

# A decade of SimPEG

## connecting research & industry through open-source software

Lindsey Heagy  
Assistant Professor  
UBC GIF

# challenges & opportunities for EM in mineral exploration

challenges:

- need to see deeper
- rugged topography
- highly conductive targets
- multiple surveys types (diff. resolution, physical properties)
- other factors: infrastructure, ...

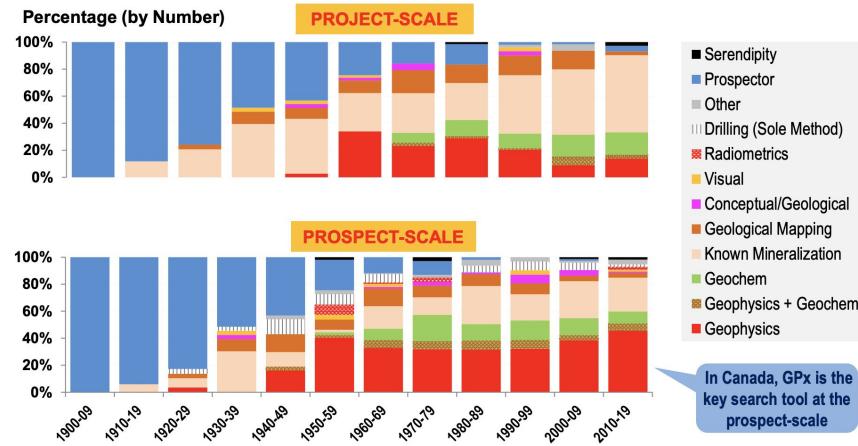
but have

- high quality data
- multiple data types (geophysical, geology, geochemistry...)
- computational machinery

opportunities...



Primary search methods used by Country  
Non-Bulk mineral discoveries in CANADA : 1900-2019



# GIF over the decades

(we can even invert Geotech data!)

- developed codes to invert individual types of geophysical data
- goal: extract information about physical properties, generate models relevant to mineral exploration
- connections with industry core to research
  - data & research problems
  - transfer of software & technology

Geophysical Journal International

*Geophys. J. Int.* (2010) 182, 168–182

doi: 10.1111/j.1365-246X.2010.04634.x

## Three-dimensional inversion of ZTEM data

Elliot Holtham and Douglas W. Oldenburg

UBC-Geophysical Inversion Facility, Dept.  
E-mail: [eholtham@eos.ubc.ca](mailto:eholtham@eos.ubc.ca)

Accepted 2010 April 17. Received 2010

Geophysical Journal International

*Geophys. J. Int.* (2014) 196, 1492–1507

Advance Access publication 2013 December 5

doi: 10.1093/gji/ggt465

## 3-D inversion of airborne electromagnetic data parallelized and accelerated by local mesh and adaptive soundings

Dikun Yang, Douglas W. Oldenburg and Eldad Haber

Geophysical Inversion Facility, Department of Earth, Ocean and Atmospheric Sciences, University of British Columbia, Vancouver

## 3D parametric hybrid inversion of time-domain airborne electromagnetic data

Michael S. McMillan<sup>1</sup>, Christoph Schwarzbach<sup>1</sup>, Eldad Haber<sup>2</sup>, and Douglas W. Oldenburg<sup>1</sup>

the Earth's  
measurements  
and modeling

doi: 10.1093/gji/ggw256

Geophysical Journal International

*Geophys. J. Int.* (2016) 207, 174–196  
Advance Access publication 2016 July 13  
GJI Marine geosciences and applied geophysics

## On recovering distributed IP information from inductive source time domain electromagnetic data

Sogei Kang and Douglas W. Oldenburg

Department of Earth, Ocean and Atmospheric Sciences, University of British Columbia, BC V6T 1Z4, Canada. E-mail: [doug@eos.ubc.ca](mailto:doug@eos.ubc.ca)

