

Investigating water stratification during the Messinian Salinity Crisis with an advection-diffusion model

Timo Millenaar

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Abstract

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1 Introduction

1.1 Models of the Messinian Salinity Crisis

Below and around the present day Mediterranean sea, vast deposits of evaporites have been discovered and described (Hsu, 1972; Krijgsman et al., 1999; Rouchy and Caruso, 2006; Roveri et al., 2014; García-Veigas et al., 2018). These evaporites were deposited during the Messinian in what is generally referred to as the Messinian Salinity Crisis (MSC). Marginal basins around the mediterranean, some of which are now found well above sea level, generally contain selenitic gypsum deposits (García-Veigas et al., 2018). These marginal basins were of shallow depth at the time of deposition, generally not surpassing 200m. In Sicily, a succession containing Halite is thought to be deposited in a basin of intermediate depth (up to a kilometre). In deeper Mediterranean basins, evaporite deposits up to two kilometres thick are found. In some locations, these deposits consist primarily of halite, while in others a gypsum-halite-gypsum trilogy is found (Roveri et al., 2014). The up to a kilometre thick halite deposit in this succession is often termed the 'Mobile Unit'. Of this Mobile Unit, only the top few metres have been drilled (Hsu, 1972). This lack of data on the deeper halite deposits leaves room for speculation about its depositional setting. Several conceptual models have been proposed to explain the origin of and developments

during the MSC, the most prominent of which can be grouped into the following categories:

1) The 'desiccated, deep-basin model', initially proposed by Hsu (1972), where drill cores taken from the top of the deep basin evaporites were discussed. These sediments were interpreted to be shallow water deposits. They proposed that a basin with a depth near that of the current Mediterranean was closed off from the world's oceans to subsequently desiccate, giving rise to extensive playa lakes that resided well below global sea level.

2) The 'shallow-water, shallow-basin model', proposed by Nesteroff (1973), based on the similar interpretation to that of Hsu (1972), that sediments in the deeper parts of the basin are of shallow water origin. Since these shallow water evaporites were deposited both in the deep basins as well as in the shallow marginal basins, they reasoned that the deep basin floor must have been brought up to a depth of 200-500 metres by vertical tectonic movements. Then a period followed where a cyclic opening and closing of the gateway would allow ocean waters to enter the basin during which evaporites were deposited. Finally these vertical tectonic movements would have brought the basin floor down to its present depth.

3) The 'shallow-water, deep-basin model', described by Rouchy and Caruso (2006), where limited inflow through narrow or shallow straits, combined with high evaporation, kept the water from fully filling the basins. This resulted,

episodically, in shallow brine seas where evaporites were deposited.

4) The 'deep-water, deep-basin model', described by [Schmalz \(1969\)](#), where dense surface water was formed as a result of evaporation. As the dense and salty water sank, it cooled, increasing its density further until it arrived at the basin floor. There it would, due to its high density, only be replaced if even denser water were to take its place. This stagnation of a dense brine would result in the precipitation of evaporites there.

With the recovery of drill cores from the deeper parts of the Mediterranean ([Hsu, 1972](#)), the 'deep-water, deep basin model' lost support as these sediments were interpreted to be deposited in shallow water conditions. A point of caution however, also noted by [Nesteroff \(1973\)](#), is that the deep basin cores represent only the terminal part of the MSC, since only the top few metres of the evaporite succession were drilled. The 'shallow-water, shallow-basin model' was not widely adopted, given the lack of a mechanism capable of inducing these large vertical movements in the Mediterranean, especially in the time span of a few million years. The 'desiccated, deep-basin model' became the widely accepted model. Later on, a model describing the relation between the marginal and deeper basins was proposed by [Rouchy and Caruso \(2006\)](#). In this model the deepest parts of the basin are, depending on the relative sea level and the size of the gateways, either filled with deep water or with a brine sea. Evaporite deposition shifts, in this model, from one basin to another. This shift is, however, in contradiction with the findings of [Krijgsman et al. \(2002\)](#), who showed that the onset of the MSC was synchronous across the Mediterranean. More recent research questioned the interpretations concerning shallow water deposition, suggesting that the data was either inconclusive or could even be indicative of a deep water setting ([Hardie and Lowenstein, 2004](#); [Roveri et al., 2001](#)). On the basis of this, and the termination of erosional surfaces towards the deeper parts of the Mediterranean, [Roveri et al. \(2014\)](#) argued that deep water conditions may have persisted throughout the MSC. A variation of the 'deep-water, deep-basin model' was proposed by [García-Veigas et al. \(2018\)](#), where geochemical indicators in evaporites were examined. These indicated that the Mediterranean was highly stratified, with high salinity and dense, anoxic bottom waters. This led them to propose a variant of the 'deep-water, deep-basin model' where deposition of the marginal gypsum would be synchronous with the

deposition of halite from a dense brine in the deeper parts of the Mediterranean. In order to investigate the plausibility of the formation of a deep brine during the MSC, an analogue is made between the present day Dead Sea and the MSC. This analogy will be tested using a linear diffusion-advection box model.

1.2 The Dead Sea as reference

The Dead Sea is a present day land-locked, hypersaline basin. The physical processes in this basin are modeled and tested against measurements by [Heiden et al. \(2017\)](#).

1.2.1 Vertical stratification in the Dead Sea

Before 1978, the basin was stratified and did not mix. The top of the basin contained a warm and fresh epilimnion. Below that, a cooler and more saline hypolimnion resided. Increasing irrigation combined with a changing climate lead to a steady decrease in fresh water runoff. In 1978, a threshold was reached where the density distribution of the water column became unstable. Since then, the water column undergoes vertical mixing during winter, resulting in a holomictic situation during winter. [Heiden et al. \(2017\)](#) used a linear diffusion model to replicate this annual mixing pattern. From this study, a conceptual model describing the annual mixing was formulated. In this model, as both the temperature and salinity of the epilimnion increase during summer, the water column remains stable since the high temperature of the epilimnion keeps it more buoyant than the hypolimnion. In this scenario, the epilimnion can be more saline than the hypolimnion, but still be stable. As the epilimnion cools however, it remains more saline than the hypolimnion and the density of the epilimnion increases until the water column becomes unstable and mixing occurs.

1.2.2 Applicability of the Dead Sea model to the MSC

While a permanent inflow of solutes from a parent ocean is required to explain the vast volume of evaporite deposits ([García-Veigas et al., 2018](#)), a connection the likes of which is absent from the present day Dead Sea, the Dead Sea model will be used as a first order analogue for the MSC. In both systems, we are dealing with a (largely) isolated, hypersaline basin where evaporite deposition is partially controlled by climatic oscillations ([Manzi et al., 2012](#)). Among present day hyper saline basins, the Dead Sea is one of the most prominent, making it one of the

better candidates to study hyper saline lake dynamics. This study adds to the model of [Heiden et al. \(2017\)](#) to apply it to a marginal basin - deep basin configuration of the MSC.

1.3 Marginal basins as agent for stratification

The model of a stratified Messinian Mediterranean as presented by [García-Veigas et al. \(2018\)](#), illustrated in their Figure 12, contains a dense, anoxic brine at the bottom, normal marine waters at the top and a gypsum saturated layer in between. This does not combine well with the holomictic Dead Sea model, where the top layer is generally the most saline part of the column. A possible solution to bring these scenarios together, inspired by [Wählin and Cenedese \(2006\)](#), is to involve the marginal basins of the Mediterranean. Water from the top of the main Mediterranean basin could flow into the marginal basins. Here, evaporation will increase the salinity of this water. As the water cools in the fall, like in the Dead Sea model, the water column in the marginal basin would mix. This would then result in an increased density of the water at the bottom of the marginal basin. While this effect would likely also take place in the deeper Mediterranean column, the increased density of the bottom water would be more pronounced in the marginal basins, precisely because of their shallower depth. If the density in the bottom of the marginal basin is higher than that in the bottom of the deeper basin, water could, due to buoyancy, flow from the marginal basin into the deeper parts of the main basin. There it would only be replaced by even denser water. This hypothesis resembles the 'Deep water, deep basin model' as described by [Schmalz \(1969\)](#), but with marginal basins essentially acting as evaporative pools.

In this study, a linear diffusion-advection box model, consisting of a deep and a marginal basin is used as a first order approximation to investigate the possibility of stratification in the Mediterranean basin during the MSC.

