Climate Change - Geological Perspective at Stockholms Universitet

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1 Reading assignment: Earth's Climate Chapter 4

1.1 Key terms

Greenouse era: times when no ice sheets are present

Icehouse era: tmes when ice sheets are present

Faint young Sun paradox: the mystery why the Earth's climate has remained relatively stable throughout most of the planet's history, even though the Sun shone 25% to 30% more faintly 4.55 Byr than today.

Thermostat: thermostat's role is to mitigate extreme temperature by reacting to hot temperature with cooling down the system (e.g. house) and to cold by heating up. We don't know what the Earth's thermostat was through the history, recompensating for the faint young Sun. Candidates include chemical weathering and life

Silicate materials: examples include quartz and feldspar. Silicate materials typically are made of positively charged cations (Na⁺¹, K⁺¹, Fe⁺², Mg⁺², Al⁺³ and Ca⁺²) that are chemically bonded to negatively charged SiO₄ (silicate) structures.

Chemical weathering feedback: chemical weathering creates a negative feedback in the climate. Since chemical weathering is strongly correlated to temperature and precipitation, we can distinguish two causal chains:

- initial change \rightarrow warmer climate \rightarrow increased temperature, precipitation, vegetation \rightarrow increased chemical weathering \rightarrow increased CO₂ removal by weathering \rightarrow reduction of initial warming
- initial chage \rightarrow colder climate \rightarrow decreased temperature, precipitation, vegetation \rightarrow decreased chemical weathering \rightarrow decreased CO₂ removal by weathering \rightarrow reduction of initial cooling

Gaia hypothesis: in its weakest and commonly accepted form, it states that as life-forms gradually developed in complexity, they played a progressively greater role in chemical weathering and its control of Earth's climate. In its most extreme version, it states that life evolved for the purpose of regulating Earth's climate.

Snowball Earth hypothesis: the hypothesis that Earth was once nearly frozen, around 715 to 640 million years ago. Climate scientists have found evidence that glaciers existed on several continents during that time. Some believe these continents were located in the tropics then, but its hard to locate them back in time.

Earth's Cli- 1.2 Review questions

1.2.1 Why is Venus so much warmer than Earth today?

Its atmosphere has 96% CO₂ (compared to Earth's 0.2%), creating a much stronger greenhouse effect, trapping much more heat.

1.2.2 What factors explain why Earth is habitable to-day?

Small greenhouse effect adding only $32^{\circ}\mathrm{C}$ to average temperature in Earth's atmosphere.

1.2.3 Why does the faint young Sun pose a paradox?

Astrophysical models of the Sun's evolution indicate it was 25% to 30% weaker early in Earth's history. Climate model simulations show that the weaker sun would have resulted in a completely frozen Earth for more than half of its early history if the atmosphere had the same composition as it does today.

Primitive life forms date back to at least 3.5 Byr ago, and their presence on Earth is incompatible wit a completely frozen planet at that time.

1.2.4 What evidence suggests that Earth has always had a long term thermostat regulating its climate?

The faint young Sun paradox, the specific evidence being prevalence of water-deposited sedimentary rocks throughout Earth's early history.

1.2.5 Why is volcanic input of CO₂ to Earth's atmosphere not a candidate for its thermostat?

Volcanic processes are diven by the heat sources located deep in the Earth's interior and are well removed from contact with (and reactions to) climate system.

1.2.6 What climate factors affect the removal of CO₂ from the atmosphere by chemical weathering?

Temperature: weathering rates roughly doubl for each 10°C increase in temperature.

Precipitation: increased rainfall boosts the level of groundwater held in soils, and the water combines with CO₂ to form carbonic acid and enhance the weathering process.

Vegetation: plants extract CO₂ from the atmosphere through photosynthesis, and deliver it to soils, where it combines with groundwater to form carbonic acid. It enhances the rate of chemical breakdown of minerals. Presence of vegetation is estimated to increase the rate of chemical weathering by a factor of 2 to 10.

1.2.7 Where did the extra CO_2 from Earth's early atmosphere go?

Sediments and rocks.

1.2.8 What arguments support and oppose the Gaia hypothesis that life is Earth's true thermostat?

Critics say that too many of the active roles played by organisms in the biosphere today are relatively recent developments in Earth's history. The also point out that the very late appearance of shell-bearing oceanic organisms near 540 million years ago means that life had played no obious role in transferring the products of chemical weathering on land to the seafloor for the preceeding 4 Byr.

Supporters claim that critics underestimate the role of primitive life-forms such as algae in the ocan and microbes on land in Earth's earlier history.

Marine organisms that created oxygen through photosynthesis long ago are believed to have enabled the development of oxygen-rich atmosphere 2.4 Byr.

2 Lecture 1: The controls of climate on geological timescales

Time imbalance: Coal takes hundreds of millions of years to accumulate from fossils, but takes decades of burning to release. Accumulation happens on **geological** timescale and release at **antropogenic** timescale.

Average Earth surface temperature is around 15°C.

