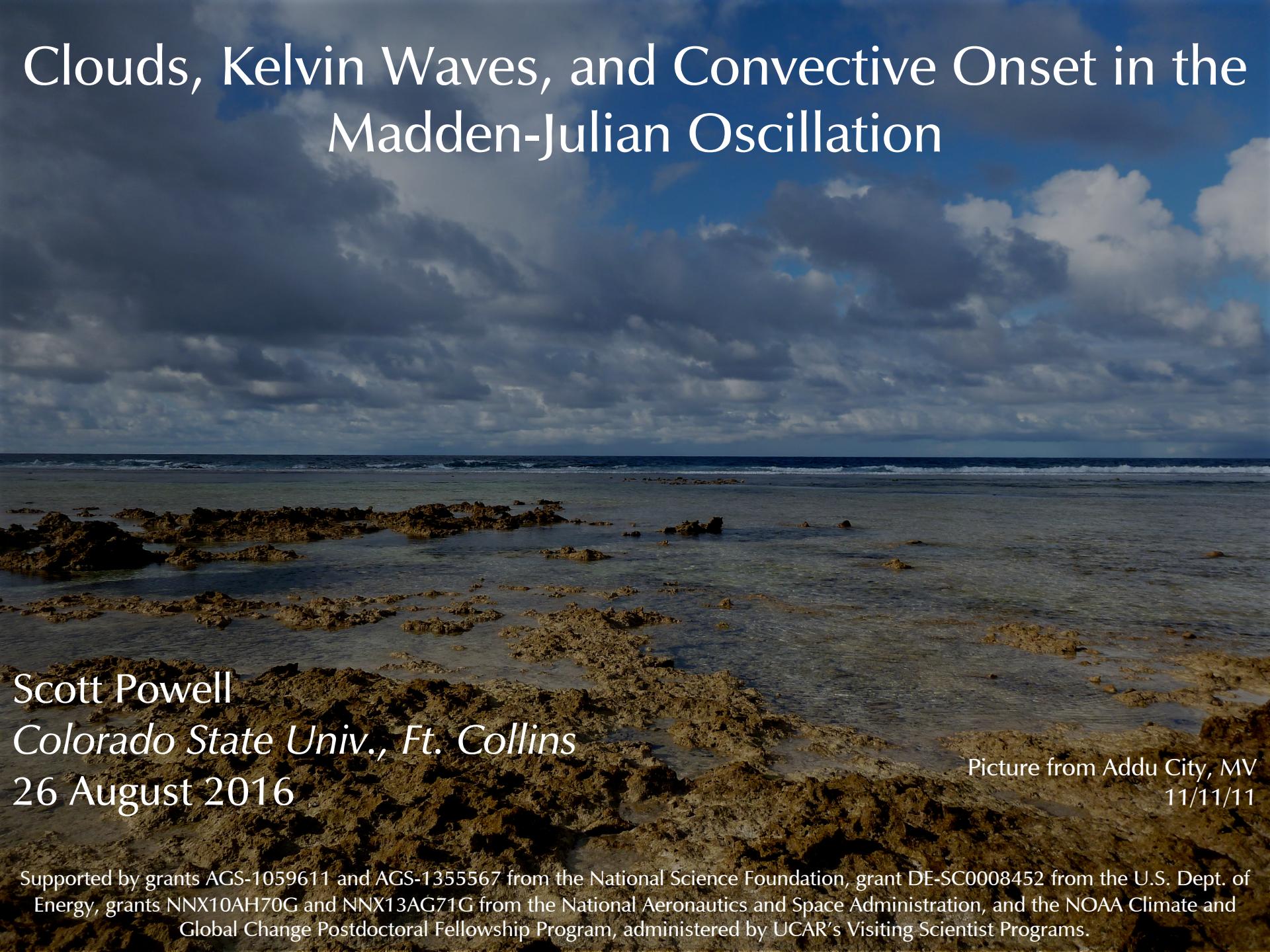


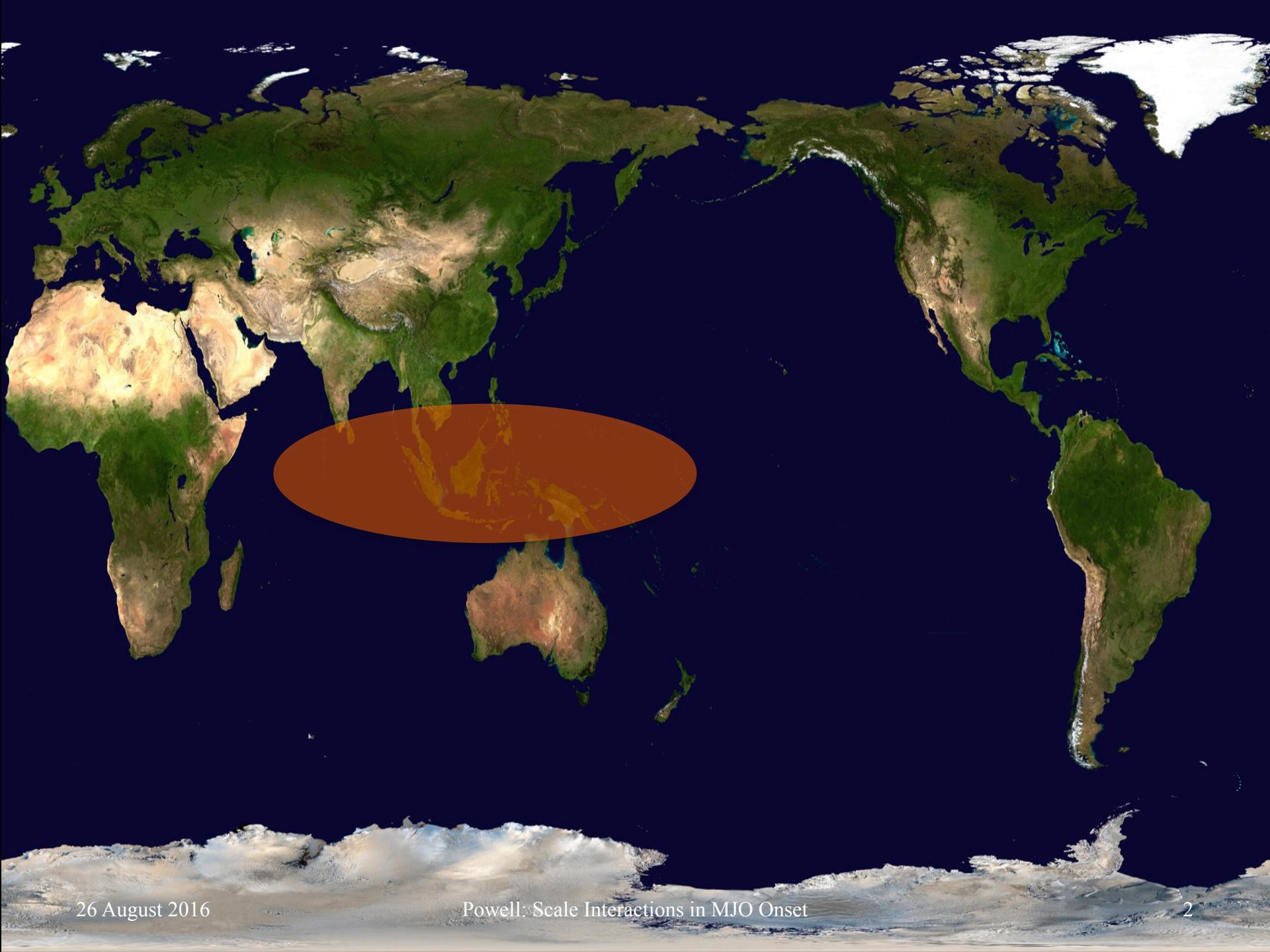
Clouds, Kelvin Waves, and Convective Onset in the Madden-Julian Oscillation



Scott Powell
Colorado State Univ., Ft. Collins
26 August 2016

Picture from Addu City, MV
11/11/11

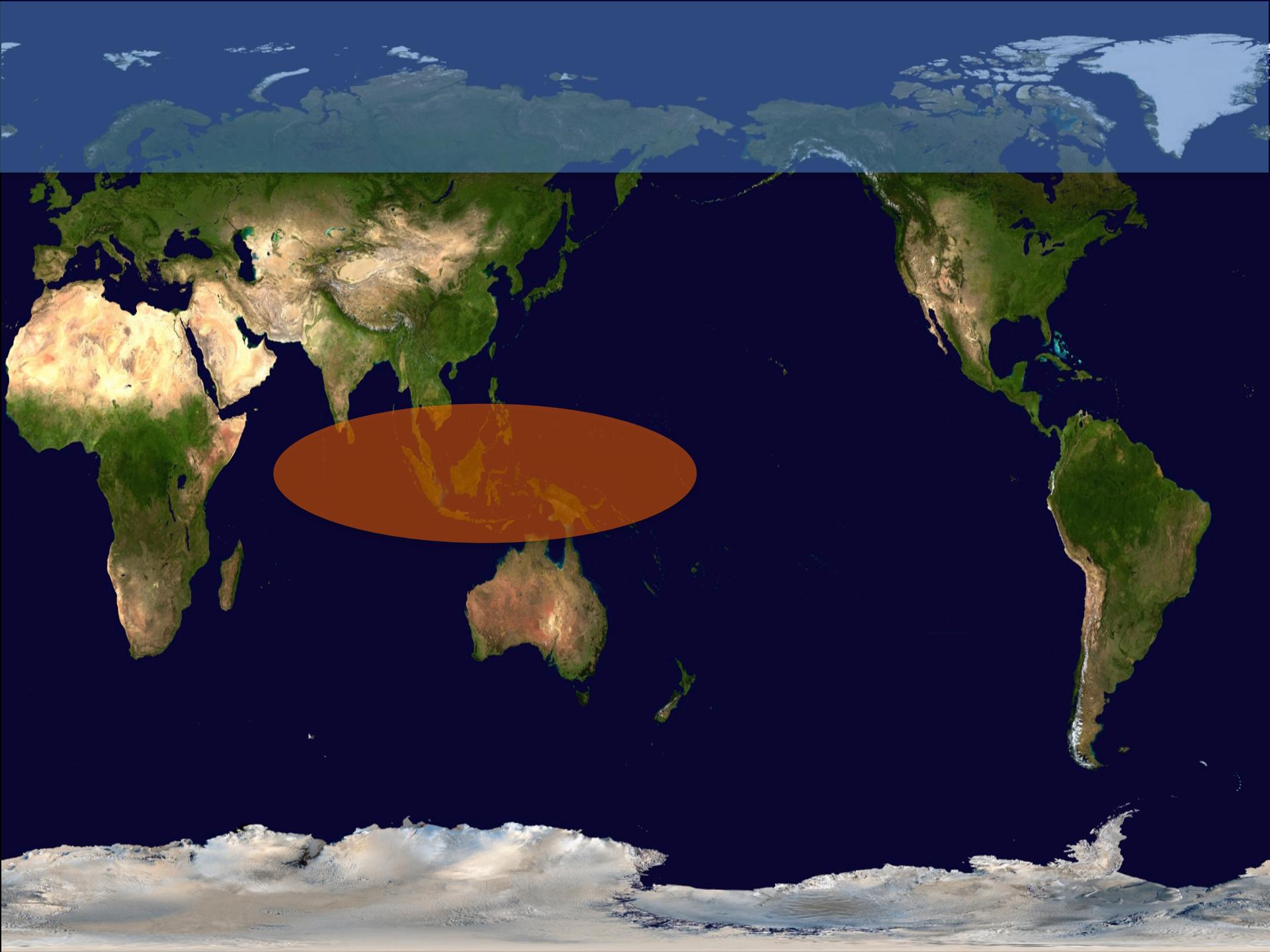
Supported by grants AGS-1059611 and AGS-1355567 from the National Science Foundation, grant DE-SC0008452 from the U.S. Dept. of Energy, grants NNX10AH70G and NNX13AG71G from the National Aeronautics and Space Administration, and the NOAA Climate and Global Change Postdoctoral Fellowship Program, administered by UCAR's Visiting Scientist Programs.

