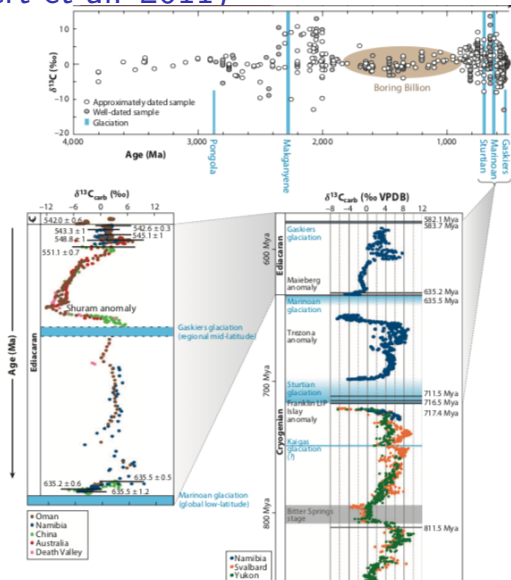
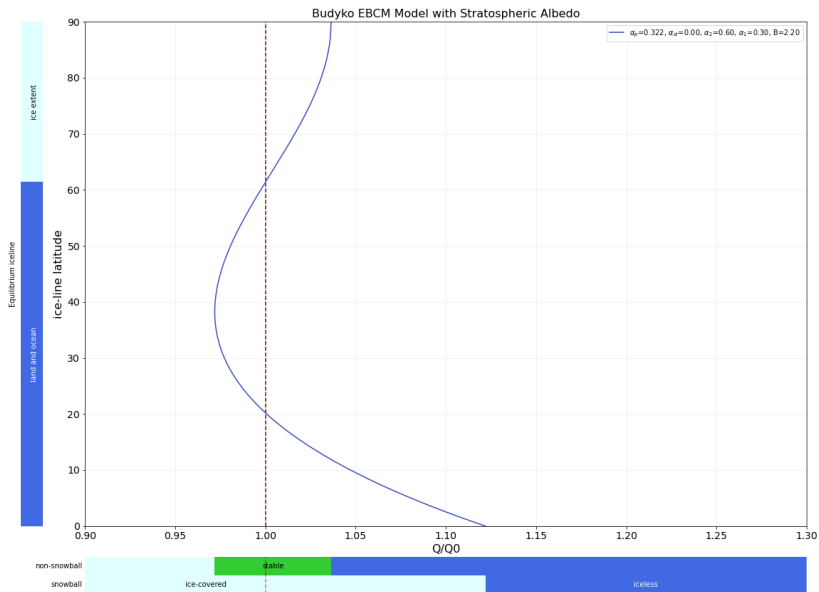


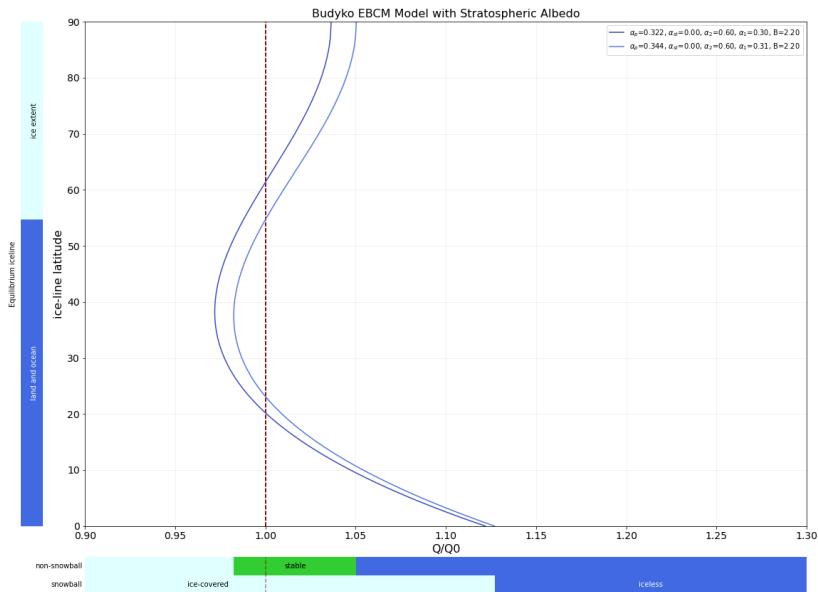
Could albedo changes explain Neoproterozoic glaciations? (Pierrehumbert et al. 2011)



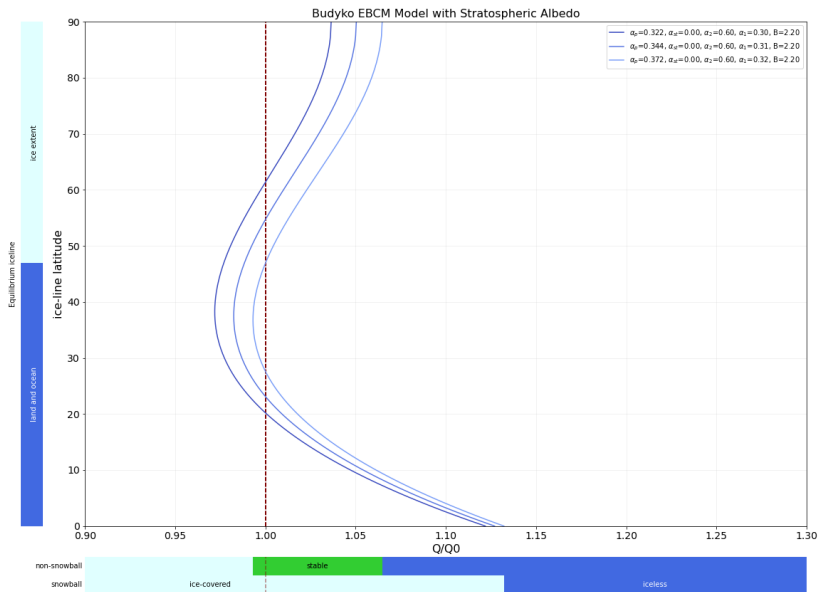
Increasing land and ocean albedo can cause a snowball



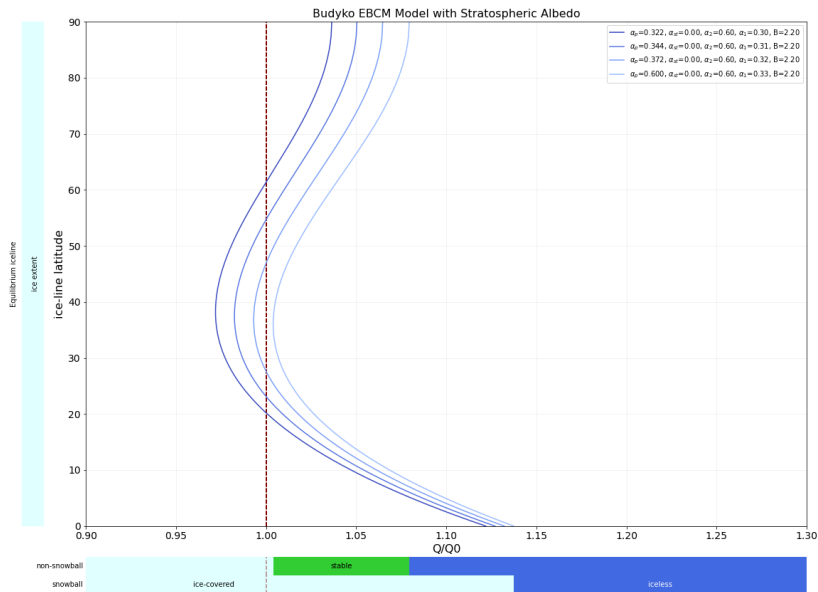
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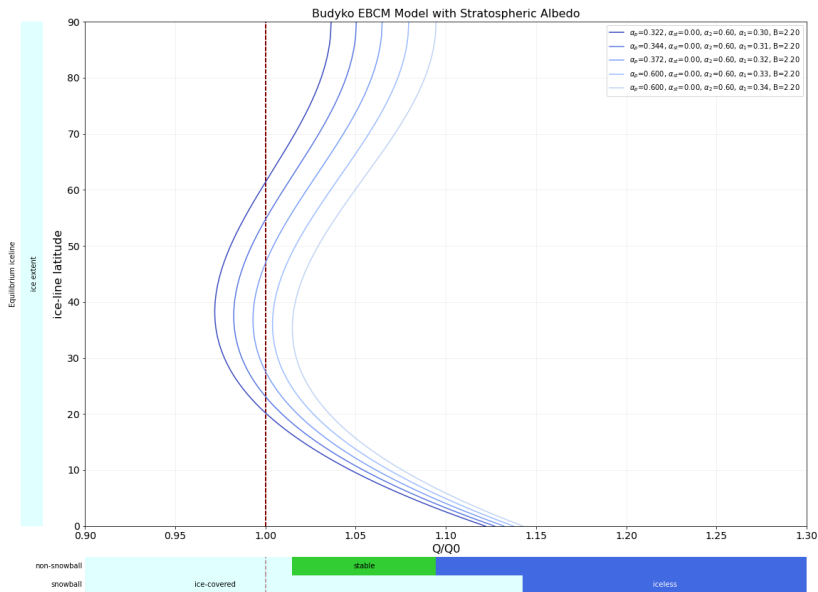
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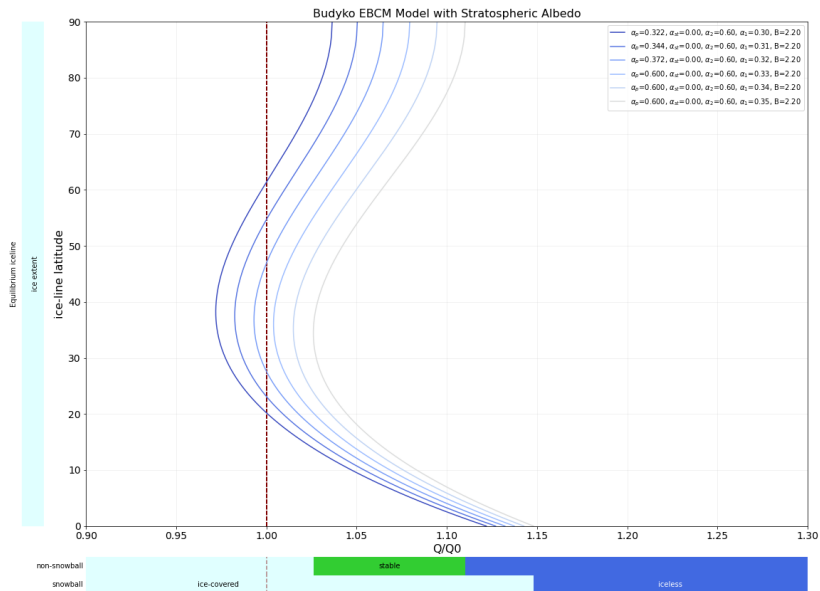
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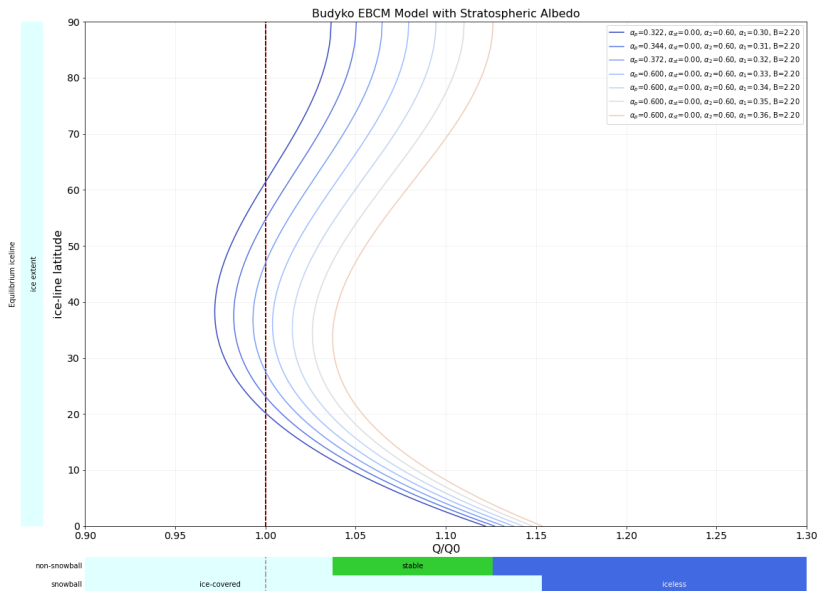
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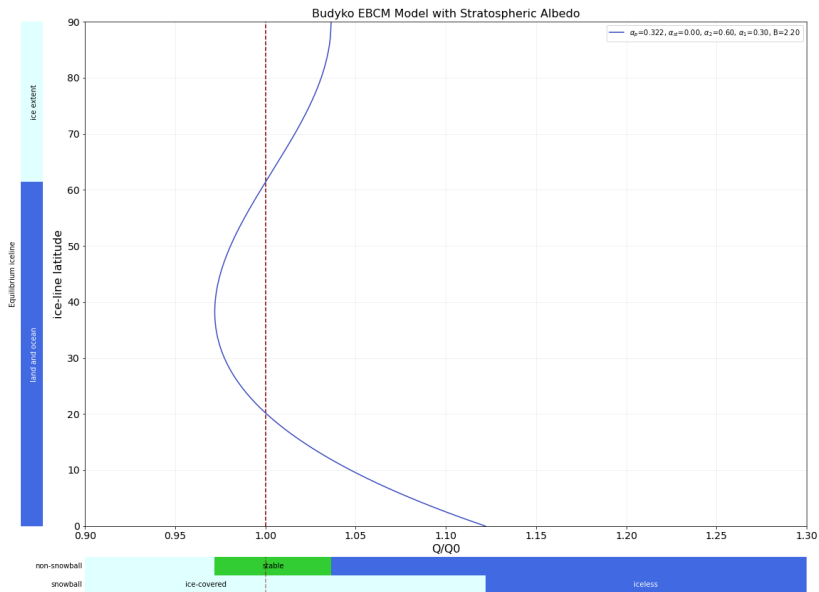
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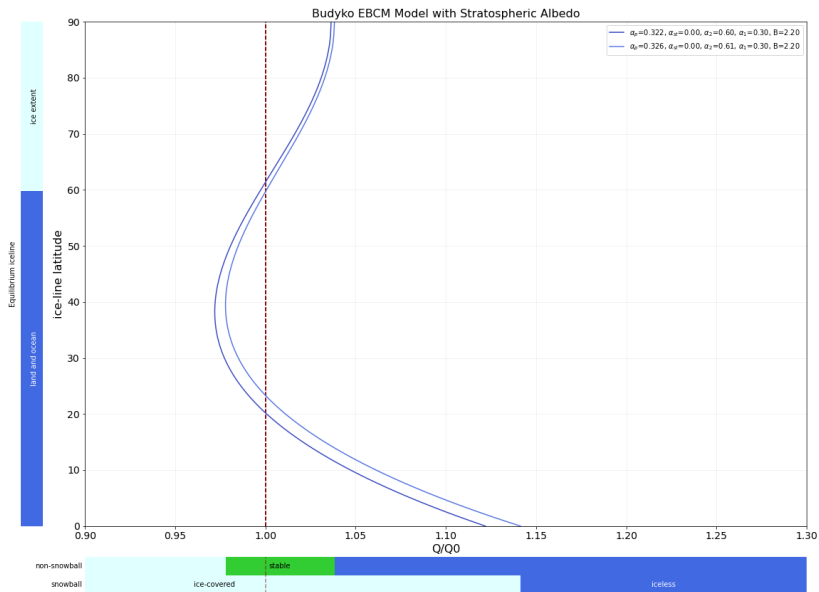
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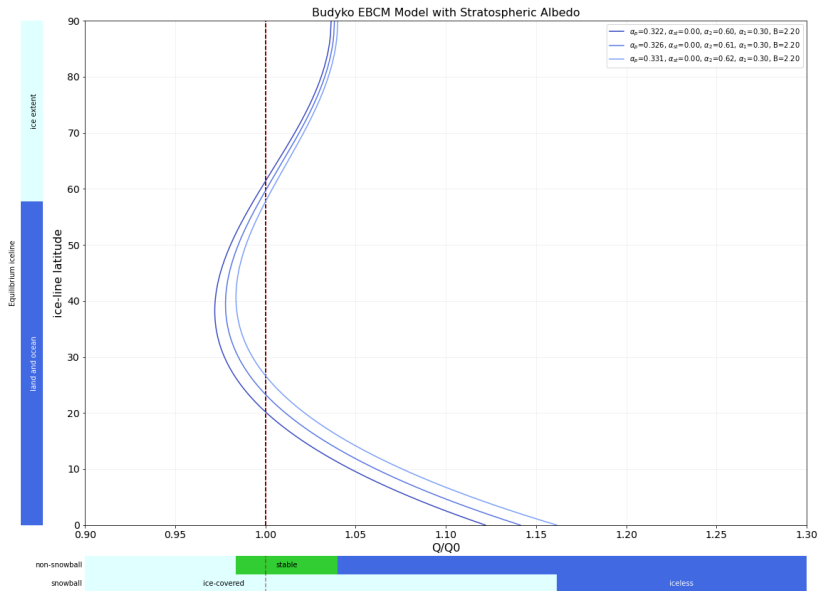
So can increasing ice albedo.



