

**1 Remote Sensing Methods for Identifying Ephemeral
2 Water Bodies Across Arizona**

3 Travis Zalesky¹

4 ¹University of Arizona,

5 **Abstract**

6 A variety of remote sensing methods are investigated for their ability to detect
 7 ephemeral water bodies across the state of Arizona. The strengths and weaknesses of
 8 each are discussed.

9 **Plain Language Summary**

10 Results and methods for detecting transient water bodies using satellite data.

11 **1 Introduction**

12 As part of the Arizona Tri-University Recharge (ATUR) project, multiple methods
 13 for identifying and quantifying ephemeral water bodies across the state of Arizona
 14 (AZ) have been investigated. For reasons that will be discussed, accurate identifi-
 15 cation of ephemeral, or temporary, water bodies is a challenge, and efforts to-date
 16 have shown limited success. Methods, results, and challenges are discussed, and
 17 potential paths forward shall be explored.

18 Ephemeral water bodies are common features in the arid American southwest. Typ-
 19 ically very shallow, these water bodies range in size from a few cm rock pools to
 20 sheet flow associated with flash floods, and playas up to several km in diameter
 21 (Crawford, 1981; Whitford & Duval, 2020). While playas and intermittent streams
 22 are well mapped across AZ, there is no consensus on the volume of water which
 23 flows through these systems. More to the point, there is no widely accepted estimate
 24 of the volume of water which either evaporates or recharges from these internally
 25 draining basins.

26 It is not clear what fraction of water which flows onto a playa is evaporated vs. recharged.
 27 On the one hand, the nature of playas results in the slow accumulation of surface
 28 salts and fine sediments (clays and silts), which are generally impervious and detri-
 29 mental to infiltration and subsequent recharge (Whitford & Duval, 2020). Addi-
 30 tionally, playas may be an oasis for drought resistant flora, increasing local evapo-
 31 transpiration, even after shallow surface soil layers have dried out (Crawford, 1981).
 32 Alternately, these mineral rich surface crusts may be prone to expansion when wet-
 33 ted, and are frequently characterized by deep cracks which provide avenues for water
 34 infiltration below the impervious layers (Whitford & Duval, 2020). While the ma-
 35 jority of water flowing onto most playas will be lost to evaporation, there may be
 36 specific playas where water loss by infiltration may exceed evaporation (Whitford
 37 & Duval, 2020). At least initially, we will assume that all water flowing onto playas
 38 shall be evaporated, subject to later reevaluation and refinement.

39 The primary target of this study is to estimate and quantify the volume of water
 40 flowing into small to medium sized playas, on the order of ~50 m to ~1 km in dia-
 41 meter, although identification of sheet flow may also be incidental. The general method
 42 outlined is to use remote sensing technology to identify standing water broadly
 43 across the study area, across a time series. By identifying standing water, and as-
 44 suming a constant water depth, water volume can be estimated. Additionally, an
 45 approximation of “continuously wetted” days can provide an initial estimate of evap-
 46 oration rates, and through comparison with alternate evapotranspiration estimates,
 47 may be useful for estimating infiltration rate (e.g. if water loss rate is 2 cm/day and
 48 the evapotranspiration rate is 1.5 cm/day then the implied infiltration rate would be
 49 0.5 cm/day).

50 The study area used for method development is primarily Hualapai Playa (a.k.a.
 51 Red Lake), located in northwest AZ, approx. 60 km south-southwest of Lake Mead
 52 (Figure 1). This ephemeral playa is roughly 8 km in diameter, and is one of the
 53 largest playas in AZ. Furthermore, local landowners surrounding Hualapai Playa
 54 have expressed sincere interest in our project, creating additional public-relations

55 incentives related to this particular playa. Lastly, it's proximity to Lake Mead and
 56 Lake Mohave provide a convenient region for testing these methods against known
 57 areas of permanent surface water. Additional playas are investigated as needed.