Accepted 2016 July 8. Received 2016 June 22; in original form 2015 September 3

### SUMMARY

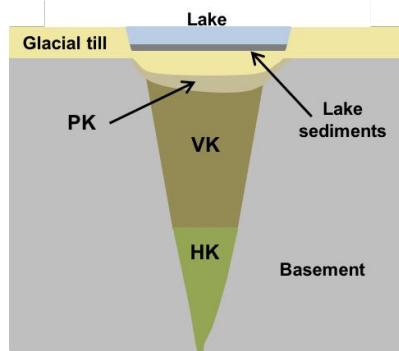
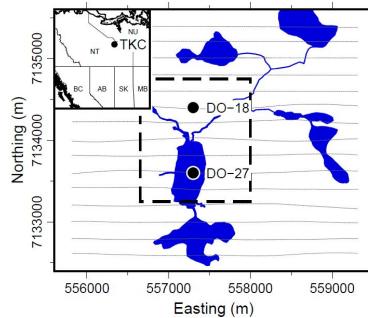
We develop a procedure to invert time domain induced polarization (IP) data for inductive sources. Our approach is based upon the inversion methodology in conventional electrical IP (EIP), which uses a sensitivity function that is independent of time. However, significant modifications are required for inductive source IP (ISIP) because electric fields in the ground do not achieve a steady state. The time-history for these fields needs to be evaluated and then used to define approximate IP currents. The resultant data, either a magnetic field or its derivative, are evaluated through the Biot-Savart law. This forms the desired linear relationship between data and pseudo-chargeability. Our inversion procedure has three steps: (1) Obtain a 3-D background conductivity model. We advocate, where possible, that this be obtained by inverting early-time data that do not suffer significantly from IP effects. (2) Decouple IP responses embedded in the observations by forward modelling the TEM data due to background conductivity and subtracting these from the observations. (3) Use the linearized

### ABSTRACT

We have developed a method to invert time domain induced polarization (IP) data for inductive sources. Our approach is based upon the inversion methodology in conventional electrical IP (EIP), which uses a sensitivity function that is independent of time. However, significant modifications are required for inductive source IP (ISIP) because electric fields in the ground do not achieve a steady state. The time-history for these fields needs to be evaluated and then used to define approximate IP currents. The resultant data, either a magnetic field or its derivative, are evaluated through the Biot-Savart law. This forms the desired linear relationship between data and pseudo-chargeability. Our inversion procedure has three steps: (1) Obtain a 3-D background conductivity model. We advocate, where possible, that this be obtained by inverting early-time data that do not suffer significantly from IP effects. (2) Decouple IP responses embedded in the observations by forward modelling the TEM data due to background conductivity and subtracting these from the observations. (3) Use the linearized

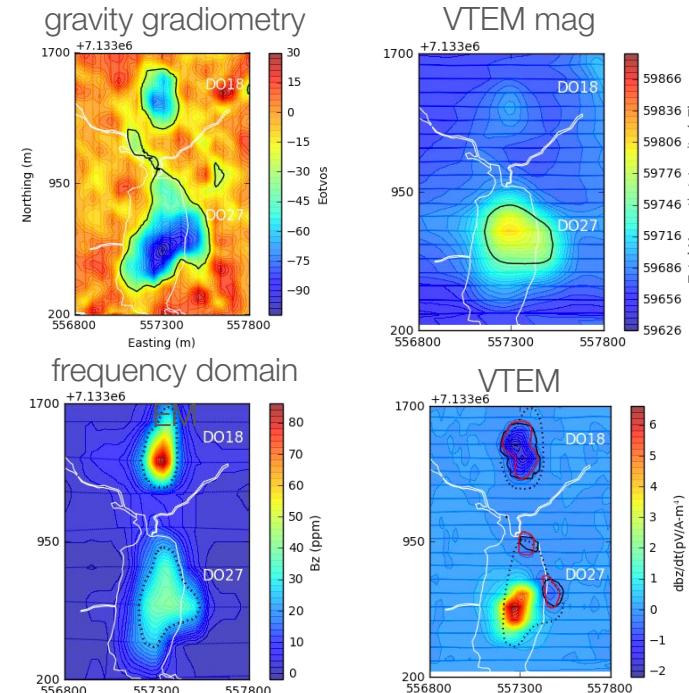
# an example: Tli Kwi Cho Kimberlite (TKC) complex

- Geophysical discovery in 90's
- airborne mag and EM
- 2 kimberlite pipes



Rock type	Glacial till	Host rock	HK	VK	PK
Density	Moderate	Moderate	Low	Low	Low
Susceptibility	None	None	High	Low-moderate	Low-moderate
Conductivity	Moderate-high	Low	Low-moderate	Moderate-high	Moderate-high
Chargeability	Low	Low	?	?	?

