

# What Controls Timing of Deep Convective Onset in the Tropics? *(Hint: Much more than only an environmental thermodynamic profile)*

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# How do we define buoyancy?

$$B \approx g \left( \frac{\theta^*}{\theta_0} + \left( \frac{R_v}{R_d} - 1 \right) q_v^* - q_{lf} \right)$$

$$B \equiv g \left( \frac{T'_v}{\bar{T}_V} \right)$$

$$\frac{Dw}{Dt} \approx - \frac{1}{\rho} \frac{\partial p'_D}{\partial z} \boxed{- \frac{1}{\rho} \frac{\partial p'_B}{\partial z} + B}$$

Vertical Pressure  
Gradient  
Accelerations

“Effective buoyancy”  
Davies-Jones (2003)  
Doswell and Markowski (2004)

# How do we define buoyancy?

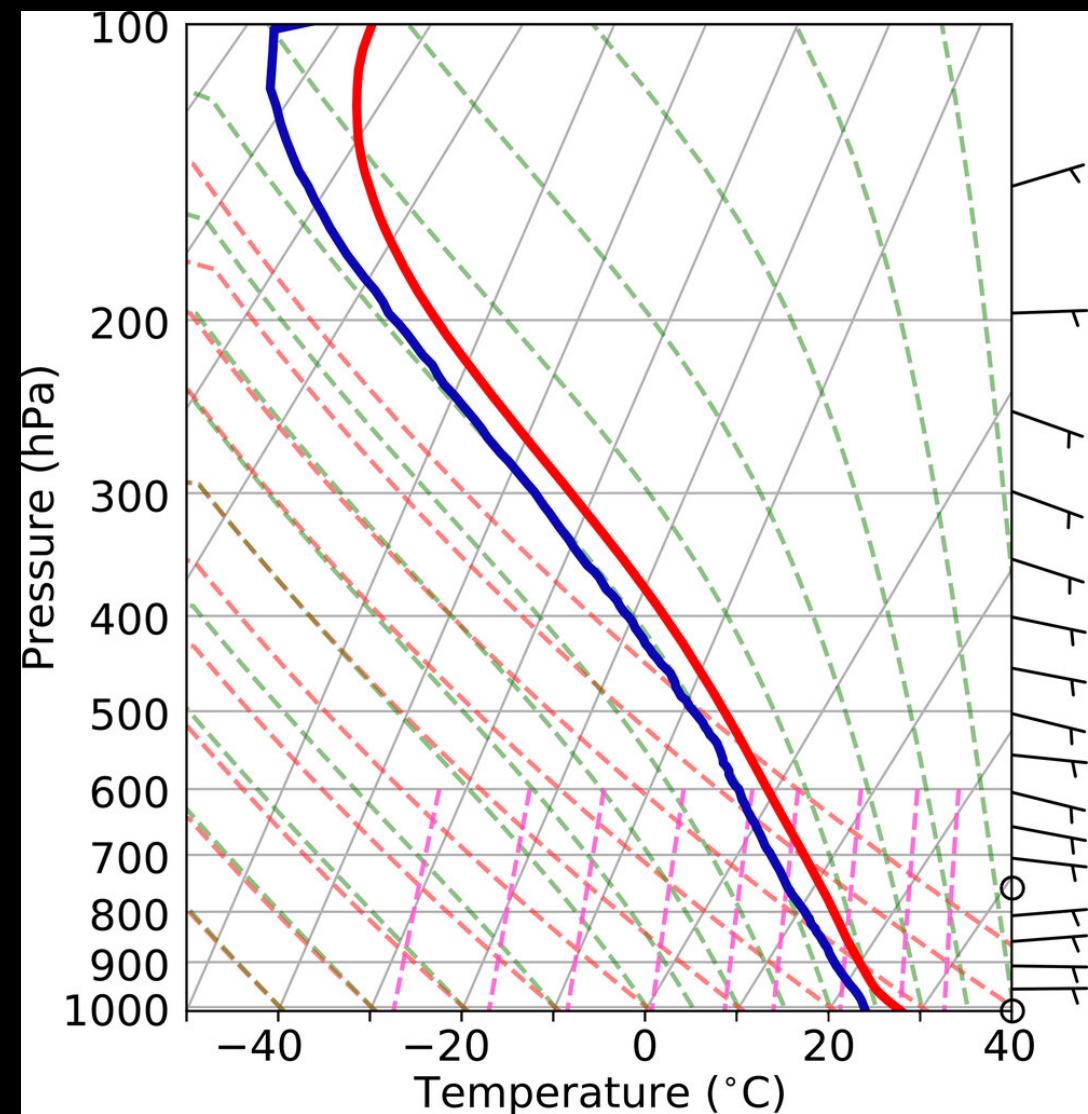
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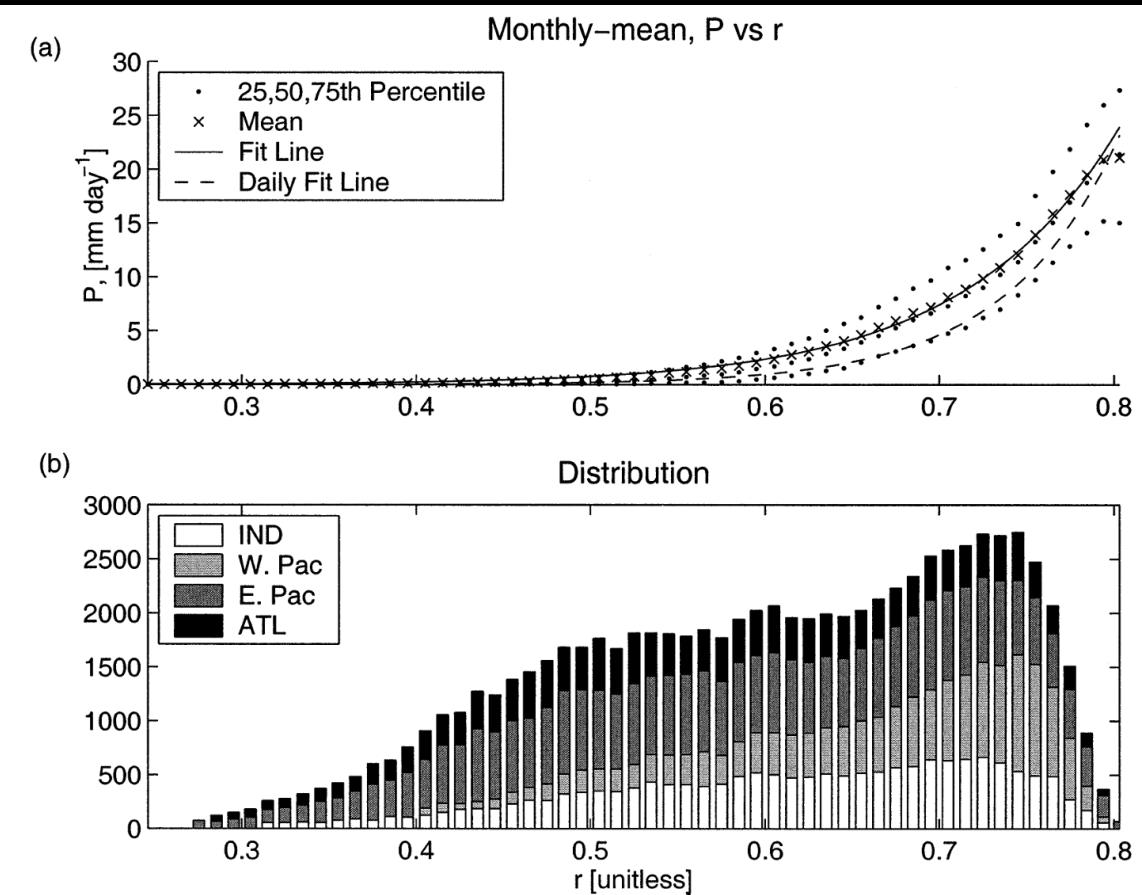
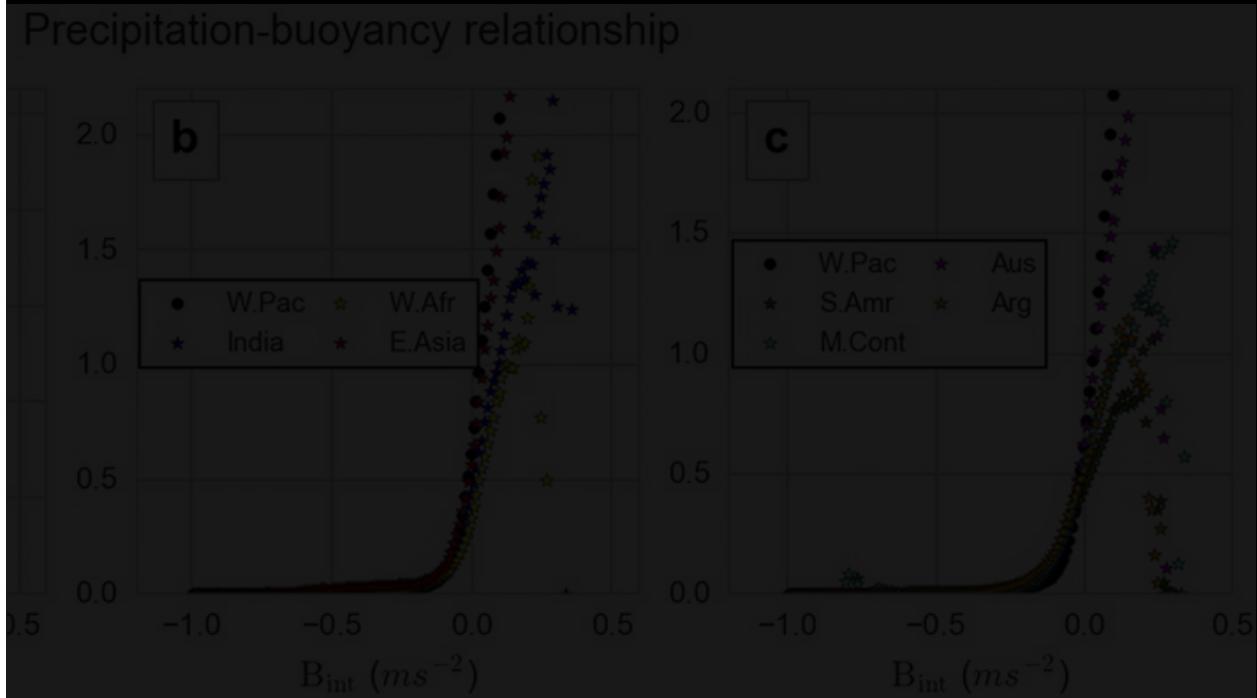


FIG. 6. (a) Distribution of monthly precipitation  $P$  in 1% bins of column-relative humidity  $r$  for all tropical ocean grid points in all months of 1998–2001. Dots show the 25th, 50th, and 75th percentiles of precipitation in each bin. The Xs show the bin-mean precipitation. The solid curve is the exponential fit (2) to the monthly data. For comparison, the dashed curve is the daily mean exponential fit (1). (b) Number of observations in each bin in the four regions.

