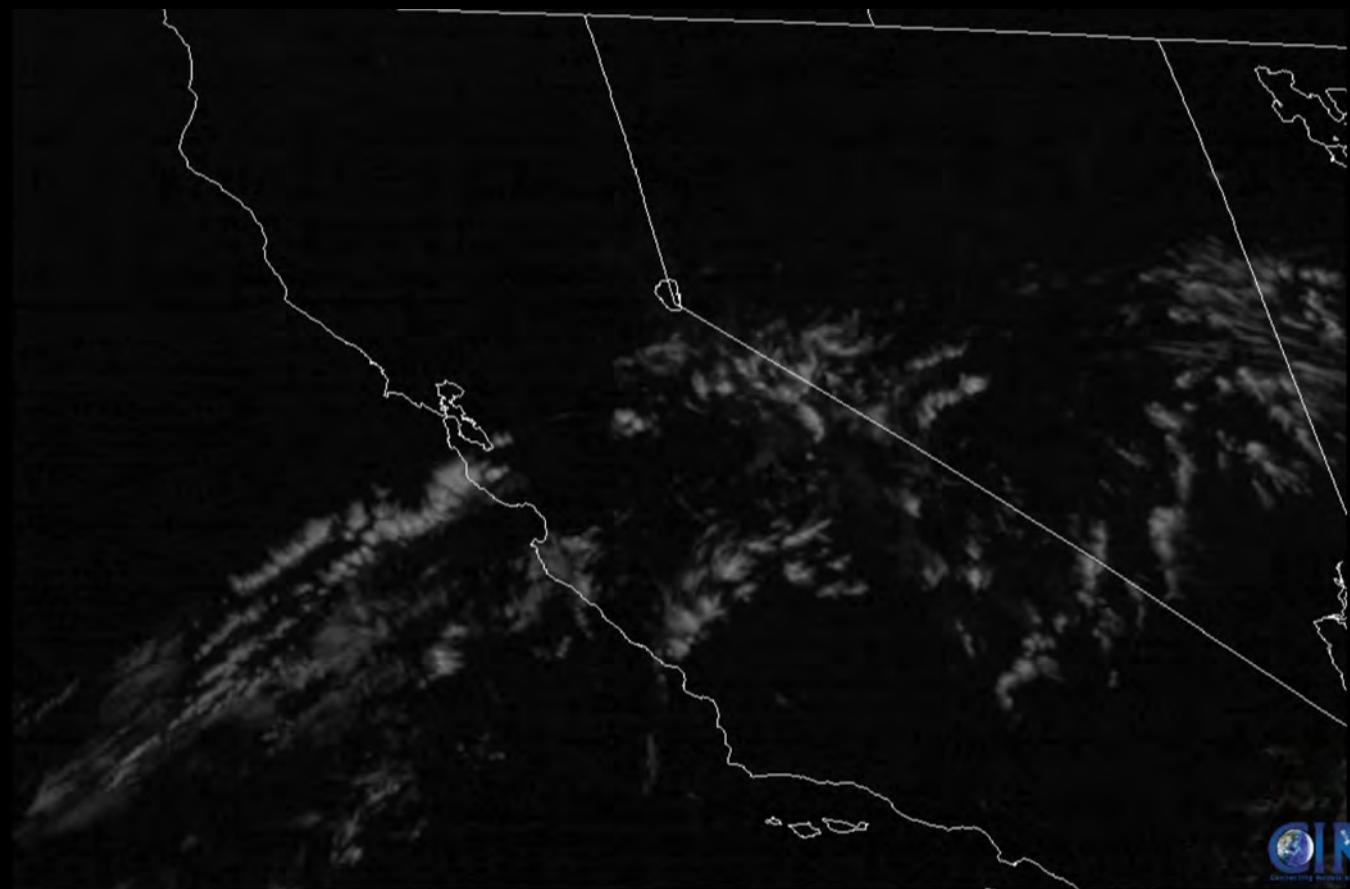


## Band 2 (640 nm)



2020-07-10 15:22:25 UTC

Band 4 (1380 nm)



2020-07-10 15:19:25 UTC



The general radiative transfer solution for a scattering atmosphere is:

$$L_t(\lambda, \theta, \phi) = L_0(\lambda, \theta, \phi) e^{-\delta(\lambda)/\mu} + \int_0^{\delta(\lambda)} \frac{\int_{4\pi} \gamma_s(\mathbf{r}, \mathbf{r}', \lambda, \mathbf{X}) L(\mathbf{r}', \lambda, \mathbf{X}) d\Omega'}{\sigma_e(\lambda, z)} e^{-\delta(\lambda, z)/\mu} \frac{d\delta}{\mu}$$

Where the surface radiance ( $L_0$ ) is due to reflection  
of the downward radiative flux at the surface  
(read section 3.6 - Kidder & Vonder Haar)

This is for  
reflected solar  
radiation only!

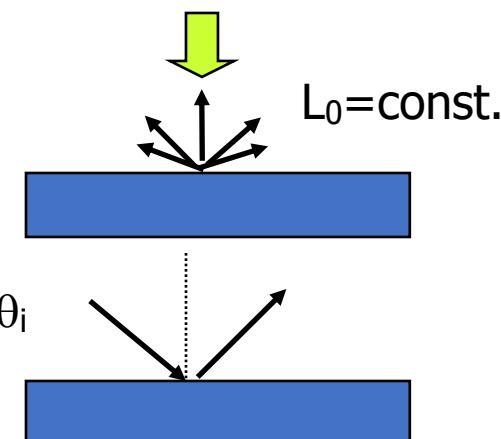
The reflection properties of various surfaces  
can be a complex function of incoming and outgoing directions

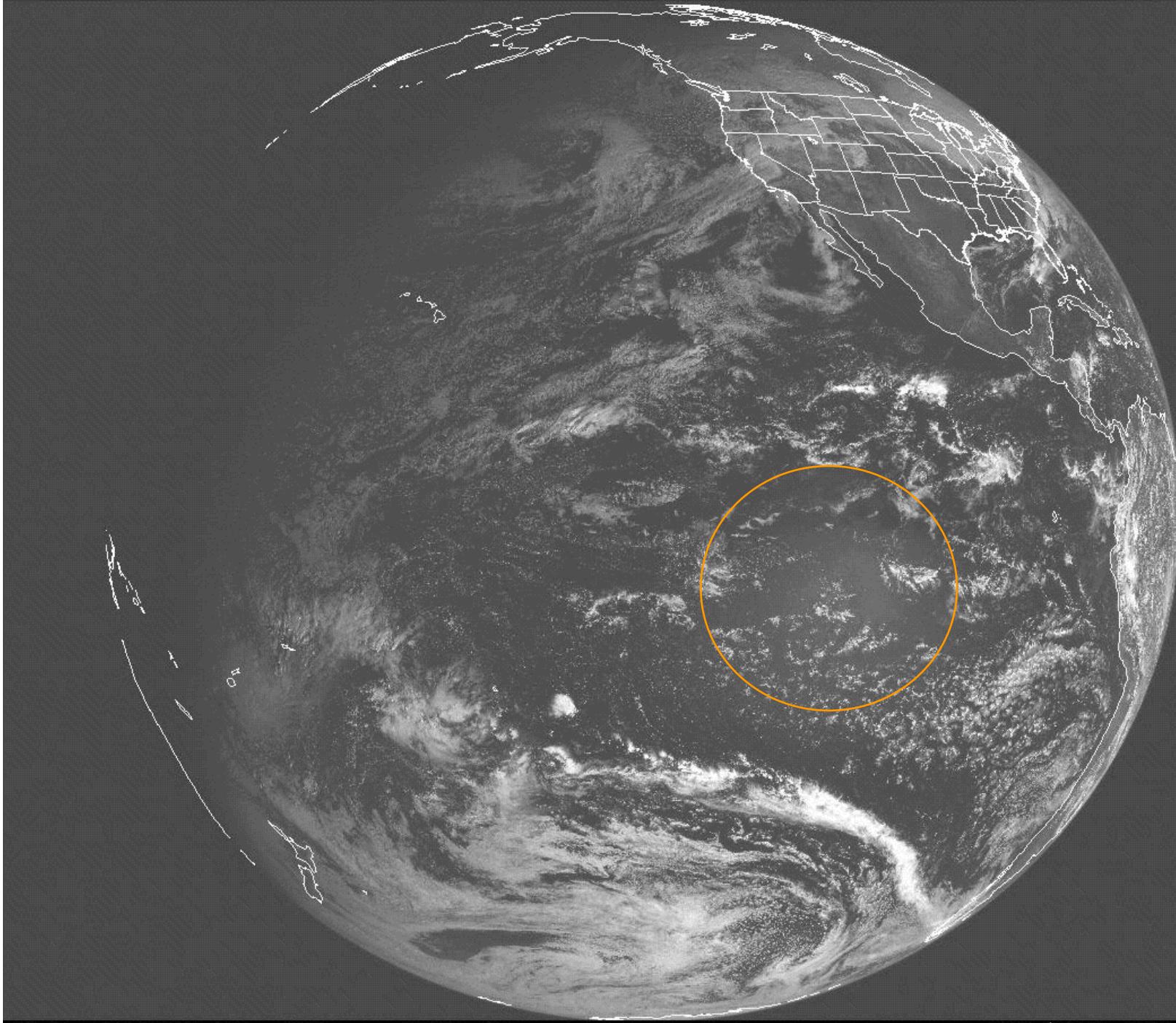
... commonly called the bidirectional reflectance,  $\gamma_r(\theta_r, \phi_r; \theta_i, \phi_i)$

Common approximations to surface reflectance include:

Lambertian or isotropic reflectance

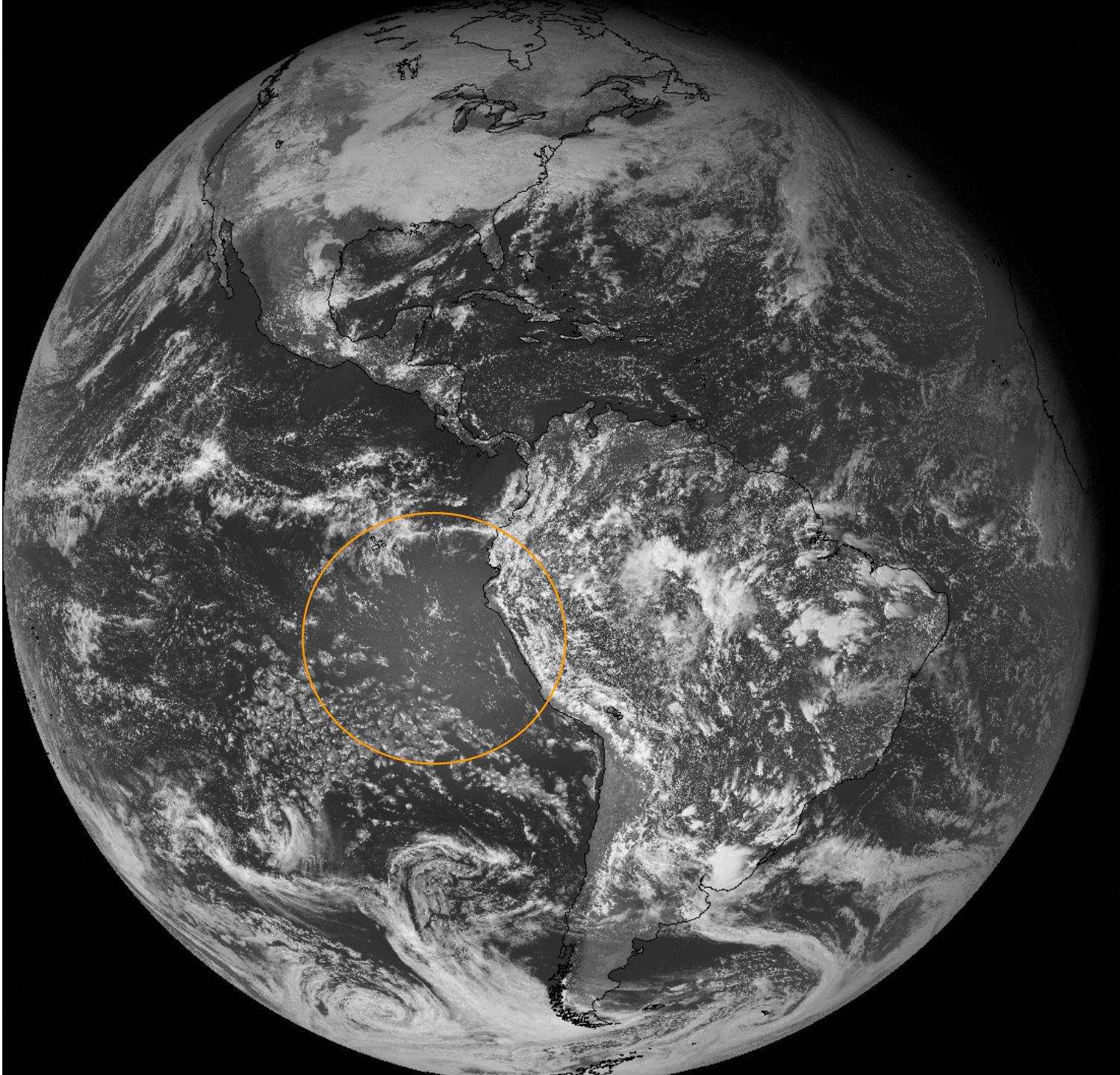
Specular or “mirror-like”



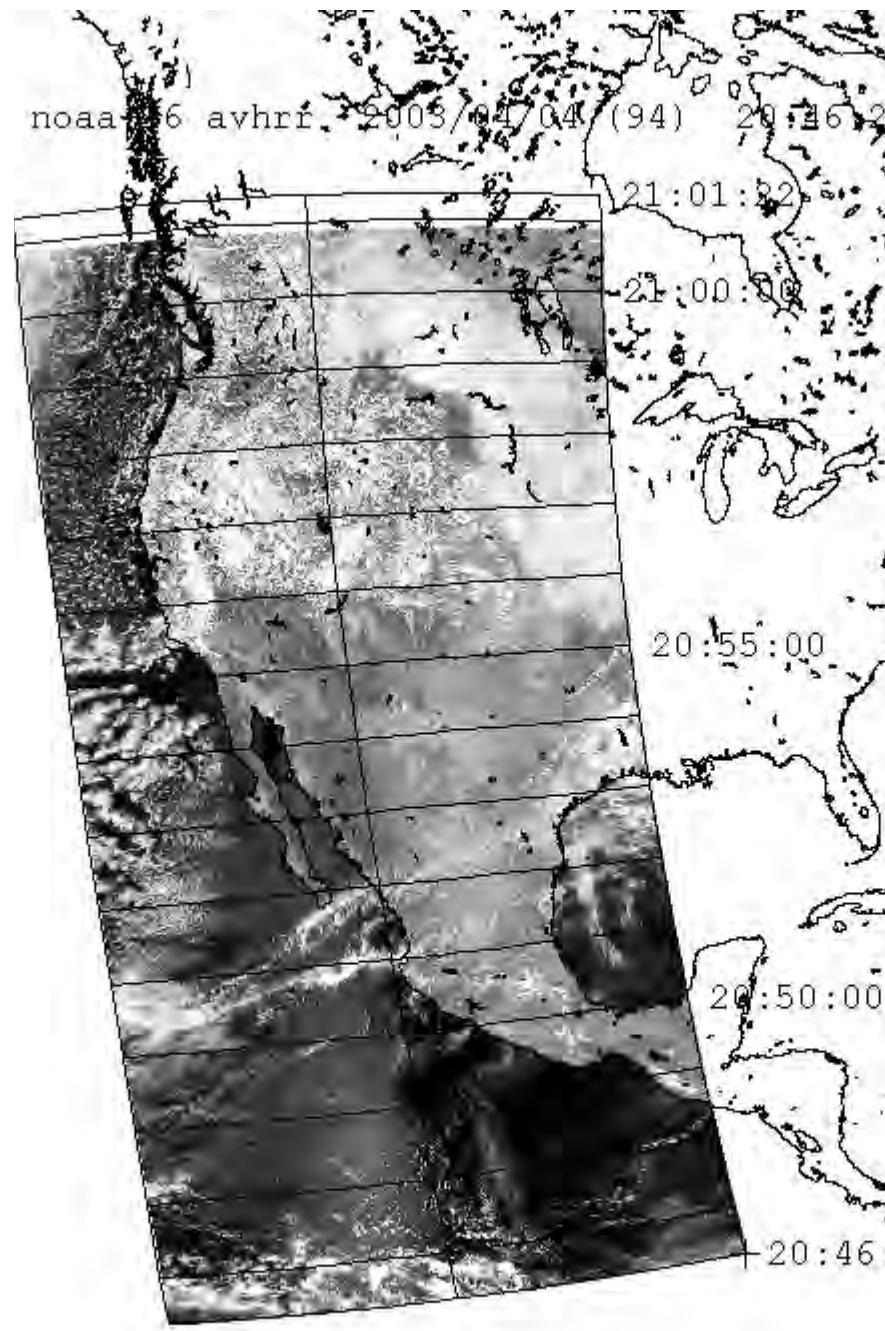


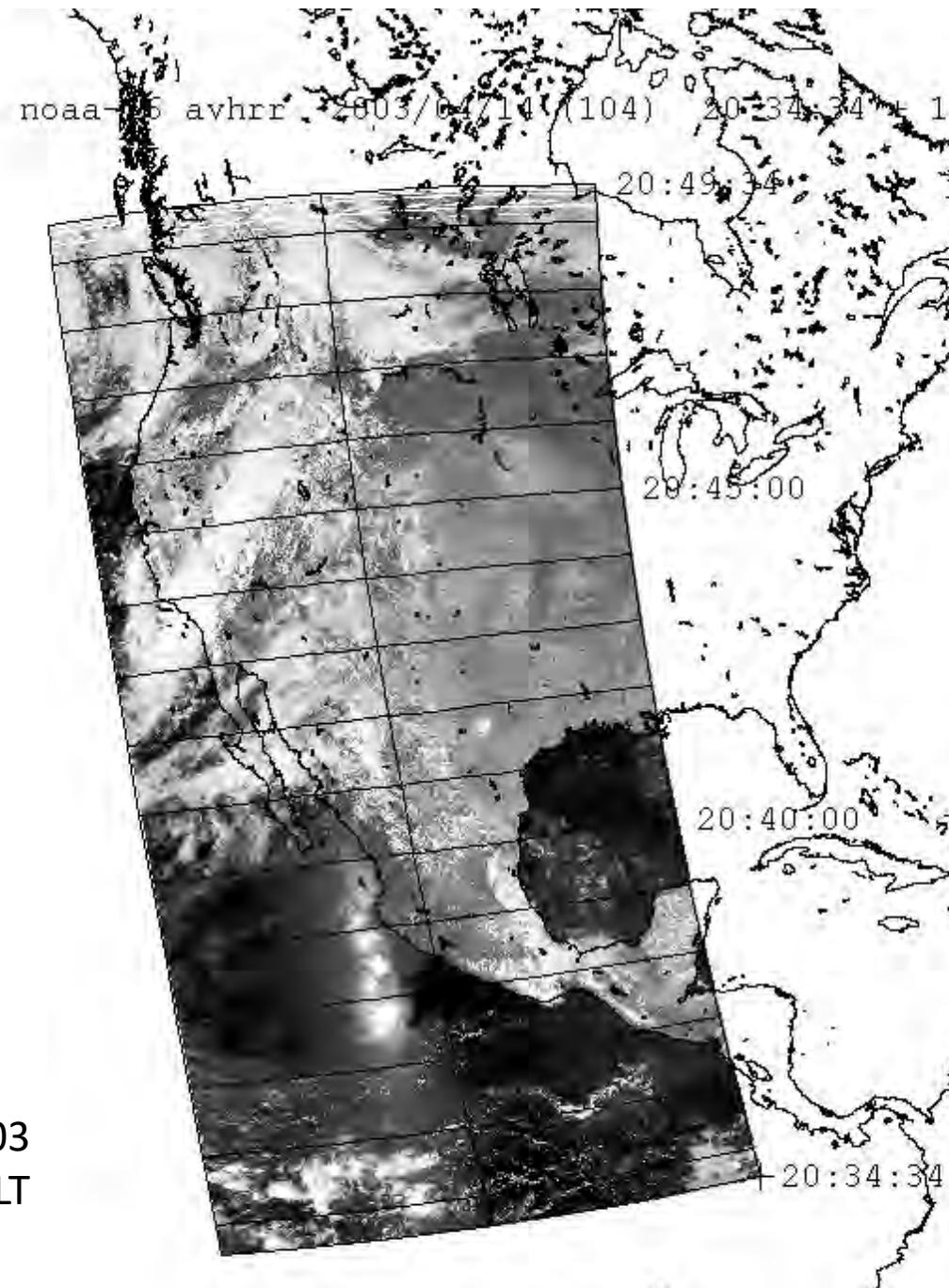
18 Jan 99 18:00:14Z 0.65 um GOES-10

CIRA/NOAA-CSU

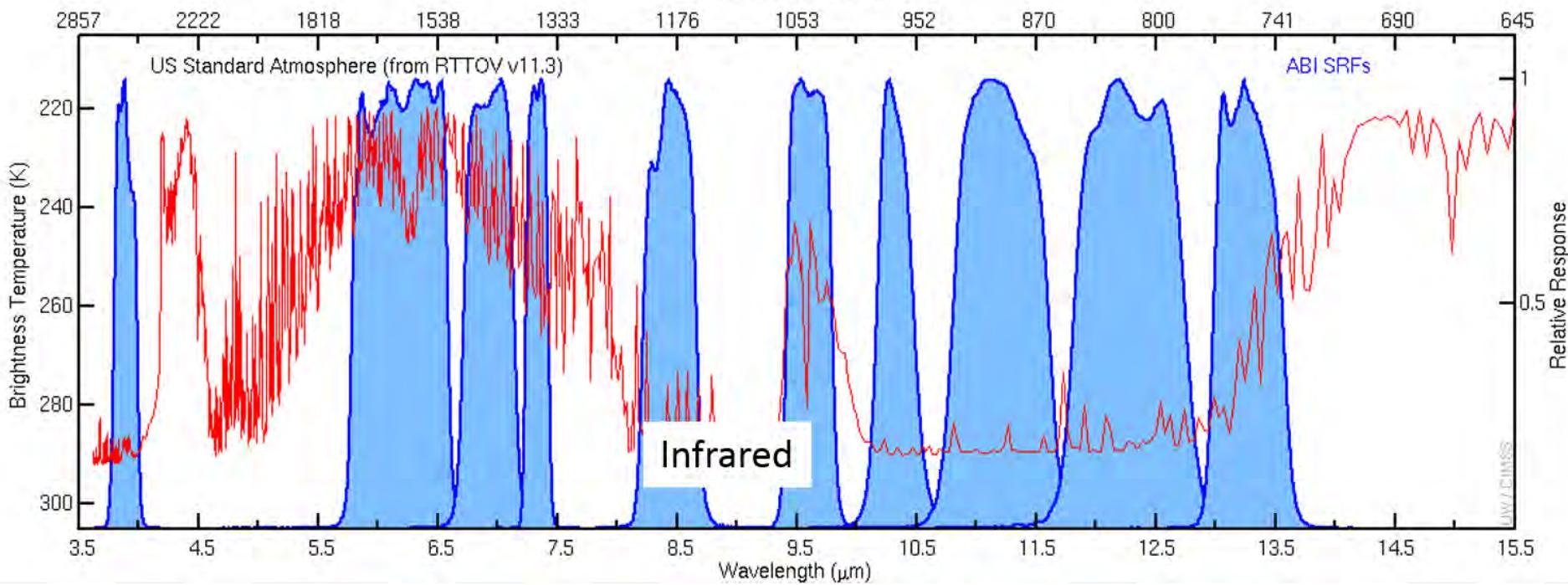
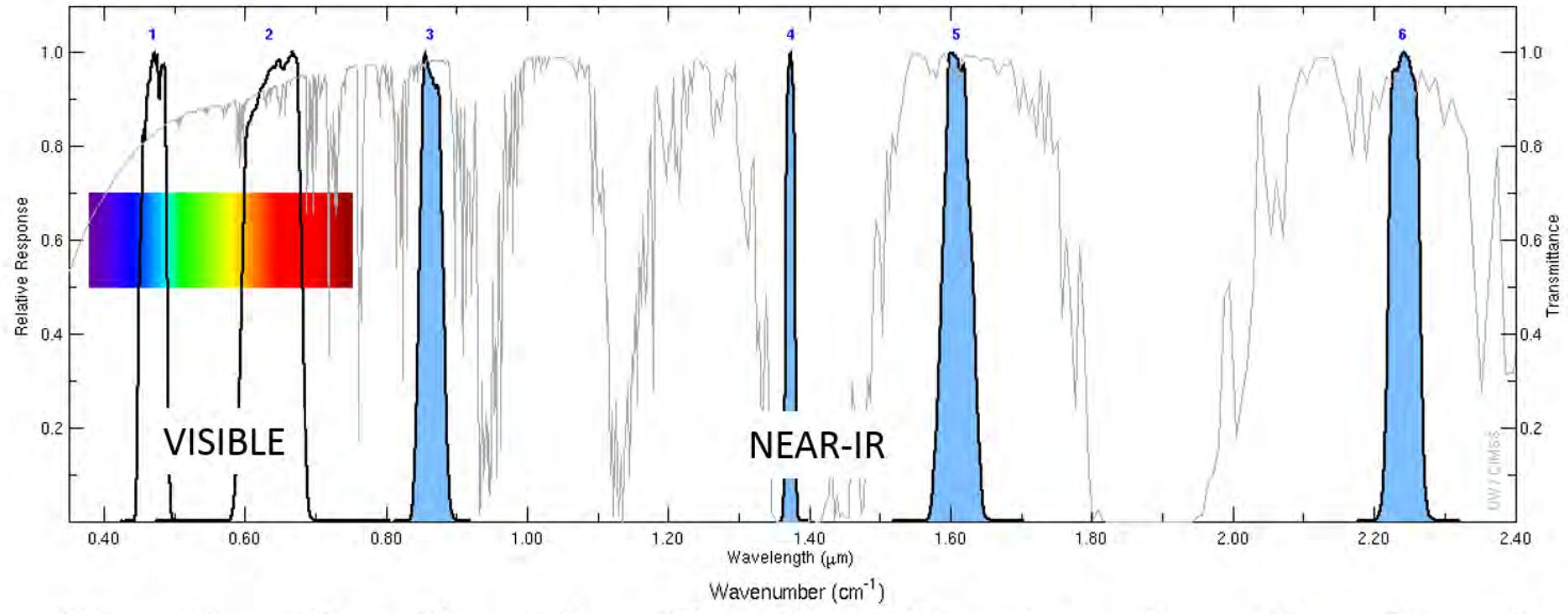


Apr 4, 2003  
13:50 LT





Apr 14, 2003  
13:40 LT



## Band 2 (640 nm)

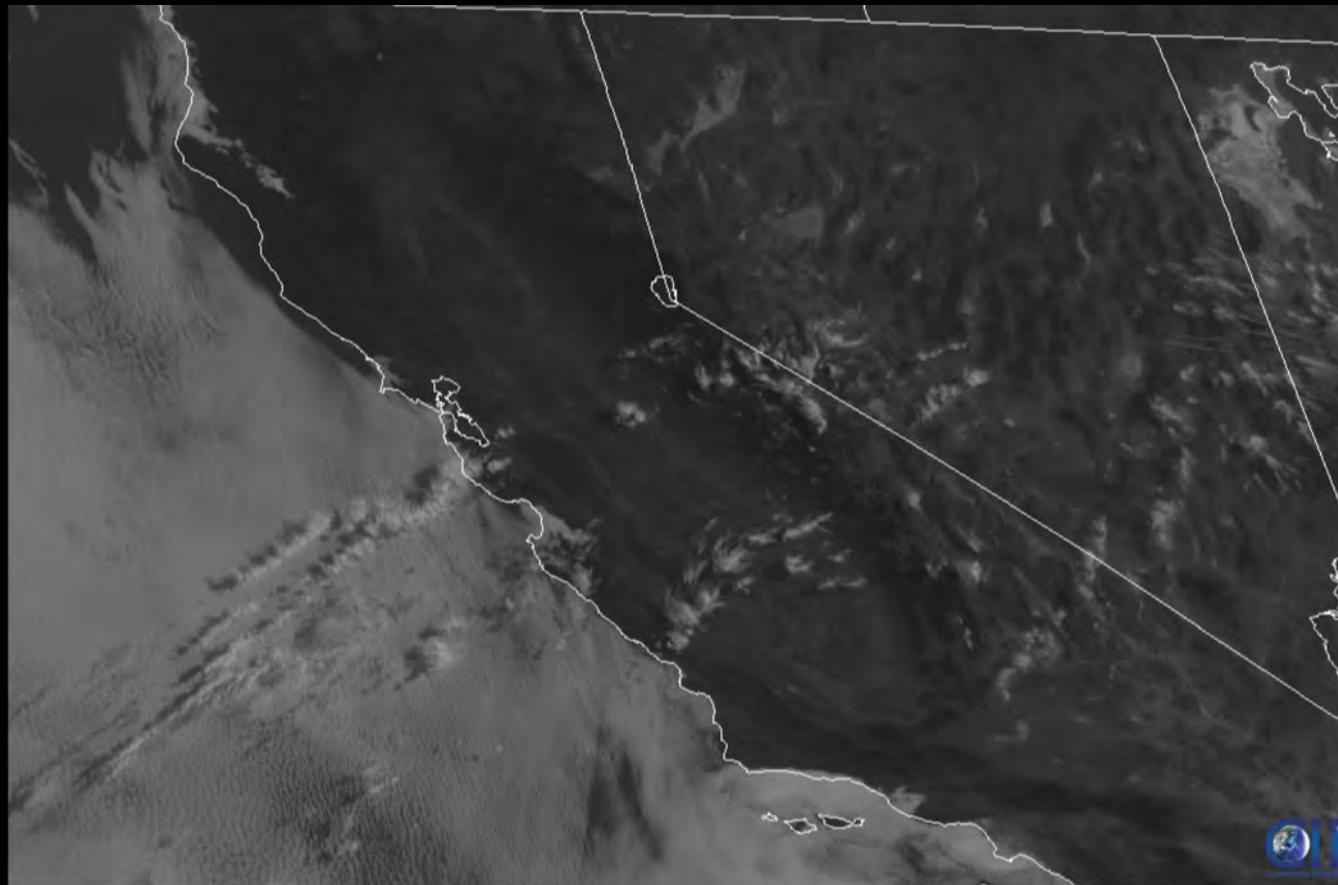


2020-07-10 15:22:25 UTC



[http://cimss.ssec.wisc.edu/goes/OCLOFactSheetPDFs/ABIQuickGuide\\_Band02.pdf](http://cimss.ssec.wisc.edu/goes/OCLOFactSheetPDFs/ABIQuickGuide_Band02.pdf)

