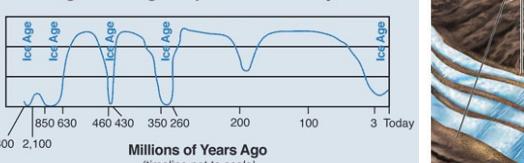


Glacial Ratio **Aerosol**: a colloidal suspension of particles dispersed in air or gas. **Firn**: the intermediate state between snow and glacial ice. **Accumulation**: Accumulation is when glaciers gain more mass through snowfall, windblown snow, avalanches, rime ice (freezing water vapor), refreezing meltwater. **Ablation**: Ablation is when glaciers loose mass through surface melt, surface meltwater runoff, sublimation, avalanching and windblown snow. **Global Connections** **Atmosphere** Past 200 years: Past 200 yr: CO₂ went up by 40% and Methane by 200% - 300%. Glaciers Reflect heat from the sun; increased dust and soot from grazing, farming, and burning of fossil fuels and forests, are also causing glacier retreat by. • Past 200 yr: CO₂ went up by 40% and Methane by 200% - 300%, which glaciers have the ability to combat. • Reflect heat from the sun, increased dust and soot from grazing, farming, and burning of fossil fuels and forests, are also causing glacier retreat (albedo) • layers of dust and soot are darkening the color of glaciers and snowpacks, causing them to absorb more solar heat and melt more quickly, and earlier in spring. • **Albedo**, or "whiteness", is a scientific term meaning reflectivity. • Cooking stoves (biomass stoves) darken snow and ice in mountainous regions. In the himalayas this is bad because the Yangtze, Yellow, Mekong, and Ganges rivers all feed from glaciers (~90% of Himalayan Glacier Melting Caused by Aerosols and Black Carbon). • **Aerosol**: a colloidal suspension of particles dispersed in air or gas, reducing albedo. **Ocean**: If glacier melted sea level would rise by ~ All of Greenland (7.2m); West Antarctic Ice Sheet (3.2m). All of Antarctica (3.7m). • sea level has risen by 4 to 8 inches over the past century. • Rise of temperature over the past 20 years has been 0.13 inches (3.2 millimeters) a year. **Lithosphere**: When glaciers erode the rock underneath them, they release carbon gases trapped in the lithosphere. Also, when ice sheets weigh down on the sea floor, the cause depression in the earth's lithosphere, and the edges are called fore bulges, which are massive hills that areas like America's east coast lie upon. When these sink, the depressions left rise, causing a reshuffling of the earth's lithosphere. This is called glacial isostatic adjustment, **basal sliding**, when the ice slides over the land with a layer of water acting as a lubricant and reducing the friction between land and ice. pressure from the weight of the ice reduces the melting point at the base of the glacier which allows the ice to melt, allowing water to be present. glaciers can move in even the coldest of climates. **Ice Core**: a core sample drilled from the accumulation of snow and ice over many years that have recrystallized and have trapped air bubbles from previous time periods, the composition of which can be used to reconstruct past climates and climate change; typically removed from an ice sheet. **Oxygen Isotopes**: common isotopes O₁₆ and O₁₈. Water w/ O₁₆ is lighter, water with O₁₈ is heavier; it tends to evaporate easier, causing it to accumulate in oceans and it to end up in water and ice. • During constant climate conditions the O₁₈ lost to evaporation returns to the oceans by rain and streams, so that the ratio of O₁₈ to O₁₆ (O₁₈ / O₁₆) is constant. • But, during a glaciation, some of the O₁₈ gets tied up in glacial ice and does not return to the oceans. Thus during glaciations the O₁₈ / O₁₆ ratio of sea water increases. • During an interglaciation, on the other hand, the O₁₈ that was tied up in glacial ice returns to the oceans causing a decrease in the O₁₈ / O₁₆ ratio of seawater. • Thus, we expect that during glaciations the O₁₈ / O₁₆ ratio in seawater will be high, and during interglaciations the O₁₈ / O₁₆ ratio in seawater will be low. **Info from Ice Cores**: • Accumulation rate - The thickness of the annual layers in ice cores used to derive a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • Melt Layers - Ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core. • Past air temperatures - It is possible to measure past air temperatures from ice cores. This can be related directly to concentrations of carbon dioxide, methane and other greenhouse gasses preserved in the ice. **Rate of flow**: Controlled by: 1. The severity of the slope. 