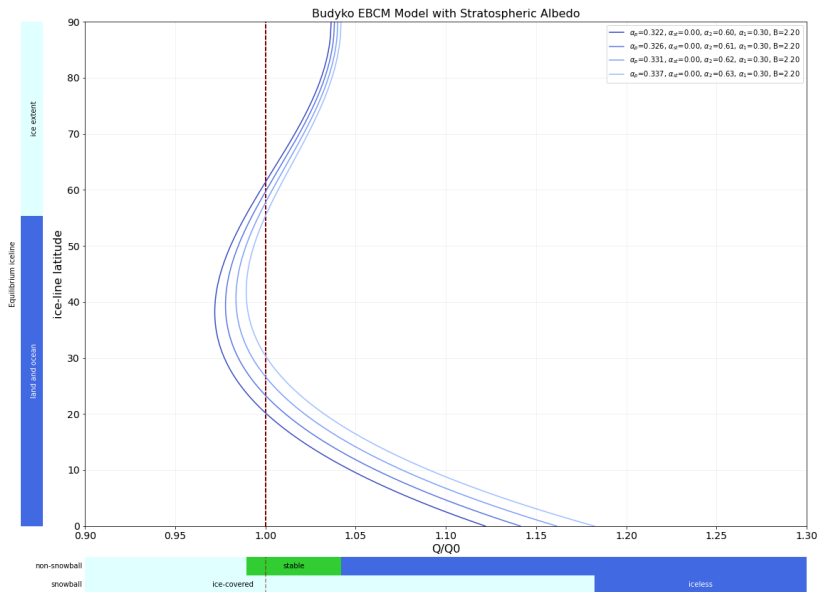
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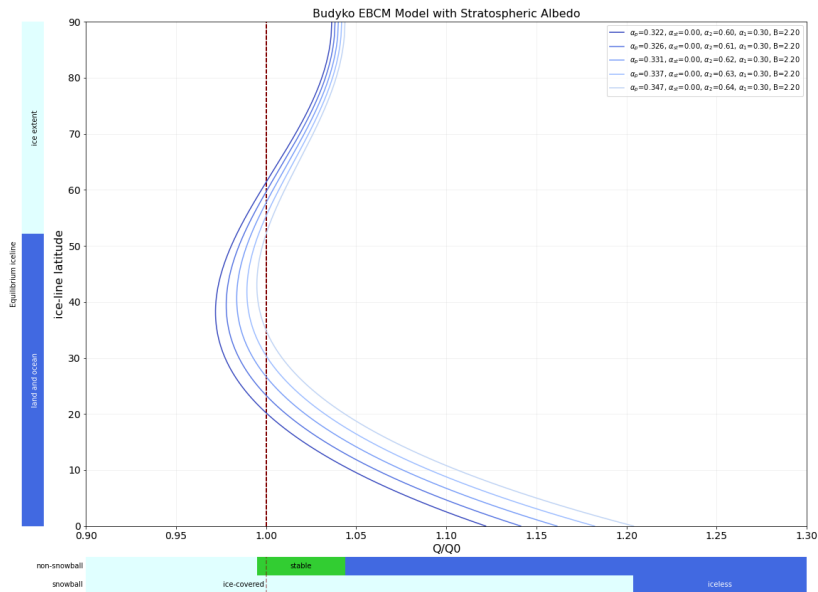
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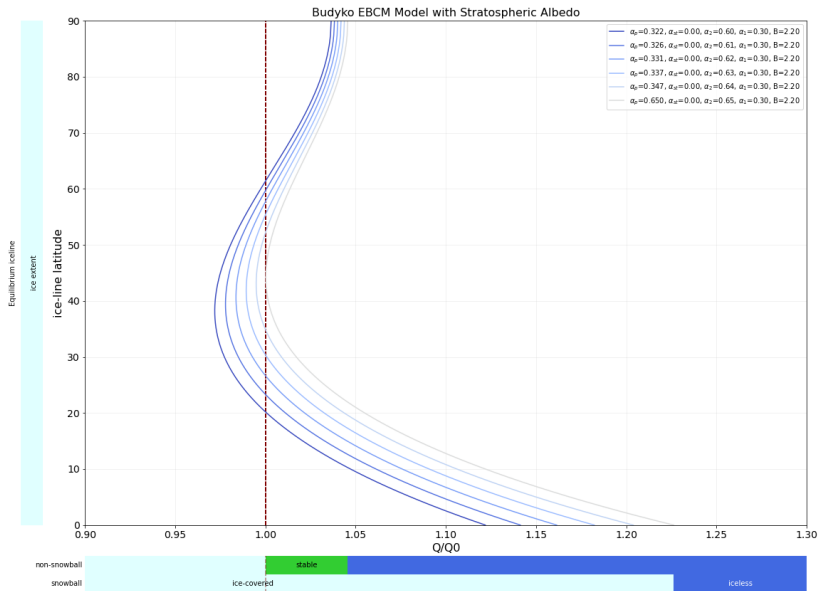
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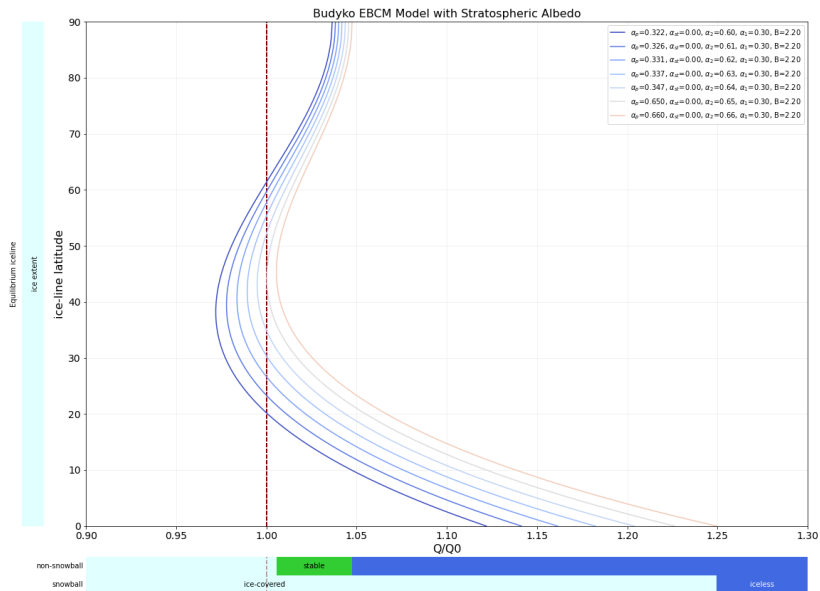
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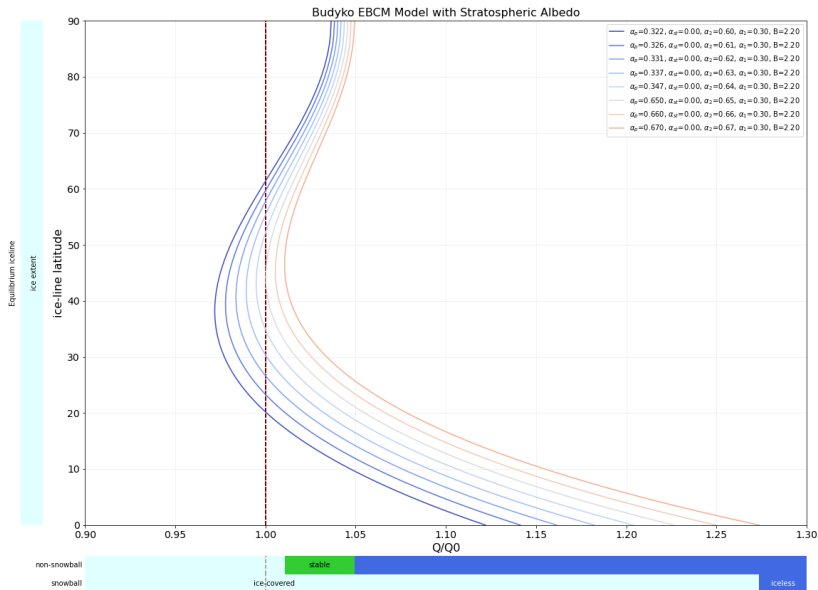
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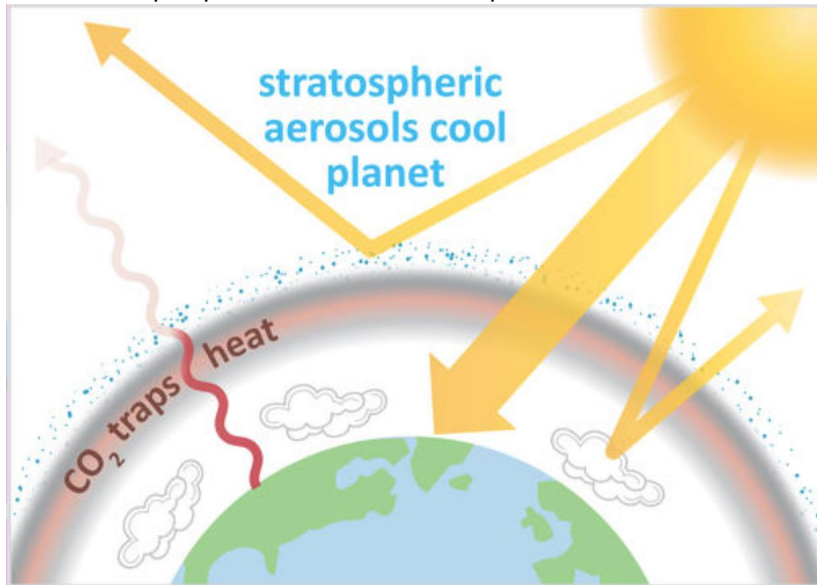