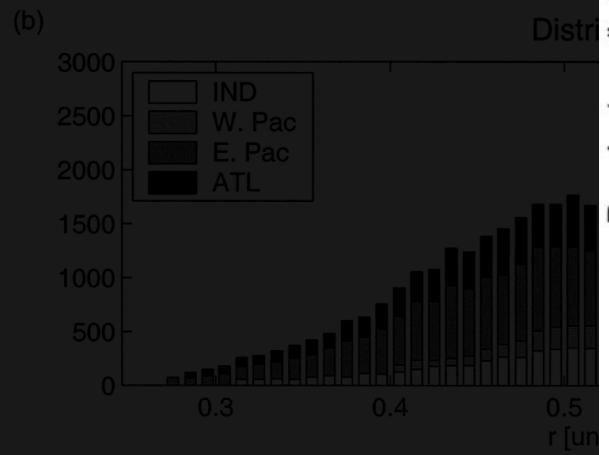
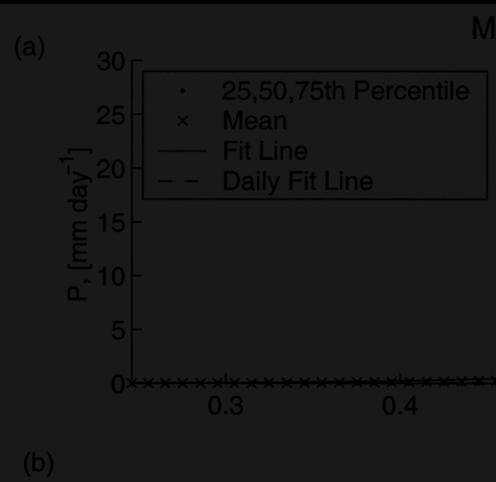
Precipitation vs. tropospheric water vapor

Ahmed and Neelin (2018)



Precipitation vs. lower-tropospheric buoyancy

Bretherton et al. (2004)



Monthly-mean,  $P$  vs  $r$

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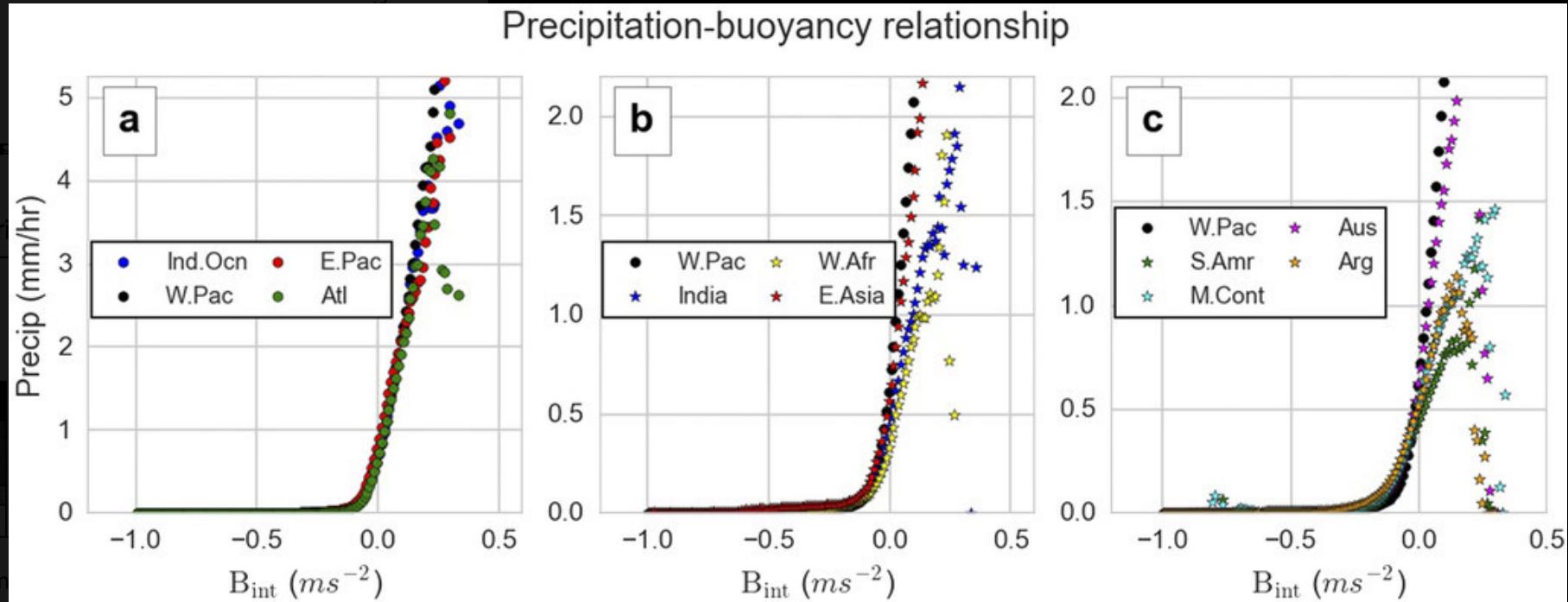
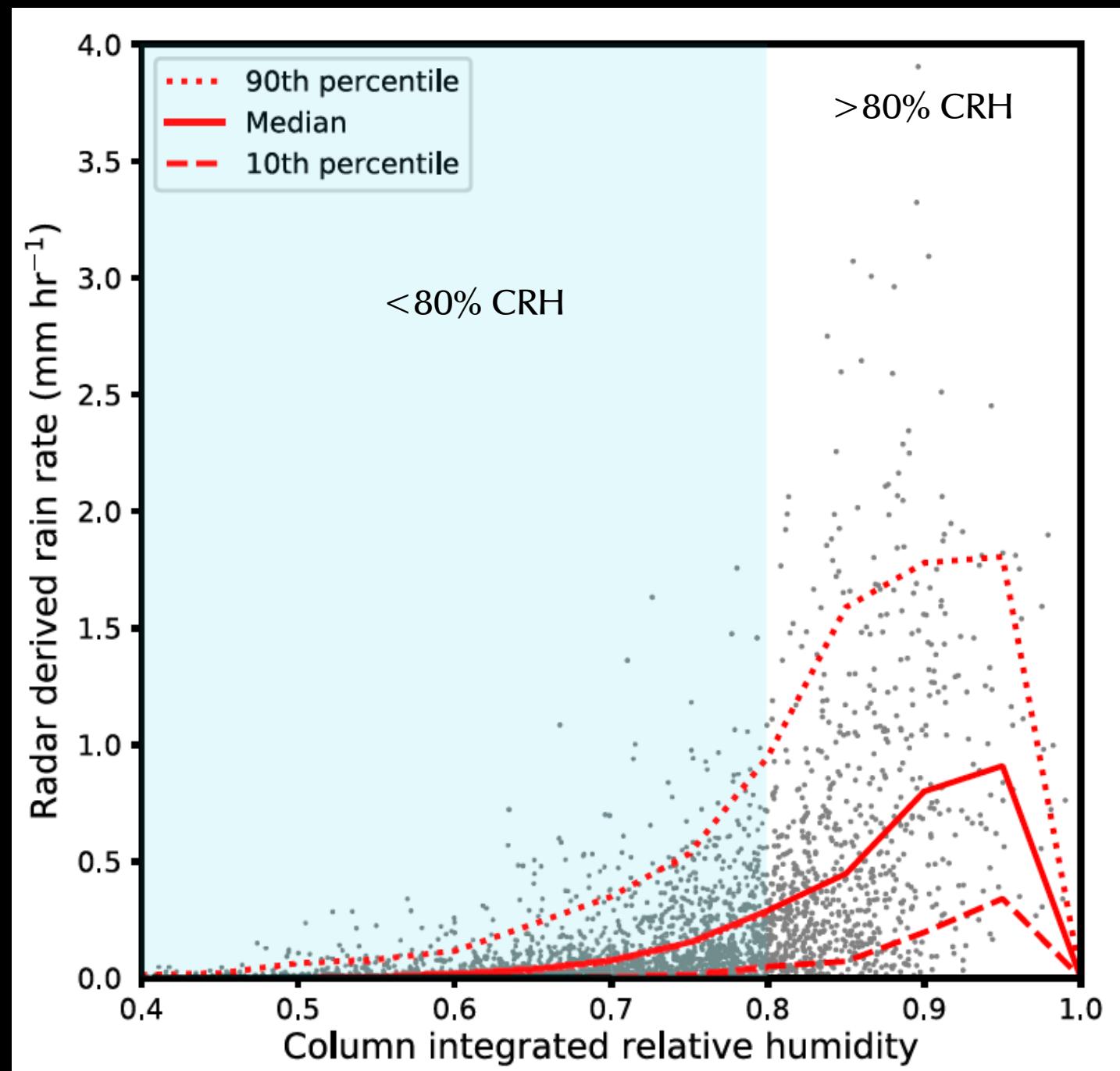
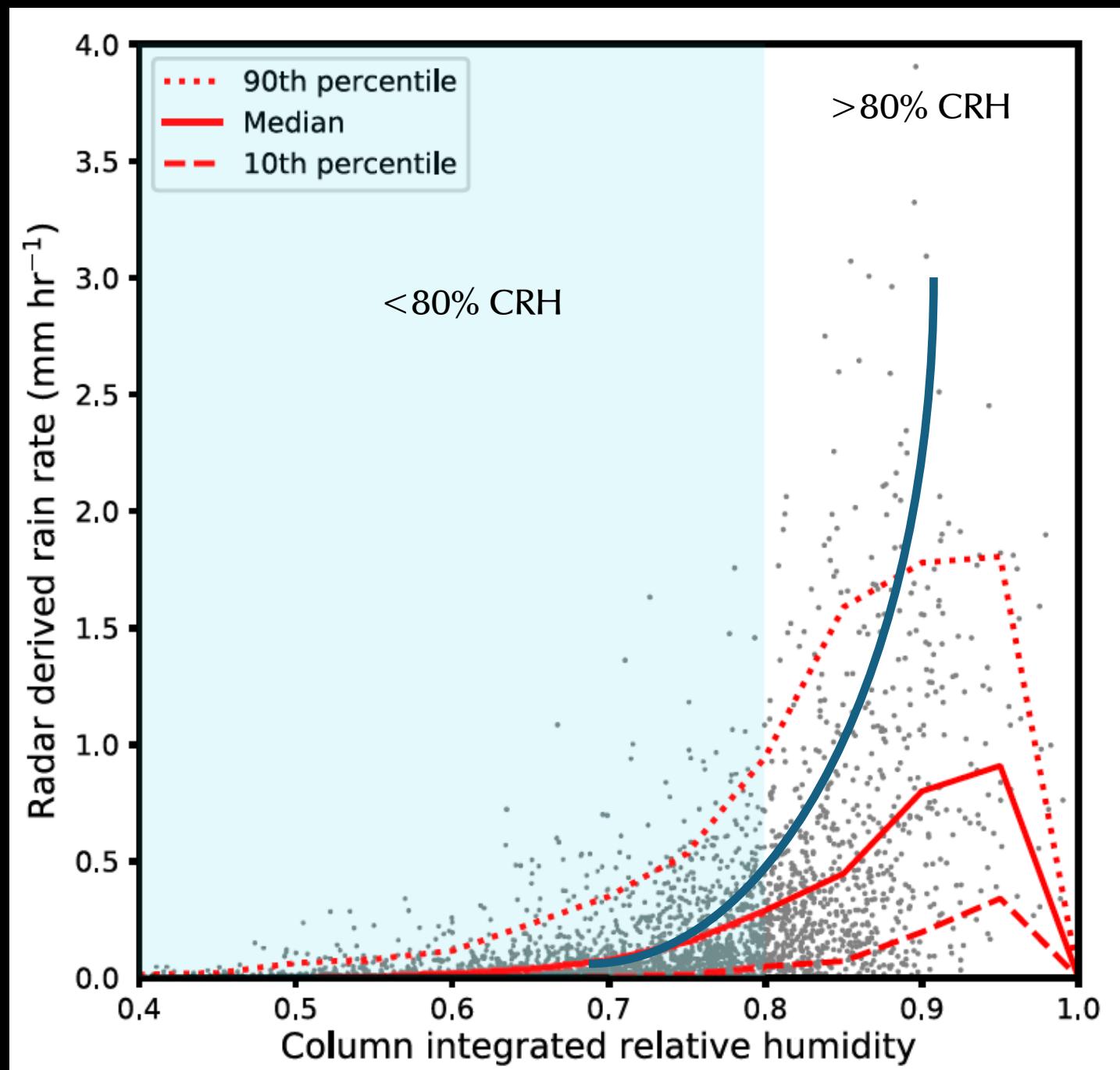