2. Basal water= wet bottom= faster flow. 3. location within glacier= greater velocity in ice center. **Sedimentary Sequences**: • Sedimentary environments are areas where sediments are deposited; glaciers are an example of this. **Supraglacial (ice marginal)**: Readily be observed along glacial margins. • A dark, dirty ice zone is not uncommon at a glacier's leading edge. • The supraglacial environment is a very unstable place because material deposited on top of ice is going to move when the ice melts. • Till-like mixtures of material with a wide range of particle sizes, called "diamicton". • Reflect a complex history of deposition. **Subglacial**: Most difficult to observe. Relies on ice cores and down-hole cameras. • Glaciers grind up and mix rock and soil debris in and beneath their base forming a mixture of material (rocks, sand, silt, and clay) that is called till. • Till is the most common subglacial material. • But river and lake deposits also occur. **Proglacial**: even more dynamic than the subglacial one. • Glacial meltwater and summer rains carry debris away from the glacier or deposit it in lakes that come and go as the force of the water causes natural dams to give way and lakes to drain; sometimes catastrophically sweeping material away in the water. • Include materials sorted by water or wind, river sediment (called outwash), lake sediment, windblown sand, and windblown silt called loess. • **Milankovitch Cycles**: describe the collective effects of changes in the Earth's movements on its climate over thousands of years. variations in eccentricity, axial tilt, and precession of the Earth's orbit resulted in cyclical variation in the solar radiation reaching the Earth, and that this orbital forcing strongly influenced climatic patterns on Earth. **Eccentricity**: refers to the earth's orbit and its shift from being circular to more elliptical over time. **Axial Tilt (Obliquity)**: Tilt of earth's axis of rotation. A greater tilt means more drastic seasons; The angle varies between 22.1° and 24.5°, over a cycle of about 41,000 years. The current tilt is 23.44°. We are currently on a downward trend, meaning warmer climates. **Axial Precession**: is a gravity-induced, slow, and continuous change in the orientation of an astronomical body's rotational axis. The cycle is relative to fixed stars, with a period of 25,771.5 yrs. **Axial Precession** changing of the line between the sun and the earth that changes. Tilt of the orbit itself. **Solar Forcing**: changes in these movements of the Earth, which alter the amount and location of solar radiation reaching the Earth. **Perihelion**: closest to the sun; **Aphelion**: farthest from the sun; The semi-major axis is a constant, therefore when the earth orbit becomes more eccentric, the semi-minor axis shortens. Increase in solar insolation: at closest approach to the Sun (perihelion) compared to the insolation at the furthest distance (aphelion) is slightly larger than four times the eccentricity. **Milutin Milankovic** Serbian geophysicist and astronomer. • **how long**: 100,000 year long cycle. **History of Ice on Earth** **Recent History**: There have been five or six major ice ages in the past 3 billion years. The Late Cenozoic Ice Age began 34 million years ago, its latest phase being the Quaternary glaciation, in progress since 2.58 million years ago. **Neoproterozoic Snowball on Earth**: Snowball earth around 650 mya - biological activity in the ocean surface collapsed for millions of years. Ended when volcanic outgassing raised CO₂ to 350x modern level; Ocean was virtually covered by thin sea ice + continents were covered in patchy ice due to hydrologic cycle; Sir Douglas Mawson proposed this. **Late Paleozoic Ice Ages**: conventional view: paleozoic ice age was a long ice age for 100 million yrs w/ some internal waning + waning of glaciers. Recent research: series of shorter glacial events separated by periods of warmth. • Expanded from South America to southern Africa to Australia. • The ending constitutes turnover to greenhouse state. • Sea level response (glacio eustatic) to ice age may be less extreme than one thought. **Eocene-Oligocene Transition**: and the impact of opening oceanic seaways. • Marked by large scale extinction. • Most affected organisms were marine or aquatic in nature. • Major cooling on land and in ocean. • Causes include volcanic activity + meteorite impacts + decrease in atmospheric CO₂. • Sea level changes mark transition - in NE Italy, sea level fell 20 m and then 50-60 m in the Oligocene Isotope Event. • Extinctions could have been caused by volcanic explosions or meteorites. • Extinction caused by climate change and major fall in sea levels. **Pleistocene onset of Northern Hemisphere glaciation**: lead to reorganization and relocation of species as associations and may have enhanced species turnover. • Changes in CO₂ could have helped to lead to glaciation.