So can increasing ice albedo.



A volcanic mechanism to increase albedo

Volcanoes put particulates into *stratosphere* that increase albedo



τ is Optical Depth, $\alpha_{st}(\tau, x_s)$ is Stratospheric Albedo

Adapted equation from Budyko (1969)

$$\frac{Q_0}{4} S(x_s) (1 - \alpha_{st}(x_s, \tau)) (1 - \alpha_{su}(x_s)) = A + BT(x_s) + C(T(x_s) - \bar{T})$$

τ is Optical Depth, $\alpha_{st}(\tau, x_s)$ is Stratospheric Albedo

Adapted equation from Budyko (1969)

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Assumption: Uniform effect over surface.

$$1 - e^{-\tau m(x)} = \alpha_{st}(x, \tau) = \alpha_{st}(\tau) = 1 - e^{-\tau}$$

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Assumption: Uniform effect over surface.

$$1 - e^{-\tau m(x)} = \alpha_{st}(x, \tau) = \alpha_{st}(\tau) = 1 - e^{-\tau}$$

Allowing us to calculate a new planetary albedo (surface + stratosphere)

$$\alpha'_p = \int_0^1 S(x) [\alpha_{st}(\tau) + \alpha_{su}(x_s) - \alpha_{st}(\tau) \alpha_{su}(x_s)] dx,$$

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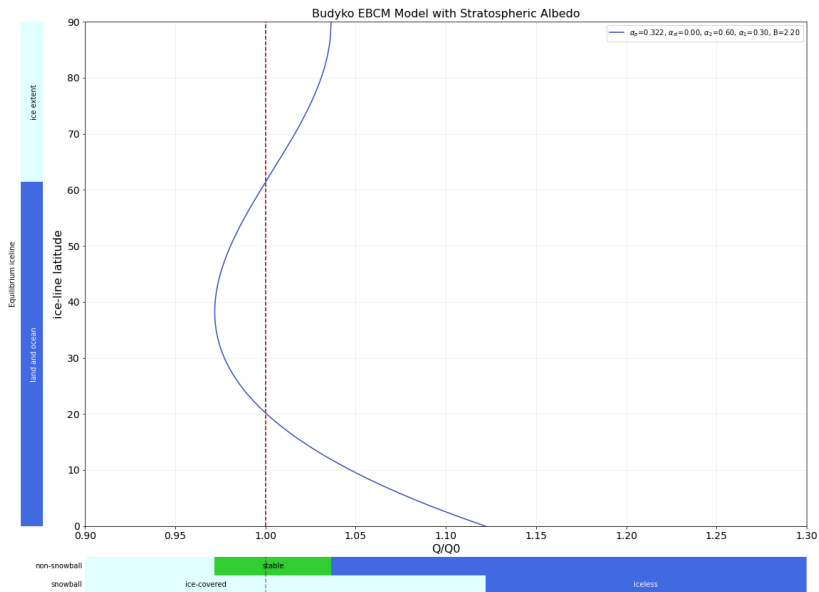
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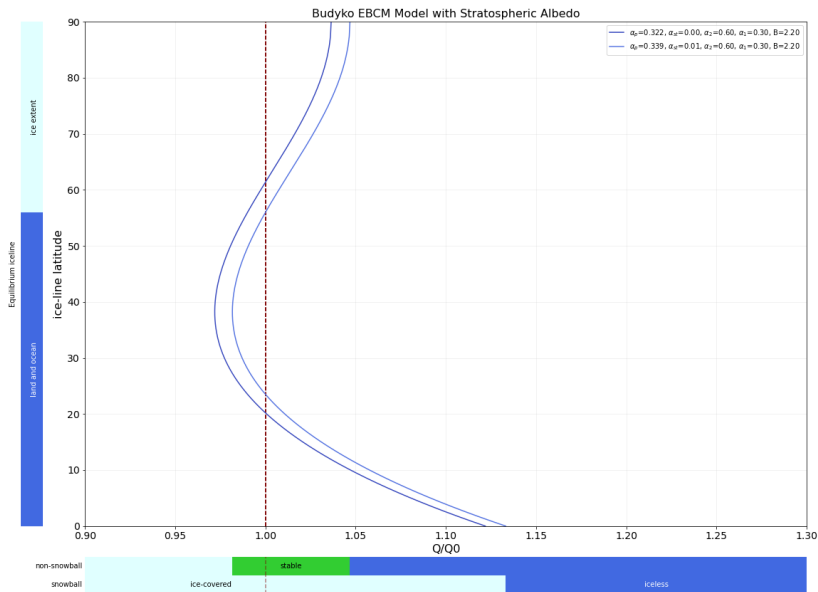
Resulting in a new equation to solve (Roe and Baker 2010)

$$\frac{Q}{4} S(x_s) (1 - \alpha_{su}) (1 - \alpha_{st}) + \frac{QC}{4B} (1 - \alpha'_p) = k$$