- invert to obtain physical property models
- interpret to build quasi-geology model

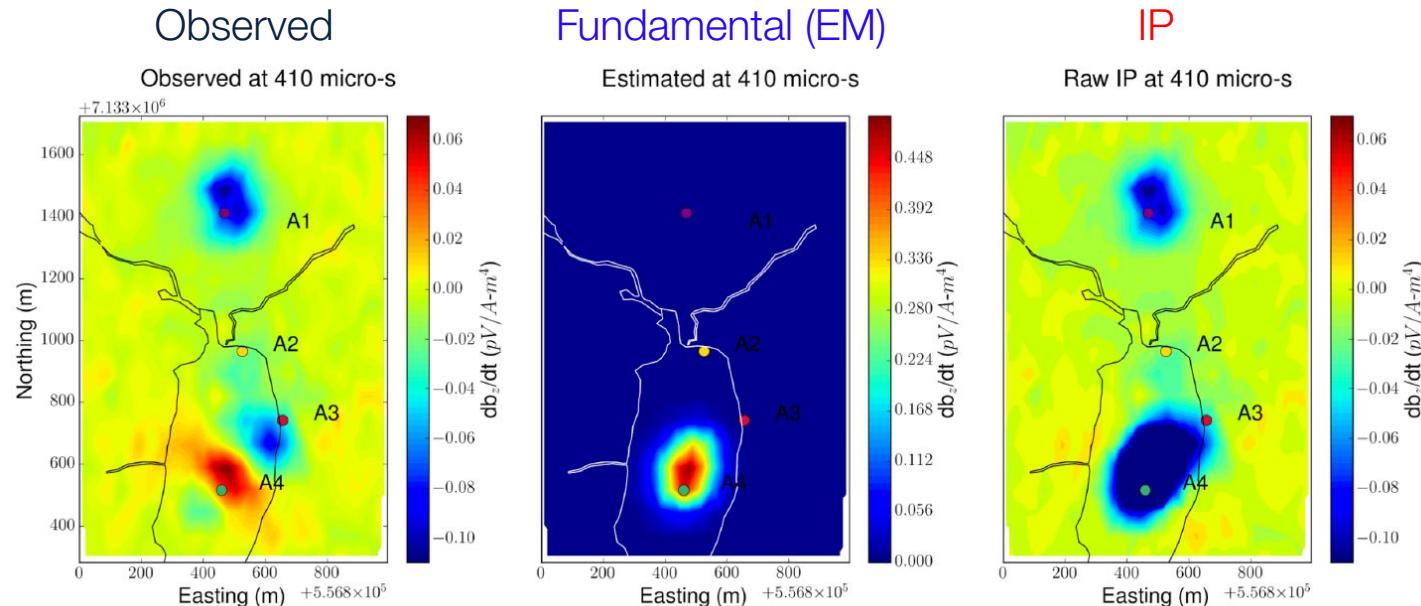




# IP effects in time domain EM data

Negative transients in VTEM presents a challenge → motivates research

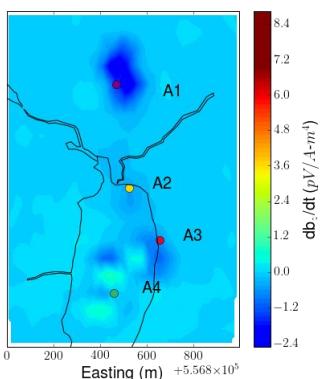
EM-decoupling: **IP** = Observation – Fundamental (EM)