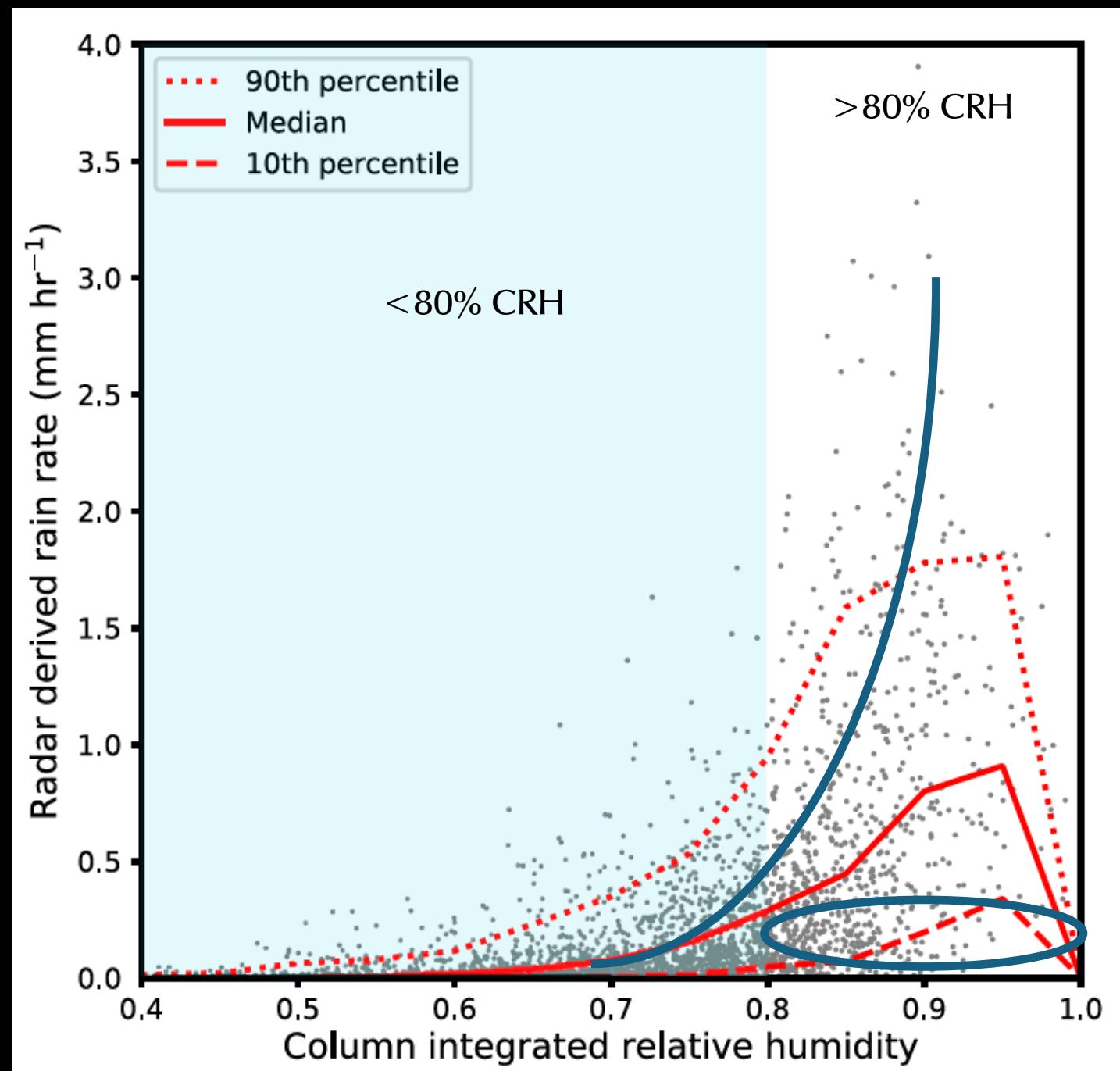
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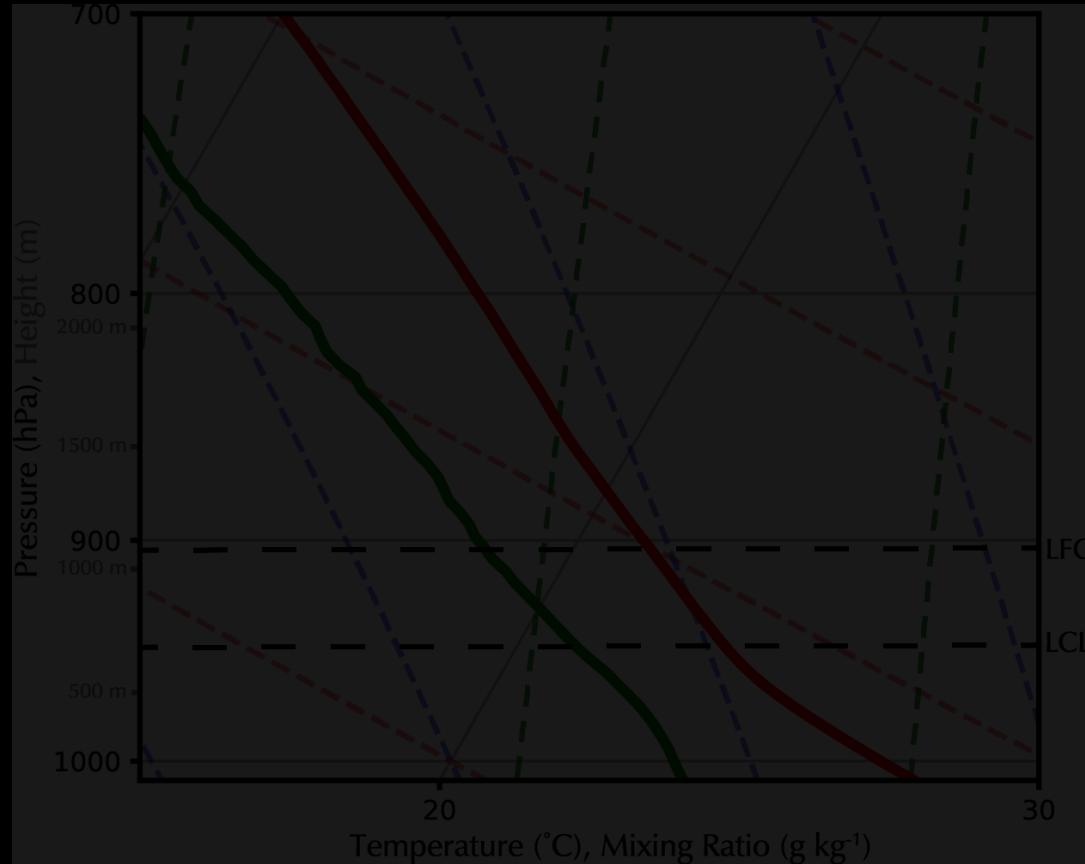
### Precipitation vs. lower-tropospheric buoyancy

Precipitation vs. tropospheric water vapor









## Lagrangian analysis of “parcel” trajectories emerging from the boundary layer in CM1 LES

- 10,240,000 parcel trajectories analyzed
- Over 5 Tb of output for just one simulation
- Trajectories were split into two categories based on whether they entered convection that grew deep or did not:
  - Growers: Trajectories reach at least 6000 m.
  - Non-growers: Stopped between 1500 m and 3000 m.

### Cloud Model 1 (CM1) configuration:

Domain size: 64 km x 64 km x 20 km

Horizontal grid spacing: 100 m

Vertical grid spacing: 50 m in boundary layer, stretched to 250 m above 3500 m.

Time step: 1 s

Output written: every minute

Boundary conditions: Periodic

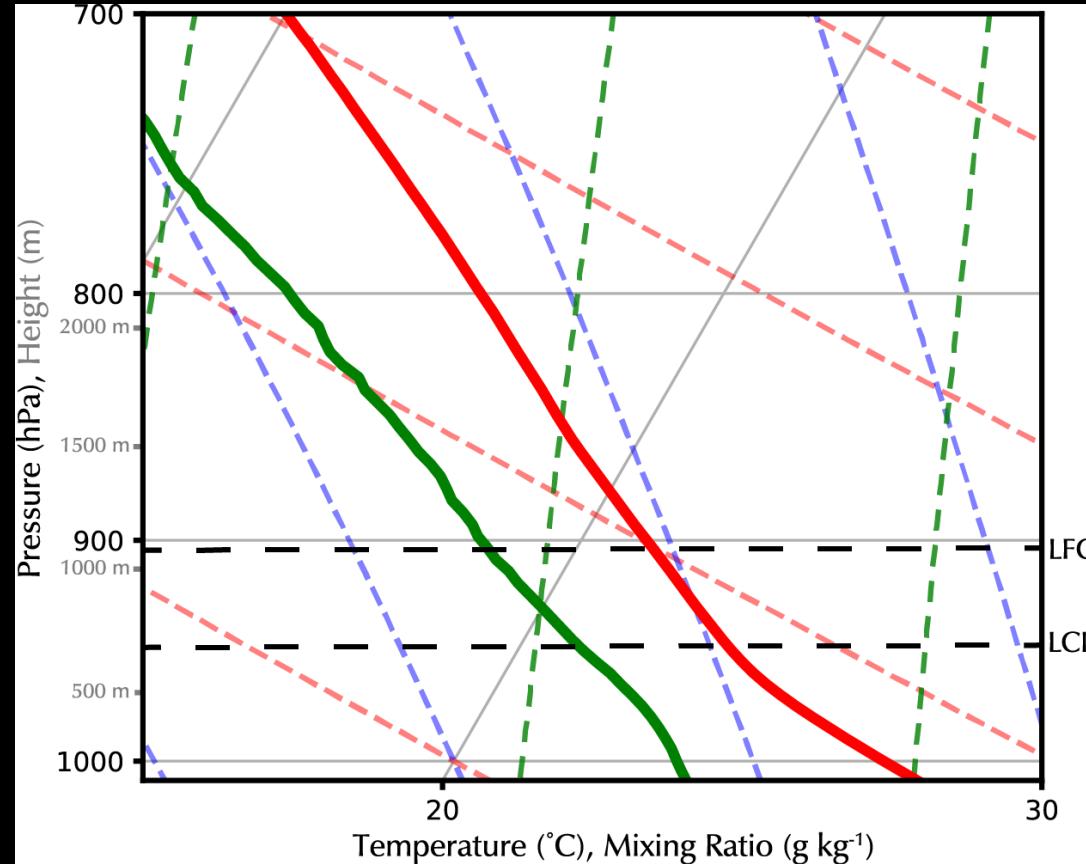
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Microphysics: Morrison

Radiation: RRTM-G (shortwave turned off)

*No cumulus or boundary layer parameterization*

Turbulence promoted by random temperature perturbations up to 0.25 K.