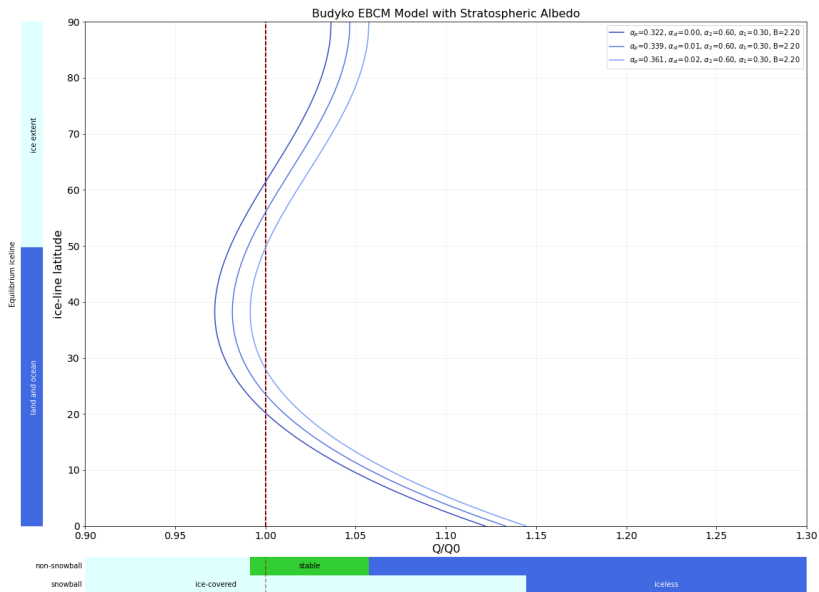
Increasing τ or α_{st}



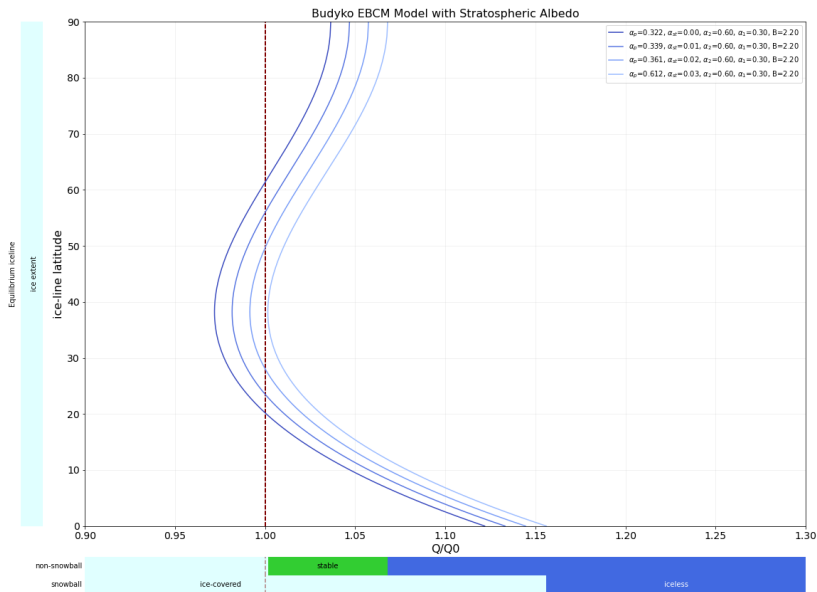
Increasing τ or α_{st}



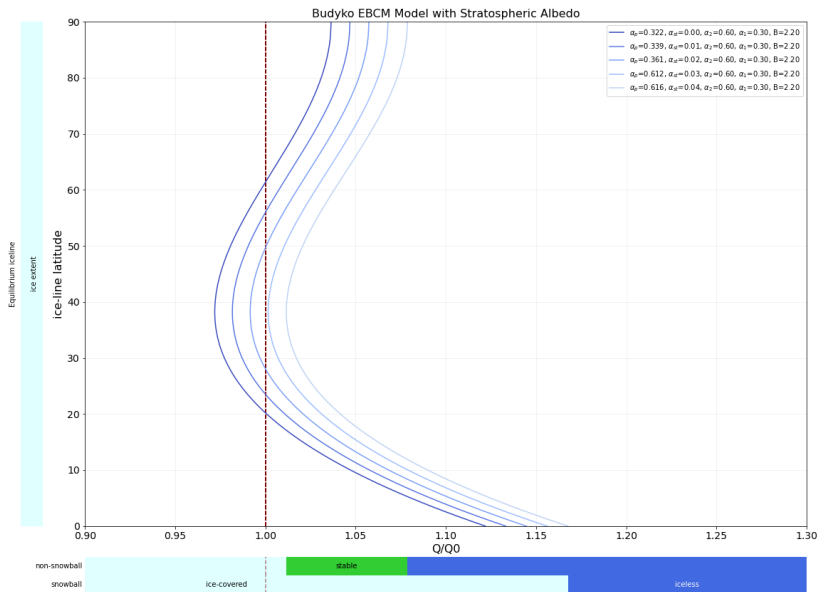
Increasing τ or α_{st}



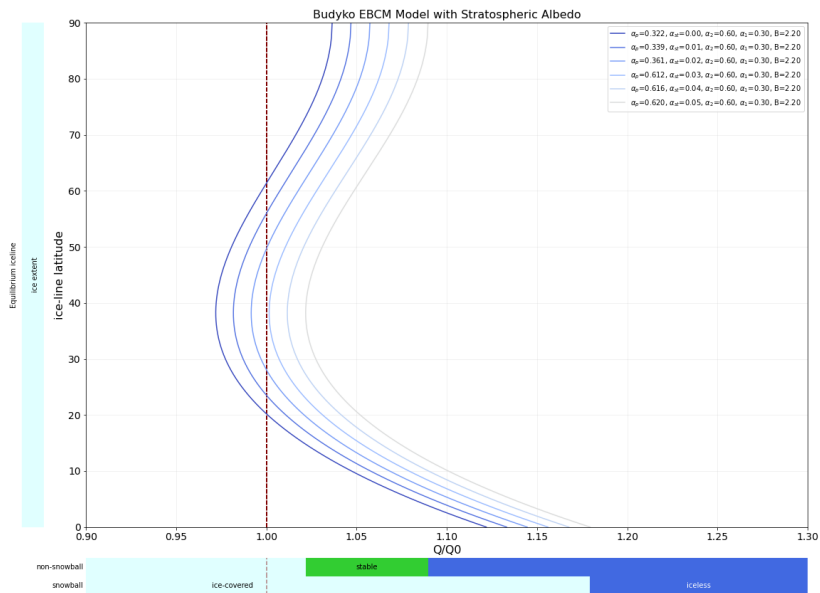
Increasing τ or α_{st}



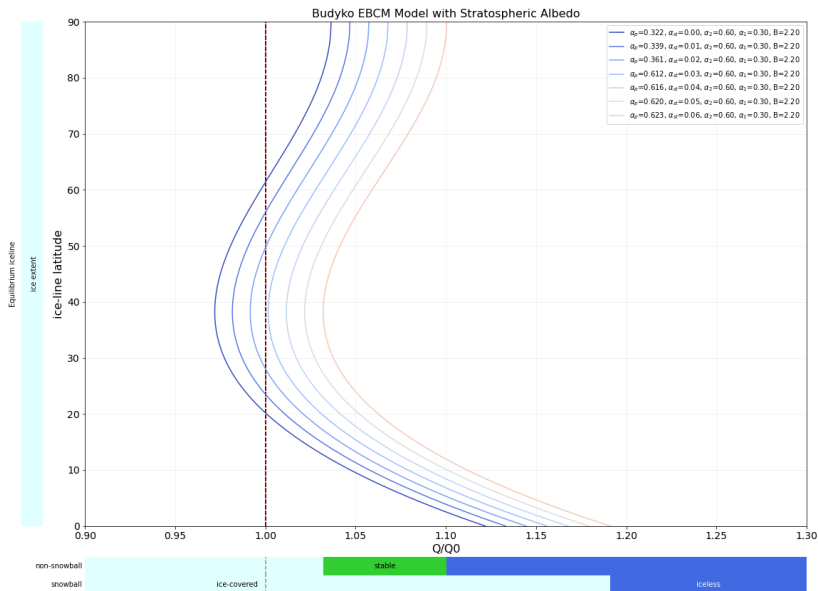
Increasing τ or α_{st}



Increasing τ or α_{st}

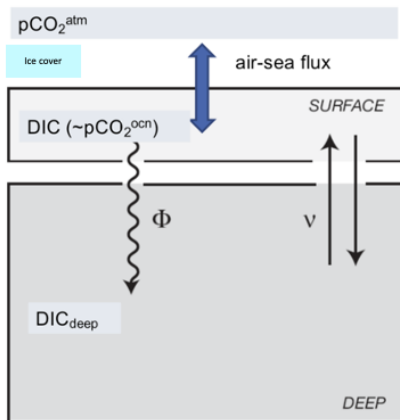


Increasing τ or α_{st}



Increasing $\text{CO}_2 \approx$ increasing B

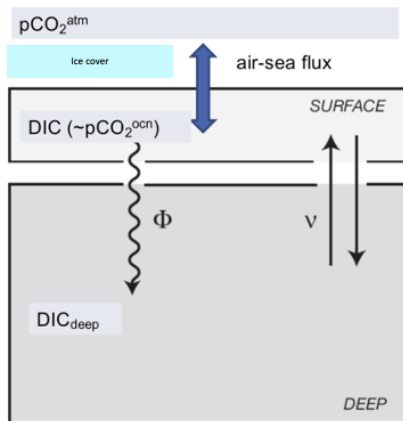
Ice sheets block the ocean sink for CO_2



Increasing $\text{CO}_2 \approx$ increasing B

Ice sheets block the ocean sink for CO_2

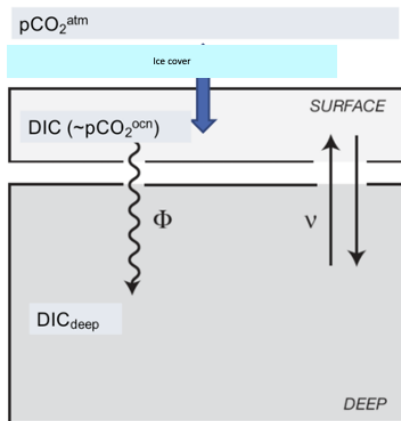
Increasing CO_2 increases \bar{T} , increasing OLR or B in the Budyko model.



Increasing $\text{CO}_2 \approx$ increasing B

Ice sheets block the ocean sink for CO_2

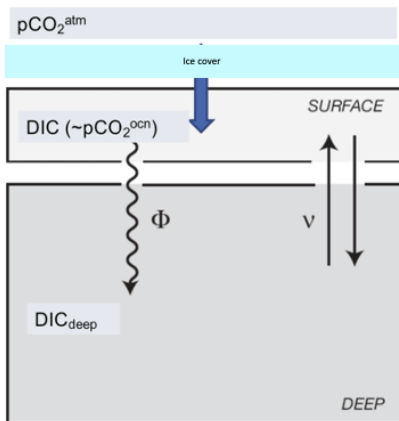
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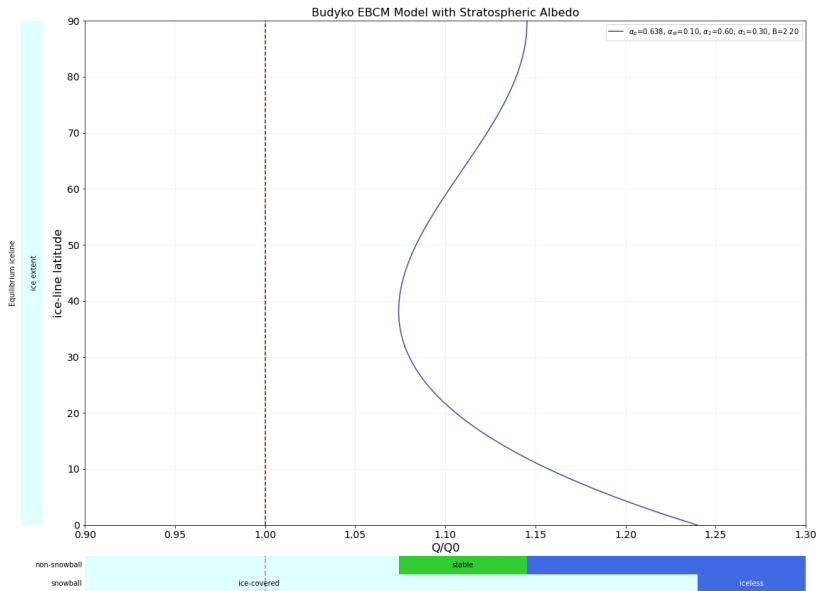
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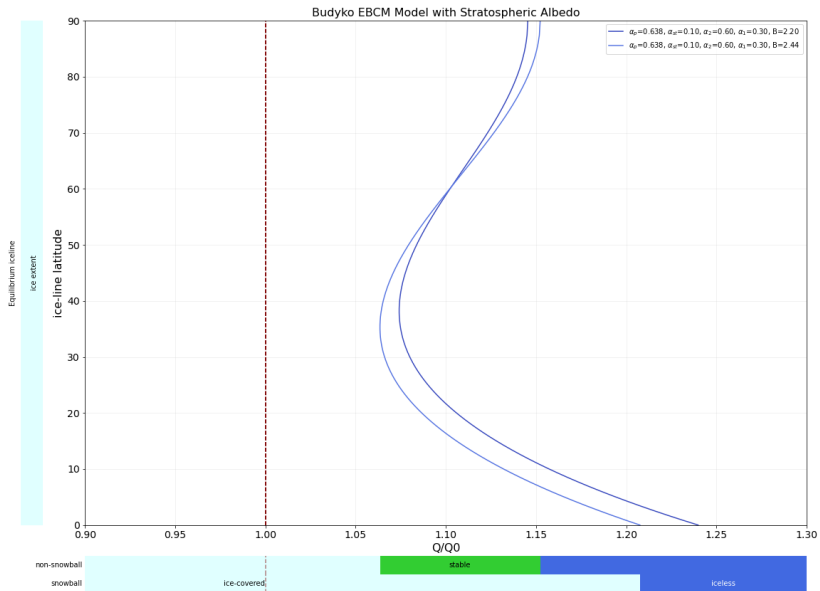
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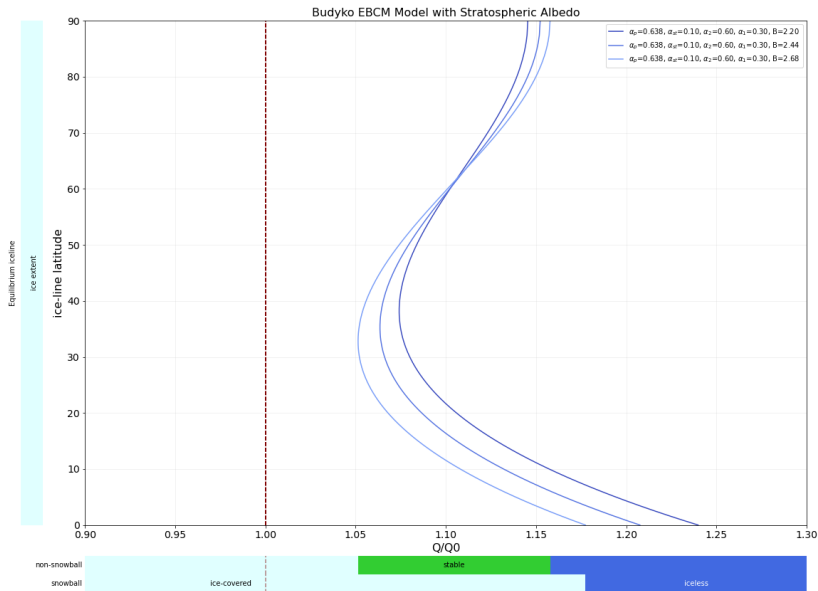
Increasing $\text{CO}_2 \rightarrow$ can get us back to a partial-ice world.