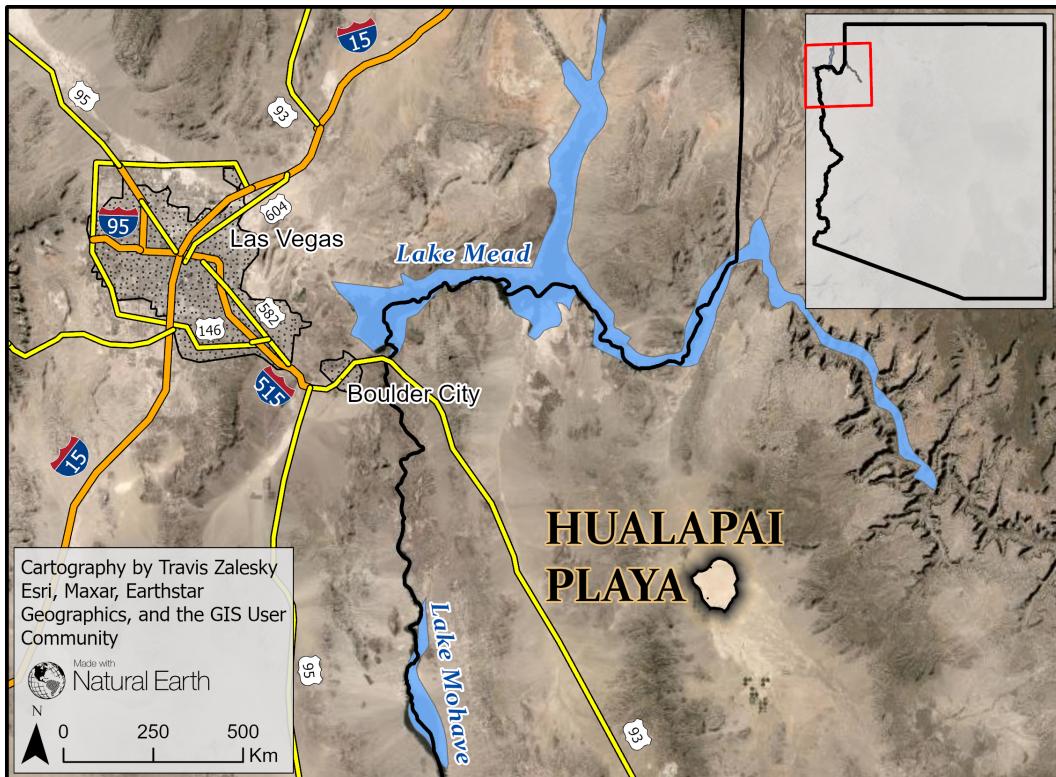


Figure 1: Hualapai Playa locator map.

58 One of the most challenging aspects of developing these methods is the lack of
 59 quality ground-truth data. Without existing data regarding date and duration of
 60 standing water within these ephemeral features, it is extremely difficult to validate
 61 our methods. While reasonable assumptions can be made about method accuracy
 62 (or inaccuracy as the case may be), any method, either those presented herein, or
 63 future methods not yet considered, will have to be validated using a known and well
 64 characterized ephemeral water body, ideally within AZ, or elsewhere in a similar
 65 arid environment.

66 2 Methods & Results

67 2.1 Method 1: Otsu Thresholding with Synthetic Aperture Radar

68 Otsu thresholding is an unsupervised machine vision technique which seeks an optimal
 69 threshold for segmenting an image into a binary classification based on a simple
 70 histogram classification (Otsu et al., 1975). The method was initially developed in
 71 the late '70s, and is widely applicable to a range of computer vision classification
 72 problems. It assumes a bi-modal histogram distribution and determines an optimal
 73 thresholding value through the integration of the histogram, rather than simply re-
 74 lying on local properties "such as valley[s]" (Otsu et al., 1975). In recent years, Otsu
 75 thresholding has been combined with Synthetic Aperture Radar (SAR) satellite im-
 76 agery to map surface water, such as flood monitoring in the Mekong Delta (Tran et
 77 al., 2022). The properties of surface water and SAR make an ideal pairing for use
 78 in Otsu thresholding. The smooth surface characteristics of surface water create low

SAR back-scatter, appearing as very dark pixels (Figures 2-3). Compared to the surrounding land surface, water typically has a very high contrast in SAR imagery, which can be leveraged for Otsu thresholding. Additionally, the cloud penetrating characteristics of SAR offer unparalleled all-weather monitoring capabilities.

