

Supporting Information for “Rheologic constraints on the upper mantle from five years of postseismic deformation following the El Mayor-Cucapah earthquake”

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Introduction

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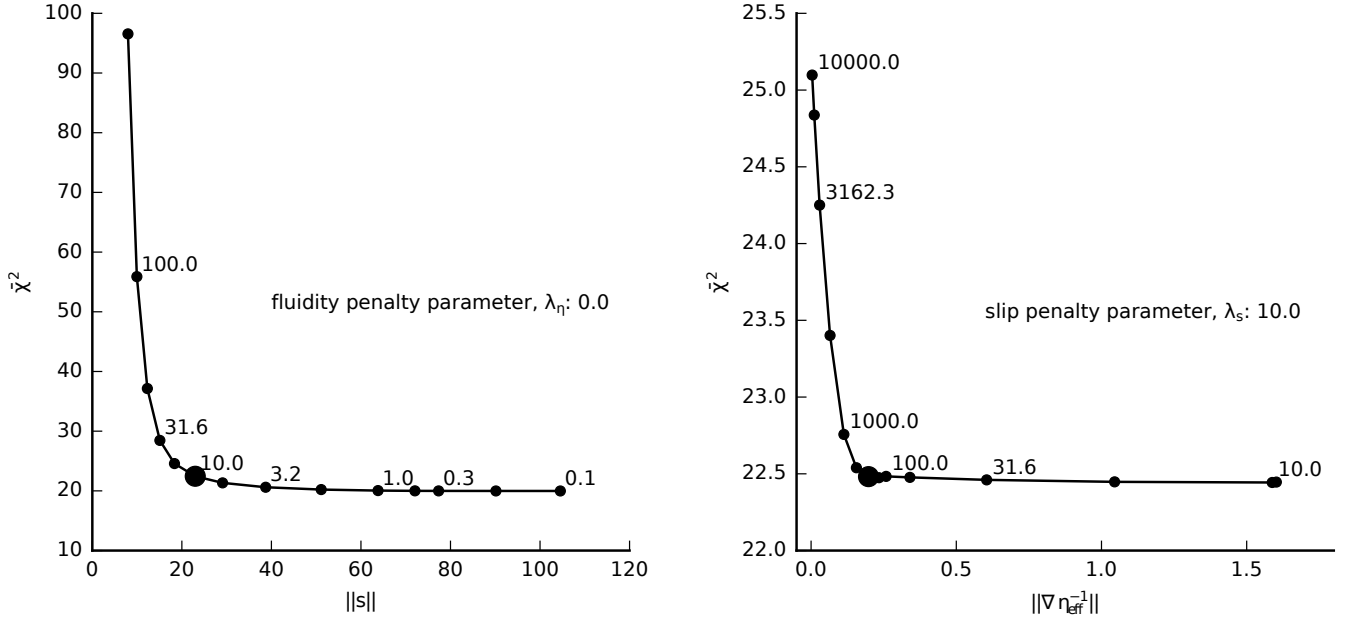


Figure S1. Trade-off curves used to determine the damping parameters λ_s and λ_η in eq. (15) of the main text. The left panel shows the trade-off curve for the fault slip penalty parameter, λ_s . We pick λ_s while keeping the penalty parameter for fluidity, λ_η , fixed at zero. The right panel shows the trade-off curve for selecting λ_η , where we fix λ_s at the chosen value from the left panel. Chosen values are indicated with the larger marker. When picking λ_s , we try to find a good balance between the mean chi-squared value, $\bar{\chi}^2$, and the size of the slip parameters, $\|s\|$. Our choice of λ_η is a balance between $\bar{\chi}^2$ and the size of the Laplacian of fluidity, $\|\nabla \eta_{\text{eff}}^{-1}\|$.

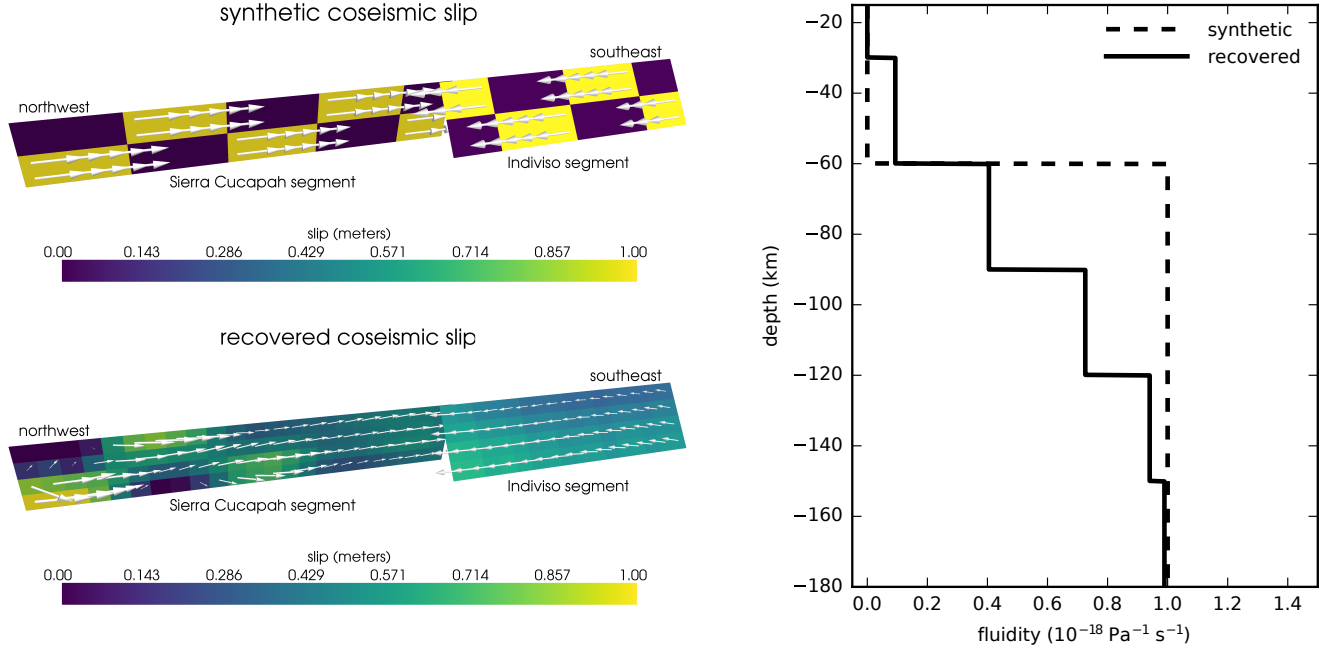


Figure S2. Checkerboard test used to assess the resolving power of the inversion in Section 3.2 of the main text. We create synthetic data at all of the GPS stations considered in this study by evaluating eq. (14) with the synthetic coseismic slip distribution and fluidity distributions. Our synthetic fluidity model has a jump from 0.0 to $10^{-18} \text{ Pa}^{-1} \text{ s}^{-1}$ at 60 km depth. Our synthetic slip model does not include afterslip, although we estimate afterslip along with coseismic slip and fluidity in this test. We estimate these values in the same way as described in the main text and we also use the same penalty parameters. We do not add any noise to our synthetic data so that the recovered model just indicates how much the regularization influences the solution. Note that our ability to recover slip decreases towards the southern end of the fault, farthest from the available data. Also note that the smoothing constraint on fluidity largely obscures the jump in the synthetic model.