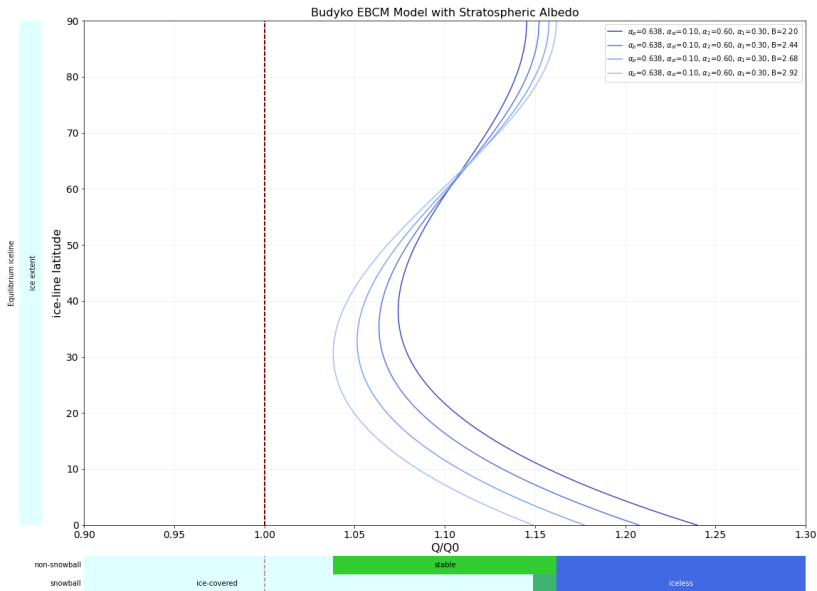
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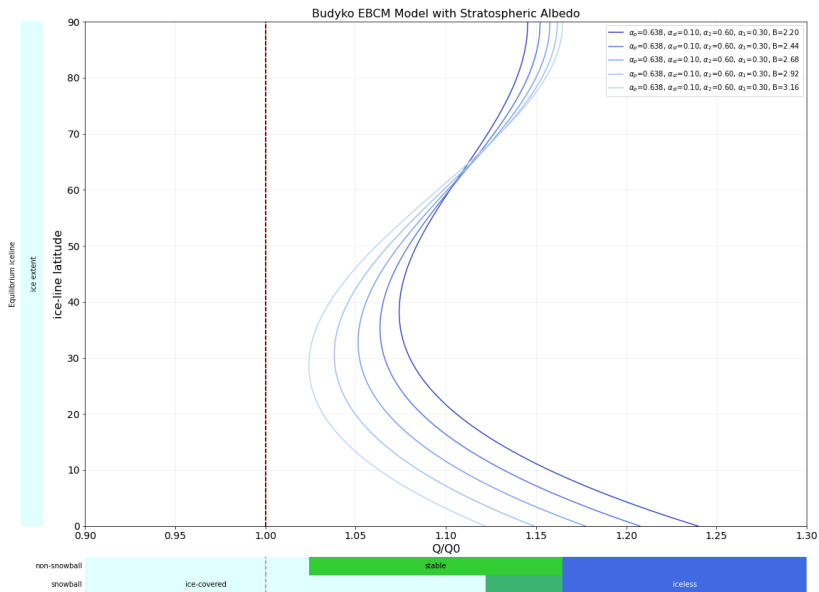
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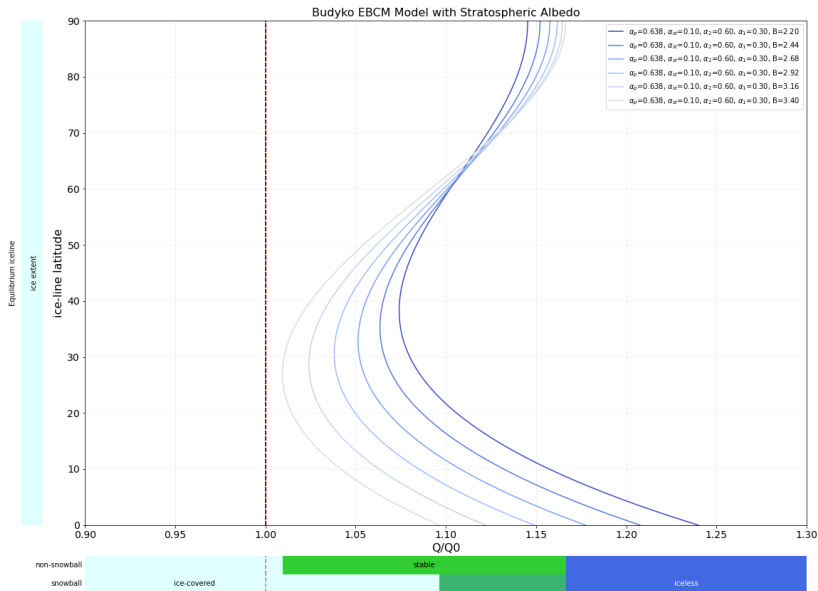
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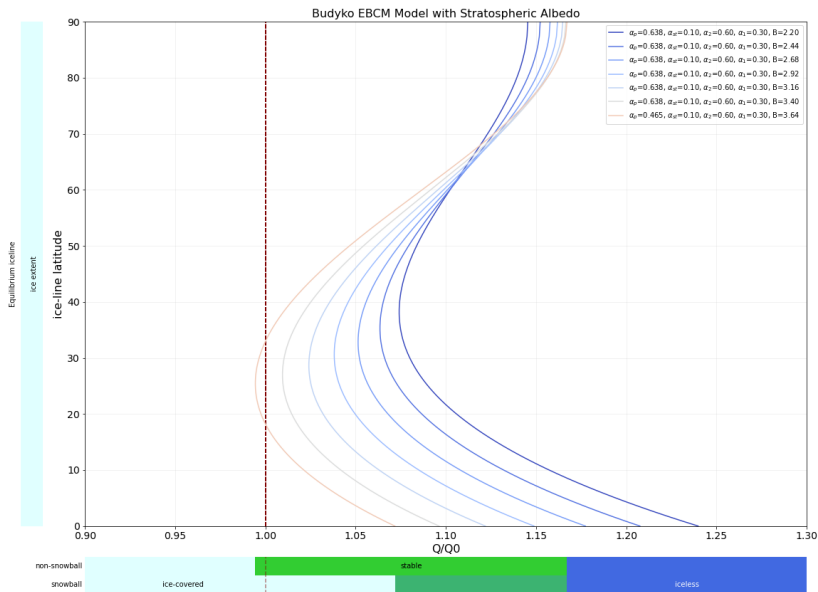
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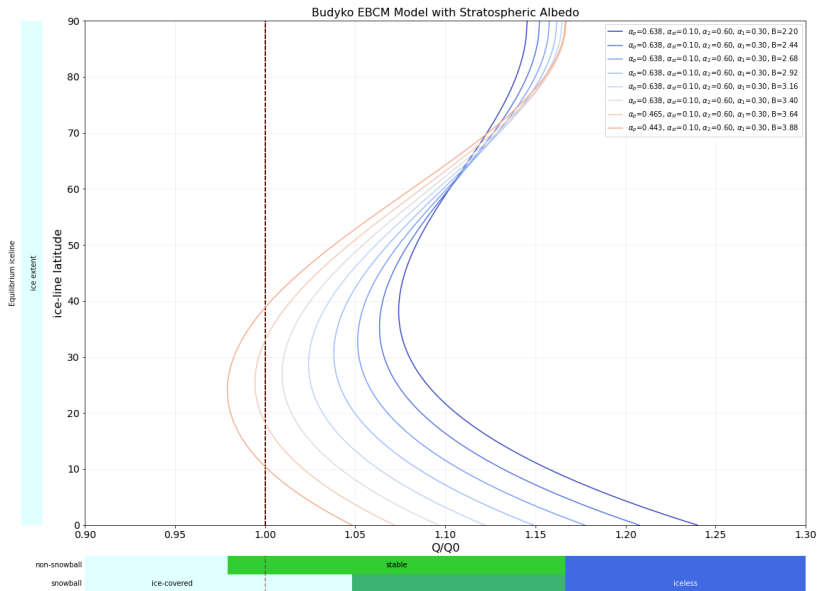
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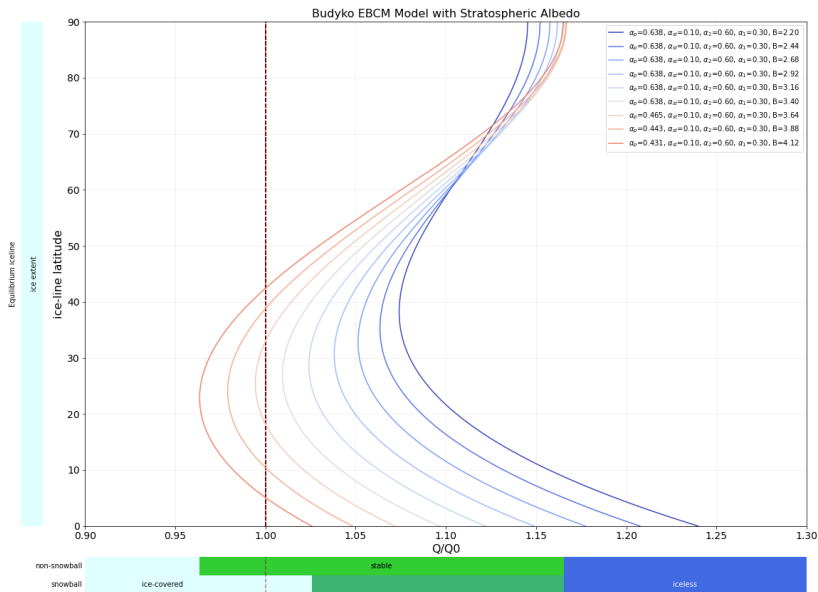
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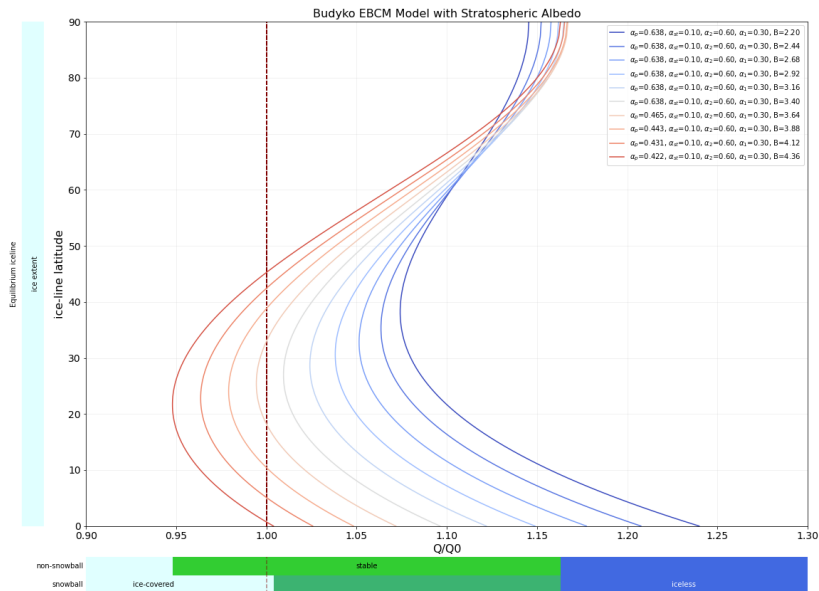
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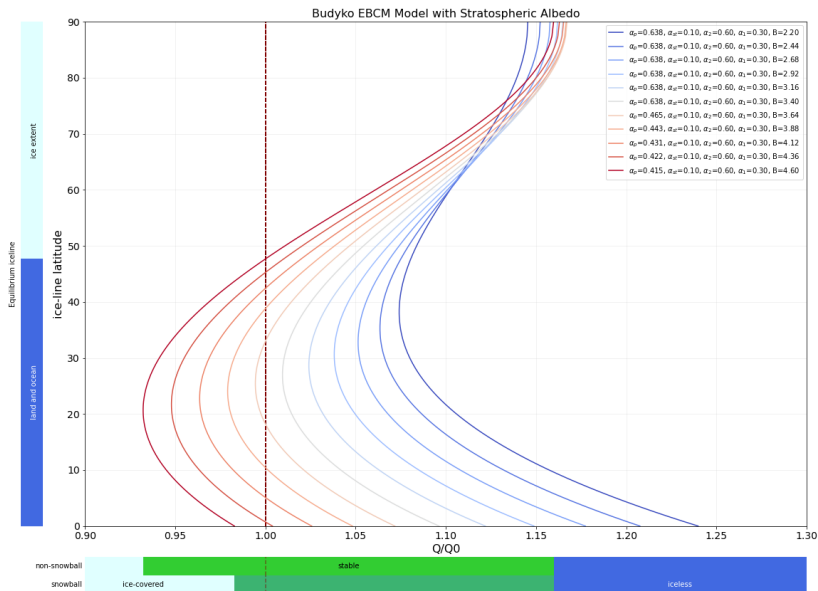
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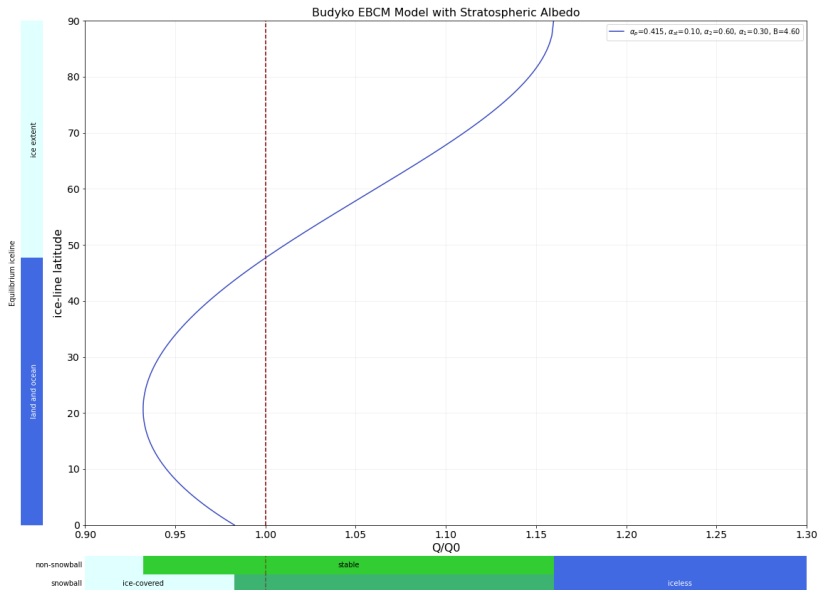
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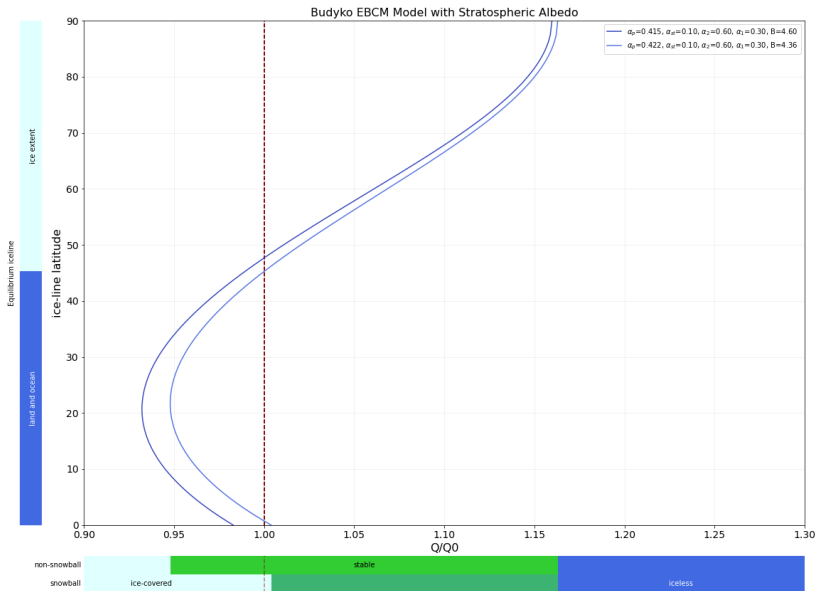
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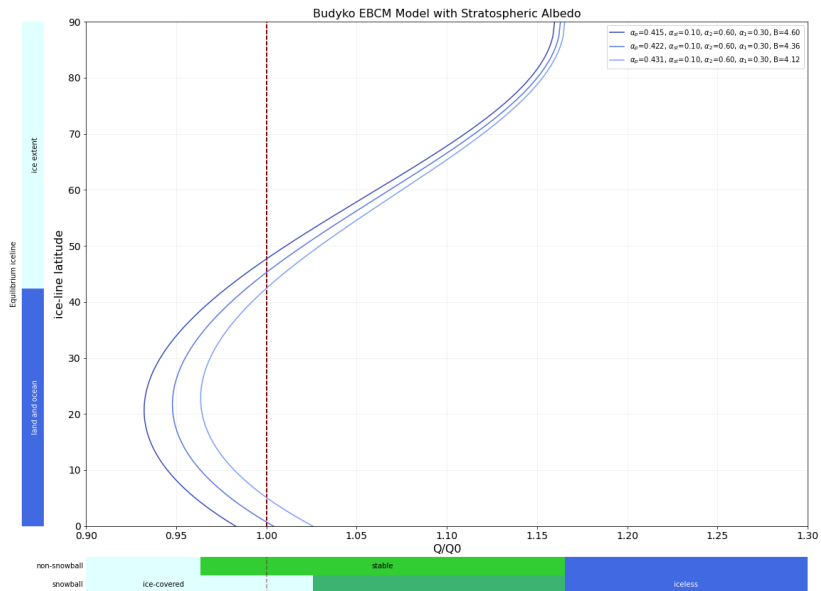
CO2 sink reactivates, causing hysteresis