tion • Began a unique period in Earth's history where both poles have remained ice locked for 10 and 6 Ma but did not gain momentum until 3.5-3 Ma. • Northern Hemisphere glaciation occurred in episodes after Greenland froze. • Tectonic changes might have triggered more extensive NH glaciation

Ice Ages during the past 2.4 billion years



Glacial periods are characterized with large ice sheets and are normally known as ice ages. The periods between these are known as interglacial periods and currently we are in the Holocene interglacial period. The last glacial period was between 120,000 to 11,500 years ago and was during the Pleistocene Epoch.

• Glacial periods are times in the Earth's history where average global temperatures were approximately 6°C lower and glaciers covered much of the planet's surface. The last of these periods ended approximately 10,000 years ago. There are 6 main factors that contribute to global climate and can cause glacial periods: **solar variability**, **insulation**, **dust**, **atmospheric composition**, **ocean current circulation**, **sea ice**, and **atmospheric circulation**. All of these are natural processes and the one that is affected by humans is atmospheric composition. • **Interglacial periods** are caused by shifts in the Earth's orbit and this causes a change in the amount of solar radiation that hits the Earth. When the amount of solar radiation increases, this is when the Earth shifts from a glacial to interglacial period. The Quaternary has had multiple shifts between glacial and interglacial periods and in the middle of this period, the change cycle between glacial and interglacial changed every 100,000 years. • Antarctica is a great indicator if Earth is in a glacial or interglacial period because the amount of ice and snow on it indicates the amount of solar radiation that is hitting the Earth as well as the average temperature of the Earth. • CO₂ is also an indicator of the changing from a glacial to interglacial period or vice versa. As the CO₂ levels increase the Earth's average temperature will increase and it will move into an interglacial period whereas if the CO₂ levels are to fall, the average temperature would fall and the Earth would change to a glacial period.

Glacial Formation **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as **new**; after two winters (one melt season), snow turns into **firm**; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