2 Methods

2.1 Model setup and boundary conditions

2.1.1 The box model

In order to investigate the influence of a shallow marginal basin on the conditions at which stratification could occur in a deep water basin, a box model was created. The box model ex-

pands on the linear diffusion code as used by [Heiden et al. \(2017\)](#). In this box model, the two basins are each represented by a vertical column. For each of these columns, diffusivity and advection calculations are solved for both temperature and salinity. From these, the density is inferred. The top nodes of the two columns are connected via the source-sink term. The amount of these connected nodes can be varied. When the bottom nodes of the marginal basin column are denser than those at the bottom of the deep basin, flow is initiated from the bottom of the marginal basin to that of the deep basin. At this time, flow from the top of the deep basin to that of the marginal basin is also initiated, as is the advection within the columns. Advection goes up in the main basin and down in the marginal basin.

2.1.2 Boundary conditions

Initially, the temperature and salinity in each entire column are set to a user modifiable value, resulting in straight vertical profiles. A flux on the top node is then initiated. This flux is sinusoidal and is linearly proportional to the inverse of the diffusivity in the top two nodes. The sinusoidal flux for salinity can be assigned an above zero mean, resulting in net evaporation. The bottom node is not subject to any boundary conditions but is determined based on the node above it. At first, mixing, source-sink and advection are not active in the model. There are two types of forcing in the model, a heat flux on the temperature and an evaporative flux on the salinity. The former is dependent on the surface water density and the heat capacity, the latter is dependent on the salinity of the surface water. Both are sine functions with a period of a year that affect only the top node. During the first four years, the amplitude of the forcing will be increased from zero to 100%, after which the amplitude will remain 100%. After the first ten years, mixing, source-sink and advection will be turned on. I will refer to this ten year period as the spin-up period and it ranges from April year 0 until April year 10. Most scenarios in this research contain a marginal basin of 150m, for most marginal basin were shallower than 200m. While the deep basins are often over two kilometres deep, the scenarios in this model generally contain a deep basin of a mere 800m, for this reduces run time while conserving the general outcome of the model.

2.2 Diffusion discretization

The principal diffusion equation for a diffusive physical quantity (Q) in differential form, with

advection and source sink is shown in equation 1.

$$S = \frac{\delta Q}{\delta t} + w \frac{\delta Q}{\delta z} - D \frac{\delta^2 Q}{\delta z^2} \quad (1)$$

Where \mathbf{w} is the advection velocity in ms^{-1} , D is the diffusivity in ms^{-1} , Δt is the incremental timestep in s , Δz is the spatial increment in m and S is the source-sink term in $[Q]s^{-1}$. This is discretized to shown to equation 2, formulated after the explicit Forward Time Central Space (FTCS) method as described in Slingerland and Kump (2011) p. 143.

$$S = \frac{Q_i^{n+1} - Q_i^n}{\Delta t} + w \frac{Q_{i+1}^n - Q_{i-1}^n}{2\Delta z} - D \frac{Q_{i-1}^n - 2Q_i^n + Q_{i+1}^n}{(\Delta z)^2} \quad (2)$$

Here, Q_i^n refers to the value of a diffusive physical quantity (Q) for time step n at depth i . Rewriting equation 2 yields the desired diffusive physical quantity (Q) for the following timestep at depth i , Q_i^{n+1} as shown in equation 3.

$$Q_i^{n+1} = Q_i^n - \frac{w\Delta t}{2\Delta z} (Q_{i+1}^n - Q_{i-1}^n) + \frac{D\Delta t}{(\Delta z)^2} (Q_{i-1}^n - 2Q_i^n + Q_{i+1}^n) + S\Delta t \quad (3)$$

This method of discretization does bring with it the following conditions:

$$\Delta t \leq \frac{\Delta x^2}{2D} \quad \text{and} \quad \Delta t \leq \frac{2D}{w^2} \quad (4)$$

These conditions are used to dynamically adjust the size of the time step, according to 0.9 times the lowest Δt between the various basins and criteria.

2.3 Column stability

The model runs with a standard, minimal diffusivity that is always present and affects both entire columns. The diffusivity can be set individually for temperature and salinity. When part of the column is unstable, which can be the entire column, the unstable part is assigned a diffusivity orders of magnitudes higher to imitate the mixing of the column. This high diffusivity is assigned to both temperature and salinity.

In the Dead Sea model as described in Heiden et al. (2017), the presence of an unstable column is determined by assessing if the top node is more dense than the one below it. For their model this is sufficient, but this creates the possibility of a dense water body at a certain depth that is considered stable by the model if there is less dense

water above it. Therefore, an adaptation from Heiden et al. (2017) is required, where the bottom of the column is also checked for instability. Two methods are implemented, between which a choice has to be made before the run starts. In one method, the column is checked for instability both from the top down and from the bottom up. In the second method, any location is checked for instability and assigned an appropriate diffusivity. The consequences of these methods will be discussed in section ??.

2.4 Density calculations

The density at a given node is determined using the equation of state as described in Ivanov et al. (2002). They calculate the water density as a density anomaly from the reference density of $1e3 \text{ kgm}^{-3}$. In this research, we will show the density as absolute values, therefore adding $1e3$, resulting in equation 5. Their values, determined for the Dead Sea, will be applied to the MSC in this research.

$$\rho = \rho_{ref} + r_0 - \alpha T + \beta S \quad (5)$$

Where $r_0 = -15.654$, $\alpha = 0.4309$ and $\beta = 0.936$, as determined by Ivanov et al. (2002). ρ is the density in kgm^{-3} , ρ_{ref} is the reference density of $1e3 \text{ kgm}^{-3}$, T is the temperature in $^{\circ}C$ and S is the salinity in kg m^{-3} .

2.5 Source-sink and advection

The volume of water that moves every time increment through the strait that connects the two basins, is determined by multiplying its width, depth and the lateral flow velocity through the strait. When the density of the bottom waters in the shallow basin is greater than that of the bottom waters in the deep basin, the flow of salt water from the shallow basin to the deep basin is initiated. The total amount of salt that is moved in this exchange is determined by multiplying the volume of water moved, with the salinity of the source. The impact this amount of transported salt has on a basins is dependent on the total volume of the basin in question. This exchange of salt is implemented through the source-sink term (S), as shown in equation 2. S is calculated using equation 6 and is either positive or negative depending on whether it represents a sink or a source, respectively.

$$S = \pm h_{source} \frac{z_{strait} x_{strait} U_{lat}}{A_{sink} z_u} \quad (6)$$

Where h_{source} is the salinity of the source in kg m^{-3} , U_{lat} is the lateral flow velocity in ms^{-1} , A_{sink} is the surface area of the sink basin in m^2 ,

z_u is the vertical extent of the flow in m and x_{strait} and z_{strait} are the strait width and depth, respectively, in m .

The physical replacement of the bottom waters as a result of the lateral flow is captured in the advection, where advection is set to be upward in the deep basin and downward in the shallow basin. The value of the advection is equal to the volume of lateral flow per second, divided by the surface area of the basin in question.

3 Understanding the model

In order to get a grasp on the basics of the model, some results from simplified scenarios are shown in this chapter. The basic boundary conditions are discussed as well as the relation between the temperature, salinity and density.

3.1 Forcing

The last year of the ramp period is shown in figure 1. Here, the amplitude of the forcing is at its maximum. Only the diffusion and the boundary conditions are active in this stage. The mean, as well as the amplitude of the forcing differ between the temperature and the salinity, but their general shape is identical. Both the temperature and salinity of the surface water increase from march until June and decrease from September until December.

During the ramp period, the temperature, salinity and density profiles show a symmetrical overall shape of the plot (see figure 1). This shape follows from the perfectly sinusoidal cyclicity of the forcing that is centered around zero. Below about forty metres there is no influence of the forcing, for the background diffusion is not quick enough to reach that depth within a year.

3.2 Density profile

The shape of the density plot is identical to those of the temperature and salinity, but it shows a reverse correlation, where the density of the surface waters is highest in January and February and lowest in July and August. The behaviour of the density profile is fully dependent on the values of the temperature and salinity at that depth as can be deduced from equation 5 and can also be reversed. This is shown in figure 2, where the salinity has a diffusivity that is an order of magnitude lower than that of the temperature. In this scenario, the inverse correlation of the density holds from the bottom up

to a depth of about seven metres, where a turning point is reached. For the top few metres, the correlation is no longer inverse.