# TKC: IP inversion (early time)

Raw IP at 130 micro-s

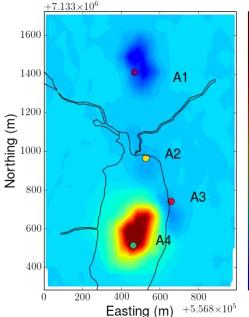
IP data



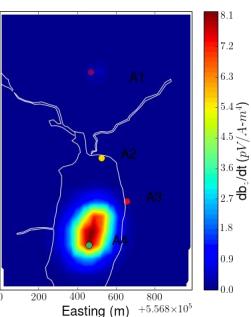
Observation

Fundamental

Observed at 130 micro-s



Estimated at 130 micro-s



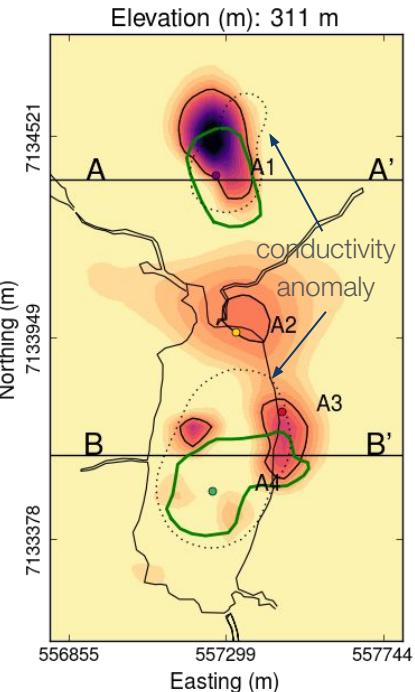
invert

$$d^{IP}(t) = G\tilde{\eta}(t)$$

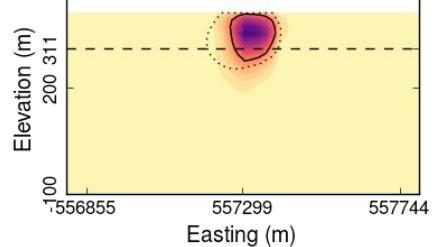
$G(\sigma_\infty)$ : Sensitivity function  
 $\tilde{\eta}$ : Pseudo-chargeability

Kang et al. (2016)

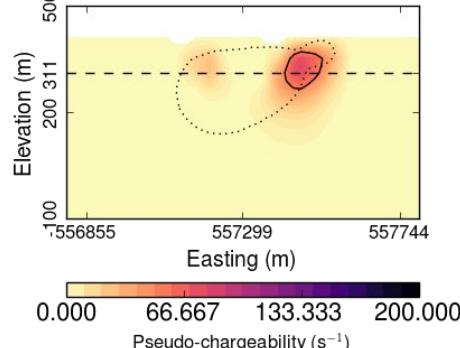
Recovered 3D model



DO-18 A-A'

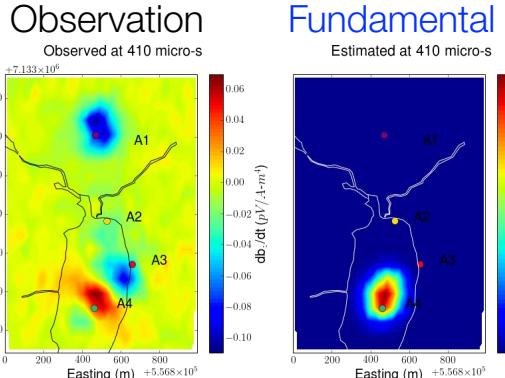
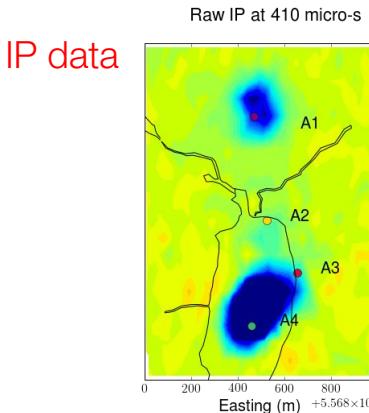


DO-27 B-B'



|IP| = Observation – Fundamental (EM)

# TKC: IP inversion (late time)



|IP| = Observation – Fundamental (EM)