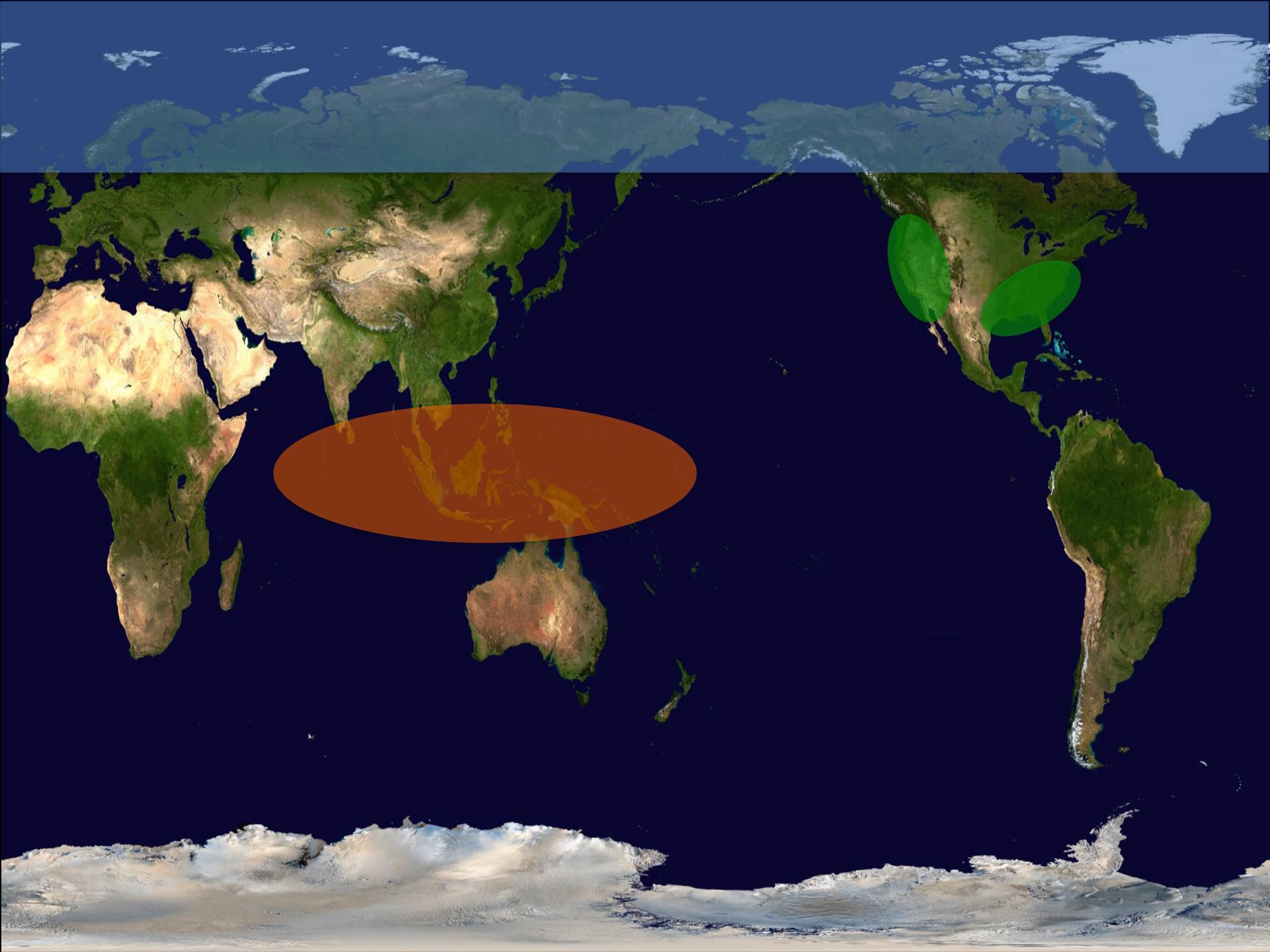


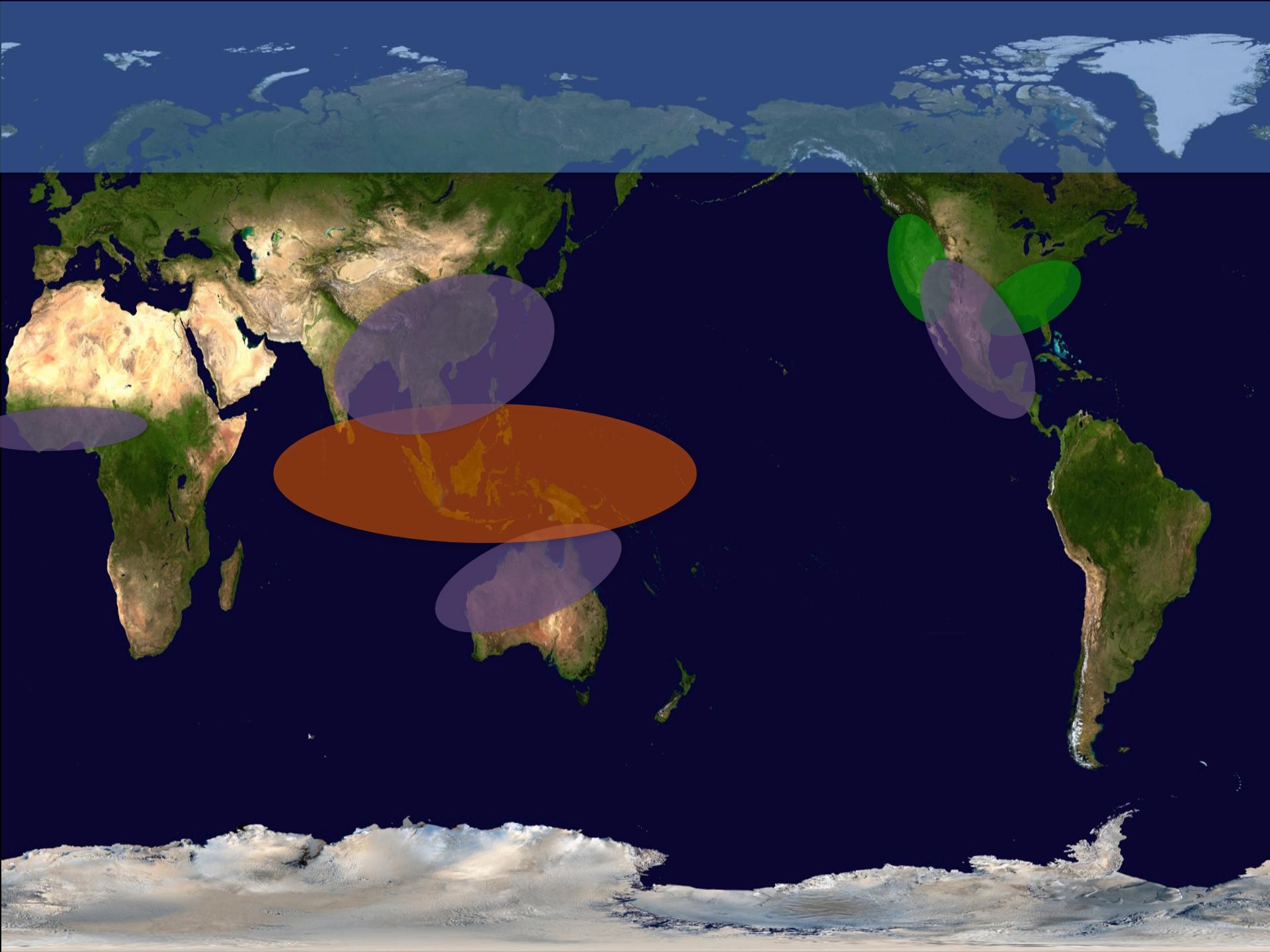
26 August 2016

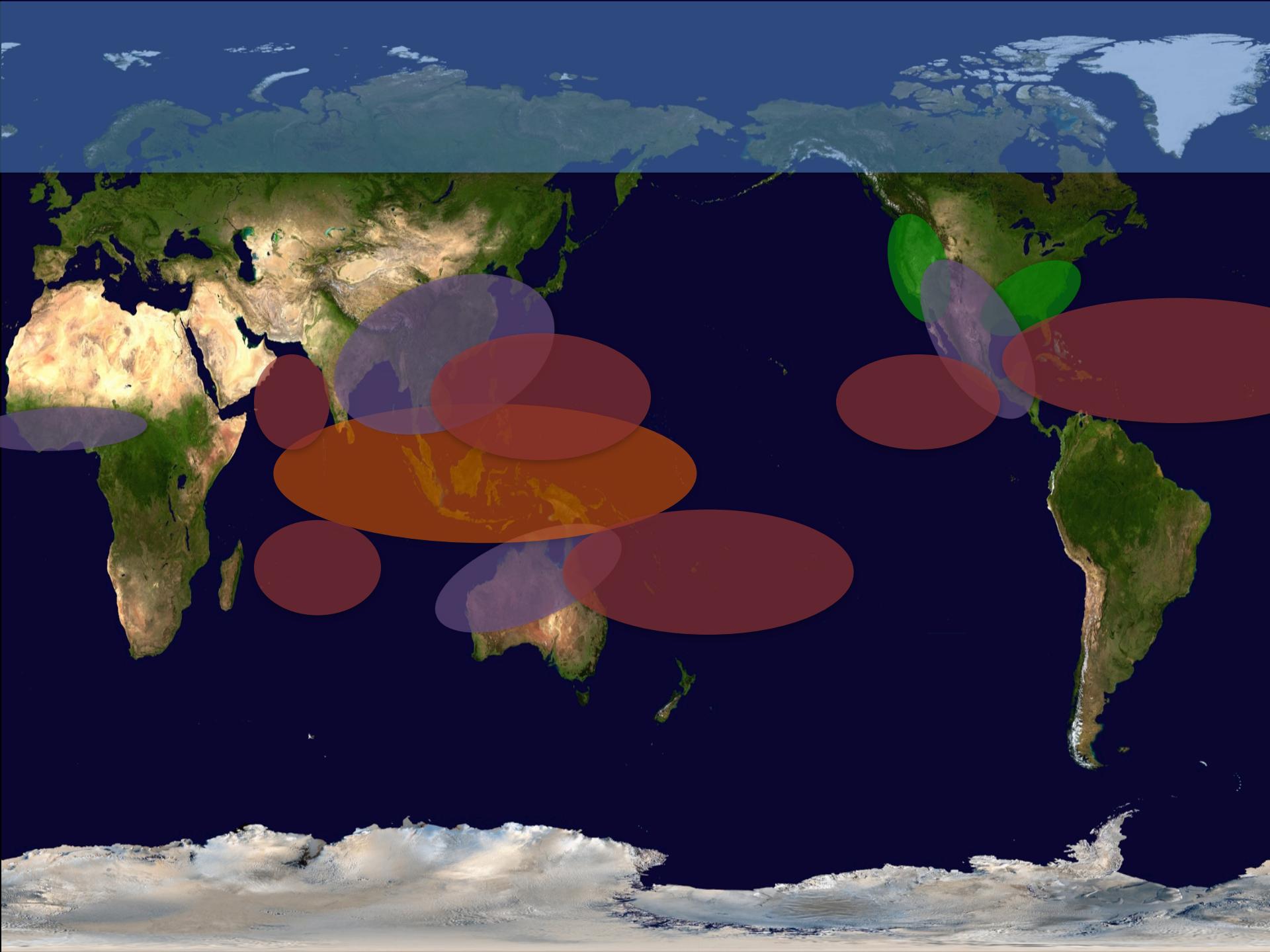
Powell: Scale Interactions in MJO Onset

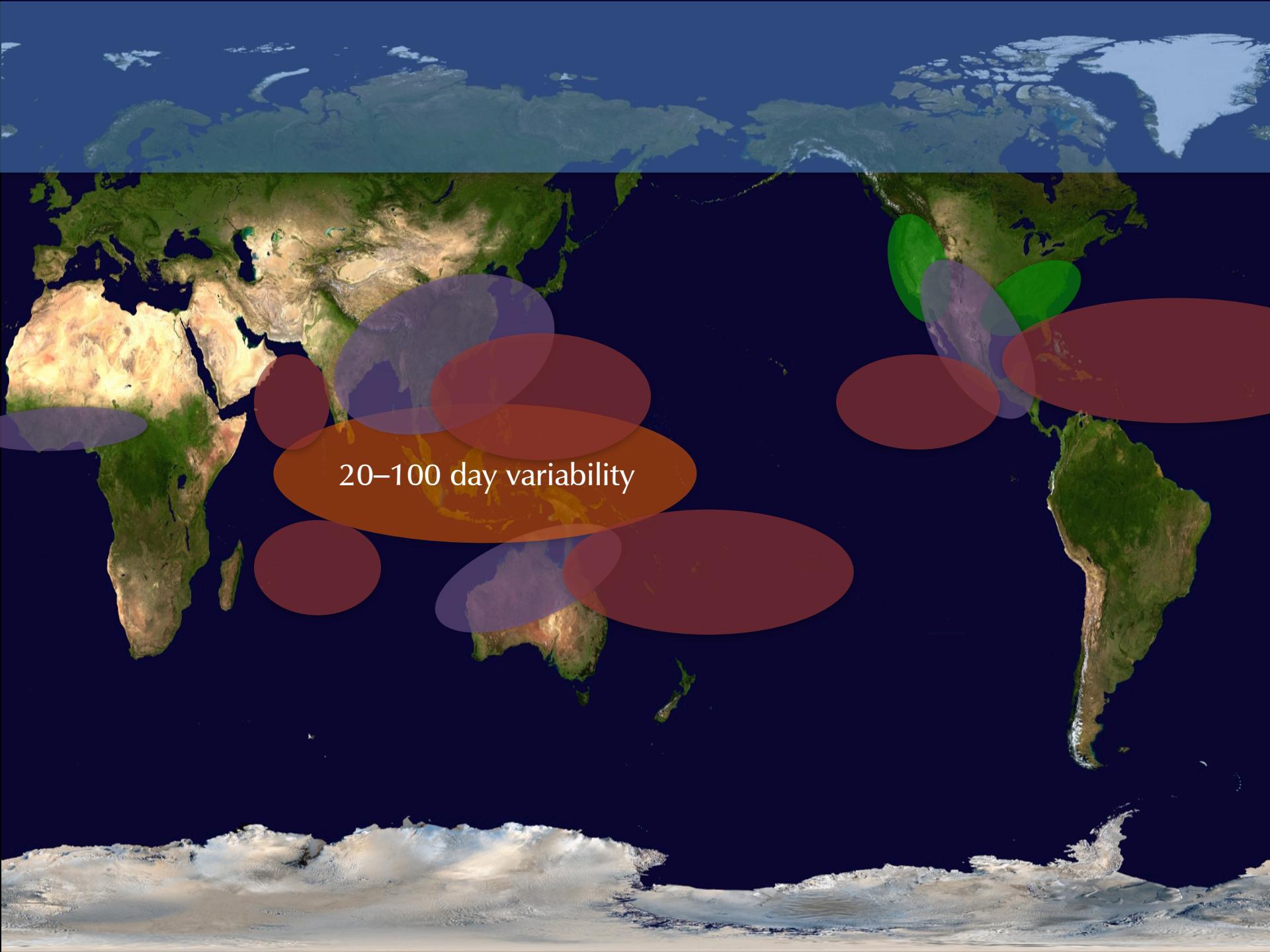
2



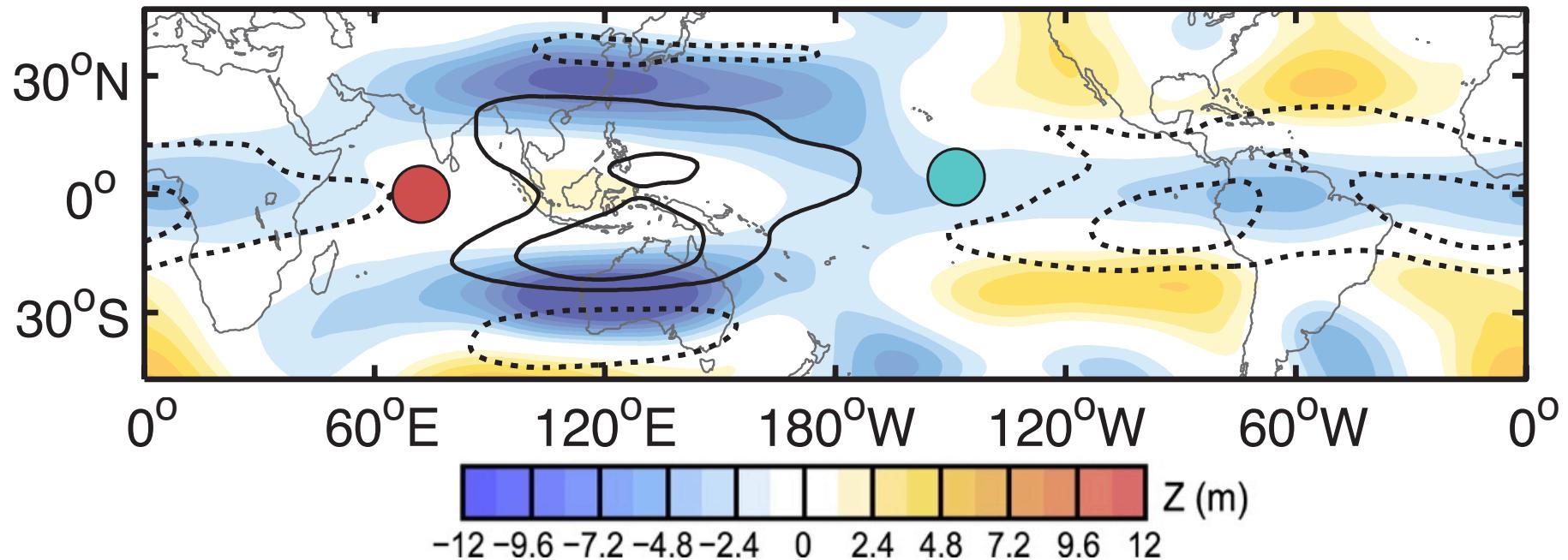


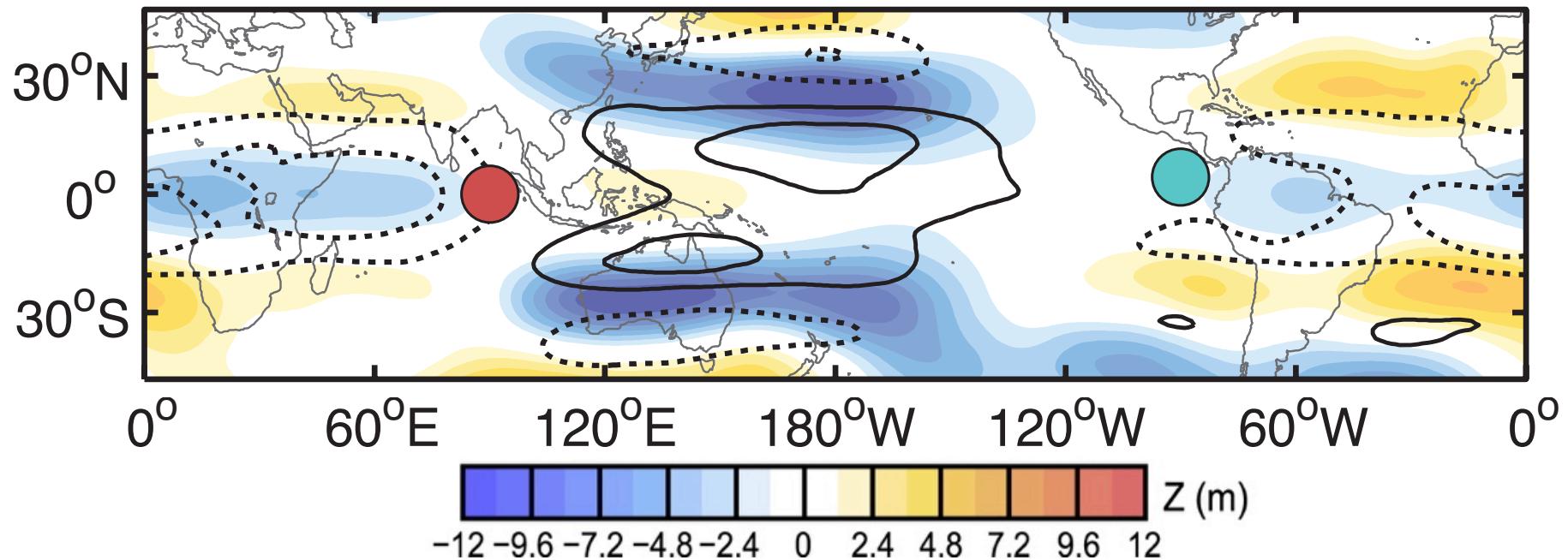


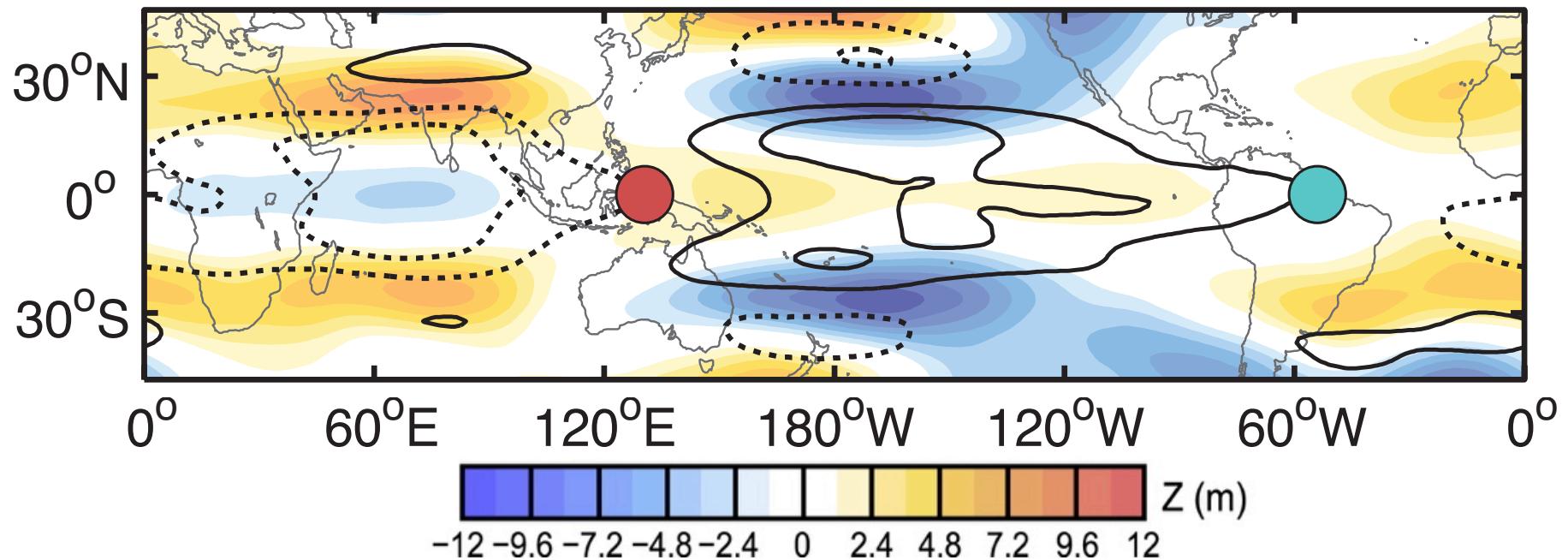


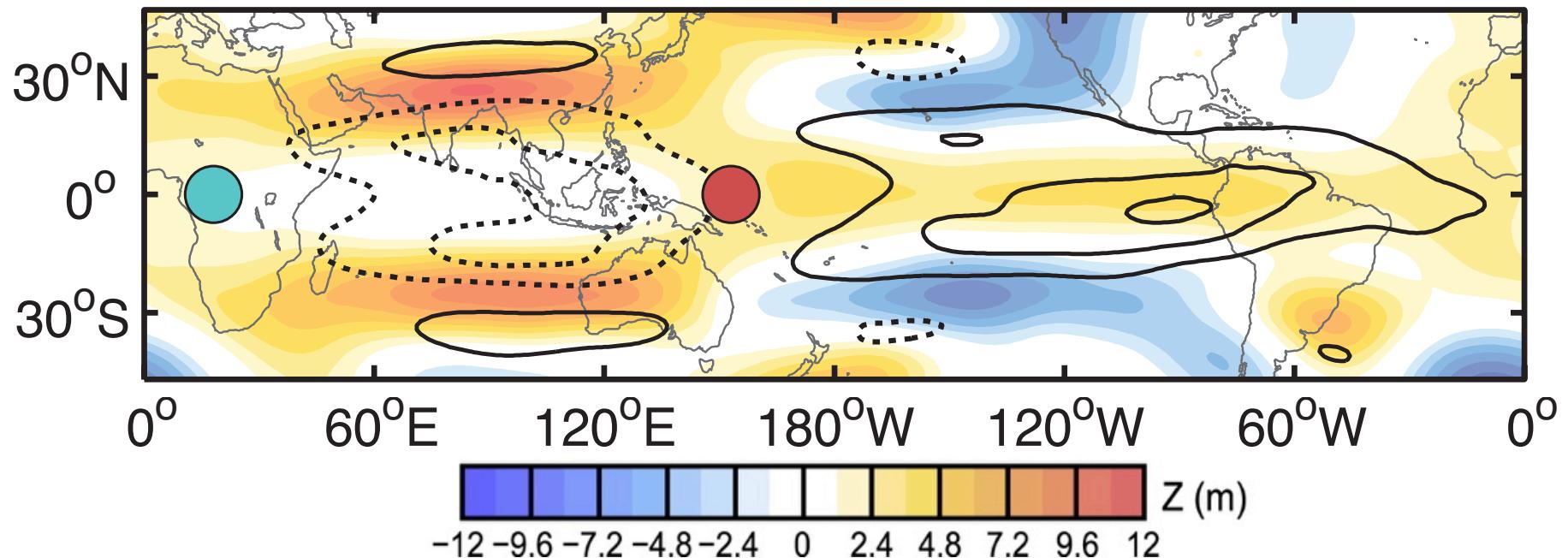


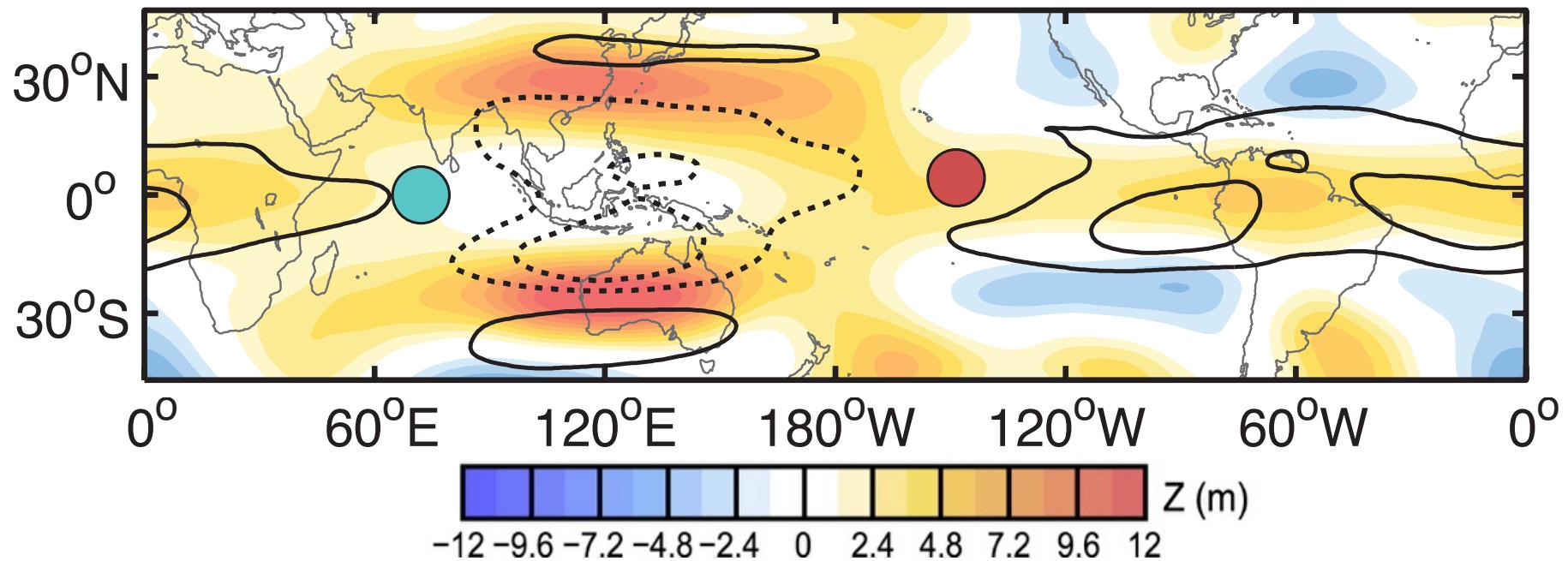
20–100 day variability

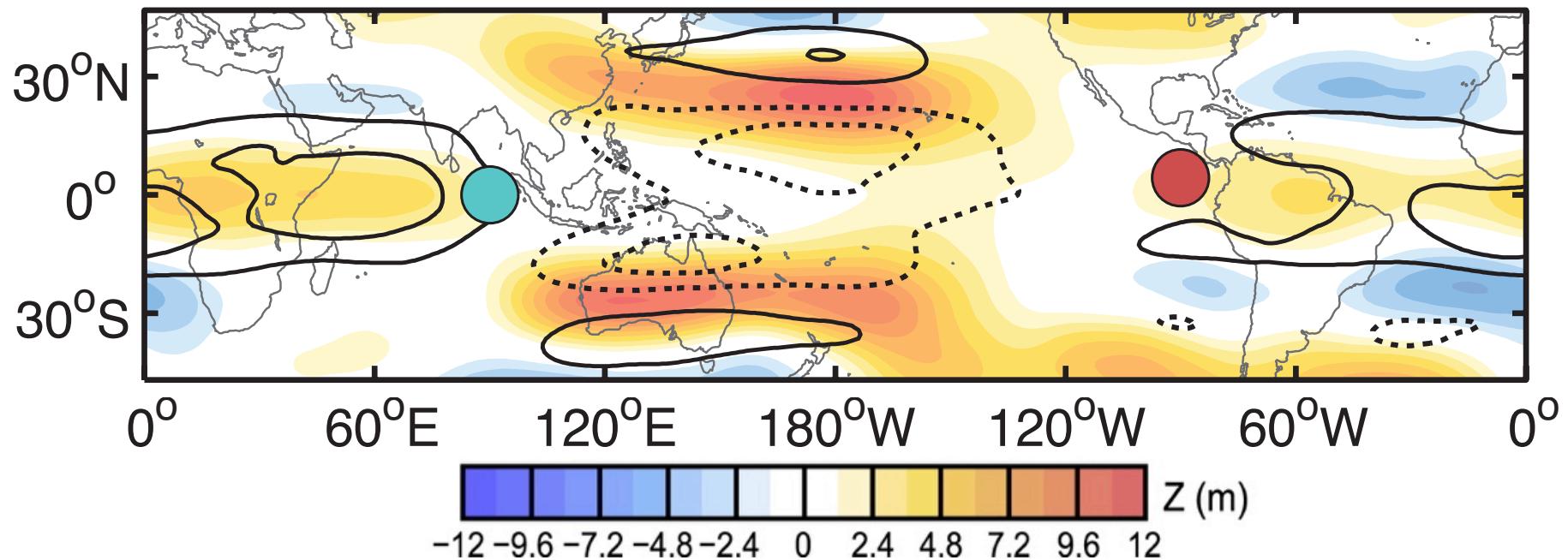


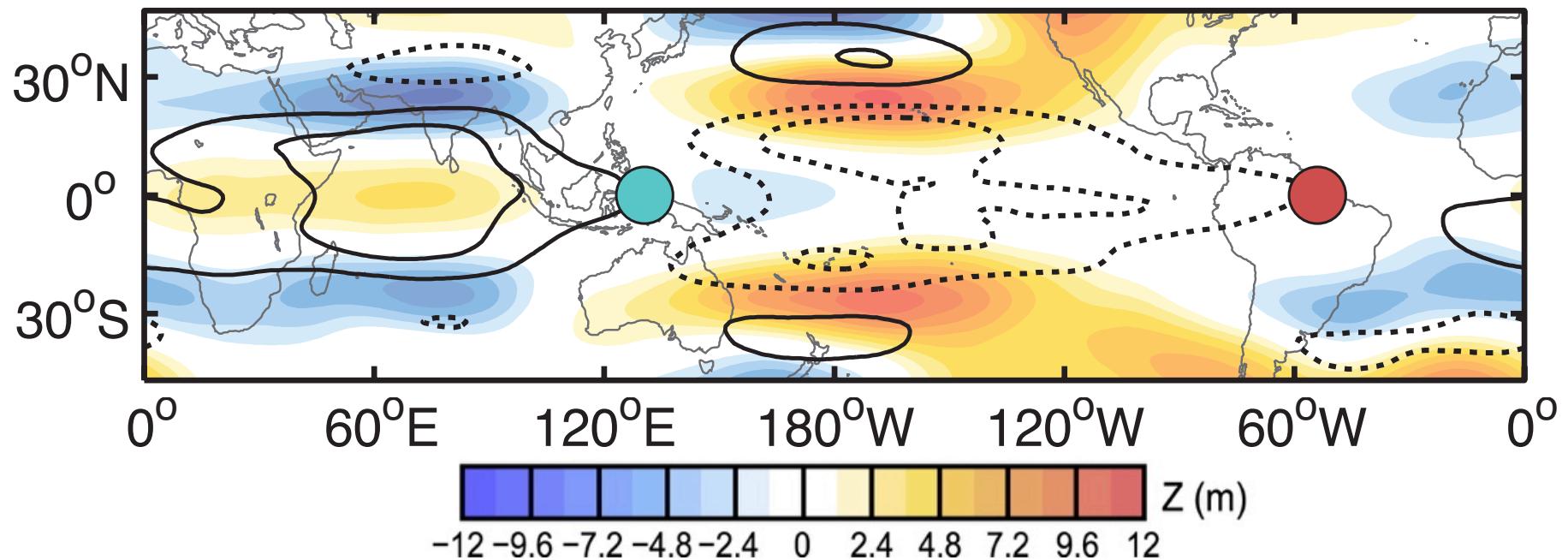


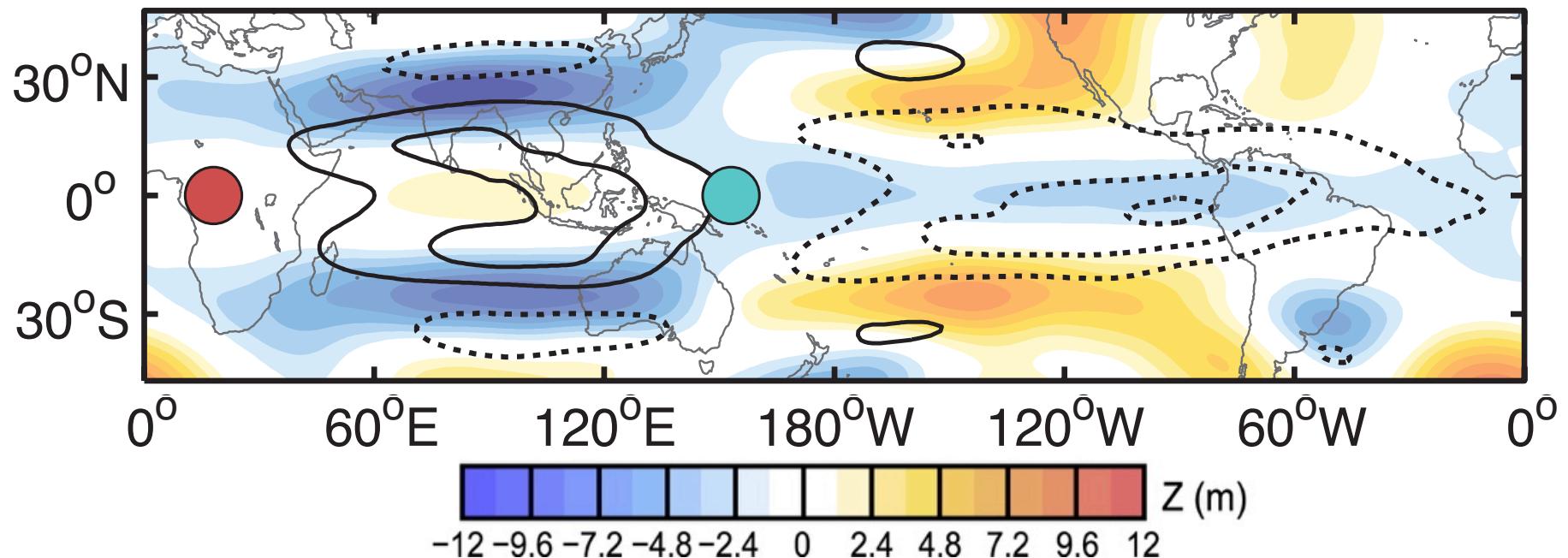


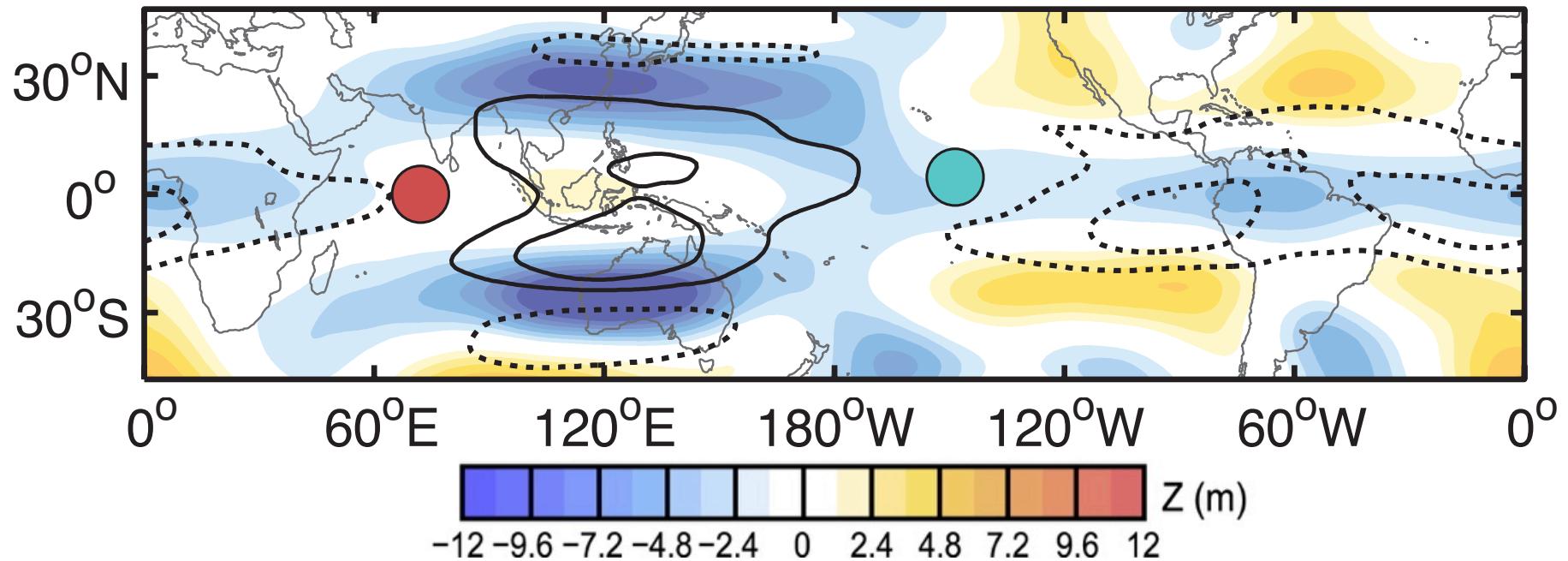


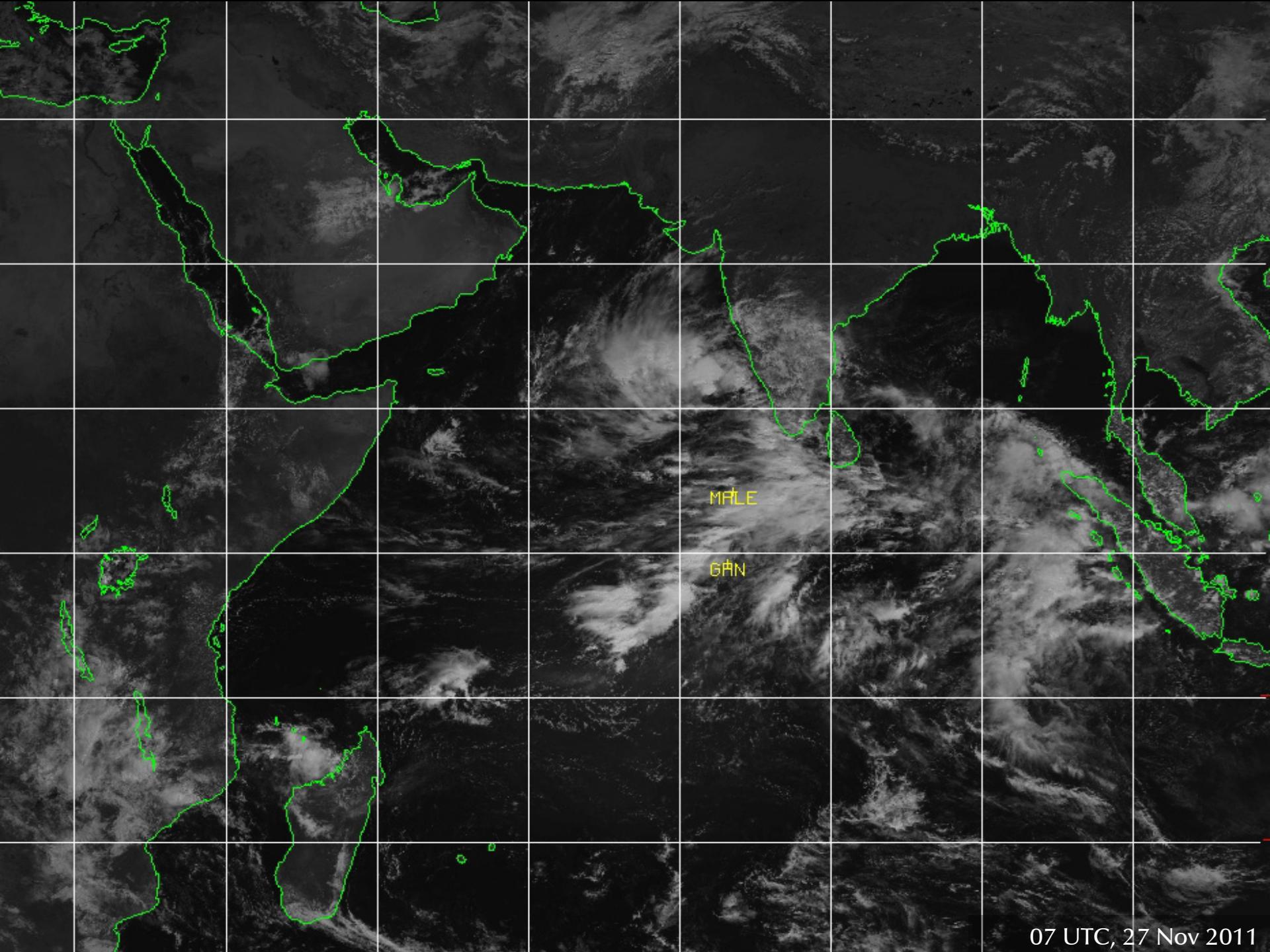


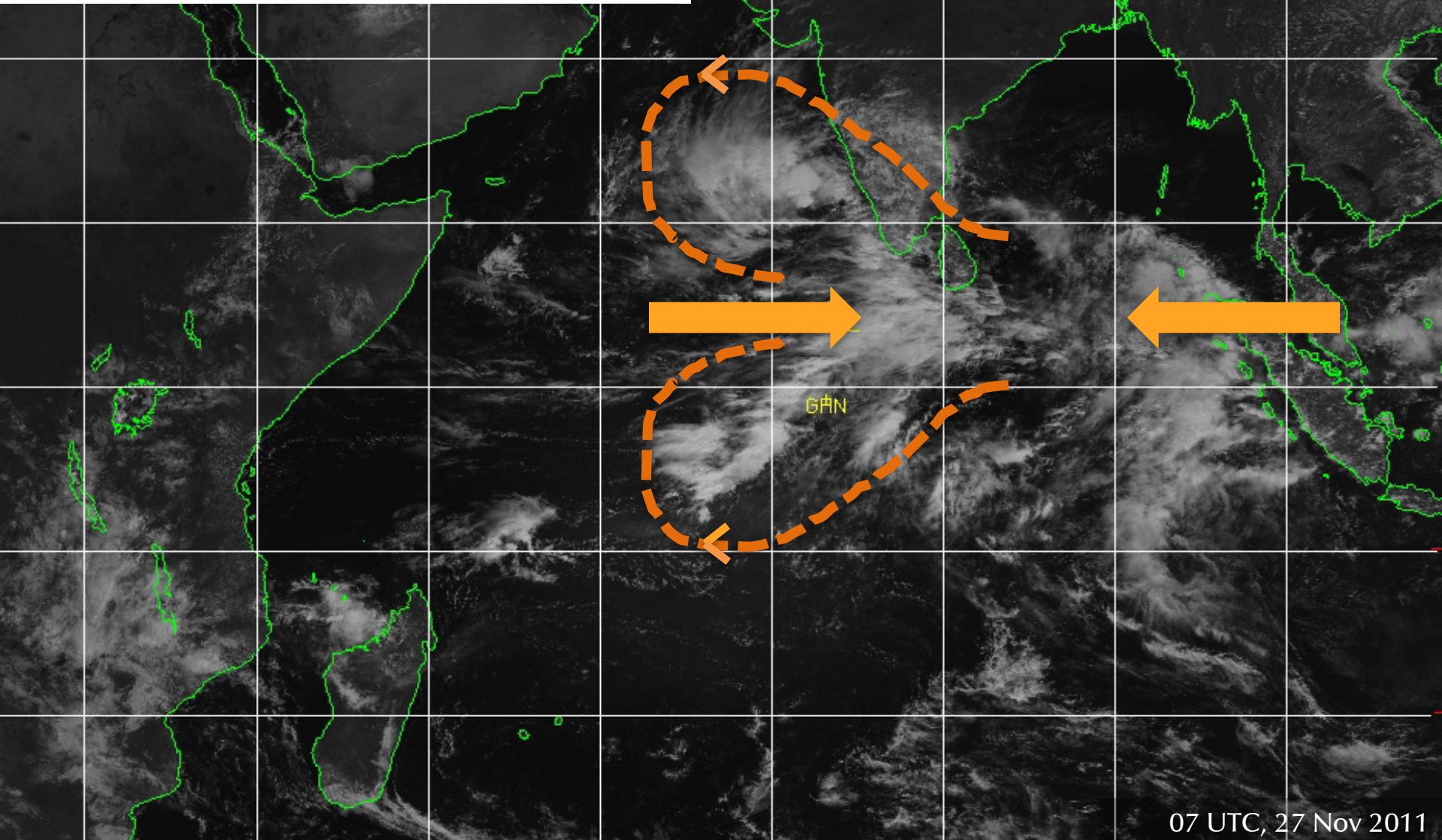
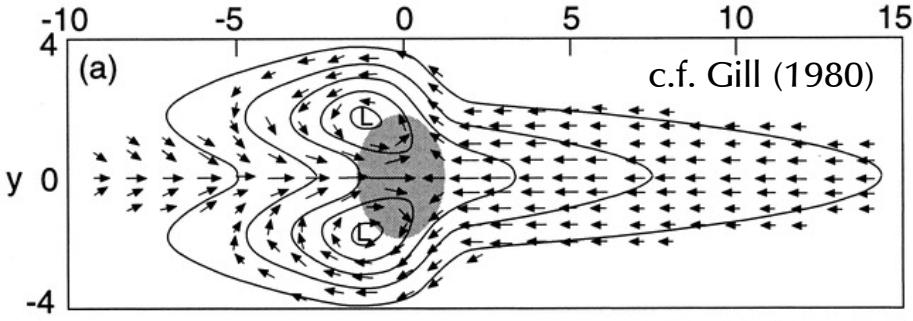


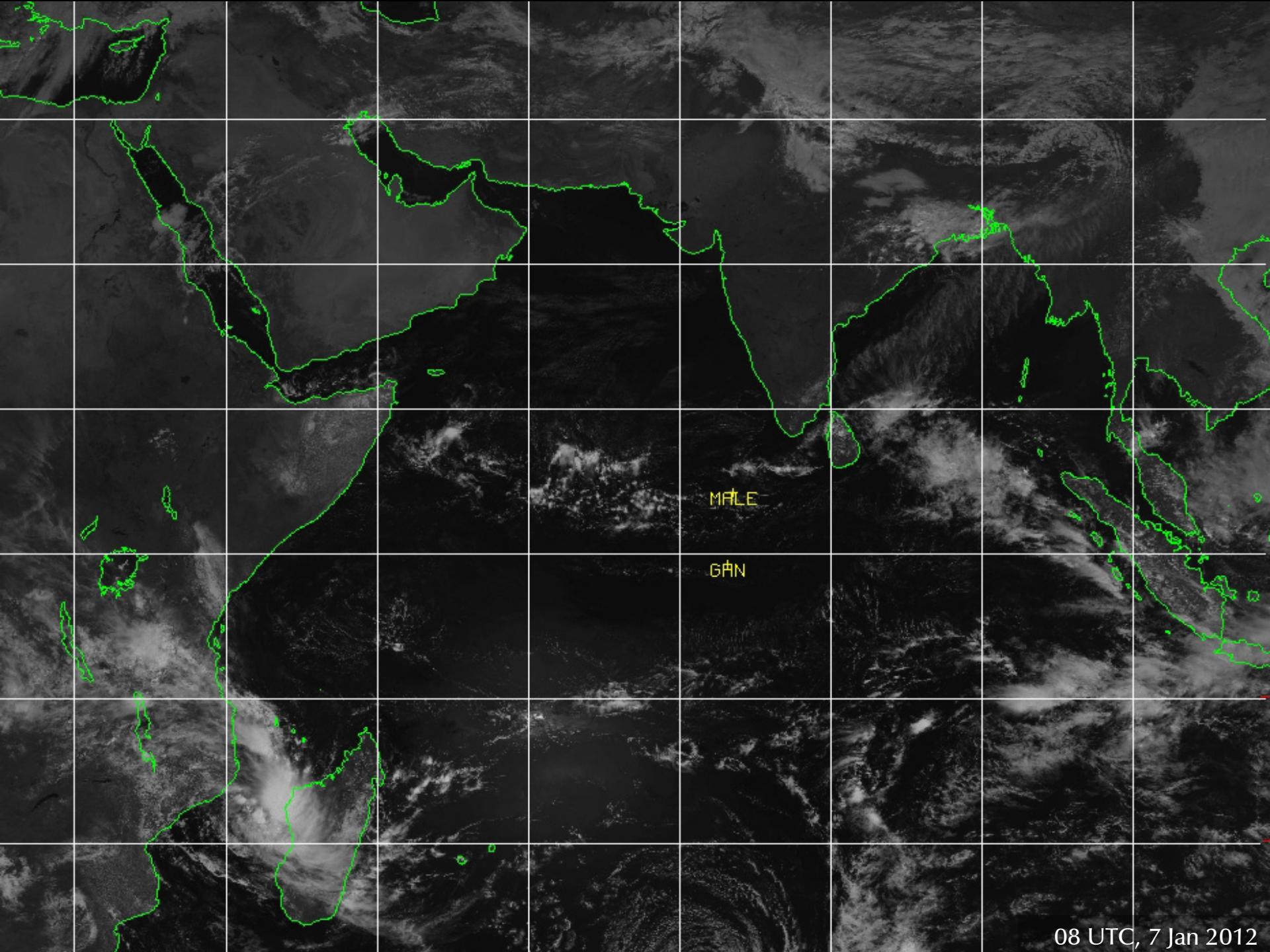




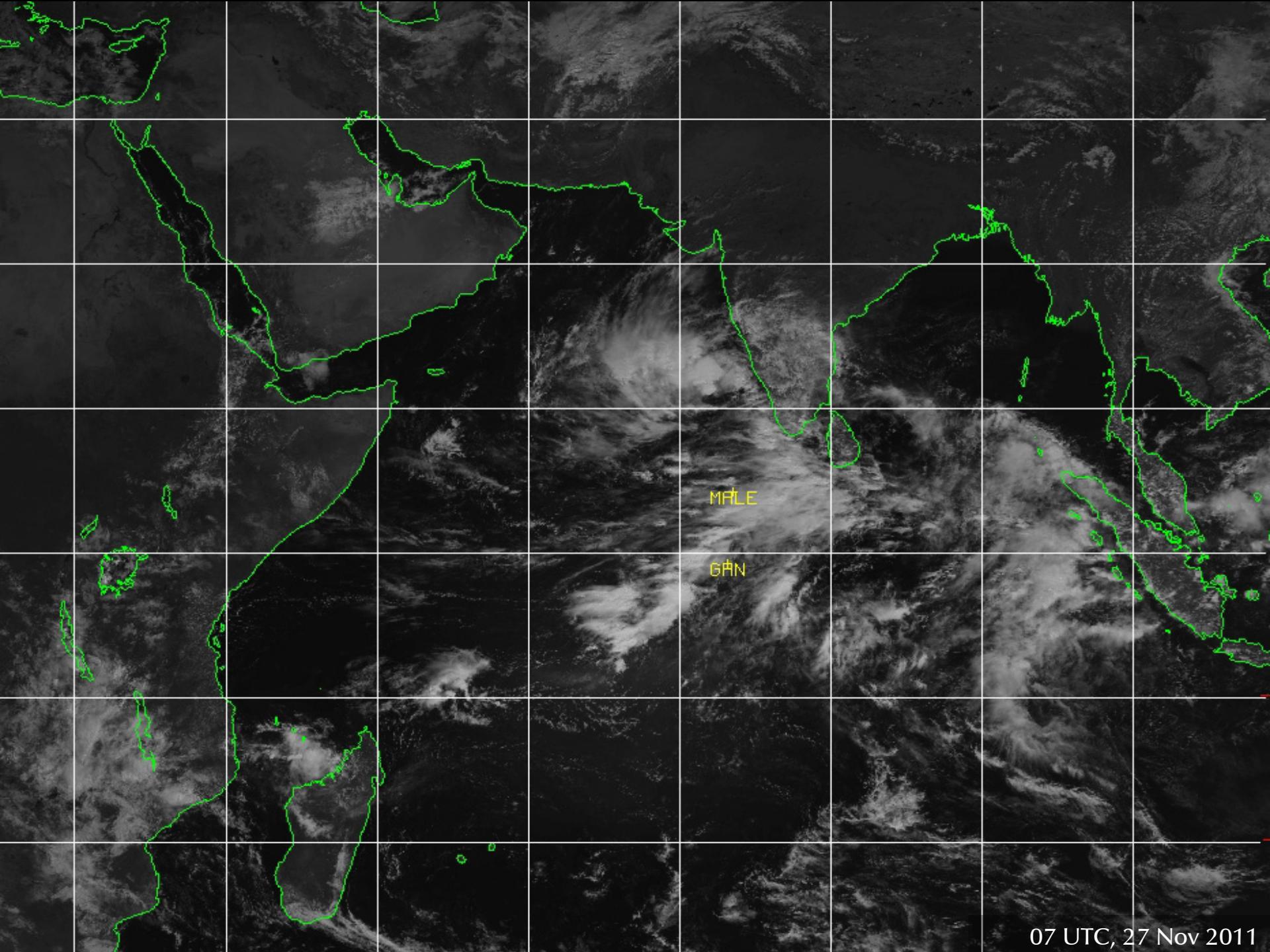




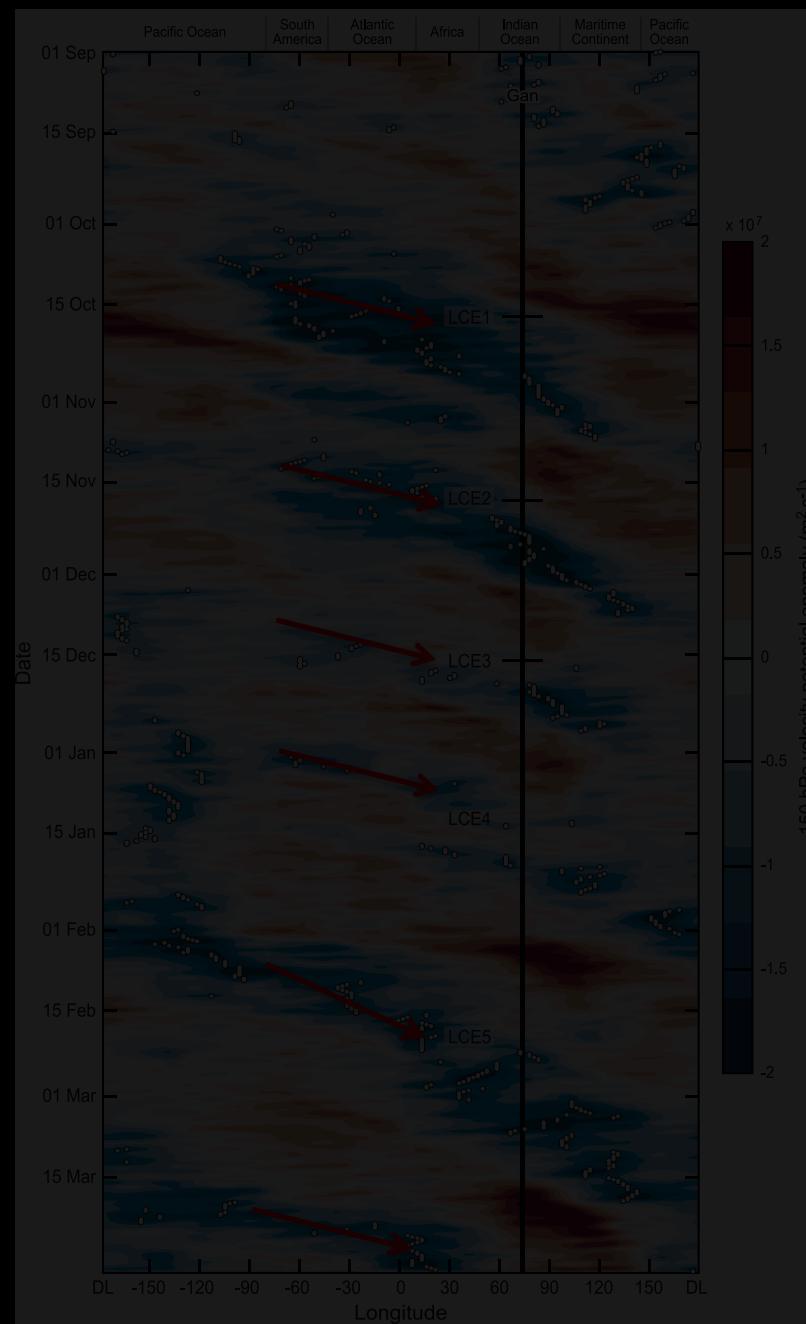




08 UTC, 7 Jan 2012



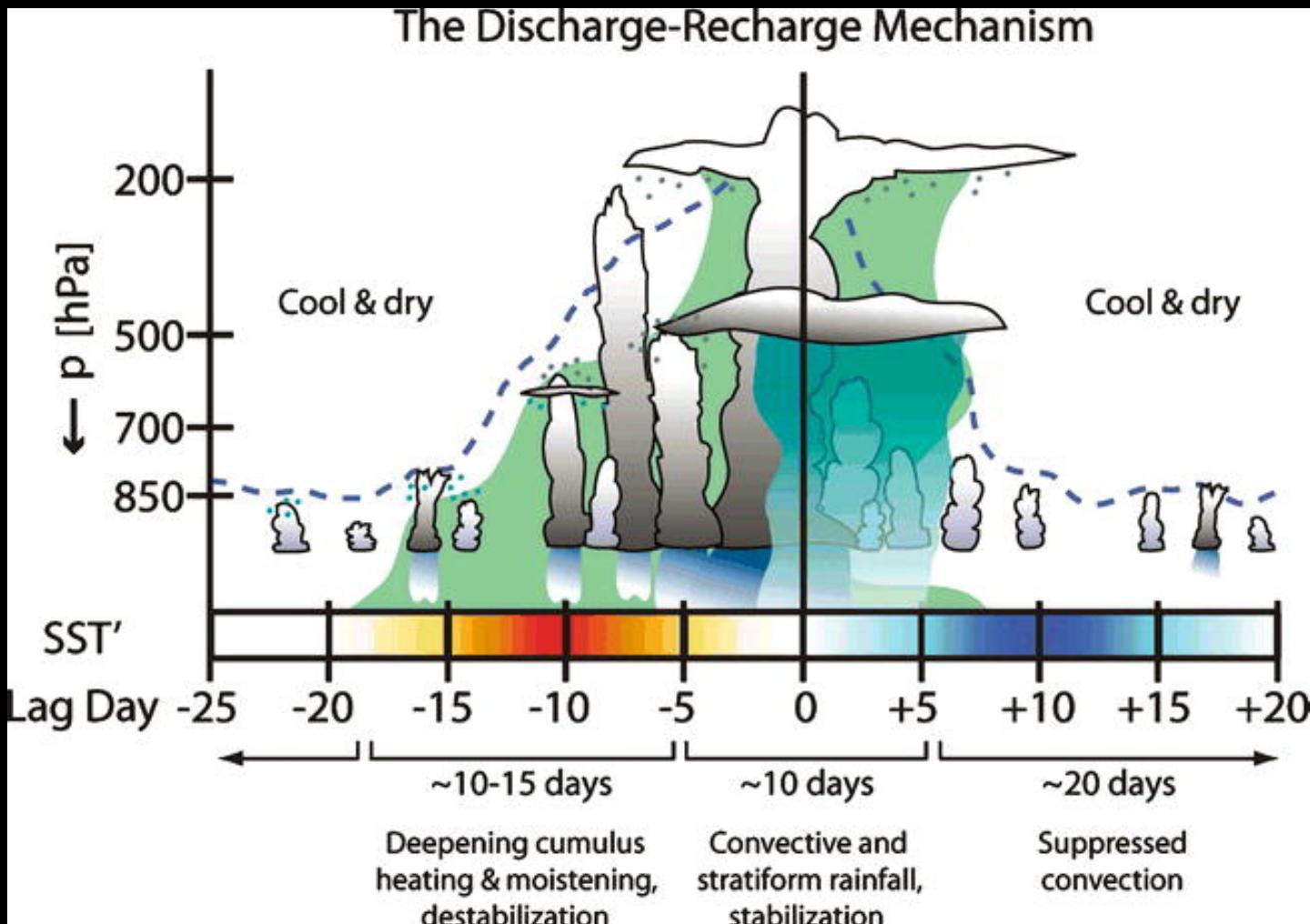
Hypothesis: Convection passively responds to changes in the large-scale environment.