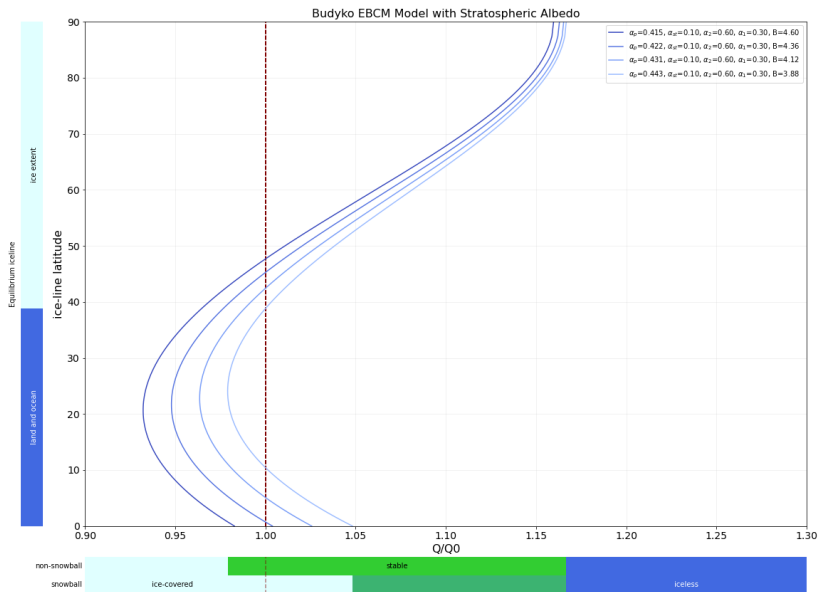
CO₂ sink reactivates, causing hysteresis



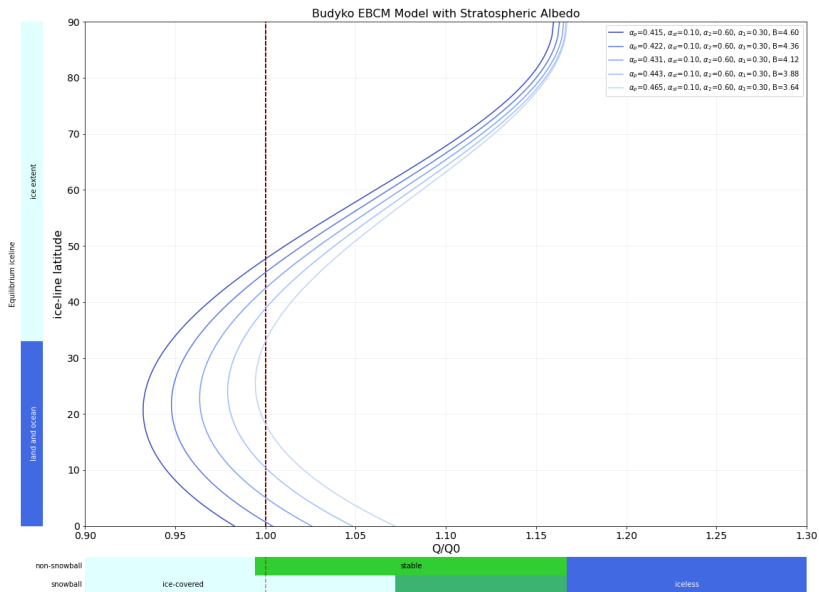
CO2 sink reactivates, causing hysteresis