## Band 1 (470 nm)

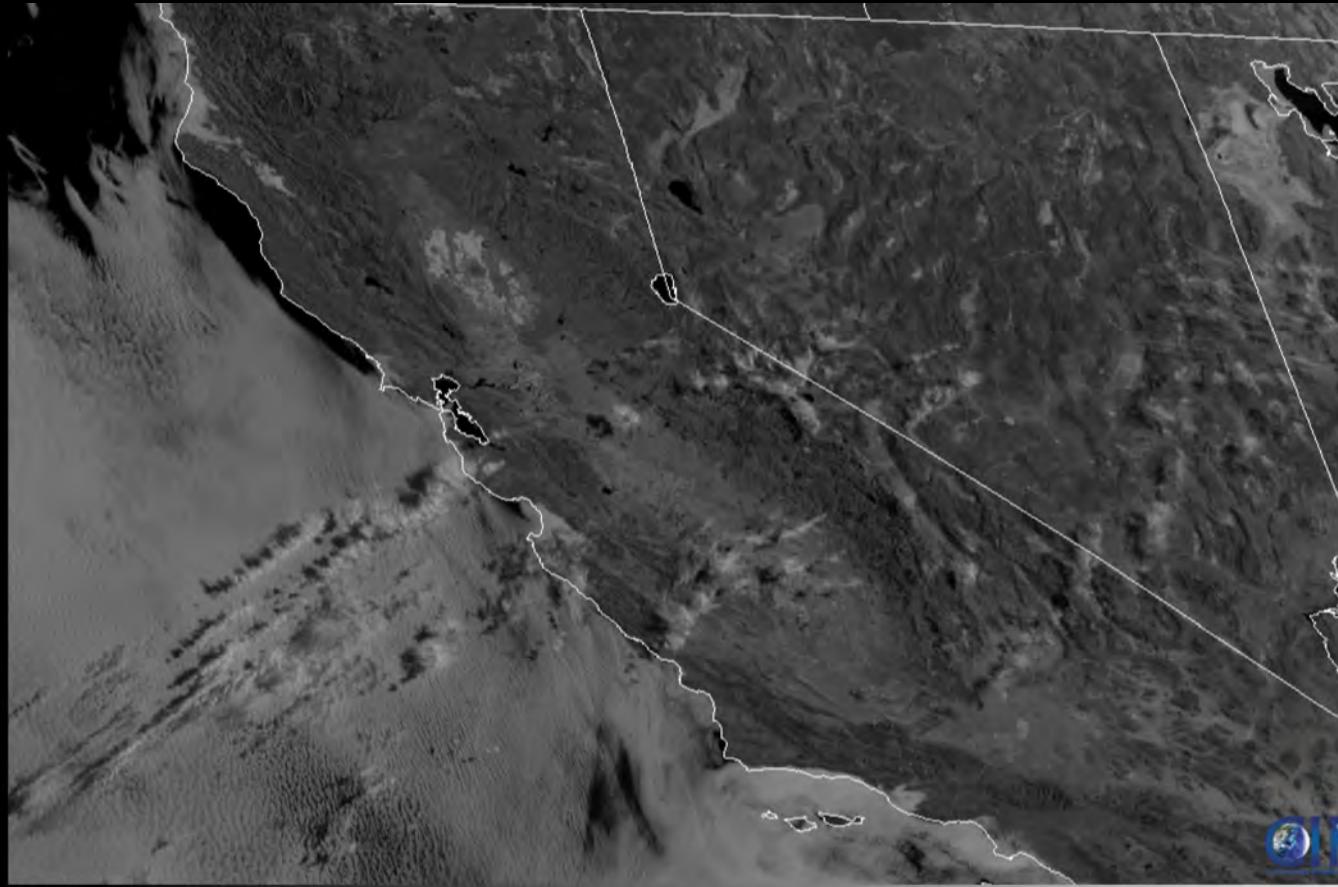


2020-07-10 15:25:25 UTC

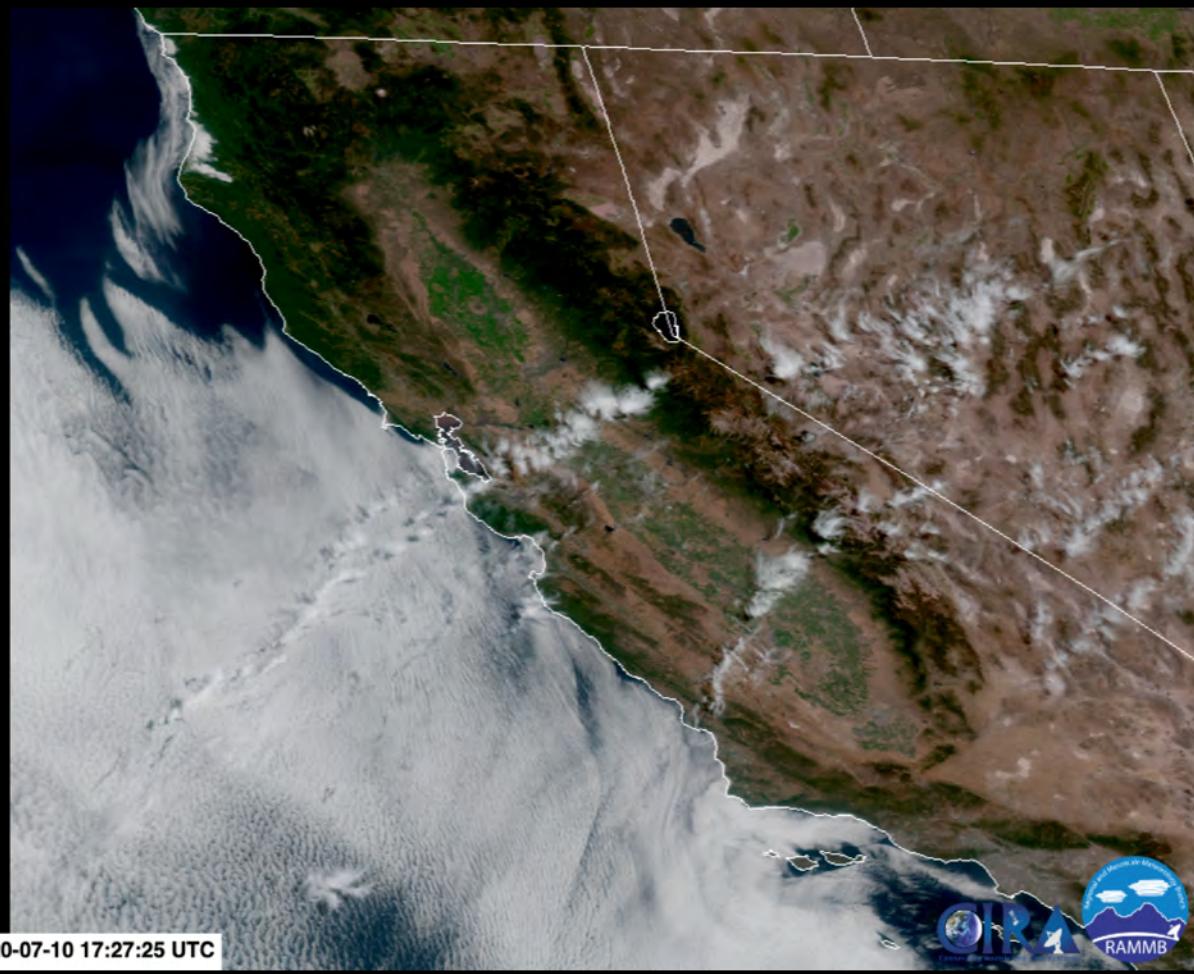


[http://cimss.ssec.wisc.edu/goes/OCLOFactSheetPDFs/ABIQuickGuide\\_Band01.pdf](http://cimss.ssec.wisc.edu/goes/OCLOFactSheetPDFs/ABIQuickGuide_Band01.pdf)

Band 3 (860 nm)



[http://cimss.ssec.wisc.edu/goes/OCLOFactSheetPDFs/ABIQuickGuide\\_Band03.pdf](http://cimss.ssec.wisc.edu/goes/OCLOFactSheetPDFs/ABIQuickGuide_Band03.pdf)



2020-07-10 17:27:25 UTC



Band 5 (1.61  $\mu$ m)



2020-07-10 15:16:25 UTC



Band 6 (2.24  $\mu$ m)



2020-07-10 15:21:25 UTC

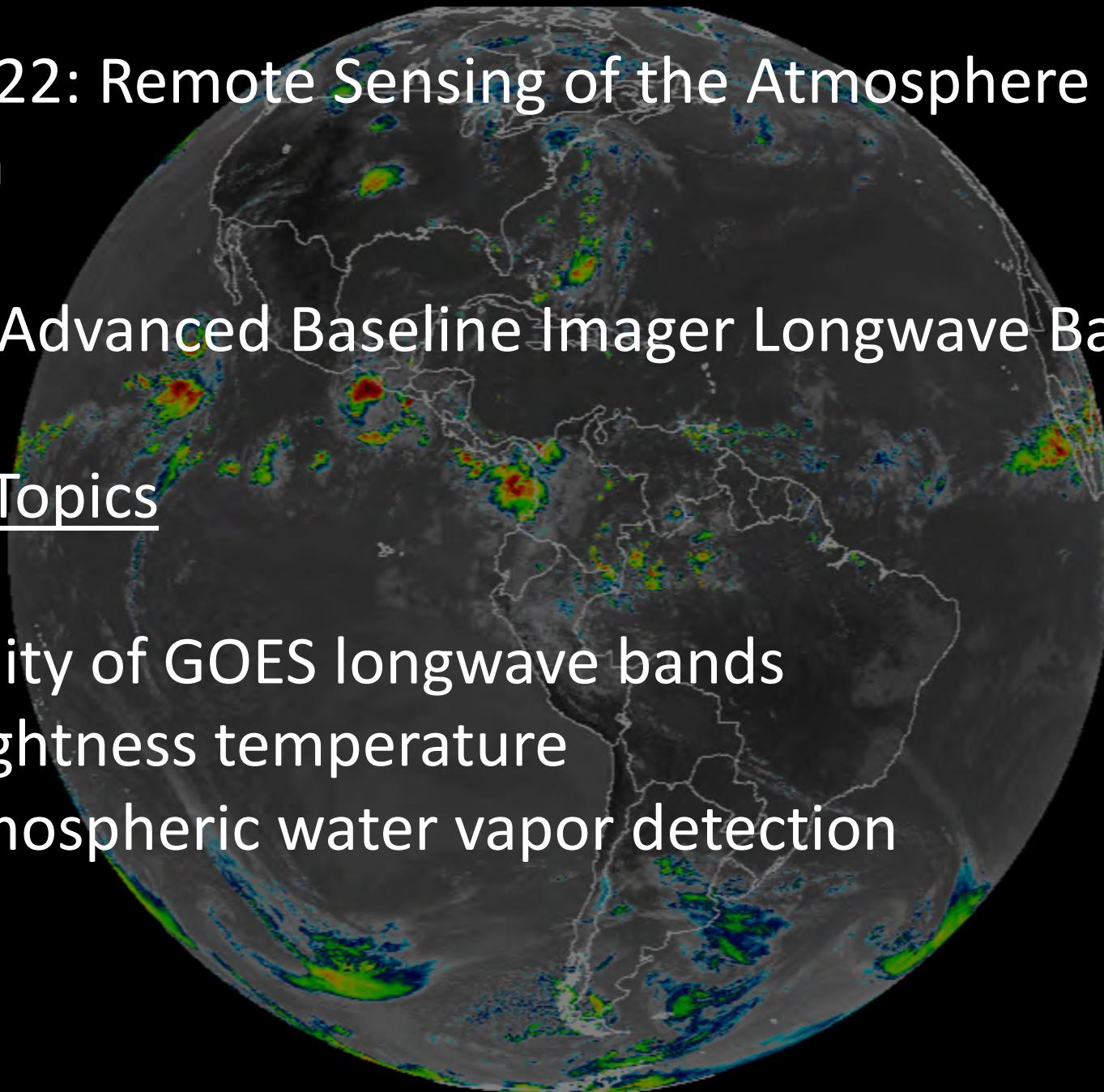


# MR3522: Remote Sensing of the Atmosphere and Ocean

## GOES Advanced Baseline Imager Longwave Bands

### Main Topics

- Utility of GOES longwave bands
- Brightness temperature
- Atmospheric water vapor detection



## Idealized Case #2

Emitted Path Radiance Only

Here emission is the only source of photons and there is **no scattering**, so  $\sigma_e = \sigma_a$ .

$$J_\lambda(z) = \sigma_{a,\lambda}(z) B_\lambda(T(z))$$

$$\varepsilon_{s,\lambda} = \sigma_{a,\lambda} \longrightarrow L(\delta_t; \mu, \phi) = \varepsilon_{s,\lambda} B_\lambda(T_s)$$



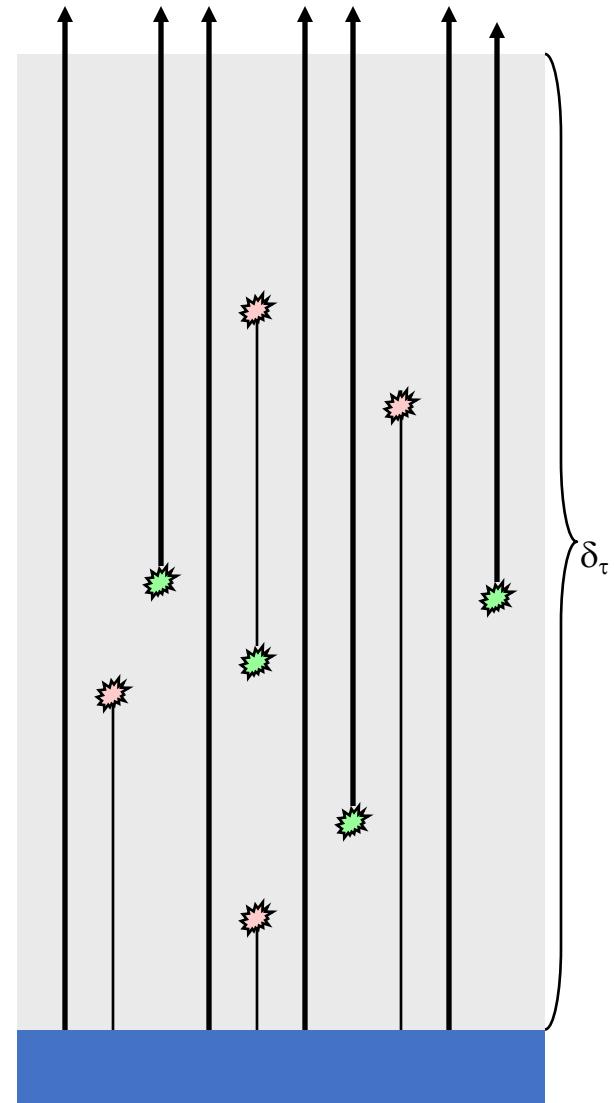
Kirchhoff's Law

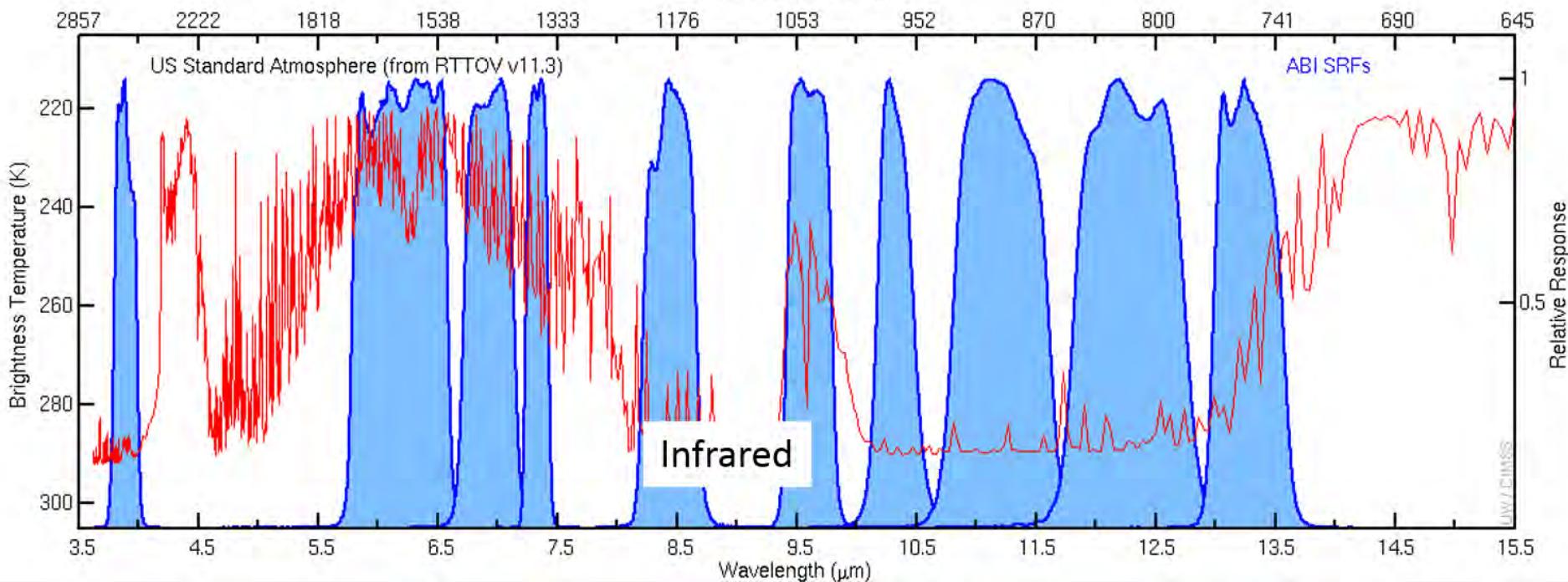
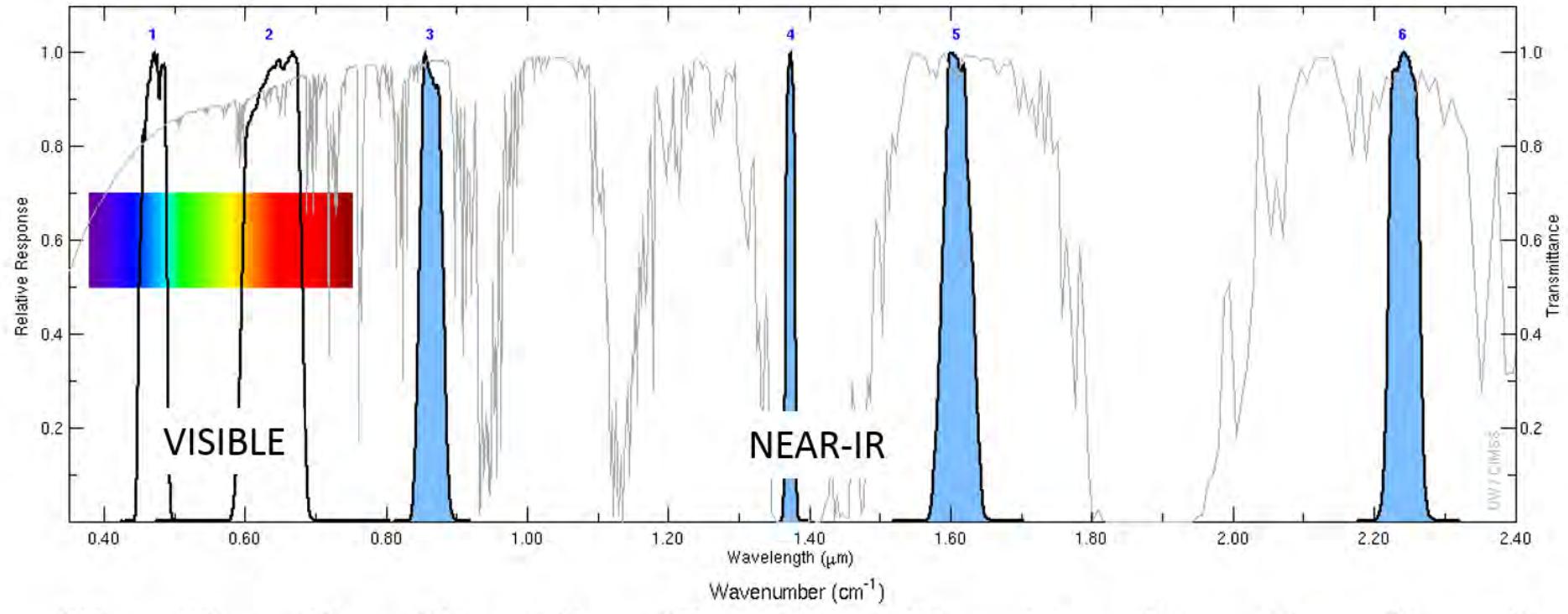
Solution:

$$L(0; \mu, \phi) = L(\delta_t; \mu, \phi) e^{-\delta_t/\mu} + \int_0^{\delta_t} \frac{J(\delta'; \mu, \phi)}{\sigma_e(\delta')} e^{-\delta'/\mu} \frac{d\delta'}{\mu}$$



$$L(0; \mu, \phi) = \varepsilon_{s,\lambda} B_\lambda(T_s) e^{-\delta_t/\mu} + \int_0^{\delta_t} B_\lambda(T(z)) e^{-\delta'/\mu} \frac{d\delta'}{\mu}$$

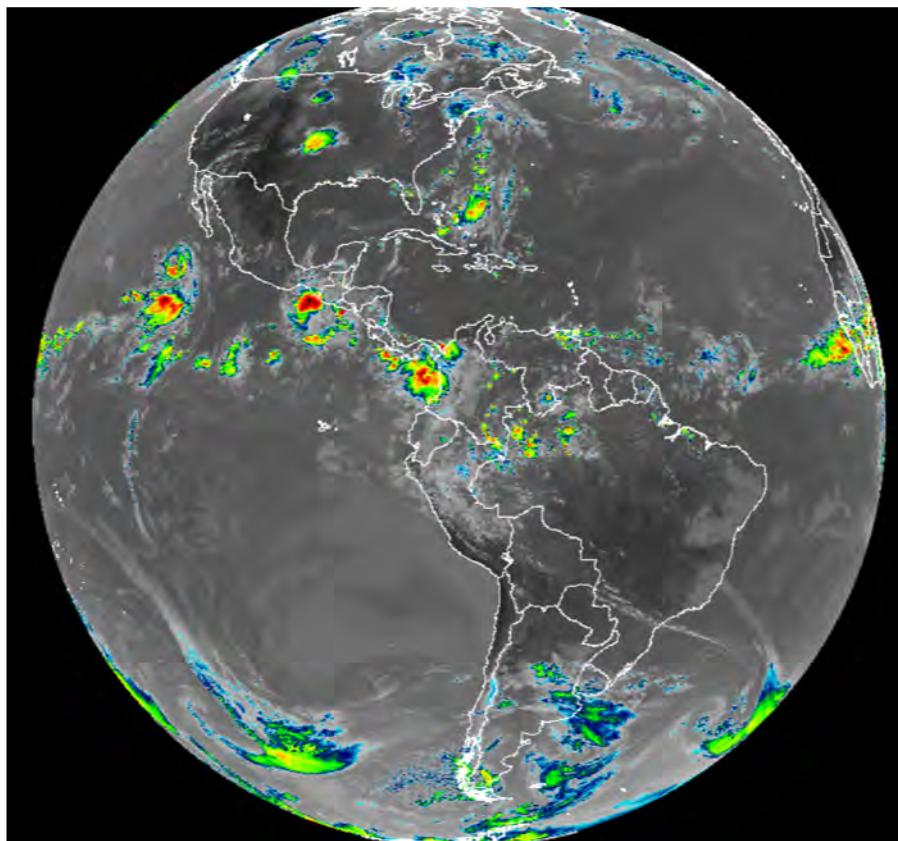
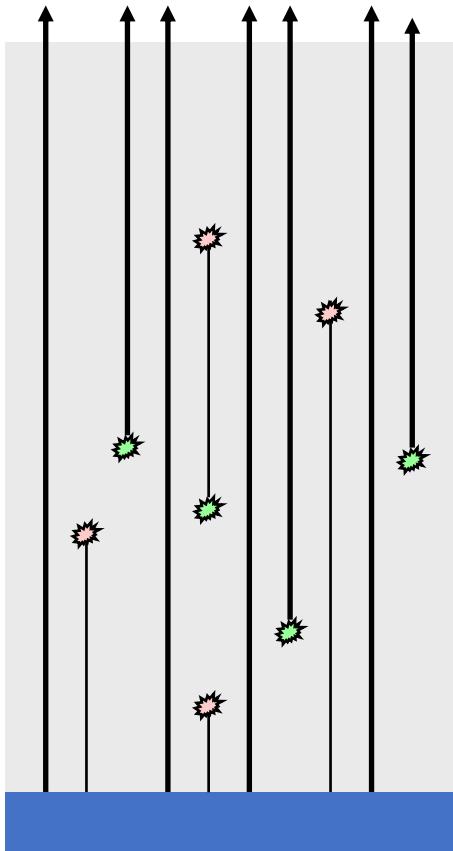




Brightness temperature is defined as  
the temperature a blackbody would need  
to emit the radiance measured by a satellite ( $L_t$ )

That is, if  $L_t$  was emitted by a blackbody, what would its temperature be?

$$L_t \rightarrow B(\lambda, T_B) = \frac{2\hbar c^2}{\lambda^5(e^{\hbar c/\lambda kT_B} - 1)} \quad \text{and} \quad T_B = \frac{\hbar c}{\lambda k \ln \frac{2\hbar c^2}{\lambda^5 L_t}}$$



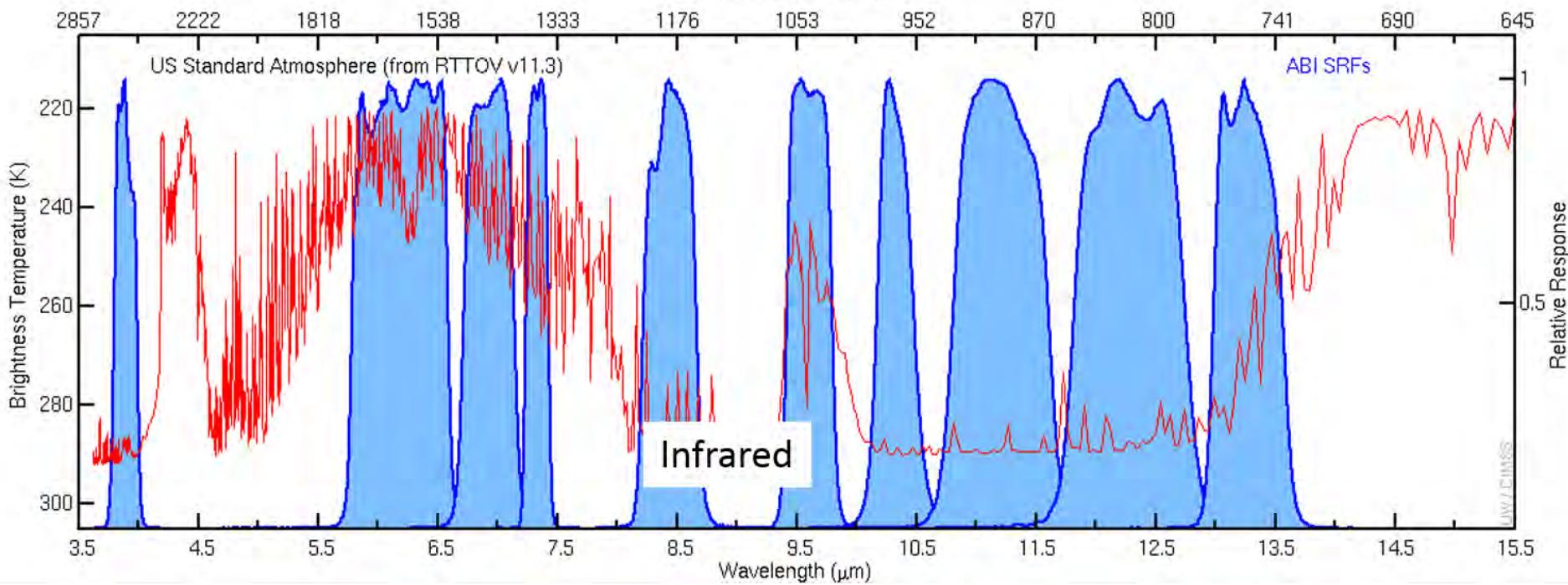
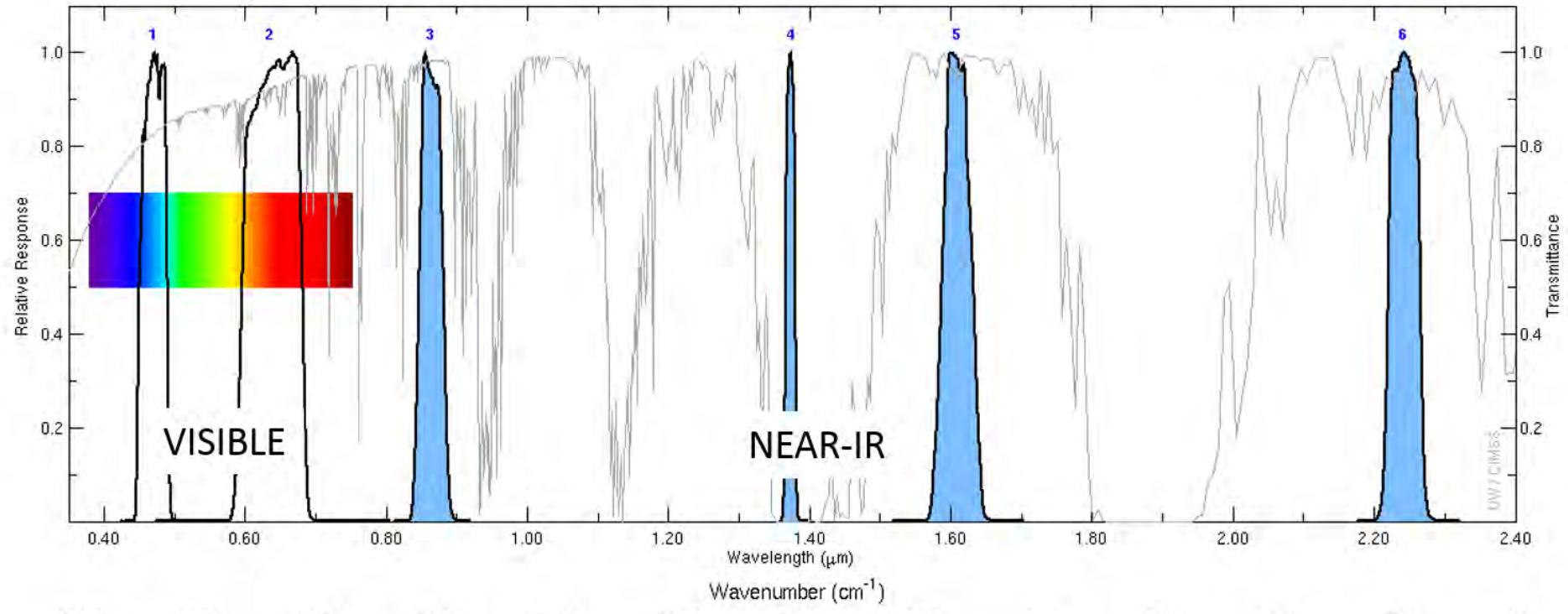
$T_B$   
(from GOES Ch 13)

## Brightness temperature is not necessarily actual temperature!

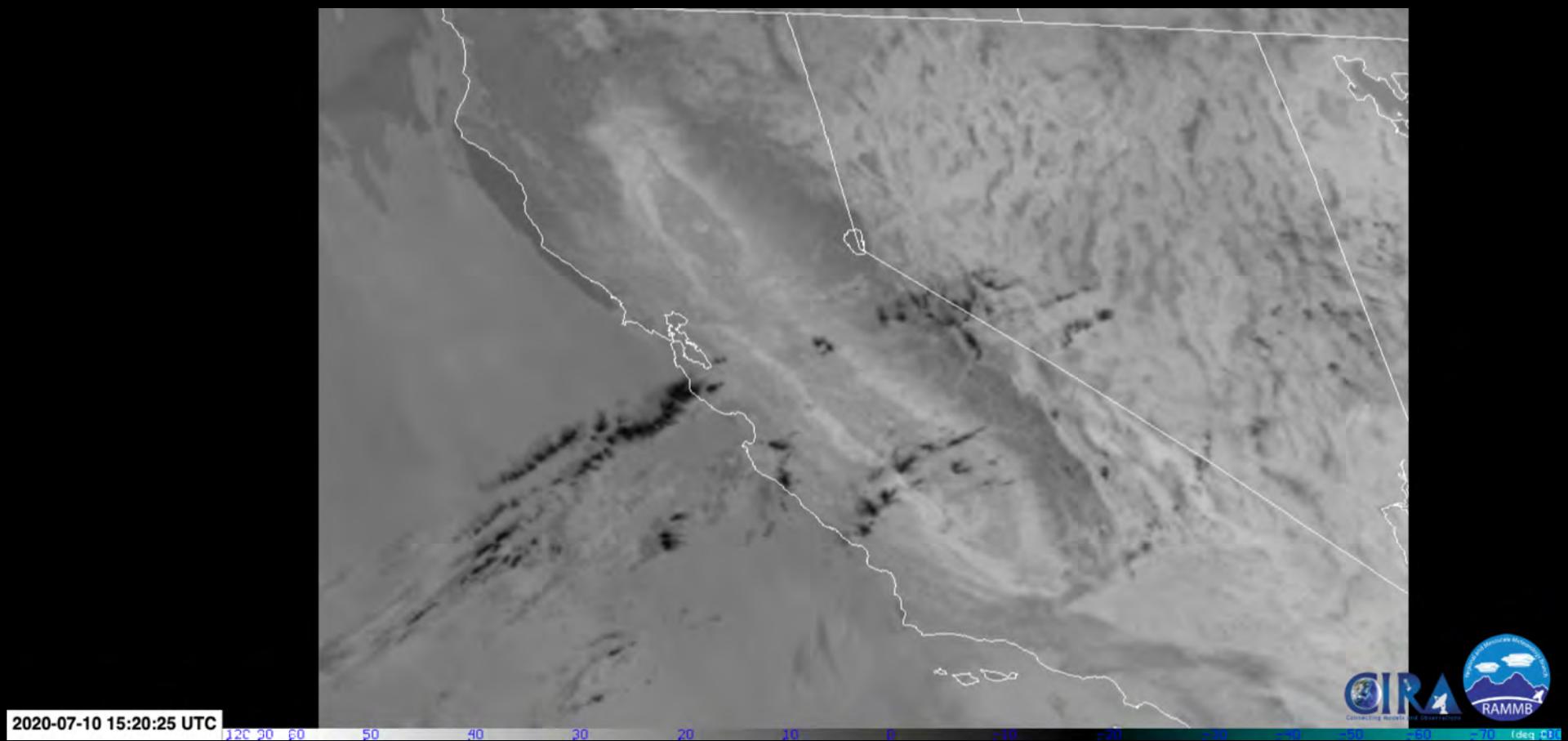
1. The emitting body might not be a blackbody.
2. Radiation could be depleted along the path to the sensor.

$$L_t(\lambda, \theta, \varphi) = \varepsilon_s(\lambda, \theta)B(\lambda, T_s)\tau_d(\lambda) + \int_p^0 B(\lambda, T(p)) \frac{d\tau_d(\lambda, p)}{dp} dp$$

For a cloud (an efficient absorber and emitter if  $\sim 100$  meters or more thick) with an optically thin atmosphere (to the wavelength of interest) above it, the brightness temperature is approximately the cloud top temperature. Knowledge of the atmospheric sounding could help determine cloud top height.