Originally: Knutson
and Weickmann
(1987)

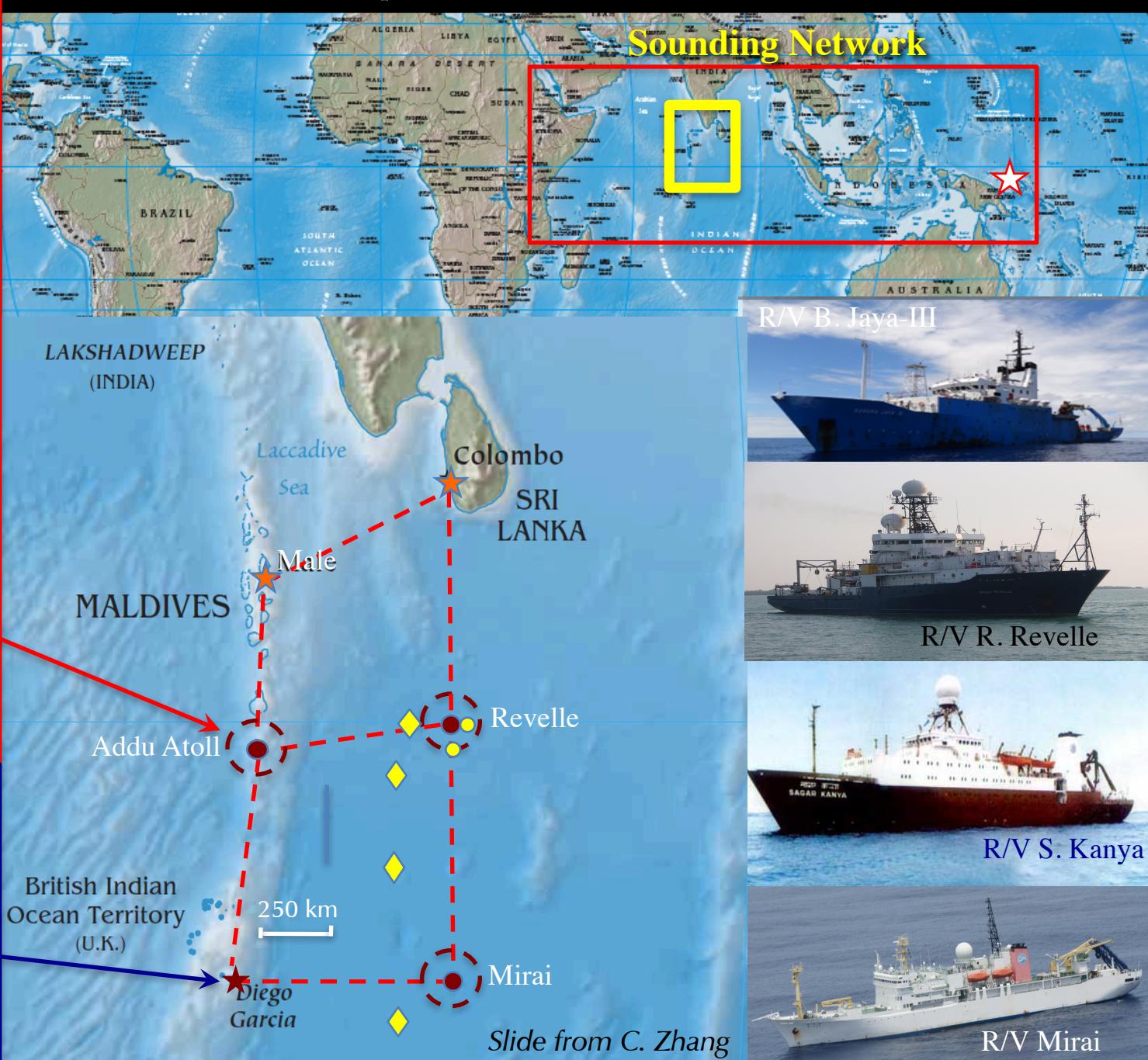
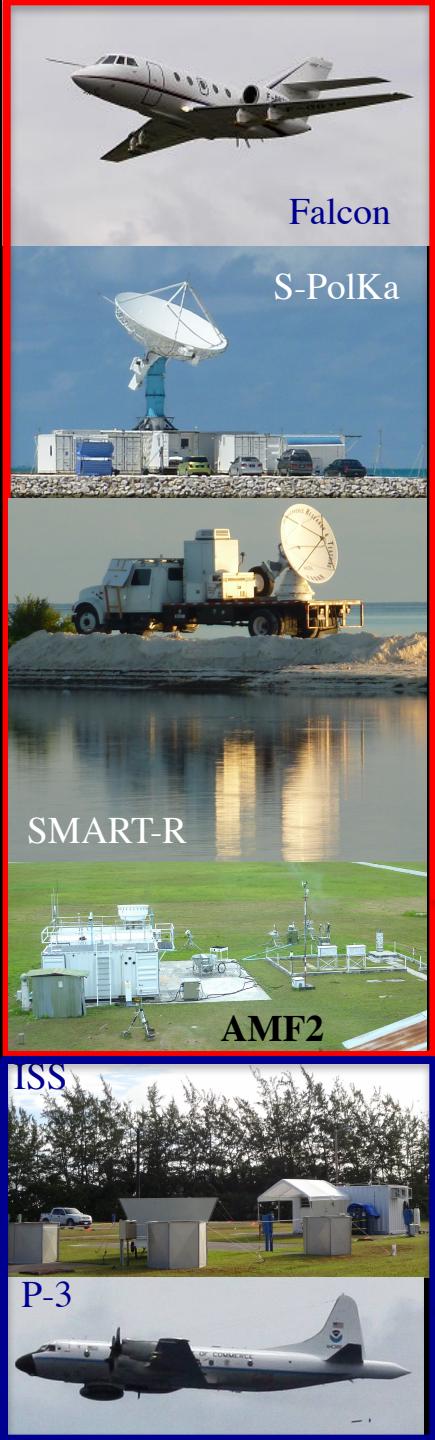
Figure: Powell and
Houze (2015b)

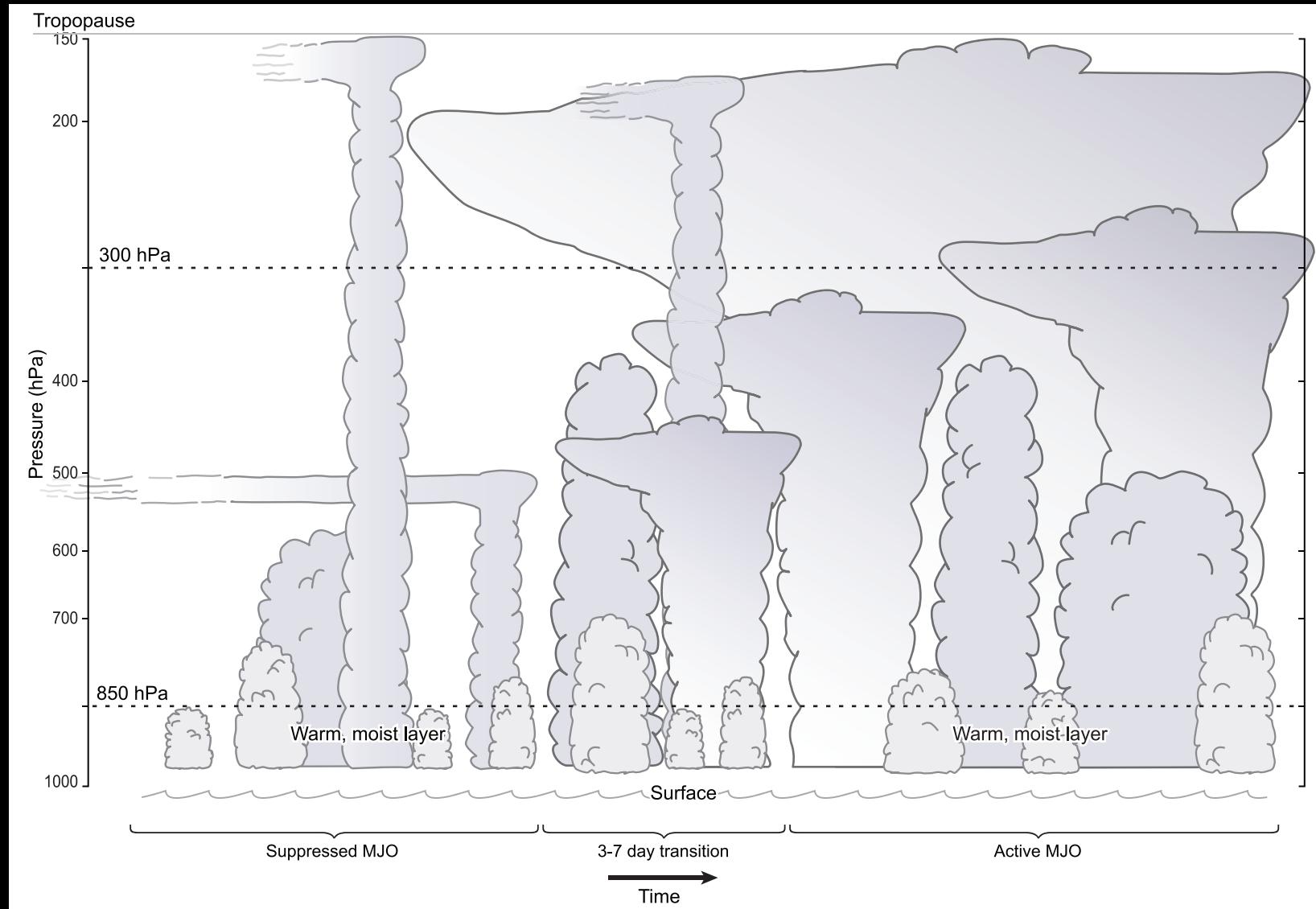
Hypothesis: Clouds are actively involved in “preconditioning” environment for MJO.

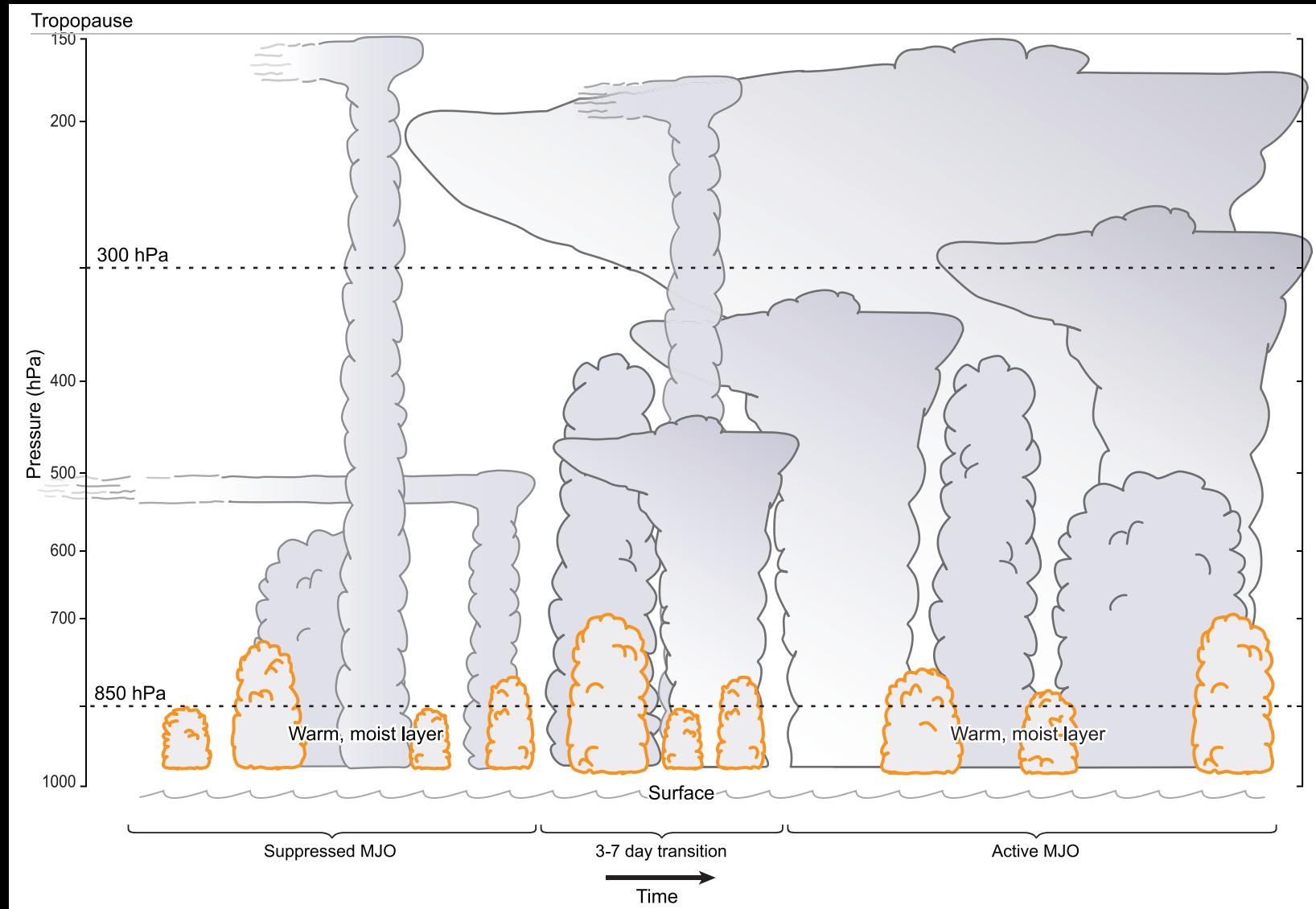


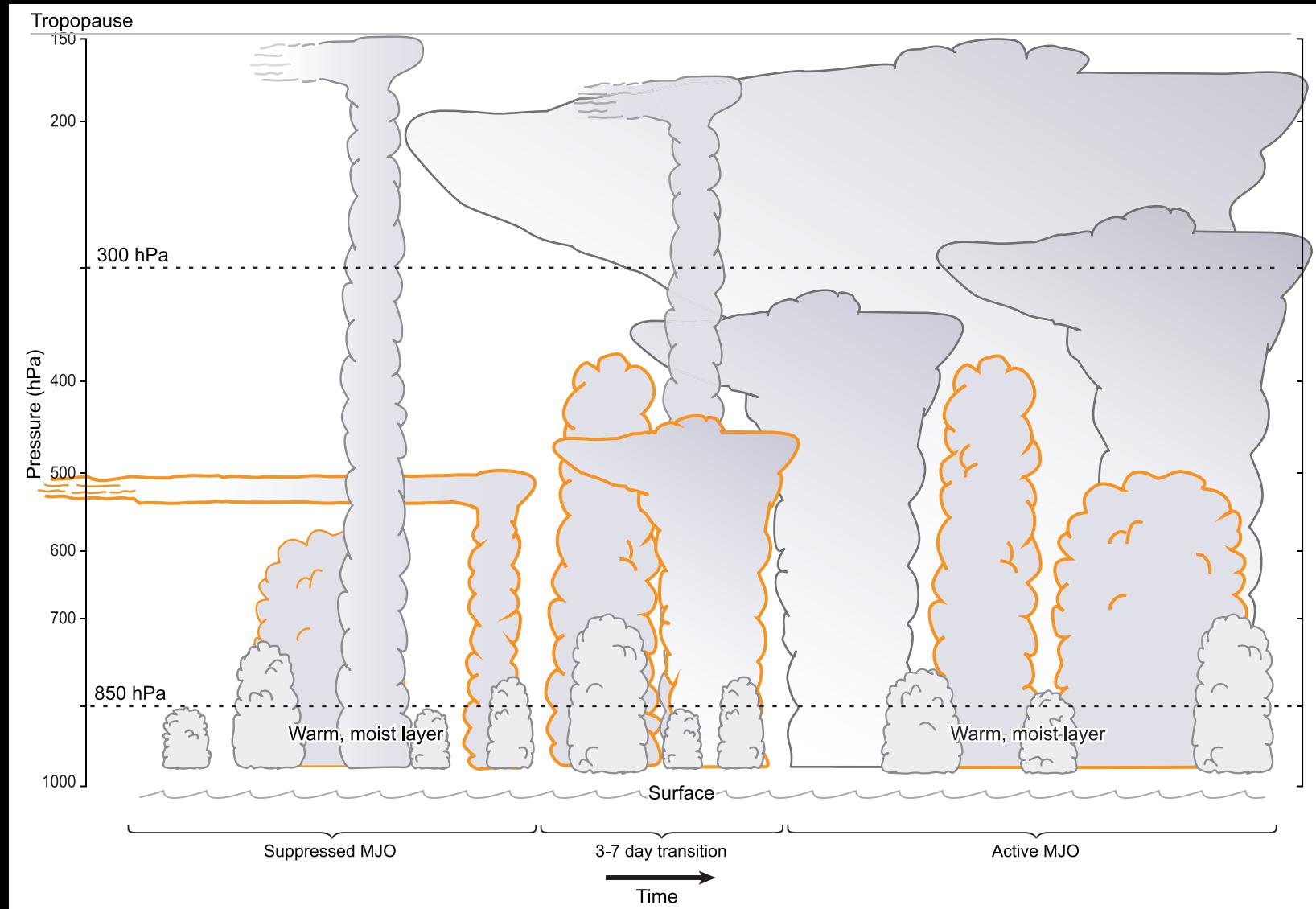
Benedict and Randall (2007), following Bladé and Hartmann (1993) and Kemball-Cook and Weare (2001)

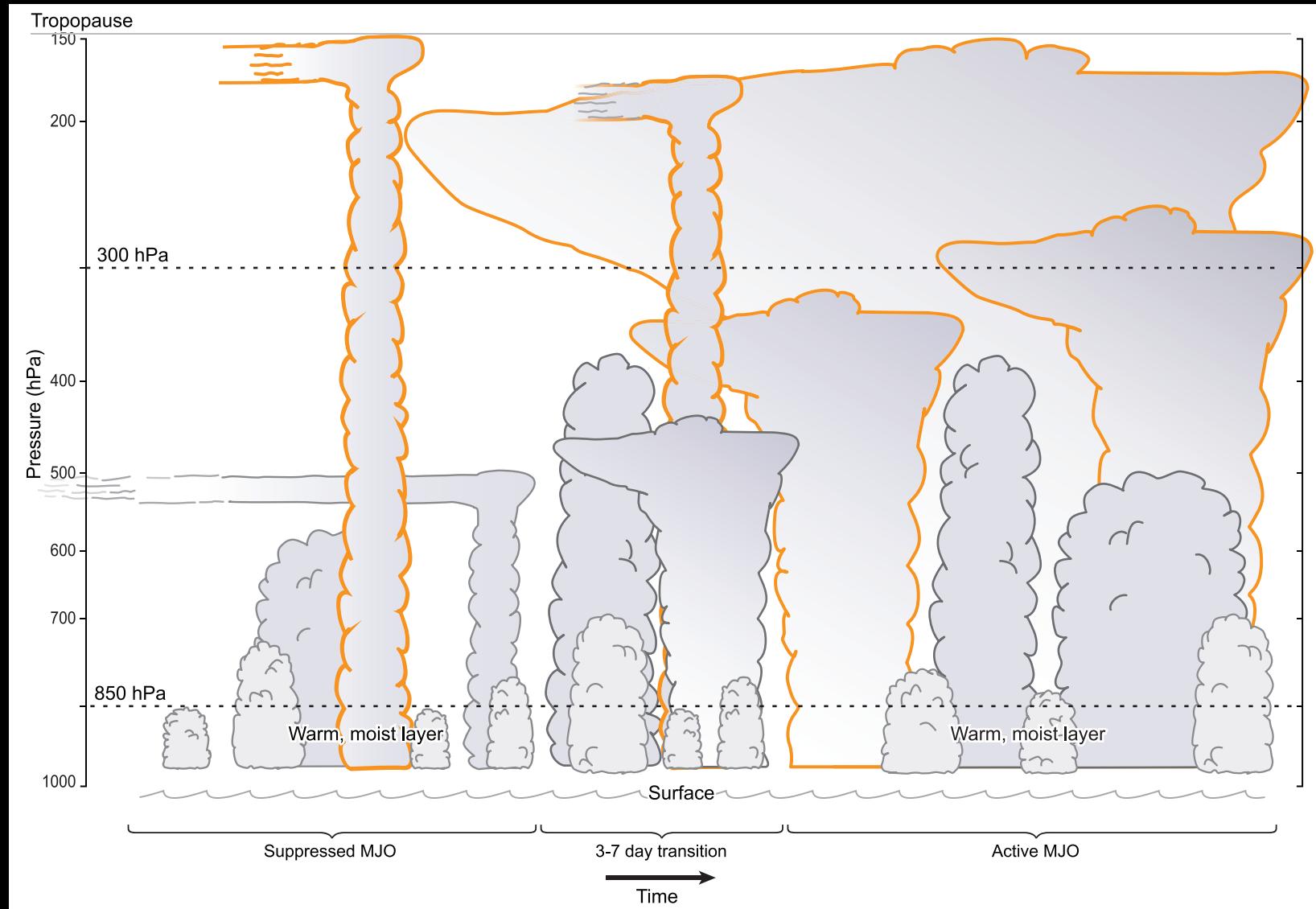
DYNAMO Field Experiment (October 2011 – March 2012)