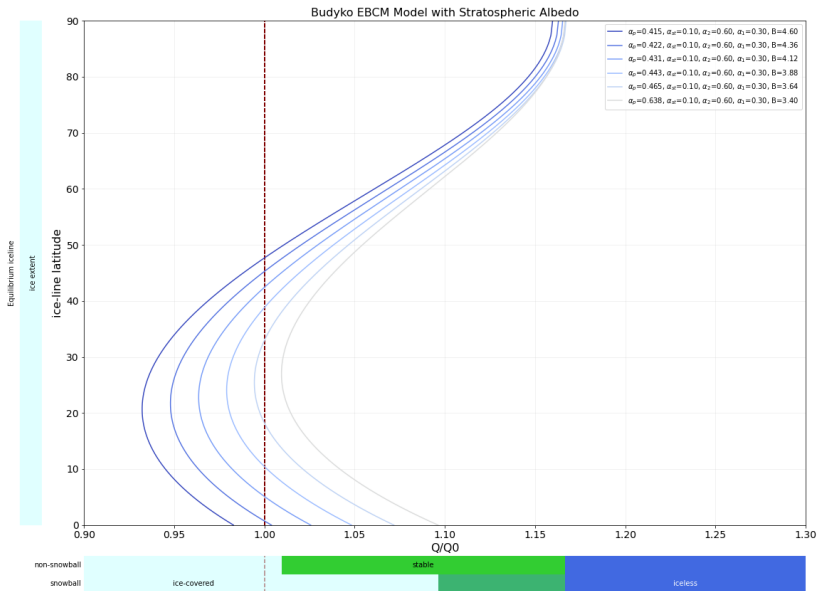
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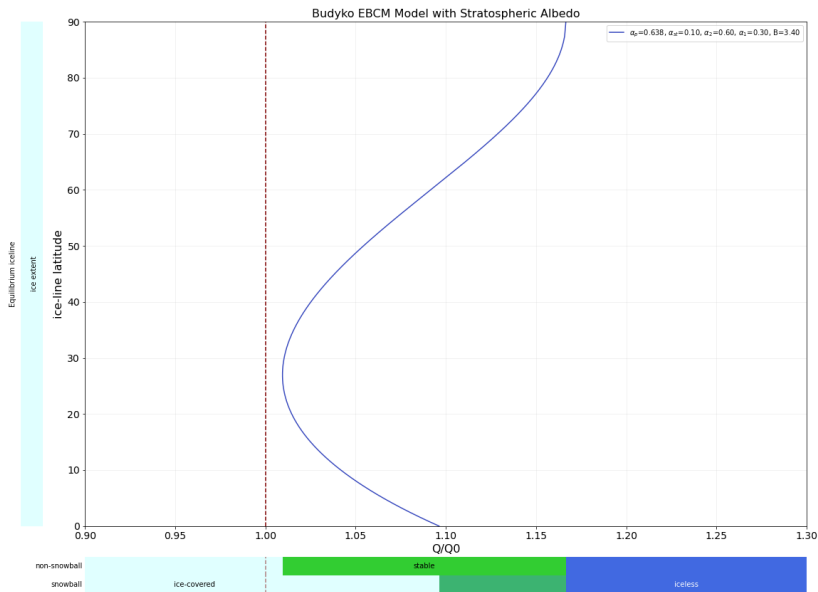
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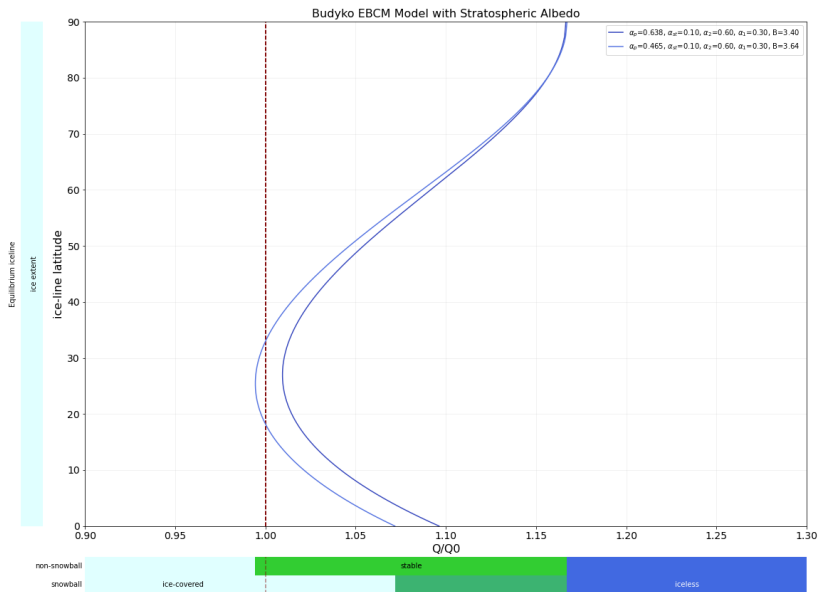
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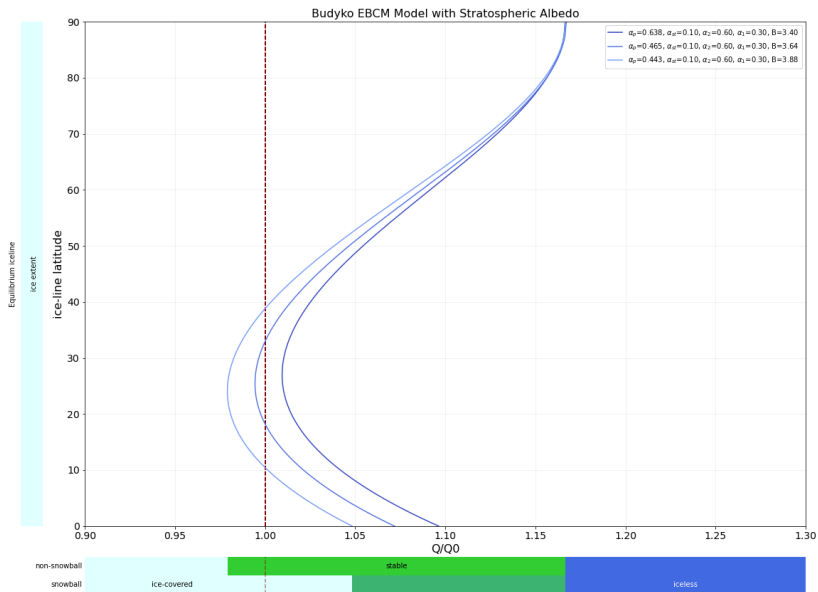
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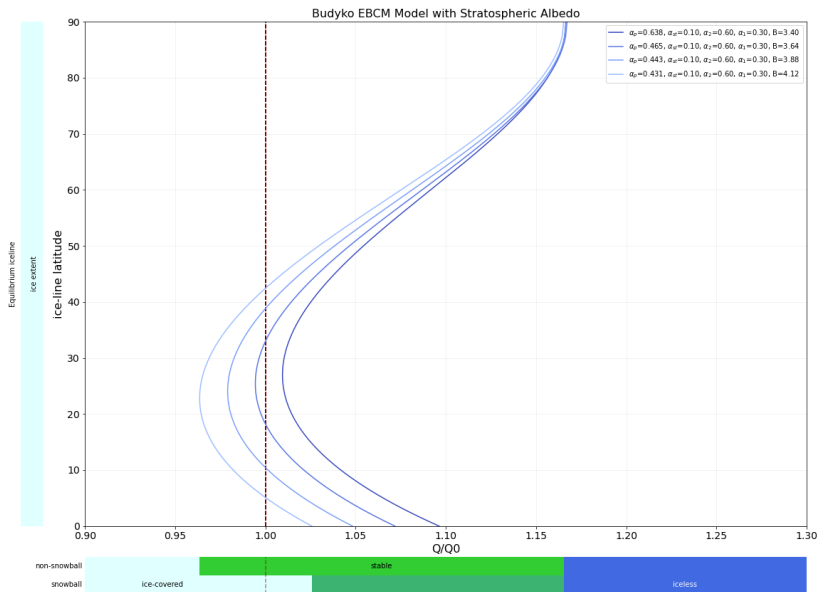
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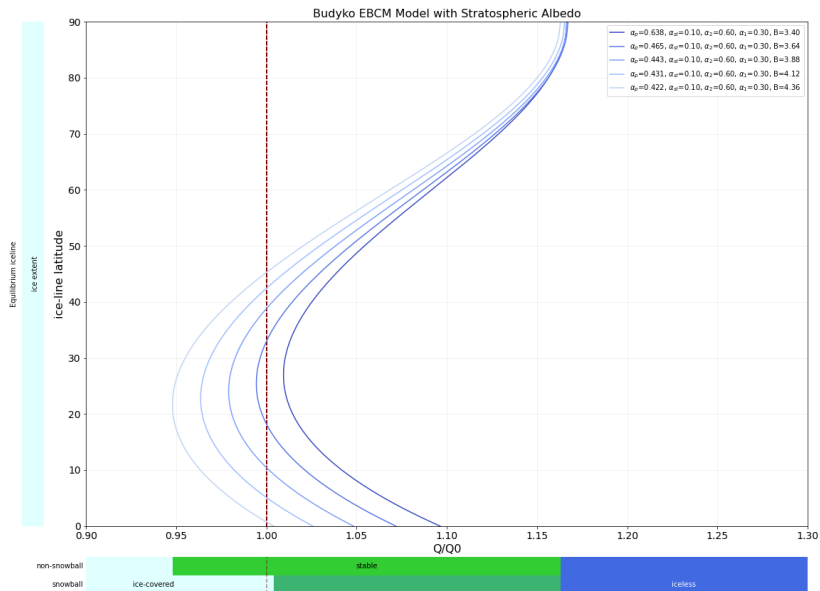
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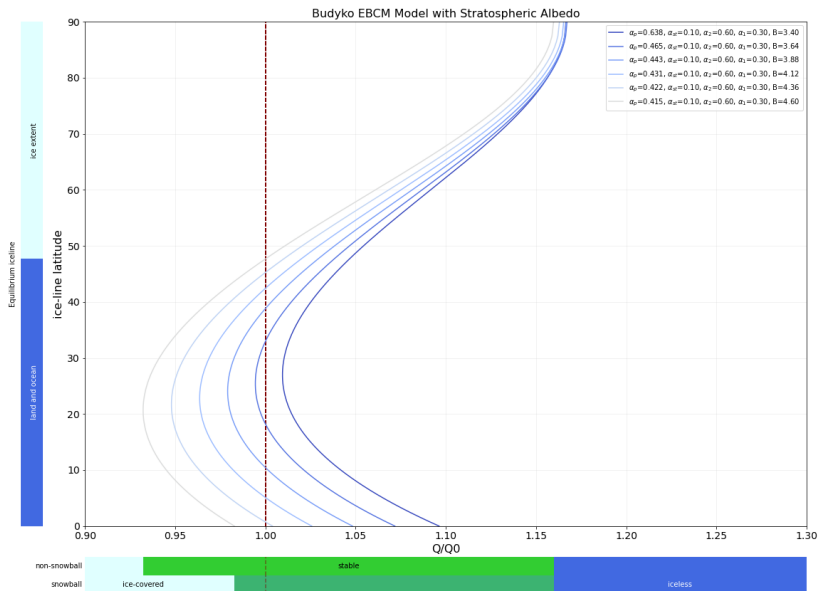
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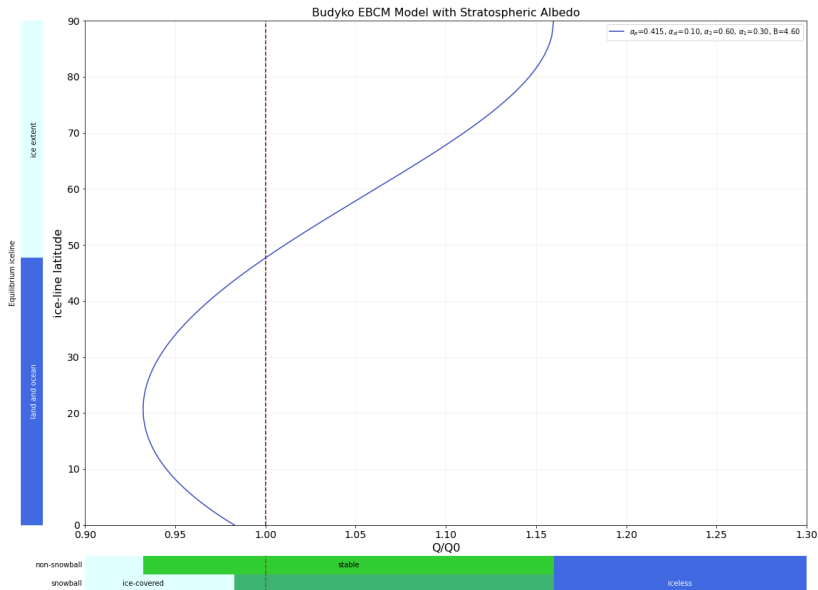
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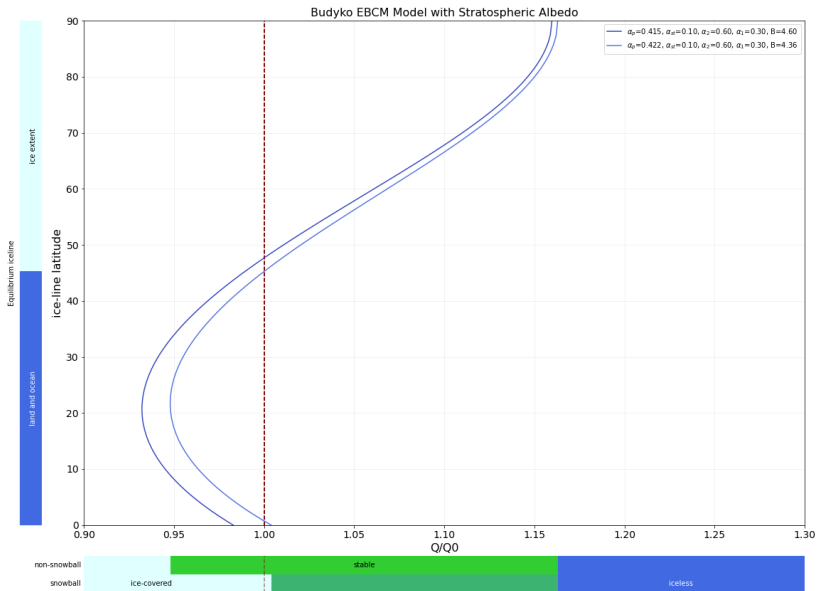
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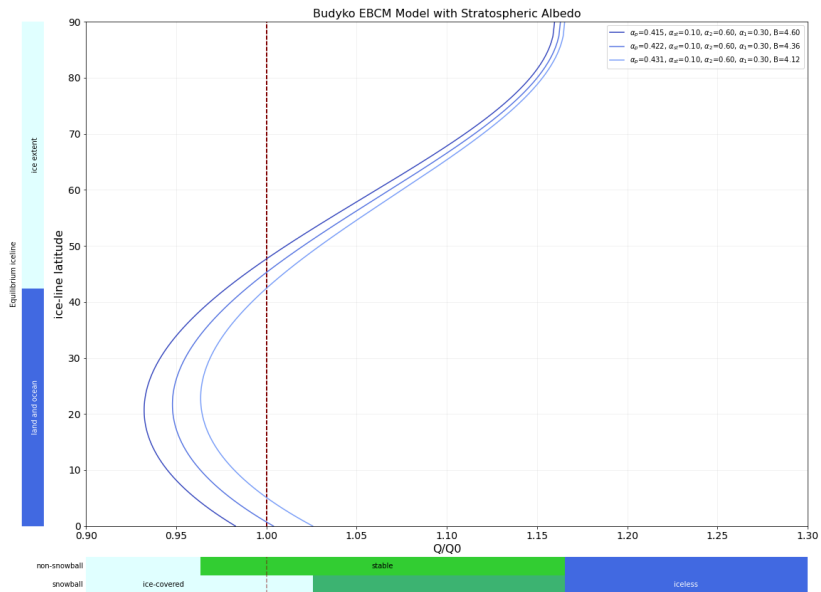