Band 7 (3.90 μm)

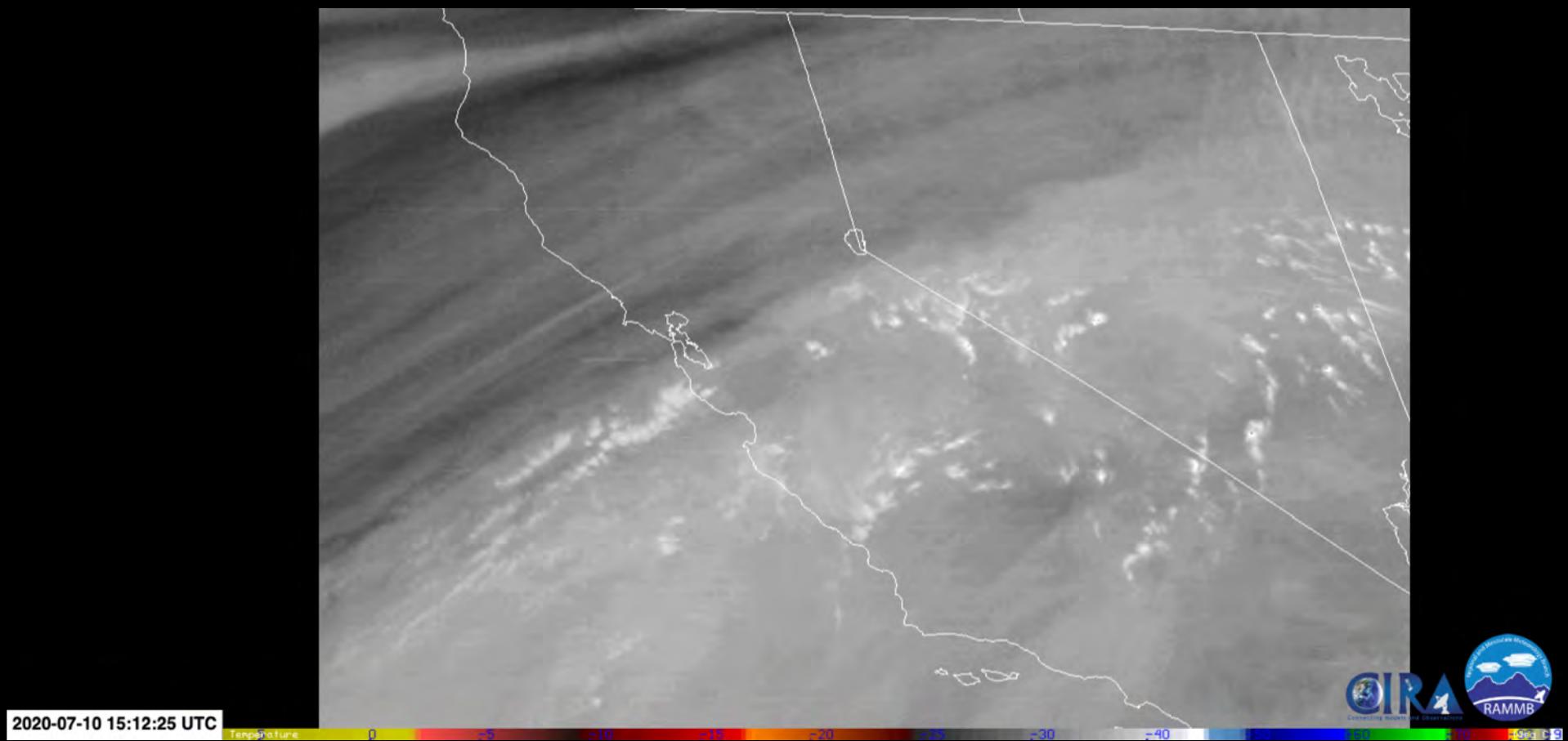


2020-07-10 15:20:25 UTC

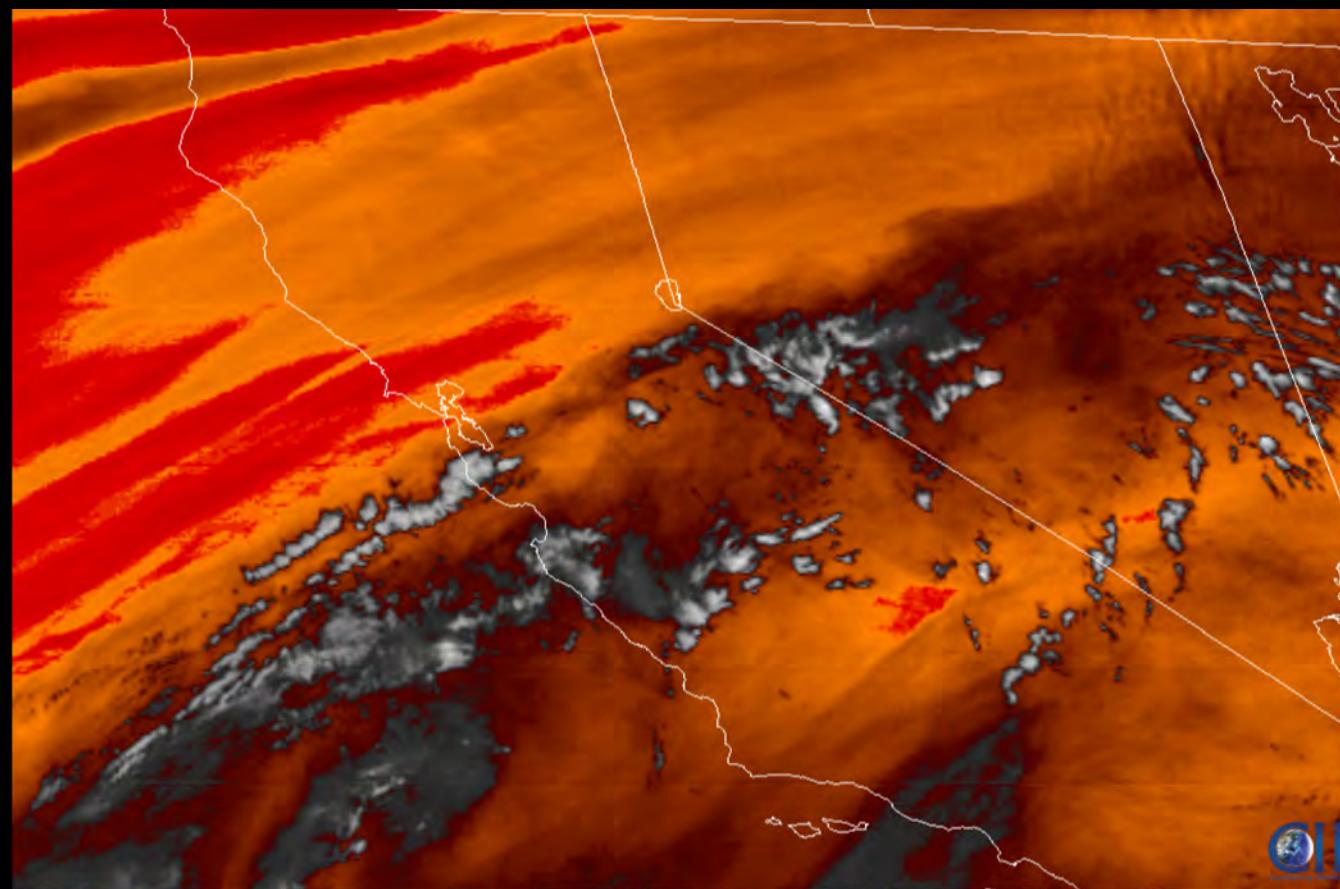
12C 20 20 50 30 20 10 0 -10 -20 -30 -40 -50 -60 -70 (deg C)



Band 8 (6.19 μm)



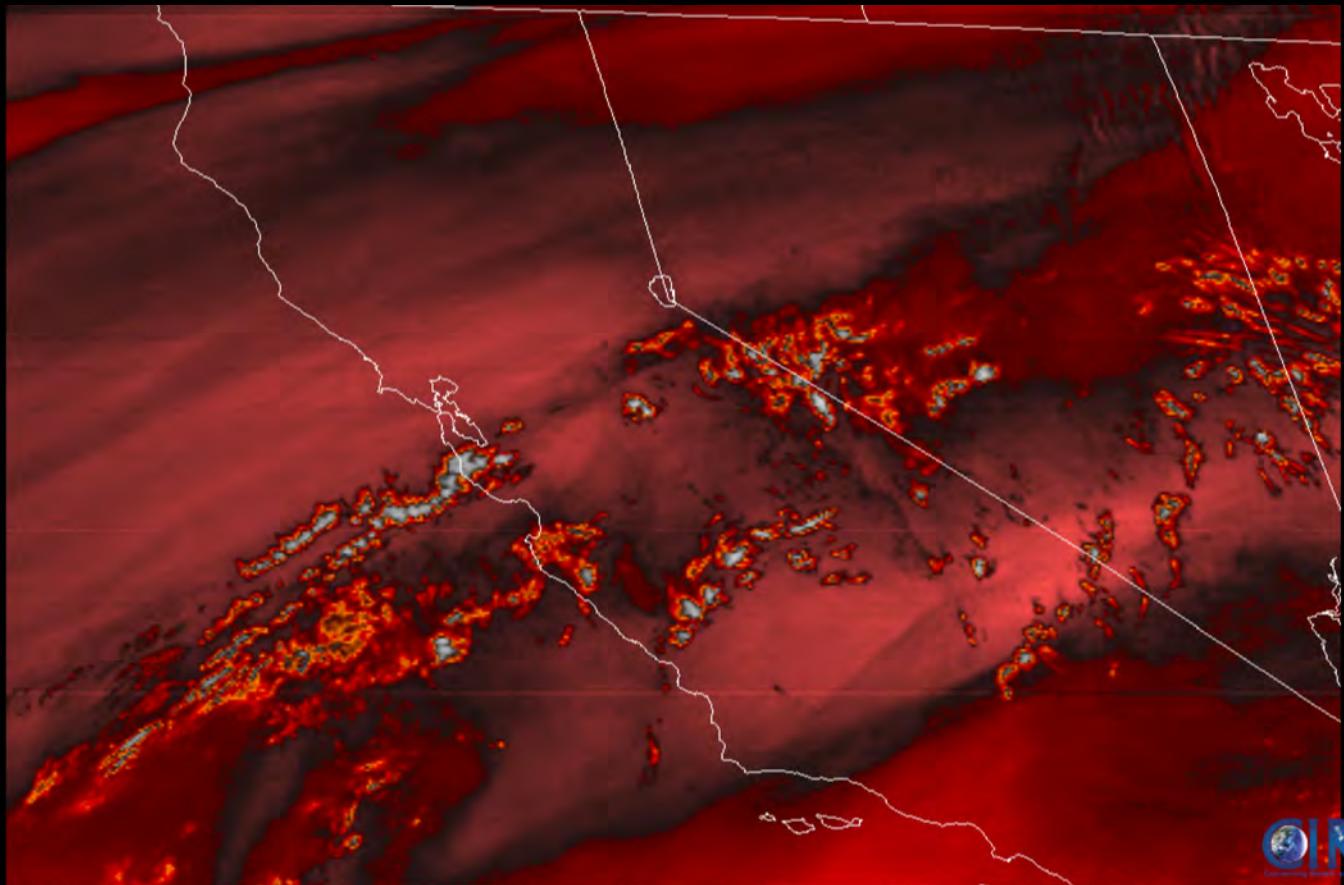
Band 9 (6.95 μm)



2020-07-10 15:31:25 UTC



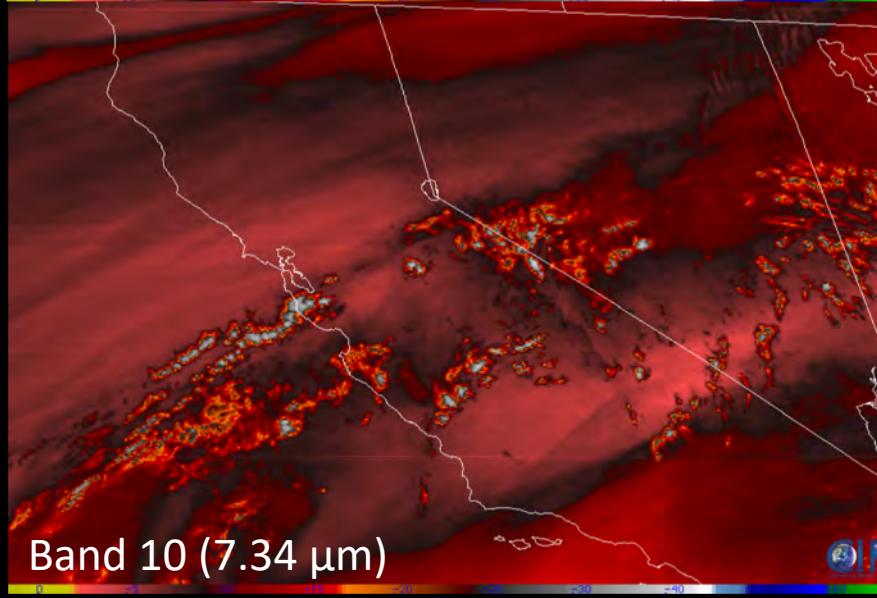
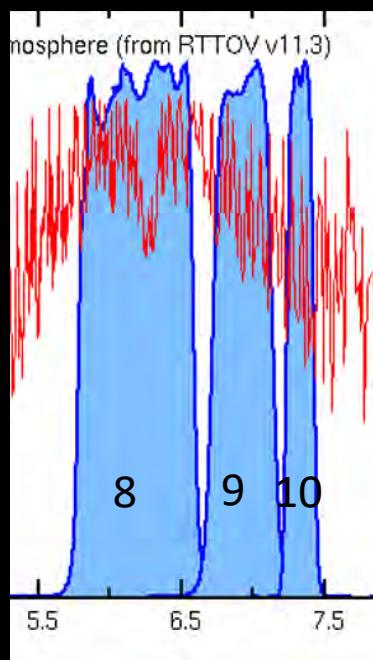
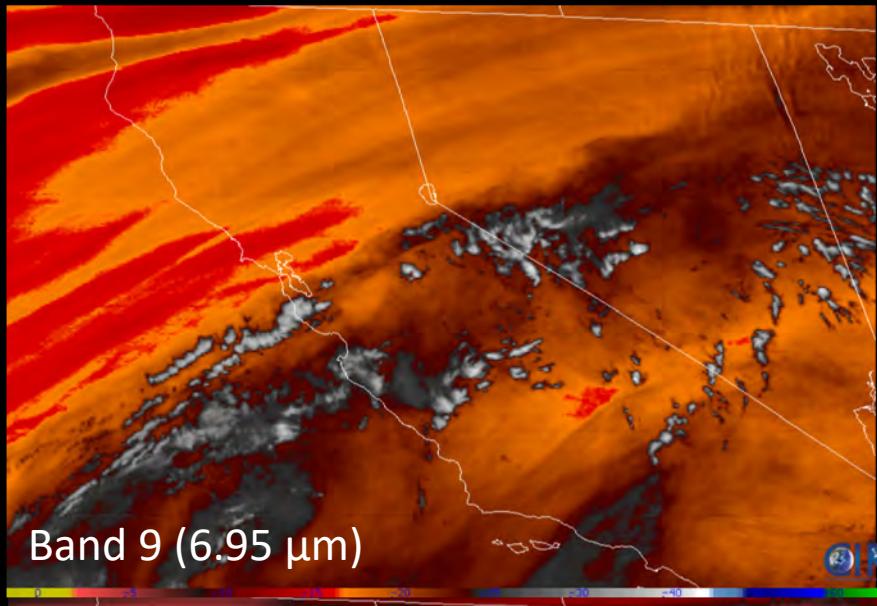
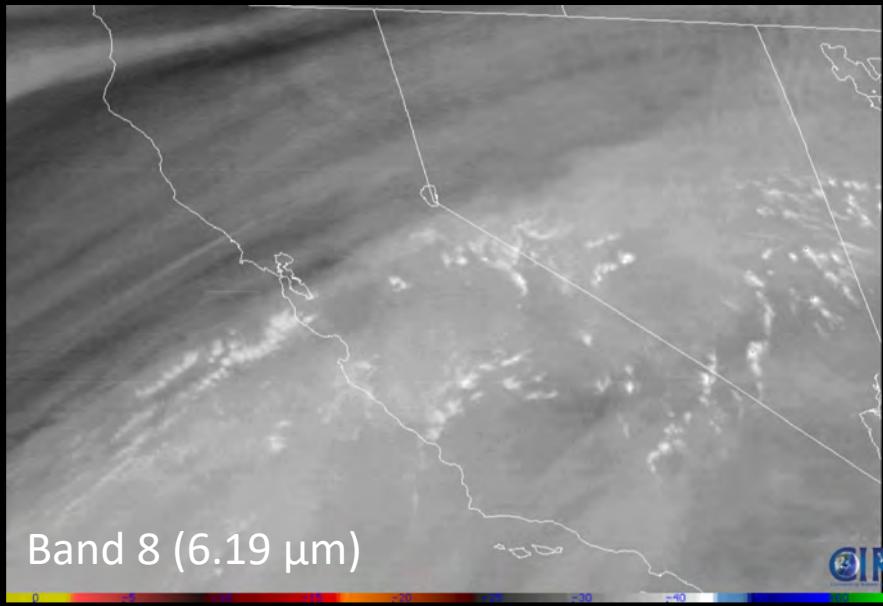
Band 10 (7.34 μm)



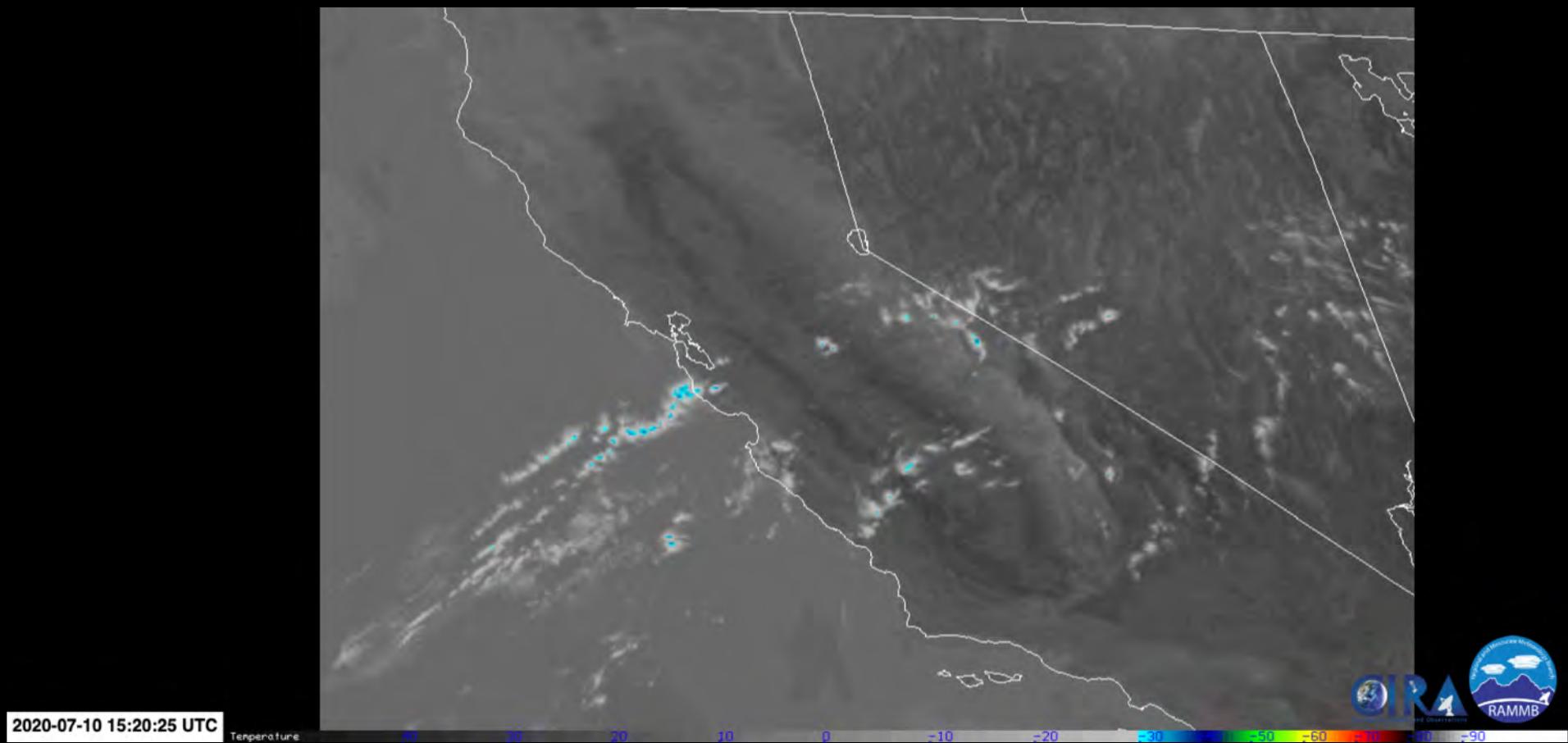
2020-07-10 15:33:25 UTC

Temp (°C) 0 -5 -10 -15 -20 -25 -30 -35 -40 50 60 70 80 90

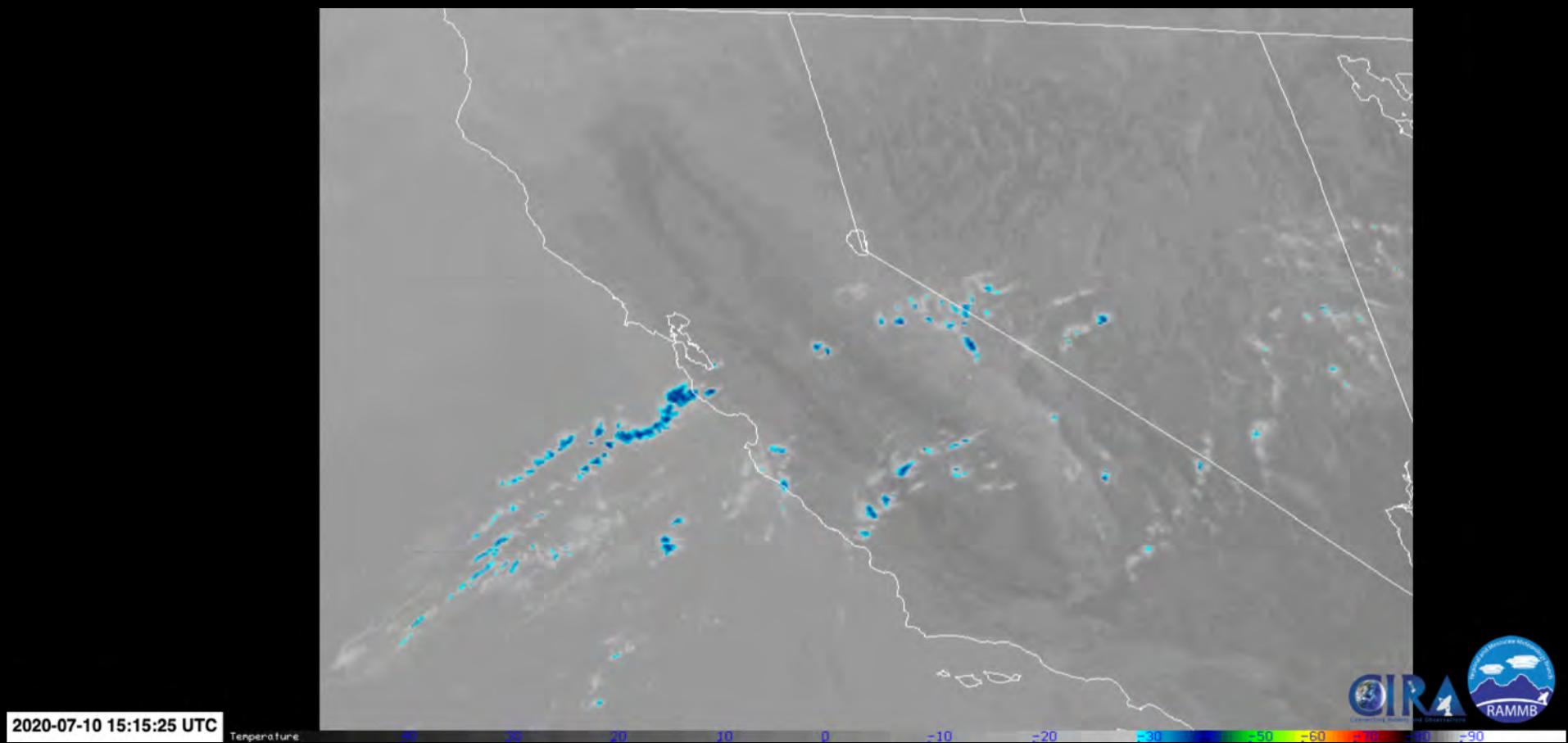




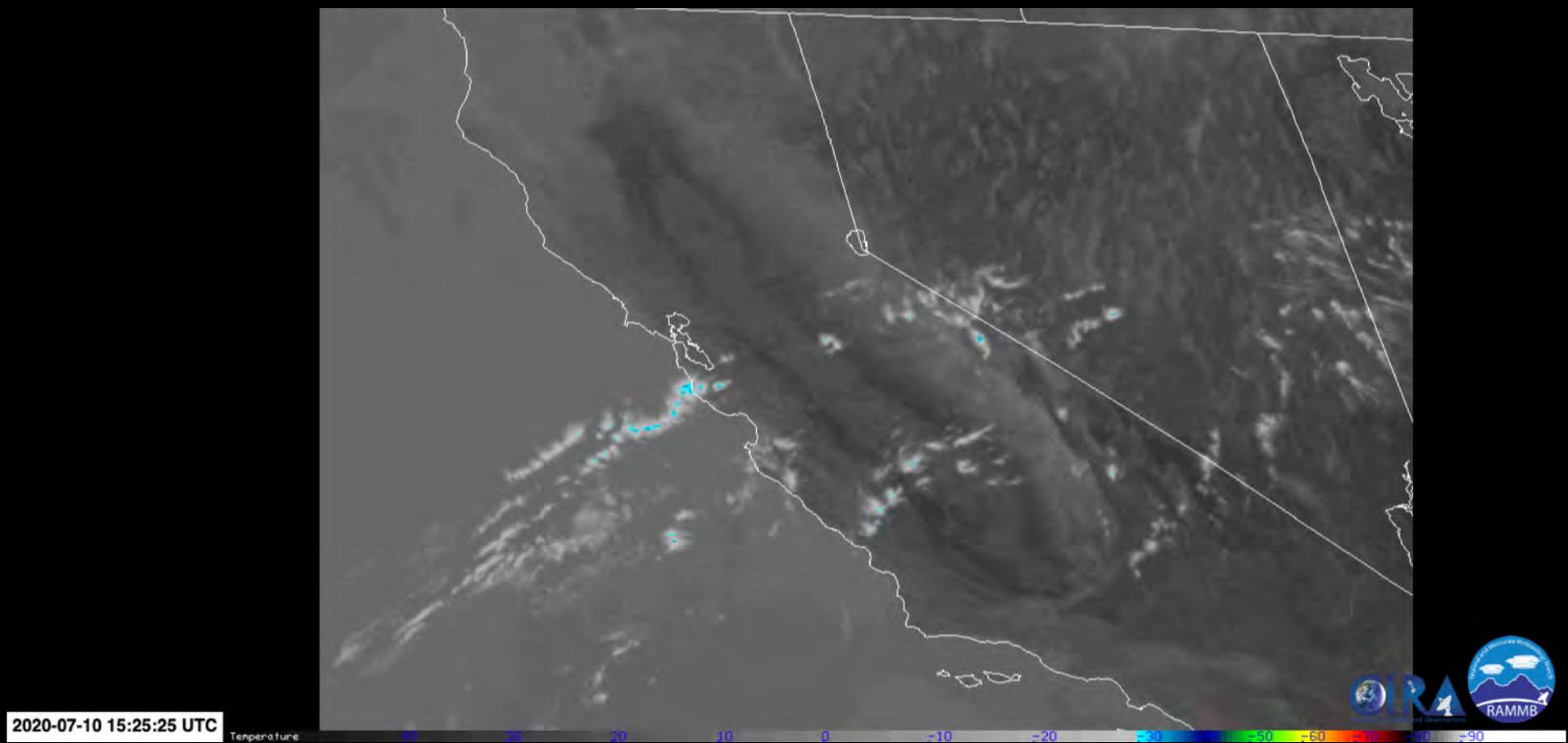
Band 11 (8.5 μm)



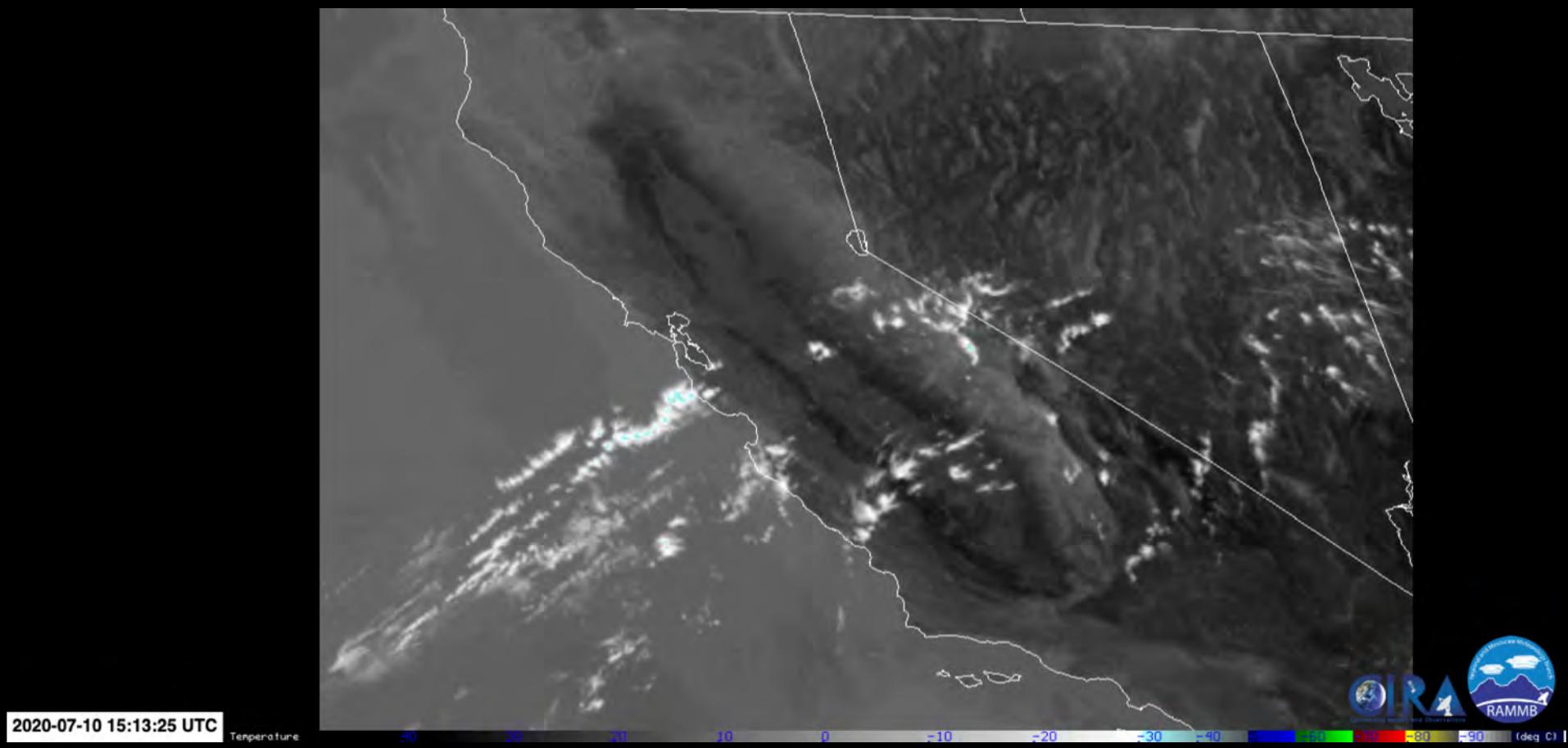
## Band 12 (9.61 μm): Ozone band



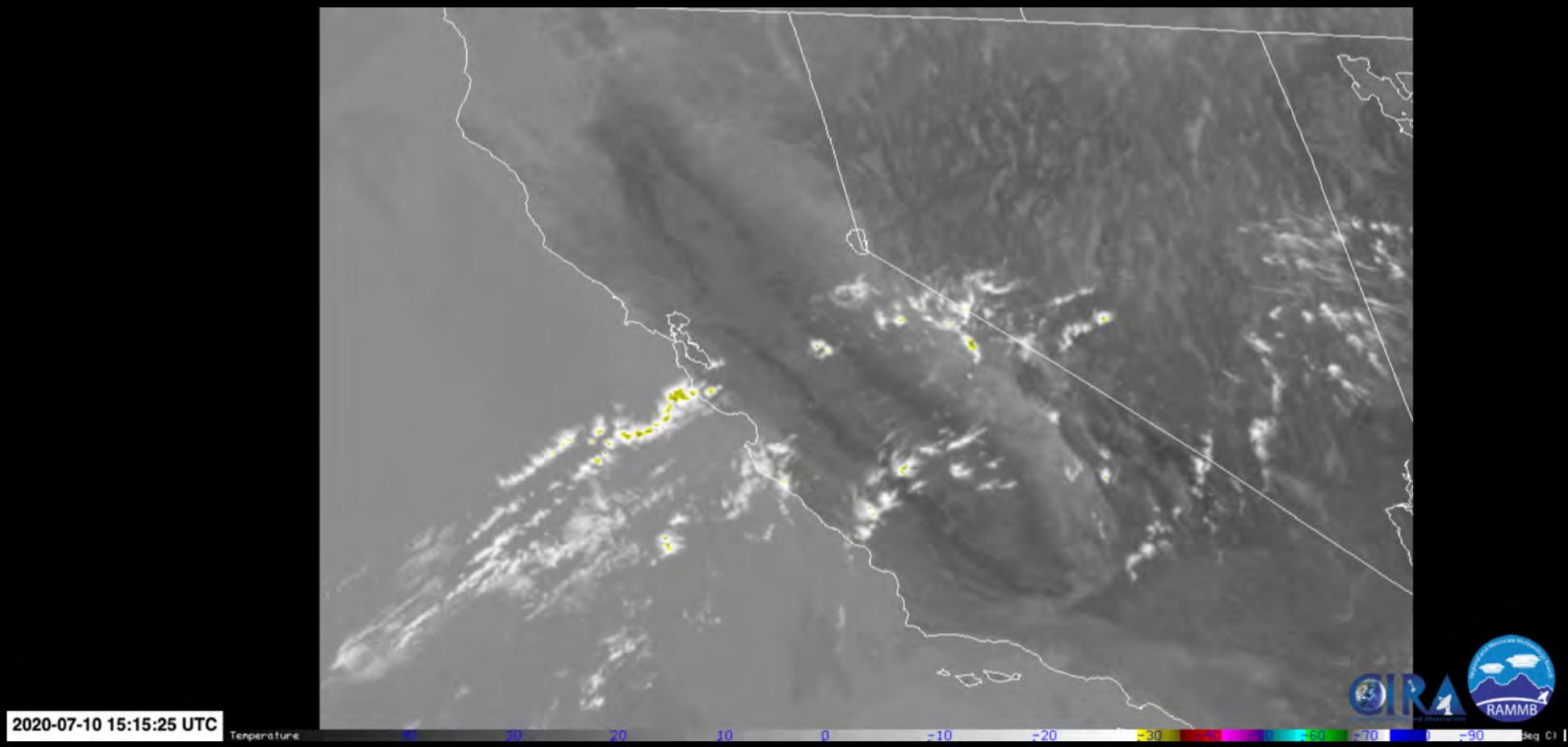
Band 13 (10.35  $\mu\text{m}$ )