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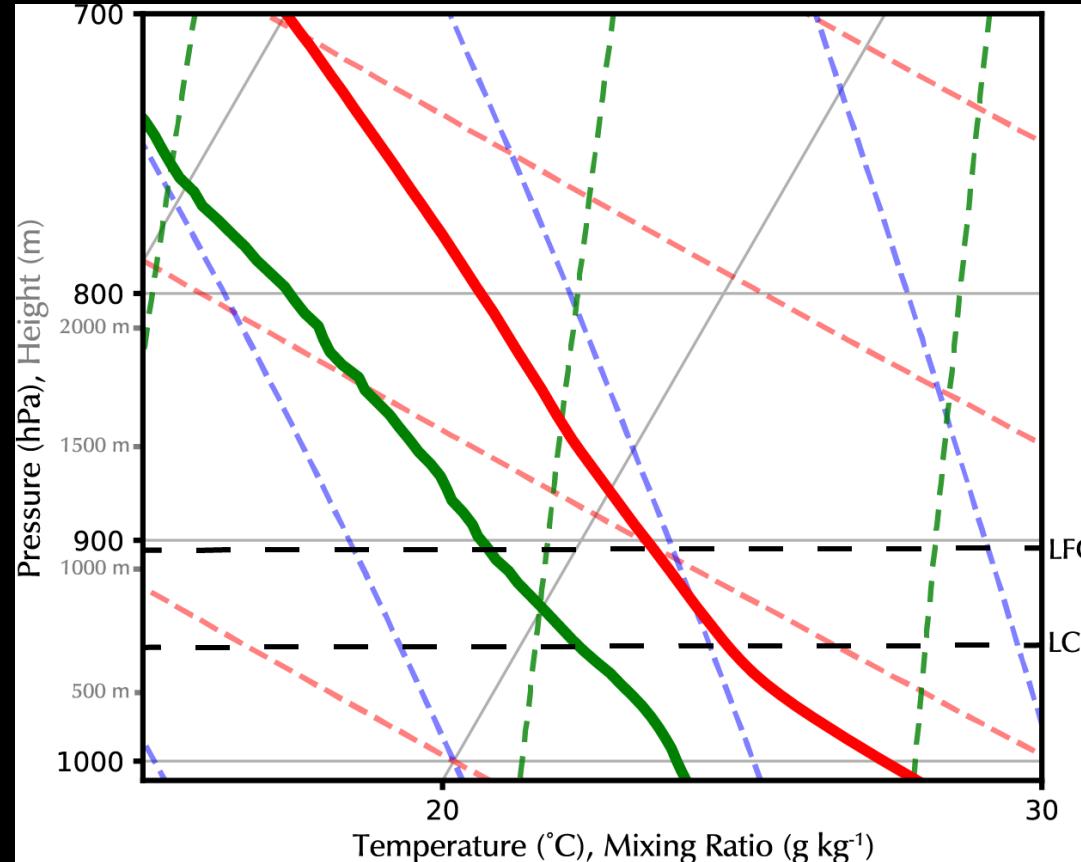
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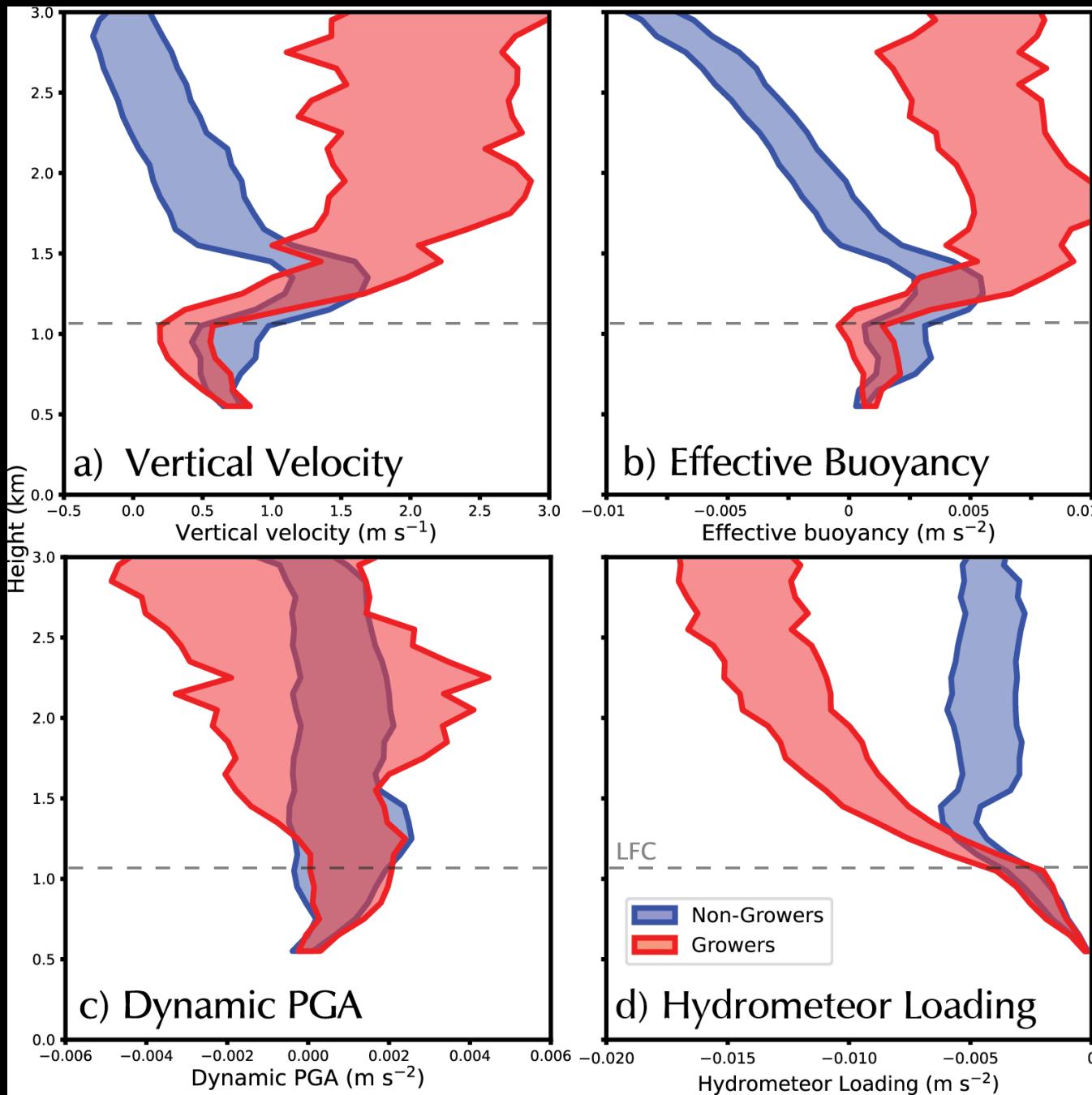
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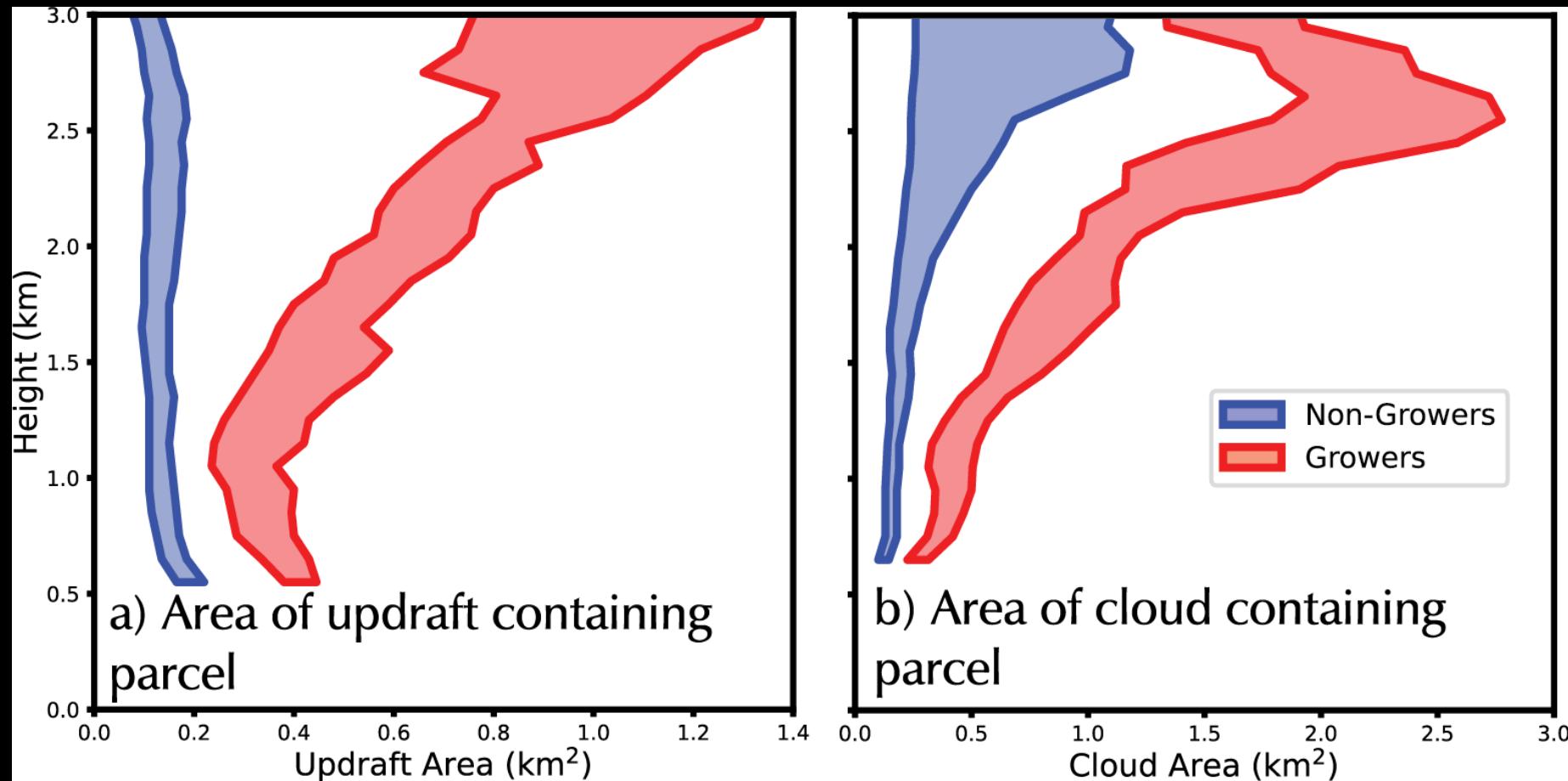
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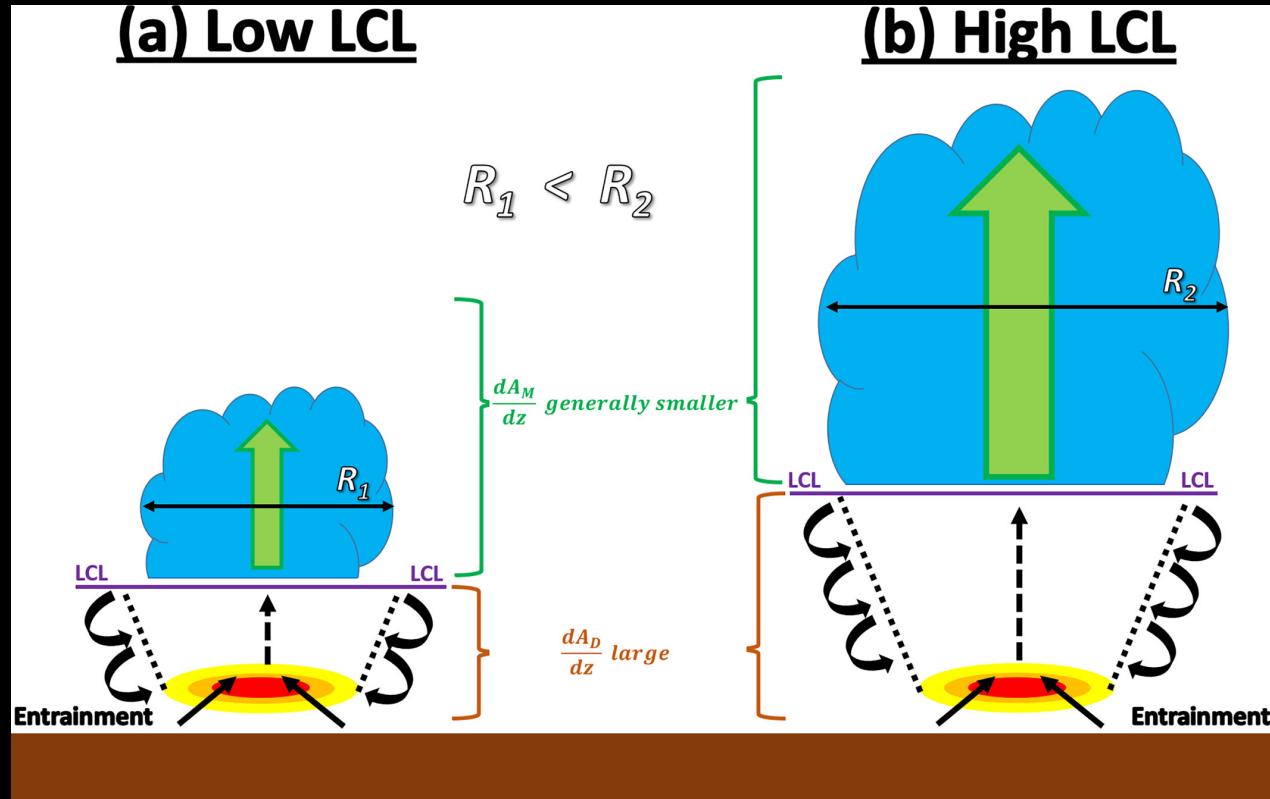
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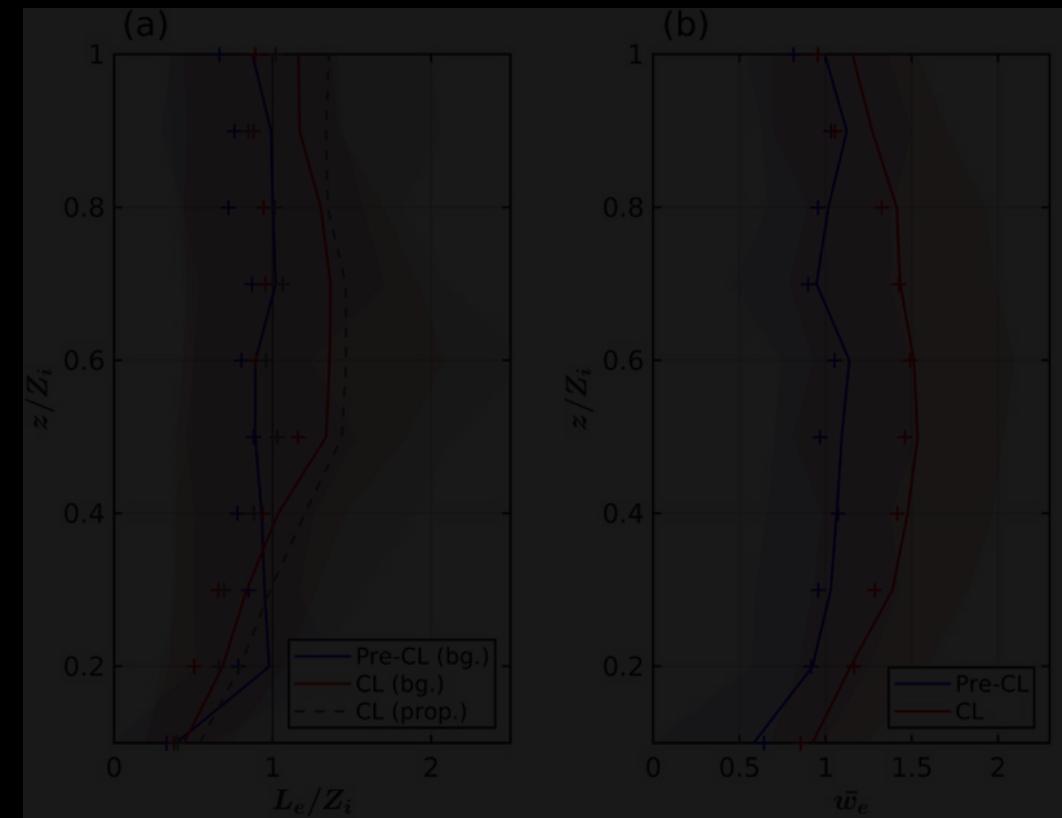
*Shaded regions indicated 95% confidence interval of median.*