The lower diffusivity of the salinity compared to that of the temperature, has two significant consequences. The first of these is the reduced depth to which the salinity forcing can penetrate within a given year. The second consequence is the dependence of the forcing on the diffusivity in the model. Lower diffusivity results in higher forcing on the salinity, as mentioned in section 2.1.2. A higher forcing, results in a higher amplitude in the salinity values. This means that, during winter, the surface water is less saline than for example that in figure 1, resulting in a stable density distribution during winter and a regular linear relation between the density and the temperature and salinity. Since this effect only penetrates the water column for the first few metres, deeper in the column the temperature has more impact and the inverse balance of the density curve as seen in figure 1 is restored. This can be deduced from the equation of state as well (equation 5), where the density decreases in a given location if αT increases more rapidly than βS or βS decreases more rapidly than αT . The latter is the case in figure 2, where the value of βS decreases significantly faster with depth than αT . If the amplitude of evaporation forcing is significantly higher than the amplitude of the heat flux forcing, the salinity of the water will control the density profile. If that is the case, and assuming there is no net evaporation, the surface water lacks the salinity to become dense enough during winter to become unstable, while the water becomes so saline during summer that the water column becomes unstable, despite its high temperature.

3.3 Transition

3.3.1 Results

The transition around the tenth year is shown in figure 3. From April year 10, convection is turned on. Around September, the top node is the same density as the one below it. From this moment onward, mixing takes place. In October, the top 30 or so metres are mixing, after which this is the case for the entire column. From march onward the mixing has stopped. This pattern of mixing only in winter results in an asymmetric profile, as shown in figure 4. During the winter mixing, the profile is almost straight, though slightly inclined due to the finiteness of the diffusivity constant. From April until September, figure 3 resembles figure 1. As the fall progresses, an unstable column develops at the surface and increases in depth as

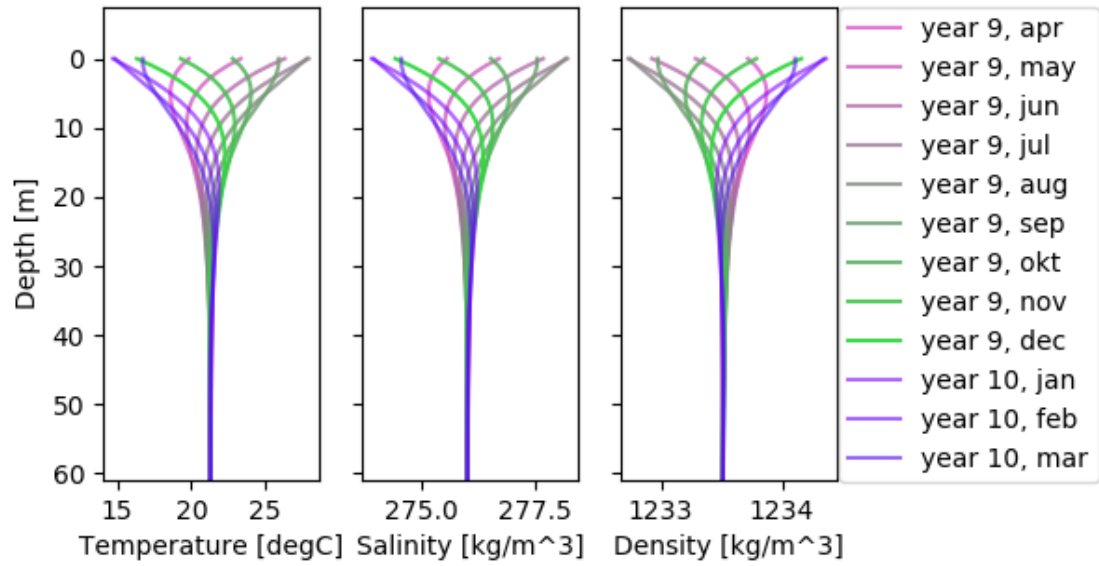


Figure 1: Last year of the 10-year Spin-Up period.

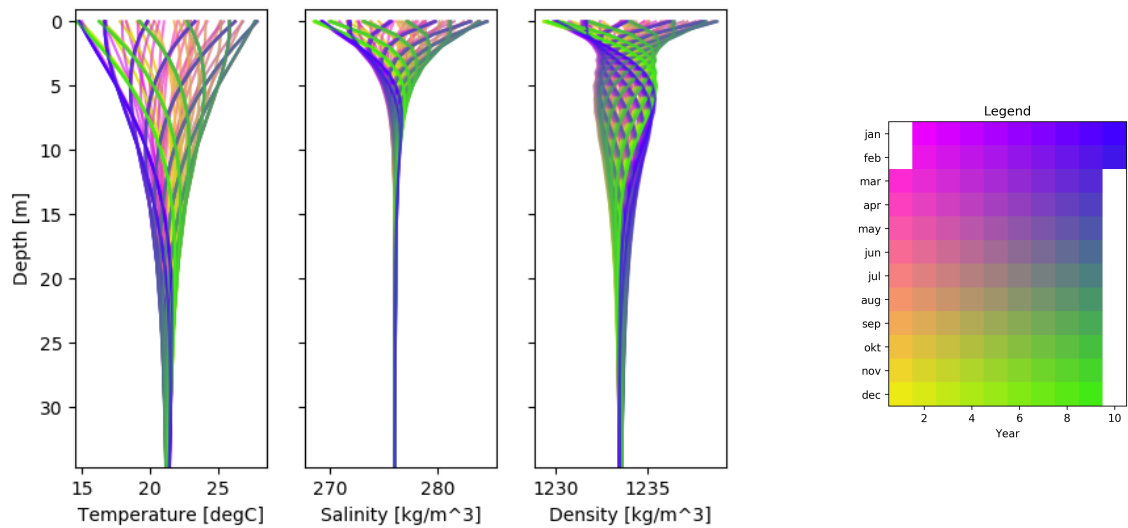


Figure 2: Spin-Up period with a salinity diffusivity one tenth of the temperature diffusivity

the surface water cools.

3.3.2 Discussion

Concerning the transition, as shown in figure 3, the mixing should start in April. This is namely the moment when convection is switched on, and the density in the top node is higher than the rest of the column. The mixing does however not occur until October, as read from figure 3. The cause of this is that the code begins the instability check by checking whether the top node is more dense than the one below it. Since, in April, year 10, the top node is as dense as the one below it, is considered stable. Even if the top twenty metres are more dense than the water below it. This situation generally only occurs during the transition, for the top nodes are always the first to move. The exception to this is a source-sink pulse, pulling salt from the bottom of the marginal basin, without directly affecting the water in the middle of the column. In order to have mixing in the bottom of the column to resupply salt to the bottom of the marginal basin and restore the density instability, a more rigorous check for instability is required, mentioned in section 2.3. When checking the column for instability both from the top node down and from the bottom node up, the salt deprivation in the bottom of the marginal basin after flow is solved in this way. The artifact that arises during the transition period as shown in figure 3 does however remain. The alternative method was to check each individual node for stability. This method solves both the artifact during the transition and the salt deprivation in the bottom of the marginal basin. In practice, the whole column check can take up to an order of magnitude longer to run. This is not only due to the increase in effort to check for the whole column, but mainly due to the limiting factors shown in equation 4, where the highest diffusivity in the column determines the size of the time step. A few scattered unstable nodes can therefore significantly impact the run time. An unforeseen issue arises when the whole column check is used, where certain scenarios result in an ever increasing salt content of both columns. An example of this is shown in appendix. A scenario run with the input parameters as shown in appendix and a marginal basin depth of 150m will produce a stable outcome, but a scenario with 149m marginal basin depth results in instability like that shown in appendix. This indicates that there is some stability threshold for the code, but this has not been narrowed down yet. When performing a column stability check from the top down and bottom up, this scenario can also occur, but for

this to happen the marginal basin would need to be a mere 25m or so, with high U_{lat} .

3.4 Summer mixing

The stability of the column is mainly determined by the density of the surface waters and therefore, ultimately determined by the balance between αT and βS (equation 5) at the surface. By altering either the heat flux or the evaporation, the stability behavior of the column can be investigated.