2.1 Climate factors

Earth absorbs sunlight and radiates heat energy back into space. These 3 factors control the process:

- solar radiation
- albedo effect
- greenhouse effect

Solar radiation

Some prerequisites for calculations:

Stefan-Boltzmann law describes the intensity of the thermal radiation emitted by atter in terms of tat matter's temperature. Formula is $E = \sigma T^4$, where $\sigma = 5.670367 \times 10^{-8} W.m^{-2}.K^{-4}$

Solar radiation constant, in other words, the amount of energy emitted by the Sun is $3.87 \times 10^{26} W^1$.

Solar constant S_0 describes the amount of energy received by a given area one astronomical unit² away from the Sun. Let's calculate it:

 $d_{Earth} = 149, 597, 870, 700m$

Solar constant $S_0 = \frac{Q}{4\pi d^2} = 1362W.m^{-2}$. Since Earth is not flat, but is a rotating sphere, this number is divided by 4, so the effective energy received from Solar radiation is $342W.m^{-2}$.

Now from Stefan-Boltzmann's law, we can calculate the temperature:

$$\begin{split} E &= \sigma T^4 \\ E &= 342 W.m^{-2} \\ T &= (E.\sigma^{-1})^{1/4} = 6^\circ \end{split}$$

Now let's compare with values for Venus:

 $\begin{aligned} d_{Venus} &= 108 \times 10^9 m \\ E_{Venus} &= 658 W.m^{-2} \\ T_{Venus} &= -55^{\circ} \end{aligned}$

Albedo

Black seat: low albedo, white cat: high albedo

Venus has albedo effecto of $\alpha = 77\%$ Earth has albedo effecto of $\alpha = 30\%$

Of course, Earth's albedo is much harder to calculate because the terrain varies a lot, compared to Venus which has a relatively uniform surface.

Venus radiates back to space $658W.m^{-2}\cdot77\%=504W.m^{-2}$. Earth radiates back to space $342.m^{-2}\cdot30\%=103.m^{-2}$.

Taking into account albedo effect, Venus' surface temperature should be -46° and Earth's -18° .

Greenhouse effect

Earth: greenhouse effect increases temperature by 32°.

Let's calculate how much the temperature increased due to greenhouse effect since the preindustrial era, knowing that CO_2 's content in atmosphere increased from 285ppm to 425ppm.

$$\Delta T = 4.38 \ln \frac{CO_{\rm 2present\ day}}{CO_{\rm 2preindustrial}} = 4.38 \ln \frac{425 \rm ppm}{285 \rm ppm} = 1.75^{\circ}$$

2.2 Earth's temperature summary

$$6^{\circ}$$
 + -24° + 32° Solar radiation + Albedo Greenhouse cases

2.3 Faint Young Sun paradox

We have fossils from 3.5 Byr ago. Earliest fossils are stromatolites 3 .

Assuming the same percentage of CO_2 in the atmosphere, the average temperature on Earth at that time (3.5 Byr ago) should have been around 0° (due to lower solar radiation), meaning no running water, which precludes the possibility of life.

2.4 Source of CO₂ on geological timescales

Volcanoes

 $^{^1}$ When an object's velocity is held constant at one meter per second against a constant opposing force of one newton, the rate at which work is done is one watt: $1W=qkg\cdot m^2\cdot s^{-3}$

²roughly equal to average distance Sun-Earth

 $^{^3{\}rm Stromatolites}$ are layered sedimentary formations created mainly by photosynthetic microorganisms such as cyanobacteria, sulfate-reducing bacteria and Pseudomonadota (formerly proteobacteria).

2.5 Earth's thermostate – chemical weathering

Hydrolysis is the main mehcanism for removing CO_2 from the atmosphere. Three key ingredients are minerals that make typical continental rocks, water derived from rain, and CO_2 derived from the atmosphere.

The central equation for chemical weathering is:

$$\begin{array}{c} CaSiO_3 \ + \ H_2CO_3 \\ \text{Silicate rock} \end{array} \\ \begin{array}{c} + \ H_2CO_3 \\ \text{Carbonic acid (soil)} \end{array} \\ \rightarrow CaCO_3 + SiO_2 + H_2O \\ \text{Shells of organisms} \end{array}$$

2.6 Chemical weathering of silicate rocks

First stage of chemical weathering happens under the influence of rain:

Granite +
$$H_2O + 2CO_2 = \text{Clay} + 2K^+ + 2HCO_3^-$$

During the weathering, carbon dioxide switches from being a greenhouse gas to being a sollute.

The bicarbonate ions are then carried by rivers and eventually end up in seas and oceans. In the ocean, bicarbonates find calcium which they react with, and make limestone, which is calcium carbonate.

$$2HCO_3^- + Ca^{2+} = CaCO_3 + H_2O + CO_2$$
 limestone

We took 2 molecules of carbon from the atmosphere, and return only one, the other one is deposited in limestone. Thus the precipitation limestone is a sink.

Q: Can chemical weathering of silicate rocks compensate for anthropogenic CO₂ emissions?

A: No, it is way too slow.