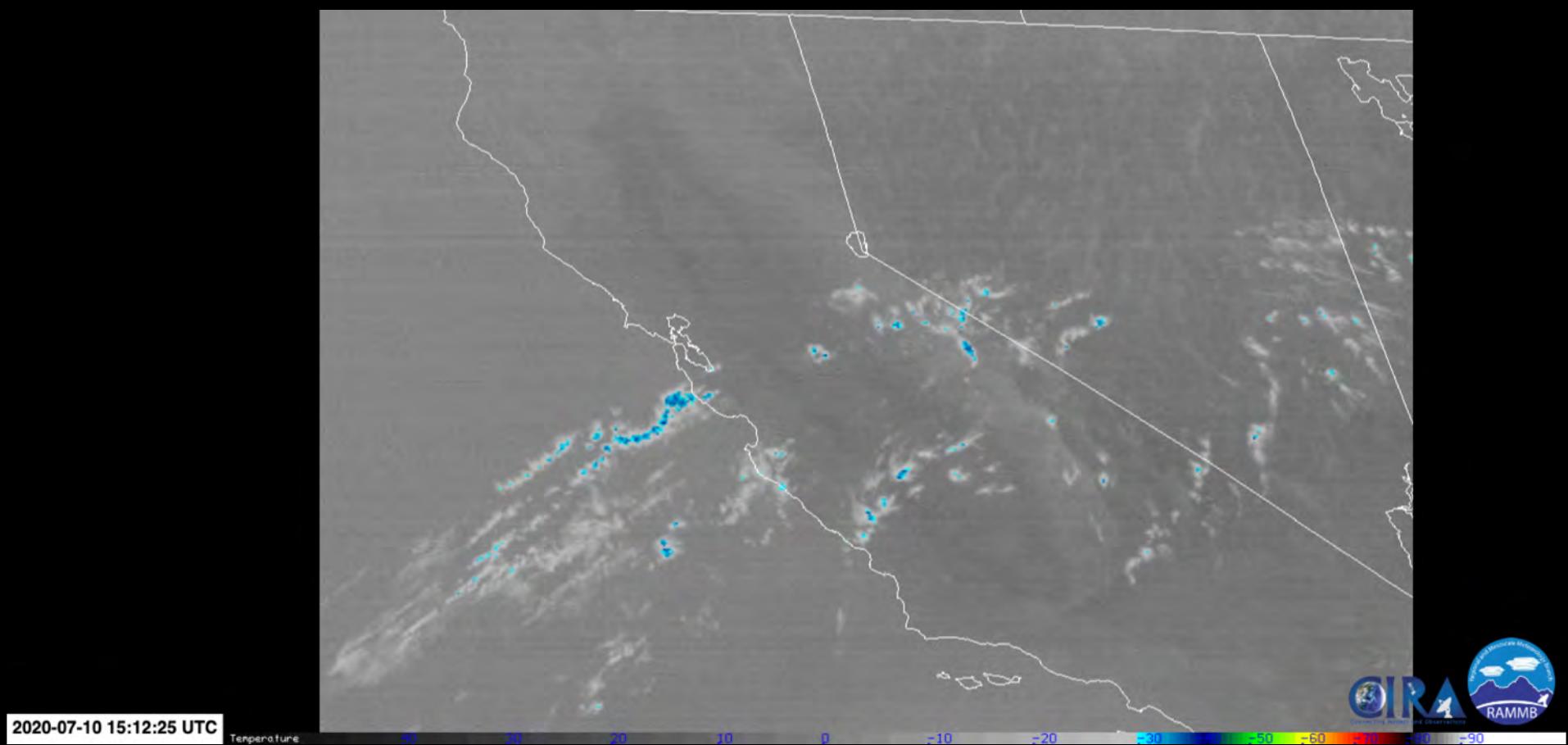
Band 14 (11.2  $\mu$ m)



Band 15 (12.3  $\mu$ m)



Band 16 (13.3 μm)



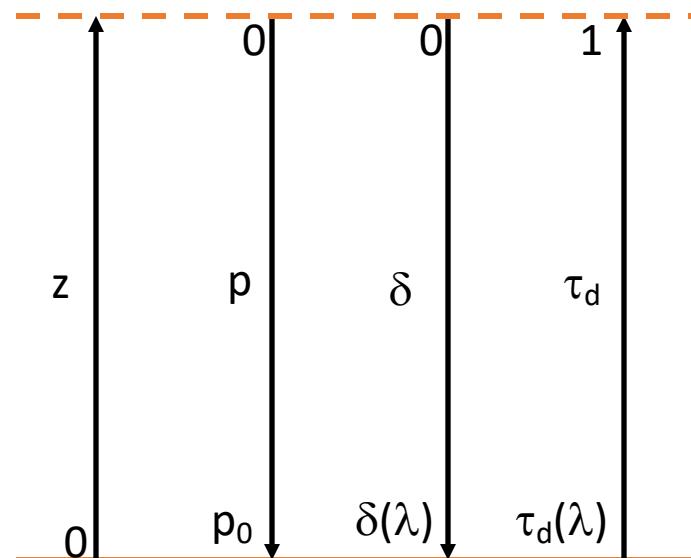
# MR3522: Remote Sensing of the Atmosphere and Ocean

## Weighting Functions

### Main Topics

- Interpreting weighting functions
- Relationship between optical depth and weighting functions

We can express variance in direct transmittance with height in terms of the weighting function.



$$\text{Direct Transmittance, } \tau_d = e^{-\delta(\lambda, p)/\mu}$$

The probability that a photon of wavelength  $\lambda$  will directly (without interaction) propagate to the top of the atmosphere from vertical position  $p$ .

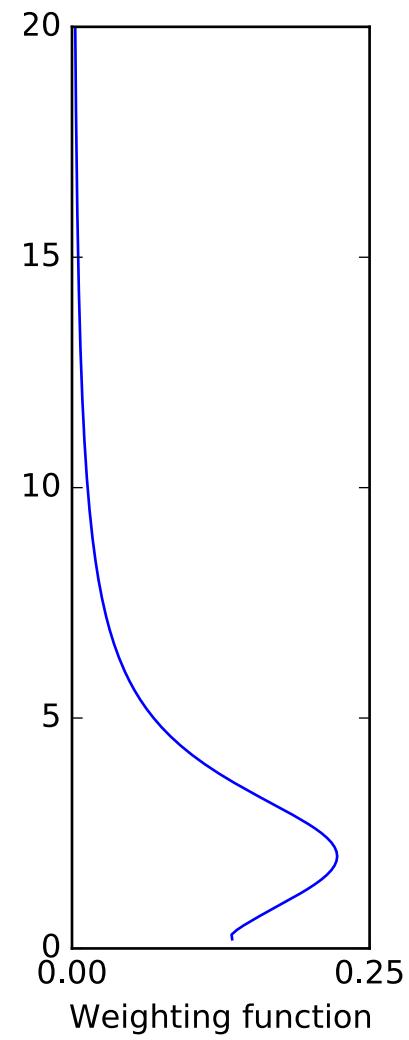
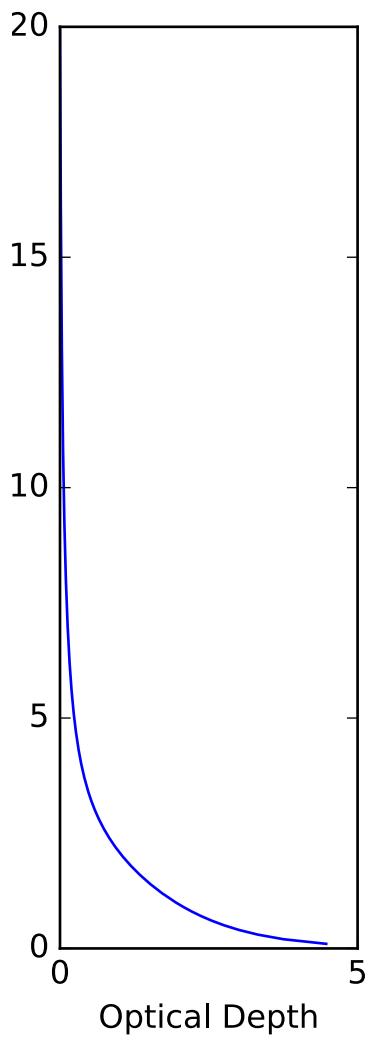
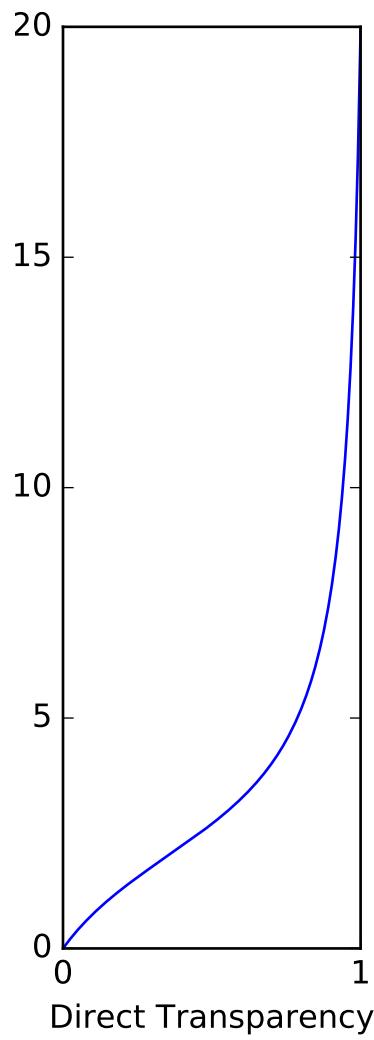
$$W = \frac{d\tau_d(\lambda, p)}{dp} = e^{-\delta(\lambda, p)/\mu} \left( -\frac{d\delta}{\mu dp} \right)$$

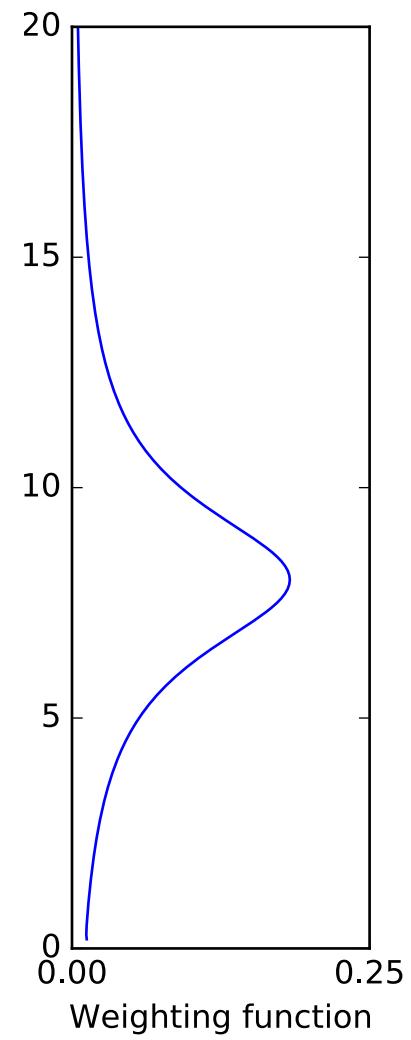
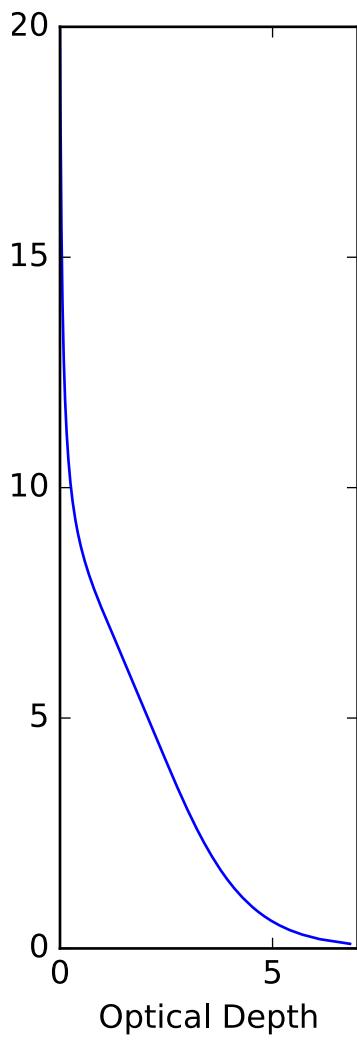
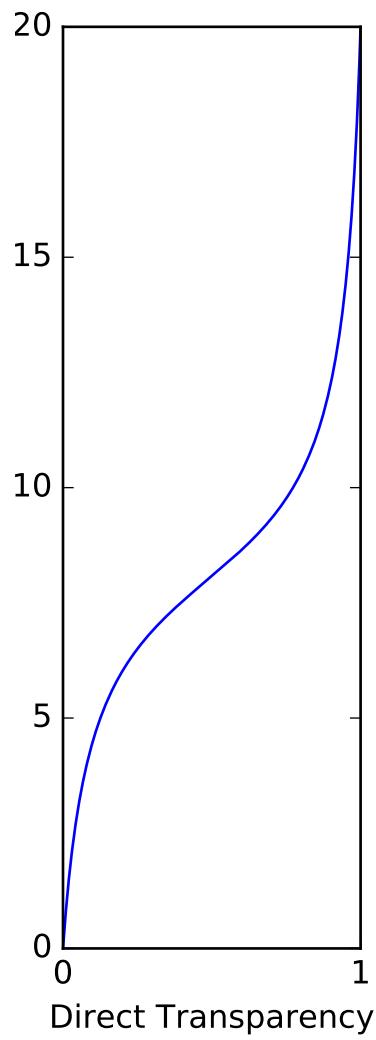
$$L_t(\lambda, \theta, \varphi) = \varepsilon_s B(\lambda, T_s) e^{-\delta(\lambda)/\mu} + \int_0^{\delta(\lambda)} B(\lambda, T(z)) e^{-\delta(\lambda, z)/\mu} \frac{d\delta}{\mu}$$

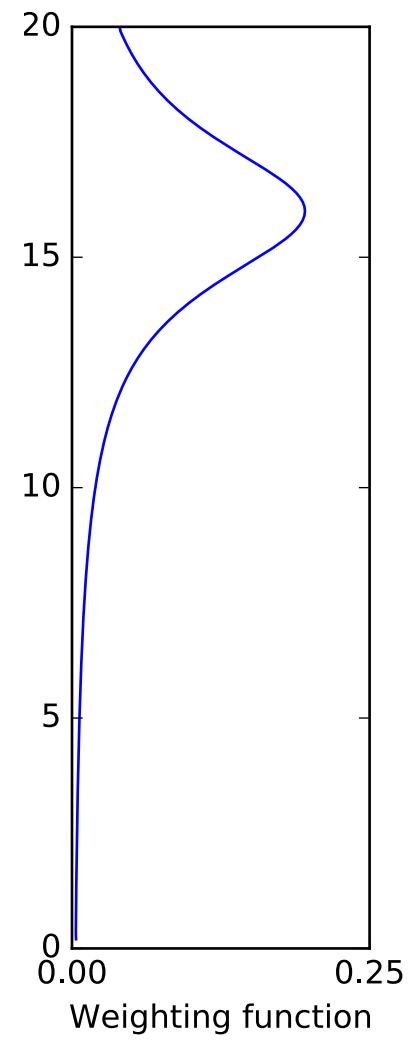
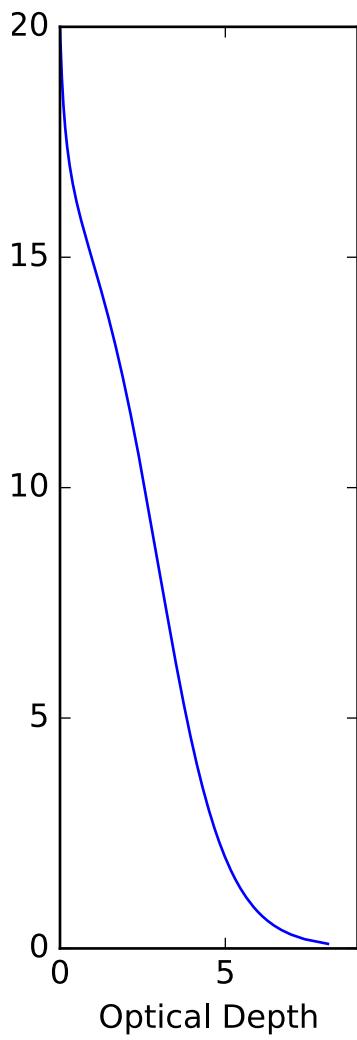
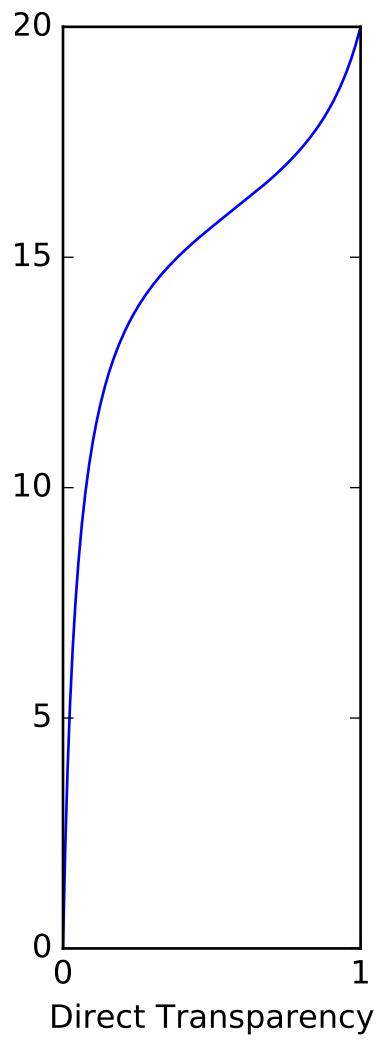
So,

$$L_t(\lambda, \theta, \varphi) = \varepsilon_s(\lambda, \theta) B(\lambda, T_s) \tau_d(\lambda) + \int_p^0 B(\lambda, T(p)) \frac{d\tau_d(\lambda, p)}{dp} dp$$

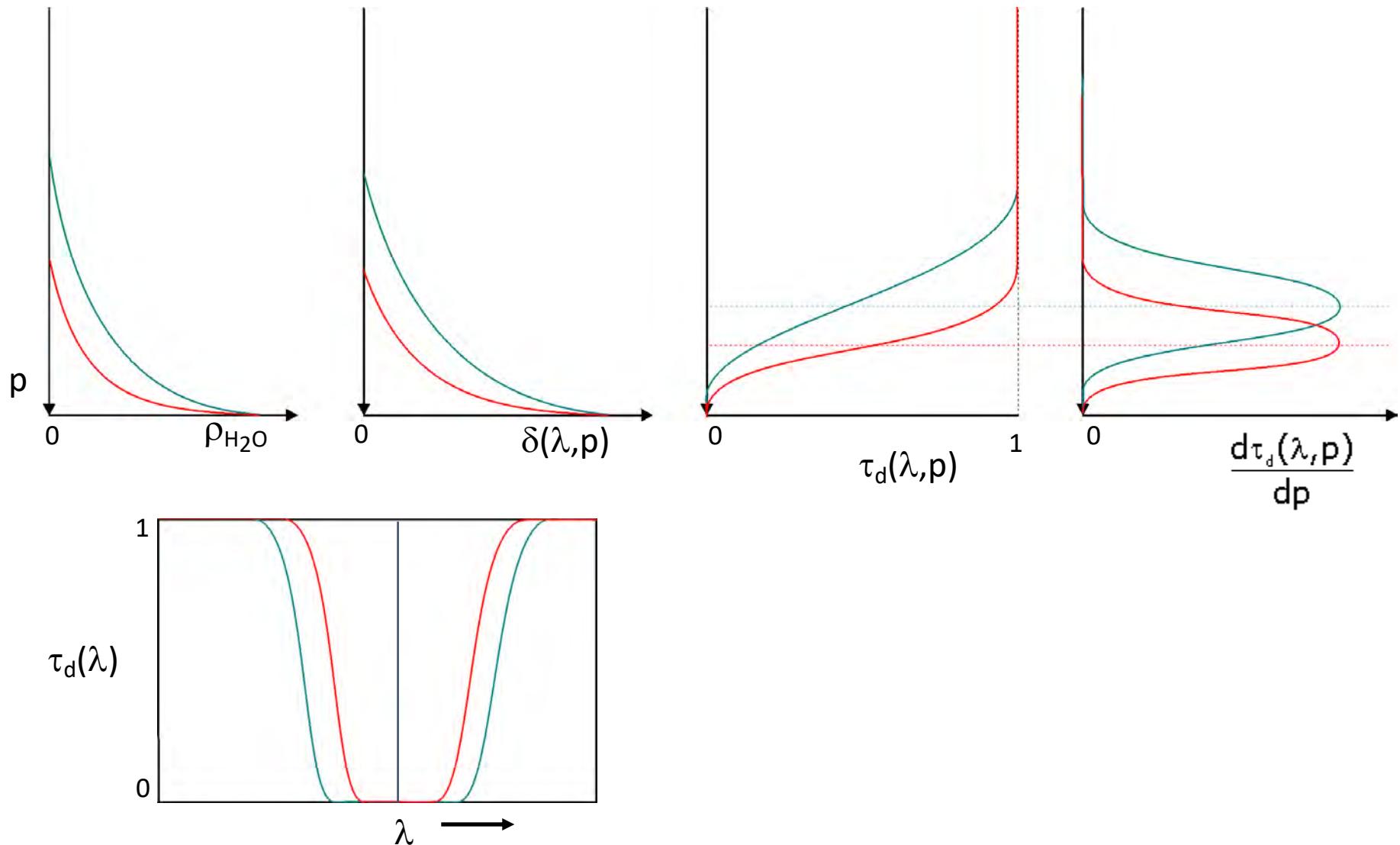
$\frac{d\tau_d(\lambda, p)}{dp}$  = Weighting Function  $\rightarrow$  Peak of  $W$  is at  $p$  where the path radiance contributes most to  $L_t$





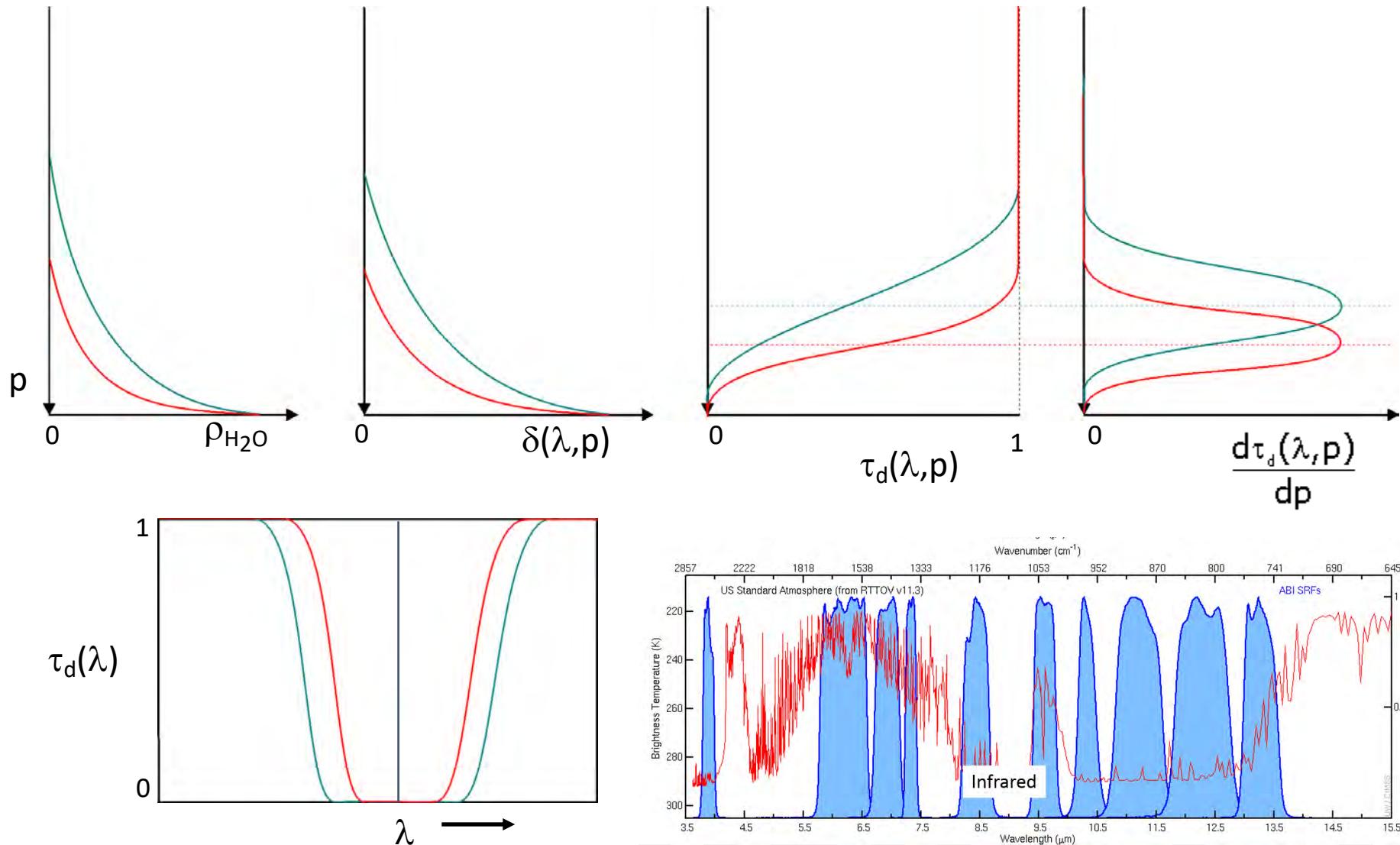


One wavelength, two concentration profiles, two different weighting functions



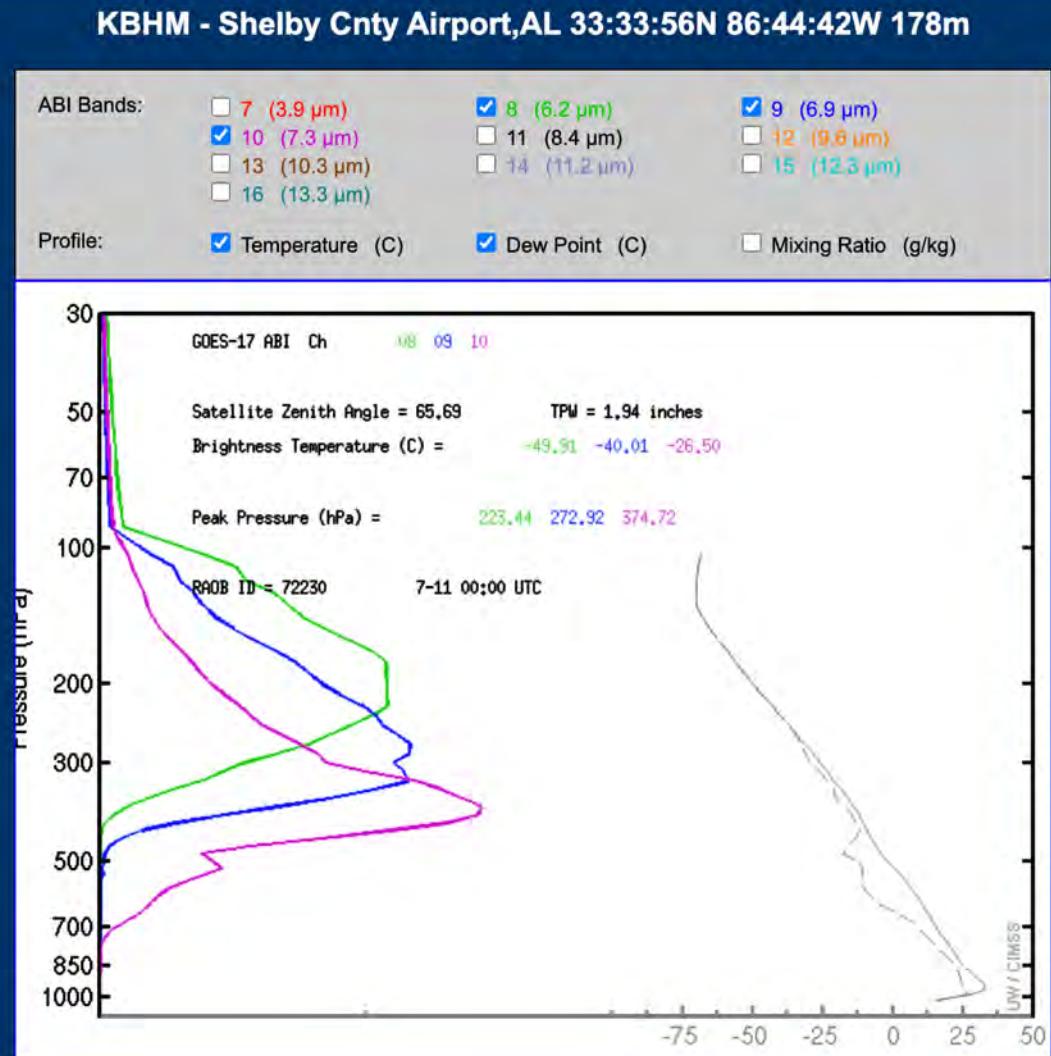
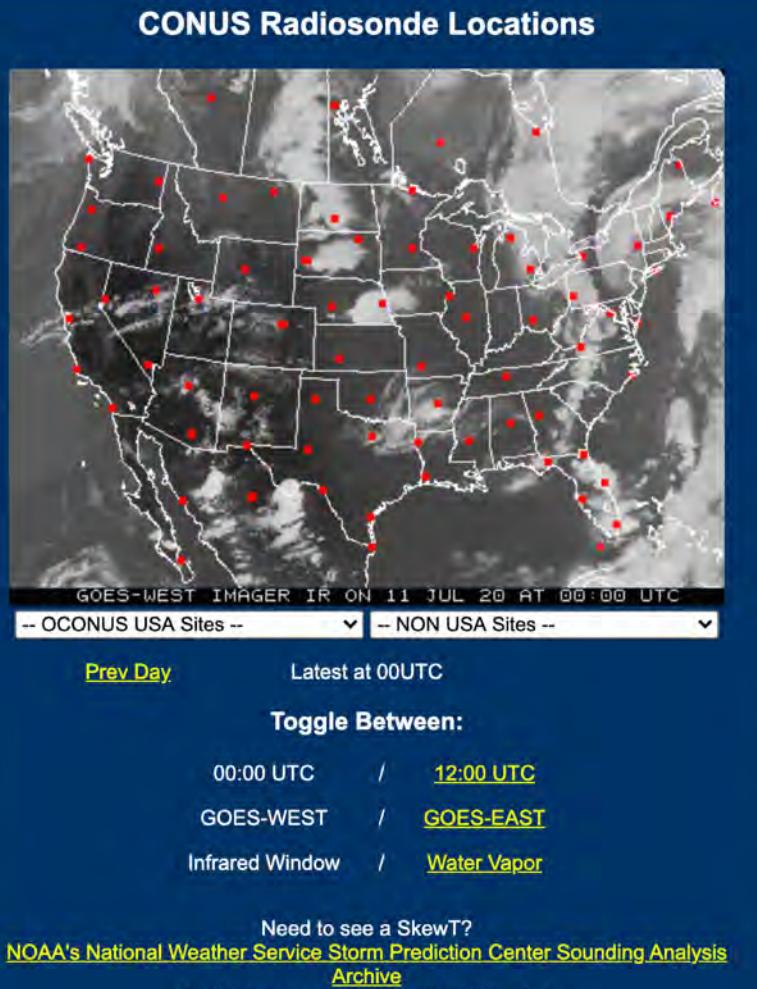
$$L_t(\lambda, \theta, \varphi) = \varepsilon_s(\lambda, \theta) B(\lambda, T_s) \tau_d(\lambda) + \int_p^0 B(\lambda, T(p)) \frac{d\tau_d(\lambda, p)}{dp} dp$$