3.4.1 Results

Figure 5 shows the results of a run with an evaporation amplitude of 0.6 rather than the standard 0.3. The mixing now occurs in summer while the column is stable in winter. The overall shape of the profile is rather the same, but the density is now almost a direct image of the temperature and salinity rather than a mirror image, as is the case in figure 4.

3.4.2 Discussion

When αT is dominant over βS , as is the case in figure 4, the column mixes in winter, for the density is inversely dependent on the temperature. In figure 5 however, βS is dominant over αT . The subsequent extra saline surface waters in summer, result in an unstable column, despite its high temperature. In winter the opposite is true, where the surface waters are fresh enough to remain stable despite the low temperatures.

The periods of mixing in the dead sea are determined by the temperature profile. For lack of knowledge on the periods of mixing during the MSC, the assumption is made here that this was also the case for the MSC. Keep in mind that the foundation for this assumption is not very solid. Periods of low river discharge into the Mediterranean could even have resulted in almost year-round mixing with high salinity creating instability in summer and low temperatures creating instability in winter.

3.5 Source sink and advection

3.5.1 Results

A result of a standard model case (appendix ??) with no advection and an increased exchanging flow from $1e-5$ SV to $5e-3$ SV is shown in figure 6. The result of the same case but *with* advection is shown in figure 7. In both instances, part of the column mixes during fall and the entire column mixes from January to March, in both the deep and marginal basins. The marginal basin of figure 6 very strongly resembles figure 4, as

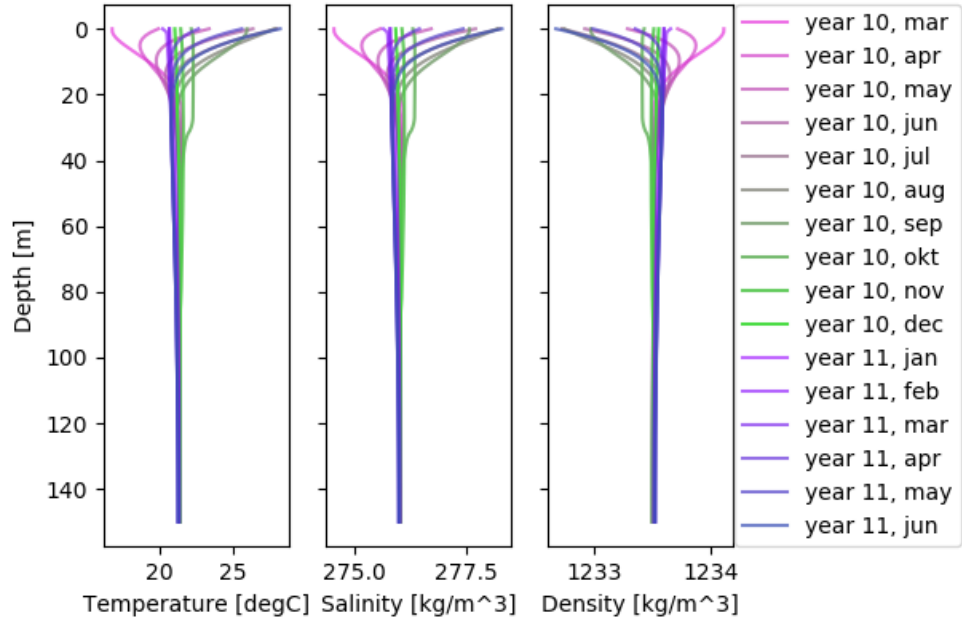


Figure 3: The transition from the Spin-Up period. The possibility of mixing is activated in April. This data was generated using an instability check in the top node instead of a whole column check for instability.

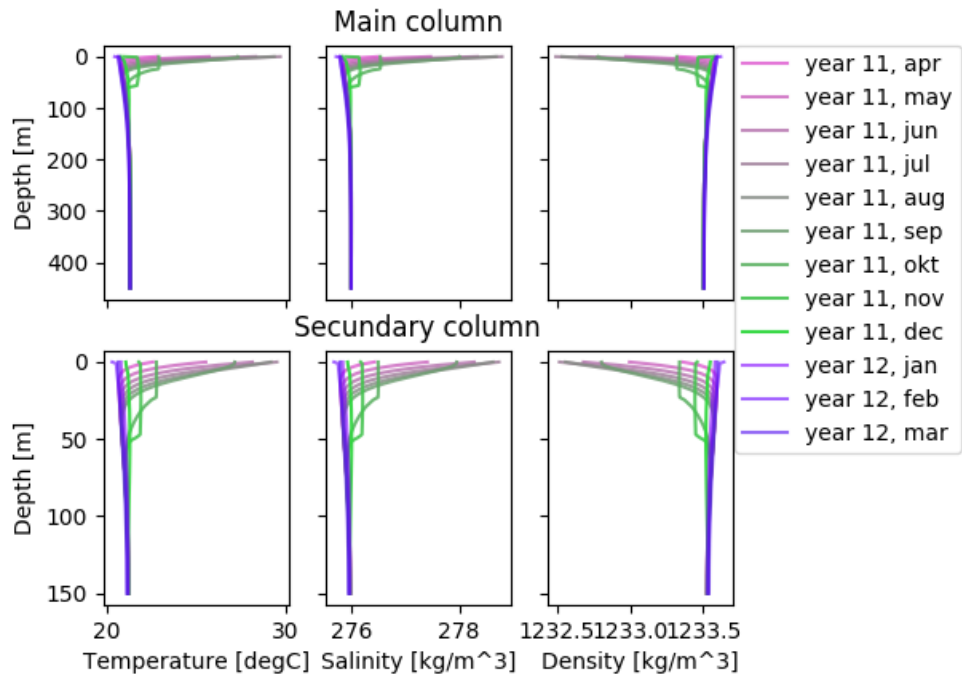


Figure 4: A typical year of a run with convection.

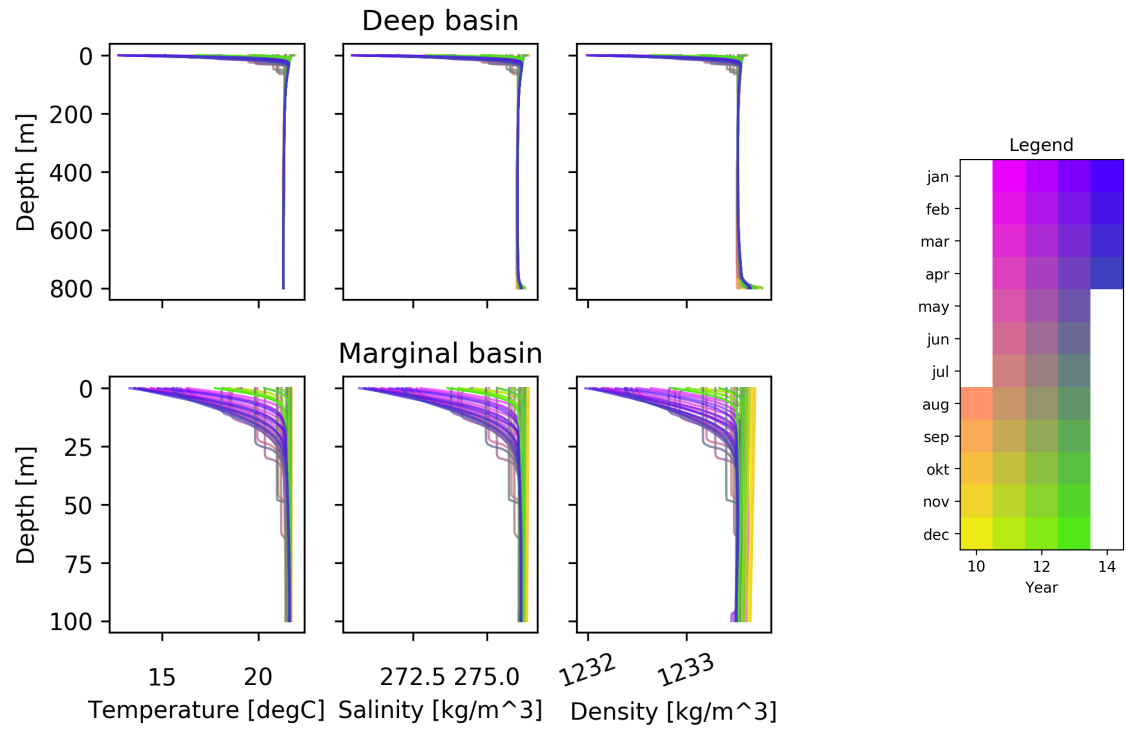


Figure 5: Model run with advection, a flow of $5e-5$ Sv and a marginal basin depth of 100 metres. The evaporation amplitude was doubled, inducing mixing in summer.