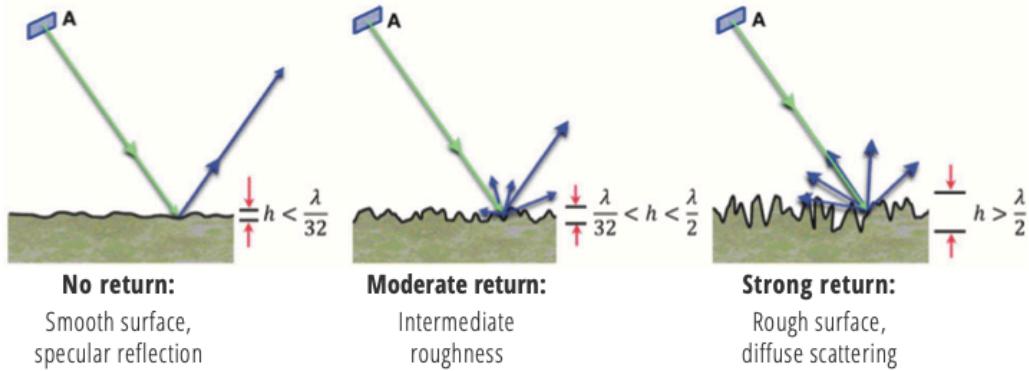


Figure 2: Synthetic Aperture Radar signal return strength (i.e. back-scatter) is a function of the roughness of the ground surface. Credit: Meyer (2019), Figure 2.8. For educational purposes only.

This method initially seemed like an ideal solution. High resolution SAR imagery is freely available from the European Space Agency (ESA) Sentinel-1 satellite, with a global coverage from late 2014 to present, and a 6-day revisit time. Google Earth Engine (GEE) code for Otsu thresholding was easily adapted from Markert et al. (2023), and preliminary results appeared to be promising. Initial analysis appeared to show surface water ponding on Hualapai Playa in mid-Jan 2020. These preliminary results were circulated internally, and purported to be both evidence of surface water, and proof-of-concept for method validation. However, further research and wider application of this method revealed two critical flaws.

Firstly, the method, as applied in GEE, was highly sensitive to the size and characteristics of any given playa. For larger playas, the method was generally suitable, but required a substantial amount of manual parameter optimization to function properly. For small to medium sized playas, the method was largely unstable, resulting in persistent errors. As a result this method is not particularly scalable, and at best could be used for a selection of the states largest playas, such as Hualapai and Willcox Dry Lake (southeast AZ).

More critically, this method is sub-optimal for arid regions, as flat sandy surfaces also result in low SAR back-scatter, appearing similar or indistinguishable from surface water. The hard-pan surface of a dry playa is particularly apt to be misidentified by this method, resulting in a high likelihood of false positives. It is currently unclear if the preliminary results purporting to have identified surface water on Hualapai Playa in Jan. 2020 are true or false positives, as there is no associated ground-truthing data against which to compare. Regardless of the truthiness of these results it is evident that this method can not be used reliably within the study area without widespread and comprehensive testing and validation of results. Given the scheduling and budgetary constraints of this project widespread survey data collection is not feasible, and while further development of this method may be