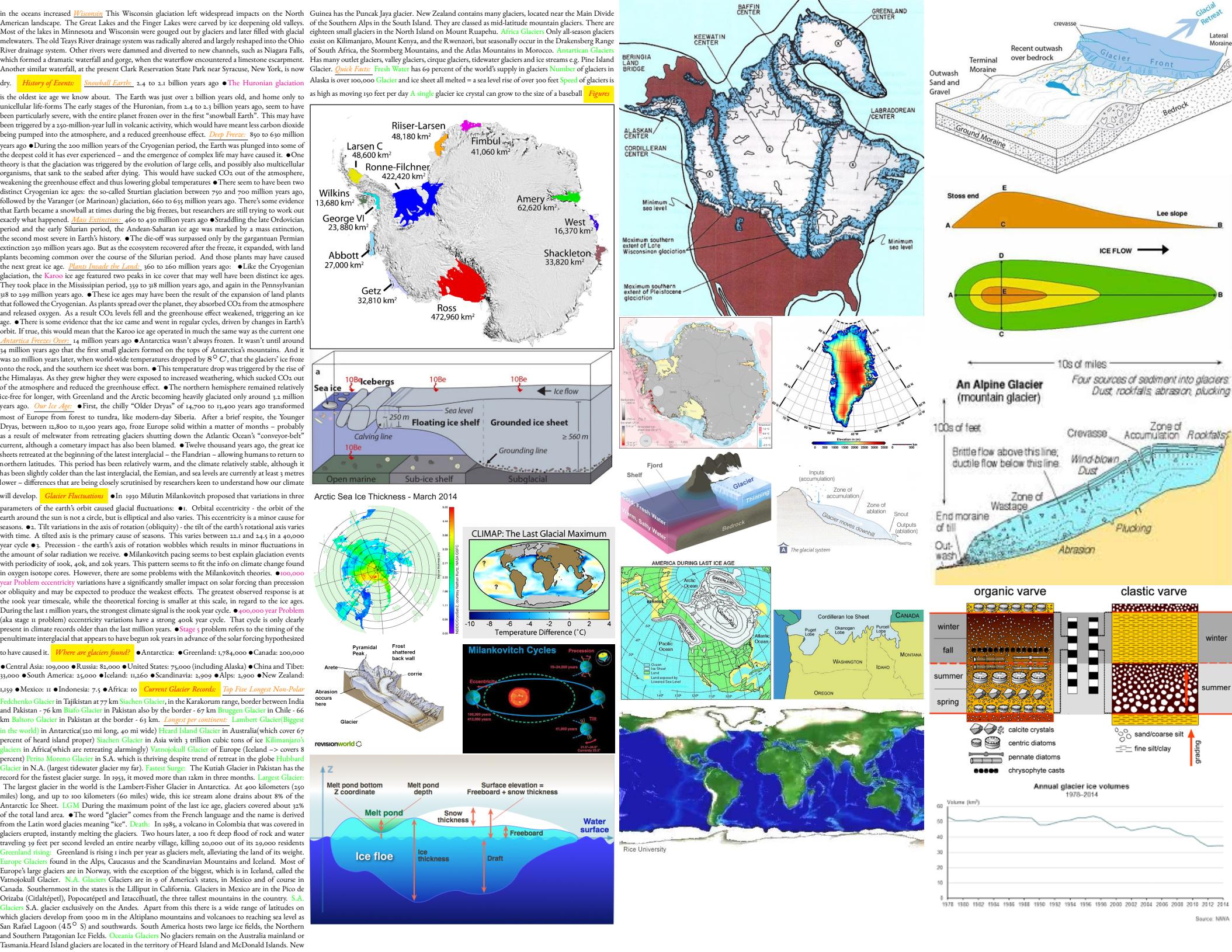
• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new; after two winters (one melt season), snow turns into firm; the ice increases in density over years. **Ice Crystal Structure**: Commonly takes the shape of sheets or planes of oxygen atoms joined in a series of open hexagonal rings; ice can form 18 different crystalline phases; tacked in a lamellar structure that occasionally deform by gliding. When this gliding deformation occurs, the bonds between the layers break, and the hydrogen atoms involved in those bonds must become attached to different oxygen atoms. **Properties of Ice**: The albedo is 0.5 to 0.9 for snow, 0.3 to 0.6 for firm, and 0.15 to 0.35 for glacial ice. **Albedo**: lowers the melting point of the glacier due to hydrostatic pressure, where deeper parts of the glacier are colder. **Moraine**: Material a glacier picks up or pushes as it moves forms moraines along the surface and sides of the glacier. As a glacier retreats, the ice literally melts away from underneath the moraines, so they leave long, narrow ridges that show where the glacier used to be. Glaciers do not always leave moraines behind, however, because sometimes the glacier's own meltwater washes the material away. **Drunlin**: long, tear-shaped depressions formed by a precipitation rate (after correcting for thinning by glacier flow). Past precipitation rates are an important paleoenvironmental indicator, often correlated to climate change, and it's an essential parameter for many past climate studies or numerical glacier simulations. • **Melt Layers**: ice cores provide us with lots of information beyond bubbles of gas in the ice. For example, melt layers are related to summer temperatures. More melt layers indicate warmer summer air temperatures. Melt layers are formed when the surface snow melts, releasing water to percolate down through the snow pack. They form bubble-free ice layers, visible in the ice core.

• **Glacier Formation**: **Glacial Ice**: Glacier formation - snow in same area year long and accumulates into masses of ice; after the first winter this is known as new;



ary • Ablati

Area: The area of a glacier where more glacier mass is lost than gained in the snow surface caused by the sun or warm, gusty wind. **Ablation moraine** in the ablation zone of a glacier; the rock has been plucked from the glacier and is melting out on the ice surface. **Ablation Season:** Period when mass than gain; usually coincides with summer. **Ablation Zone:** snow and ice ablation exceed accumulation. **Abrasions:** rocks within the mouth and polish the surface below; pulverized rock produced is called rock flour. The bottom of a glacier contains large rock fragments, and long scratches of travel. **Accommodation Space Equation:** represents a simple volume of the left controlling the amount of space that can be occupied by sediments in the right describing how much water or sediment fills the accommodation space of a glacier; also known as the glacier balance equation.

glacier or a glacier where more mass is gained than lost. **Accumulation Season:** seasons more mass than it loses usually coincides with winter. **Accumulation:** more mass is gained than lost. **Advance:** When a mountain glacier's terminus advances before; glacial advance occurs when a glacier flows down valley at its terminus. **Alpine Glacier:** A glacier that is confined by surrounding mountain glacier. **Arête:** Sharp, narrow ridge is formed as a result of glacial erosion. **Cirque:** A U-shaped depression in a mountain side, usually formed by glacial erosion.