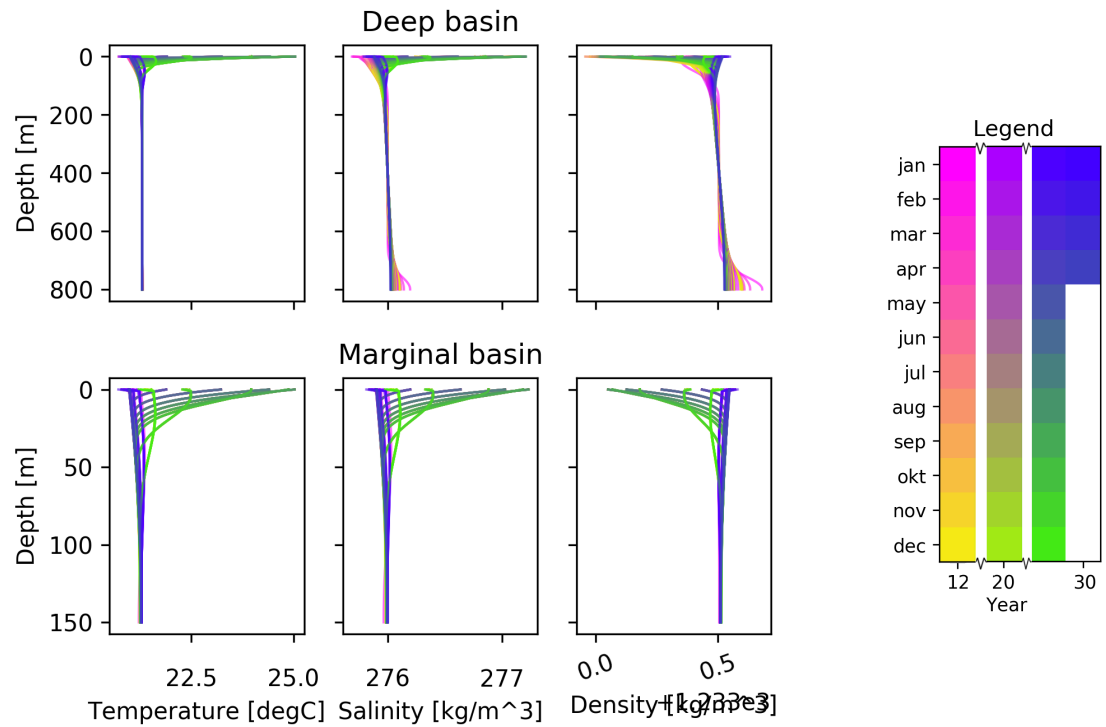


Figure 6: Model run with source-sink and no advection. The situation shown is right after flow between the two basins occurred.

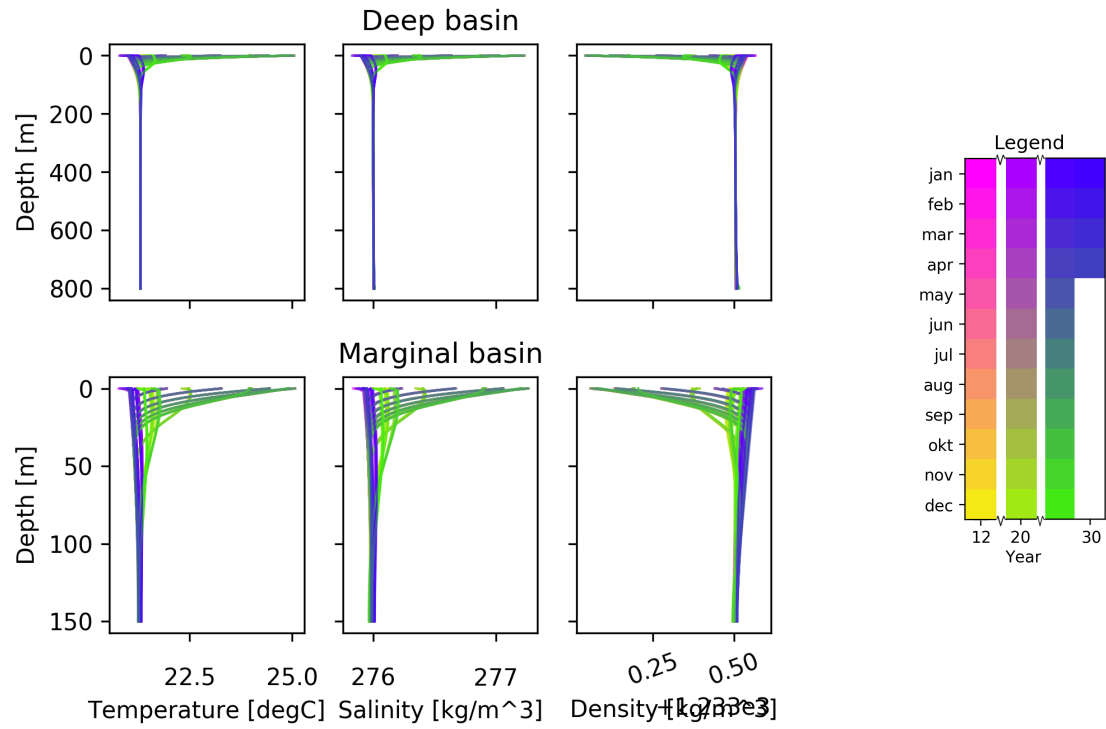


Figure 7: Model run with advection. The same input parameters as the run shown in figure 6 were used.

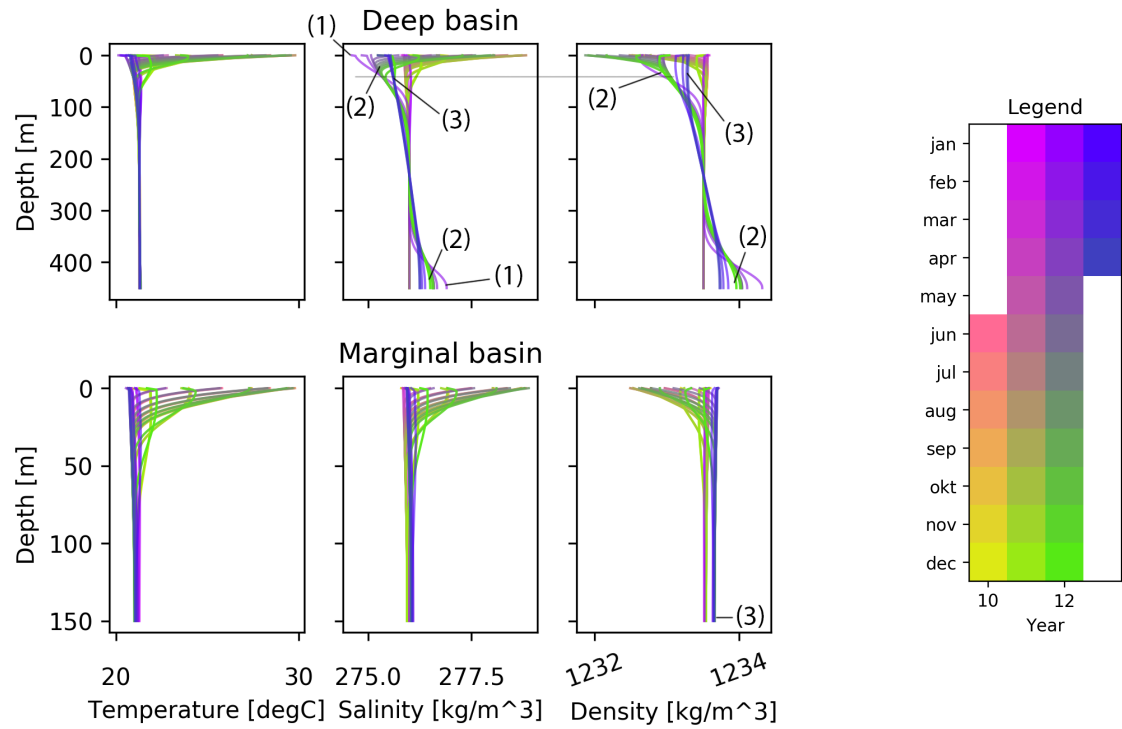


Figure 8: Model run with advection. (1): initiation of the exchanging flow (2): slow decrease of anomaly (3): mixing.

does the deep basin with the exception of the increased salinity at its bottom. Only one pulse of exchanging flow occurred. This pulse was large enough to create a clear salinity access at the bottom of the deep basin in one exchange. In a case with advection, the dense water body at the bottom of the deep basin is generally a lot less pronounced or not present at all, as is the case in figure 7. A model run with advection and a flow of 40 Sv is shown in figure 8. With this large exchanging flow, the workings of the code can be easily analyzed in report format, for the effects are large enough to visualize without zooming.