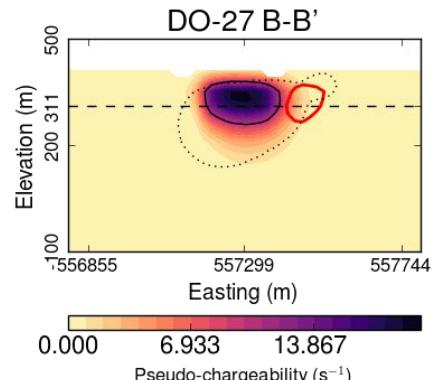
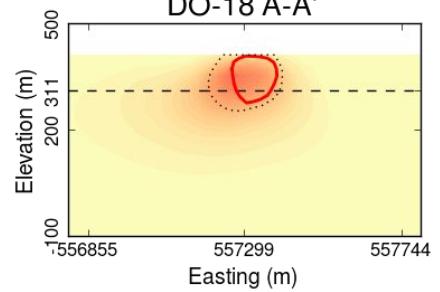
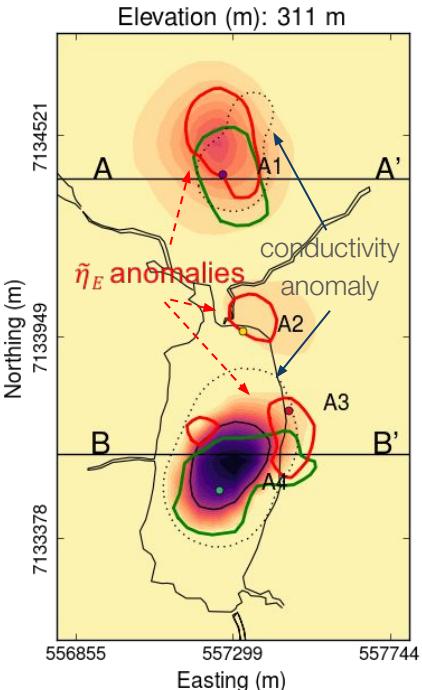
invert

$$d^{IP}(t) = G\tilde{\eta}(t)$$

$G(\sigma_\infty)$ : Sensitivity function  
 $\tilde{\eta}$ : Pseudo-chargeability

Kang et al. (2016)

Recovered 3D model

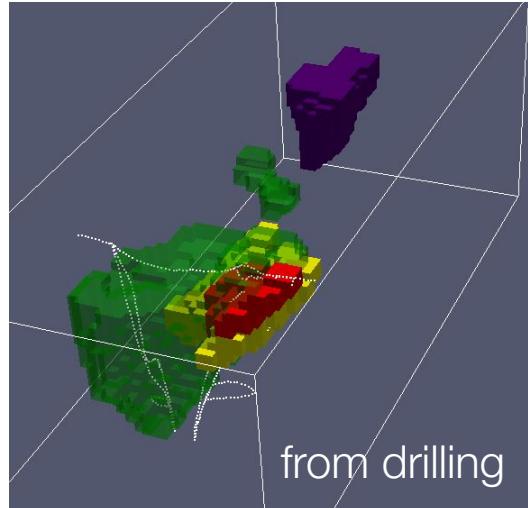
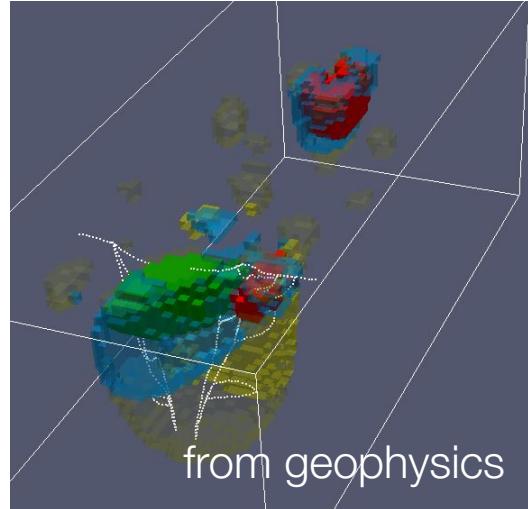


# A quasi-geology model from physical properties

Rock type	Glacial till	Host rock	HK	VK	PK
Density	Moderate	Moderate	Low	Low	Low
Susceptibility	None	None	High	Low-moderate	Low-moderate
Conductivity	Moderate-high	Low	Low-moderate	Moderate-high	Moderate-high
Chargeability	Low	Low	?	?	?

small time constant  
large time constant

- Independently inverted multiple airborne geophysical data sets in 3D, built a representative 3D rock model
- Importance of conductivity, chargeability & related computational tools
- Research grounded in data: helps prioritize questions
- Challenges...?