Dges. Alternate bands of light and dark on a glacier; usually found below the result of different flow and ablation rates between summer sliding of a glacier over bedrock; melting point of ice decreases with pressure; that separates bedrock from stagnant ice at the head of a glacier that has one or more tributary glaciers that flow into it; distinguished by a central, white, meltwater-losing **Head Zone**. The uppermost section of a glacier is called a **subnival zone**.

It has only a single tributary glacier **Brittle**. The upper 50 meters or more budget of Glacier as terminus, or bottom of glacier, retreats, zone of glacial influence will be reached eventually between accumulation and wastage, and finally the glacier disappears; but no matter how margin is moving ice within the glacier continues to flow, but not enough to stop ablation **Bubble Rock**. Name of this is the Calving process by which a block of glacier breaks off and falls into the glacier that has a valley bottom relatively higher than nearby valleys formed by larger glaciers **Headwall**: a steep cliff, usually the uppermost part of a cirque. **Horn**: a peak or pinnacle thinned and eroded by three or more glacial cirques **Hoarfrost**: a deposit of interlocking ice crystals (hoar crystals) formed by direct sublimation on objects, usually those of small diameter freely exposed to the air, such as tree branches, plant stems and leaf edges, wires, poles, etc.; the surfaces of these objects are sufficiently cooled, mostly by nocturnal radiation, to allow the direct sublimation of the water vapor contained in the ambient air.

Glacier: A semi permanent mass of firm formed by drifting snow behind a snowdrift or a drift glacier. **Chattermarks:** Striations also called a snowdrift glacier or a drift glacier. **Hummock:** Small area of raised ground which is formed as a glacier slowly retreats, leaving behind ground moraine. **Ice apron:** a mass of ice advancing to a mountainside. **Ice cap:** a dome-shaped mass of glacier ice that spreads out in all directions; an ice cap is usually larger than an ice field but less than 50,000 square kilometers (12 million acres). **Ice Cap Glacier:** Mounds of ice that submerge peaks and ridges at the crest of a mountain range. **Ice cave:** a cave of ice, usually underneath a glacier and formed by meltwater; cave

Glaciers have a characteristic, circular shape, with their width as wide or wider at the point of a ridge or saddle between two peaks, typically affording a pass through to another. **Gold Glacier**. Glacier in which most of the ice is below the surface; the glacier's surface may be susceptible to melt due to incoming solar radiation or ice interface. **Ice shelf**: flow that occurs when glacier motion is decelerating down-ice. **Flow**: flow that occurs when glacier motion is decelerating down-ice. **Ice cover**: land overlaid by snow or ice. **Ice cliff**: boundary separating opposing flow directions of ice on a glacier or ice sheet. **Ice dome**: ice surface with parabolic surface; located in accumulation zone. **Ice floe**: a shoving of ice caused by crevasse formation or jerk motion. **Ice rise**: when ice gets on top of rock in the seabed; these happen to ice shelves; they are usually dome shaped. **Ice sheet**: a dome-shaped mass of

Snow: metamorphism that adds molecules to sharpen the corners of a mineral. **Glacier**: a glacier that covers much of a continent or large island; a small glacier that is armchair shaped. **Cordilleran Ice Sheet**: the ice cap that contains in the northwestern part of North America during the Pleistocene. **Hill**: followed by a tail of **Glaciocayote**: Depression in a glacier formed by shear and melt the surrounding snow. **Crevasses**: Open fissure in the ice sheeted, these happen to ice sheets, they are usually tongue shaped **ice sheet**: a tongue-shaped mass of glacier ice that covers surrounding terrain and is greater than 50,000 square kilometers (12 million acres), the Greenland and Antarctic ice sheets. **Ice Shelves**: ice sheet attached to land, extends over sea, floats on water. **Ice stream**: (i) a current of ice in an ice sheet or ice cap that flows faster than the surrounding ice (ii) sometimes refers to the confluent sections of a branched glacier tongue (3) obsolete synonym of valley glaciers. **Tongue**: a long and narrow sheet of ice projecting out from the coastline to the ocean. **Tee-**