# One wavelength, two concentration profiles, two different weighting functions



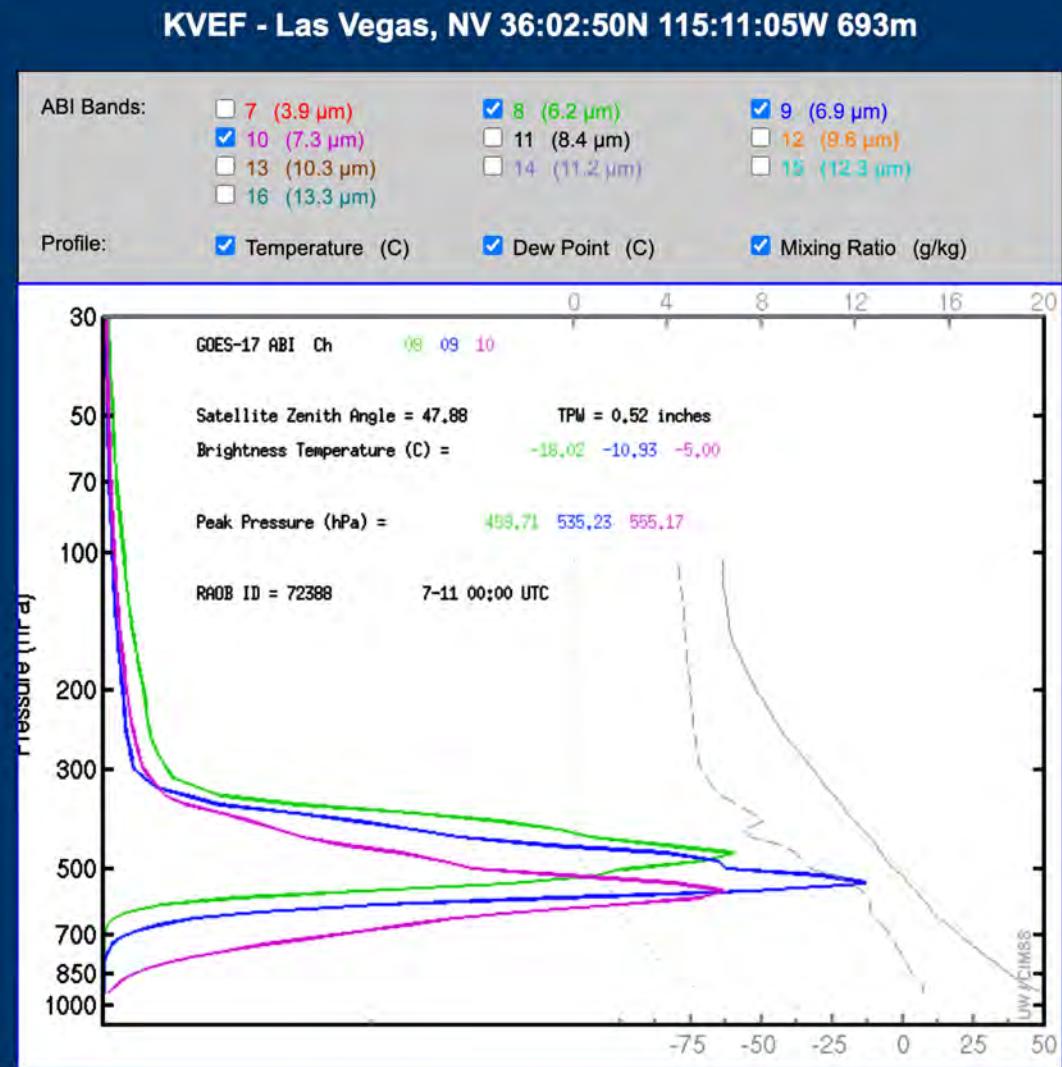
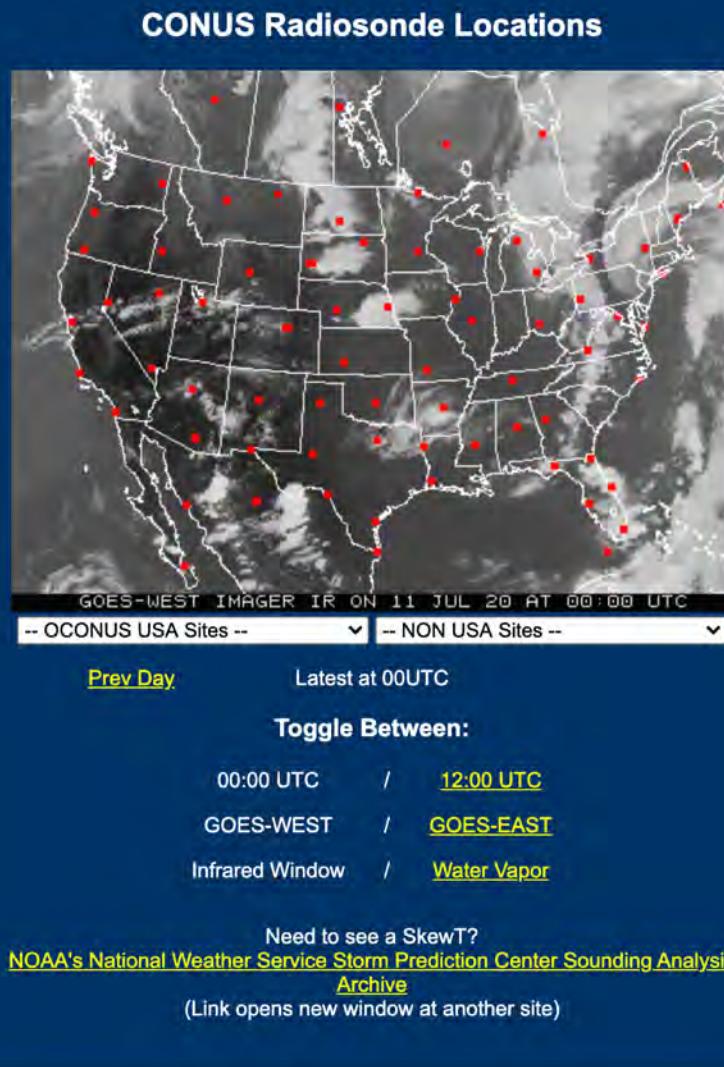
$$L_t(\lambda, \theta, \varphi) = \varepsilon_s(\lambda, \theta) B(\lambda, T_s) \tau_d(\lambda) + \int_p^0 B(\lambda, T(p)) \frac{d\tau_d(\lambda, p)}{dp} dp$$

# Moist Environment



<https://cimss.ssec.wisc.edu/goes/wf/>

# Dry Environment



<https://cimss.ssec.wisc.edu/goes/wf/>

# MR3522: Remote Sensing of the Atmosphere and Ocean

## Joint Polar Satellite System (JPSS)

### Main Topics

- NOAA-20 and Suomi-NPP
  - Orbital parameters and field of view
  - Bands/channels used

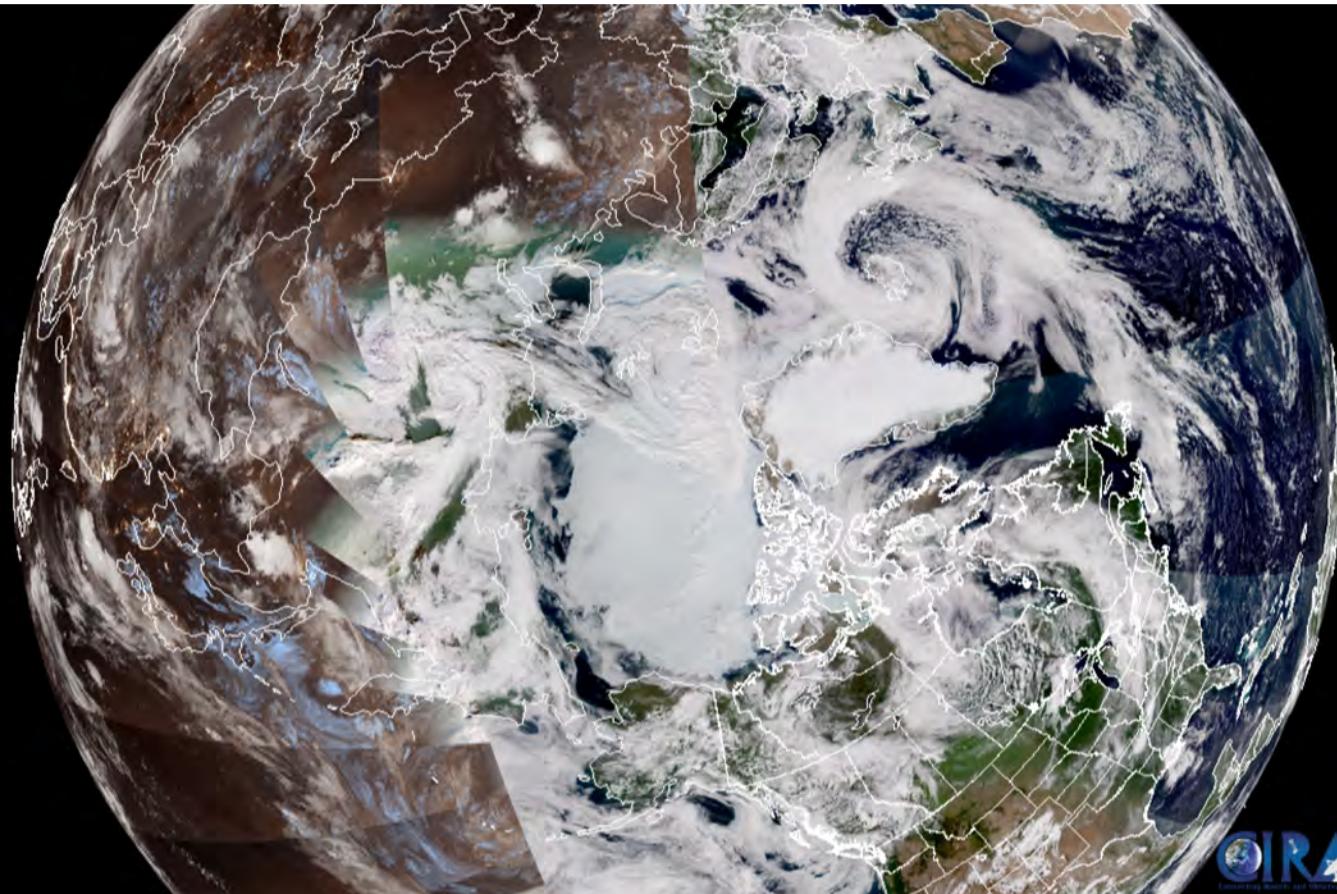
## Satellites in JPSS

- Suomi National Polar-orbiting Partnership (NPP)
- NOAA-20
- Three other satellites with future planned launch dates

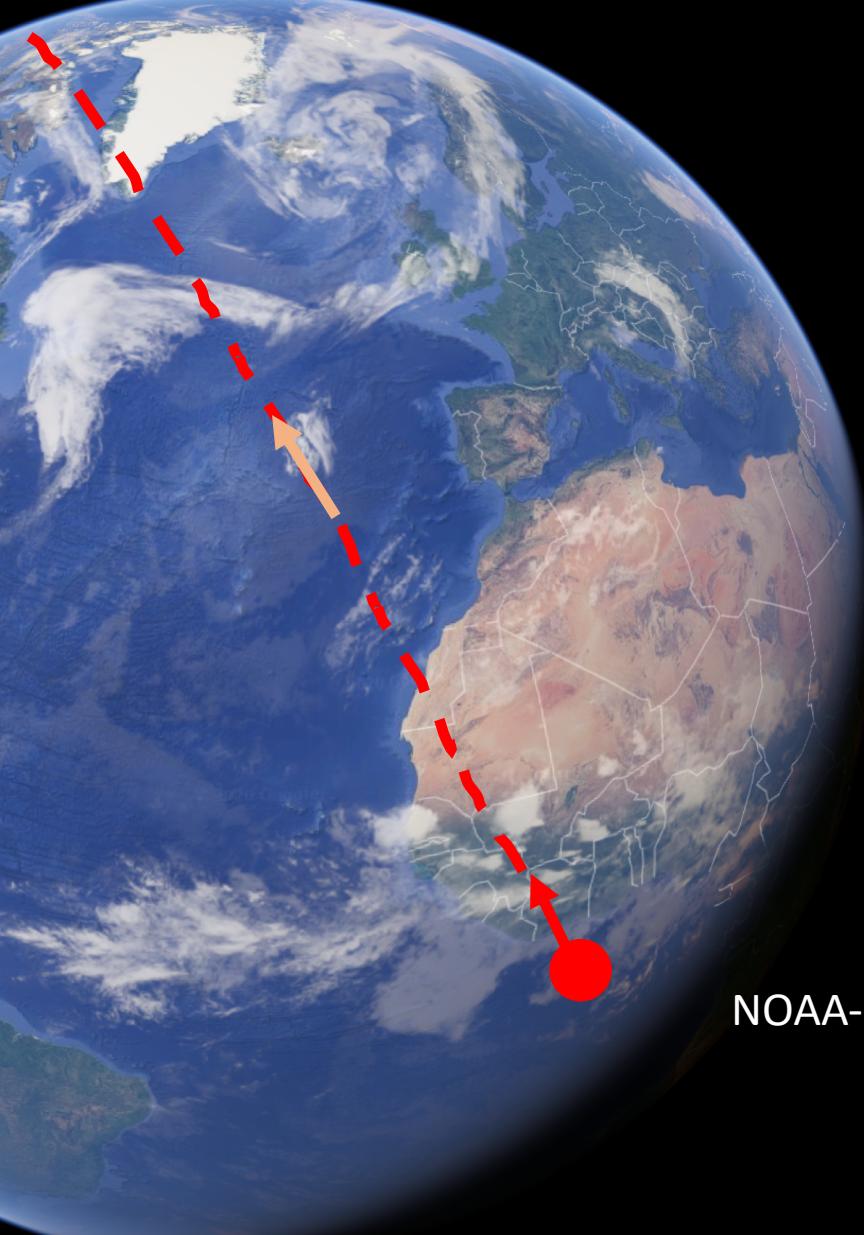
Orbit: Sun-synchronous, daytime ascending, equator crossing time of 1:30PM

- NOAA-20 is about 50 minutes behind of Suomi-NPP so crosses a little to the west.

JPSS  
Composite  
RGB Image:  
North Pole  
View



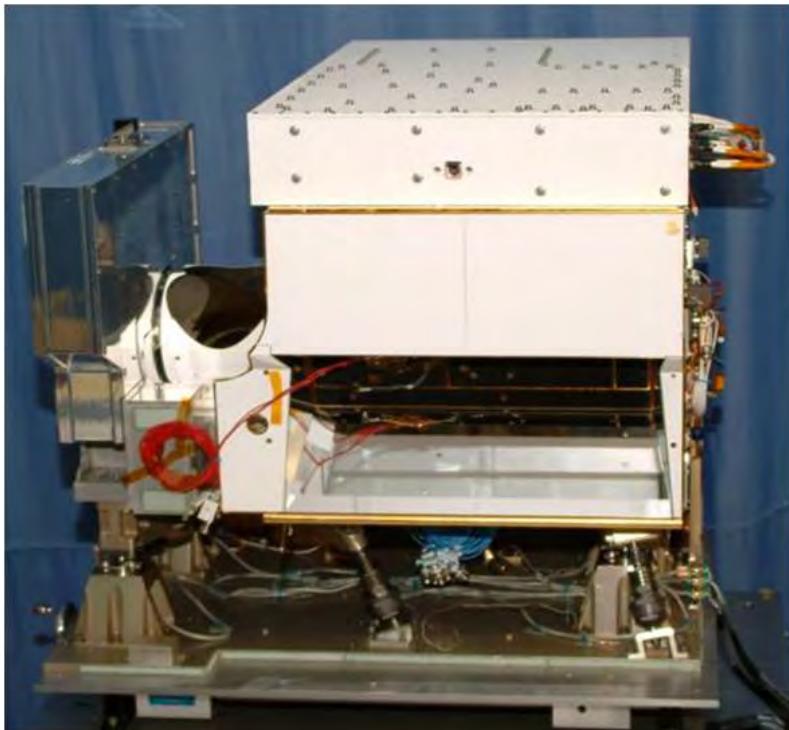
Suomi-NPP



NOAA-20

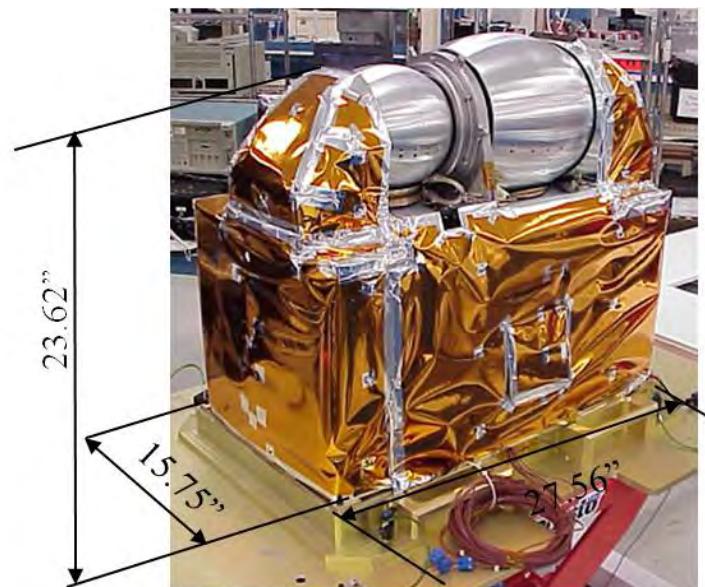
## Cross-track Infrared Sounder (CrIS)

- High spectral resolution sounder with over 3,000 spectral channels
- Estimates of atmospheric temperature and humidity profiles
- 14 km nadir data point spacing; 1 km vertical spacing
- Cannot see through clouds



## Advanced Technology Microwave Sounder (ATMS)

- 22 channels from 23.8 GHz to 183.3 GHz
- 16–75 km data point spacing
- Can see through clouds because microwave is not scattered as much



## Clouds and the Earth's Radiant Energy System (CERES) FM6 broadband radiometer

- Radiative fluxes, some cloud properties in "high", "middle" and "low" layers
- 20 km spacing of data points at nadir



[nasa.gov/image-feature/ceres-fm6/](https://nasa.gov/image-feature/ceres-fm6/)

<https://ceres.larc.nasa.gov/data/>

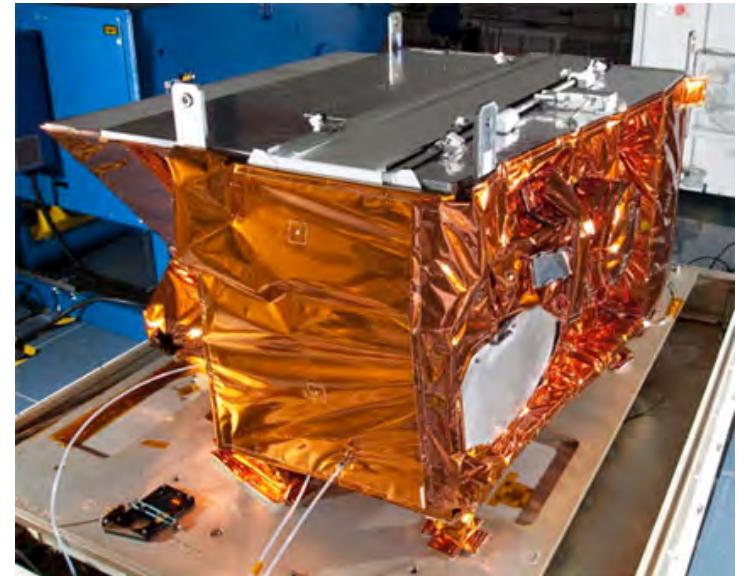
## Ozone Mapping and Profiler Suite (OMPS)

- Three spectrometers for measuring column-integrated ozone and ozone profiles
- 50 km data point spacing for mapper; 250 km for profiler



## Visible Infrared Imaging Radiometer Suite (VIIRS)

- Most similar to ABI on GOES with multiple sensors in the visible and infrared
- Up to 350 meter data point spacing
- Useful for weather monitoring, ocean color, SST estimates, aerosol detection



<https://www.star.nesdis.noaa.gov/jpss>

# VIIRS Sensor Bands

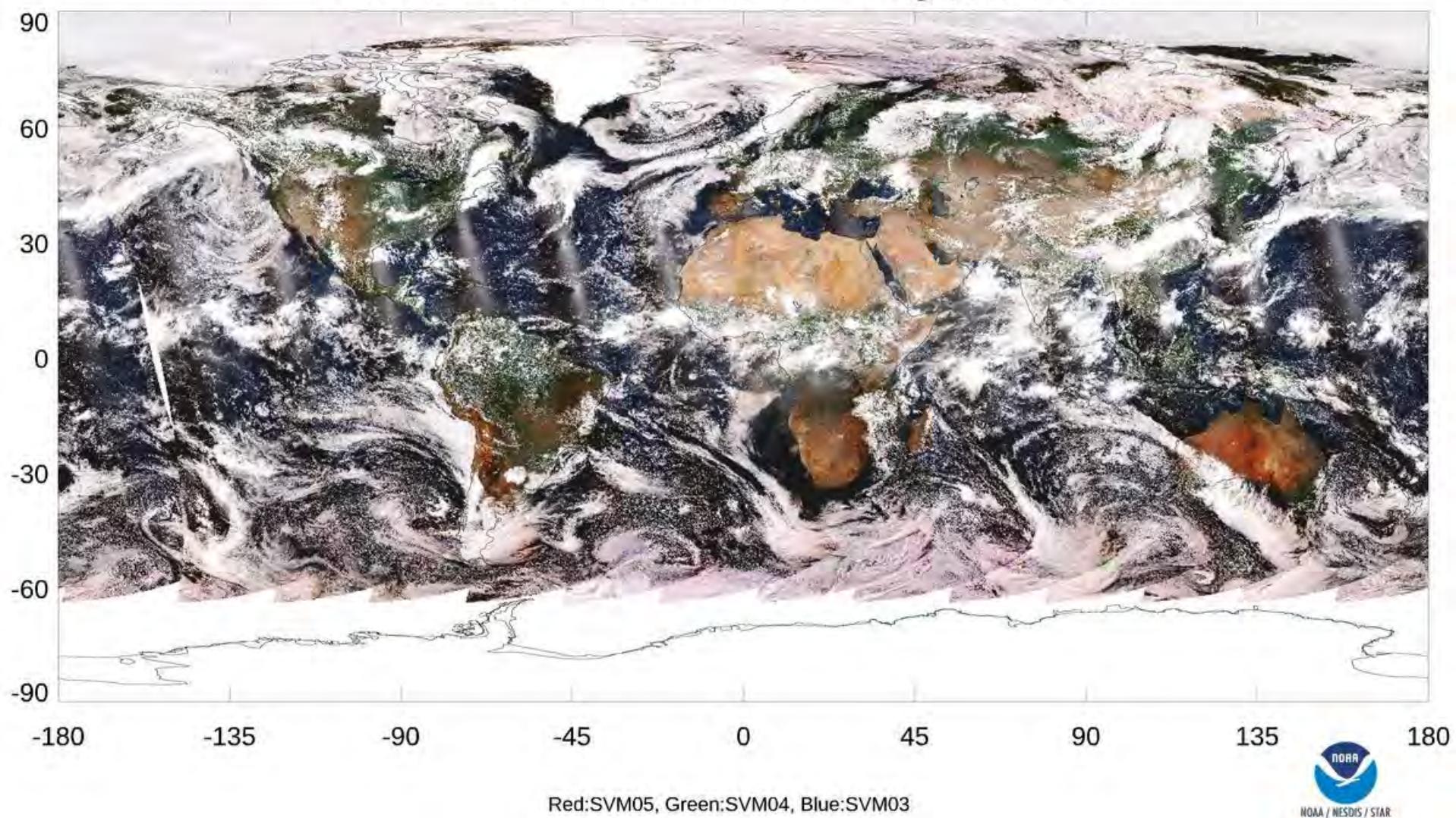
	Band No.	Wavelength (μm)	Horiz Sample Interval (km Downtrack x Crosstrack)		Driving EDRs	Radiance Range	Ltyp or Ttyp
			Nadir	End of Scan			
Vis/NIR FPA Silicon PIN Diodes	M1	0.412	0.742 x 0.259	1.60 x 1.58	Ocean Color Aerosols	Low High	44.9 155
	M2	0.445	0.742 x 0.259	1.60 x 1.58	Ocean Color Aerosols	Low High	40 146
	M3	0.488	0.742 x 0.259	1.60 x 1.58	Ocean Color Aerosols	Low High	32 123
	M4	0.555	0.742 x 0.259	1.60 x 1.58	Ocean Color Aerosols	Low High	21 90
	I1	0.640	0.371 x 0.387	0.80 x 0.789	Imagery	Single	22
	M5	0.672	0.742 x 0.259	1.60 x 1.58	Ocean Color Aerosols	Low High	10 68
	M6	0.746	0.742 x 0.776	1.60 x 1.58	Atmospheric Corr'n	Single	9.6
	I2	0.865	0.371 x 0.387	0.80 x 0.789	NDVI	Single	25
	M7	0.865	0.742 x 0.259	1.60 x 1.58	Ocean Color Aerosols	Low High	6.4 33.4
	CCD	DNB	0.7	0.742 x 0.742	0.742 x 0.742	Imagery	Var. 6.70E-05
S/MWIR PV HgCdTe (HCT)	M8	1.24	0.742 x 0.776	1.60 x 1.58	Cloud Particle Size	Single	5.4
	M9	1.378	0.742 x 0.776	1.60 x 1.58	Cirrus/Cloud Cover	Single	6
	I3	1.61	0.371 x 0.387	0.80 x 0.789	Binary Snow Map	Single	7.3
	M10	1.61	0.742 x 0.776	1.60 x 1.58	Snow Fraction	Single	7.3
	M11	2.25	0.742 x 0.776	1.60 x 1.58	Clouds	Single	0.12
	I4	3.74	0.371 x 0.387	0.80 x 0.789	Imagery Clouds	Single	270 K
	M12	3.70	0.742 x 0.776	1.60 x 1.58	SST	Single	270 K
	M13	4.05	0.742 x 0.259	1.60 x 1.58	SST	Low	300 K
					Fires	High	380 K
LWIR PV HCT	M14	8.55	0.742 x 0.776	1.60 x 1.58	Cloud Top Properties	Single	270 K
	M15	10.763	0.742 x 0.776	1.60 x 1.58	SST	Single	300 K
	I5	11.450	0.371 x 0.387	0.80 x 0.789	Cloud Imagery	Single	210 K
	M16	12.013	0.742 x 0.776	1.60 x 1.58	SST	Single	300 K

High resolution IR imagery;  
750 m compared to 2 km  
with GOES. 11.45 micron  
imagery has 370 x 387 meter  
sampling at nadir

Many more band in **blue** light part of visible spectrum;  
useful for aerosol detection and ocean color

Band	Midpoint (μm)	Bandwidth (μm)	Range (μm)	Region	Spatial Resolution at nadir
M1	0.412	0.02	0.402 - 0.422	Visible (reflective)	750 m
M2	0.445	0.018	0.436 - 0.454		
M3	0.488	0.02	0.478 - 0.488		
M4	0.555	0.02	0.545 - 0.565		
M5 (B)	0.672	0.02	0.662 - 0.682		
M6	0.746	0.015	0.739 - 0.754		
M7 (G)	0.865	0.039	0.846 - 0.885		
M8	1.240	0.020	1.23 - 1.25		
M9	1.378	0.015	1.371 - 1.386		
M10 (R)	1.61	0.06	1.58 - 1.64		
M11	2.25	0.05	2.23 - 2.28	Shortwave IR	Medium-wave IR
M12	3.7	0.18	3.61 - 3.79		
M13	4.05	0.155	3.97 - 4.13		
M14	8.55	0.3	8.4 - 8.7		
M15 <sup>1</sup>	10.763	1.0	10.26 - 11.26		Longwave IR
M16	12.013	0.95	11.54 - 12.49		
DNB	0.7	0.4	0.5 - 0.9	Visible (reflective)	
I1 (B) <sup>2</sup>	0.64	0.08	0.6 - 0.68	Visible (reflective)	375 m
I2 (G)	0.865	0.039	0.85 - 0.88	Near IR	
I3 (R) <sup>2</sup>	1.61	0.06	1.58 - 1.64	Shortwave IR	
I4	3.74	0.38	3.55 - 3.93	Medium-wave IR	
I5	11.45	1.9	10.5 - 12.4	Longwave IR	

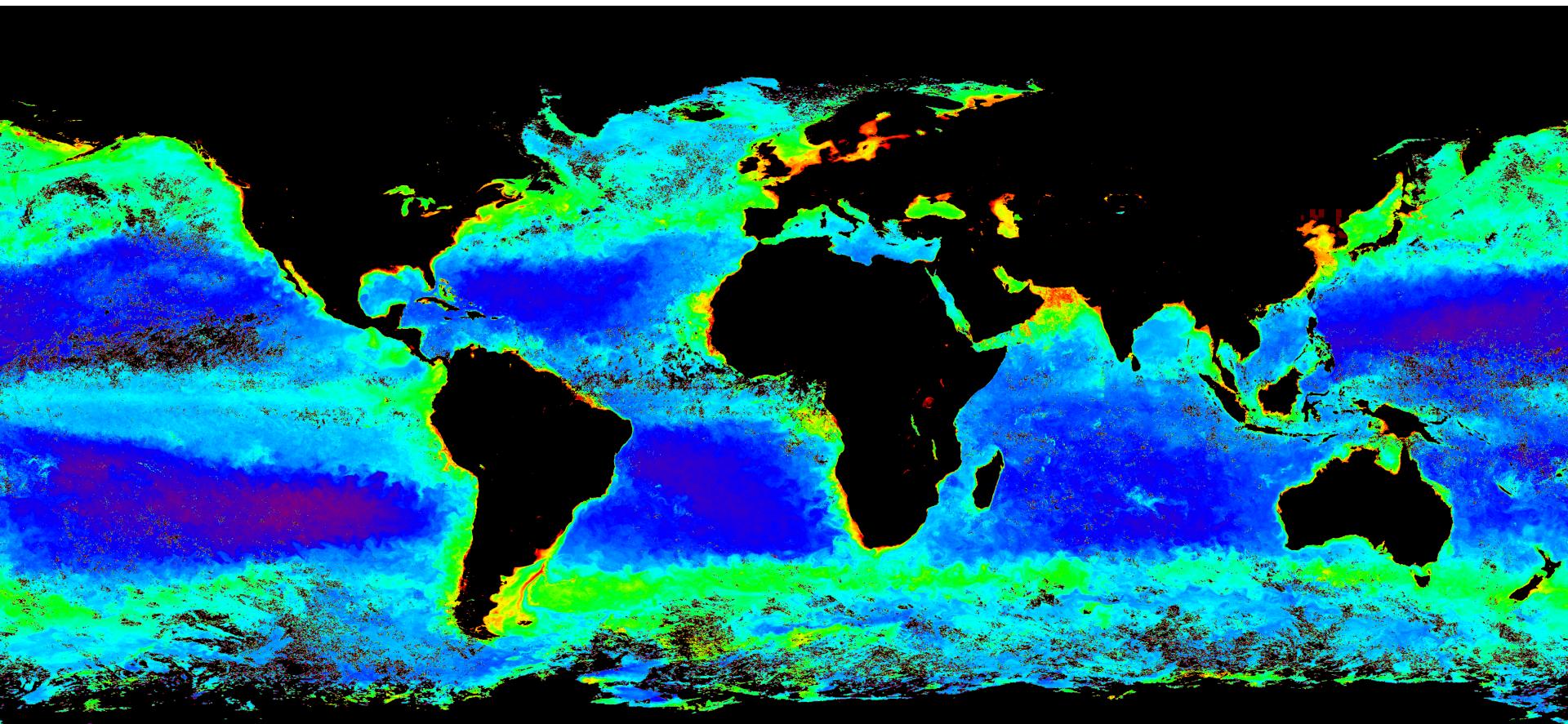
## NOAA-20 VIIRS Global True Color Image, 2020-07-17