# research model: challenges & opportunities

Limitations to research mode

- Production code (Fortran) decoupled from research code (Matlab)
  - difficult to alter (even for small changes)
- Fortran software an end-product, structured to carry out one task
  - connecting multiple physical properties would require a complete rewrite

Needs

- extensible, adaptable software → modular software in a “friendly” language
- collaboration → testing, software practices

Tools



GitHub



Python



open source  
initiative®

# motivations for open-source

develop + deploy in same programming language

→ research and tech transfer tightly coupled

community-driven development

→ users can also be contributors

→ reduces “bus-factor” risk

open license facilitates use, adaptation

→ flexibility to extend to new use cases

→ commercial and academic use

interoperate with broader open-source ecosystem

→ leverage advances in other tools

→ more integrated workflows





*Simulation and parameter estimation in geophysics*

common framework for simulations & inversions

accelerate research: build upon others work

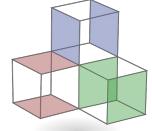
facilitate reproducibility of results

build & deploy in python

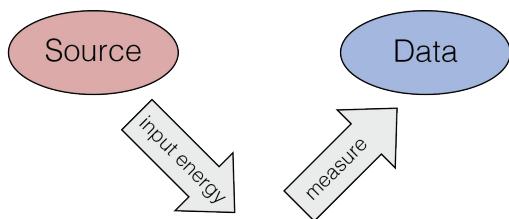
open-source

The screenshot shows a web browser displaying the SimPEG website at [simpeg.xyz](https://simpeg.xyz). The page features a green header with the SimPEG logo and navigation links for WHY, ABOUT, CONTACT, CONTRIBUTE, and a GitHub icon. Below the header is a large image of the 3D cube logo. The main content area has a white background with the title "Simulation and Parameter Estimation in Geophysics". A subtitle explains it as "An open source python package for simulation and gradient based parameter estimation in geophysical applications." A section titled "Geophysical Methods" is shown with a horizontal line, followed by a bulleted list of methods: Gravity, Magnetics, Direct current resistivity, Induced polarization, Electromagnetics (with sub-points Time domain, Frequency domain, and Natural source (e.g.)). To the right of the text is a 3D visualization of a geological model with various colored regions and data points.

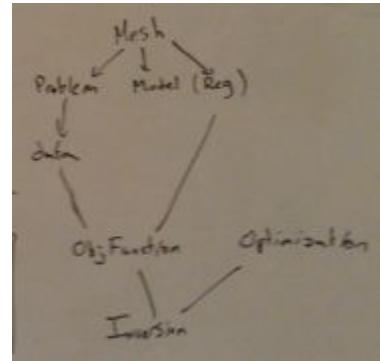
# 10 years coming soon!



goal of developing a common framework for inversions in geophysics



Subsurface:  
Physical properties & contrasts



```
In [1]: import SimPEG
In [ ]: SimPEG.|_
          SimPEG.DataMisfit
          SimPEG.Directives
          SimPEG.InvProblem
          SimPEG.Inversion
          SimPEG.Maps
          SimPEG.Mesh
          SimPEG.Models
          SimPEG.Optimization
          SimPEG.Problem
          SimPEG.Regularization
```

founders



Rowan  
Cockett



Seogi  
Kang



Lindsey  
Heagy

**A framework for geophysical inversions  
with application to vadose zone parameter estimation**

by

Archa Rowan B. Cockett

B.Sc. Applied and Environmental Geology, University of Calgary, 2010

A THESIS SUBMITTED IN PARTIAL FULFILLMENT  
OF THE REQUIREMENTS FOR THE DEGREE OF

**Doctor of Philosophy**

in

THE FACULTY OF GRADUATE AND POSTDOCTORAL

STUDIES

(Geophysics)

The University of British Columbia  
(Vancouver)