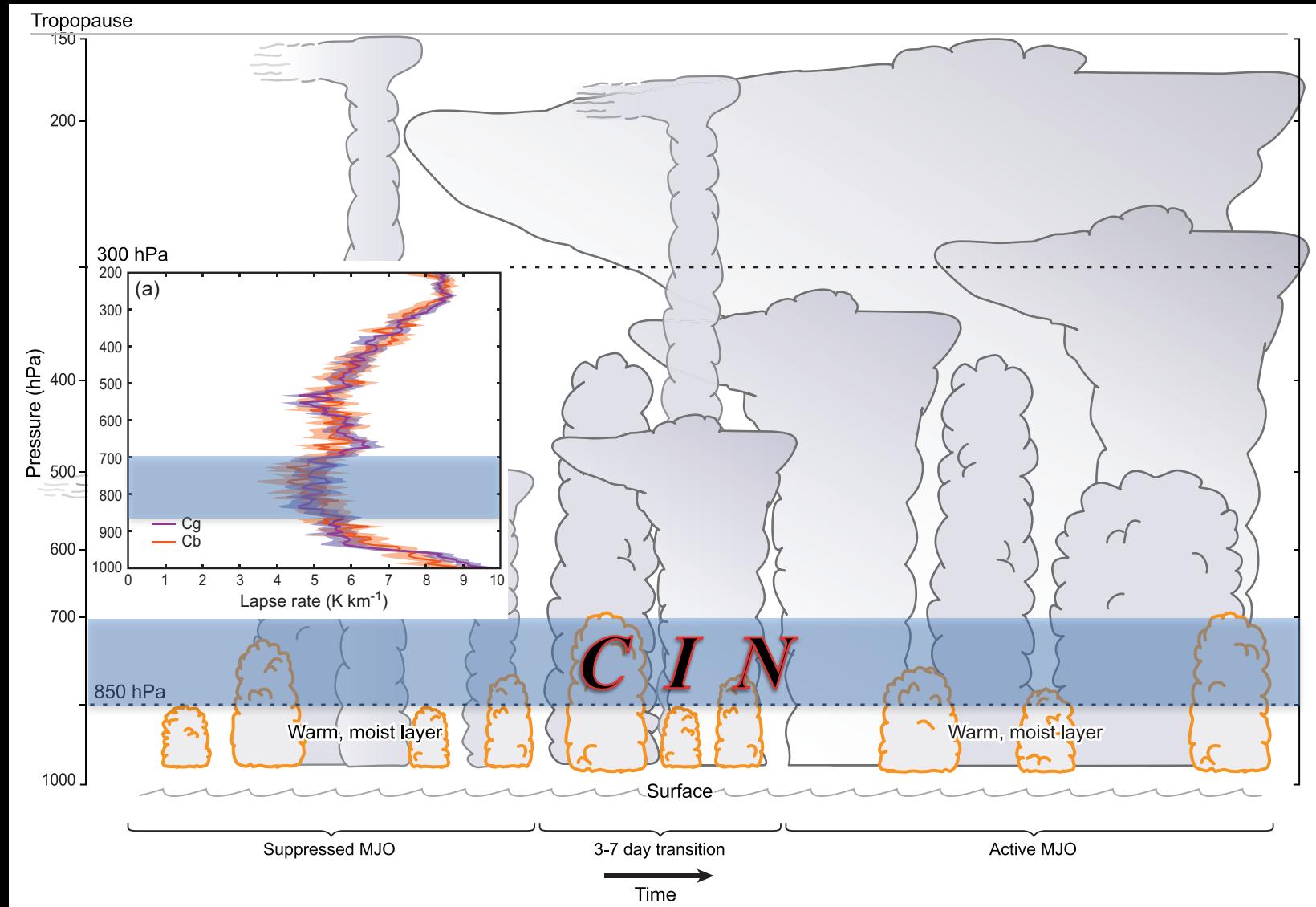


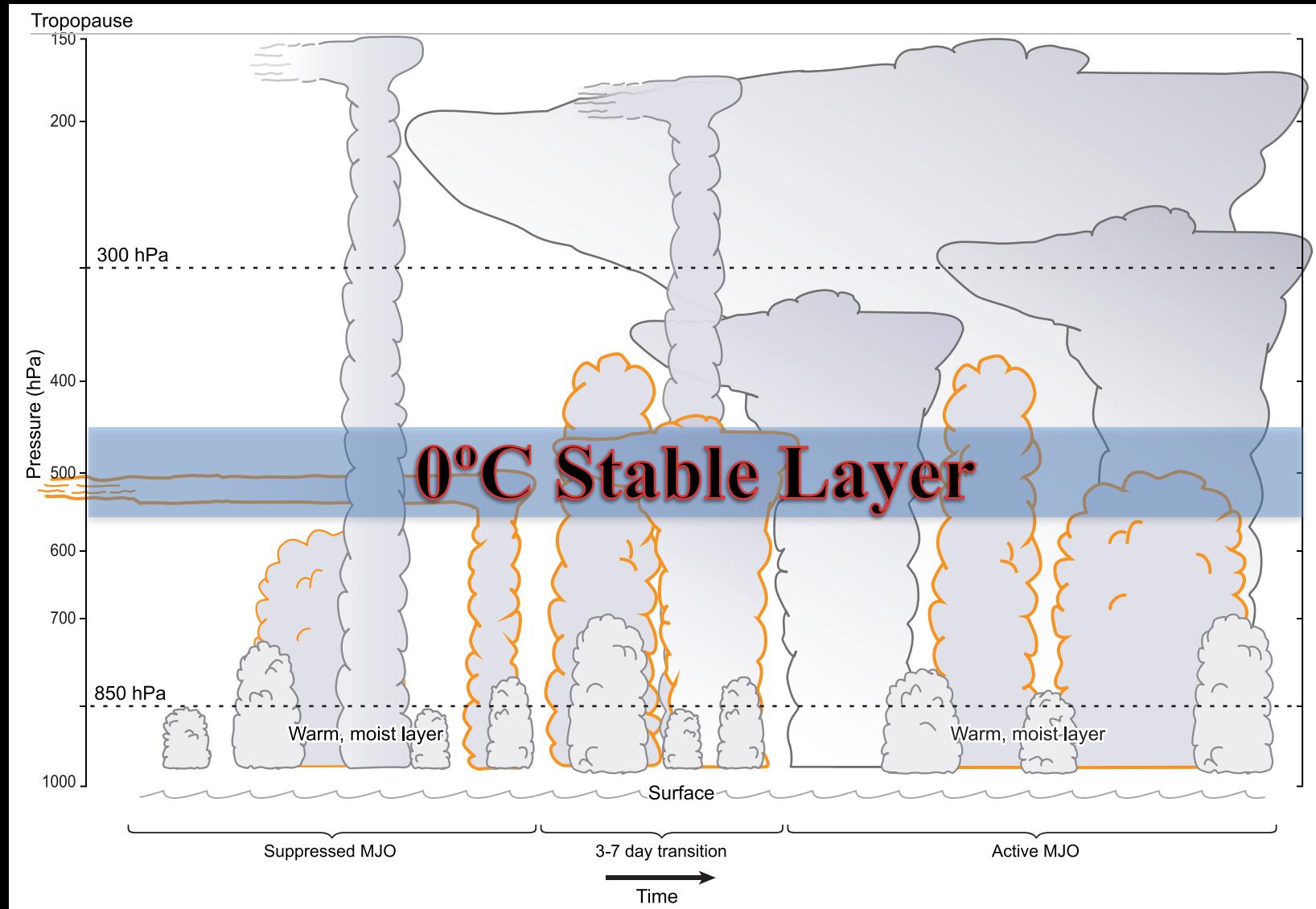


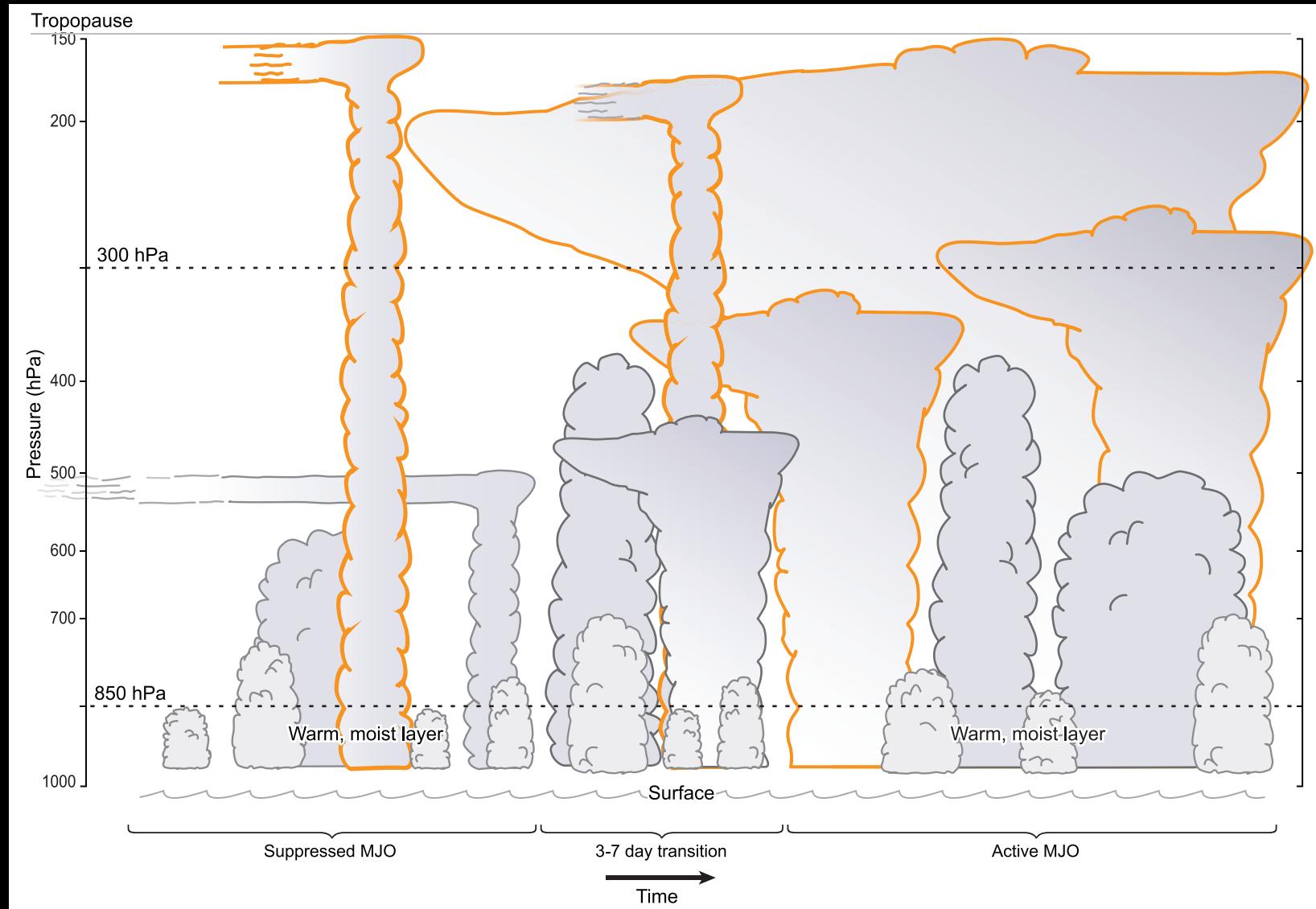


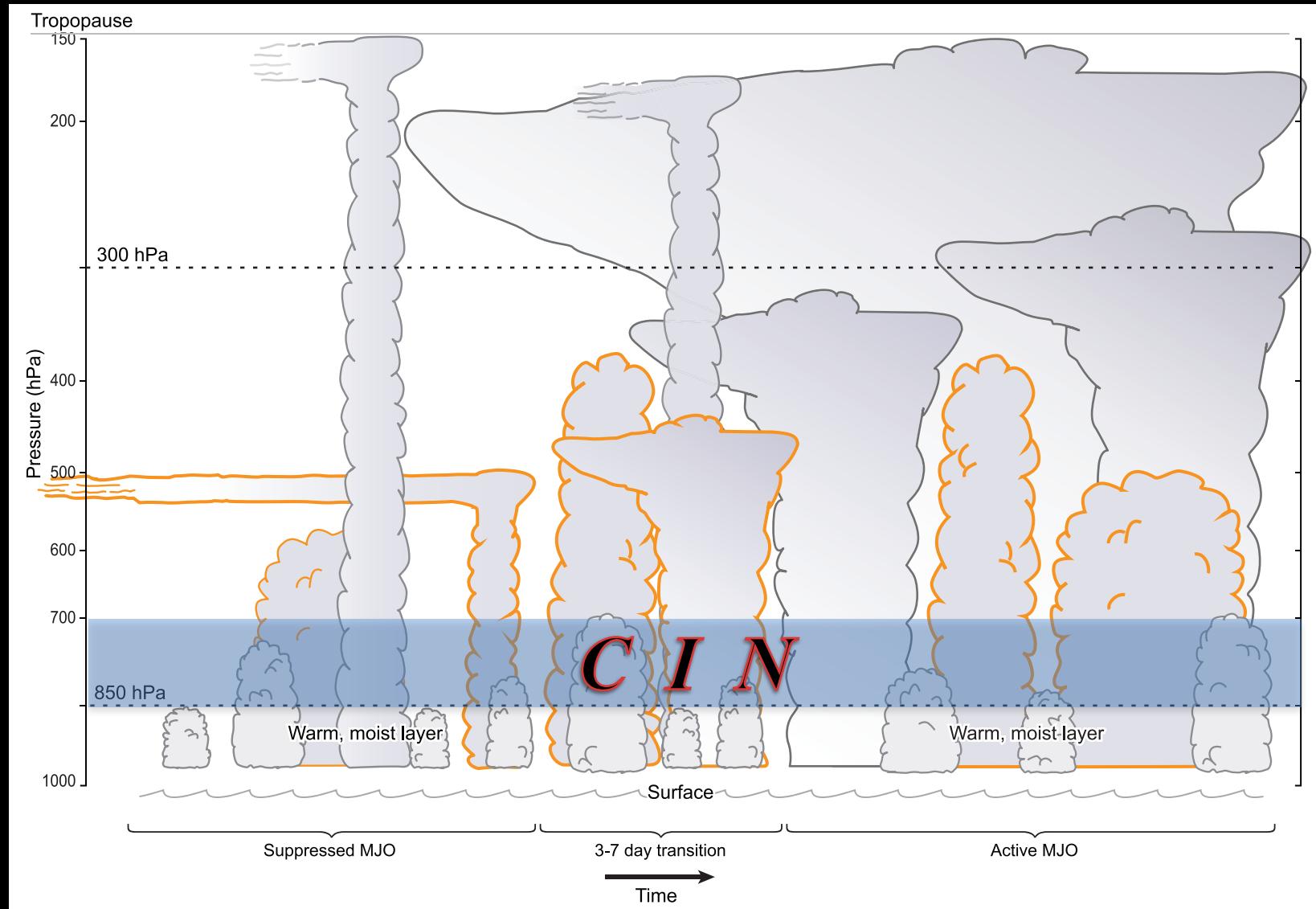










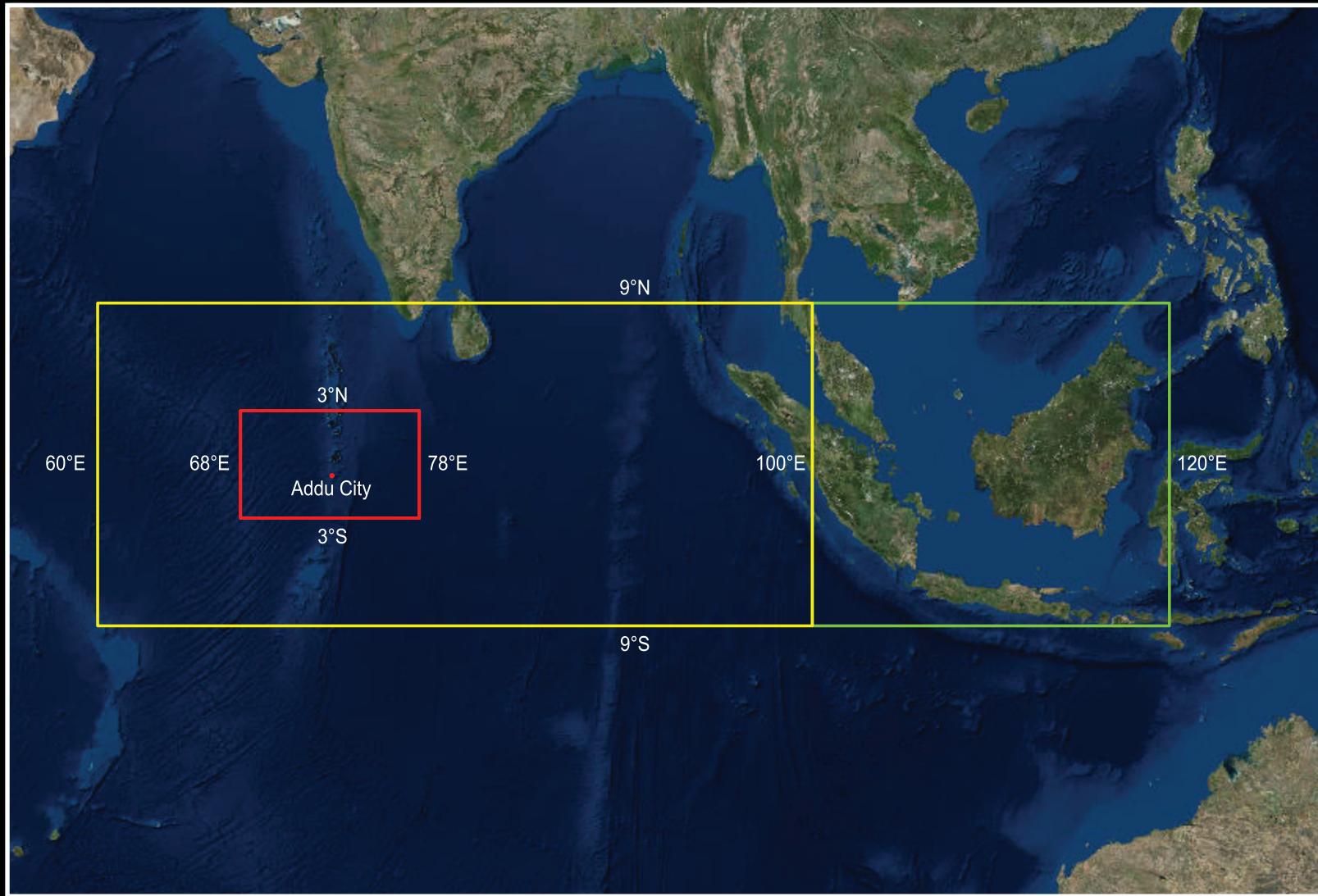


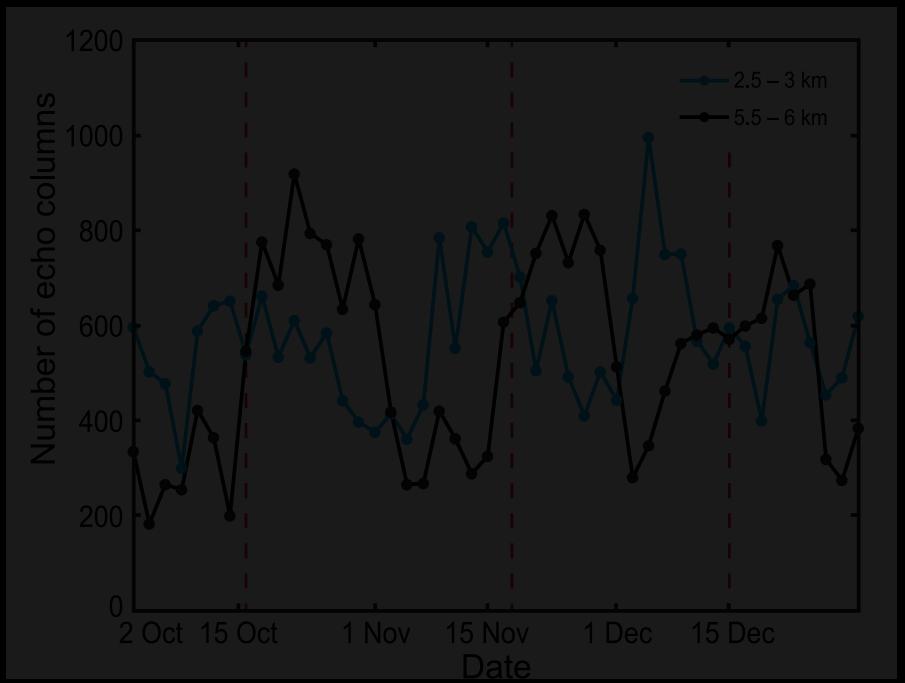
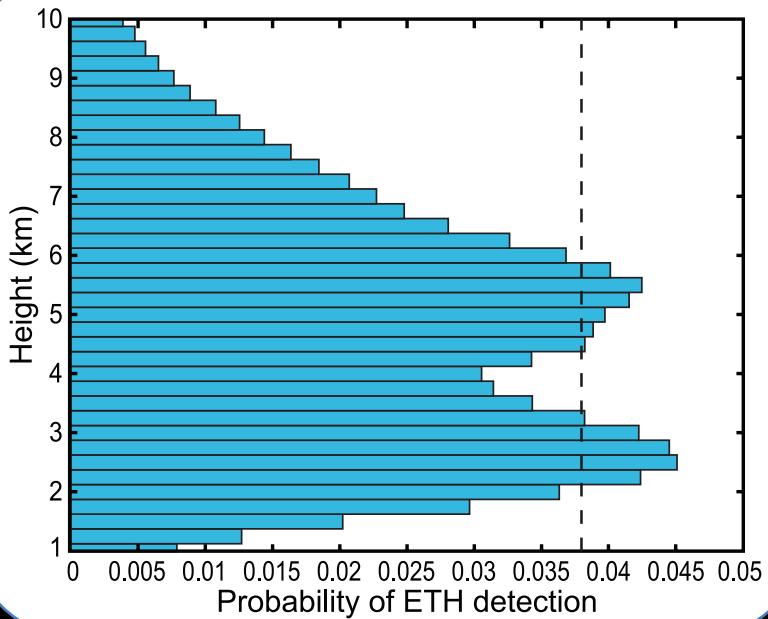
Objectives

1. Do clouds moisten environment or does something else, allowing for cloud development?
2. Role of global circulation anomalies in cloud growth

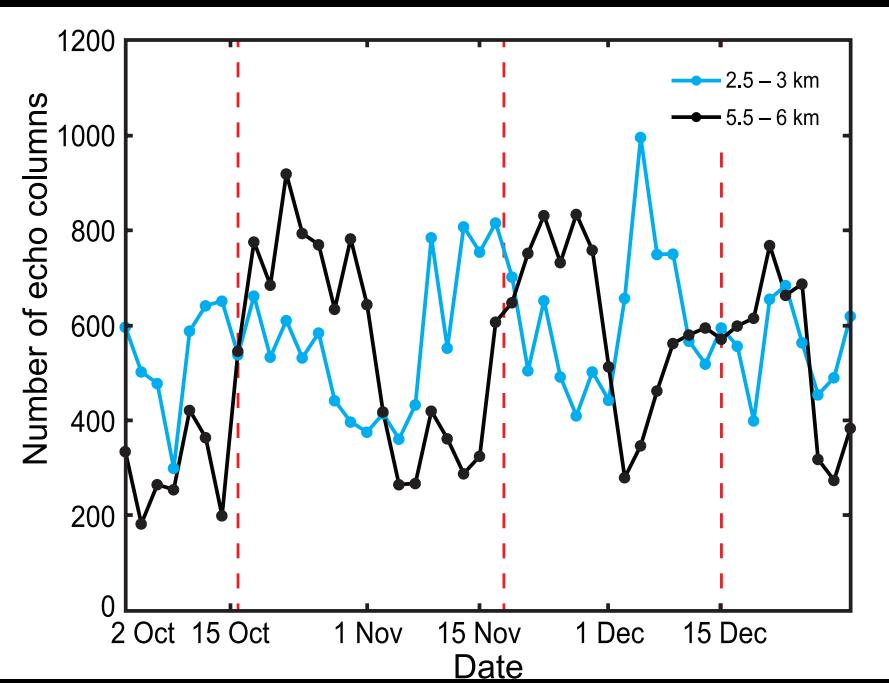
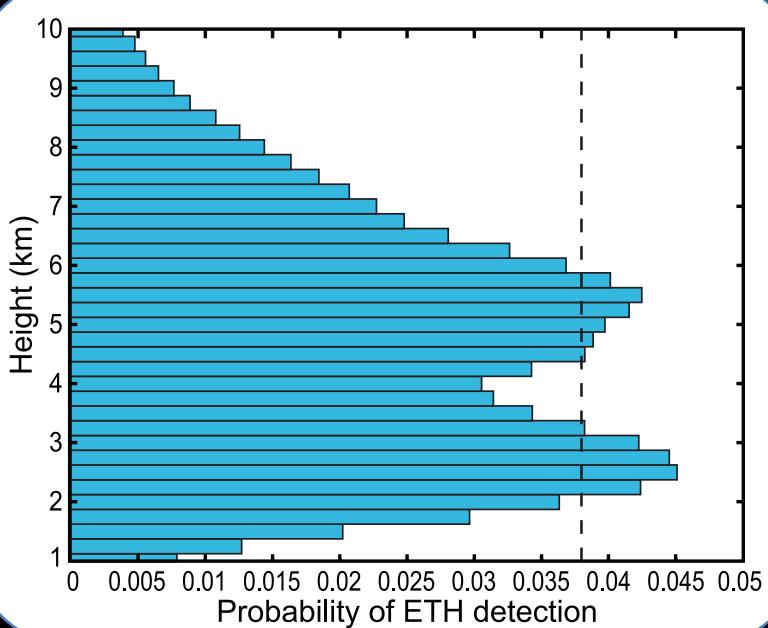
Moistening by Cumulonimbi

Do moderately deep clouds moisten the troposphere during transition periods, or does moistening permit observed cloud deepening?

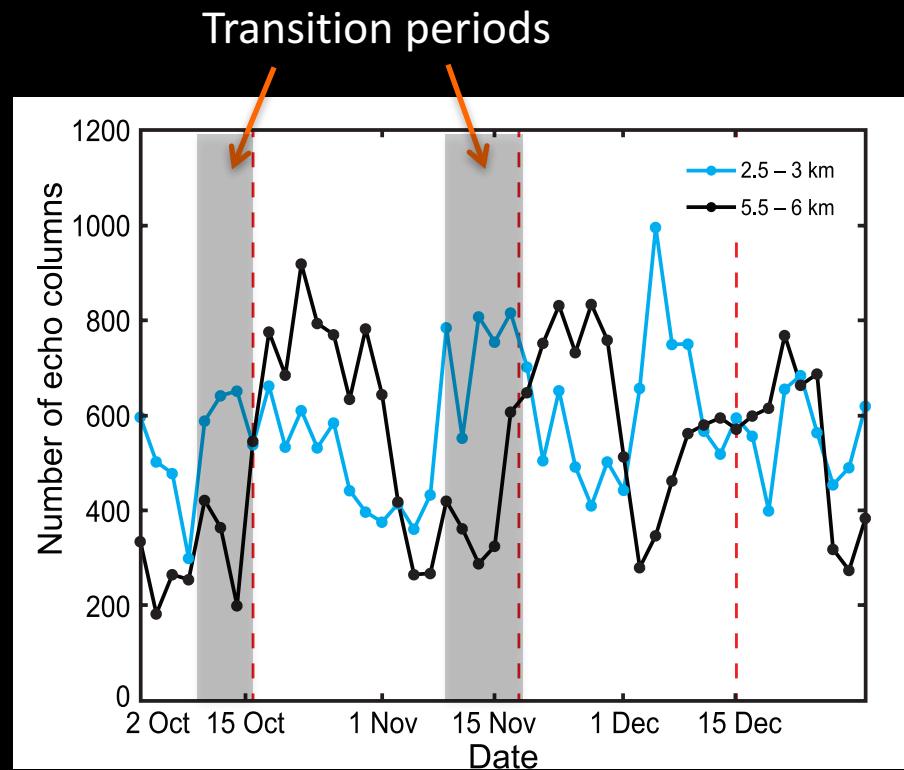
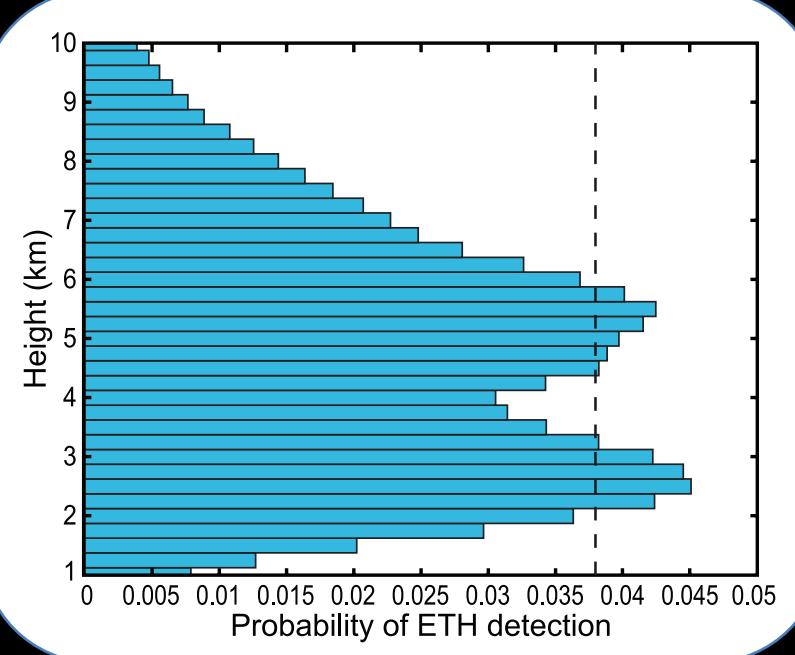




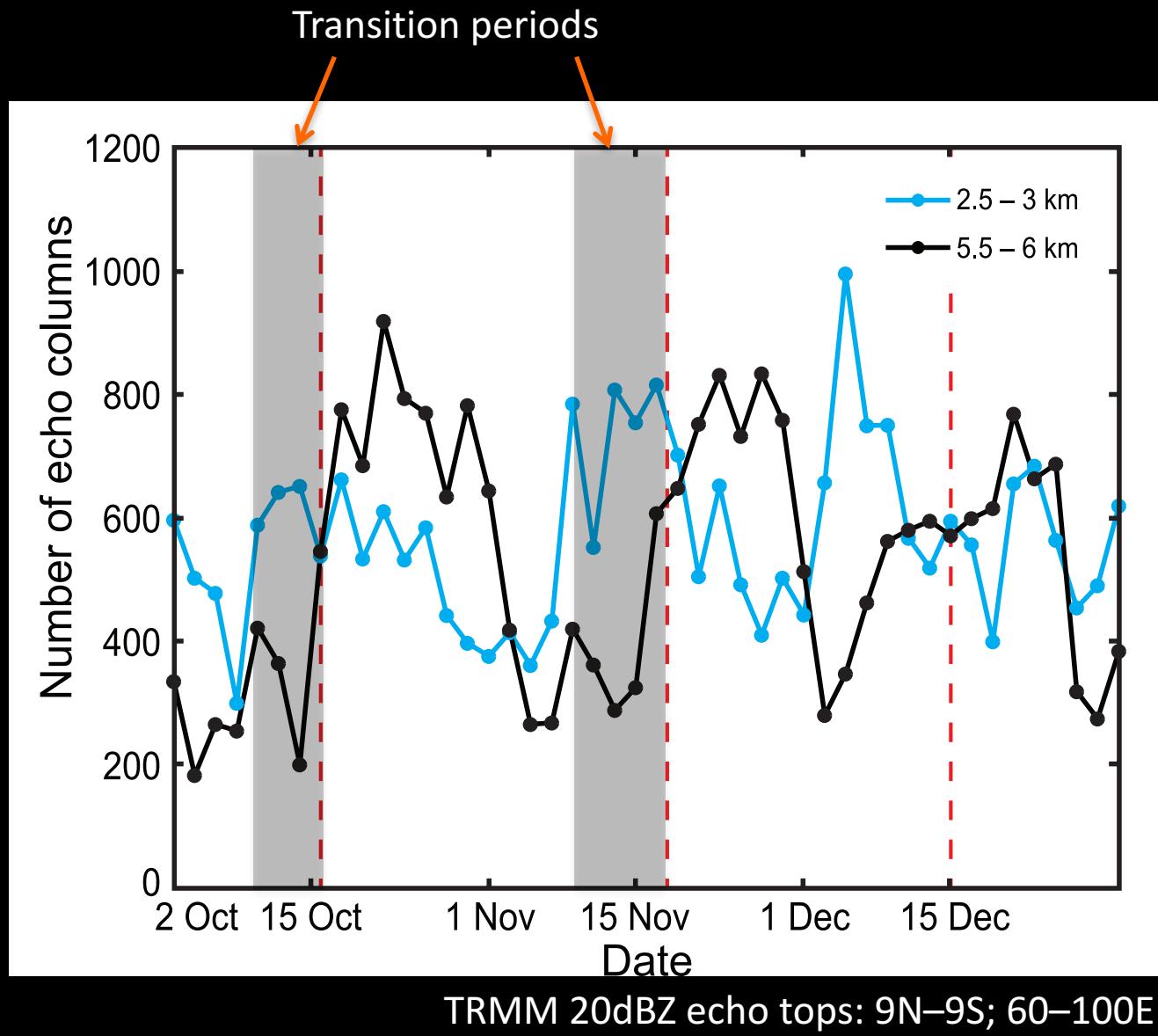
TRMM 20dBZ echo tops: 9N–9S; 60–100E



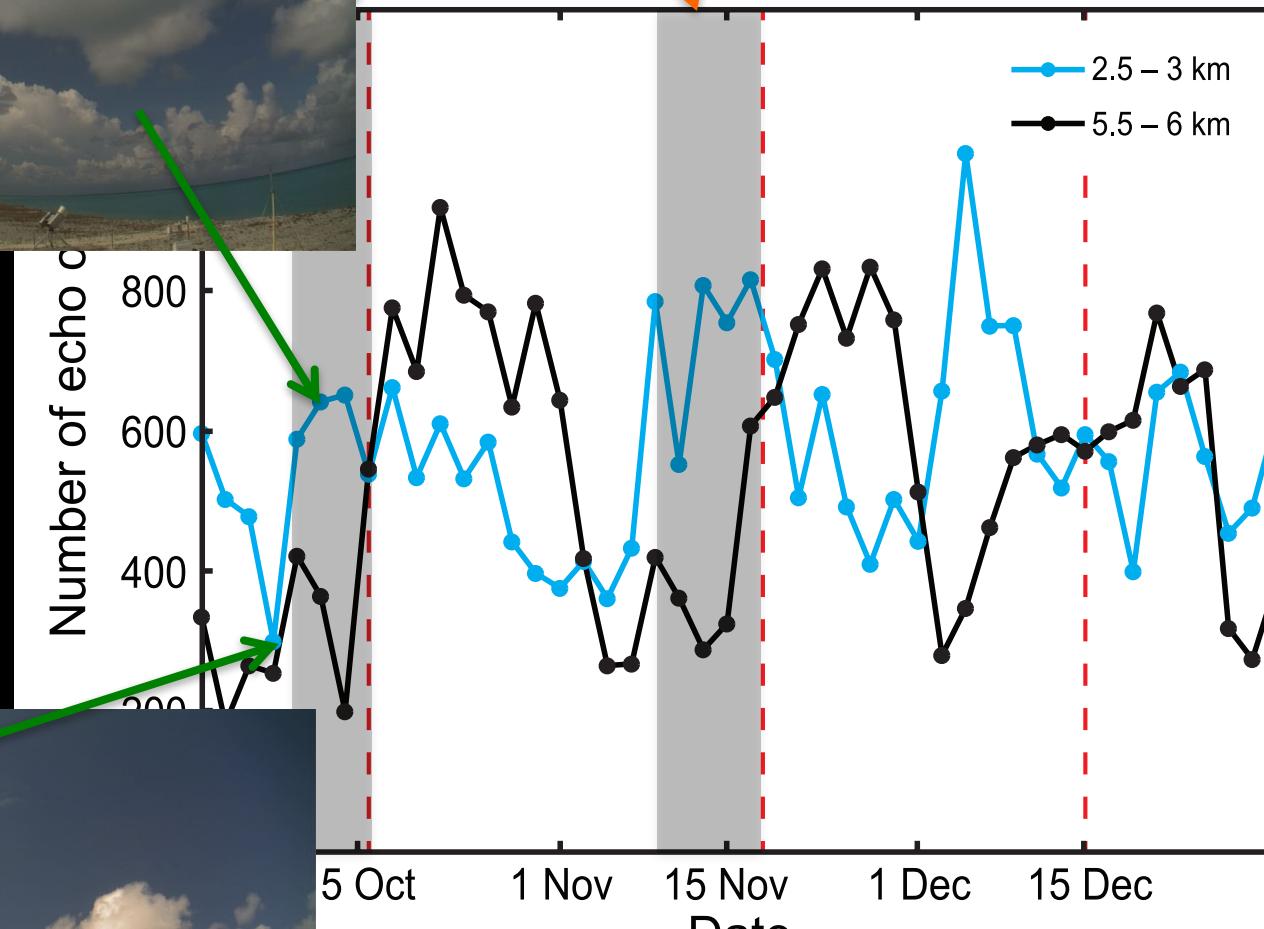
TRMM 20dBZ echo tops: 9N–9S; 60–100E



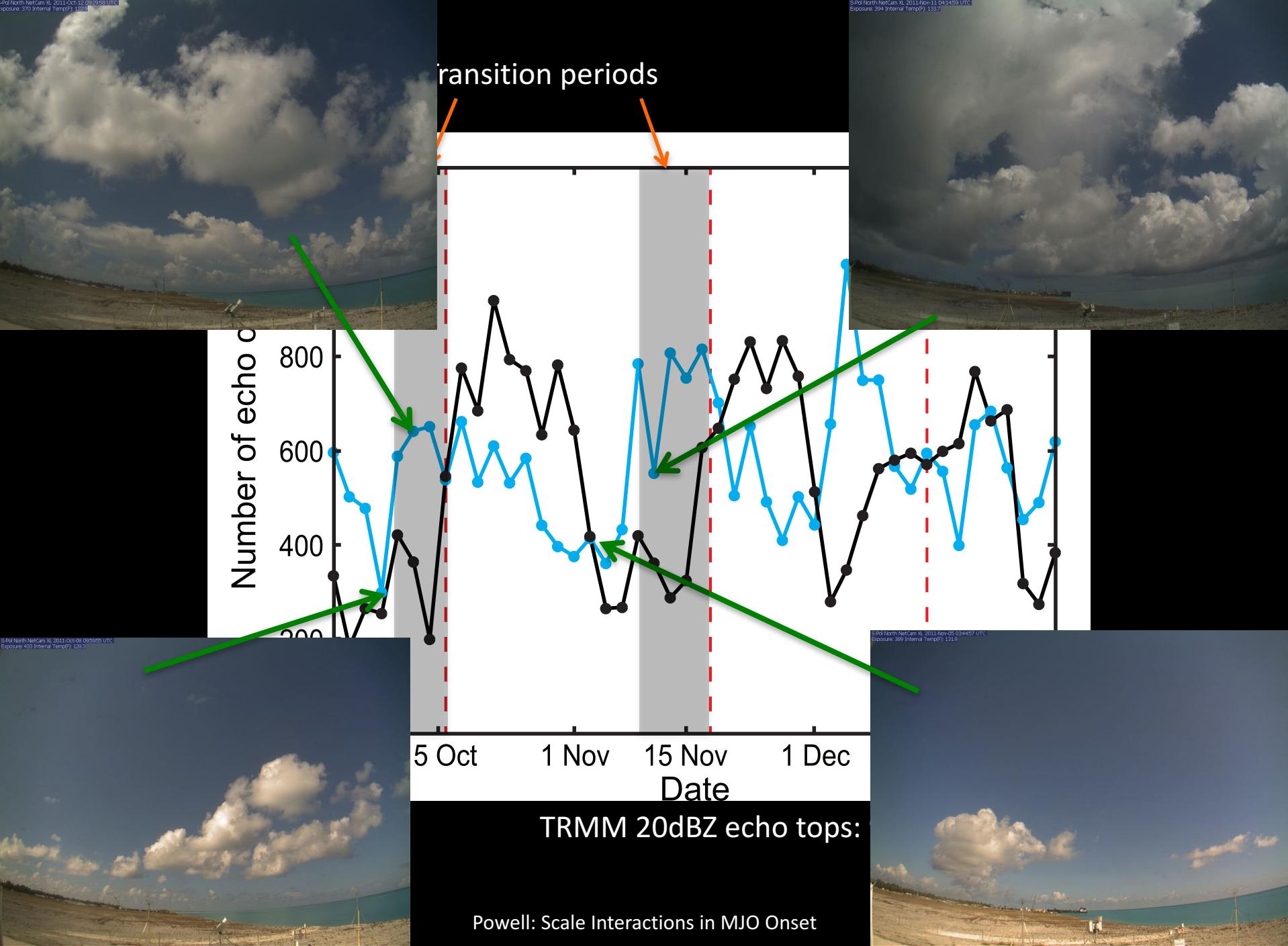
TRMM 20dBZ echo tops: 9N–9S; 60–100E

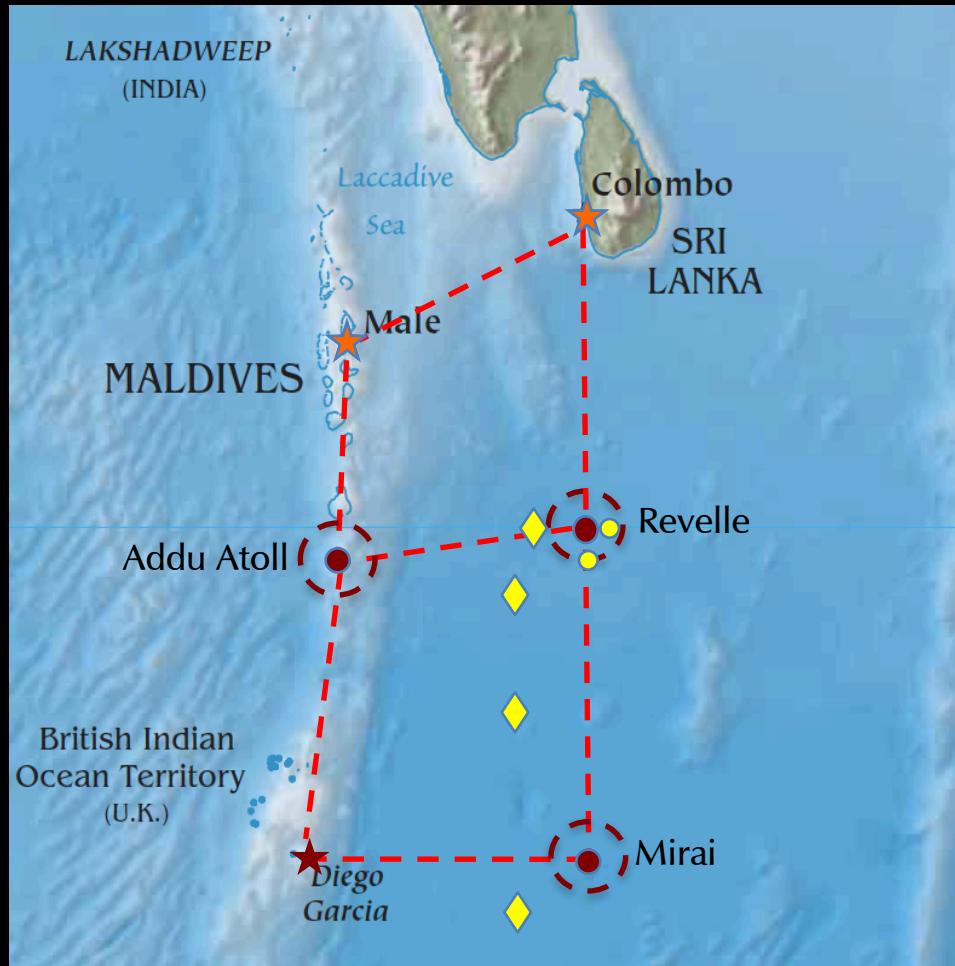


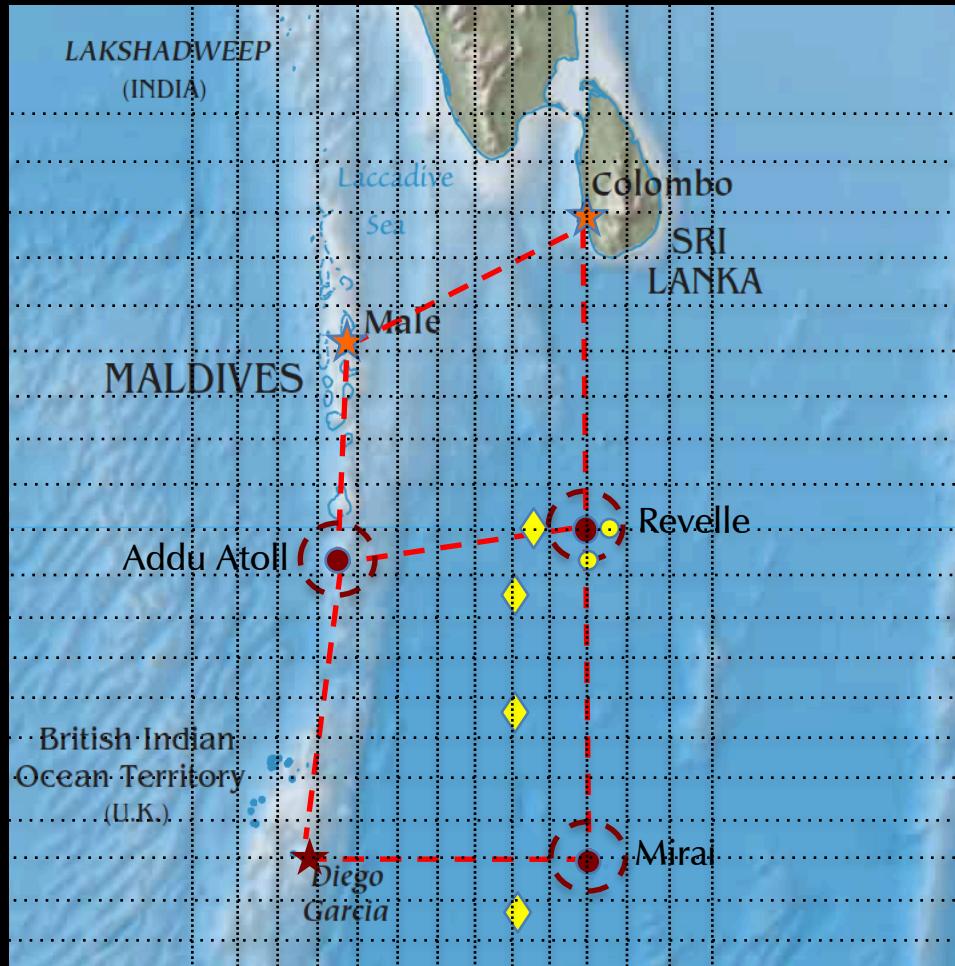
transition periods



TRMM 20dBZ echo tops: 9N–9S; 60–100E

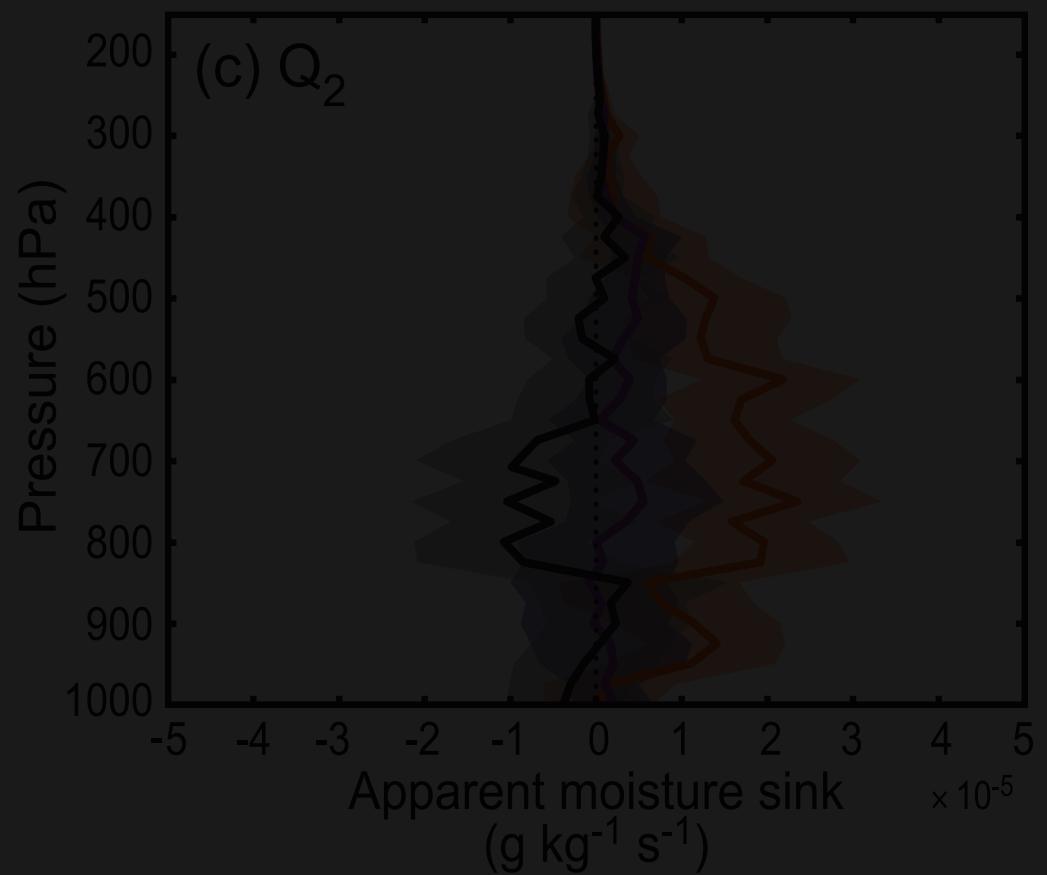






$$\frac{\partial q}{\partial t} = -\mathbf{v}_h \cdot \nabla q - \omega \frac{\partial q}{\partial p} - Q_2$$