3.5.2 Discussion

An initial pulse of exchanging flow occurs because the bottom of the marginal basin experiences slight changes in temperature and salinity whereas the bottom of the deep basin is too deep to experience changes created through surface forcing. The exchange between the marginal basin and the deep basin, shown in figure 6, slightly lowered the total amount of salt in the marginal basin, for the water that flowed out was more saline than the water that flowed in. While salt is directly resupplied to the bottom of the marginal basin due to the unstable density configuration, the excess salt in the bottom of the deep basin created a stable configuration. This excess salinity only decreases due to the background diffusion. When the excess salinity has been reduced to its original state due to the background diffusion, and the bottom of the marginal basin is sufficiently replenished, a new pulse could occur. However, since the first pulse slightly altered the average salinity of both basins, it can take years for the marginal basin bottom waters to once again become more saline than the bottom waters of the deep basin. In fact, if the difference in average salinity is larger than the yearly difference due to forcing, no second pulse will ever occur, for no other factors influence the density distribution in this model scenario. Realistically, this scenario is impossible for it would create space problems due to an incomplete flow cycle caused by the lack of advection. The scenario shown in figure 7 was initialized with the same conditions as the one shown in figure 6 except that, during the periods of exchange, advection was allowed. As is evident when comparing the two figures, advection completely removed the effect of the source-sink term. Any water that arrived in the deep basin was directly being advected up, thus not allowing for a dense water body at the bottom of the deep basin to form. In order to investigate how the model handles these advection

situations, the model was run with a flow of 40 Sv. This amount of volume is in no way realistic but makes it easier to investigate the workings of the model. Figure 8 shows an annotated plot of this model run. This figure shows that right after the initiation of flow (1), the deep basin is fully stable with denser than average waters at the bottom and the least dense water at the top. Most of any negative density anomaly at the base of the marginal basin is directly resupplied by mixing. From the initiation of flow (1) until the next mixing during winter (3), the salt anomalies dissipate slowly due to diffusion (2). When the columns mix in winter, the bottom waters of the deep column are unaffected due to their stable state. During this mixing period, some salt is resupplied to the bottom of the basin, leading to a small jump in the density plot. Aside from the large amount of flow required for this effect to be visible, an issue becomes apparent when analyzing the results. Due to the technical choice of imitating the flow with the source-sink term, combined with this large flow of 40SV, the bottom waters of the deep basin become significantly more saline than the bottom of the marginal basin ever was.

In order to obtain a dense saline brine at the bottom of the deep basin, the bottom of the marginal basin clearly needs to be significantly more saline than is achieved through the parameters used in the runs shown so far. This could be done by increasing the total amount of salt in the marginal basin through net evaporation or by allowing for faster transport of salt to the bottom of the marginal basin. The latter can be achieved by decreasing the basin depth or increasing the amount of diffusion during mixing. A combination of these is likely required to achieve the sought effect. These factors will be discussed in section 4.

3.6 Unintended increases and decreases in the total salt and heat content

The two types of column stability check, mentioned in section 2.3, reveal an error in the code. The choice between a whole column check or a bottom up/top down check should be one based on model efficiency and desired detail. Instead, it appears that the total amount of salt and heat in the system systematically increases every year when the whole column check is used, and it systematically decreases every year when the top down/bottom up check is used. The magnitude of this error seems to be dependent on the ratio between the background diffusivity and convective diffusivity. The problem affects the salinity more than it does the temperature, though the problem systematically occurs in both. For a

standard run, the error seems to be around -0.2 kg/m^3 salt per year when the top down/bottom up check is used. This is a significant decrease. When the whole column check is used, the error seems to be around 0.05 kg/m^3 salt per year. Since the error is smallest when using the whole column check, this method will be used in this research, though the effects are still rather significant when long time spans are investigated. The root cause of this problem has not yet been discovered.

4 Important factors on the formation of a deep brine

As discussed in the previous section, the most relevant parameters are those that greatly influence the amount of salt in the bottom of the marginal basin. Likely candidates are the marginal basin depth, the background diffusion and net evaporation. In order for a saline marginal basin to have a significant impact on the salinity of the deep waters, the flow between the basins would need to be sufficient but should not reach unrealistic proportions, as is the case in figure ???. When the exchanging flow is small, the comparative volumes of the basins needs to be such that the total amount of salt in the marginal basin is not negligible when compared to that of the deep basin. These factors and their impact on the possible existence of a deep salt brine will therefore be investigated in the following sections.

4.1 Results

The standard run is one with with all features activated and a 150m deep marginal basin. The surface area of the marginal basin is an order of magnitude smaller then that of the deep basin and the connecting flow is $1\text{e-}5 \text{ Sv}$. The full list of parameters is shown in appendix ??

4.1.1 Marginal basin depth

Figure 9 shows the results of a standard case with a 60m deep marginal basin. A minor increase in salinity has been achieved. Incremental steps are made towards the redistribution of salinity, rather then the one pulse in case 6.

4.1.2 Background diffusivity

Figure 11 shows the results of a standard run and a background diffusivity of $1\text{e-}3.0$ **CHECK IF CORRECT VALUE**. The depth to which

the surface forces penetrate, as well as the maximum deviation from the average temperature at the surface differ from the previous scenario. The bottom of the maringal basin experiences significance variation throughout the year. A very small increase in salinity is observed in the bottom of the deep basin, of which the largest pusle occurred in the first year after the spinup.

4.2 A depth and diffusivity hybrid

4.2.1 Net Evaporation

Figure 11 shows the results of a standard run and a net evaporation rate resulting in an increase of 0.001 kg/m^3 salt per second.

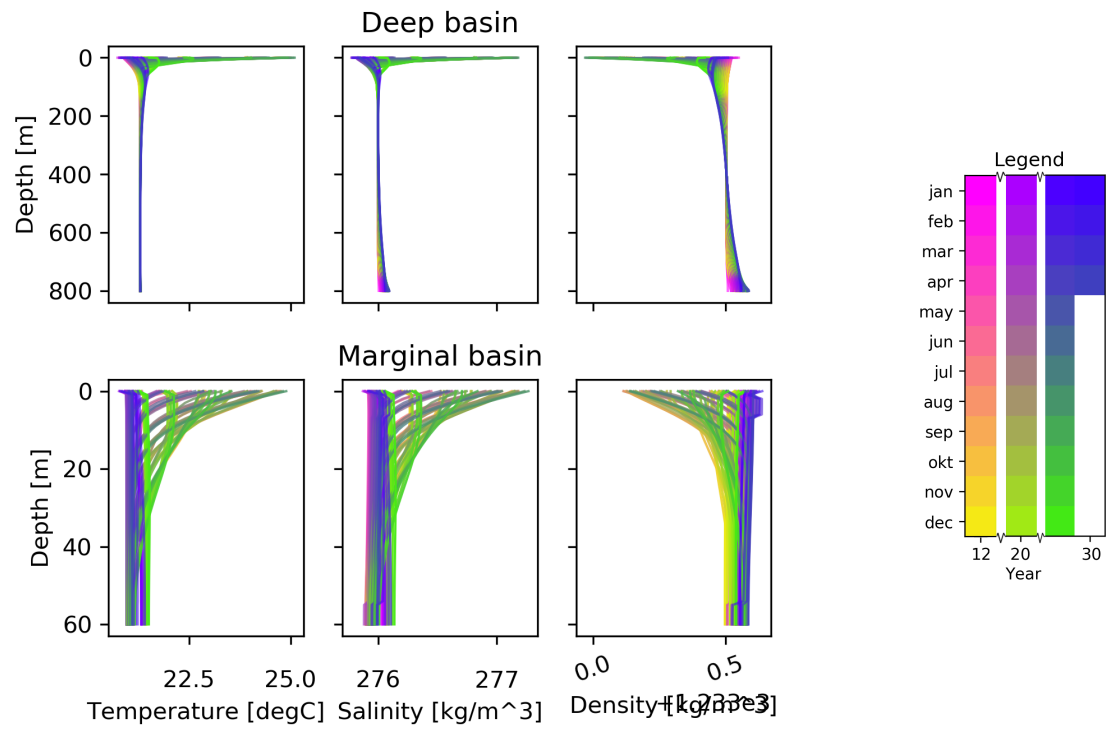


Figure 9: Model run with a very shallow marginal basin. Three periods, those at 11, 25 and 39 years, are shown.