December 2017

© Archa Rowan B. Cockett, 2017

# what simpeg solves

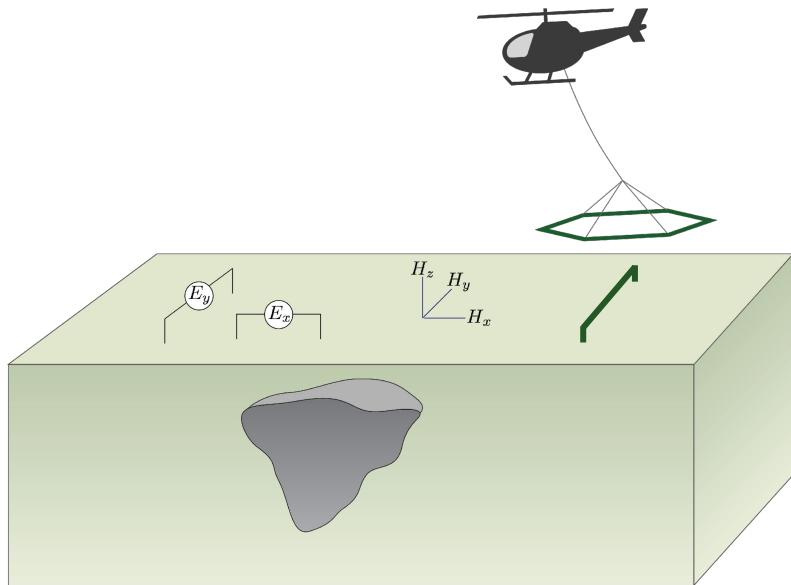
Given

- observations:  $d_j^{obs}$ ,  $j = 1, \dots, N$
- uncertainties:  $\epsilon_j$
- ability to forward model:  $\mathcal{F}[m] = d$

Inverse problem: Find the Earth model that gave rise to the data

$$\min_{\mathbf{m}} \phi(\mathbf{m}) = \phi_d(\mathbf{m}) + \beta \phi_m(\mathbf{m})$$

$$\text{s.t. } \phi_d \leq \phi_d^* \quad \mathbf{m}_L \leq \mathbf{m} \leq \mathbf{m}_U$$