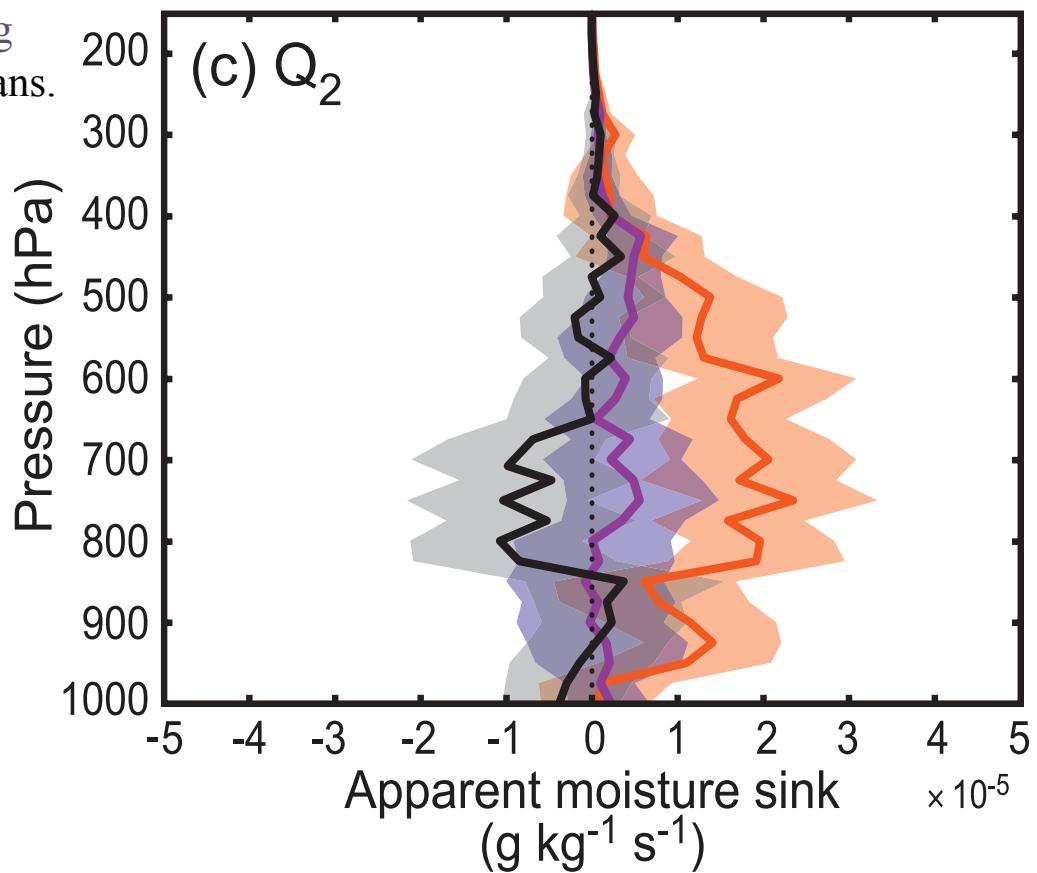
$$Q_2 = (\bar{c} - \bar{e}) + \frac{\partial}{\partial p} (\overline{\omega' q'})$$



$$\frac{\partial q}{\partial t} = -\mathbf{v}_h \cdot \nabla q - \omega \frac{\partial q}{\partial p} - Q_2$$

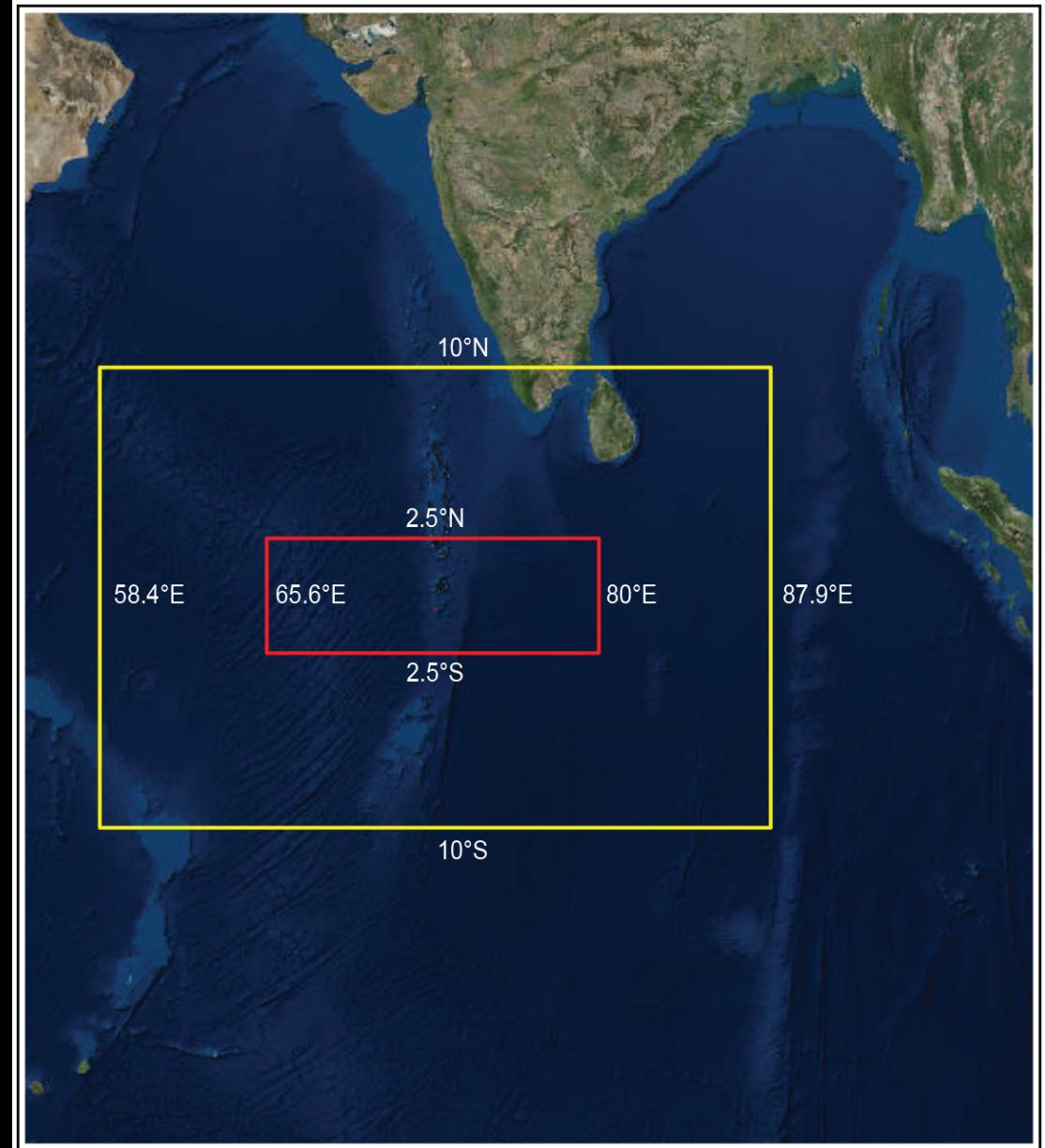
$$Q_2 = (\bar{c} - \bar{e}) + \frac{\partial}{\partial p}(\bar{\omega}' q')$$

Purple = Cg
 Black = Trans.
 Red = Cb

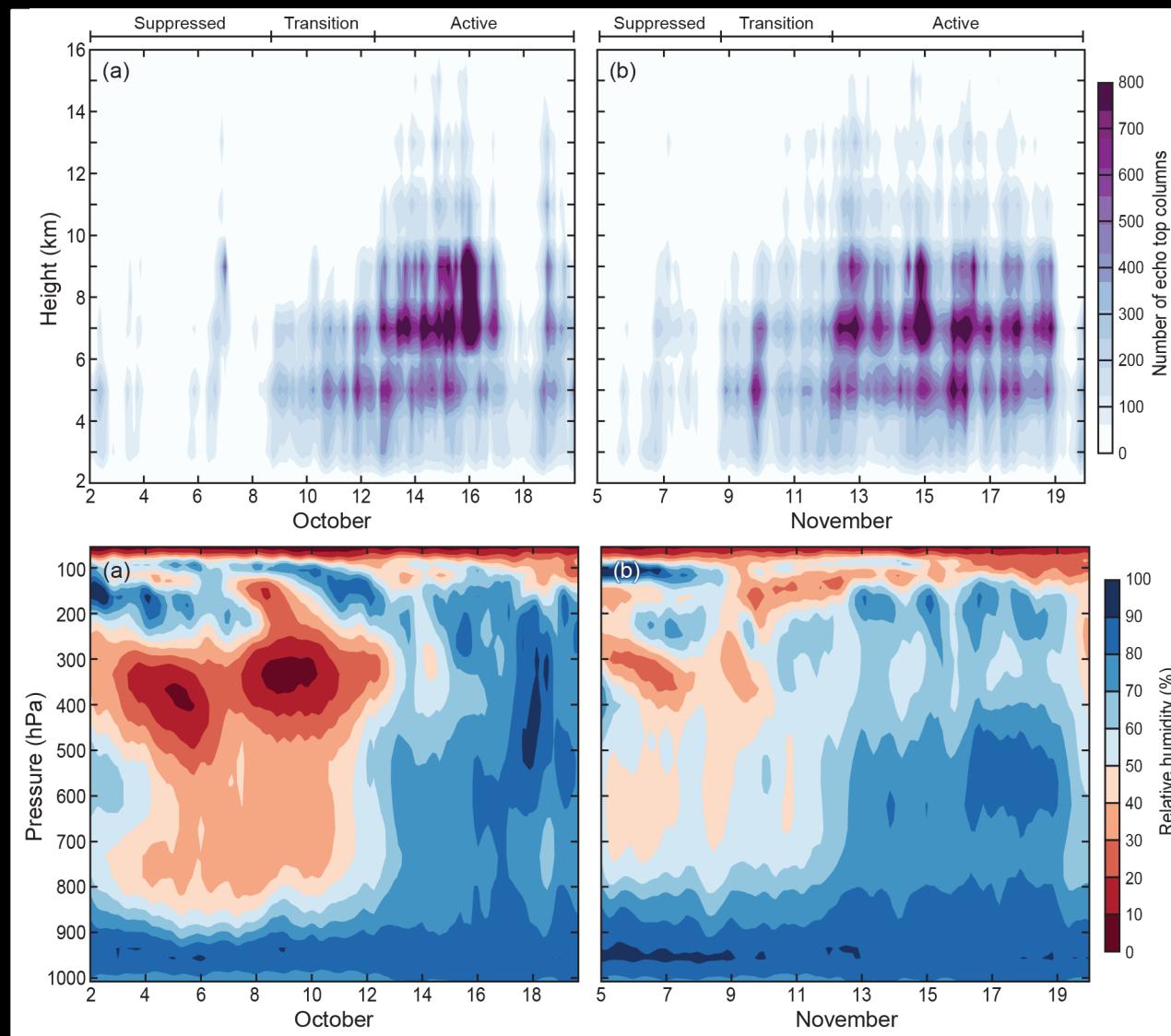


WRF V3.5.1

- 2 km grid spacing
- Thompson
microphysics
(following, e.g., Powell
et al. 2012)
- MYJ PBL scheme
- Forced with ERA-I
every 6 hours and
NOAA RTG for sea
surface temperature
- 1–20 October and 4–
20 November 2011

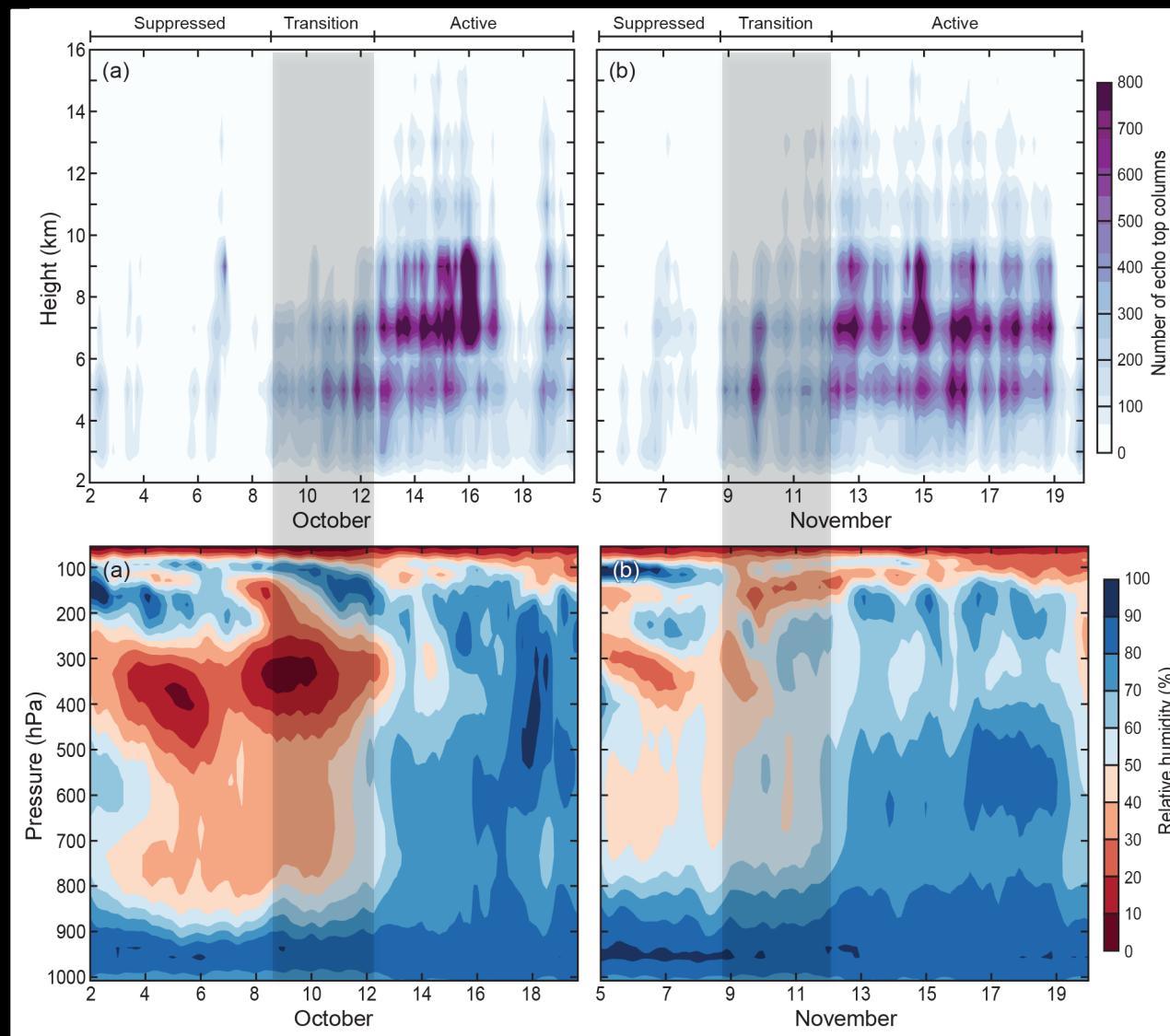


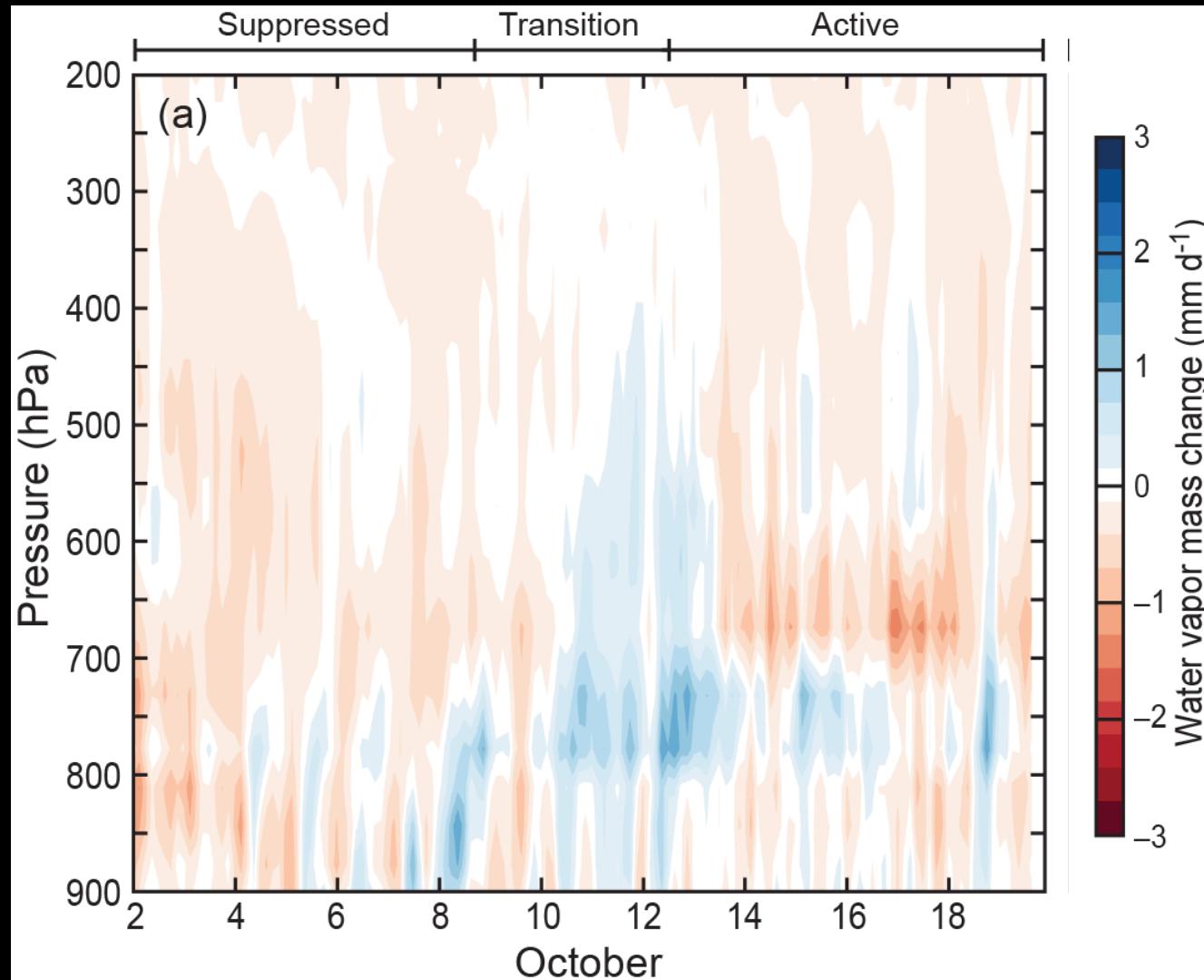
20 dBZ echo
top height
frequency



Relative
Humidity

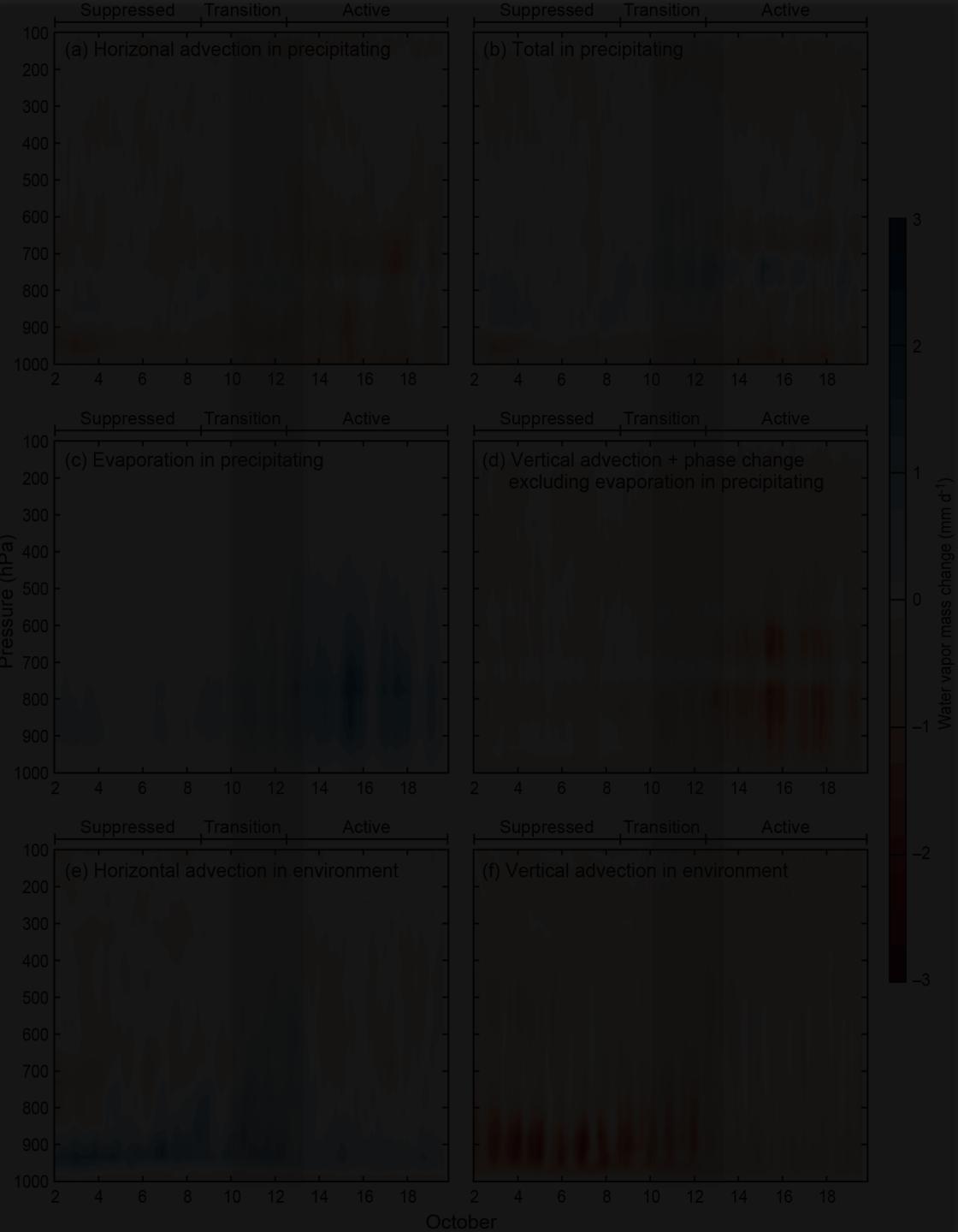
20 dBZ echo
 top height
 frequency





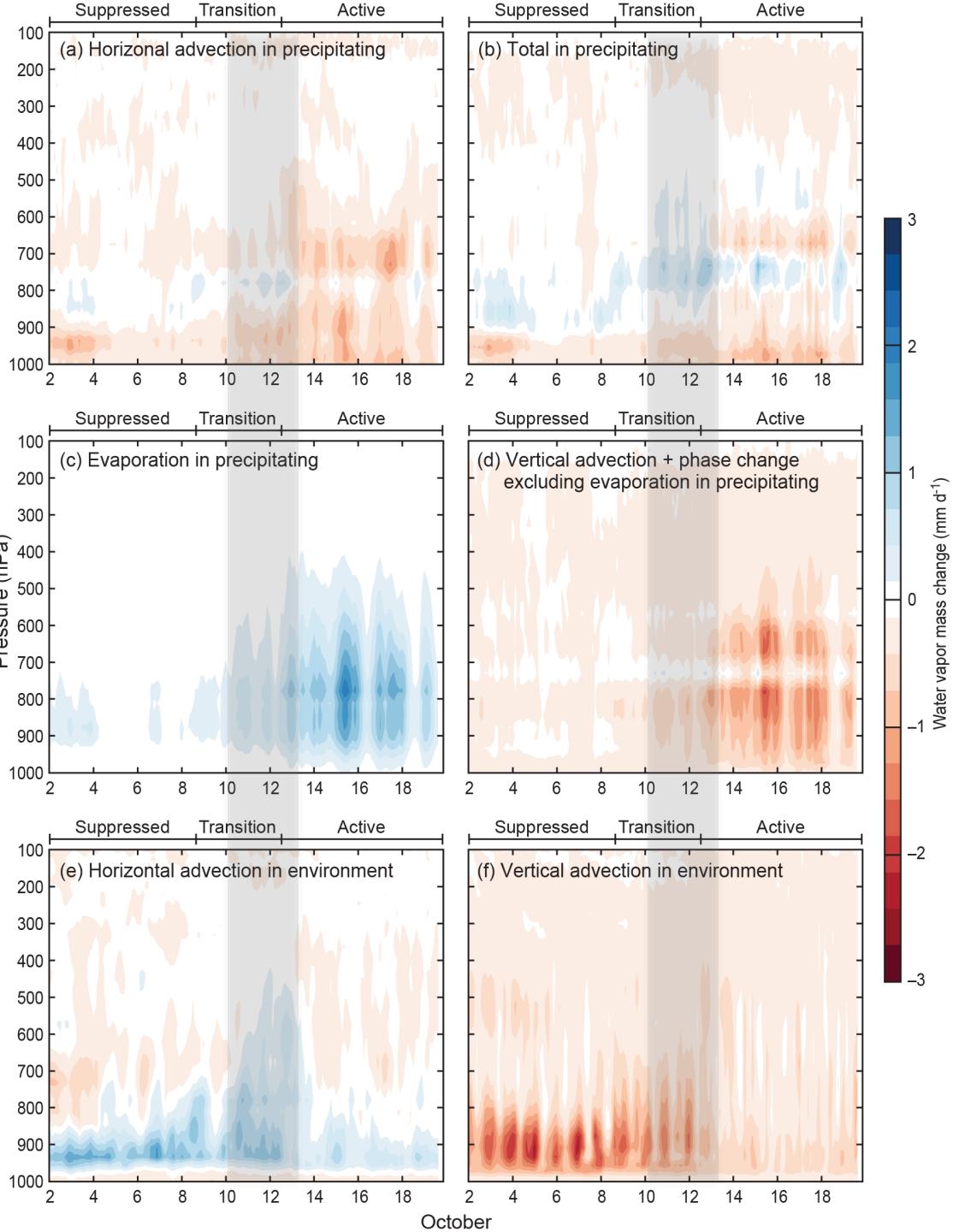
$$\frac{\partial m_{grid}}{\partial t} = -\frac{dP}{g} dx^2 (\mathbf{u} \cdot \nabla q) + M$$

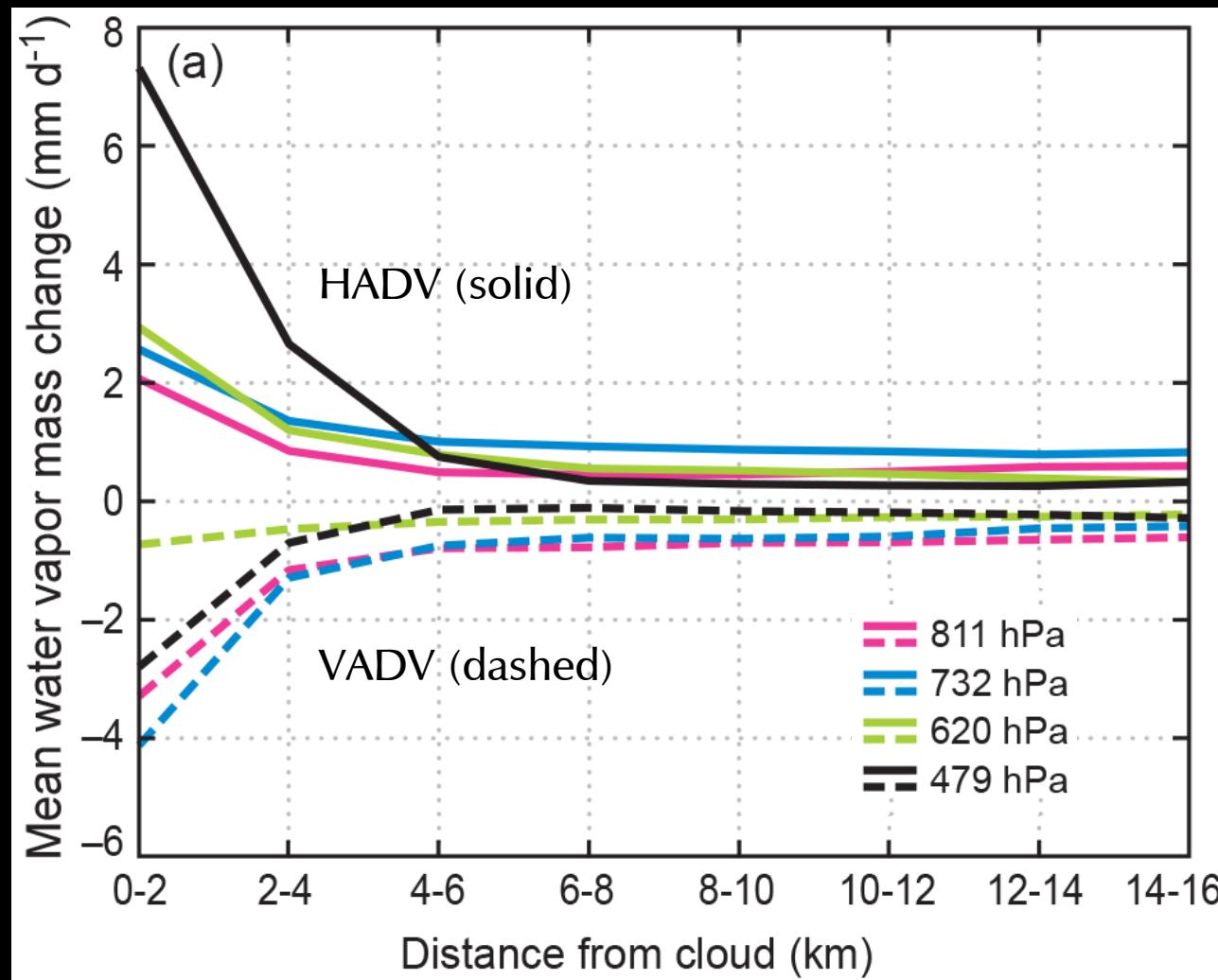
- HADV in precipitating clouds
- VADV in precipitating clouds
- Net phase change in precipitating clouds
- HADV in clear-air environment
- VADV in clear-air environment



$$\frac{\partial m_{grid}}{\partial t} = -\frac{dP}{g} dx^2 (\mathbf{u} \cdot \nabla q) + M$$

- HADV in precipitating clouds
- VADV in precipitating clouds
- Net phase change in precipitating clouds
- HADV in clear-air environment
- VADV in clear-air environment

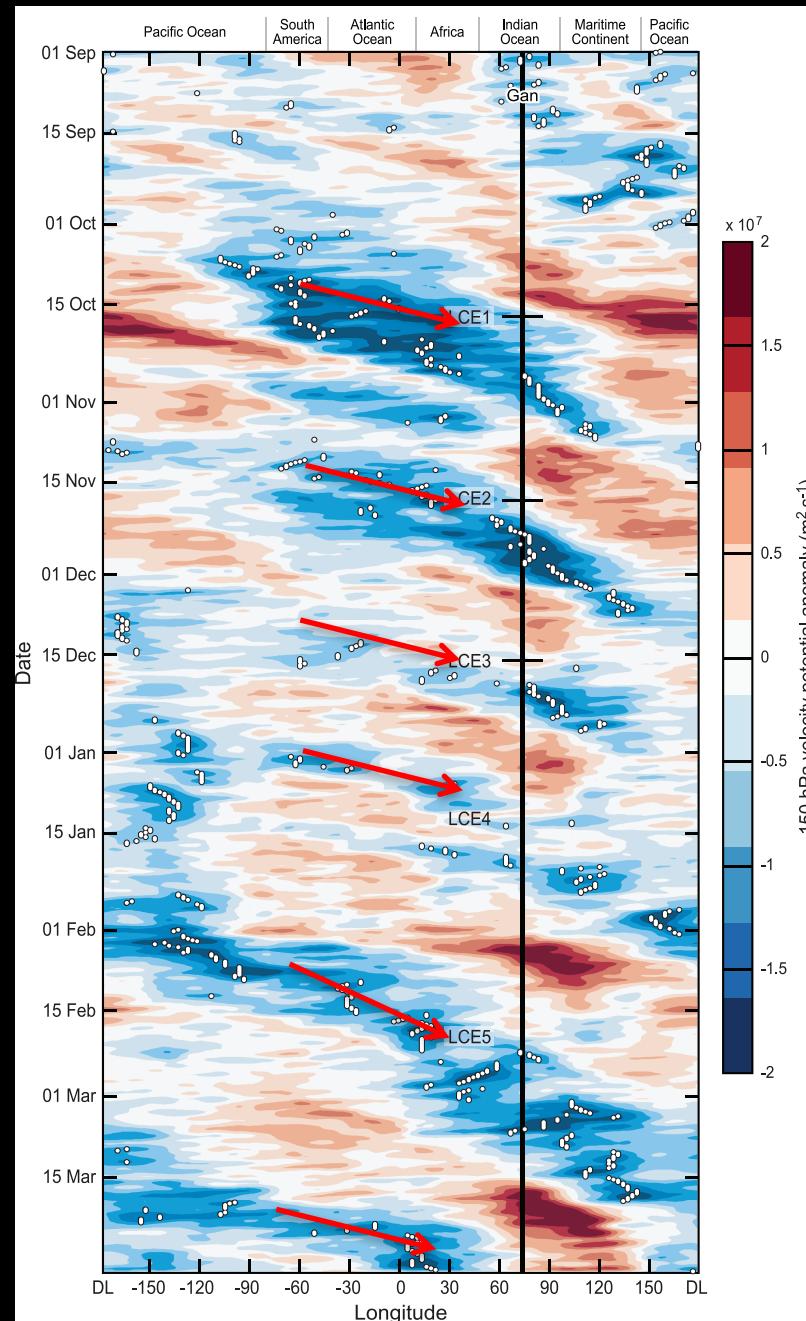




The Circumnavigating Kelvin Wave

How does LS upper-tropospheric divergence relate to convection rooted in a warm, moist boundary layer?

Hypothesis: Convection passively responds to changes in the large-scale environment.



Originally: Knutson
and Weickmann
(1987)

Figure: Powell and
Houze (2015b)

Large-scale vertical velocity anomalies are in phase with velocity potential anomalies.