*Why are some updrafts wider than others?*

Mulholland et al. (2021): LCL height



Liu et al. (2023): Convergence lines

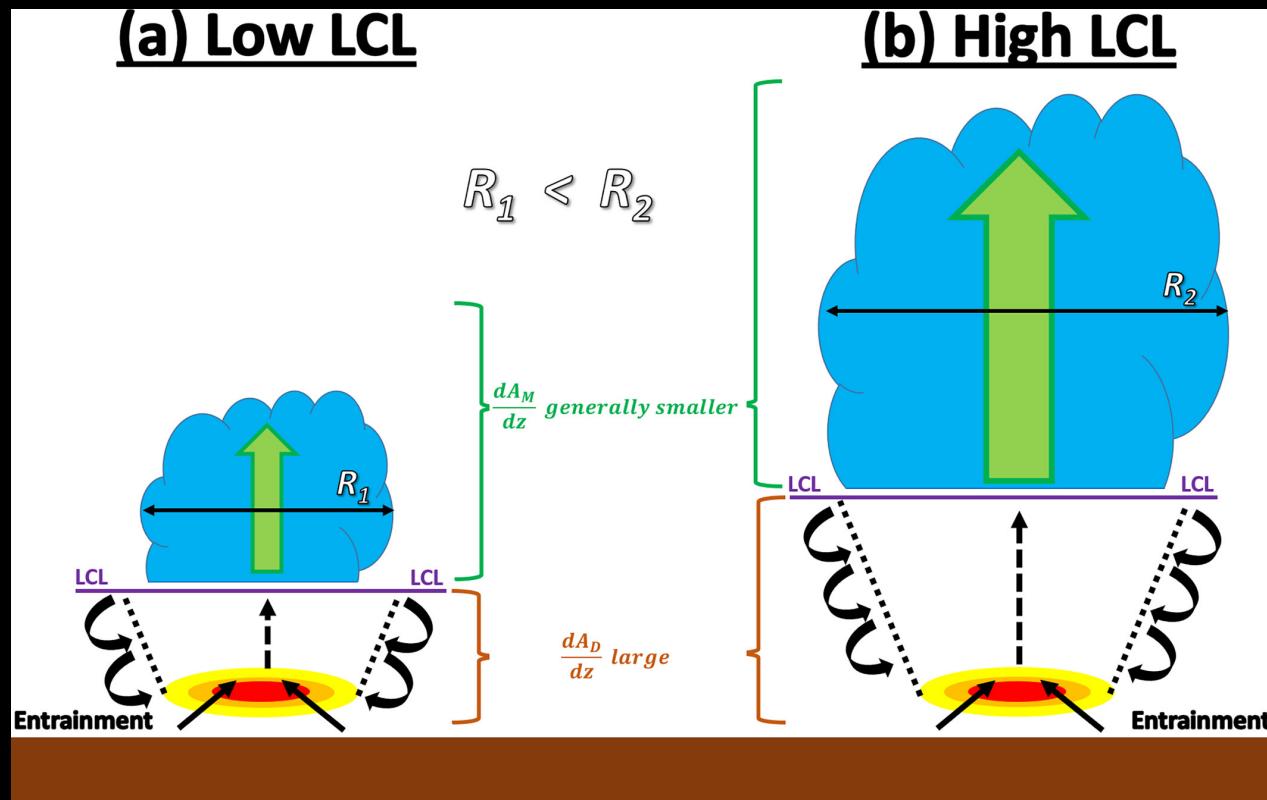


Size

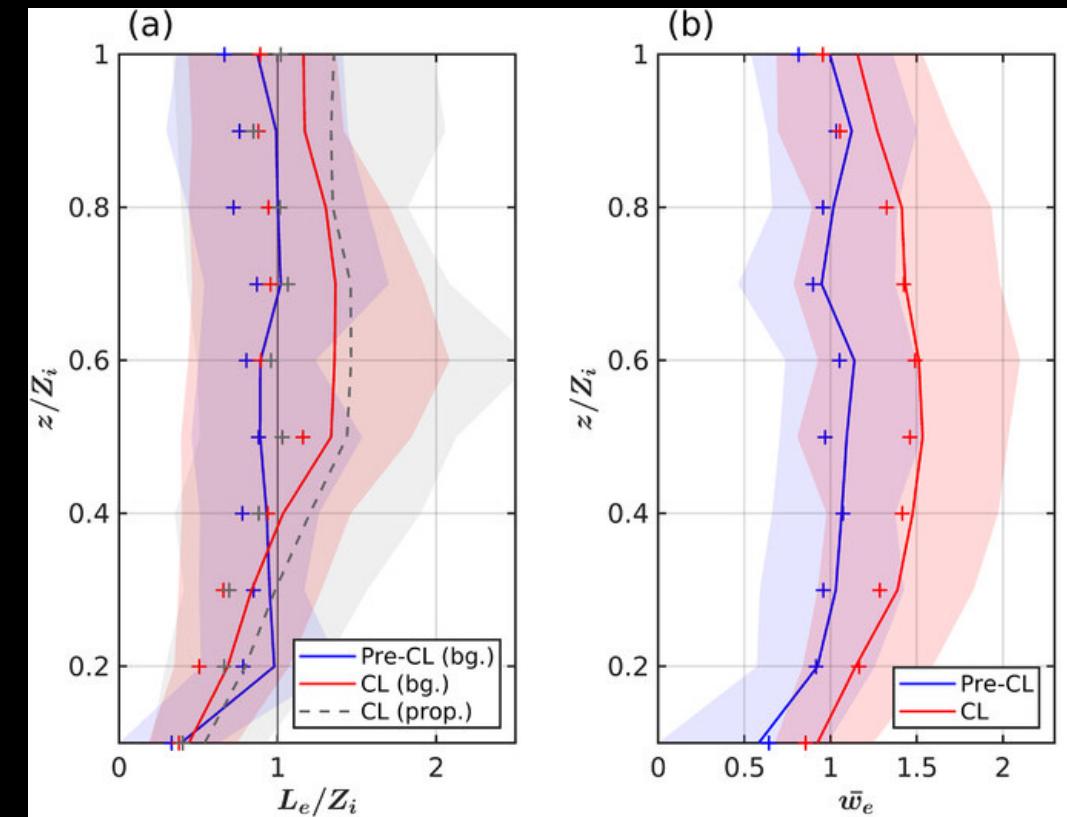
Strength

Red = In convergence line  
Blue = Not in convergence line

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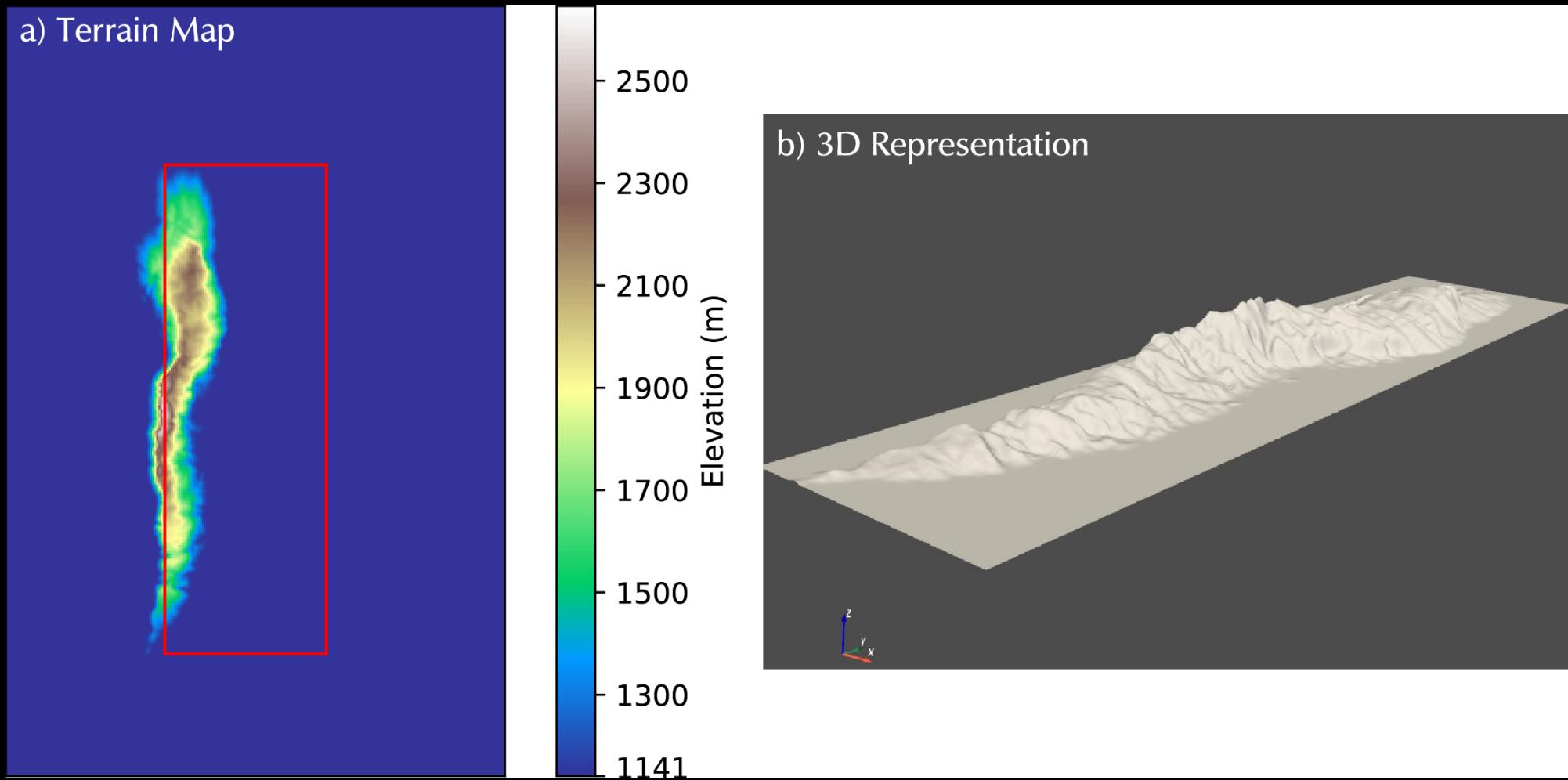