A kind of hoarfrost; ice crystals that develop by sublimation in glacial air. A snowdrift; a mass of snow and debris deposited by wind. **Cave**: a hollow in the snowdrifts and drifts of the snowounding snow. **Cavasie**: open tissue in the interior of a glacier. **Cemented glacier**: a rock glacier that has interstitial ice a meter or so below the surface. **Ice-cored glacier**: a rock glacier that has a buried core of ice. **Icefall**: part of a glacier with rapid flow and a chaotic crevassed surface; occurs where the glacier bed steepens or narrows. **Icefield**: a mass of glacier ice; similar to an ice cap, and usually smaller and lacking a dome-like shape; somewhat controlled by terrain. **Interglacial**: a period of time between two glaciations. **Jokulhlappt**: (1) a large outburst flood that usually occurs when a glacially dammed lake drains catastrophically; (2) any catastrophic release of water from a glacier. **Kame**:

a wide range of sizes derived from a broad provenance. **Dirt Cone:** parallel ridges like ripple marks. Only a few 100 ft apart. **Dirt Cone:** that is covered by dirt; a dirt cone is caused by a differential pattern of surface and bare ice. **Drain Channel:** Preferred for meltwater to snow cover. **Drift:** Any material carried and deposited by a glacier. Drift is firm formed by drifted snow behind obstructions or in the rounded hollows of a glacially dammed lake. **Catastrophic (or any catastrophic) release of water:** from a glacier. **Kame:** an irregularly shaped hill or mound composed of sand, gravel and till that accumulates in a depression on a retreating glacier, and is then deposited on the land surface with further melting of the glacier. **Kettle Hole:** A circular depression in the ground made when a block of ice calves off the toe of a glacier, becomes buried by till, and later melts. **Kerides:** irregular till thickness and depressions where large blocks of ice melted within them. **Knife Edged Ridge / Pointe D Peaks:** ridges between widening u-shaped glacial

or a snowdrift glacier **Dropstone**: A rock that was carried elsewhere by the ice, melted, the rock sinks to the bottom of the body of water today. **Drumlin**: Remnant elongated hills formed by historical glacial flow as they were formed and why they form only in some glaciated regions. **Dry** soil that it remains frozen to the substrate, also called a polar glacier.

valleys that become narrower until they rise steeply to narrow, arêtes/pointy pyramids. **Lateral moraine**: a ridge-shaped moraine deposited at the side of a glacier and composed of material eroded from the valley walls by the moving glacier. **Laurentide Ice Sheet**: The continental glacier that covered eastern Canada and parts of the northeastern United States during the Pleistocene Epoch. **Leward side**: Of a natural or man-made elements that does not receive wind. **Loess**: wind-blown silt deposit blown away from the floodplains and bars of the outwash streams that built up as sand dunes and a frostline of fine silt. **Luis**

material (typically gravel or sand) deposited by streams that cut channels in **Erosion**. Large pieces of rock that have been transported away from their sheets **Extending flow**: when glacier motion is accelerating down-slope and dark on a glacier that were formed by rock avalanching **Fjord** **Glacial Formation**: layering in glacial ice that has distinctive crystal size and/or top of, or near the side of a glacier that drains meltwater out of the glacier; usually kept open by the frictional heating of flowing water that melts the ice walls of the conduit **Moraine**: a mound, ridge, or other distinct accumulation of glacial till **Moraine shoal**: glacial moraine that has formed a shallow place in water **Moulin**: a nearly vertical channel in ice that is formed by flowing water; usually found at a relatively flat section of glacier in a region of transverse crevasses; also called a pothole **Negative mill**: a glacier that is confined by surrounding mountain terrain; also called an alpine glacier **Mountain glacier**: a

is caused by stress and deformation that a glacier experiences as it flows over or originates as a sedimentary feature. **Finger bands:** alternate bands of light and dark, steep narrow icefalls and thought to be the result of different conditions in summer and winter. **Frontal stripes:** shallow, parallel grooves on the face of a glacier; a fountain that develops when water from a conduit is forced up to the surface of a glacier. **Niche geyser:** a fountain that develops when water from a conduit is forced up to the surface of a glacier. **Niche glacier:** very small glacier that occupies gullies and hollows on north-facing slopes (northern hemisphere); may develop into cirque glacier if conditions are favorable. **Ninatac:** a rocky crag or small mountain projecting from and surrounded by a glacier or ice sheet. **Osgoo:** alternate bands of light and dark ice seen on a glacier surface. **Park-summits:** like ice-walls. They kind of bend towards the middle, indicating a flow direction.