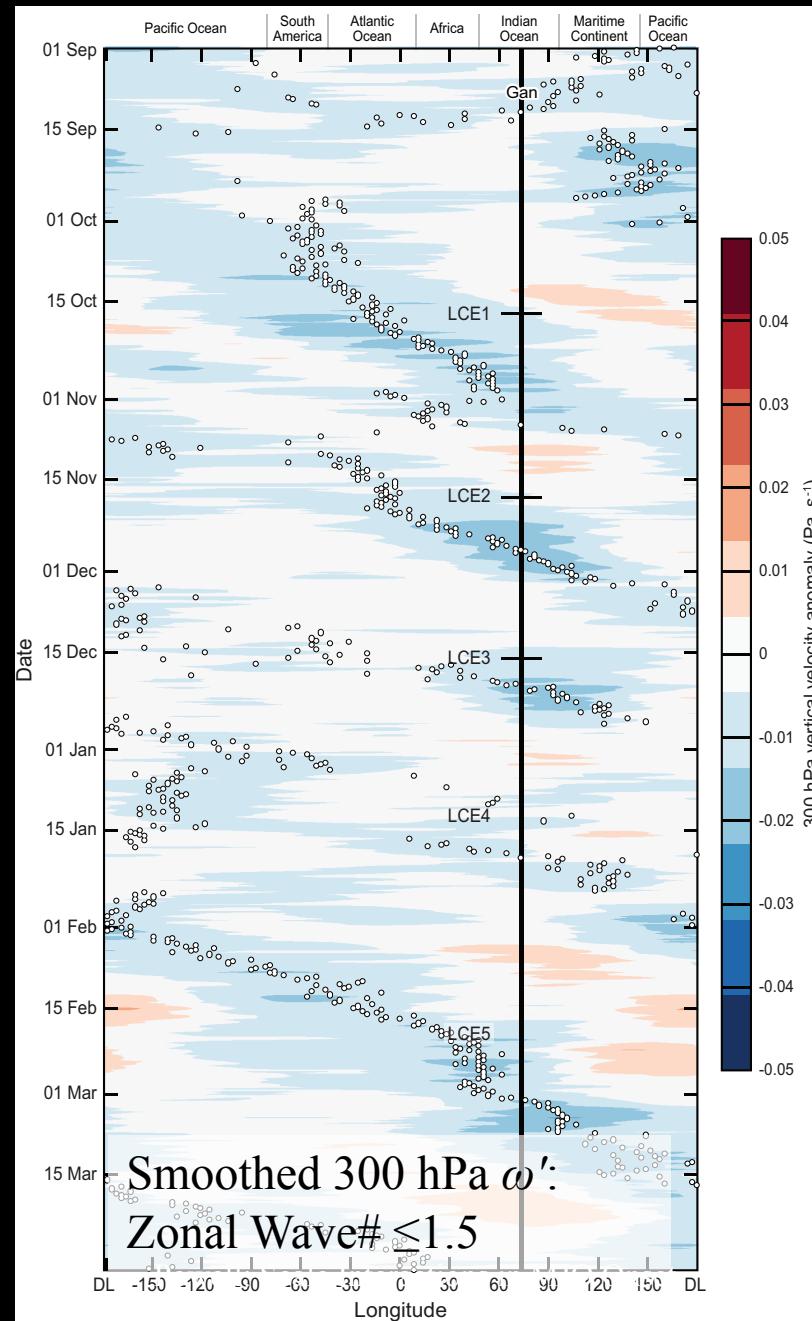
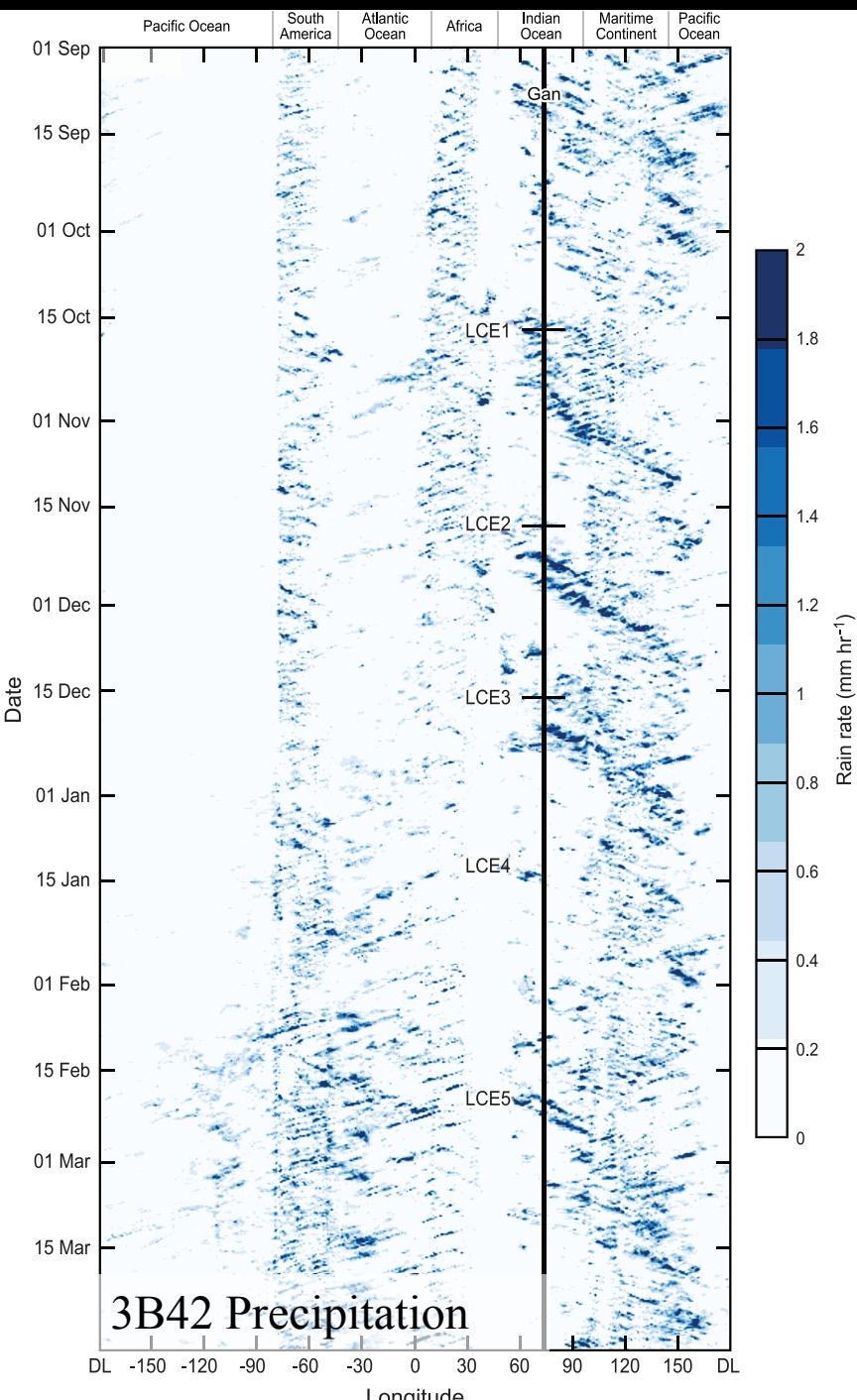
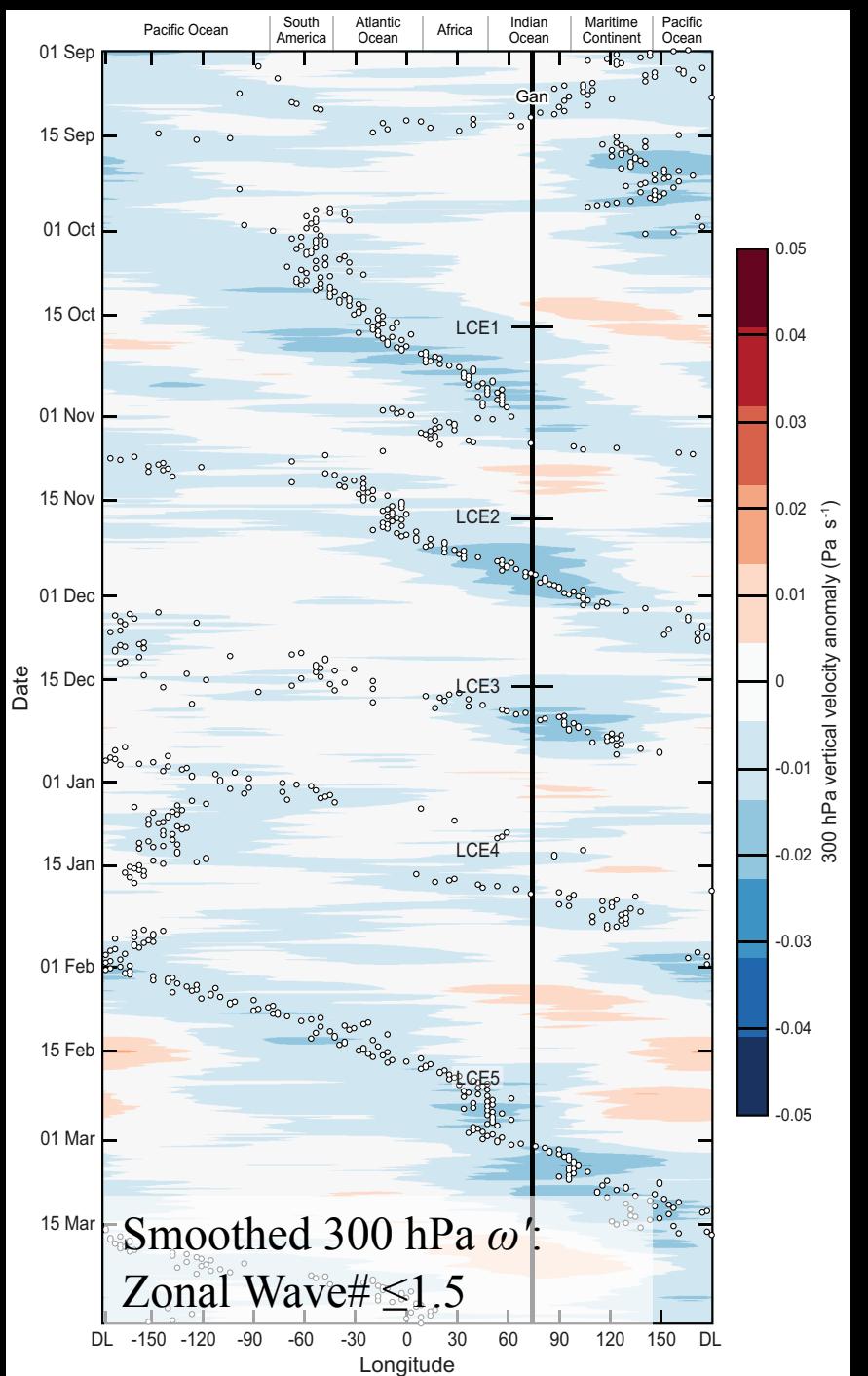
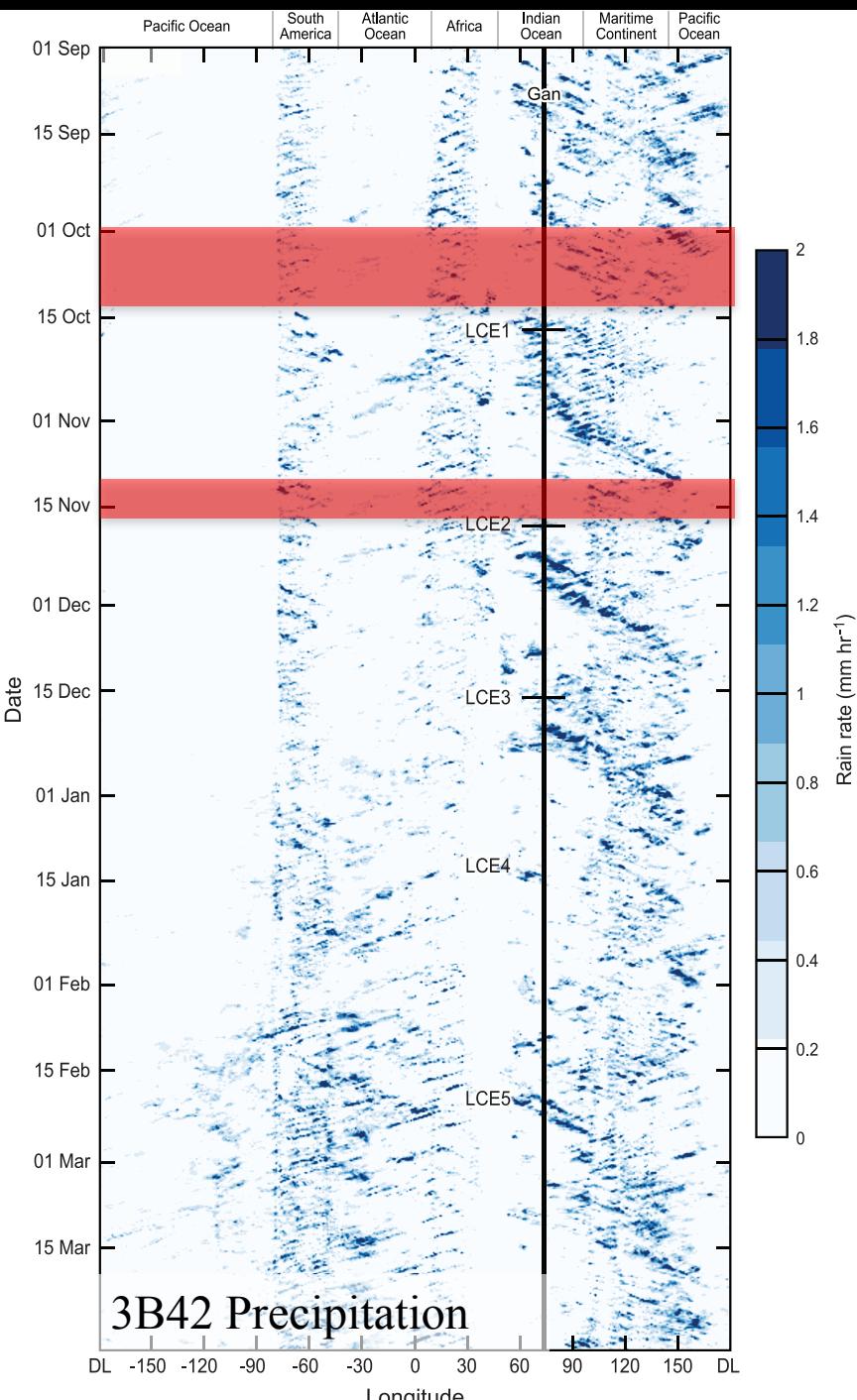
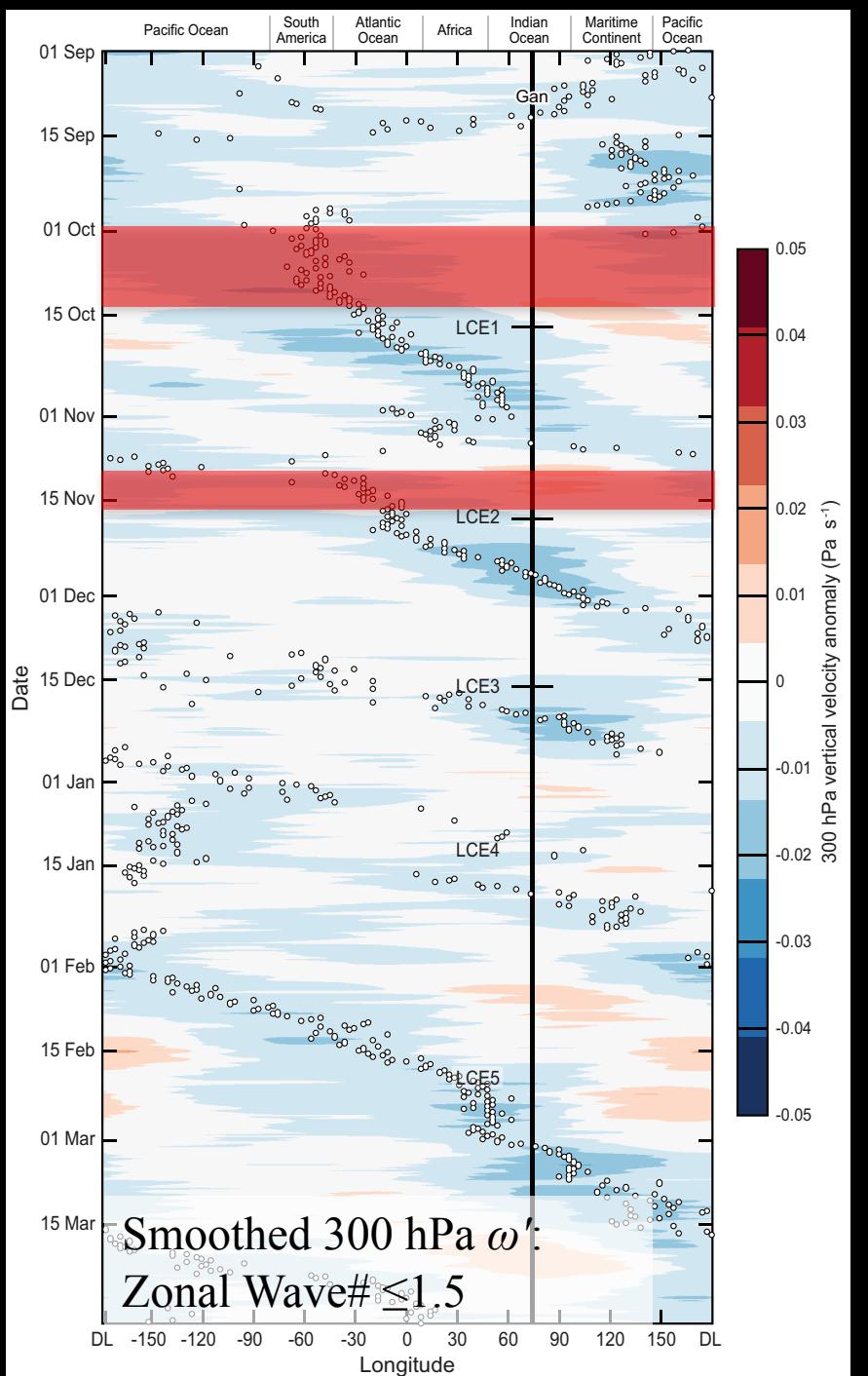
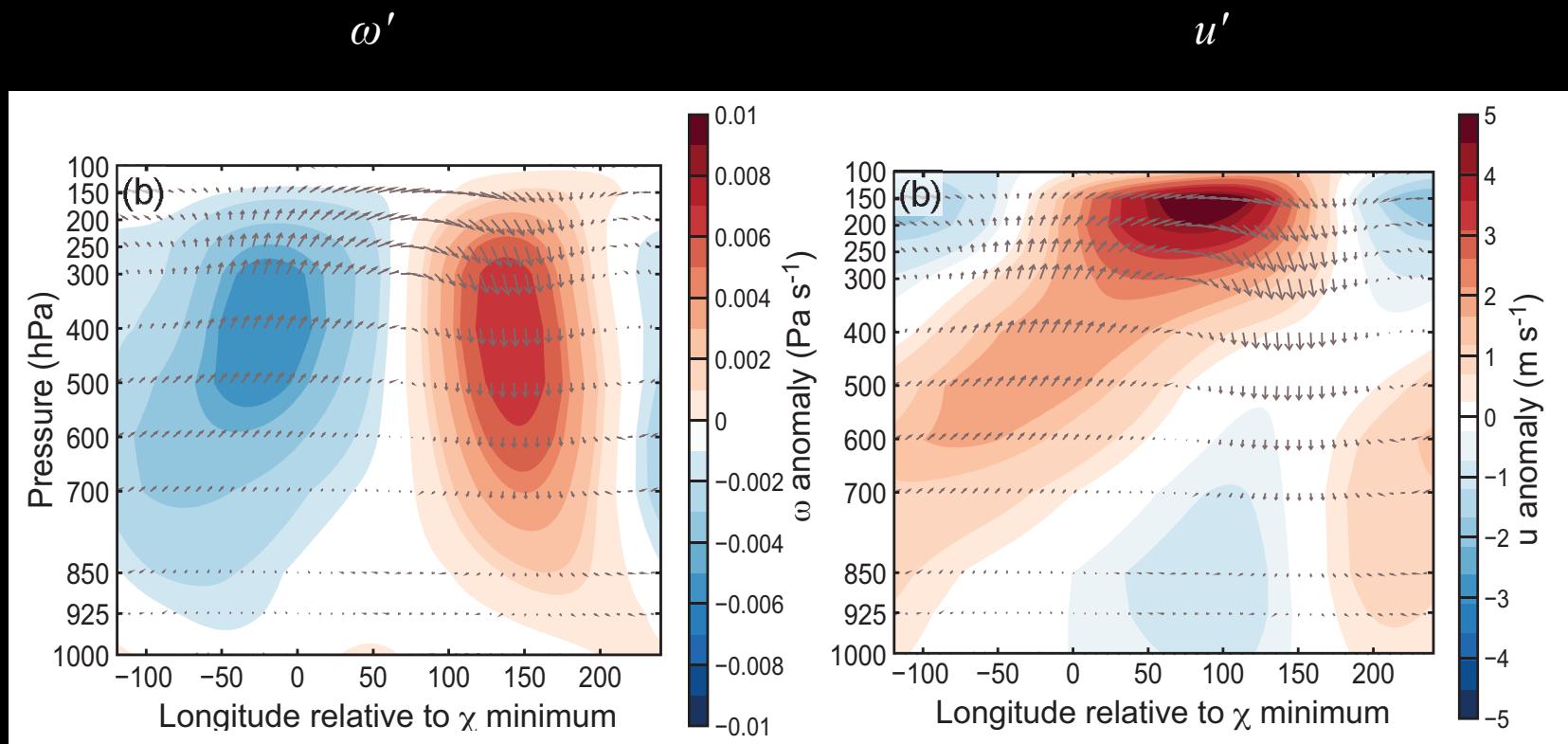
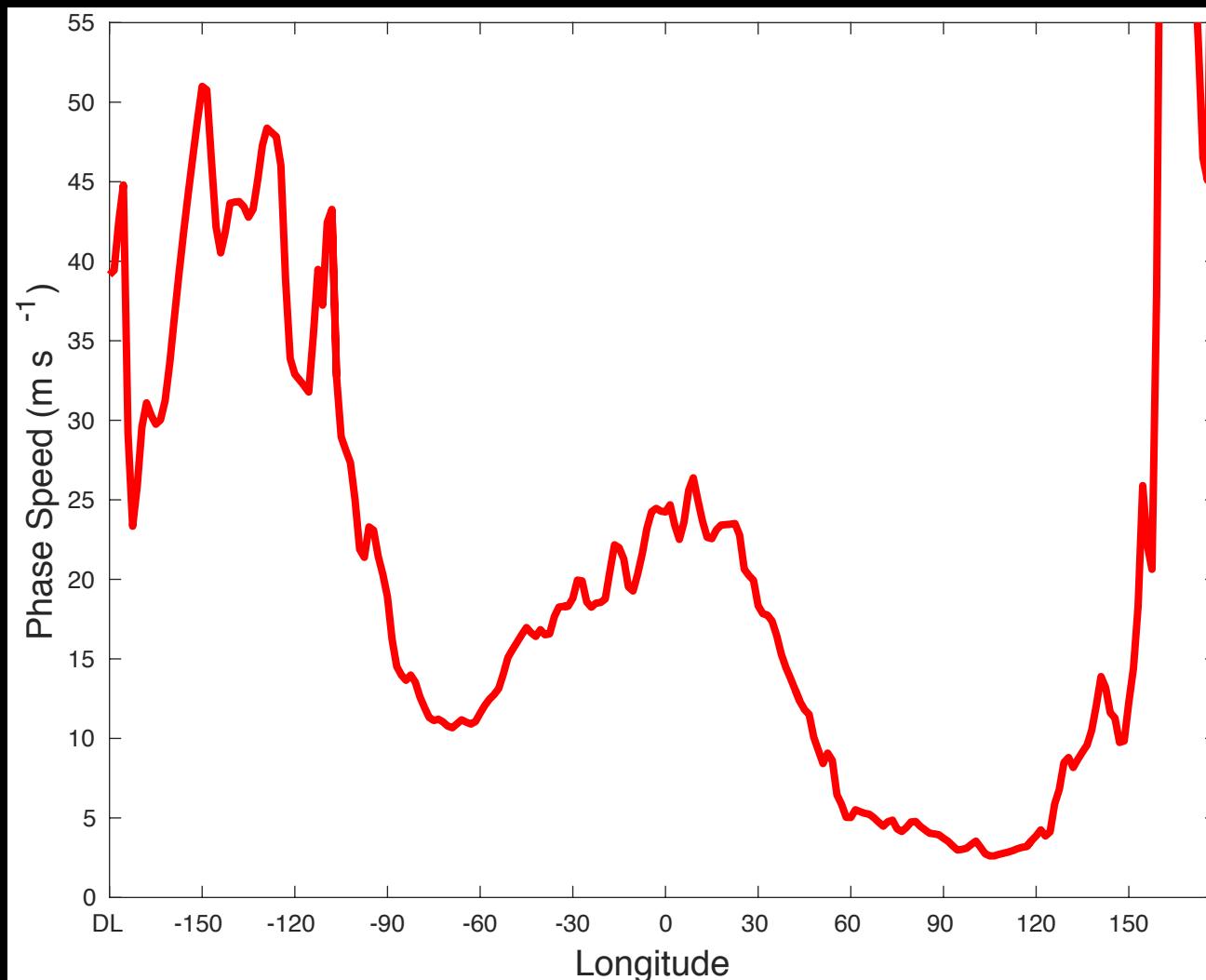


Figure: Powell and Houze (2015b)

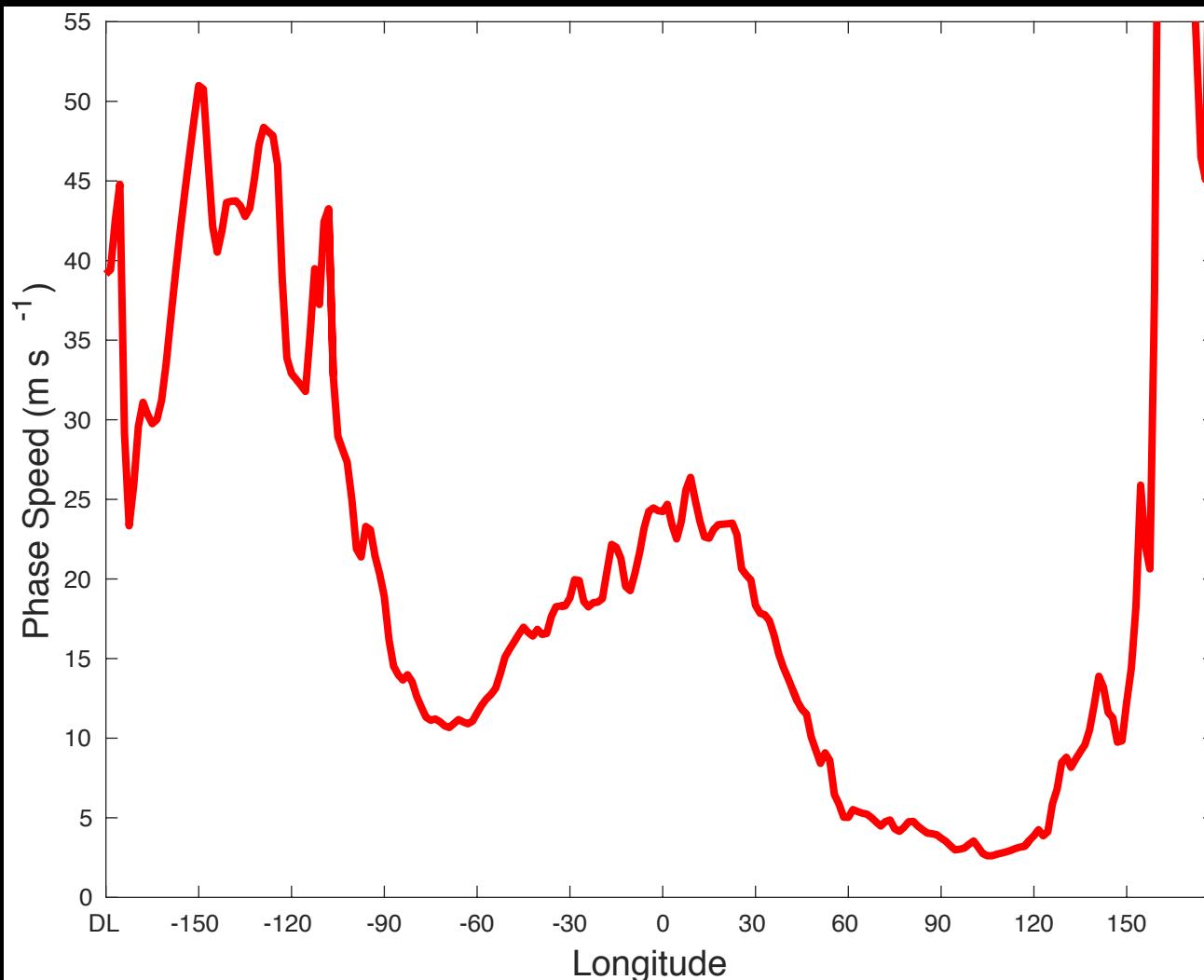




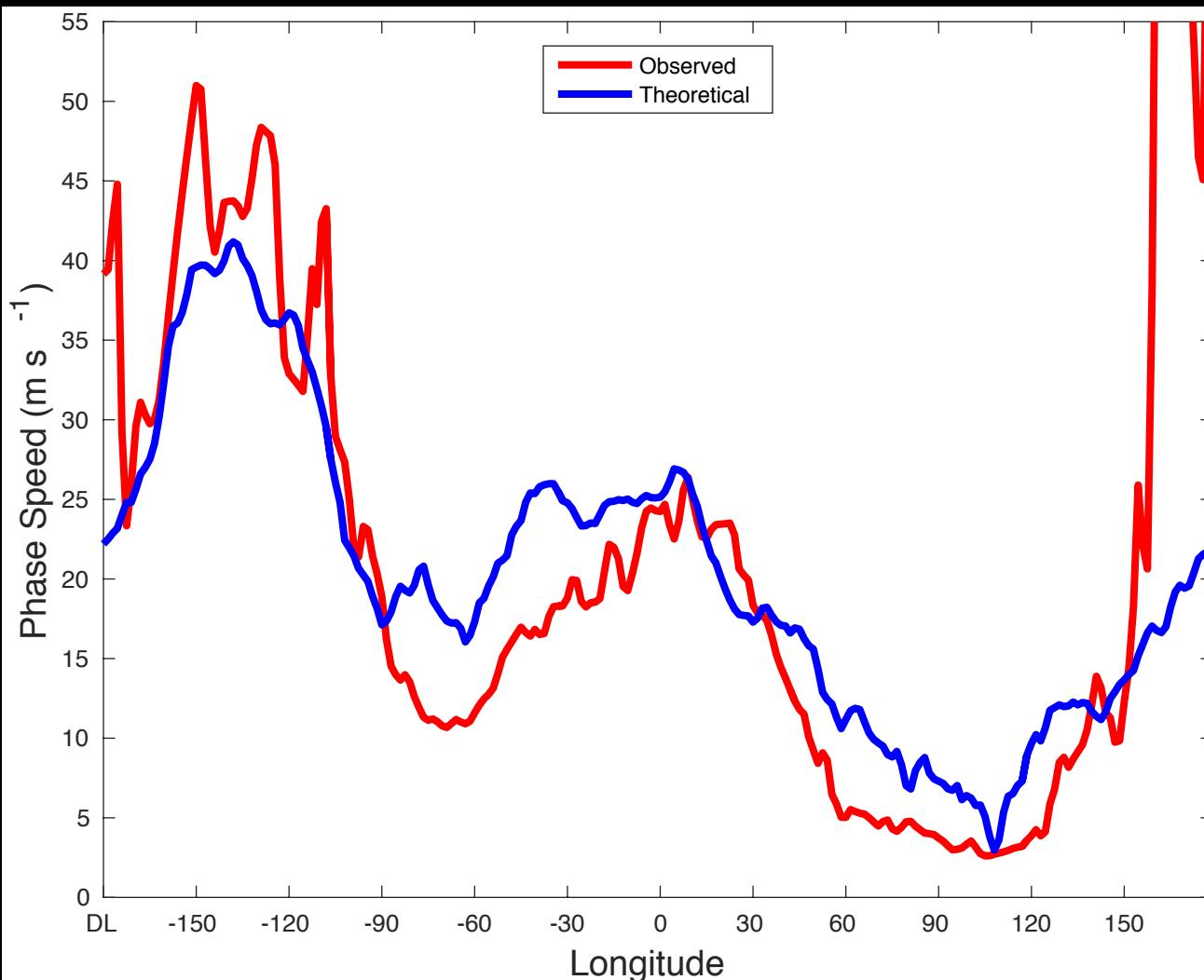




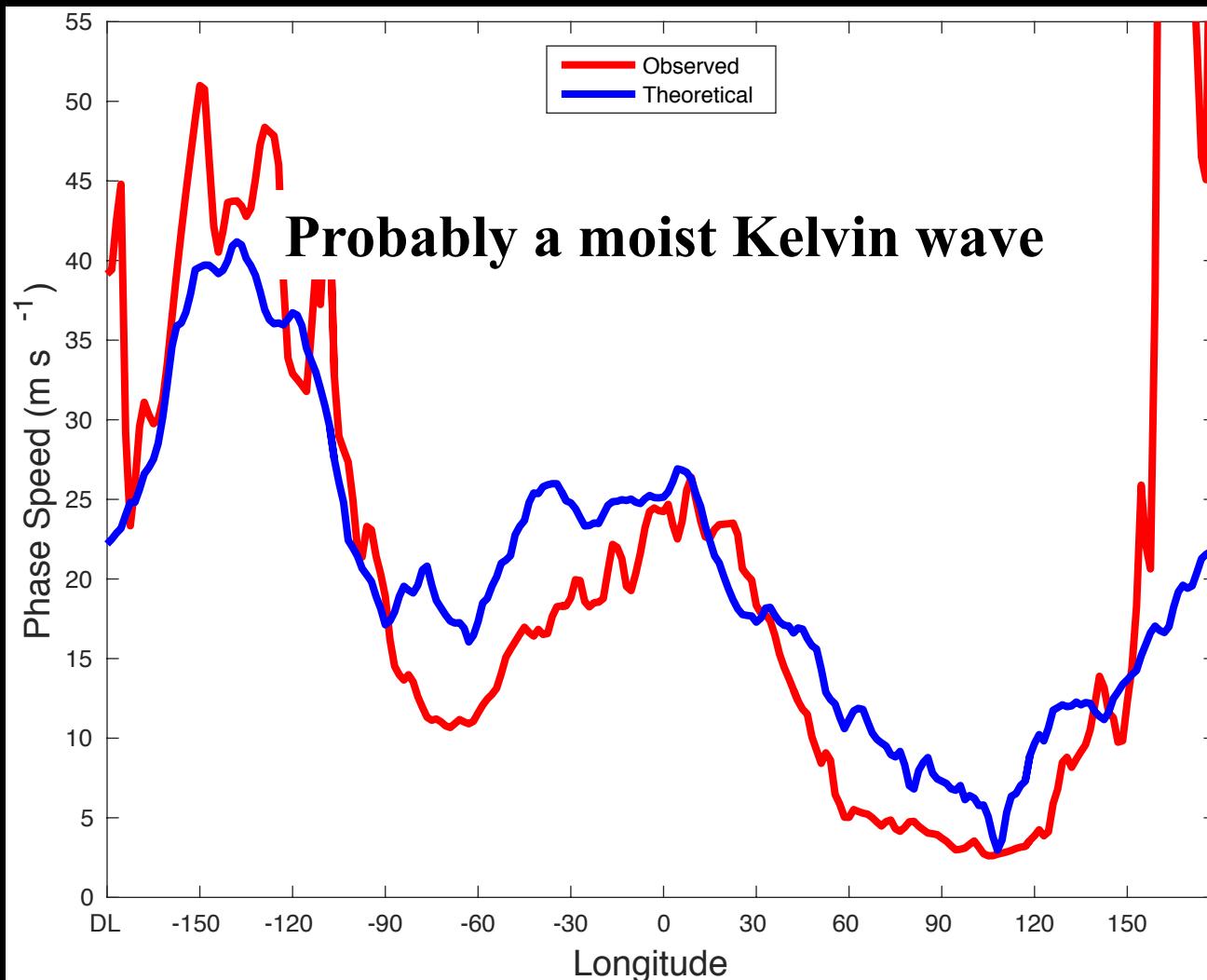
$$\frac{\partial T}{\partial t} - S\omega = Q, Q \approx -\mu S\omega, c = \sqrt{(1-\mu)gh_e}$$