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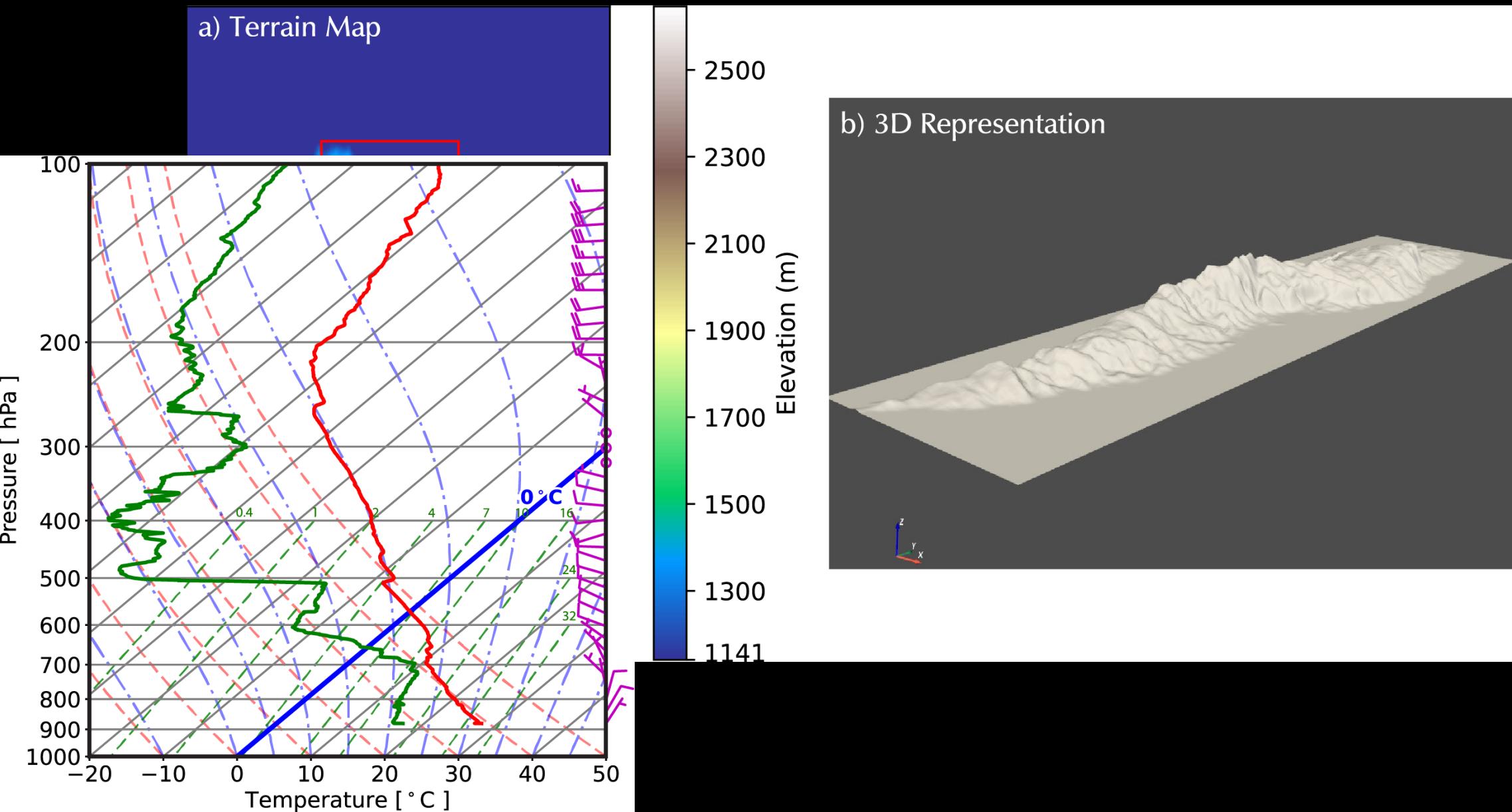


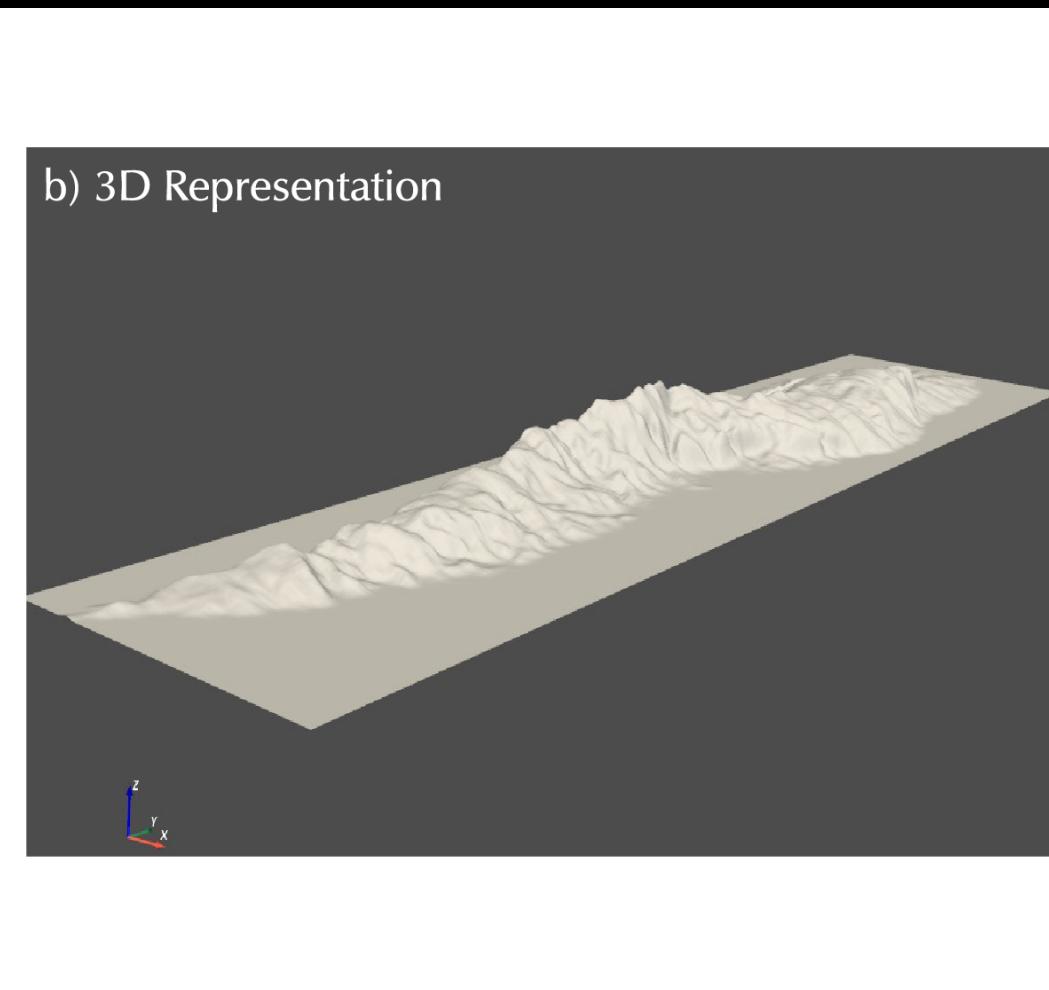
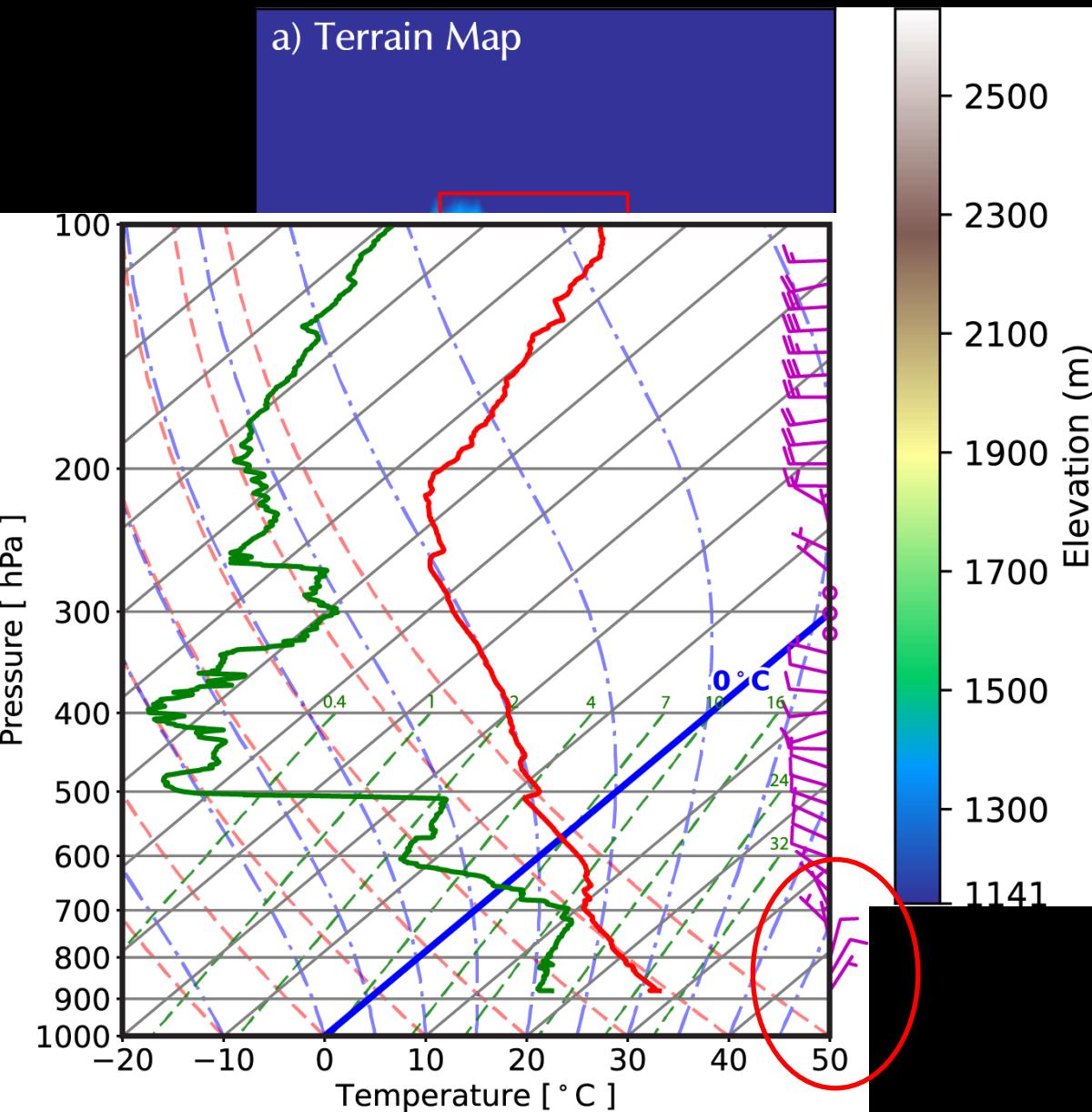
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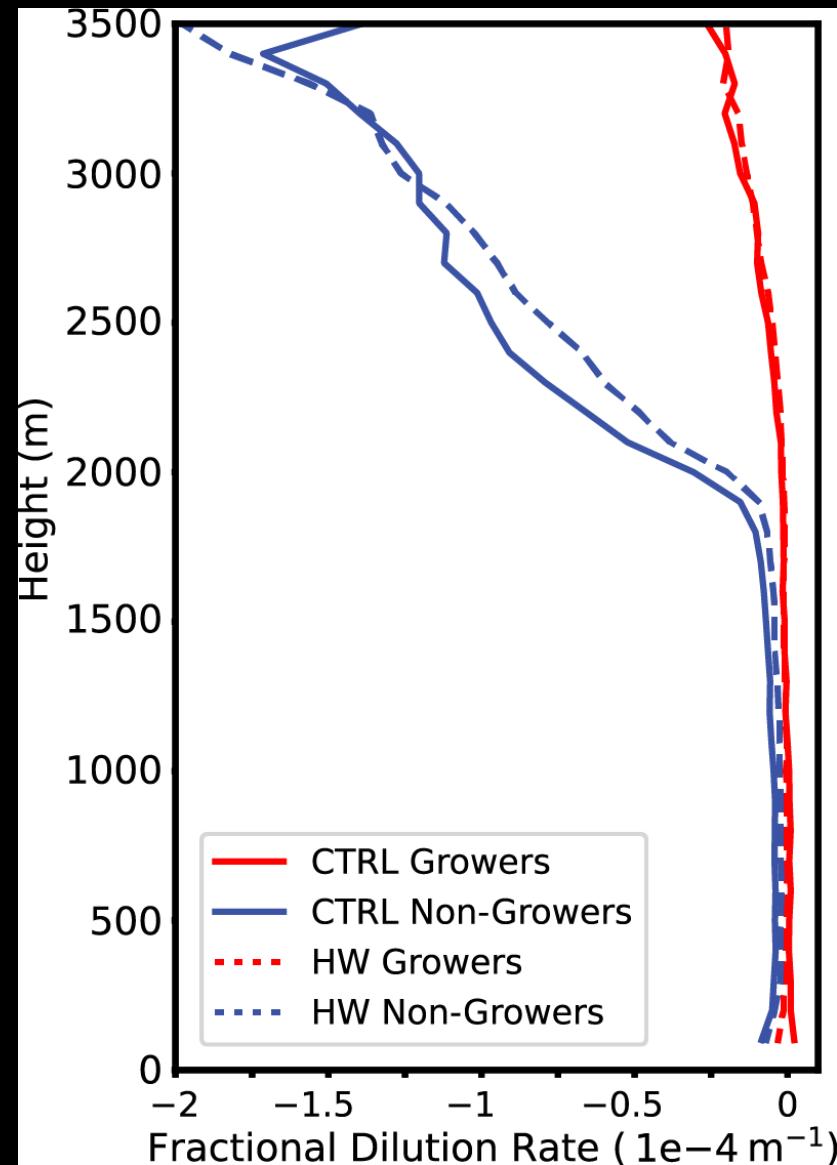
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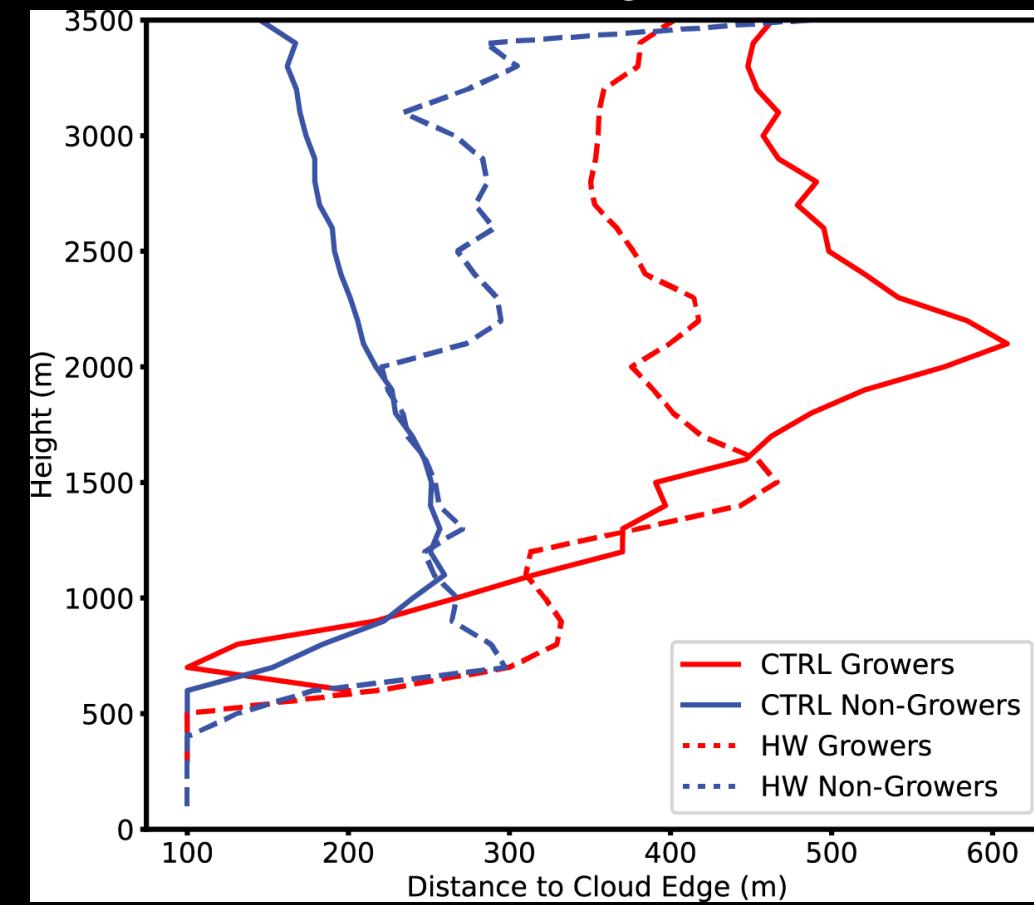