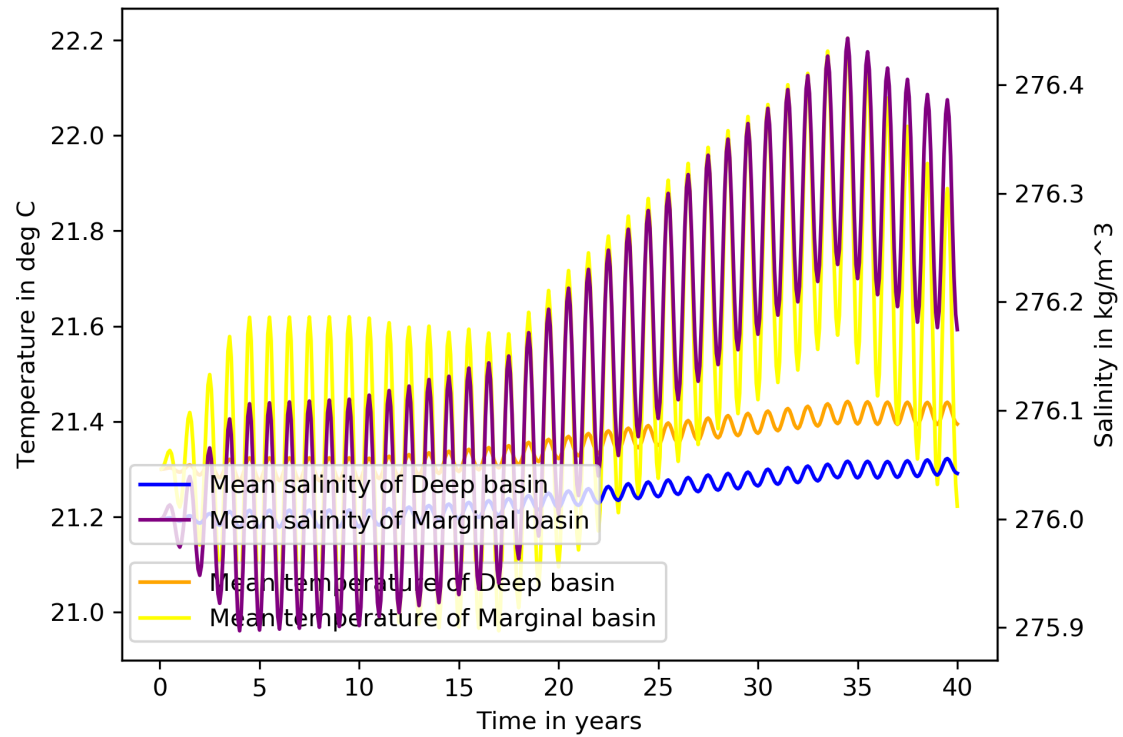


Figure 10: Mean values over time corresponding to the scenario shown in figure 9

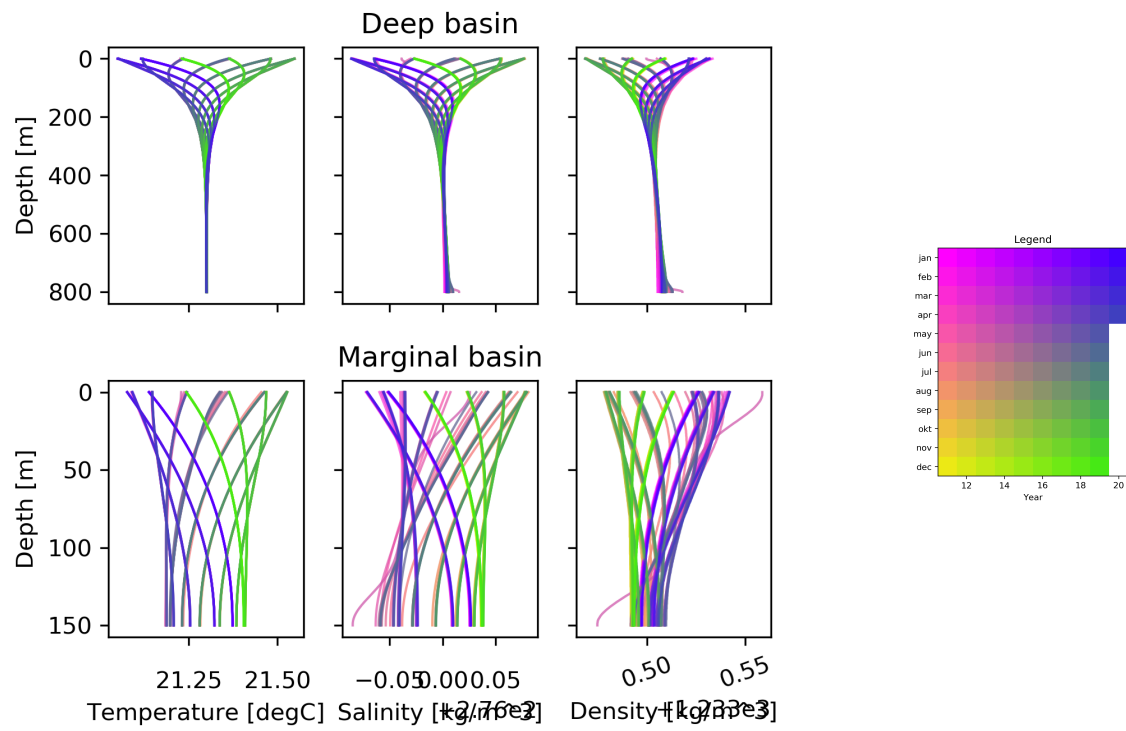


Figure 11: Model run with advection, a flow of $1e-5$ Sv and a marginal basin depth of 150 metres. The background diffusivity was increased, allowing surface processes to penetrate to greater depth

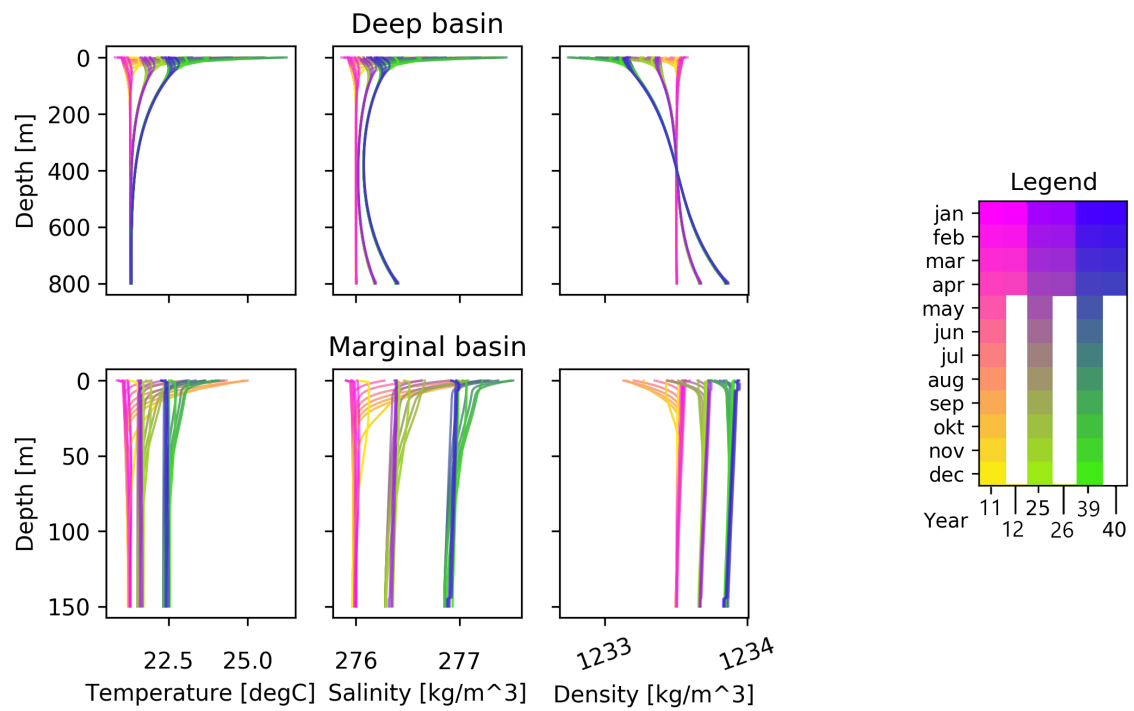


Figure 12: Model run with a net evaporation rate of 0.001 kg/m^3 salt per second.