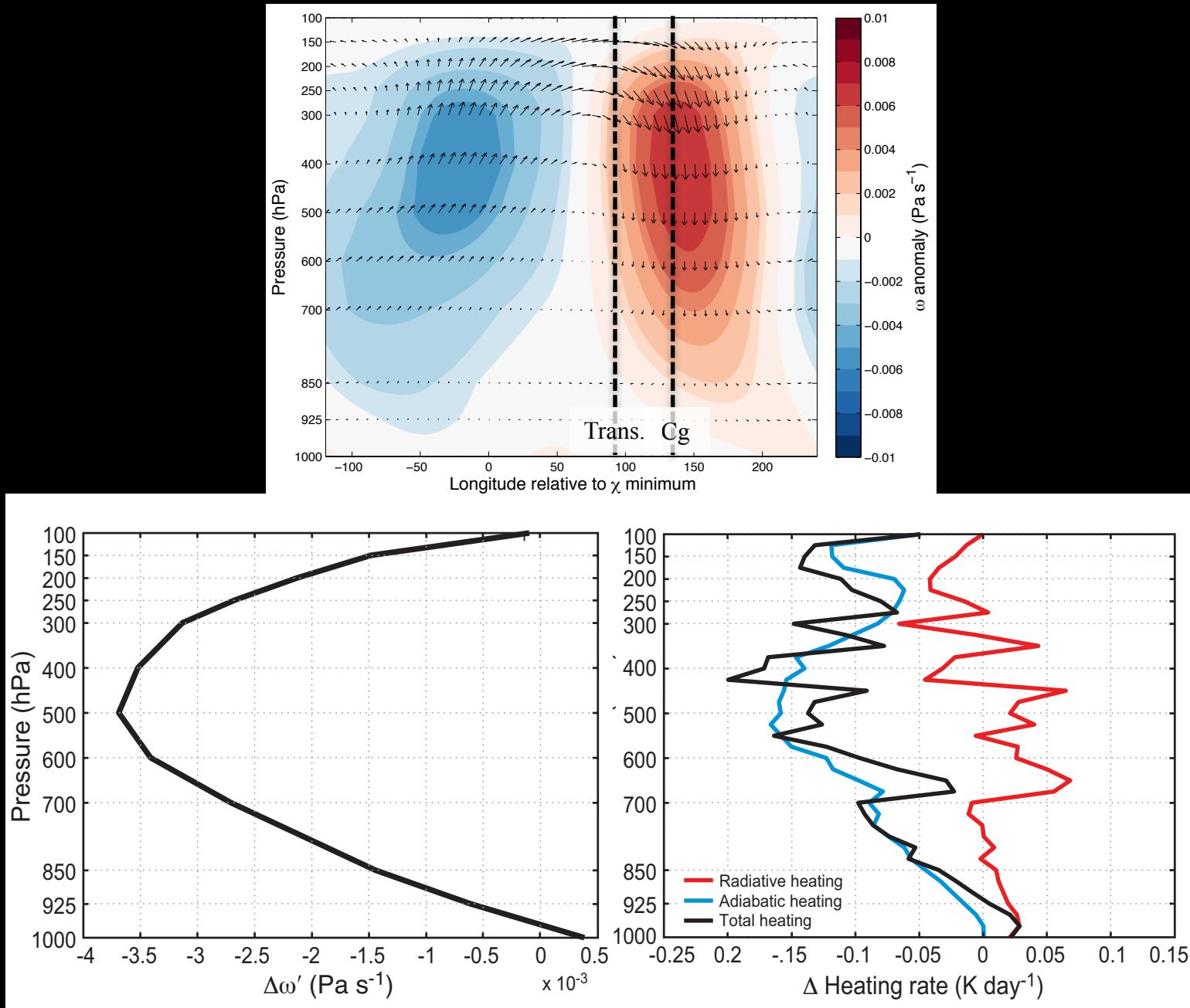


$$\frac{\partial T}{\partial t} - S\omega = Q, Q \approx -\mu S\omega, c = \sqrt{(1-\mu)gh_e}$$

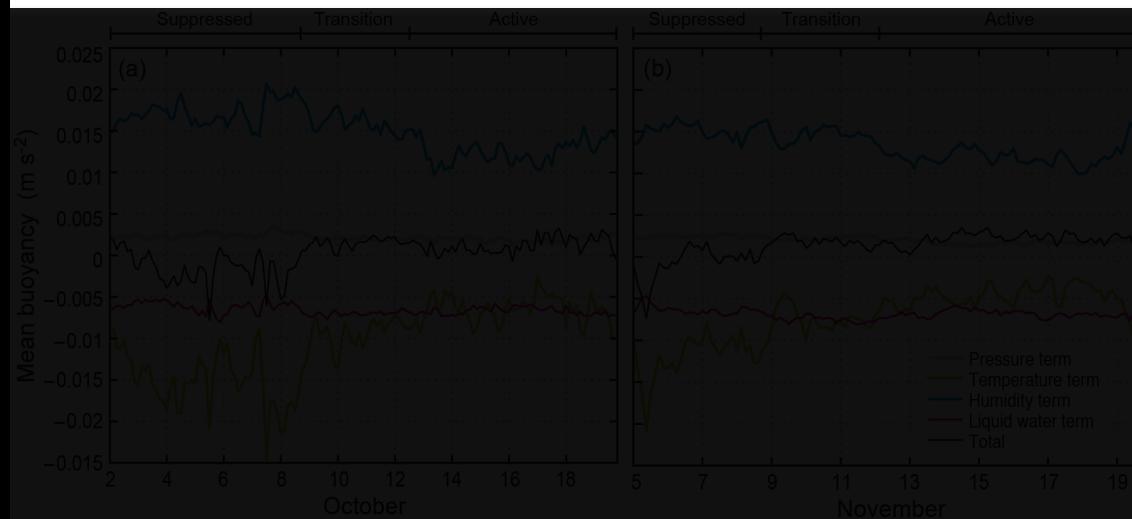
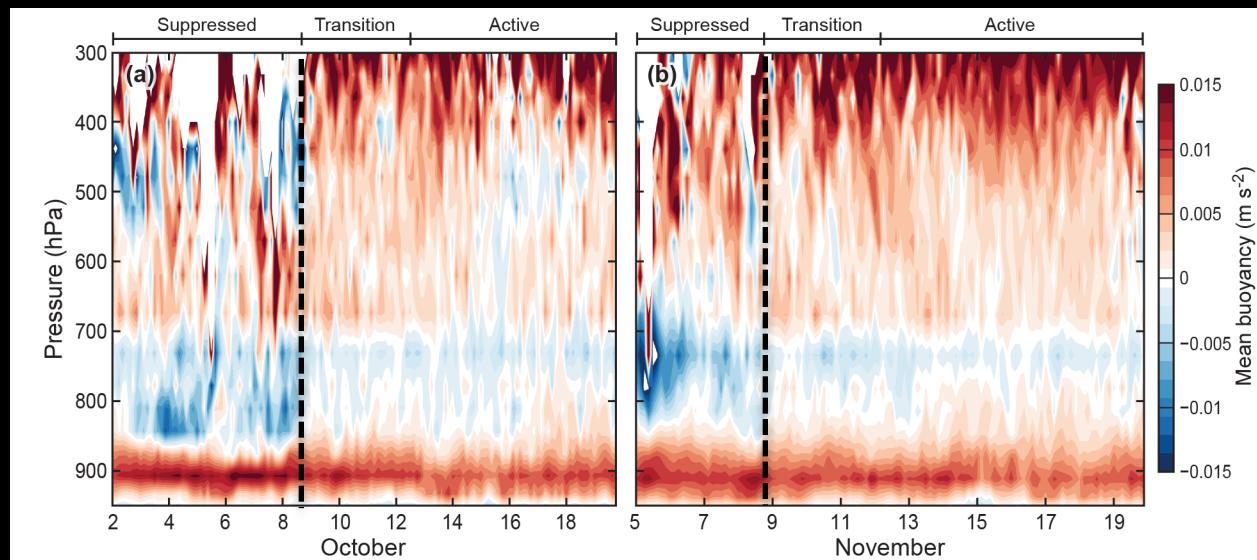


$$\frac{\partial T}{\partial t} - S\omega = Q, Q \approx -\mu S\omega, c = \sqrt{(1-\mu)gh_e}$$





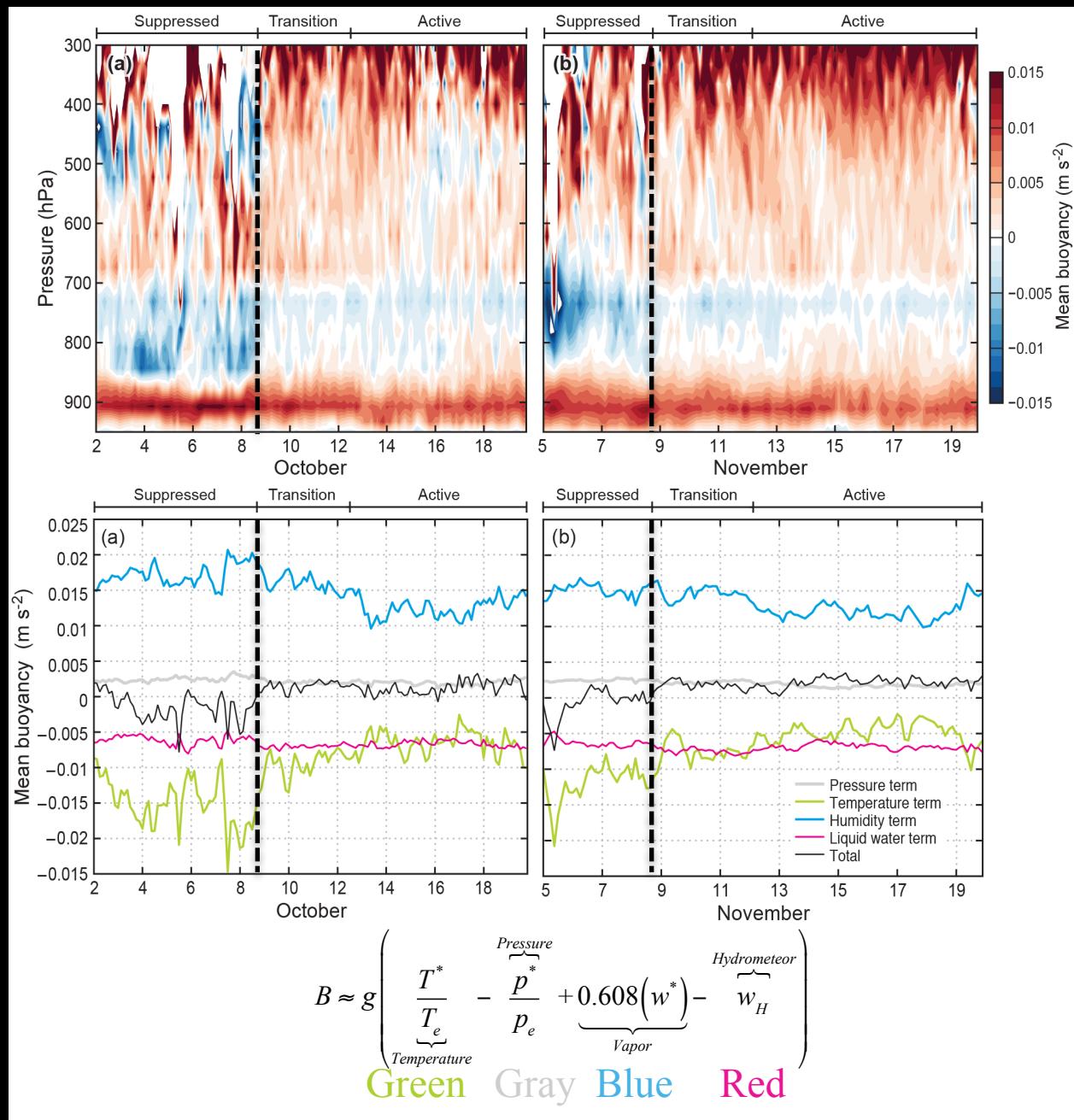
Updraft
buoyancy for
convective
echoes with
 $w \geq 0.3 \text{ m s}^{-1}$



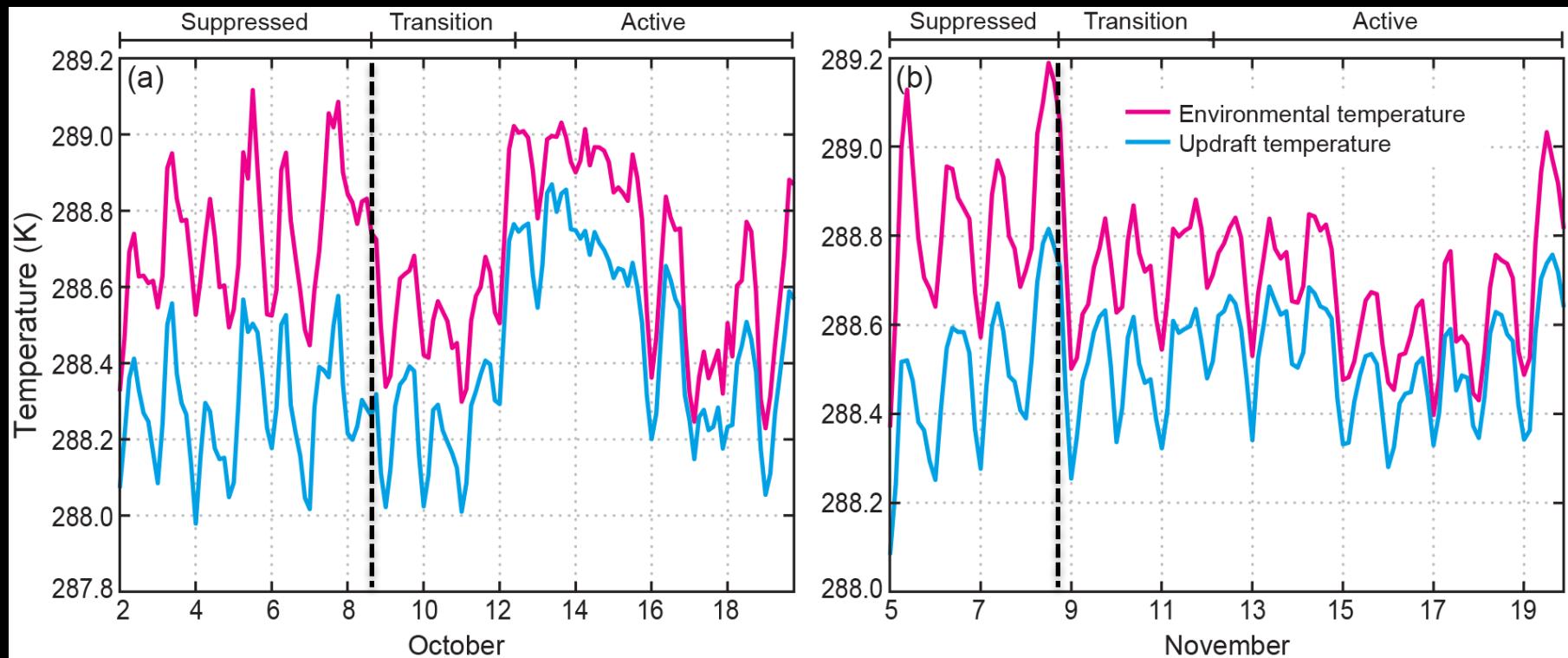
$$B \approx g \left(\underbrace{\frac{T^*}{T_e}}_{\text{Temperature}} - \underbrace{\frac{p^*}{p_e}}_{\text{Pressure}} + \underbrace{0.608(w^*)}_{\text{Vapor}} - \underbrace{w_H}_{\text{Hydrometeor}} \right)$$

Green Gray Blue Red

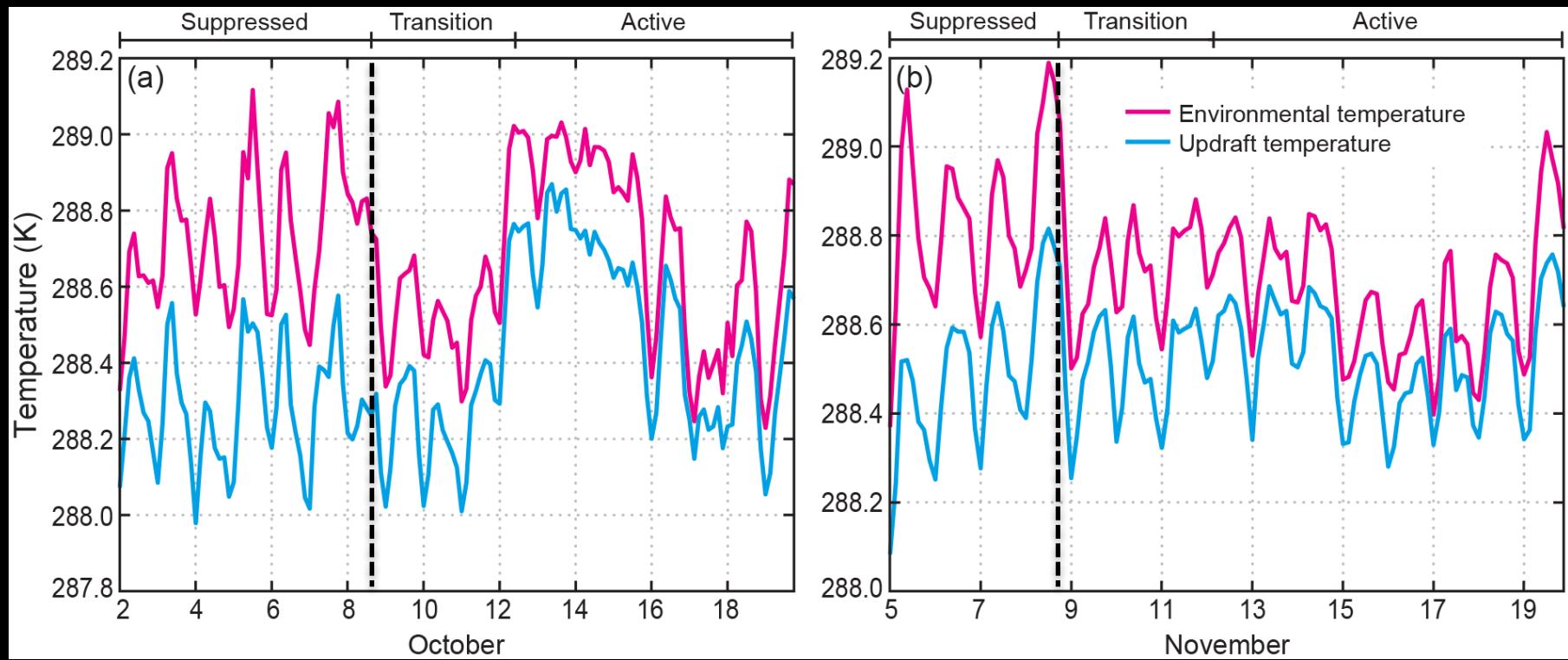
Individual
terms in
buoyancy
equation:
Mean in 700–
850 mb layer



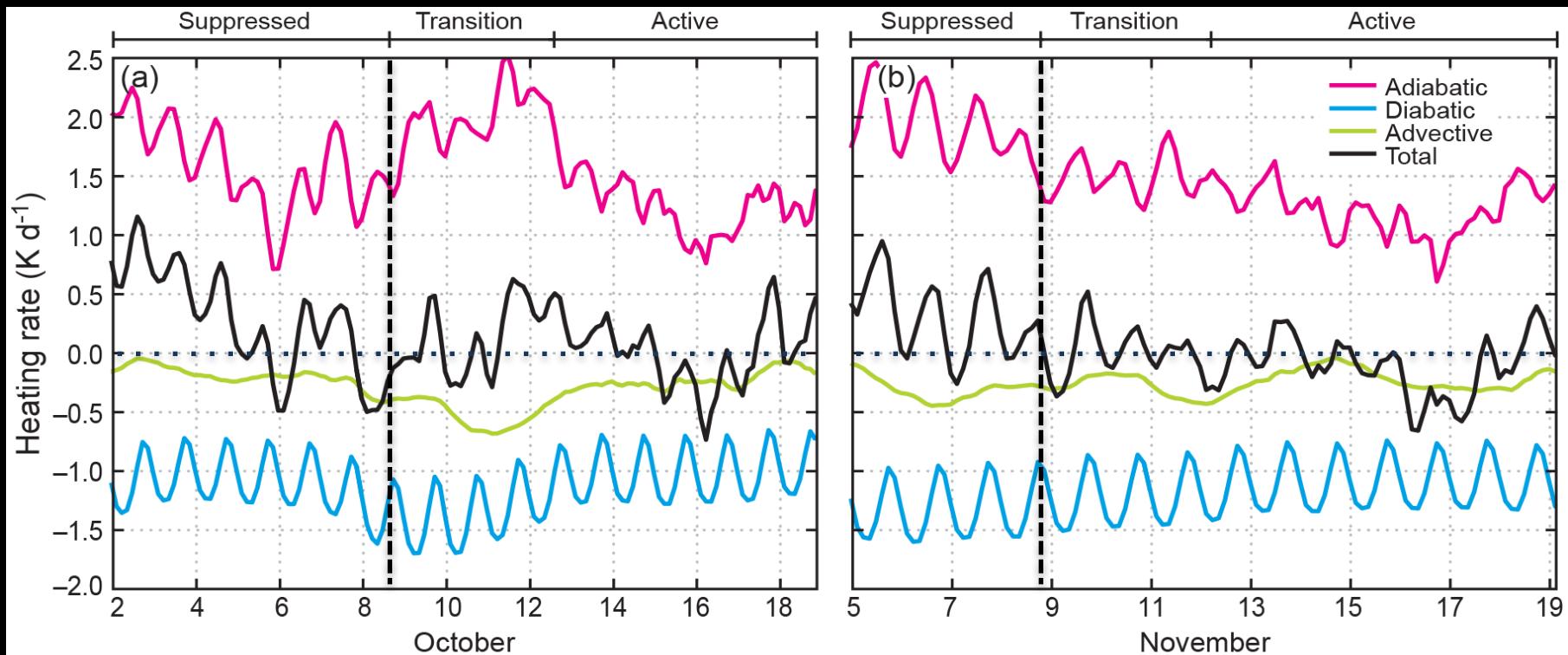
Mean 700–850 mb temperature



Mean 700–850 mb temperature



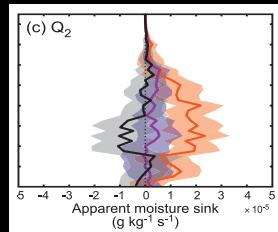
Changes in environmental temperature at start of transition periods are less than 1K!



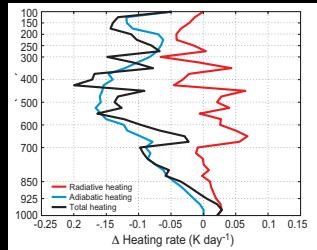
$$\frac{\partial T}{\partial t} = \underbrace{-\mathbf{u}_h \cdot \nabla T}_{\text{advective}} - w \overbrace{\left(\frac{g}{c_p} + \Gamma \right)}^{\text{adiabatic}} + \underbrace{\frac{J}{c_p}}_{\text{diabatic}}$$

Conclusions

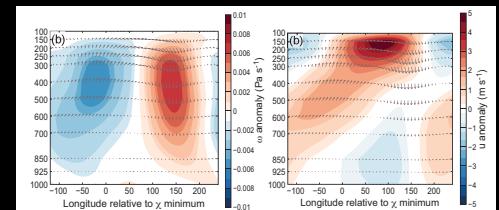
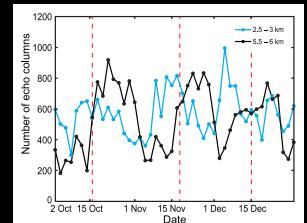
- 3–7 day build up in cloud population during transition periods prior to MJO convective onset.



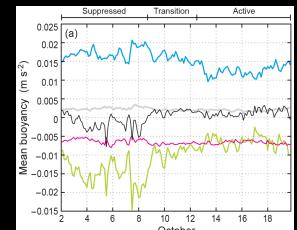
- Circumnavigating wave has impacts on low-wavenumber ω anomalies of $O(0.01 \text{ Pa s}^{-1})$.



- Small changes in environmental temperature dramatically alter mean buoyancy of cloud updrafts in 700–850 hPa layer.



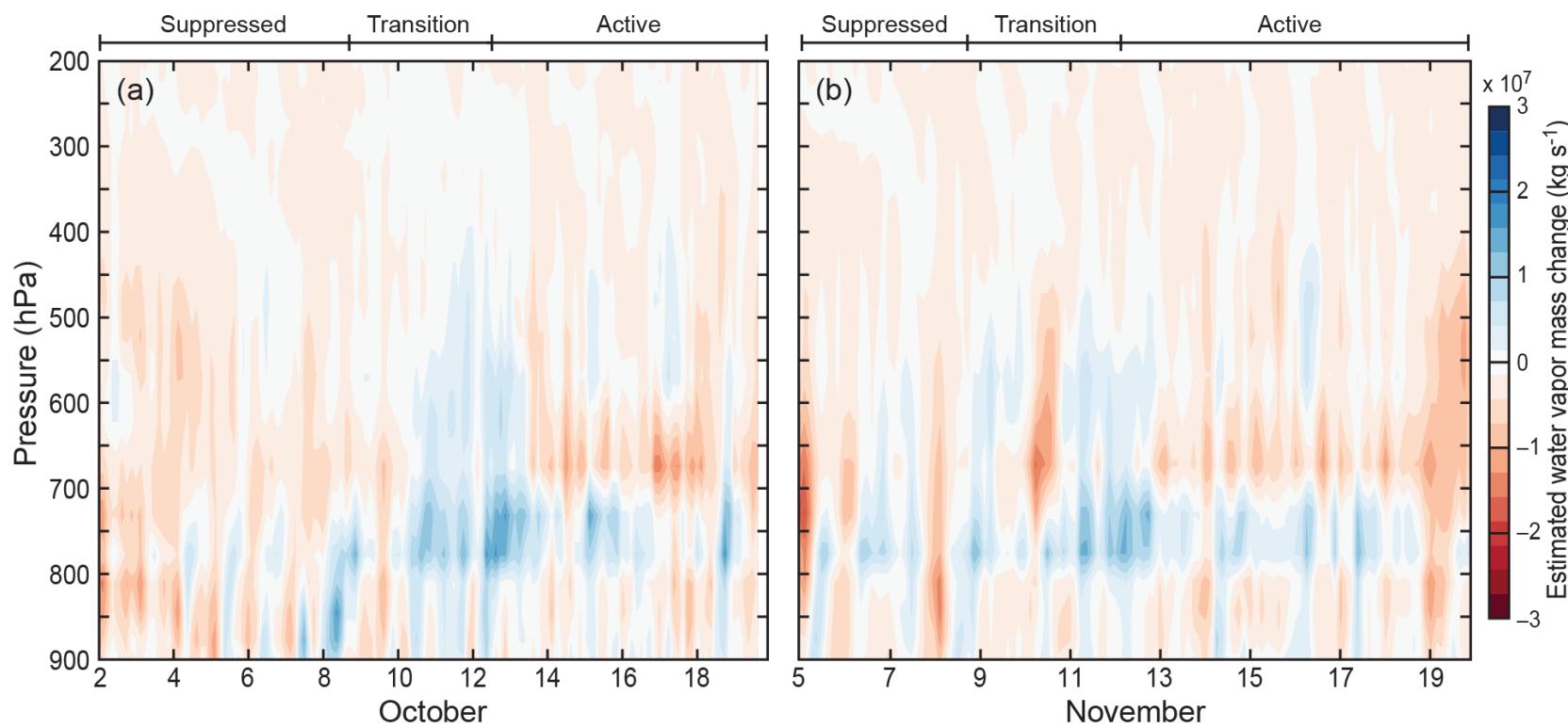
- Changes in vertical velocity cause small changes of $O(0.1\text{K})$ in tropospheric temperature below 500 hPa.

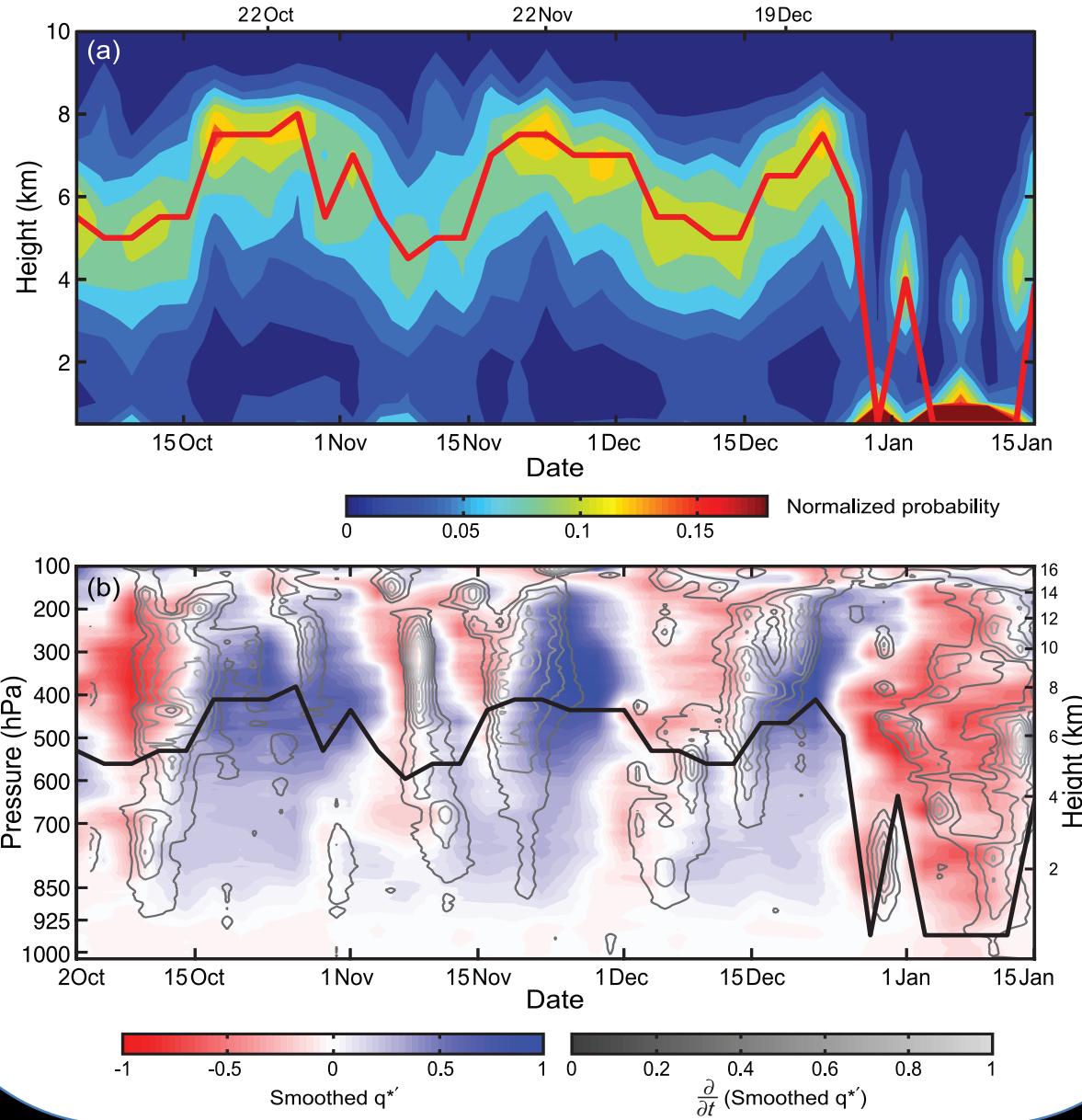


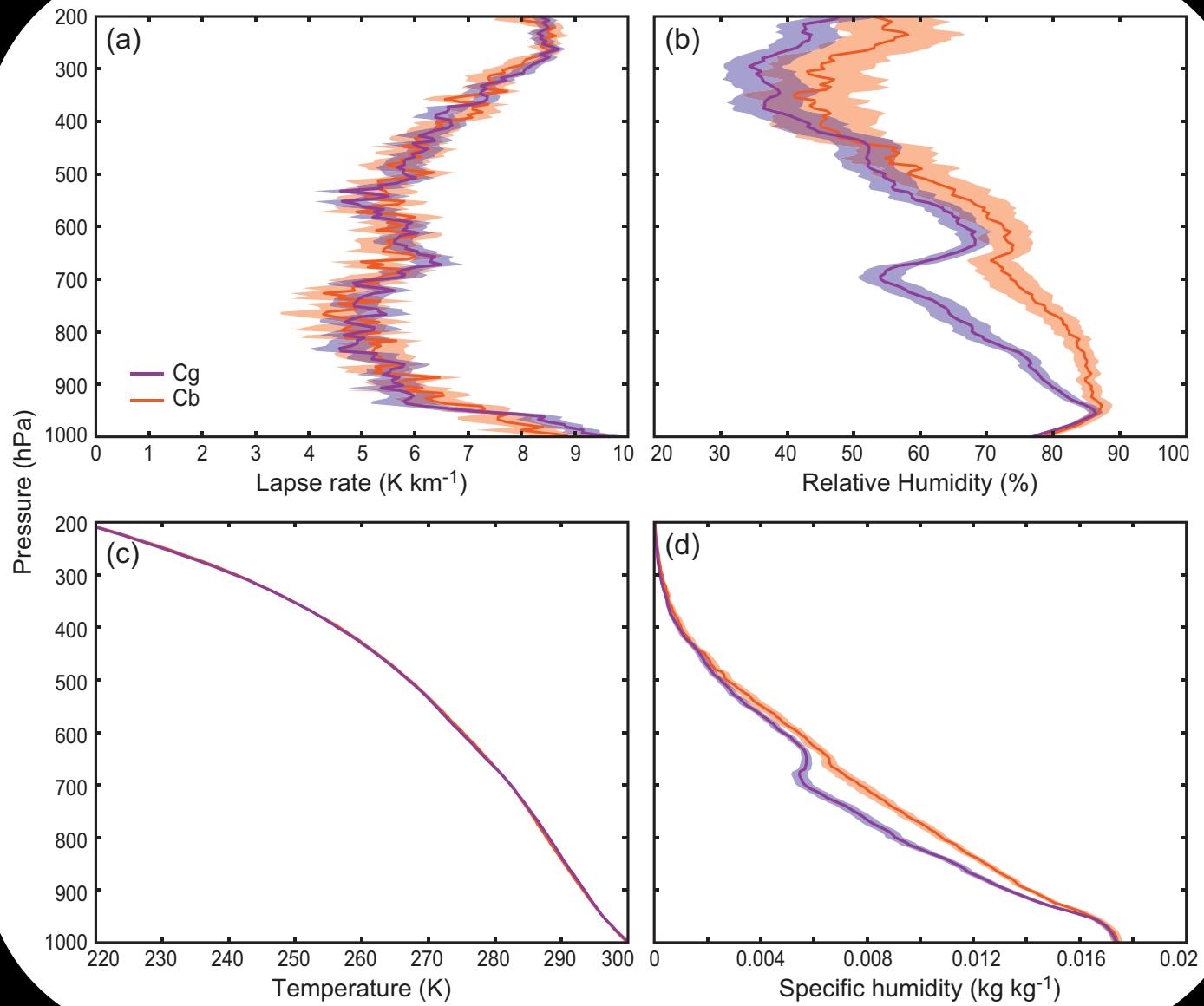
A photograph of a sunset over a calm body of water. The sky is filled with dark, billowing clouds, with patches of orange and yellow light from the setting sun visible on the horizon. The water's surface is very still, creating a clear reflection of the sky and clouds. In the foreground, dark silhouettes of rocks or low-lying land are visible.