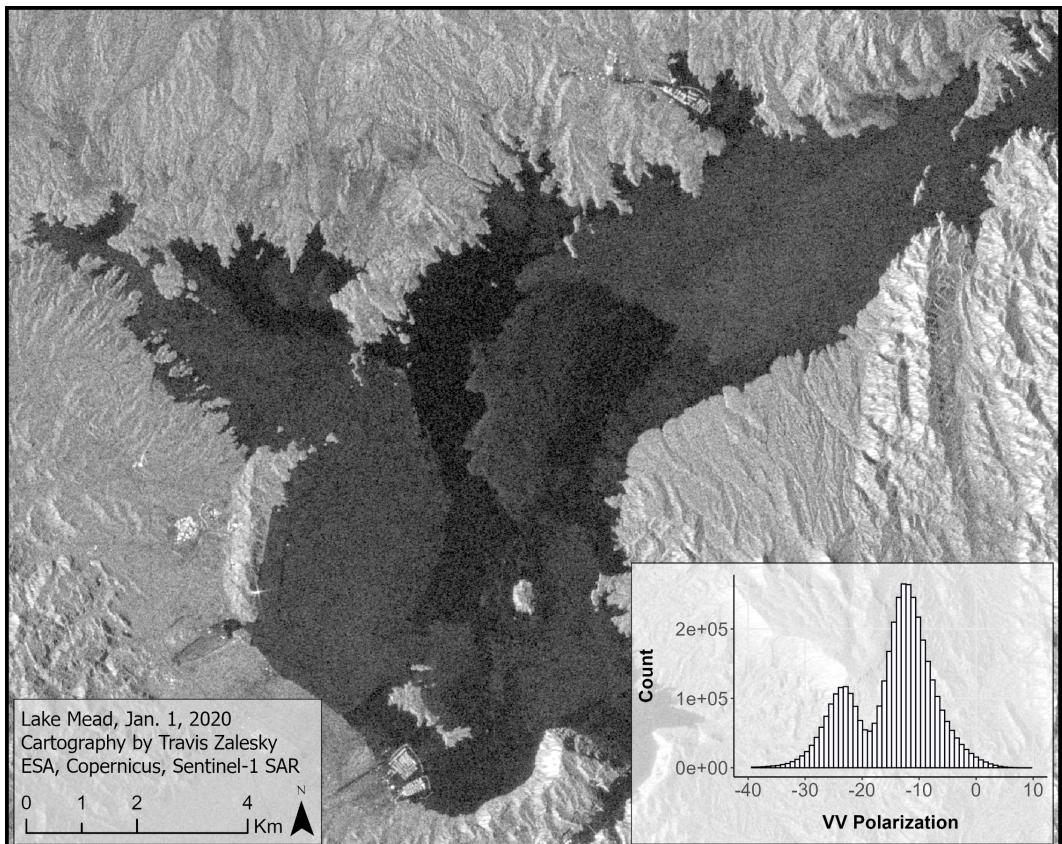


Figure 3: High resolution Synthetic Aperture Radar (SAR) imagery of Lake Mead, obtained from the European Space Agency (ESA) Sentinel-1 satellite (1-1-2020). Surface water appears dark, and contrasts strongly with the surrounding land surface, resulting in a bimodal histogram distribution, ideal for use in Otsu thresholding.

warranted (generally), it is not appropriate for our purposes. Therefore alternate methods must be explored.

112 **2.2 Method 2: Normalized Difference Water Index with Landsat Multi-** 113 **spectral Imagery**

114 Collectively, the 9 Landsat satellites provide the longest continuous record of space-
115 based earth observations, from 1973 to present. Therefore, it may seem like an
116 obvious candidate for monitoring infrequent phenomena, such as surface water pond-
117 ing on playas, which may only happen once every several years (depending on the
118 playa). However, Landsat imagery has several disadvantages compared to SAR.
119 Firstly, the 30 m resolution of Landsat imagery is nine times coarser than the 10
120 m resolution of Sentinel-1 SAR. Second, the revisit time for Landsat satellites is
121 16-days, more than double that of Sentinel-1. By combining data from multiple
122 Landsat satellites this effective revisit time may be cut to 8-days, but this may not
123 be possible throughout the full time series of interest. Additionally, the visible and
124 infrared spectral bands used by Landsat sensors is not capable of penetrating clouds,
125 and while cloud-masking can be achieved through post-processing, the presence of
126 clouds has the potential to obfuscate the phenomena of interest. Combined, these
127 three disadvantages compound, resulting in much sparser data availability for Land-
128 sat imagery in comparison to Sentinel-1.