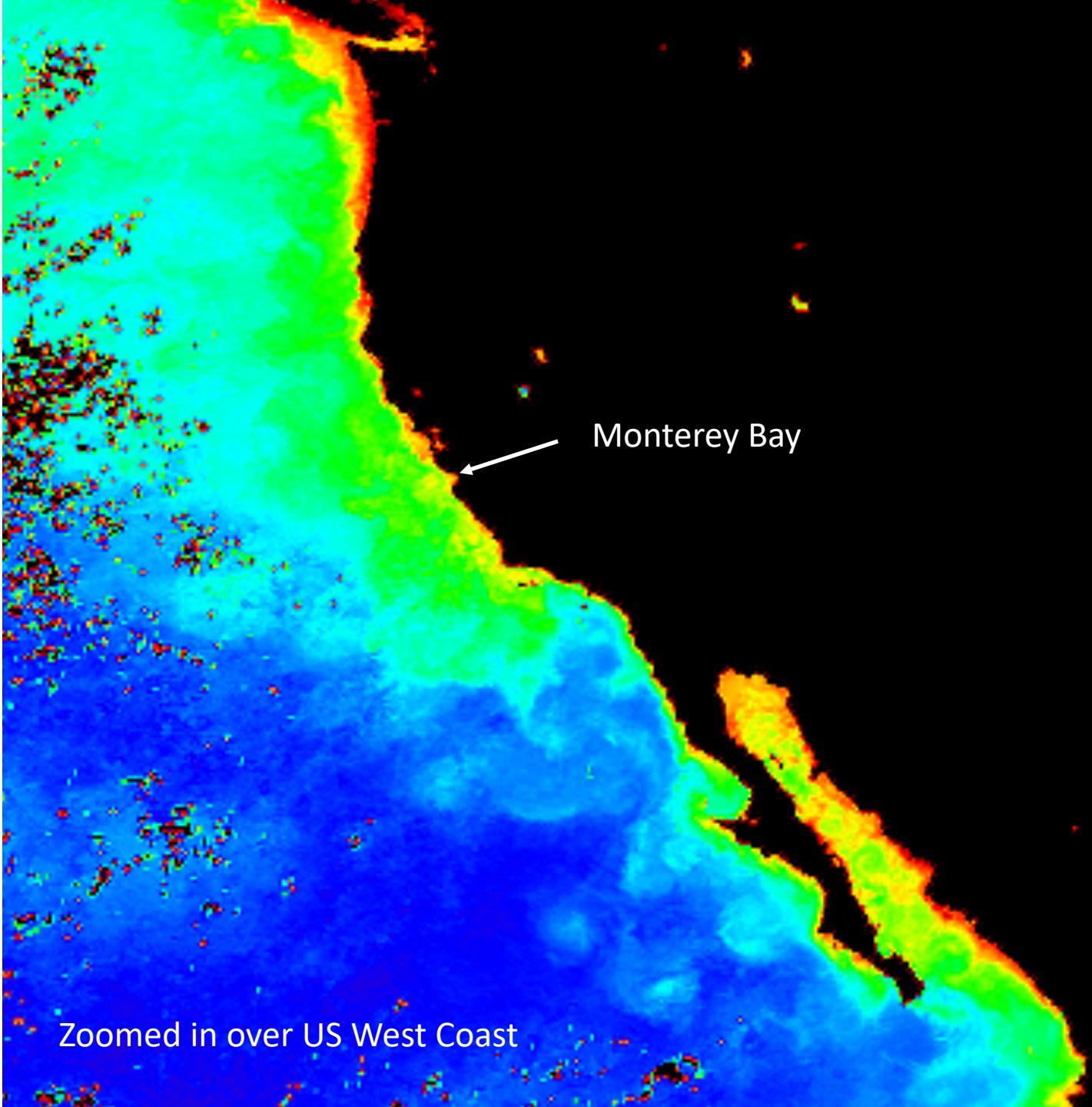
Red:SVM05, Green:SVM04, Blue:SVM03



March 2020 average chlorophyll concentration from Suomi-NPP

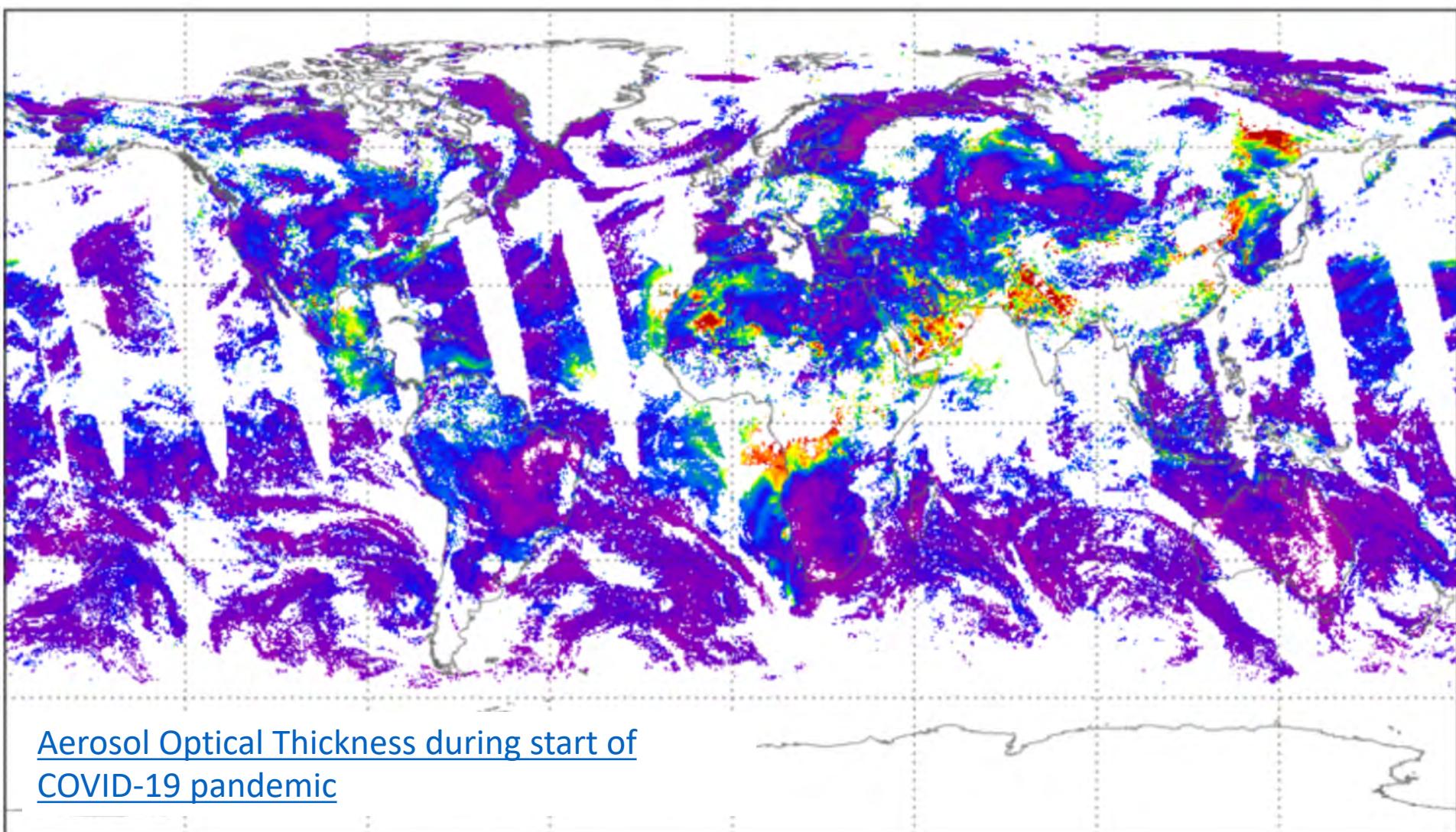


<https://oceancolor.gsfc.nasa.gov/>



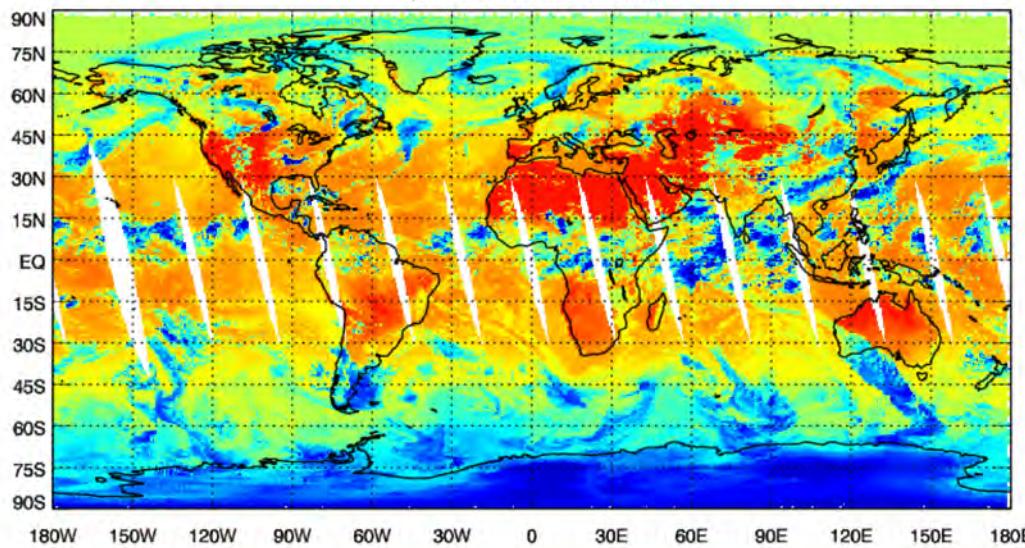
# Suomi NPP VIIRS High Quality Aerosol Optical Thickness at 550 nm JPSS EPS

17 Jul 2020

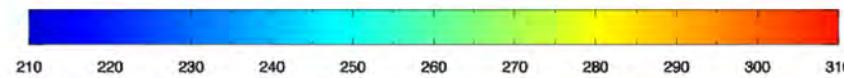
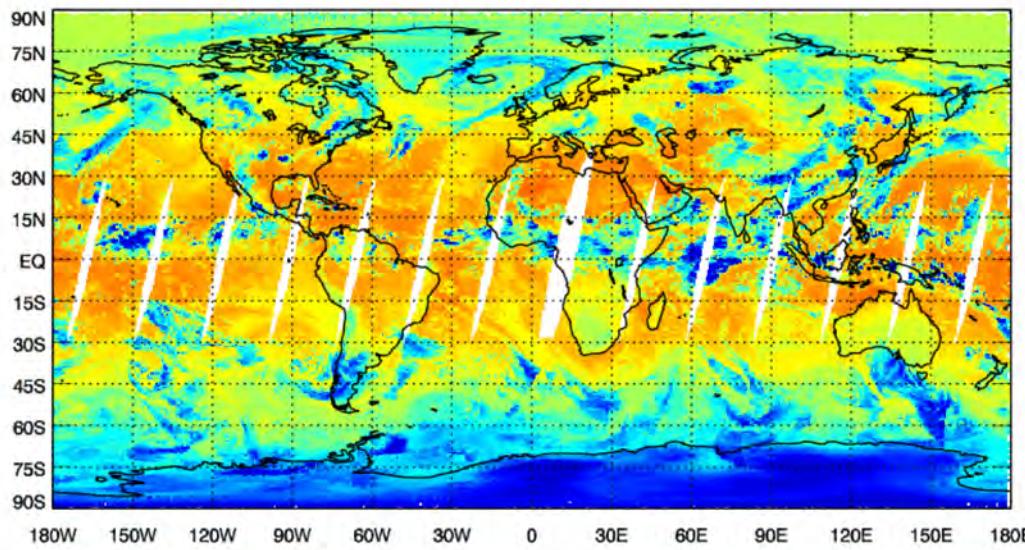


N20 CrIS FSR BT, 11  $\mu\text{m}$  (900  $\text{cm}^{-1}$ ), Mapped, Ascending, 07/17/2020

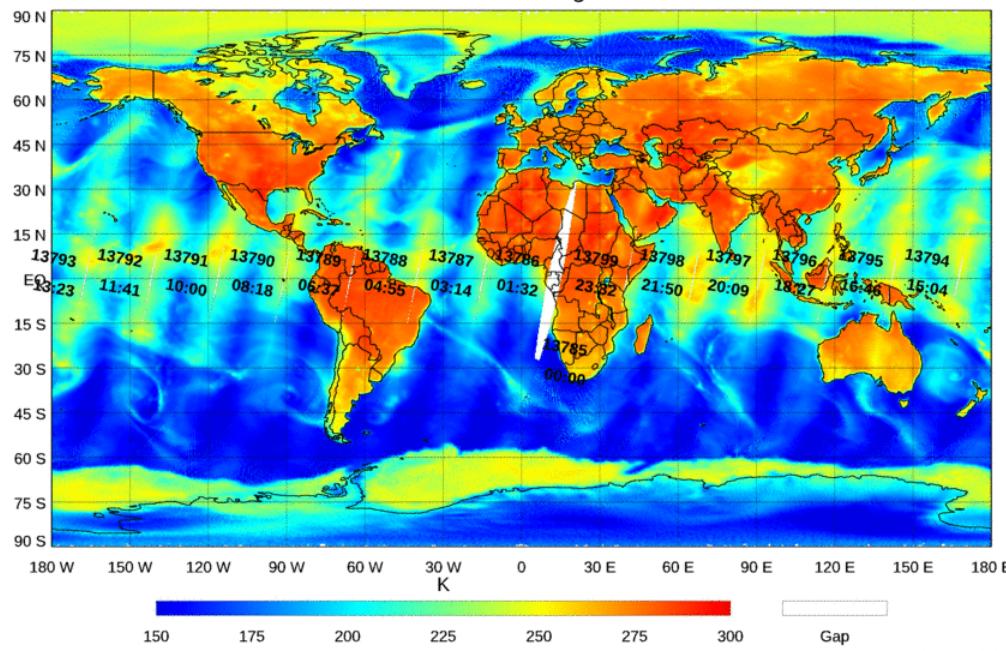
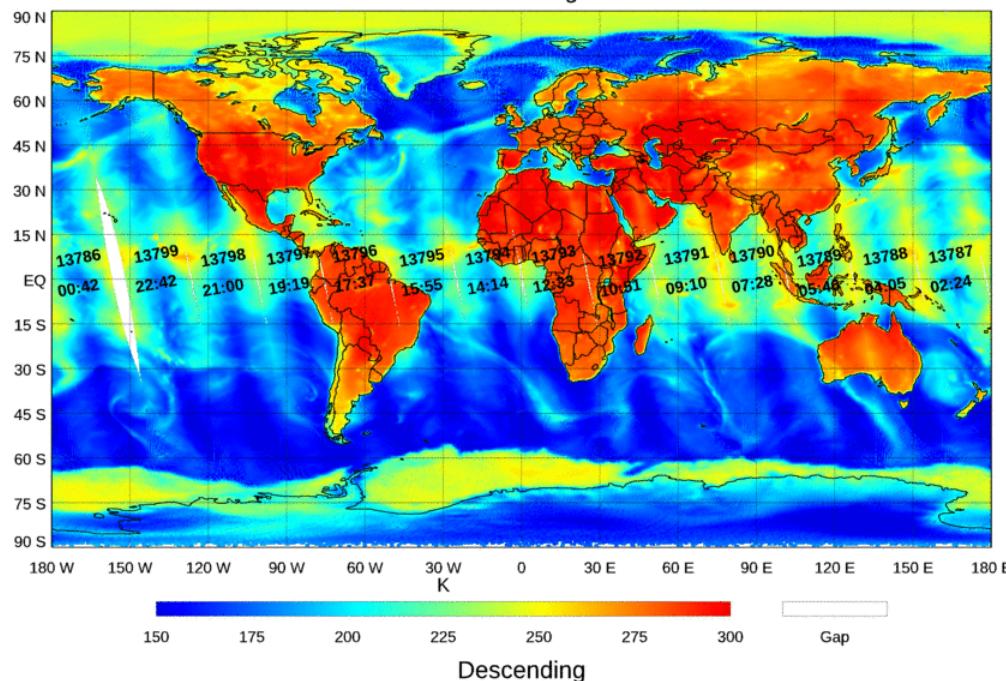
Updated at Jul 18 02:19:38 2020 UTC



N20 CrIS FSR BT, 11  $\mu\text{m}$  (900  $\text{cm}^{-1}$ ), Mapped, Descending, 07/17/2020

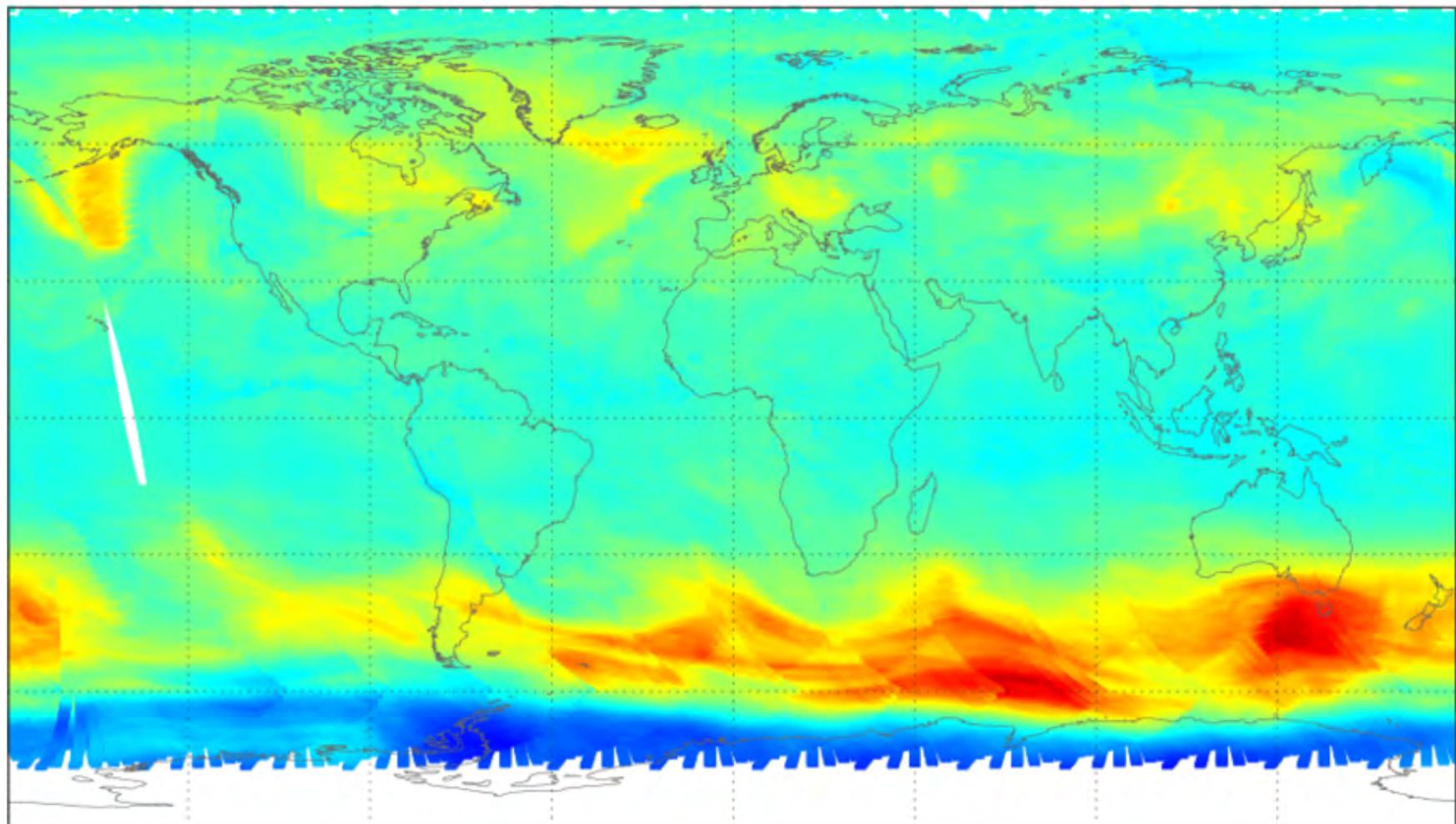


NOAA-20 ATMS SDR Ch.1 23.8 GHz QV-POL  
17 Jul 2020  
Ascending



# Suomi NPP OMPS V8 Total Ozone

4 Sep 2017



NOAA/NESDIS/STAR

# MR3522: Remote Sensing of the Atmosphere and Ocean

## Surface Temperature Estimation using Infrared Radiances

### Main Topics

- IR bands used to estimate SST
- Multi-channel SST estimates
- Land surface temperature
- Variable non-blackbody emittance by surface objects

Remember, even in an atmospheric window:

**Brightness temperature is not actual temperature!**

- 1) Small amounts of absorption could reduce radiation from surface.
- 2) Emitting surface may not be a blackbody.

$$T_B = \frac{\hbar c}{\lambda k \ln \frac{2\hbar c^2}{\lambda^5 L_t}}$$


We can compute brightness temperature from observed radiance.

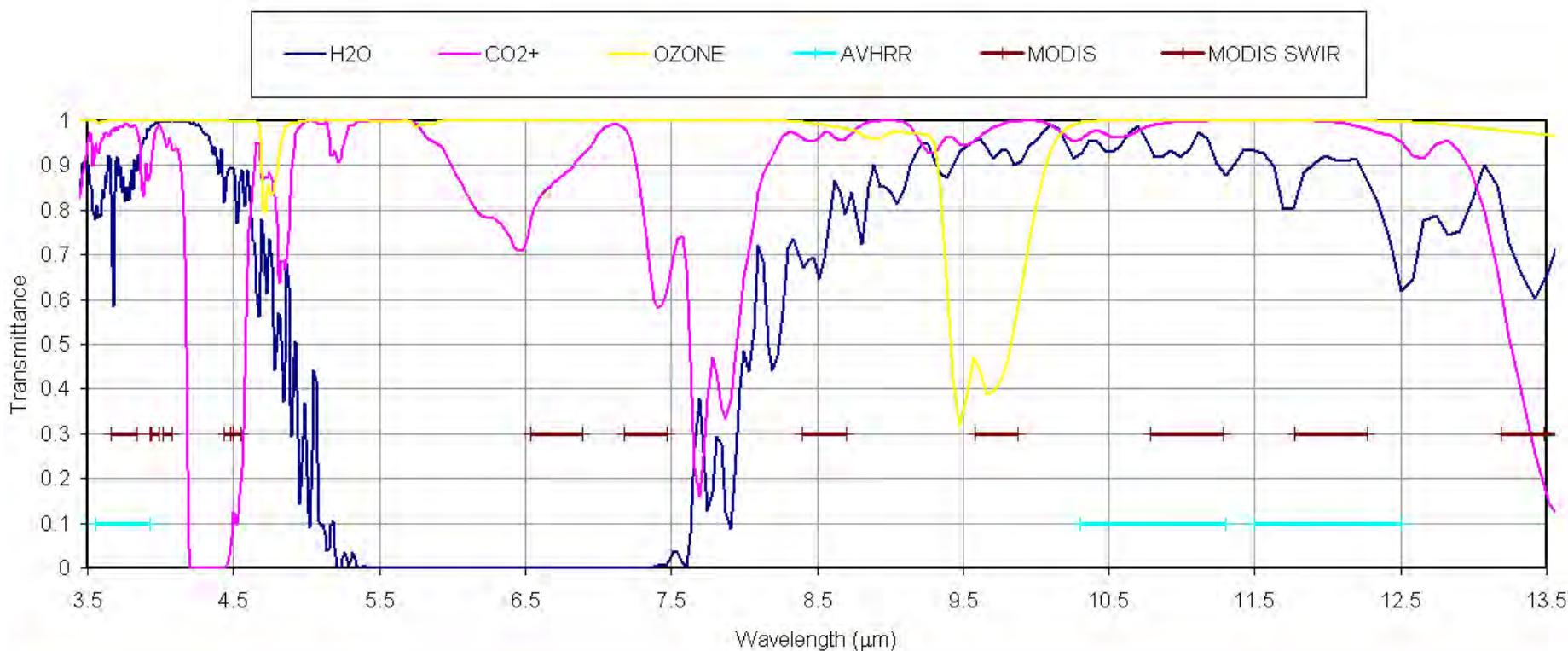
GOAL: Correct brightness temperatures using radiances in multiple bands to actual temperature at surface by accounting for water vapor absorption.

Limitation: This only works in cloud-free areas!

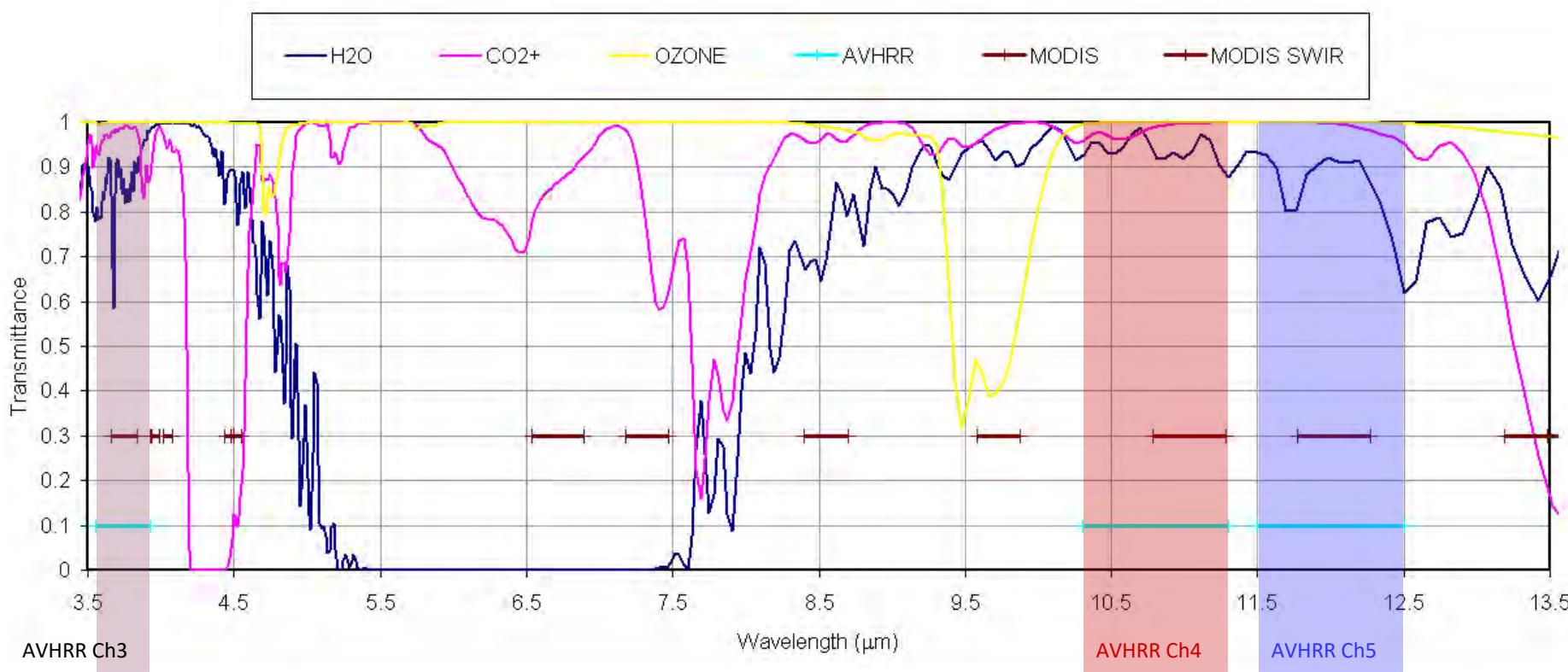
From Schwartzchild's Equation:

$$L_t(\lambda, \theta, \phi) = \varepsilon_s(\lambda, \theta)B(\lambda, T_s)\tau_d(\lambda) + \int_p^{\infty} B(\lambda, T(p)) \frac{d\tau_d(\lambda, p)}{dp} dp$$

... and for surface radiance to dominate we need a wavelength in a “window”  
so  $\tau_d(\lambda)$  is large and  $d\tau_d(\lambda, p)/dp$  is small above the surface



None of these are in perfect windows so there will always be some contribution from the path term due primarily to absorption/emission by water vapor

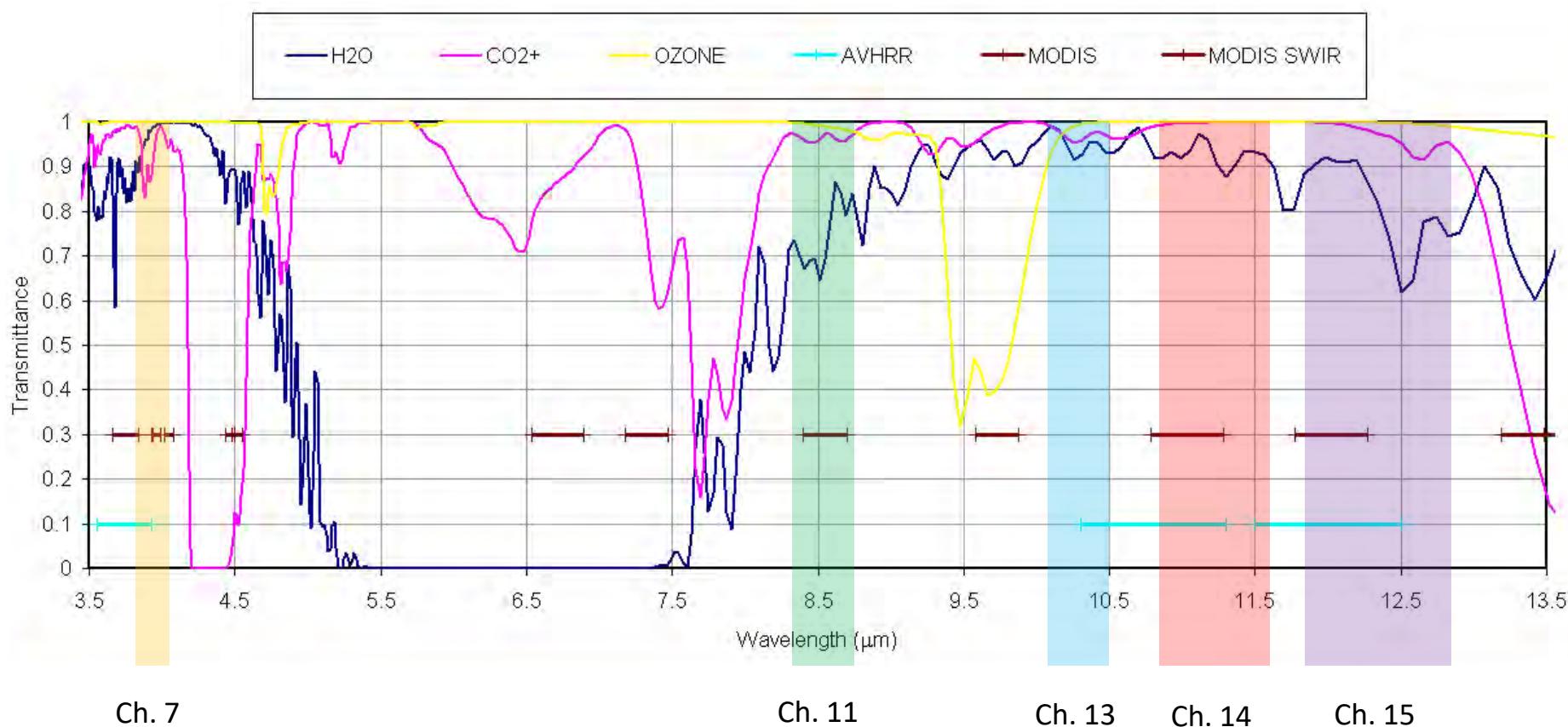


Therefore variations in the AVHRR Ch4 and Ch5 radiance/T<sub>B</sub> difference,  
for example  $T_B(11\mu\text{m}) - T_B(12\mu\text{m})$   
will be due to variations in column water vapor amount

**Is this difference > or < 0?**

AVHRR (NOAA-19) channels for SST are similar to those used by VIIRS (NOAA-20; M12, M13, M15, M16), except VIIRS has two SWIR channels.

None of these are in perfect windows so there will always be some contribution from the path term due primarily to absorption/emission by water vapor



GOES Channels used for SST algorithm

Does the addition of atmospheric water vapor **increase** or **decrease**  $T_B$ ?

Let's look at a simple case - addition of a homogeneous layer



$$L_t(\lambda, \theta, \phi) = \epsilon_s(\lambda, \theta)B(\lambda, T_s)\tau_d(\lambda) + \int_0^{\infty} B(\lambda, T(p)) \frac{d\tau_d(\lambda, p)}{dp} dp$$

Becomes,

$$L_t(\lambda, \theta, \phi) = B(\lambda, T_s)\tau_d(\lambda) + B(\lambda, T_s)d\tau_d(\lambda)$$

or,

$$L_t(\lambda, \theta, \phi) = B(\lambda, T_s)0.9 + B(\lambda, T_s)0.1$$

10% of the surface-emitted radiance

has been replaced with 10% of the “cooler” path-emitted radiance

$T_B$  **decreases** (determined by  $T_s$ ,  $T_a$ , and  $\tau_d$ )

How does the amount of water vapor affect the spectral variation of  $T_B$ ?