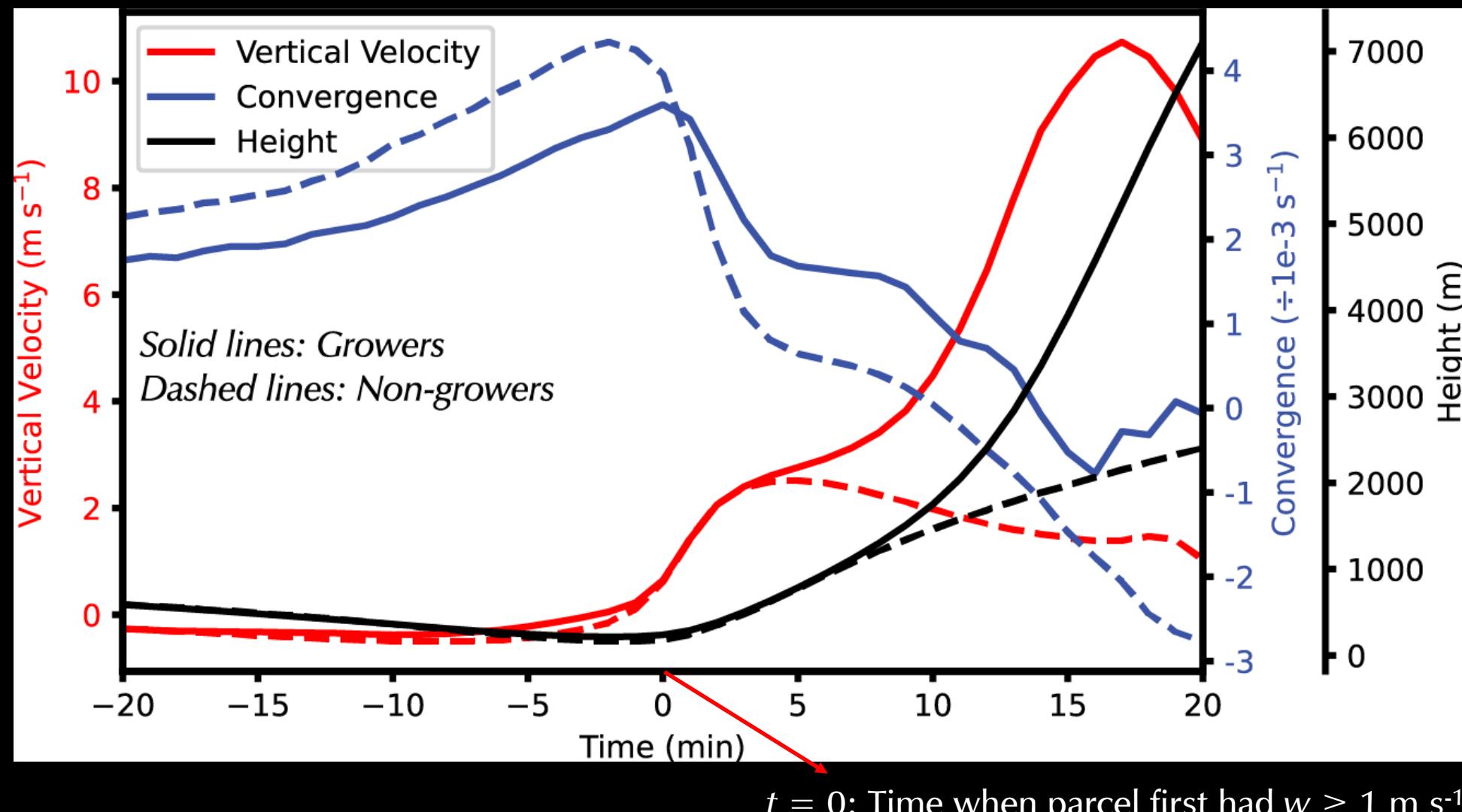


## Fractional dilution rate



## Distance to nearest cloud edge



Time series of convergence,  $w$ , and height of parcel

# Conclusions

- Obviously, thermodynamic profile is very important factor for determining if deep convection can develop, but...