129 Furthermore, while indices for surface water monitoring utilizing Landsat spectral
130 bands, such as the Normalized Difference Water Index (NDWI), are widely used,
131 they are imprecise classifiers on their own. NDWI in particular, is suspectable to
132 misclassifying shadows as water. NDWI is defined as the difference between the
133 green and near-infrared (NIR) bands divided by the sum of the green and NIR
134 bands (for Landsat 8/9, bands 3 and 5 respectively; Equation 1).

$$135 \quad NDWI = \frac{Green - NIR}{Green + NIR} \quad (1)$$

136 Despite these disadvantages, it may still be possible to characterize and quantify
137 ephemeral water bodies using this dataset.

138 Fortunately, Landsat data is freely available, easily accessible through GEE, and
139 Landsat Collection 2 datasets have added post-processing data which is directly
140 applicable to this method, including cloud and cloud shadow masks, as well as a
141 pre-calculated water mask (which appears to be largely coincident with $NDWI > 0$).
However, this water mask is over inclusive, and not suitable for use as is.

142 A large number of features identified as water in the provided Landsat layer ap-
143 peared to correspond to the northward aspect of terrain features. Therefore, the first
144 and most important step in refining the provided water mask layer was to reduce the
145 false positives as a result of natural terrain shadows. This was accomplished by cre-
146 ating a hillshade (or hill shadow) layer in GEE from a high resolution (10 m) USGS
147 digital elevation model (DEM).

148 **3 Further Recommendations**

149 **4 Conclusion**

150 **References**

- 151 Crawford, C. S. (1981). Chapter 17 - the invertebrate community of ephemeral
152 waters. In *Biology of desert invertebrates* (pp. 234–247). Springer-Verlag.
153 Markert, K., Donchyts, G., & Haag, A. (2023). Chapter A2.3 surface water map-
154 ping. In J. A. Cardille, N. Clinton, M. A. Crowley, & D. Saah (Eds.), *Cloud-
155 based remote sensing with google earth engine: Fundamentals and applica-
156 tions* (1st ed.). Springer Cham. <https://docs.google.com/document/d/>

- 157 [12cwzbNXtBQnm5switfL1v1stnp5D5gXcQeZMP_iJI/edit?tab=t.0#heading=h2gd4k88k0hkr](https://doi.org/10.25966/nr2c-s697)
- 158
159 Meyer, F. (2019). Chapter 2 - spaceborne synthetic aperture radar: Principles, data
160 access, and basic processing techniques. In A. Flores-Anderson, K. Herndon, R.
161 Thapa, & E. Cherrington (Eds.), *The SAR handbook: Comprehensive methodologies for forest monitoring and biomass estimation*. <https://doi.org/10.25966/nr2c-s697>
- 162
163 Otsu, N. et al. (1975). A threshold selection method from gray-level histograms.
164 *Automatica*, 11(285-296), 23–27.
- 165
166 Tran, K. H., Menenti, M., & Jia, L. (2022). Surface water mapping and flood mon-
167 itoring in the mekong delta using sentinel-1 SAR time series and otsu threshold.
168 *Remote Sensing*, 14(22). <https://doi.org/10.3390/rs14225721>
- 169
170 Whitford, W. G., & Duval, B. D. (2020). *Chapter 8 - consumers and their ef-*
171 *fects* (W. G. Whitford & B. D. Duval, Eds.; pp. 203–263). Academic Press.
172 <https://doi.org/10.1016/B978-0-12-815055-9.00008-4>
- 173