[referring to AVHRR Channel (4) and Channel (5)]

$$T_a, \tau_d(4) = 0.95, \tau_d(5) = 0.9$$

$$\delta(4) = -\ln(0.95) = 0.051, \delta(5) = 0.105$$

(for nadir view)

$$T_s, \varepsilon_s = 1$$

$$L_t(4, \theta, \varphi) = B(\lambda, T_s)0.95 + B(\lambda, T_s)0.05$$

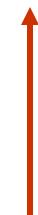
$$L_t(5, \theta, \varphi) = B(\lambda, T_s)0.9 + B(\lambda, T_s)0.1$$

$$L_t(4, \theta, \varphi) - L_t(5, \theta, \varphi) = [B(\lambda, T_s) - B(\lambda, T_s)]0.05$$

If we double the amount of water vapor, what changes?

$$\delta \text{ doubles so... } \delta(4) = 0.102, \delta(5) = 0.210$$

and,  $\tau_d(4) = 0.90, \tau_d(5) = 0.81$



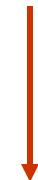
vs

Since  $T_s$  and  $T_a$  don't change with wavelength, now...

$$L_t(4, \theta, \varphi) = B(\lambda, T_s)0.90 + B(\lambda, T_s)0.10$$

$$L_t(5, \theta, \varphi) = B(\lambda, T_s)0.81 + B(\lambda, T_s)0.19$$

$$L_t(4, \theta, \varphi) - L_t(5, \theta, \varphi) = [B(\lambda, T_s) - B(\lambda, T_s)]0.09$$



## MCSST - MultiChannel Sea Surface Temperature

This technique assumes that the true surface temperature can be derived from a linear composite of the AVHRR Ch 4 and 5 brightness temperatures (accounting for water vapor variations)

$$T_s = A + B T_4 + C T_5$$

A, B, and C can then be determined empirically:

Measure  $T_s$  (ship, buoy, etc.) in many places coincident with measurements of  $T_4$  and  $T_5$  (AVHRR)

Statistically determine A, B, and C that produce a best fit of

$$T_s(x) = A + B T_4(x) + C T_5(x)$$

Or at night we can also use Ch 3 (3.7 $\mu$ m wavelength) (why?)

$$T_s(x) = D + E T_3(x) + F T_4(x) + G T_5(x)$$

The sea surface temperature lab exercise will give you an opportunity to examine MCSST calculations first hand

# AVHRR Split Window formulation for NOAA-19

Note the night versions that add the 3.7 micron window

**Day MCSST Split**

$$T_s = a_0 + a_1 * \text{band4} + a_2(\text{band4} - \text{band5}) + a_3(\text{band4} - \text{band5})(\sec(f) - 1)$$

**Night MCSST Split**

$$T_s = a_0 + a_1 * \text{band4} + a_2(\text{band4} - \text{band5}) + a_3(\text{band4} - \text{band5})(\sec(f) - 1)$$

**Night MCSST Dual**

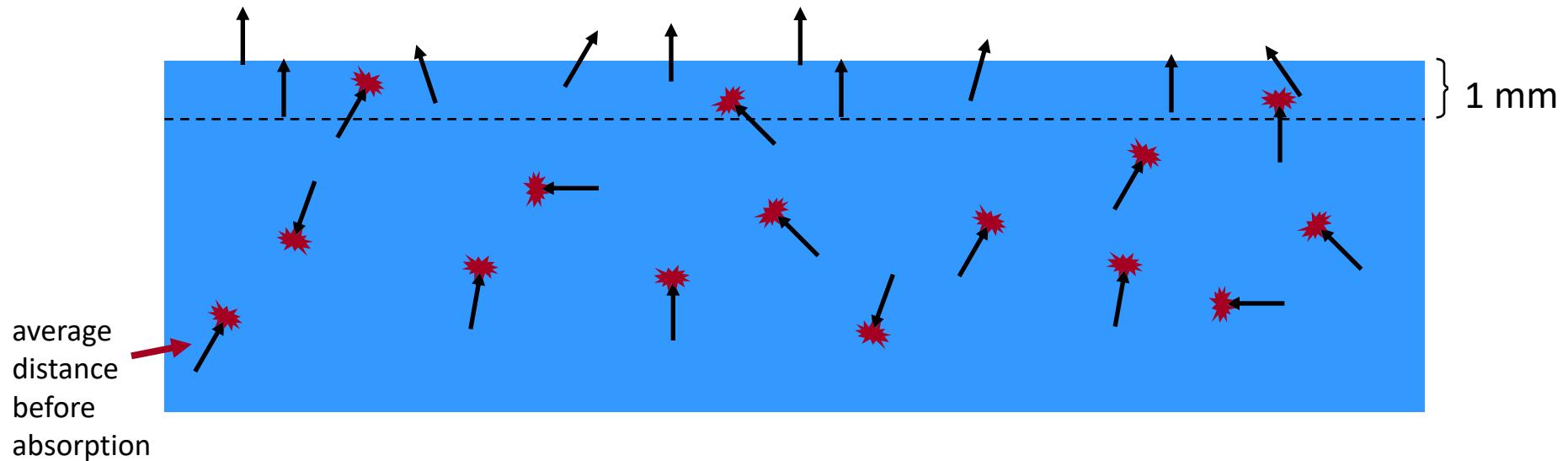
$$T_s = a_0 + a_1 * \text{band4} + a_2(\text{band3} - \text{band4}) + a_3(\sec(f) - 1)$$

**Night MCSST Triple**

$$T_s = a_0 + a_1 * \text{band4} + a_2(\text{band3} - \text{band5}) + a_3(\text{band3} - \text{band5})(\sec(f) - 1)$$

What part of the water column are we “sensing”?

How far can photons travel in water before they are absorbed?

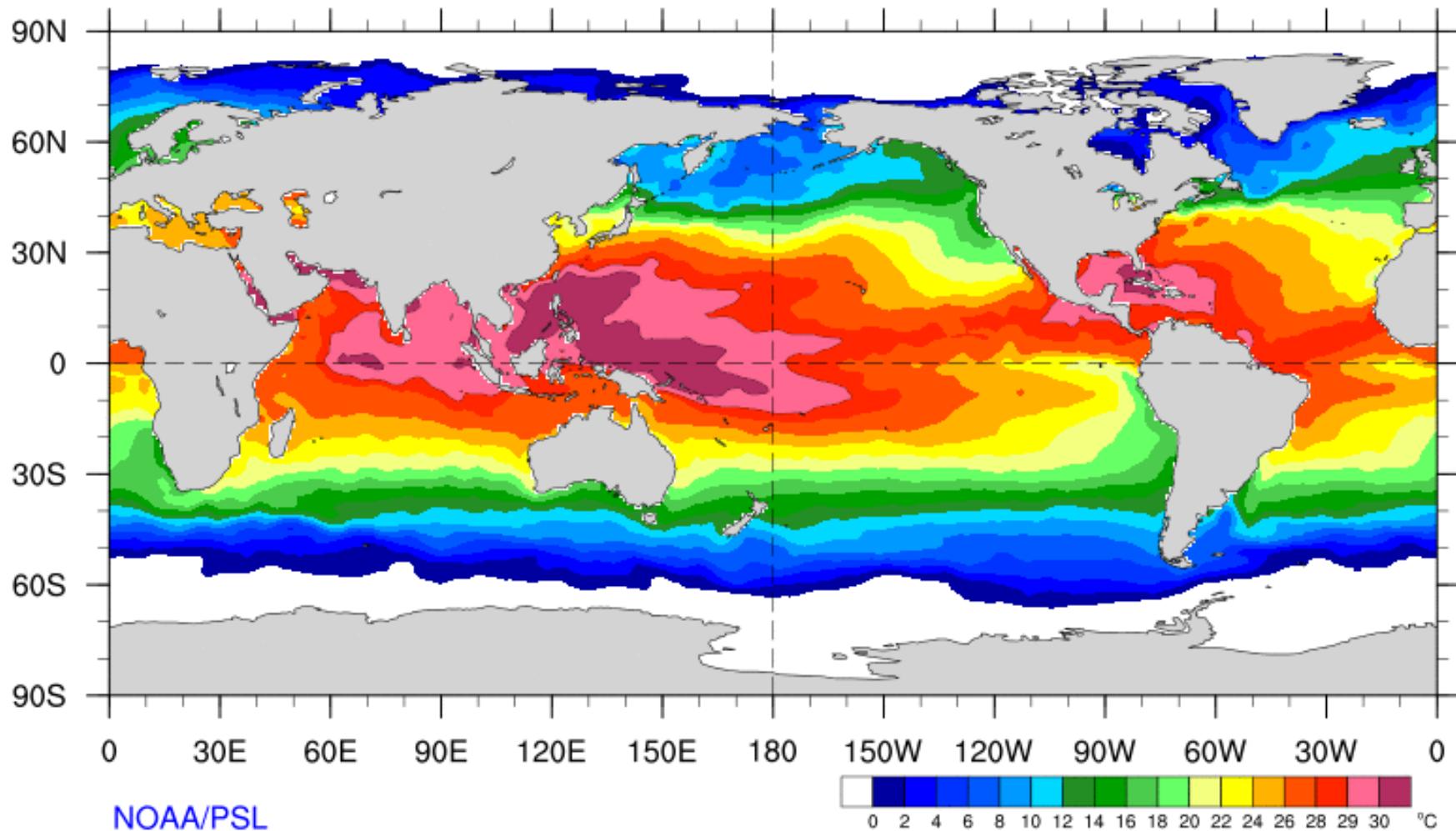


We are sensing just the “skin” temperature of the ocean

The advantage of the statistical technique is that  $T_B$  is correlated to bulk temperature measurements and the coefficients in MCSST correct for the difference between bulk and skin temperature.

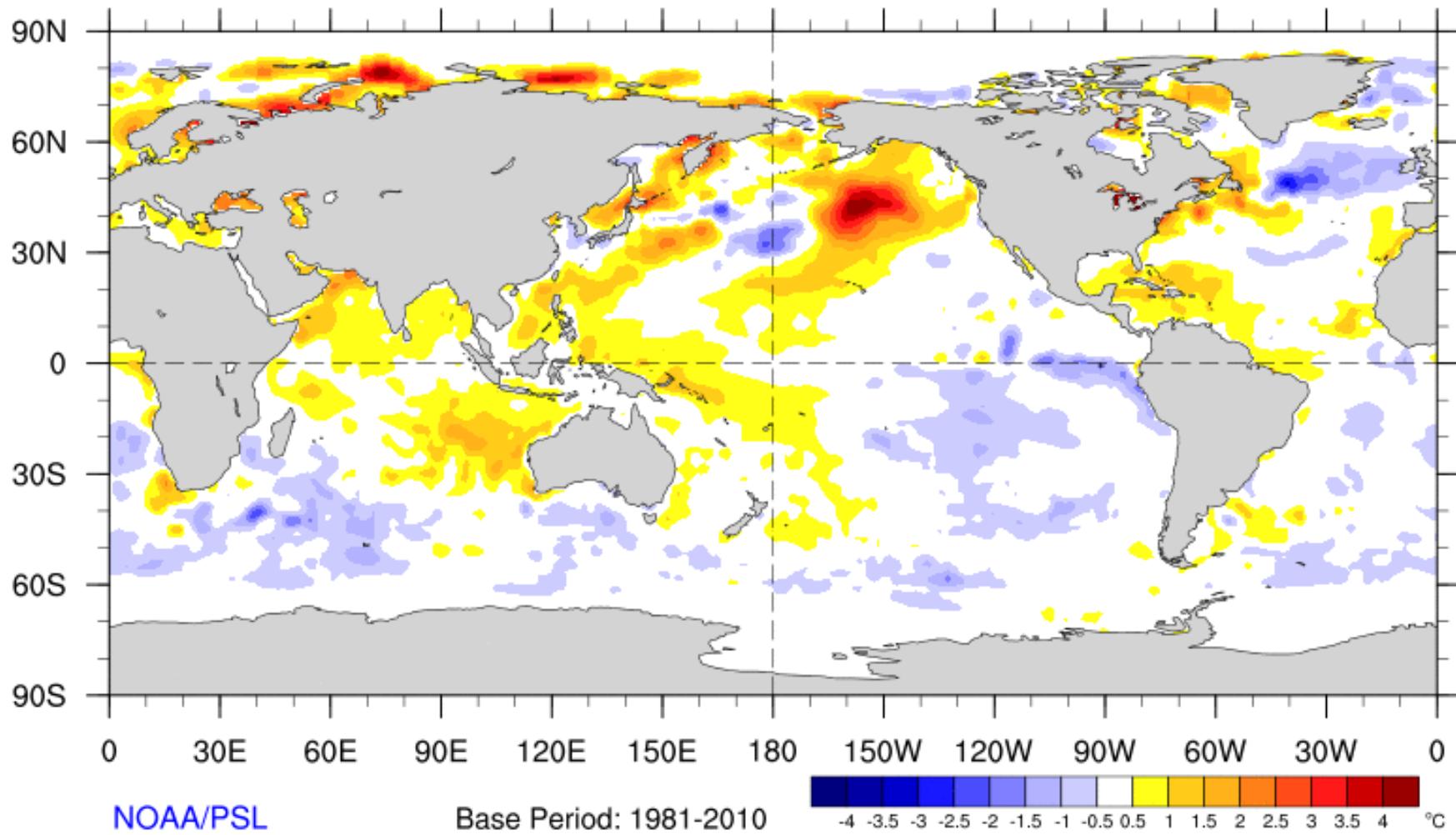
# Weekly Average SST

2020/07/05 - 2020/07/11

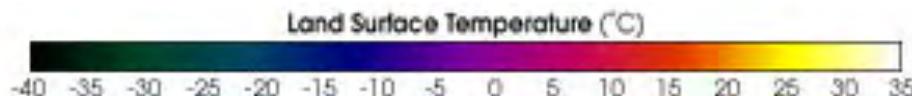
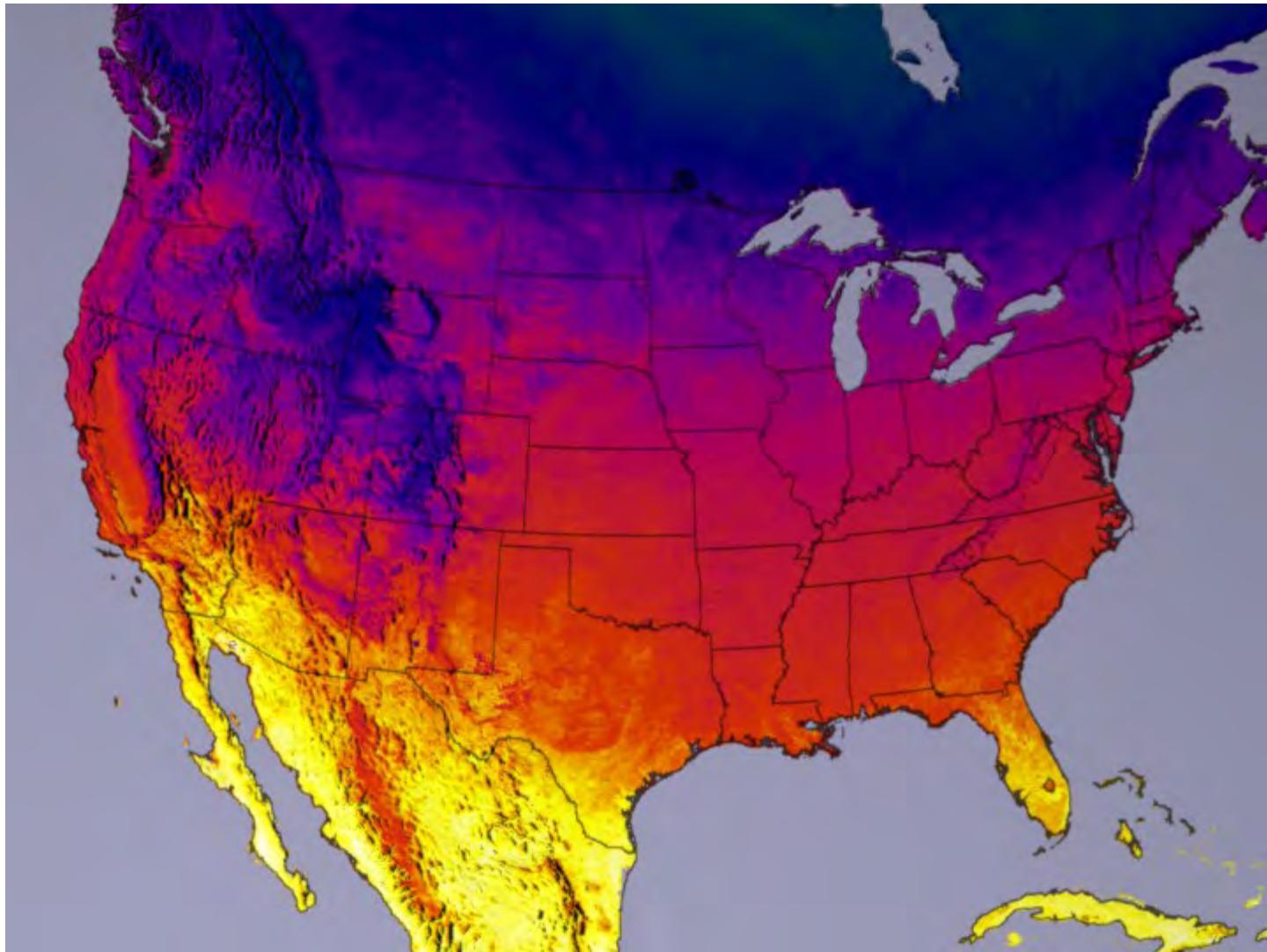


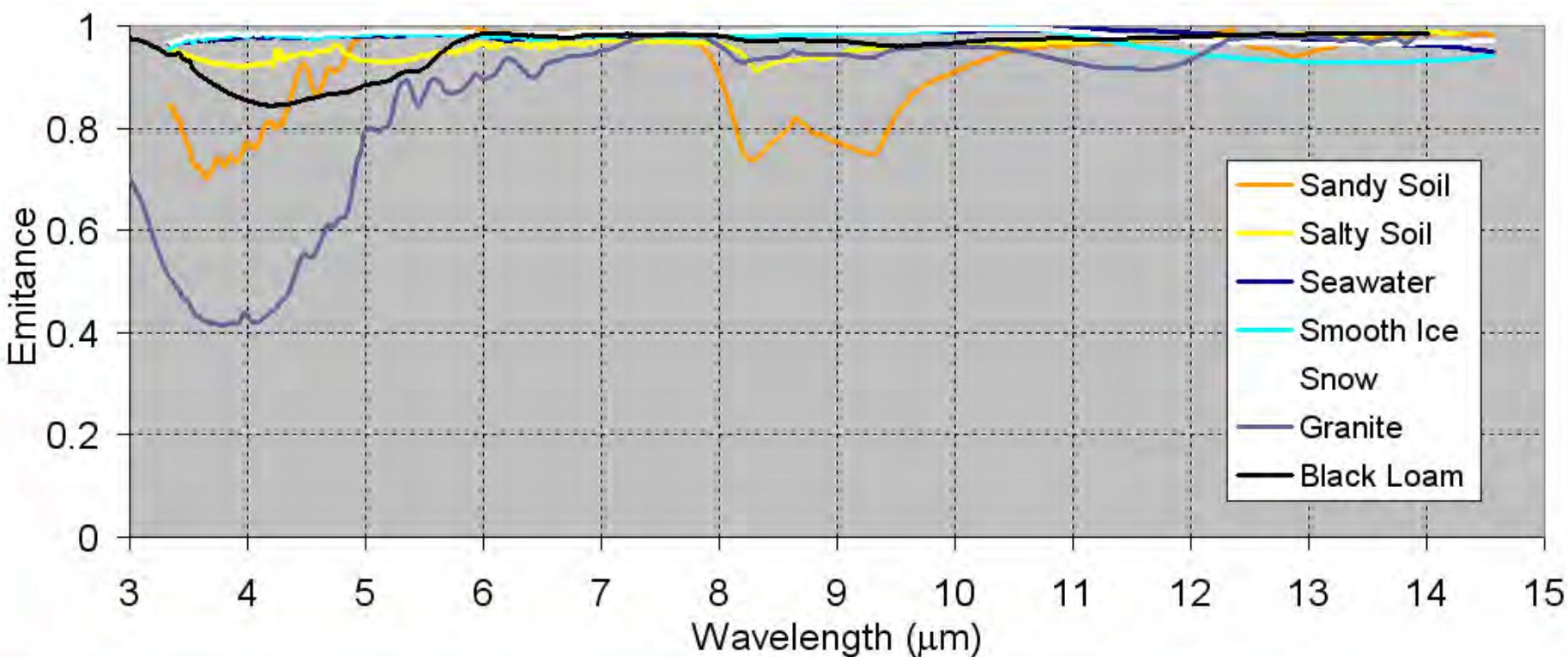
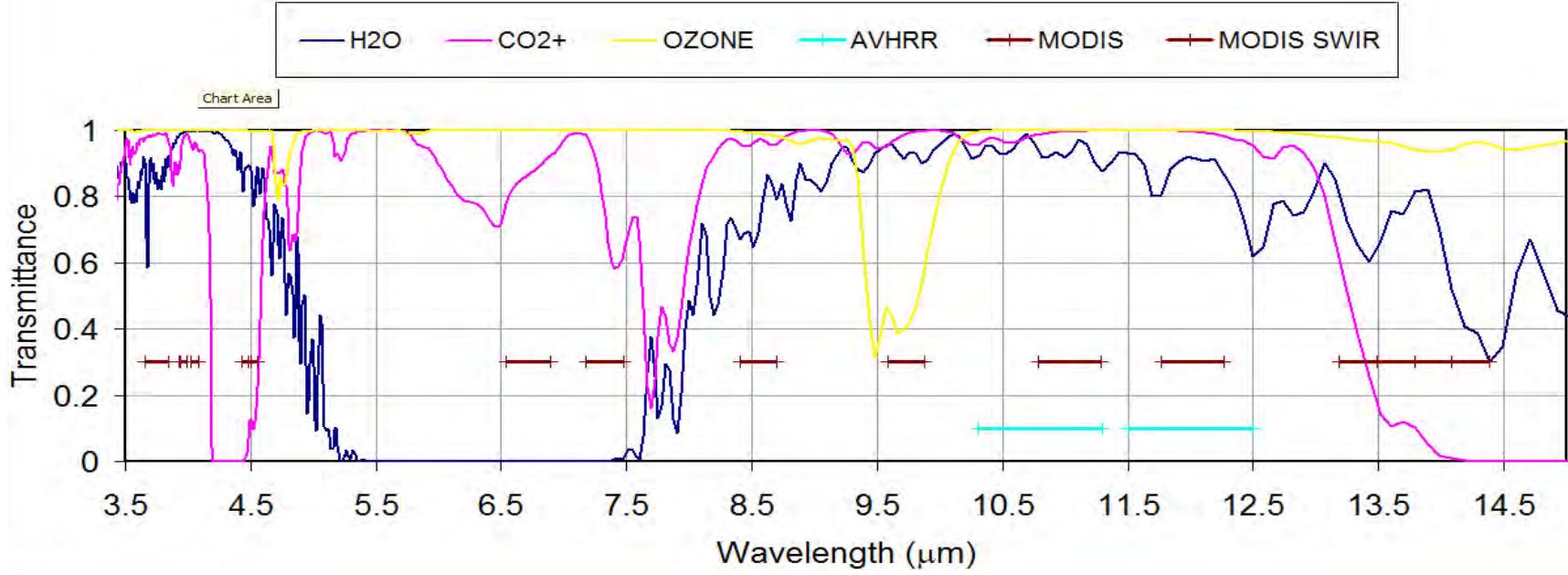
## Weekly SST Anomaly

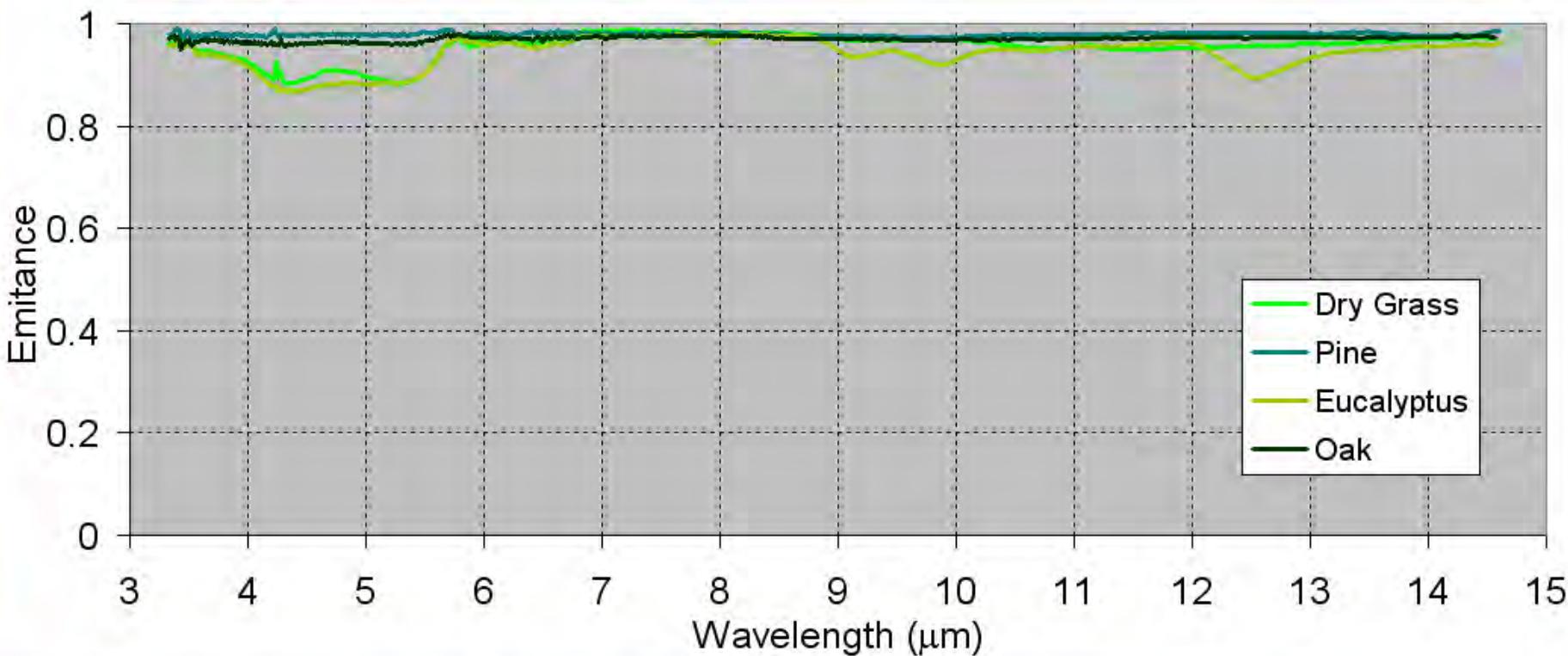
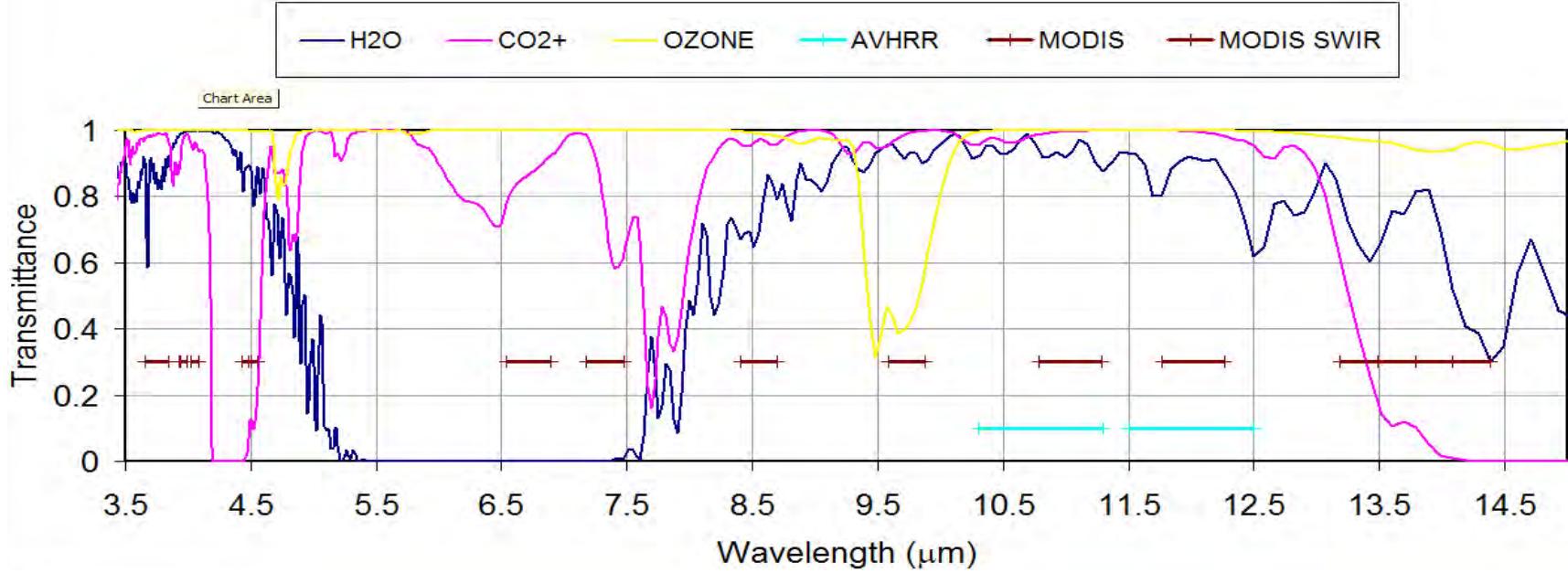
2020/07/05 - 2020/07/11

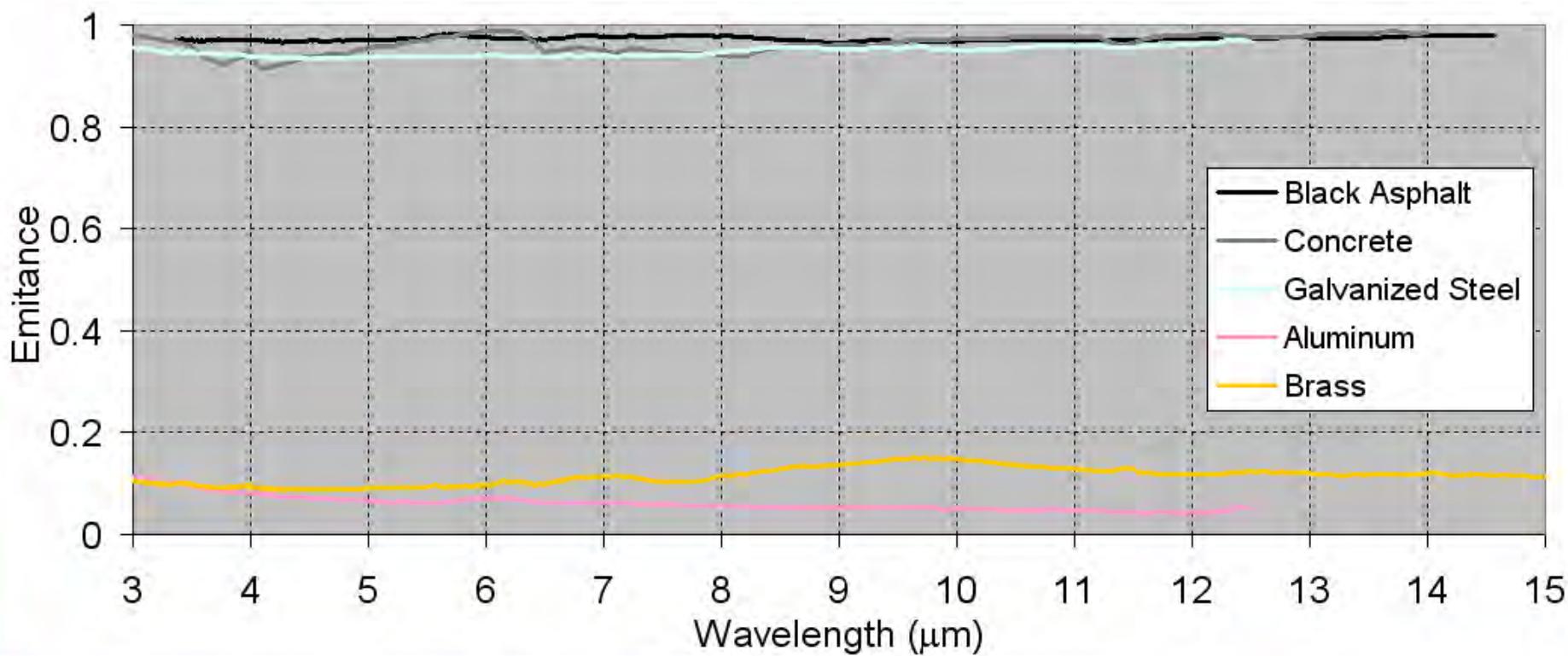
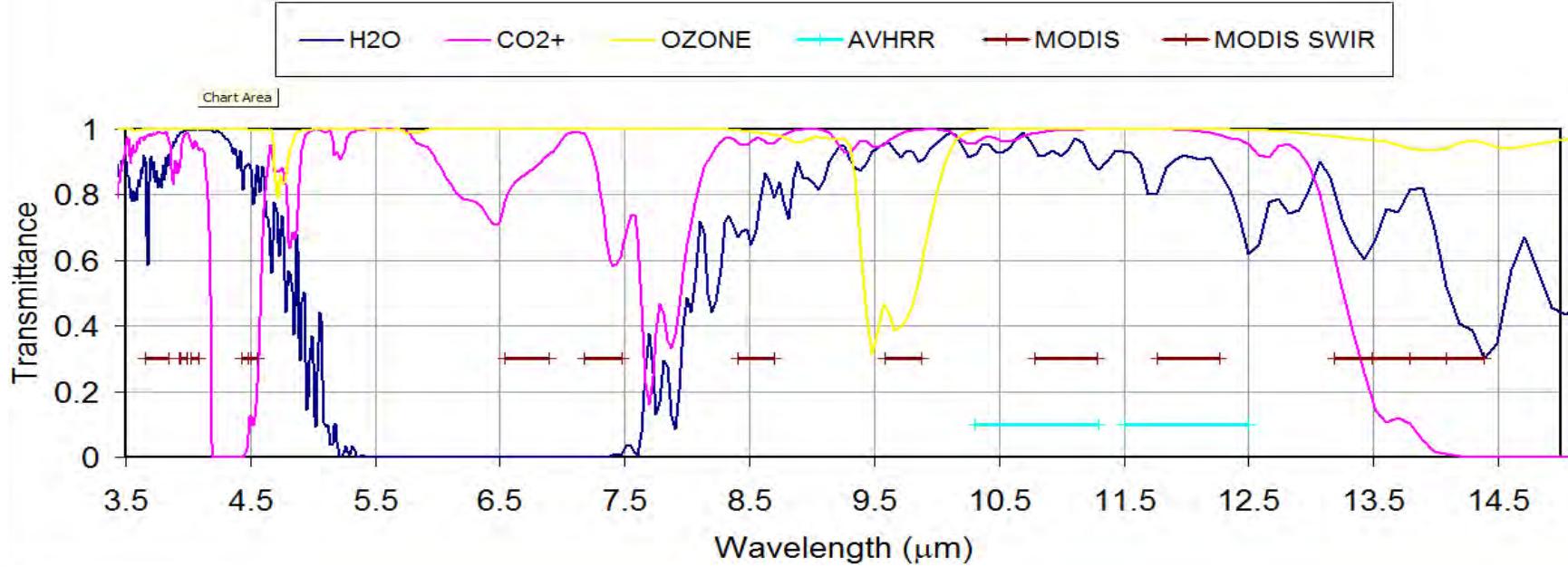