before; occurs when a glacier flows downvalley faster than the rate of ablation; a boulder swept from its place of origin by glacier advance or retreat after glacial melt; after glacial melt, the boulder might be stranded in a field of its type or size class **Glacial Formation**: (1)Loose snow (>90% air) (2) Glacial air), 4) fine-grained ice (<20% air), 5) coarse grained ice (<20% air) (6) Glacial till; consists of mostly symmetrical geometries displaced across middle of the glacier flows faster than the sites **Oubloub flood**: any catastrophic flooding from a glacier, may originate from trapped water in cavities inside a glacier or at the margins of glaciers or from lakes that are dammed by flowing glacers **Outlet glacier**: a valley glacier which drains an inland ice sheet or ice cap and flows through a gap in peripheral mountains **Outwash Plain**: Formed when sand is eroded, transported and deposited by meltwater streams from the glaciers snout and nearby till deposits to areas in front of the glacier. **Patterned grounds**: consists of mostly symmetrical geometries displaced across

it into the bedrock by gravel and rocks carried by glacial ice and meltwater. **Glacial Incorporation:** A form of glacial erosion where the ice surrounds and moves over the ice. **Glacial Rebound:** The process by which the surface of an overlying continental ice sheet melts away and the weight of the ice is removed. **Glacial Retreat:** When the position of a mountain glacier's terminus is moving back as a glacier ablates more material at its terminus than it transports.

grooves or gouges cut into the bedrock by gravel and rocks carried by glacial grooves **Glacial Subsidence**: The sinking of the surface of an overlying glacial ice sheet. **Glacial till**: accumulations of unsorted, fine, sand, gravel, and boulders; the usual composition of a moraine. **Glacial trough**: a large u-shaped valley formed from a v-gouge of a glacier. **Glacial plow**: large ice lobe spread out or otherwise subject to repeated freezing and thawing. **Plastic glacier**: large ice lobe spread out over surrounding terrain, associated with the terminus of a large mountain valley glacier. **Pingo**: also called hydroglaciation or bogsuniaah, is a mound of earth-covered found in the Arctic and subarctic that can reach up to 70 metres in height and up to 600 m in diameter. **Plastic Zone**: place where cracks cannot form in the glacier. **Plastic Flow**: slow movement of a glacier in which ice crystals slip over each other. **Plastic Deformation**: When a sufficient load is applied to a material, it will cause the material to change