4.3 Discussion

The effects of the surface heat flux and evaporation decrease with depth and can only penetrate down to a maximum depth, the extend of which is determined by the speed of the diffusivity. Shallower basins will therefore experience changes as a result of forcing more quickly and more severely then deeper basins. Increased depth of penetration of surface forcing due to increased background diffusivity goes hand in hand with a reduction in maximum deviation from the mean surface value. When more salt or heat is transported down, there will subsequently be less of it left at the surface. When the evaporation is significantly higher then the influx of fresh water, the whole column will become more saline. This has a larger effect on the marginal basin, for it contains a smaller volume of water.

The marginal basin in figure 12 is almost permanently mixing. During winter the columns are unstable due to the cooling, as is the case with most other model output in this report. During summer, the increased overall evaporation has resulted in extra saline surface waters, creating instability during summer not unlike that described in section 3.4.2. In figures 9 and 11, the marginal basin is also permanently mixing due to the net increase in salinity. Similar reasoning can be applied there, but the cause of the increased salinity in these cases partially resulted from unintended causes, as discussed in section 3.6.

Nevertheless, all three methods showed to be successful in creating a dense brine in bottom of the deep basin.

4.3.1 Equilibrium

Every year depicted in figures 9, 10 and 12, flow takes place that increases salinity in the bottom of the deep basin by a slight amount. While this excess salinity does decrease in part of the year due to the background diffusion, enough salt is supplied during the year that the deep bottom water experiences a net increase in salinity every year. The upwards diffusion of this salt creates a linear profile over time, the slope of which is mainly dependent on the background diffusivity. The linear dependence of the slope $\frac{\delta Q}{\delta t}$ on the diffusivity is shown in equation 1. This equation shows a similar relation between the slope and the advection, but the advection is generally about two orders of magnitude smaller then the background diffusion. Also, advection only occurs in periods of exchanging flow.

4.3.2 The cycle of exchanging flow

The large pulse of flow marked as (1) in figure 8 occurred in January after which no more flow was exchanged the following years. This is rather different in a scenario where the flow occurs gradually rather then pulse-wise. In all three scenarios (figures 9, 11 and 12) the exchanged volume is rather minor. In these scenarios the flow between the basins also starts in winter. The months differ slightly per scenario but the process is the same. As the winter is progressing, the marginal basin is continuously cooling and thus increasing in density. From April onward, the column is also increasing in salinity due to the increasing evaporation in spring. In the marginal basin the effect is increased since the exchanging flow still continues and the surface water that is entering the marginal basin now is warmer and more saline then the deeper water leaving it. The increase in temperature and by extension the effect it has on the density is however rather small when compared to the increase in salinity. The marginal basin is still in an unbalanced state due to the saline surface water so mixing in the marginal basin is quick in resupplying salt to the bottom. Since the exchanging flow from the marginal basin to the deep basin is now more saline then in winter, the salinity of the bottom waters of the deep basin increases rapidly during spring. In the end of spring/beginning of summer, the turning point is reached, where the marginal basin is now too warm and the bottom water of the deep basin too dense to allow further exchanging flow. During the summer and fall, salt diffuses up from the saline bottom waters in the deep basin, thereby slowly reducing its salinity. During the fall, the marginal basin is cooling again, increasing its density. At the start of winter, the density of the marginal basin has increased enough and the density of the bottom water of the deep basin has decreased enough to again allow for exchanging flow. However, the bottom waters of the deep basin have not yet lost all of their salt since the previous exchange, for the background diffusivity is slow. Also, since significantly more water enters the marginal basin then leaves it during every spring, the deep bottom water increases in salinity every year and the surface water of the deep basin decreases in salinity every year. As made evident by the decreasing yearly increase in the salinity of the marginal basin, when time approaches infinity, it seems a near equilibrium will be established, save the cyclic yearly fluctuations. This equilibrium seems to be determined by the difference, or eventual lack thereof, in salinity between the top and the bottom of the deep basin.

The equilibrium is supposedly reached when the yearly maximum salinity in the bottom of the deep basin is roughly similar to that in the bottom. The marginal basin will then reach a stable yearly average density, capping any further growth of the deep basin brine. No solid conclusions can be made on the long term outcome of the model, for as of yet unidentified imperfections in the code have undesired effects on the total salt and heat contents in the basins, as discussed in section 3.6

4.3.3 Extend of the salt brine

It was suggested by [García-Veigas et al. \(2018\)](#) that the brine would extend roughly from the base of the marginal basin to the bottom of the deep basin. From the viewpoint of the box model this would also make sense, for if flow already occurred and the bottom water of the basin is too dense to be replaced with the dense water from the marginal basin, the marginal basin water would replace water slightly higher in the deep basin instead. This could happen as long as that water is below the base of the marginal basin. This has not been implemented in the model, but may be worth investigating in further research.

4.3.4 Buoyancy

This model deals with instability by mixing the unstable columns, for this is the observed process in the Dead Sea. This model does not take into account any type of gravity interaction with the water, except for the assumption that dense water from the bottom of the marginal basin flows to the bottom of the deep basin if the density allows for it, which is hard coded into the model. When a dense water mass resides below the surface and is considered unstable, the diffusivity is increased drastically to imitate mixing, meaning that the dense mass is moved both up and down, where the water is less dense. The possibility of this dense water body to sink due to buoyancy forces rather than be diffused up as well, is not considered in this model. Future research could investigate the importance of the degree of mixing between a dense water body and the surrounding waters, combined with gravity calculations.

5 Conclusion

If the MSC evaporite deposits originated from a deep sea, a dense brine would have existed at the bottom of the basin. This deep brine may have formed as a result of the shallow marginal basins

that acted as large evaporative pools where water may have flowed in at the top and out at the bottom. This flow could have been driven by density differences between the bottom waters of the marginal basin and the water in the deep basin that resides below the depth of the marginal basin. Since small differences in salinity are quickly re-distributed due to diffusion of salt, the difference in salinity would either need to be substantial or regularly supplemented. Two methods of achieving very saline bottom waters in the marginal basin were investigated.

The first is net evaporation averaged out over the year. If the net evaporation per unit of surface area is similar in the deep basin and the marginal basins, the effect is larger in the marginal basins due to their smaller volume. Then, in periods where the water column is mixing, salt is brought to the bottom of the marginal basin from where it may flow into the deep basin.

The second method to achieve a similar effect, where no net evaporation is required, is when the effects of the increase in salinity of the basins during summer are able to reach the bottom of the marginal basin. One scenario in which this can happen is in a basin that is shallow enough for the effects to reach the bottom before summer is over. **REWRITE COMMENTED PART BASED ON NEW FINDINGS** A final scenario in which the saline summer waters can reach the bottom is through mixing in summer rather than in winter. While this scenario is not the current condition of the Dead Sea, this scenario can occur when the increase in density of the surface water in summer due to high evaporation is larger than the decreasing effect heating has on its density.

This research is a first approximation of the scenario and detailed scenarios of the climate and flow patterns during the MSC are beyond the scope of this research. This research does indicate however, that it may be possible to create deep saline brines through high evaporation in marginal basins.

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