## What about **Land** Surface Temperature (LST)?







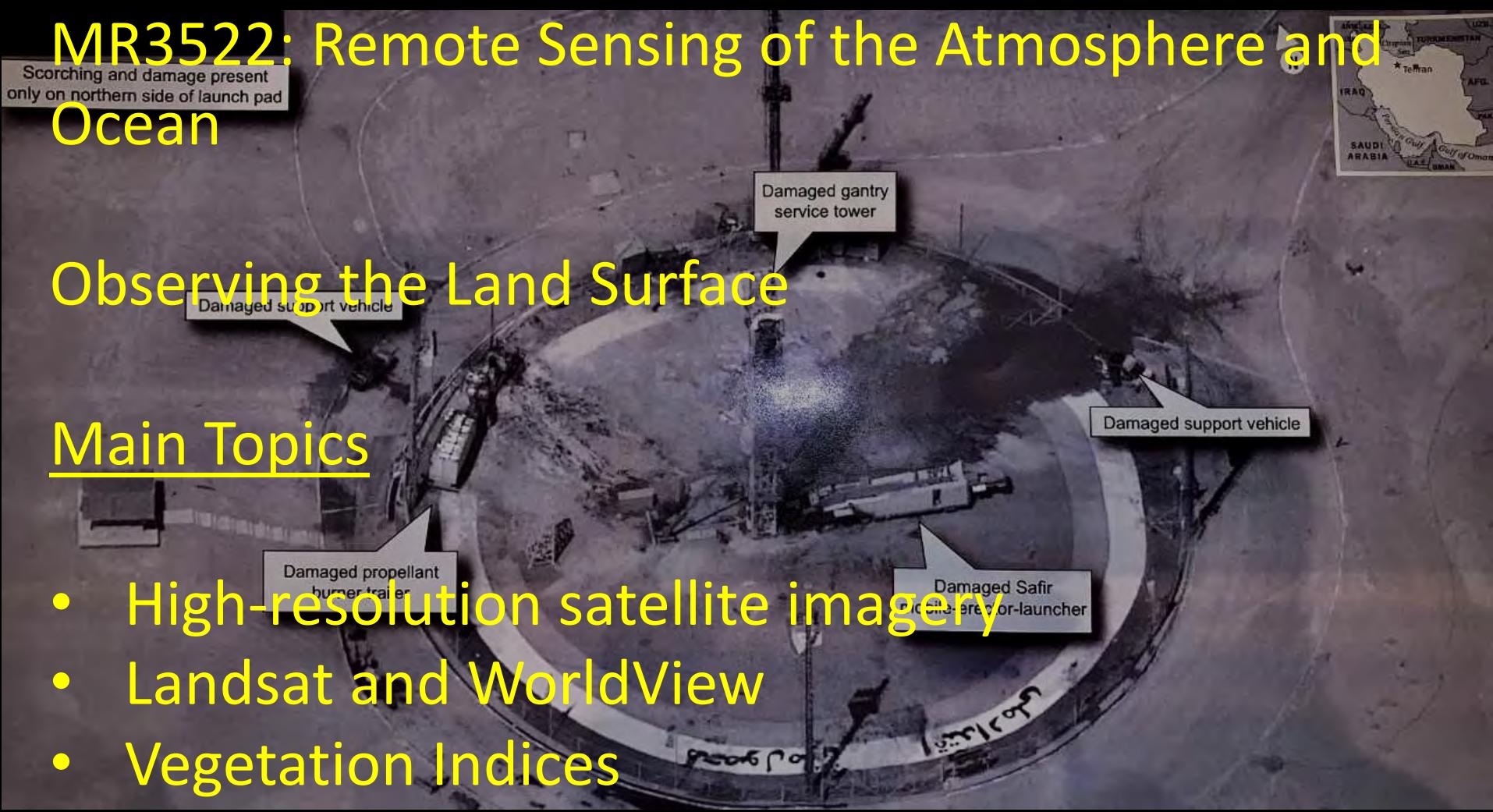


# MR3522: Remote Sensing of the Atmosphere and Ocean

## Observing the Land Surface

### Main Topics

- High-resolution satellite imagery
- Landsat and WorldView
- Vegetation Indices

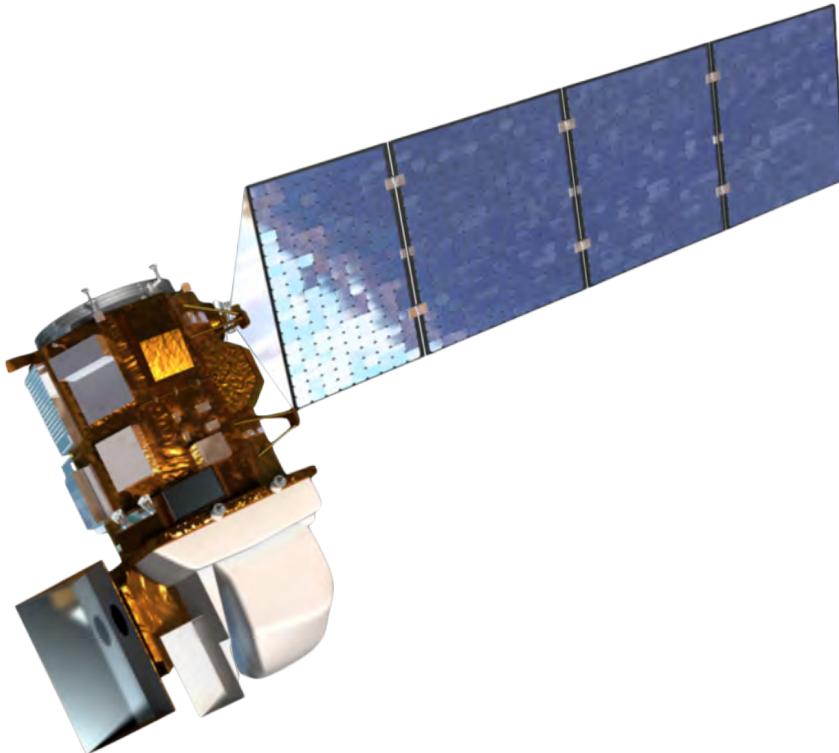


Credit: Twitter (@realdonaldtrump)

Previously classified imagery

*San Francisco, California, USA  
SPOT-5 5m Panchromatic 09-Aug-2002*





## Landsat 8

- Altitude: 705 km
- Inclination: 98.2°
- Period: 99 minutes
- Equatorial crossing (descending): 10:11am (Landsat7 at 10:00am)
- 2 instruments
  - Operational Land Imager (OLI)
    - 9 bands in VIS and near-IR
    - 30x30 meter spatial resolution (15x15 meter panchromatic)
  - Thermal Infrared Sensor (TIRS)
    - 2 bands in Earth IR
    - 100x100 meter spatial resolution

## Landsat 7 Bands (1999–present)

Band	Wavelength	Useful for mapping
Band 1 - Blue	0.45 - 0.52	Bathymetric mapping, distinguishing soil from vegetation, and deciduous from coniferous vegetation
Band 2 - Green	0.52 - 0.60	Emphasizes peak vegetation, which is useful for assessing plant vigor
Band 3 - Red	0.63 - 0.69	Discriminates vegetation slopes
Band 4 - Near Infrared	0.77 - 0.90	Emphasizes biomass content and shorelines
Band 5 - Short-wave Infrared	1.55 - 1.75	Discriminates moisture content of soil and vegetation; penetrates thin clouds
Band 6 - Thermal Infrared	10.40 - 12.50	Thermal mapping and estimated soil moisture
Band 7 - Short-wave Infrared	2.09 - 2.35	Hydrothermally altered rocks associated with mineral deposits
Band 8 - Panchromatic (Landsat 7 only)	0.52 - 0.90	15 meter resolution, sharper image definition

## Landsat 8 Bands (2013–present) also planned for Landsat-9 (launch in 2020)

<b>Band</b>	<b>Wavelength</b>	<b>Useful for mapping</b>
Band 1 – Coastal Aerosol	0.435 - 0.451	Coastal and aerosol studies
Band 2 – Blue	0.452 - 0.512	Bathymetric mapping, distinguishing soil from vegetation, and deciduous from coniferous vegetation
Band 3 - Green	0.533 - 0.590	Emphasizes peak vegetation, which is useful for assessing plant vigor
Band 4 - Red	0.636 - 0.673	Discriminates vegetation slopes
Band 5 - Near Infrared (NIR)	0.851 - 0.879	Emphasizes biomass content and shorelines
Band 6 - Short-wave Infrared (SWIR) 1	1.566 - 1.651	Discriminates moisture content of soil and vegetation; penetrates thin clouds
Band 7 - Short-wave Infrared (SWIR) 2	2.107 - 2.294	Improved moisture content of soil and vegetation and thin cloud penetration
Band 8 – Panchromatic (15m)	0.503 - 0.676	15 meter resolution, sharper image definition
Band 9 – Cirrus	1.363 - 1.384	Improved detection of cirrus cloud contamination
Band 10 – TIRS 1	10.60 – 11.19	100 meter resolution, thermal mapping and estimated soil moisture
Band 11 – TIRS 2	11.50 - 12.51	100 meter resolution, Improved thermal mapping and estimated soil moisture

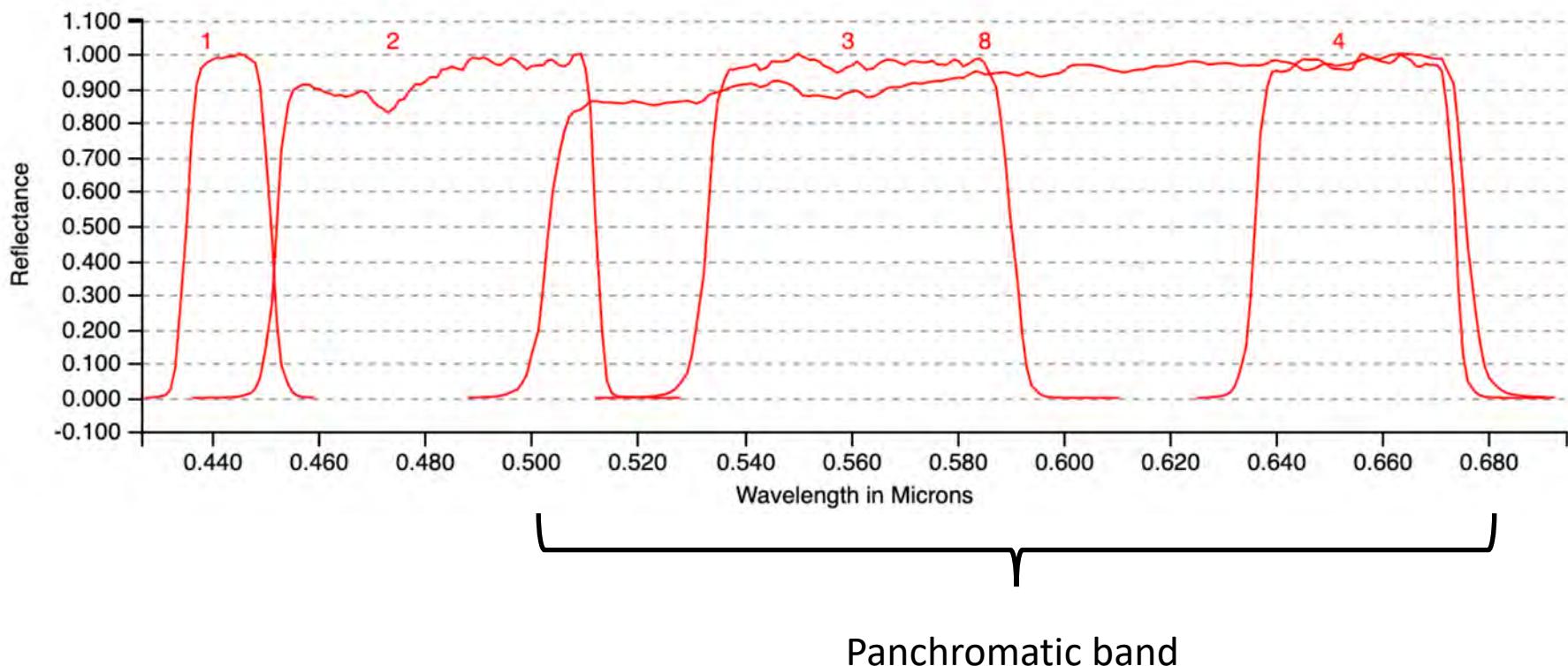
# Multispectral vs. Panchromatic Imagery

**Multispectral** imagery is derived from several discrete bands in the visible part of the EM spectrum. This is the type of true color imagery we have viewed so far.

**Panchromatic** imagery is derived from a single wide band in the visible part of the spectrum. This imagery usually has higher resolution than multispectral imagery because the large bandwidth (i.e. more total radiance detected by the sensor than in a narrow band) permits the use of smaller detectors.

**Pan-sharpening** is the process of combining multispectral and panchromatic data from a single instrument to create a single high resolution color image.

## Landsat-8 Spectral Response Functions



# WorldView-3

## Design and specifications

<b>Orbit</b>	Altitude: 617 km Type: SunSync, 10:30 am descending Node Period: 97 min.			
<b>Life</b>	Spec Mission Life: 7.25 years Estimated Service Life: 10 to 12 years			
<b>Spacecraft Size, Mass and Power</b>	Size: 5.7 m (18.7 ft) tall x 2.5 m (8 ft) across 7.1 m (23 ft) across deployed solar arrays Mass: 2800 kg (6200 lbs) Power: 3.1 kW solar array, 100 Ahr battery			
<b>Sensor Bands</b>	Panchromatic: 450 - 800 nm  8 Multispectral: Coastal: 400 - 450 nm Red: 630 - 690 nm Blue: 450 - 510 nm Red Edge: 705 - 745 nm Green: 510 - 580 nm Near-IR1: 770 - 895 nm Yellow: 585 - 625 nm Near-IR2: 860 - 1040 nm	<b>Swath Width</b>	At nadir: 13.1 km	
	8 SWIR Bands: SWIR-1: 1195 - 1225 nm SWIR-5: 2145 - 2185 nm SWIR-2: 1550 - 1590 nm SWIR-6: 2185 - 2225 nm SWIR-3: 1640 - 1680 nm SWIR-7: 2235 - 2285 nm SWIR-4: 1710 - 1750 nm SWIR-8: 2295 - 2365 nm	<b>Attitude Determination and Control</b>	Type: 3-axis Stabilized Actuators: Control Moment Gyros (CMGs) Sensors: Star trackers, precision IRU, GPS	
	12 CAVIS Bands: Desert Clouds: 405 - 420 nm Water-3: 930 - 965 nm Aerosol-1: 459 - 509 nm NDVI-SWIR: 1220 - 1252 nm Green: 525 - 585 nm Cirrus: 1365 - 1405 nm Aerosol-2: 635 - 685 nm Snow: 1620 - 1680 nm Water-1: 845 - 885 nm Aerosol-1: 2105 - 2245 nm Water-2: 897 - 927 nm Aerosol-2: 2105 - 2245 nm	<b>Pointing Accuracy and Knowledge</b>	Accuracy: <500 m at image start/stop Knowledge: Supports geolocation accuracy below	
		<b>Retargeting Agility</b>	Time to Slew 200 km: 12 sec	
		<b>Onboard Storage</b>	2199 Gb solid state with EDAC	
		<b>Communications</b>	Image & Ancillary Data: 800 and 1200 Mbps X-band Housekeeping: 4, 16, 32, or 64 kbps real time, 524 kbps stored, X-band Command: 2 or 64 kbps S-band	
<b>Sensor Resolution (or GSD, Ground Sample Distance; off-nadir is geometric mean)</b>	Panchromatic Nadir: 0.31 m 20° Off-Nadir: 0.34 m Multispectral Nadir: 1.24 m 20° Off-Nadir: 1.38 m SWIR Nadir: 3.70 m 20° Off-Nadir: 4.10 m CAVIS Nadir: 30.00 m	<b>Max Contiguous Area Collected in a Single Pass (30° off-nadir angle)</b>	Mono: 66.5 km x 112 km (5 strips) Stereo: 26.6 km x 112 km (2 pairs)	
<b>Dynamic Range</b>	11-bits per pixel Pan and MS; 14-bits per pixel SWIR	<b>Revisit Frequency (at 40°N Latitude)</b>	1 m GSD: <1.0 day 4.5 days at 20° off-nadir or less	
		<b>Geolocation Accuracy (CE90)</b>	Predicted <3.5 m CE90 without ground control	
		<b>Capacity</b>	680,000 km² per day	

## Scattering (Reflectance of solar radiation)

- Applies to Bands 1 through 9

$$L_t(\lambda, \theta, \phi) = L_0(\lambda, \theta, \phi) e^{-\delta(\lambda)/\mu} + \int_0^{\delta(\lambda)} \frac{\int_{4\pi} \gamma_s(\mathbf{r}, \mathbf{r}', \lambda, \mathbf{X}) L(\mathbf{r}', \lambda, \mathbf{X}) d\Omega' \sigma_e(\lambda, z)}{d\Omega'} e^{-\delta(\lambda, z)/\mu} \frac{d\delta}{\mu}$$

To derive land properties, we want the path radiance  
to have a small and removable contribution:

Three sources:

scatter by clouds (**irremovable**)

Not as important because of higher  $L_0(\lambda)$  at most  $\lambda_s$

→ { scatter by molecules (**calculable** as Rayleigh scatter)  
scatter by aerosol particles (**implied** from some NIR channels)

## Emission (Radiation emitted directly by objects on Earth)

- Applies to Bands 10 and 11

$$L_t(\lambda, \theta, \phi) = \varepsilon_s B(\lambda, T_s) e^{-\delta(\lambda)/\mu} + \int_0^{\delta(\lambda)} B(\lambda, T(z)) e^{-\delta(\lambda, z)/\mu} \frac{d\delta}{\mu}$$



2013

<https://landsatlook.usgs.gov/>



Landsat 8 viewing creation a man-made island out of a reef in the Spratly Islands

2015

WorldView-3  
March 2020

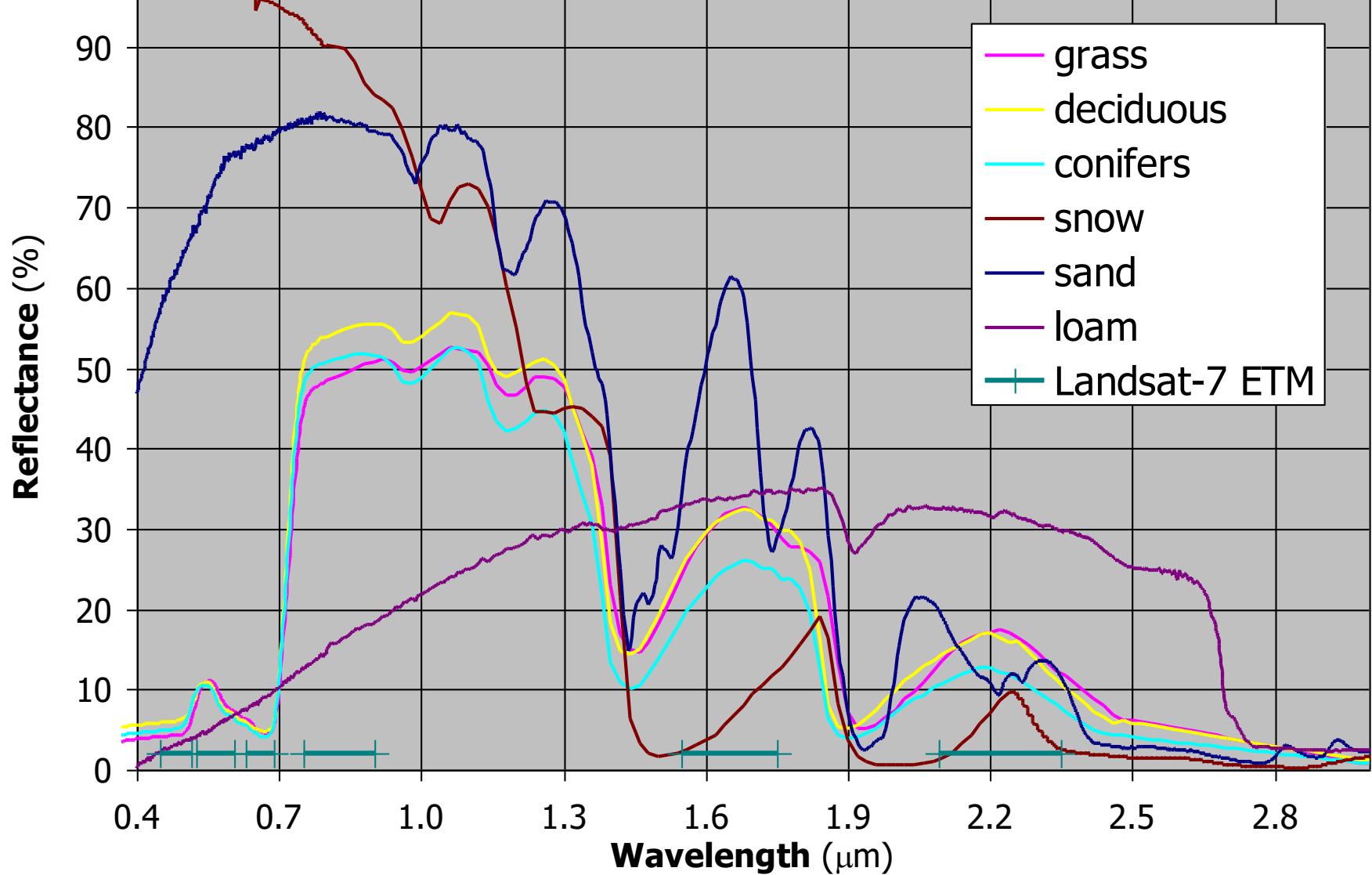




Santa Lucia  
Mountains  
(2019)

# Surface Reflectance

(data from ASTER Spectral Library - [speclib.jpl.nasa.gov](http://speclib.jpl.nasa.gov))



# Detecting Vegetation

- One widely used index is the Normalized Difference Vegetation Index (NDVI). The combination of near-IR and red-visible data make this sensitive to chlorophyll.

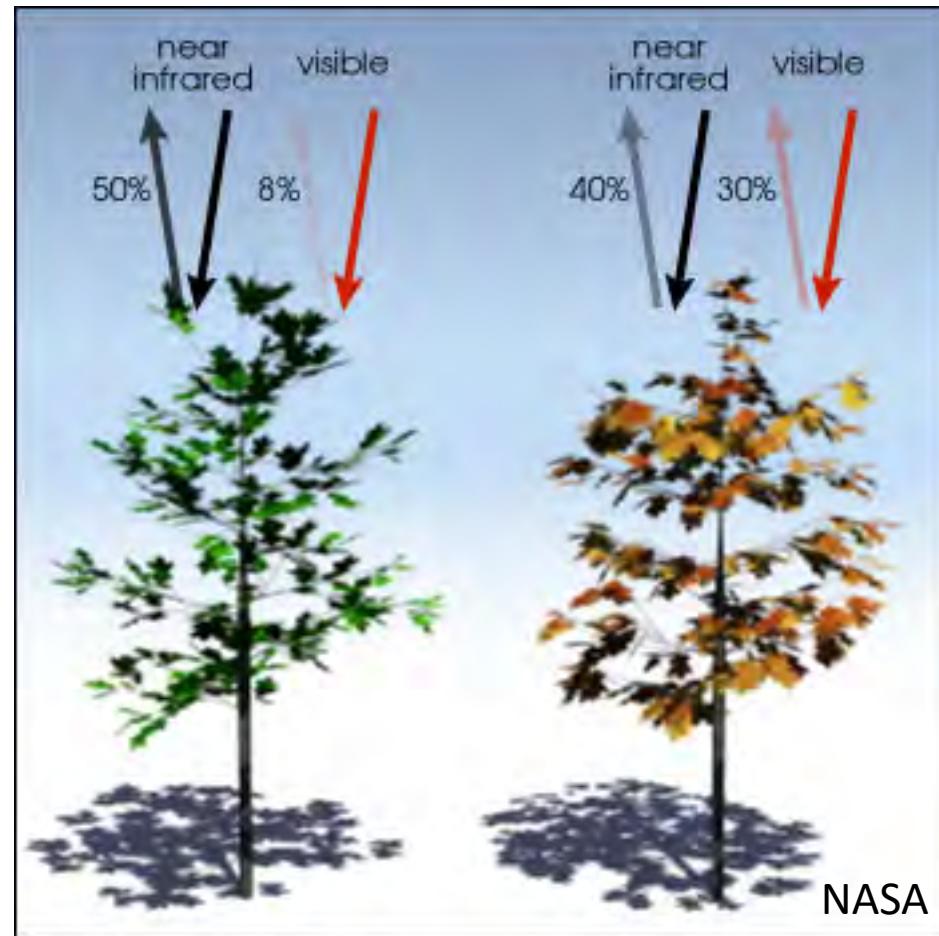
$$NDVI = \frac{NIR - VIS}{NIR + VIS}$$

NIR = Near-IR reflectance

VIS = Visible (Usually red band)  
reflectance

This can also be computed using  
GOES data. Which bands would  
you use?

Some limitations: Often has lots  
of atmospheric induced or  
background noise; needs to be  
smoothed heavily.



$$\frac{(0.50 - 0.08)}{(0.50 + 0.08)} = 0.72$$

$$\frac{(0.4 - 0.30)}{(0.4 + 0.30)} = 0.14$$

# Detecting Vegetation

- A complimentary alternative to NDVI is the EVI, or Enhanced Vegetation Index. It is more responsive to variations in canopy properties and leaf area index.

$$EVI = \frac{NIR - RED}{NIR + a * RED - b * BLUE + c}$$

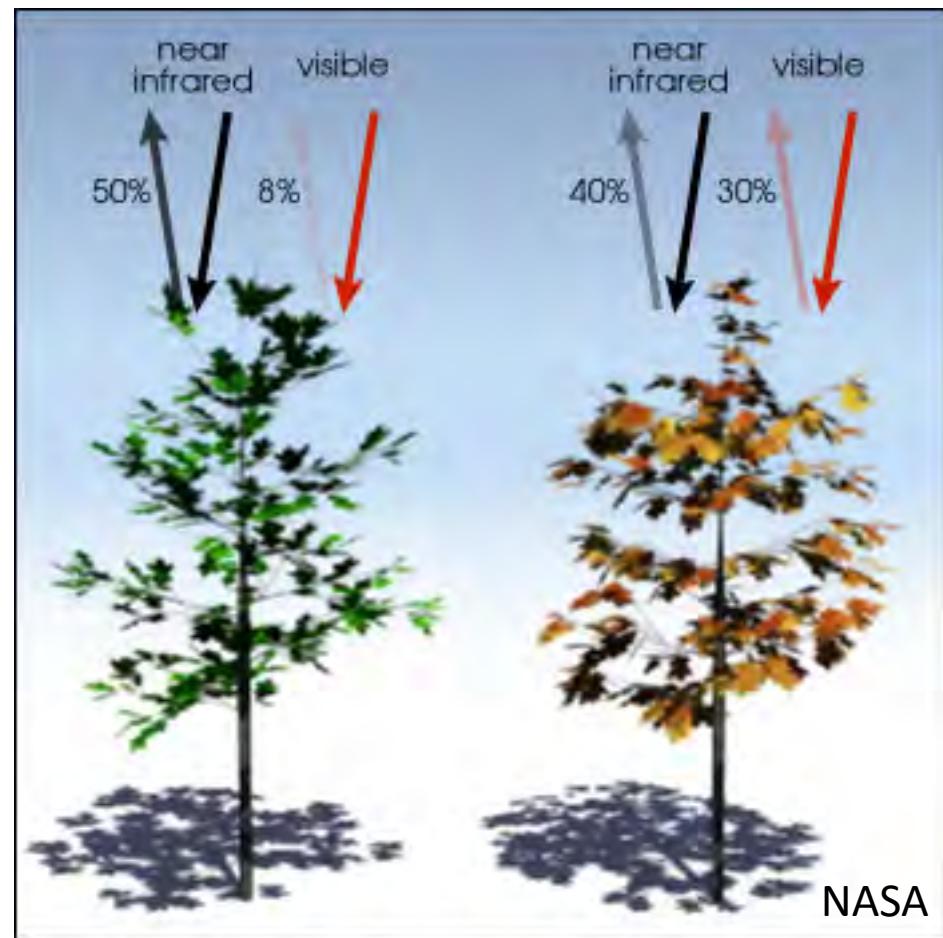
NIR = Near-IR reflectance

RED = red-band reflectance

BLUE = blue-band reflectance

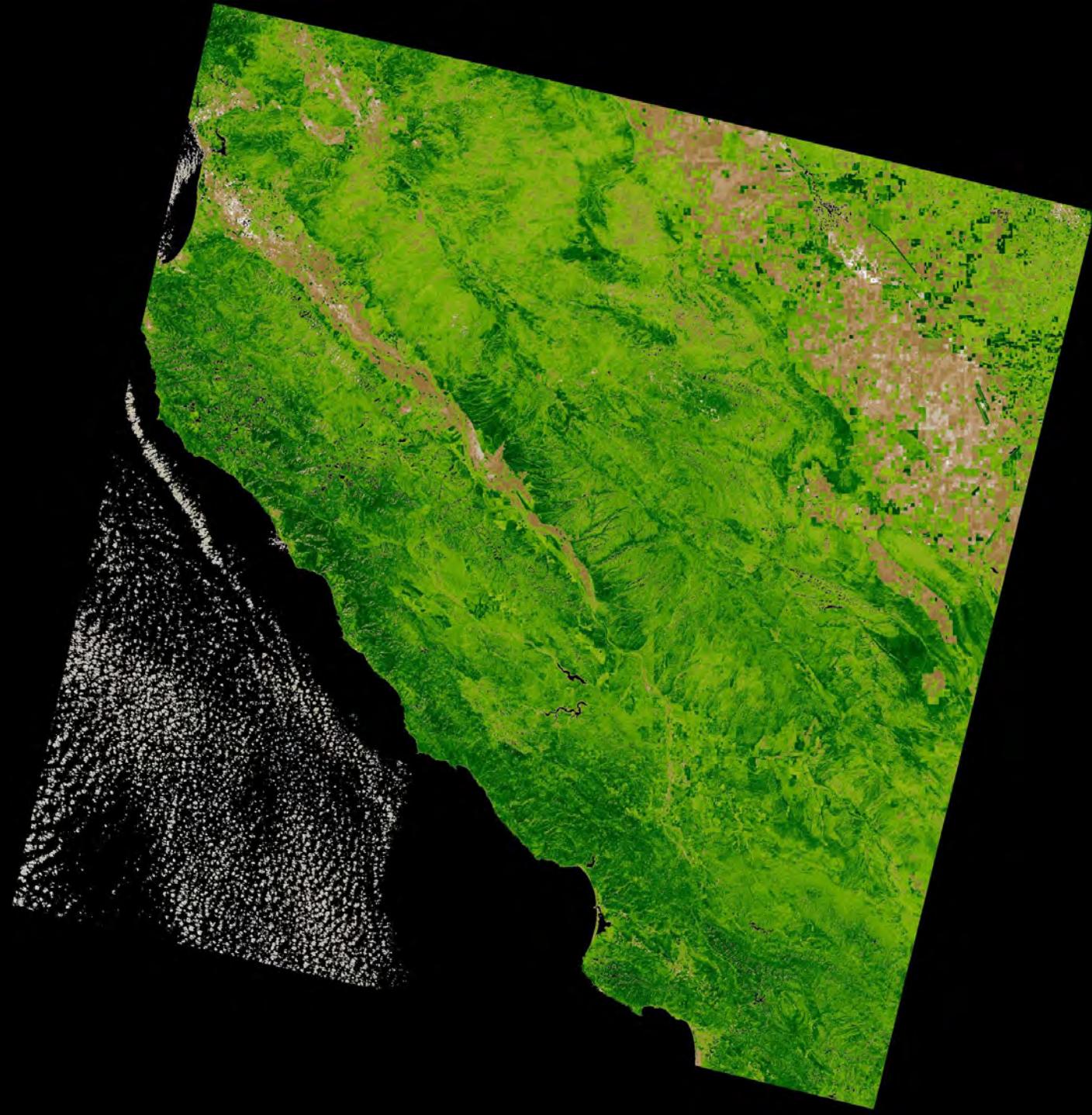
a, b, and c are coefficients that depend on satellite

Some limitations: Signal to noise ratio in blue band is often low; also some old sensors did not have a blue band.



$$\frac{(0.50 - 0.08)}{(0.50 + 0.08)} = 0.72$$

$$\frac{(0.4 - 0.30)}{(0.4 + 0.30)} = 0.14$$



NDVI