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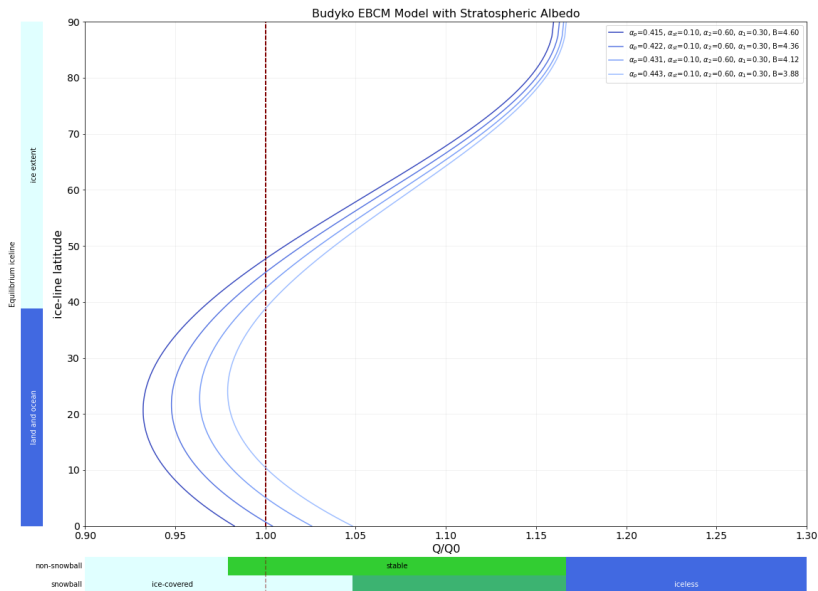
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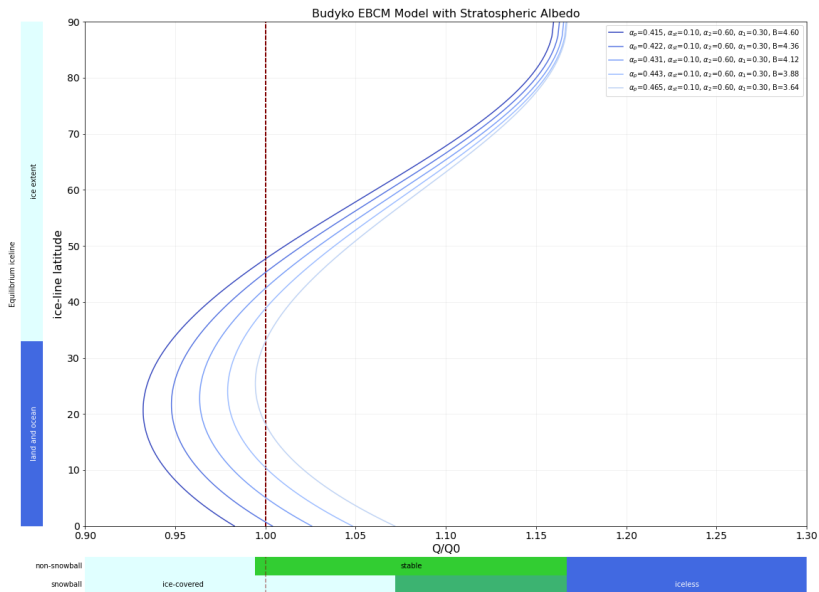
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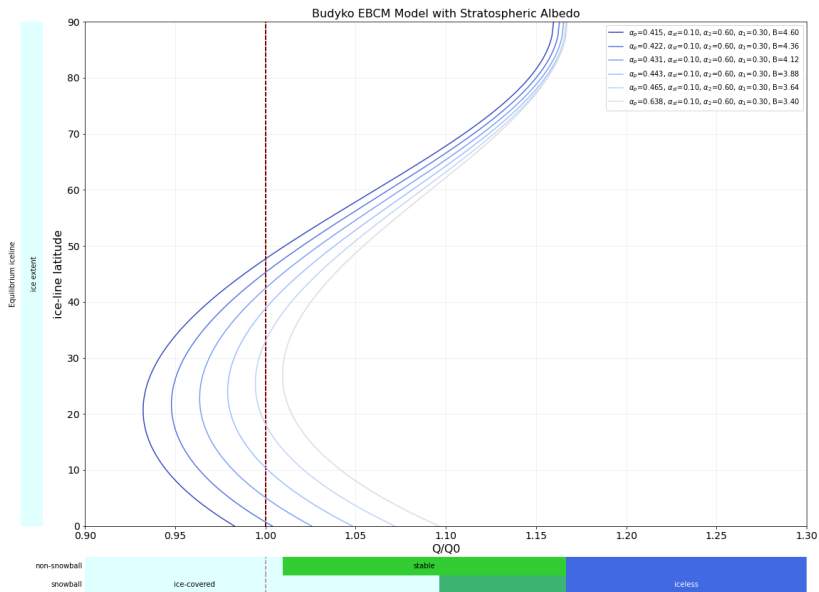
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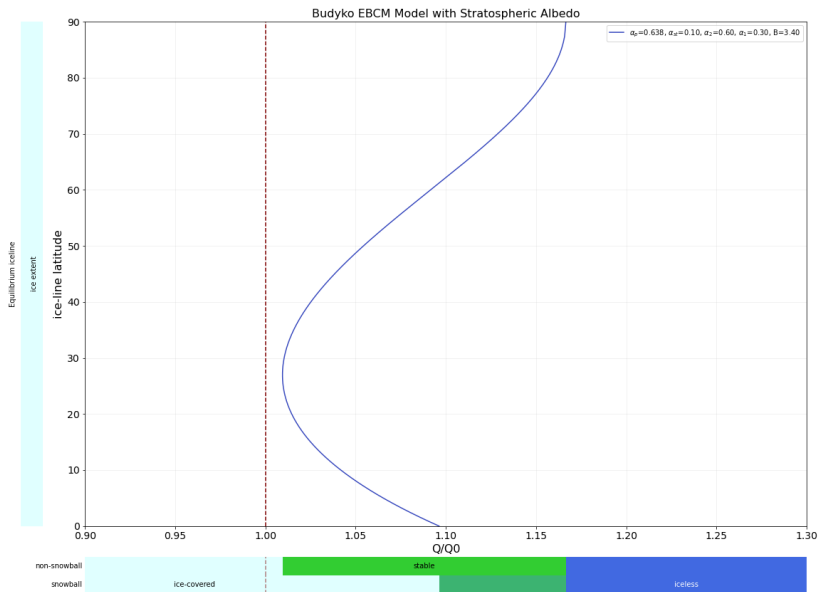
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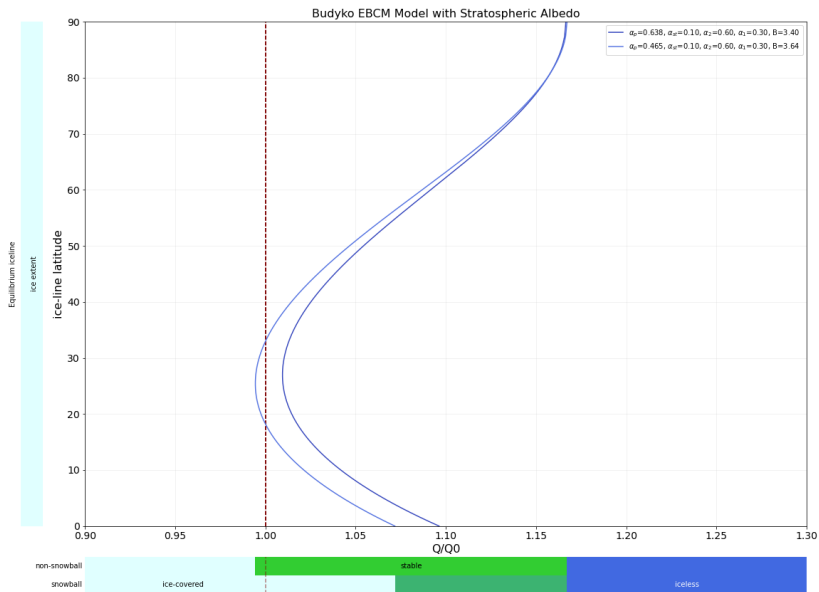
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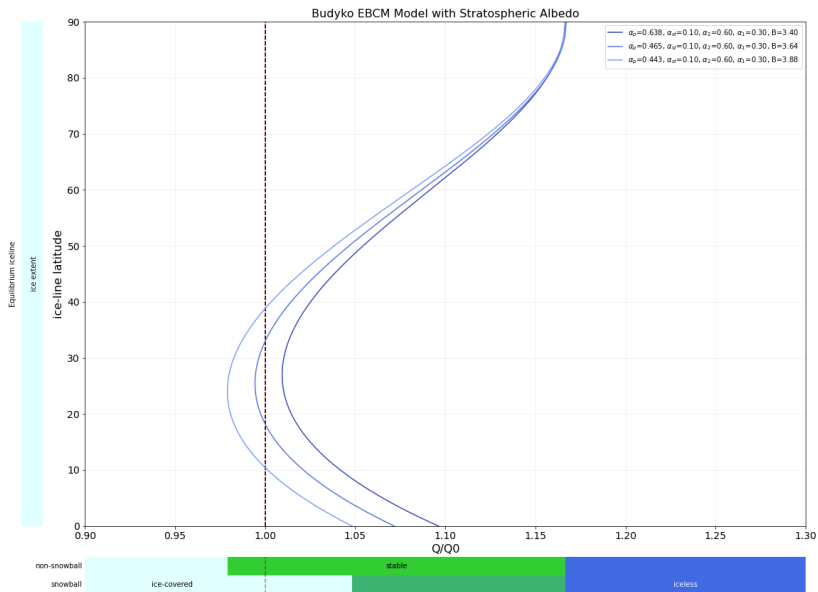
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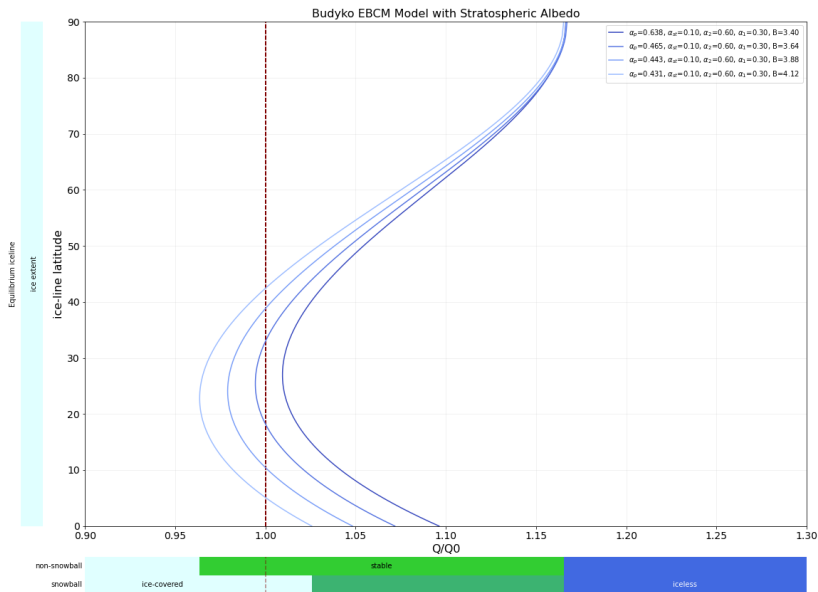
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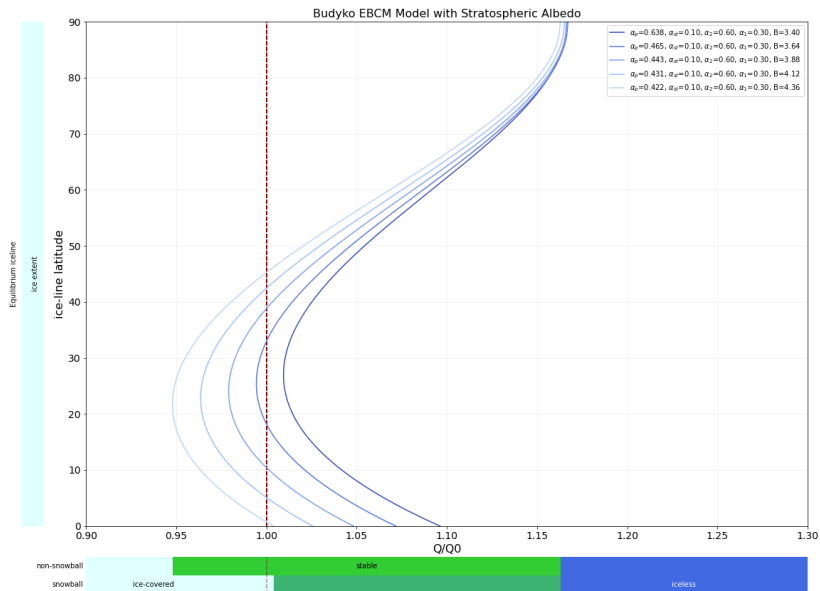
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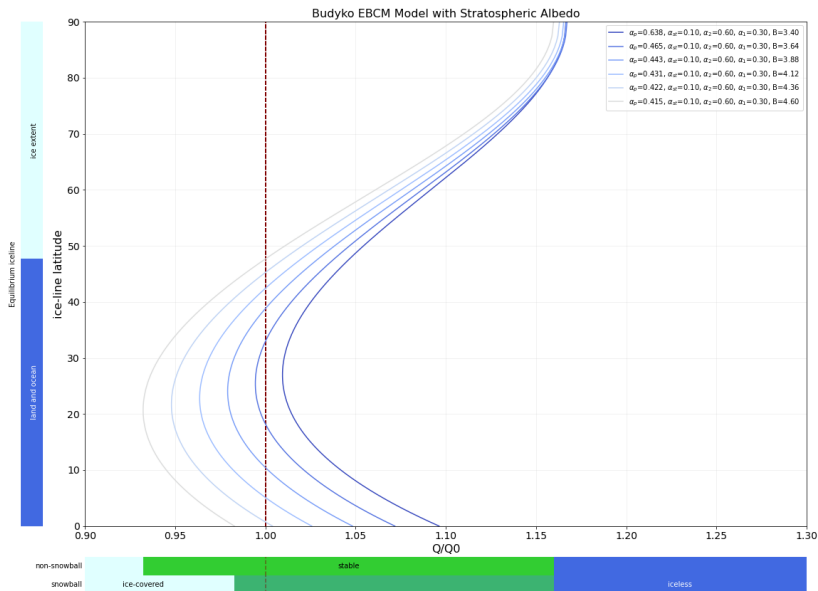
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Could albedo changes explain Neoproterozoic glaciations?