glacier entirely below freezing, except possibly for a thin layer of melt near the surface during summer near the bed; polar glaciars are found only in polar regions of the globe or at high altitudes. **Pothole:** a vertically vertical channel in ice that is formed by flowing water; usually found after a relatively flat section of glacier in a region of transverse crevasses; also called a moulin. **Push moraine:** c. moraine built ahead an advancing glacier. **Quaternary:** geological period of the last Cenozoic; two million years ago to the present. The name refers to the fourth interval of earth time, according to early geologists. **Reconstructed glacier:** a glacier that is reconstructed or reconstituted out of other glacier material; usually formed by seracs falling a hanging glacier then re-adhering; also called reconstructed glacier, regenerated glacier, or glacier re-inie. **Reconstructed glacier:** a glacier that is reconstructed or reconstituted out of other glacier material; usually formed by seracs falling from a hanging glacier then re-adhering; also called reconstructed glacier, regenerated glacier, or glacier remains. **Ressional Moraine:** Terminal moraine after it has come back and is moving forward again. **Regelation:** motion of an object through ice by melting and freezing that is caused by pressure differences; this process allows a glacier to slide past small obstacles on its bed. **Regenerated glacier:** a glacier that is reconstructed or reconstituted out of other glacier material; usually formed by forces falling from a hanging glacier then re-adhering; also called reconstructed or regenerated glacier. **Retreat:** when a mountain glacier's terminus doesn't extend as far downvalley as it previously did; occurs when ablation surpasses accumulation. **Retreating glacier:** a glacier whose terminus is increasingly retreating upvalley compared to its previous position due to a higher level of ablation compared to accumulation. **Ribbon Lake Long:** thin lakes that form after a glacier retreats that rock-laden flows. **Rock flour:** a fine powder of silt- and clay-sized particles that a glacier creates as its rock-laden scrapes over bedrock; usually flushed out in meltwater streams, causing water to look powdery gray; eels and oysters that fill with glacier flour may develop a banded appearance. **Rock glacier:** looks like a mountain glacier and has active flow; usually includes a poorly sorted mess of rocks and fine materials; may include: (1) interstitial ice a meter or so below the surface ("ice-cemented"), (2) a buried core of ice ("ice-cored"), and/or (3) rock debris from avalanching snow and rock. **Rock Mountain Rock drumlinogen Moraine:** (also called ribbed moraine) is a subglacially (i.e. under a glacier or ice sheet) formed type moraine landform, that mainly occurs in Fennoscandia, Scotland, Ireland and Canada. Landform assemblage of numerous, parallel, closely-spaced ridges consisting of glacial drift, usually TILL. The ridges formed transverse to ice flow in a subglacial position, and are usually found in the central portions former ice sheets, believed to have been the central areas of the ice sheets. Formation linked closely Drumlin. **Sastrugi:** parallel wave-like ridges caused by winds on the surface of hard snow, especially polar regions. **Sedimentary ogives:** alternating bands of light and dark at the firm limit of a glacier; light bands are usually young and lightest at the highest level up-glacier, becoming increasingly older & darker as they progress down-glacier. **Serac:** an isolated block of ice that is formed where the glacier face is fractured. **Schistwannen:** curved grooves formed by water under immense pressure at the base in side valleys. **Kinematic waves:** refer to a wave of ice moving downglacier propagated by its increased thickness. The wave of ice may move at two to six times the velocity of surrounding thinner ice. **Lahar:** is a mudflow or debris flow originating on a volcano. Jokulhlaups (see above) often become lahars when they incorporate the rock debris that lies within their path. **Lateral moraine:** a moraine formed at the side of a glacier. Piles of loose unsorted rocks along the side margins of the glacier. The rocks may be pushed there by the moving ice or dumped from the glacier's rounded surface. **Mass balance:** describes the net gain or loss of snow and ice through a given year. It is usually expressed in terms of water gain or loss. **Neoglaciation:** refers to the advances made by mountain glaciers since the great Pleistocene ice age. In the Cascades the advances have occurred since 6,600 years before present. **Ogives:** are arc-shaped features occasionally found across the glacier surface below icefalls. They may be ridges and swales in the ice or bands of darker or lighter ice. One theory of their formation suggests that the ice is stretched sometimes diluted when exposed in the icefall during the high velocities of summer; it is compressed during the winter so that bands of different ice thickness form. **Perfectly plastic solid:** is a solid that does not deform until it reaches a critical value of stress, after which it will yield indefinitely. Some glaciologists say that ice is a perfectly plastic substance. (That is, brittle and capable of cracking like a solid, yet deformable and capable of flowing at other stresses.) **Pleistocene:** is the period of earth's history, roughly two million years ago to about ten thousand years ago, characterized by the advance and recession of continental ice sheets. **Roche moutonnée:** is a small asymmetrically-shaped hill formed by glacial erosion. The upper sides are rounded and smoothed and the lower sides are rough and broken due to quarrying by the glacier. **Serrations:** are the pinnacles of ice formed where the glacier surface is torn by sets of crevasses. **Striations:** are the scratches etched into the rock at the bed of a glacier. Their presence indicates grinding of sand and rock particles into the bed under considerable pressure. In some places fine-grained debris polished the bedrock to a lustrous surface finish called glacial polish. **Sunup:** is a small depression on a glacier or firm surface formed by melting and evaporation resulting from direct exposure to the sun. **Trimlines:** are the sharp vegetative boundaries delimiting the upper margin of a former glaciation. The age differences of the ground surface are often visible because of different ages of the vegetation. **Ice Core:** Ice cores are cylinders of ice drilled out of an ice sheet or glacier. Most ice cores records come from Greenland and Antarctica, and the longest ice cores extend to 3km in depth. The oldest continuous ice core records to date extend 123,000 years in Greenland and 800,000 years in Antarctica. Ice cores contain information about past temperature, and about many other aspects of the environment. Crucially, the ice encloses small bubbles of air that contain a sample of the atmosphere – from these it is possible to measure directly the past concentration of gases (including carbon dioxide and methane) in the atmosphere.