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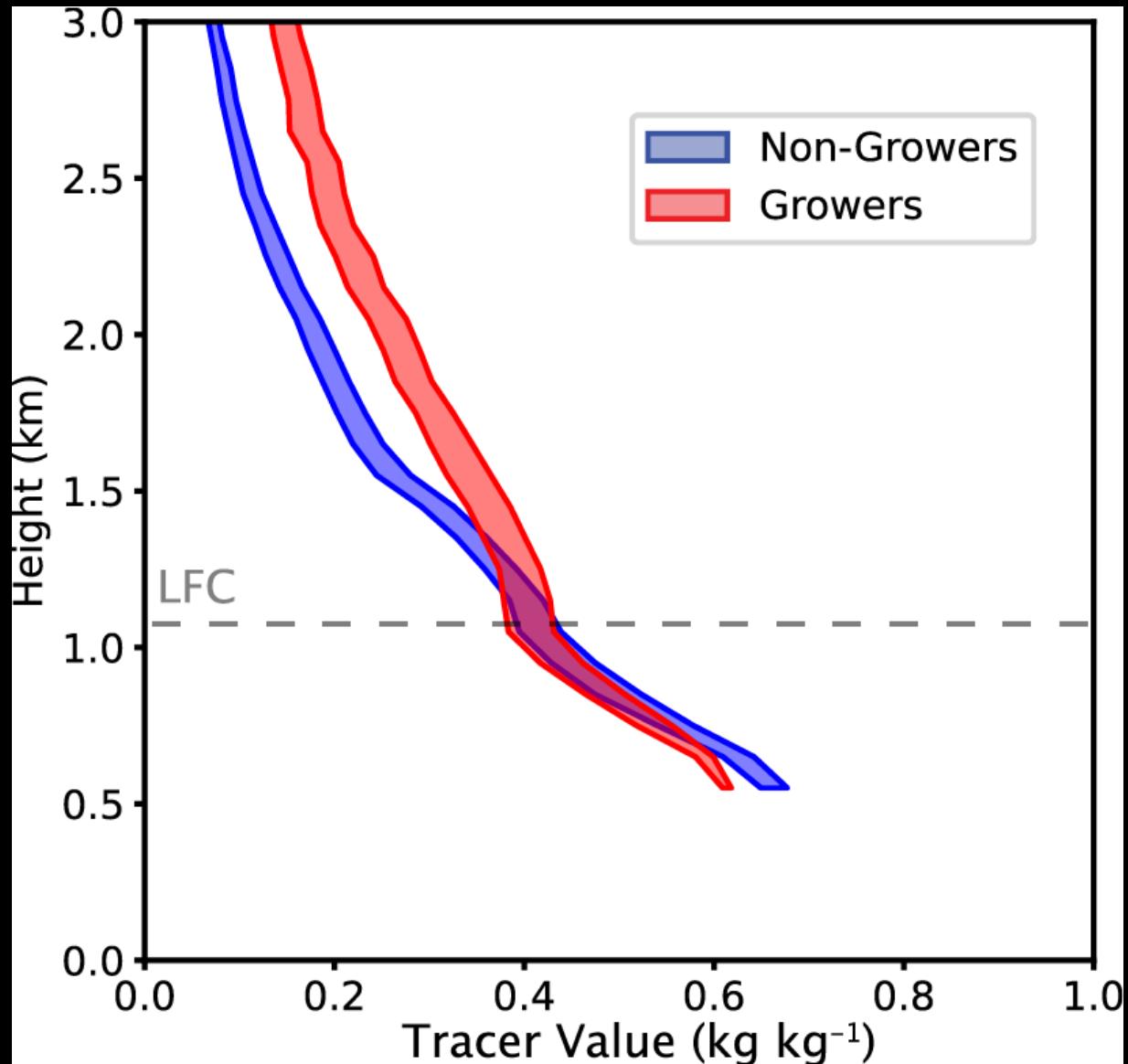
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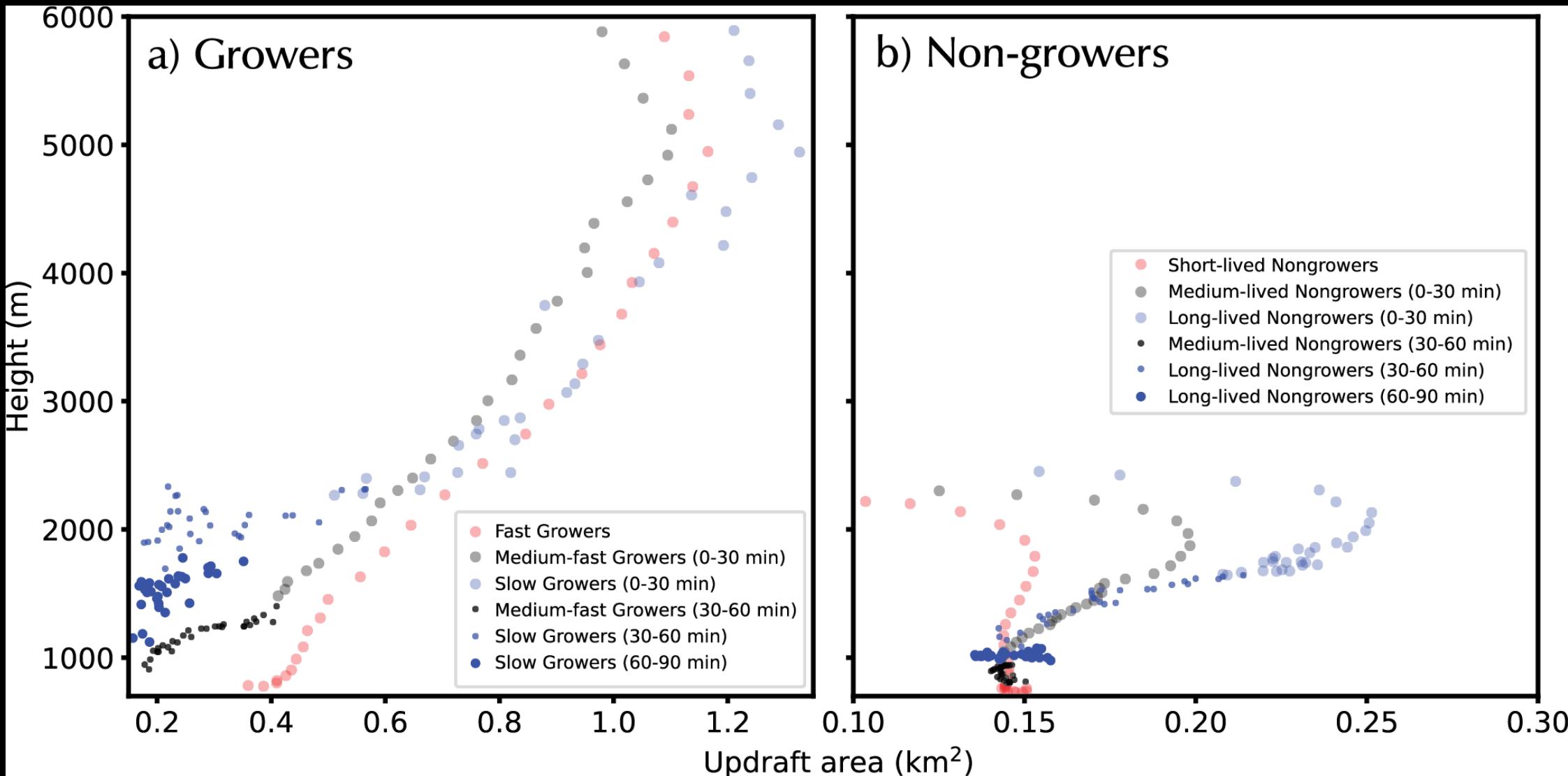
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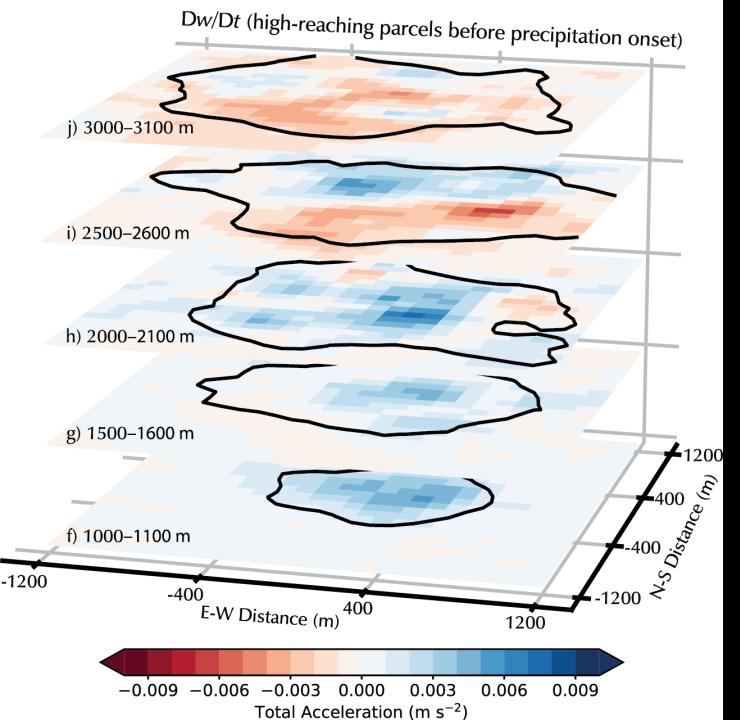
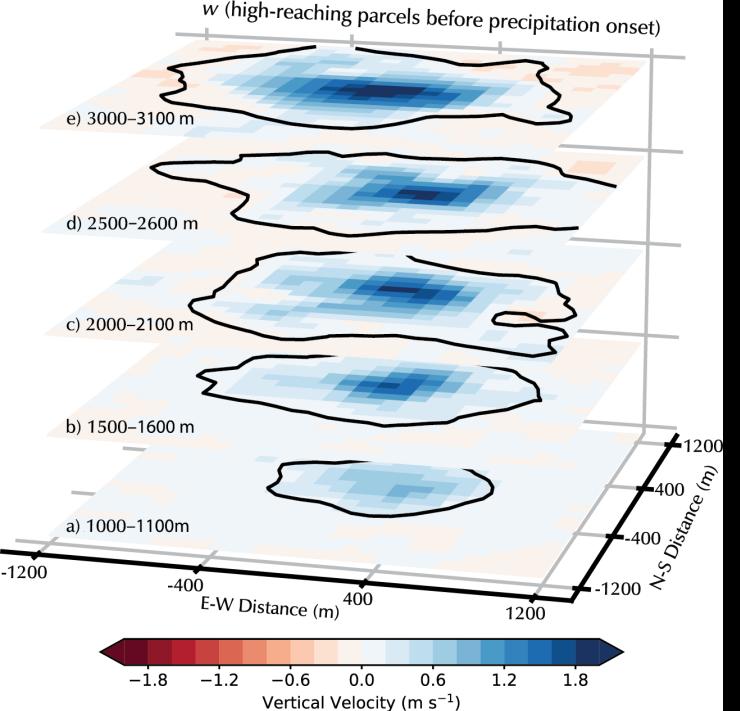
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- What controls the size?
  - Convergence seems to be important. But is it a cause or consequence of convection?
  - Moisture “preconditioning” by weak convection may actually matter on short time scales by allowing successive clouds to grow wider. Just a hypothesis for now.

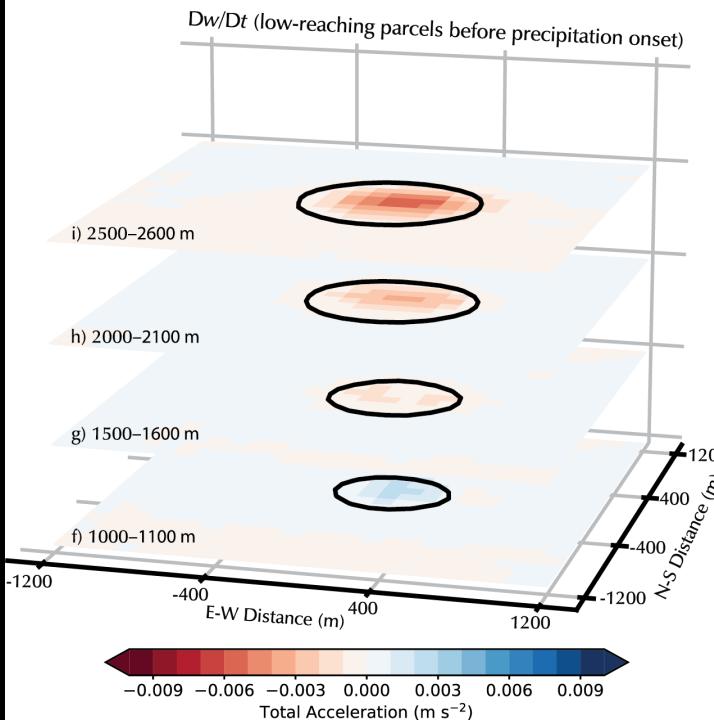
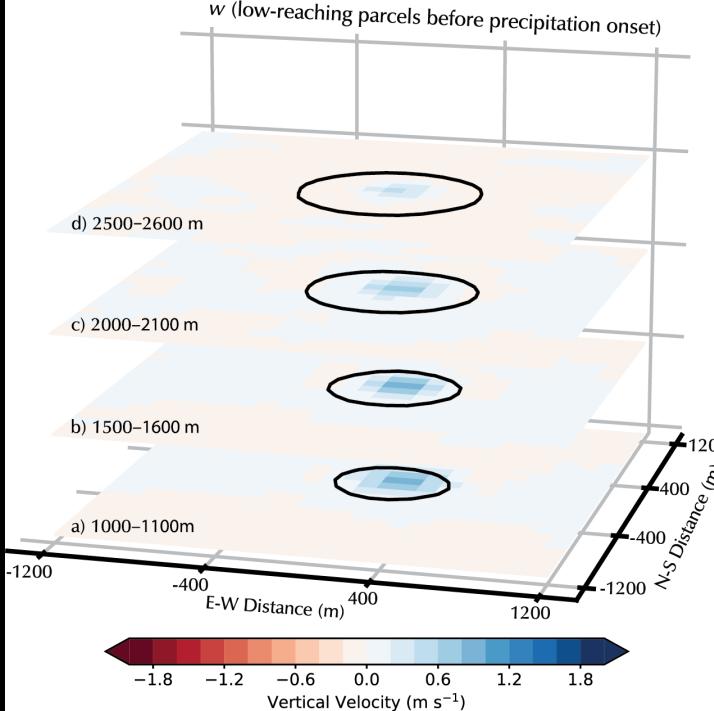




## Ascending Parcels (growers)



## w (low-reaching parcels before precipitation onset)



Non-Ascending  
Parcels  
(non-growers)