End

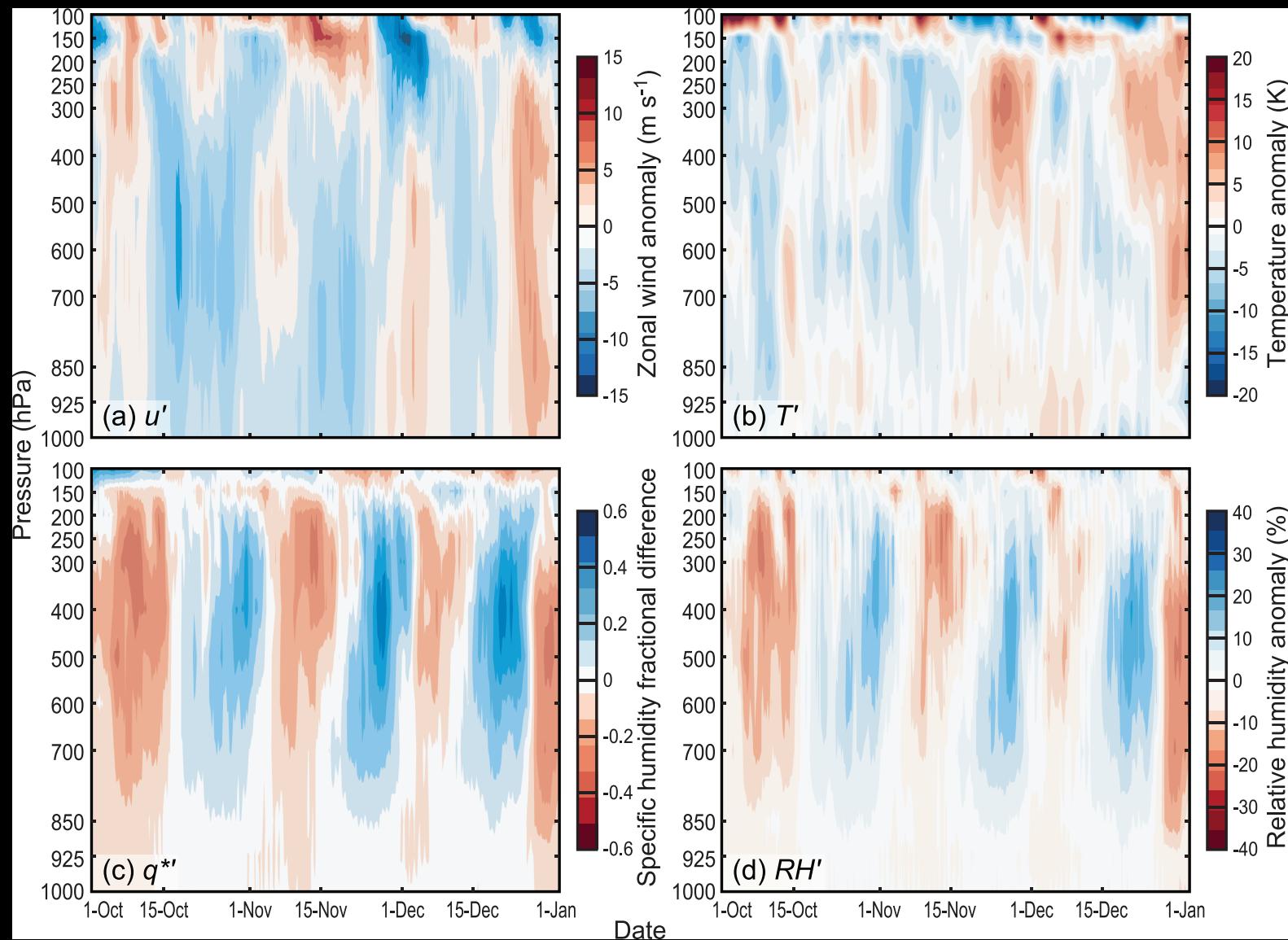
Extra Slides

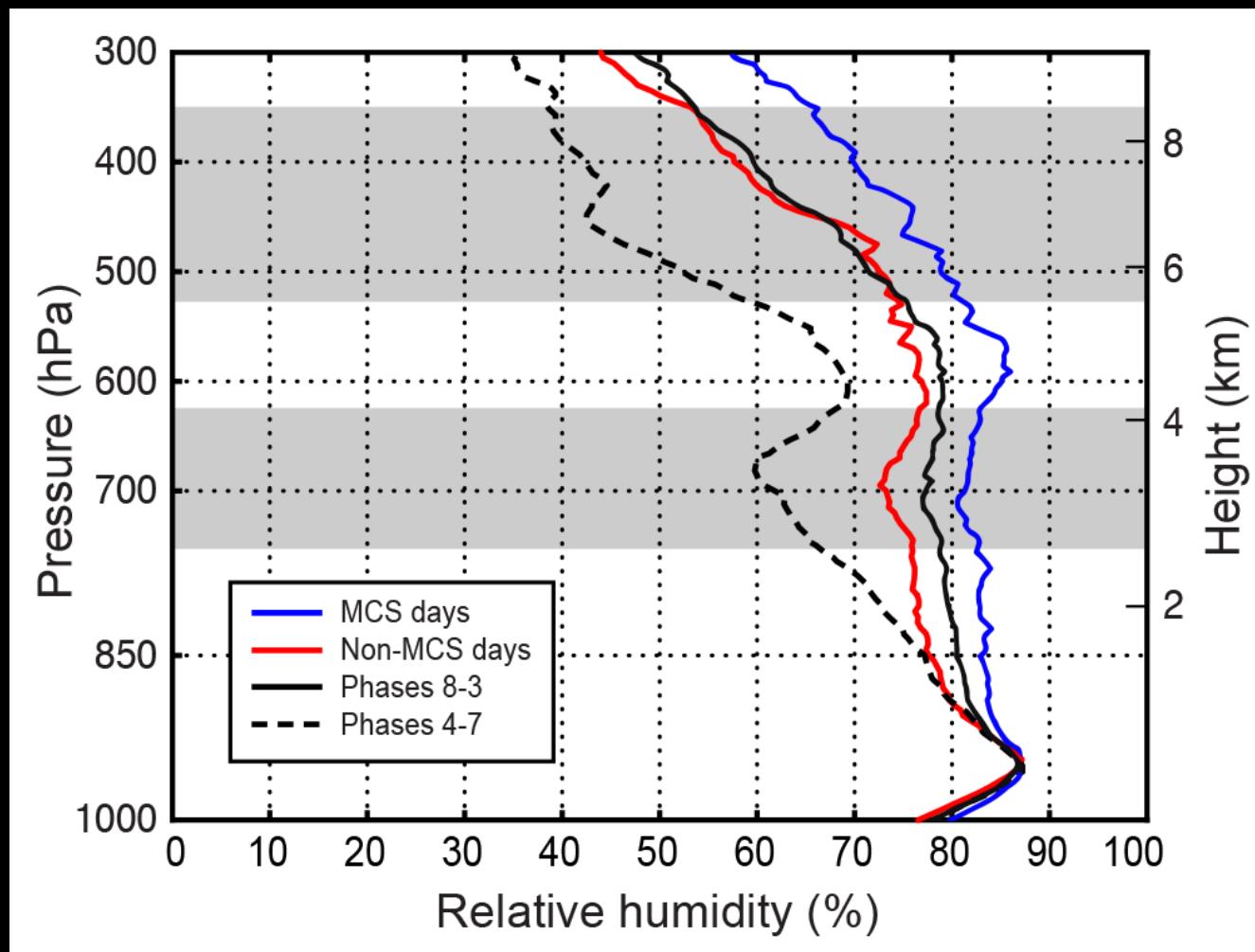






ERA-Interim



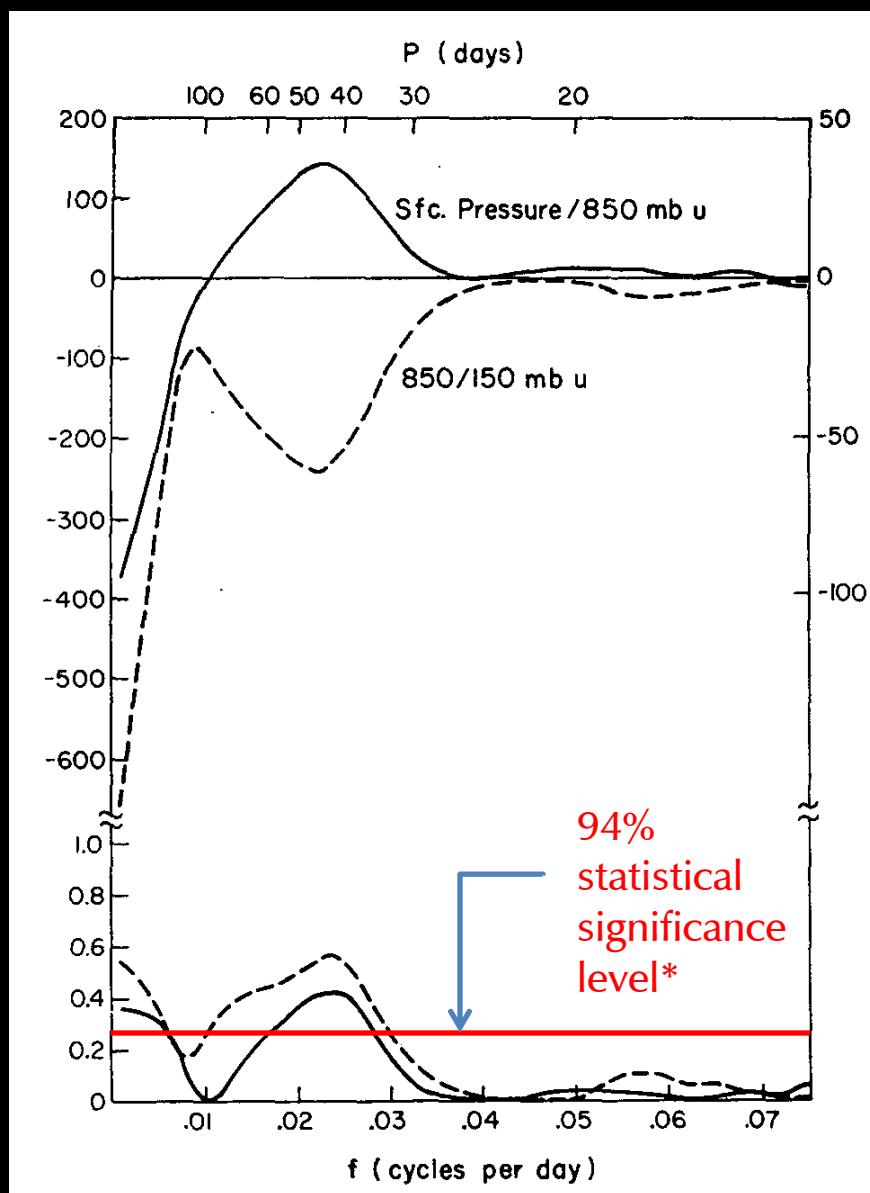


Timescale of MJO Convective Build-up

What duration is the transition from suppressed to widespread, deep convection?

WRF (V3.5.1) Specifications

- 1–20 October and 4–20 November
- ERA-I forcing with NOAA RTG High-Res SST
- 2km grid spacing, 38 vertical levels
- Microphysics: Thompson
- Radiation: RRTMG
- PBL: MYJ
- Monin-Obukhov surface layer physics
- Noah LSM



**A posteriori*. 99.9% if expected *a priori*.

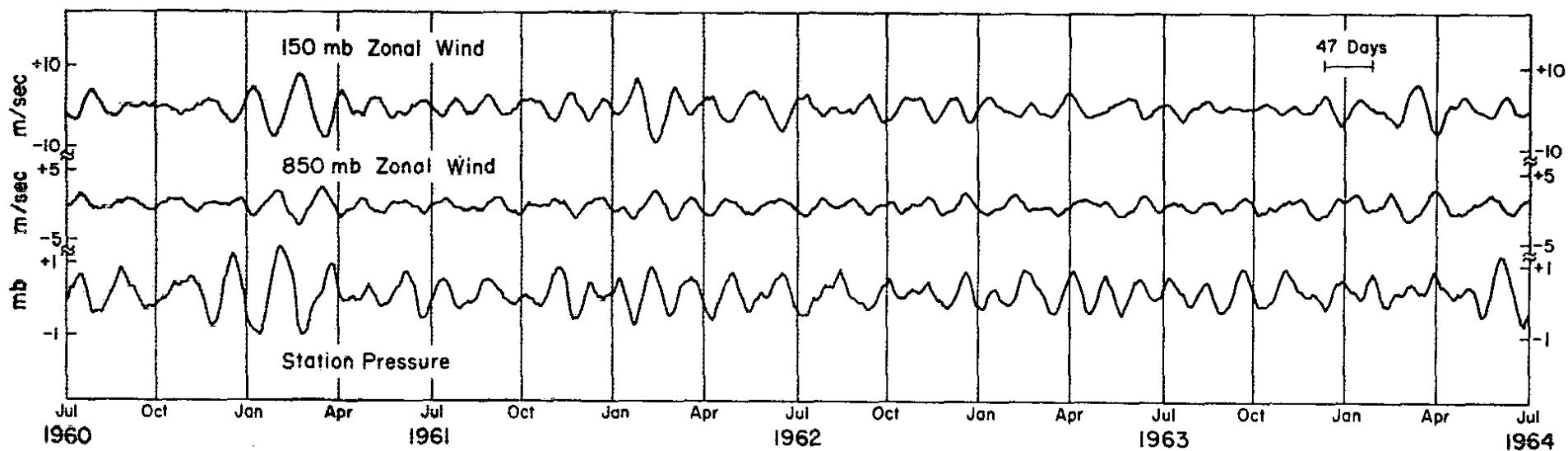
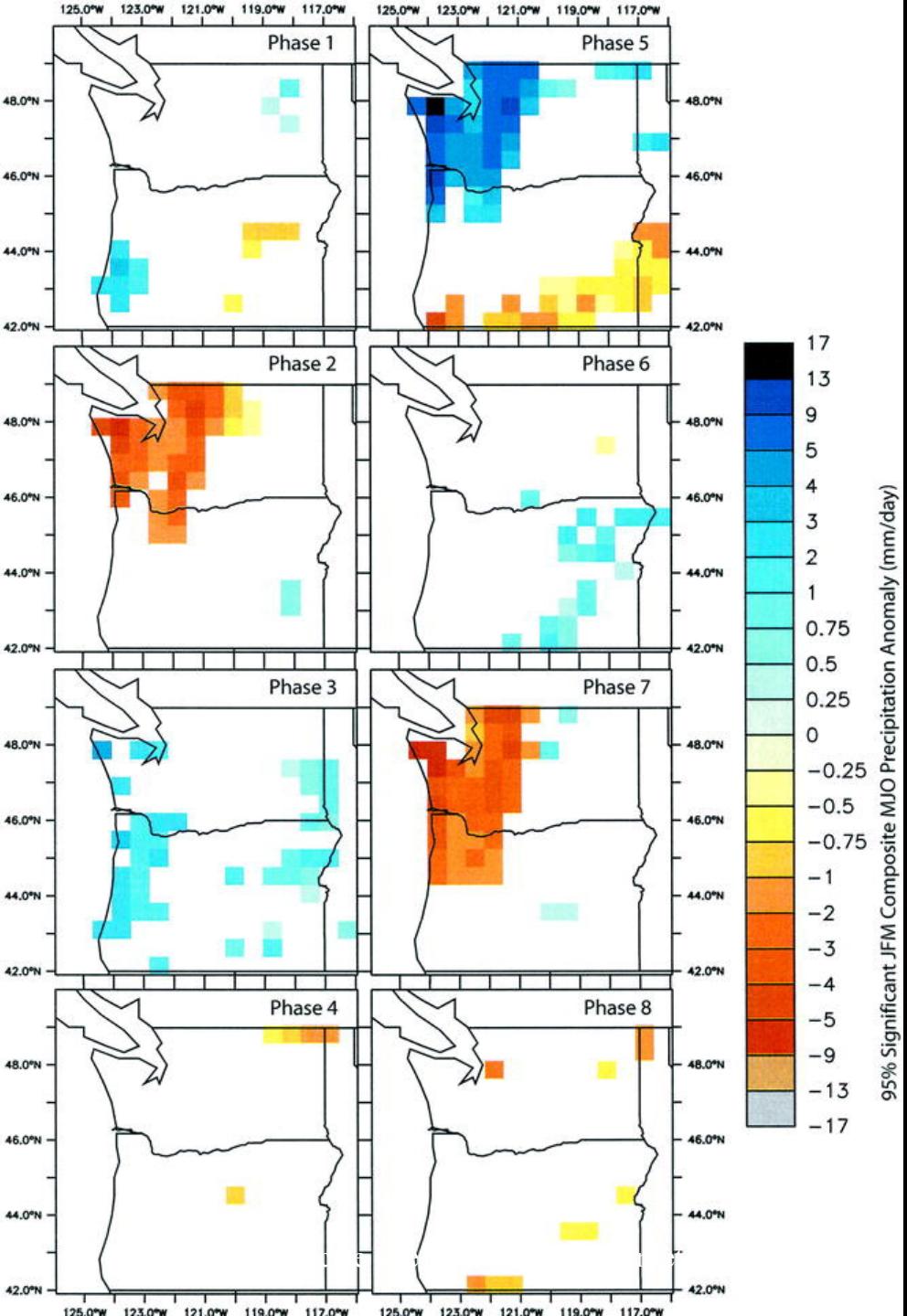
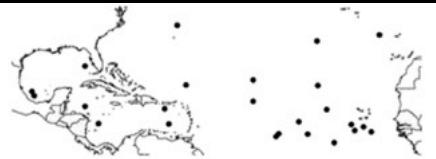


FIG. 5. The 150- and 850-mb u component and station pressure records for Canton Island from July 1960 through June 1964 treated with a 47-day band-pass filter.



Atlantic

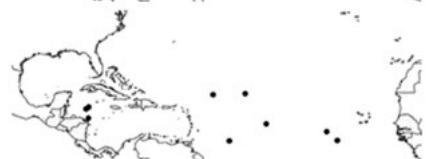
Phase 1

489 Days
23 H

Phase 2

441 Days
32 H

Phase 3

292 Days
9 H

Phase 4

301 Days
19 H

Phase 5

456 Days
16 H

Phase 6

399 Days
13 H

Phase 7

271 Days
5 H

Phase 8

273 Days
2 H**W. Pacific**

Phase 1

579 Days
42 H

Phase 2

524 Days
33 H

Phase 3

347 Days
12 H

Phase 4

379 Days
18 H

Phase 5

527 Days
49 H

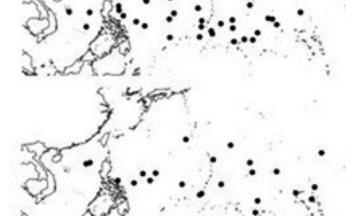
Phase 6

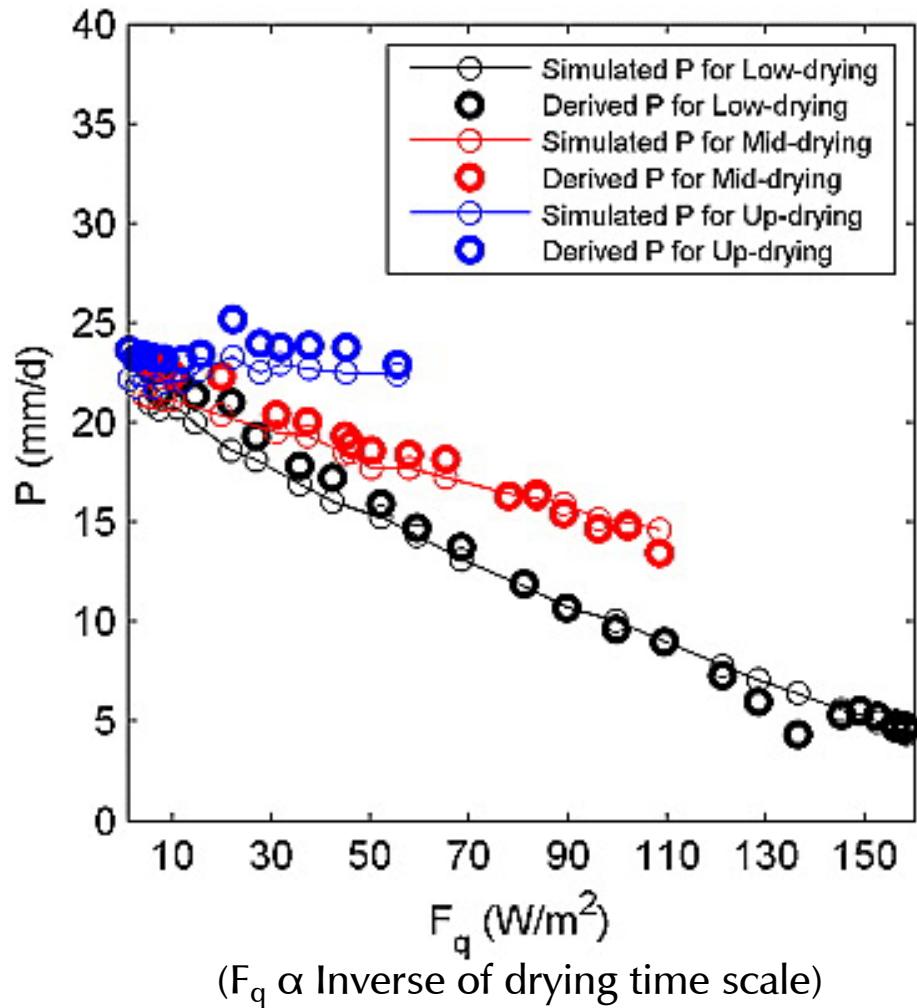
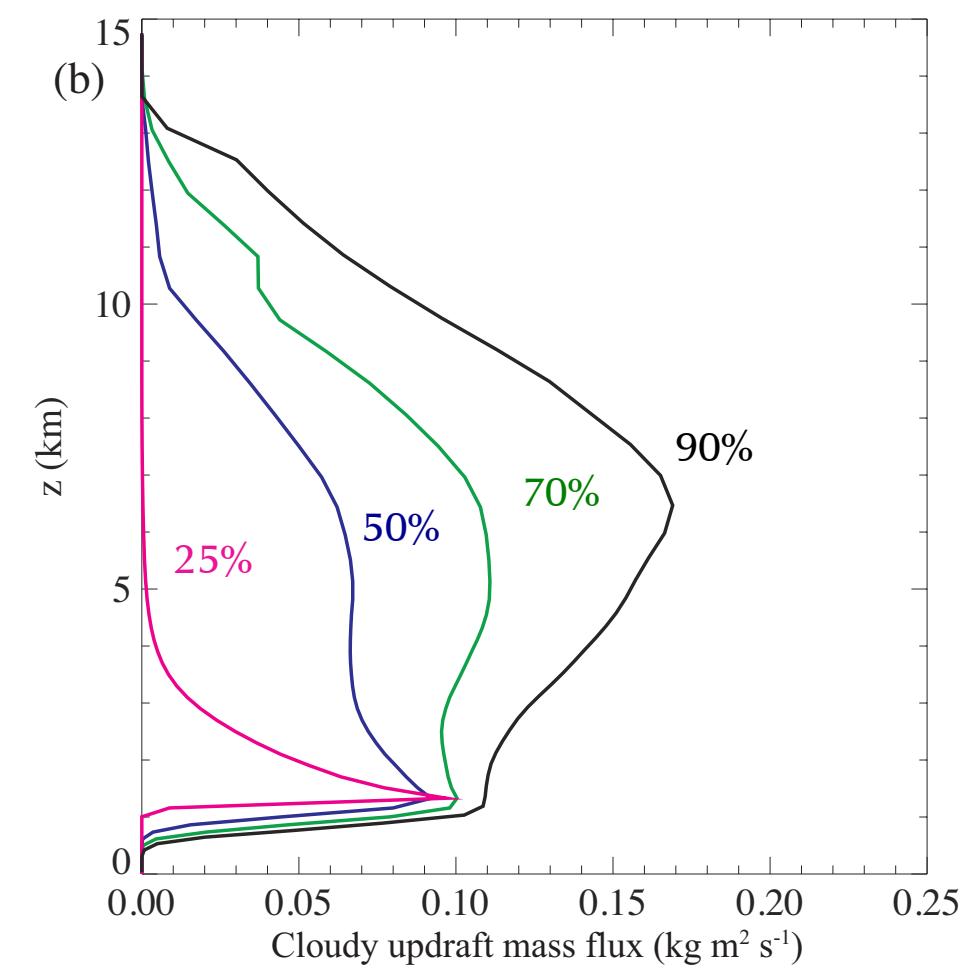
482 Days
57 H

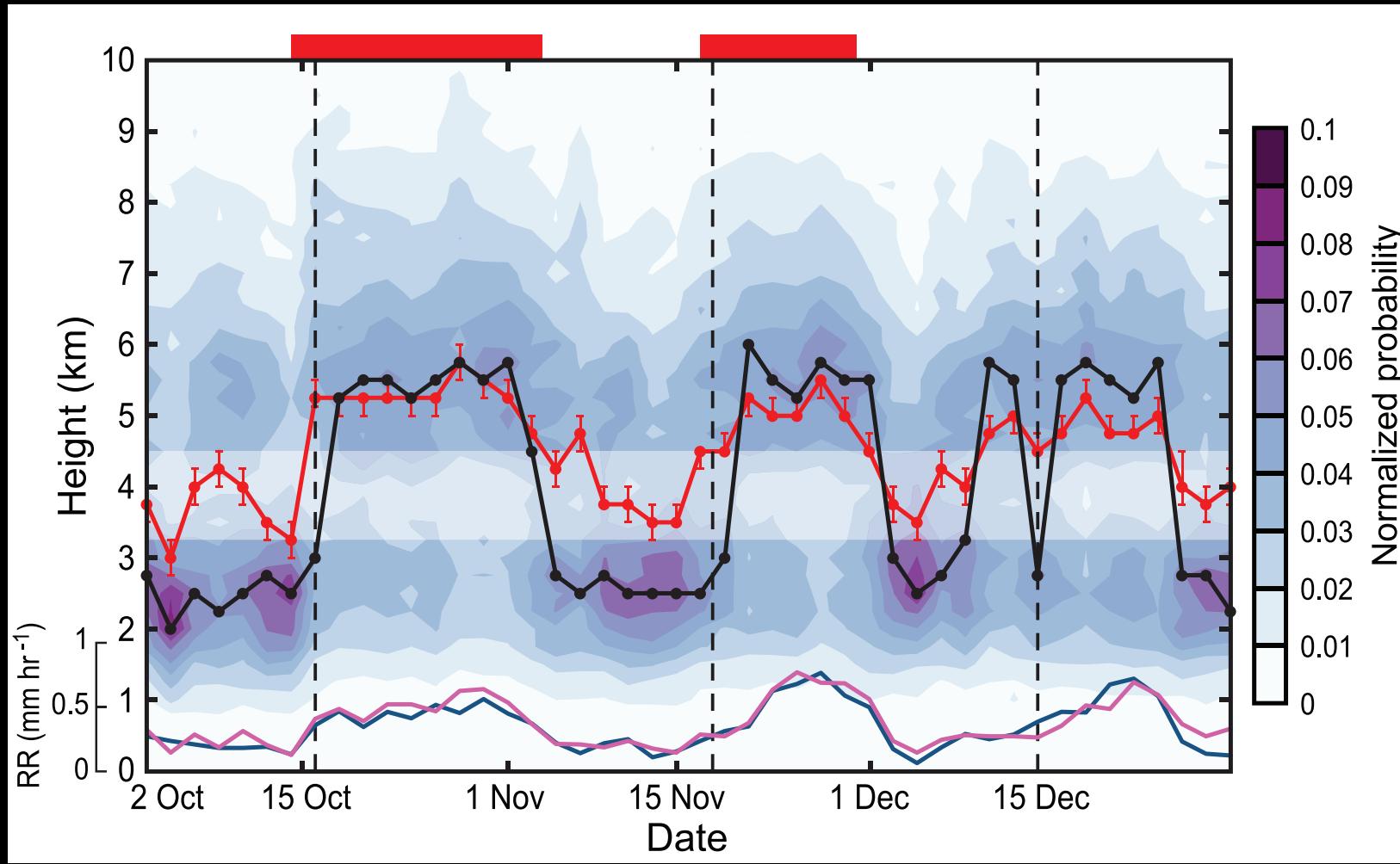
Phase 7

316 Days
35 H

Phase 8

349 Days
23 H





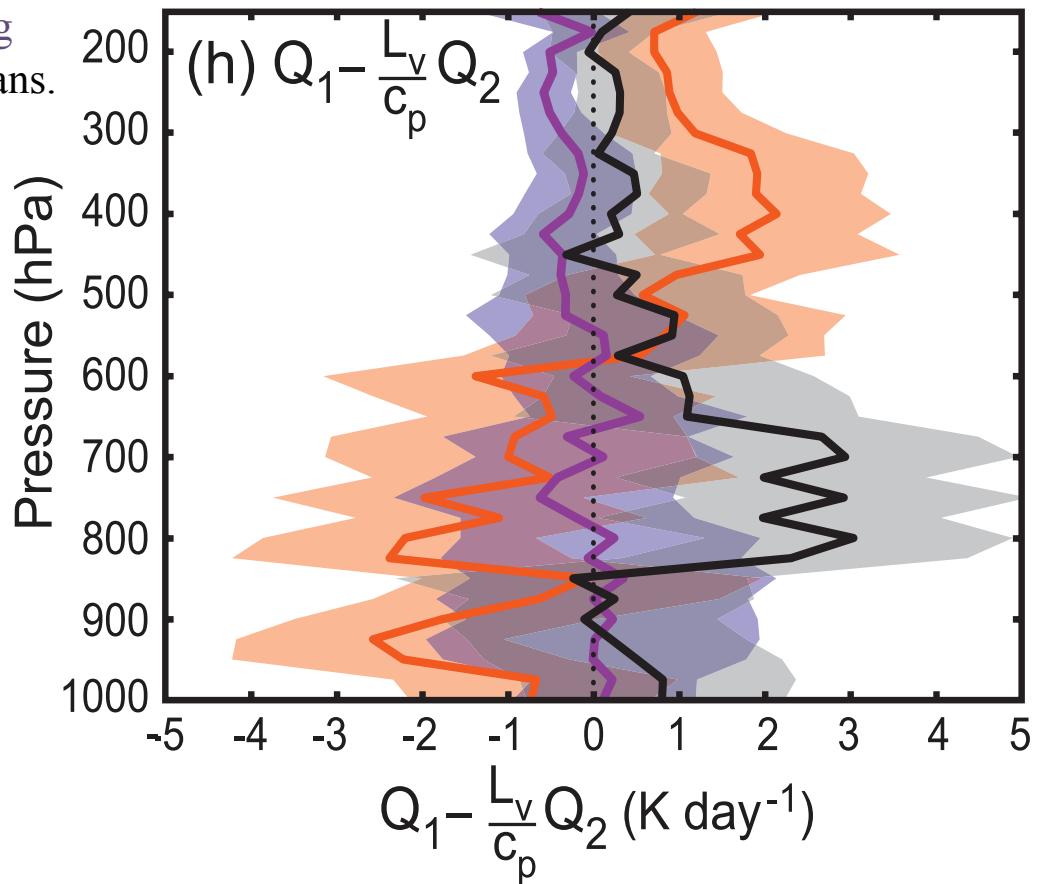
$$\frac{\partial q}{\partial t} = \mathbf{v}_h \cdot \nabla q + \omega \frac{\partial q}{\partial p} + Q_2$$

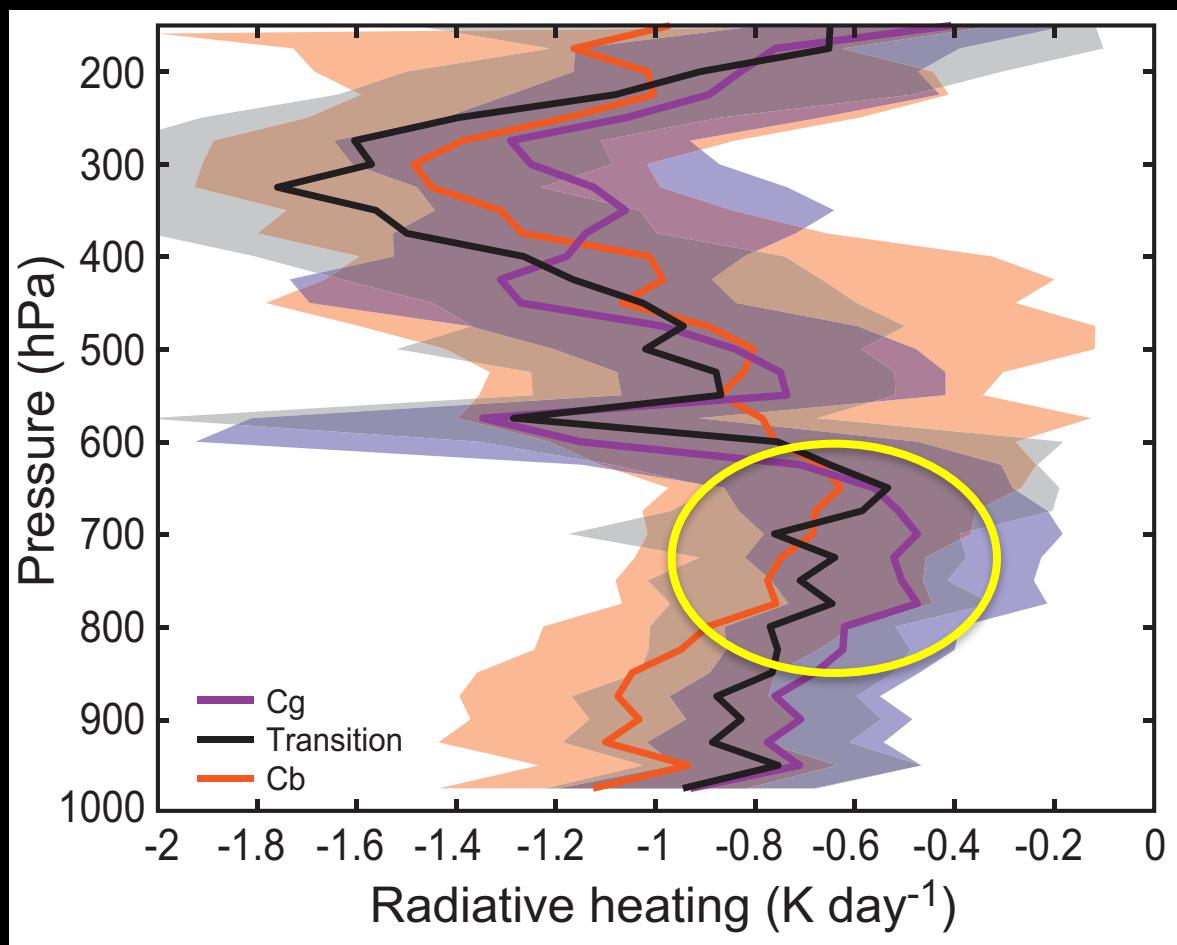
$$Q_2 = (\bar{c} - \bar{e}) + \frac{\partial}{\partial p} (\overline{\omega' q'})$$

$$Q_1 = Q_R + \frac{1}{c_p} \left[L_v (\bar{c} - \bar{e}) + \frac{\partial}{\partial p} (\overline{\omega' s'}) \right]$$

$$Q_1 - \frac{L_v}{c_p} Q_2 = Q_R - \frac{1}{c_p} \frac{\partial}{\partial p} (\overline{\omega' h'})$$

Purple = Cg
 Black = Trans.
 Red = Cb





$$\frac{\partial q}{\partial t} = \mathbf{v}_h \cdot \nabla q + \omega \frac{\partial q}{\partial p} + Q_2$$

$$Q_2 = (\bar{c} - \bar{e}) + \frac{\partial}{\partial p} (\overline{\omega' q'})$$

$$Q_1 = Q_R + \frac{1}{c_p} \left[L_v (\bar{c} - \bar{e}) + \frac{\partial}{\partial p} (\overline{\omega' s'}) \right]$$

$$Q_1 - \frac{L_v}{c_p} Q_2 = Q_R - \frac{1}{c_p} \frac{\partial}{\partial p} (\overline{\omega' h'})$$

Purple = Cg
 Black = Trans.
 Red = Cb