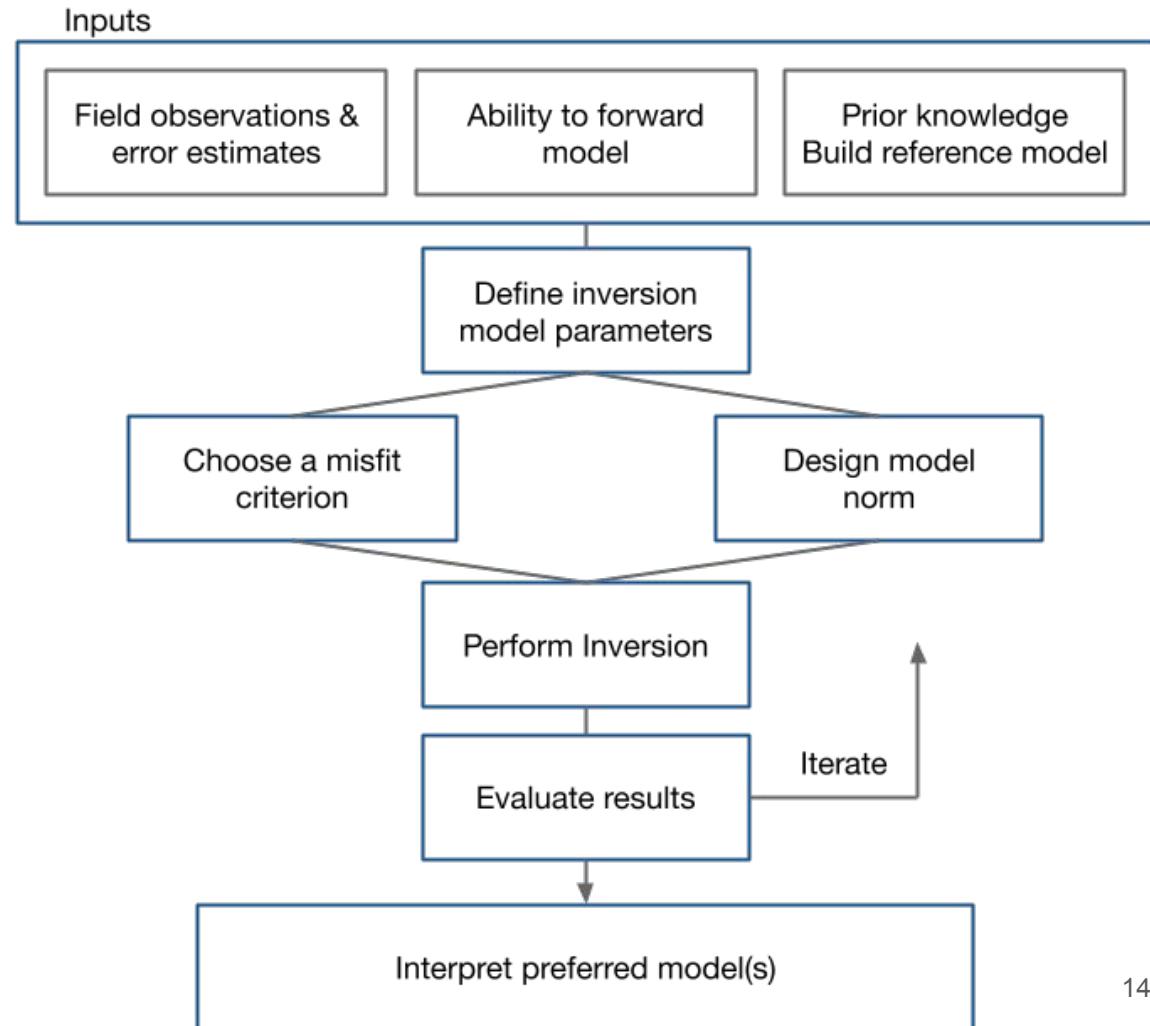


# flow chart for the inverse problem

- many components
- iterative process to obtain solution
- each component requires evaluation, adjustment by user



Fundamentals of Inversion – D. Oldenburg  
Capturing knowledge in code – L. Heagy  
<http://www.mtnet.info/EMinars/EMinars.html>



# the forward simulation: discretize and solve

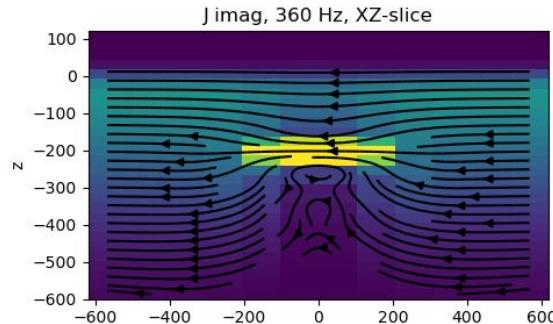
frequency domain electromagnetics

$$\nabla \times \mu^{-1} \nabla \times \vec{E} + i\omega\sigma \vec{E} = -i\omega \vec{J}_s$$

discretize (finite volume)

$$\left( \mathbf{C}^T \mathbf{M}_{\mu^{-1}}^f \mathbf{C} + i\omega \mathbf{M}_\sigma^e \right) \mathbf{e} = -i\omega \mathbf{j}_s$$

$\mathbf{A}$



```
ω = 2 * np.pi * frequency
C = mesh.edgeCurl
Mfμi = mesh.getFaceInnerProduct(1./μ₀)
Meσ = mesh.getEdgeInnerProduct(sigma)

A = C.T * Mfμi * C + i * ω * Meσ
Ainv = Solver(A) # acts like A inverse

rhs = - i * ω * js

e = Ainv * rhs
```

# forward simulation: discretize and solve

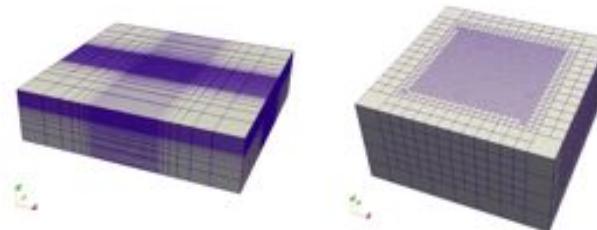
frequency domain electromagnetics

$$\nabla \times \mu^{-1} \nabla \times \vec{E} + i\omega\sigma\vec{E} = -i\omega\vec{J}_s$$

discretize (finite volume)

$$(\mathbf{C}^T \mathbf{M}_{\mu^{-1}}^f \mathbf{C} + i\omega \mathbf{M}_\sigma^e) \mathbf{e} = -i\omega \mathbf{j}_s$$

A



Tensor

OcTree



J. Capriotti

```
ω = 2 * np.pi * frequency
C = mesh.edgeCurl
Mfμi = mesh.getFaceInnerProduct(1./μ₀)
Meσ = mesh.getEdgeInnerProduct(sigma)

A = C.T * Mfμi * C + i * ω * Meσ
Ainv = Solver(A) # acts like A inverse

rhs = - i * ω * js

e = Ainv * rhs
```

# simpeg electromagnetics

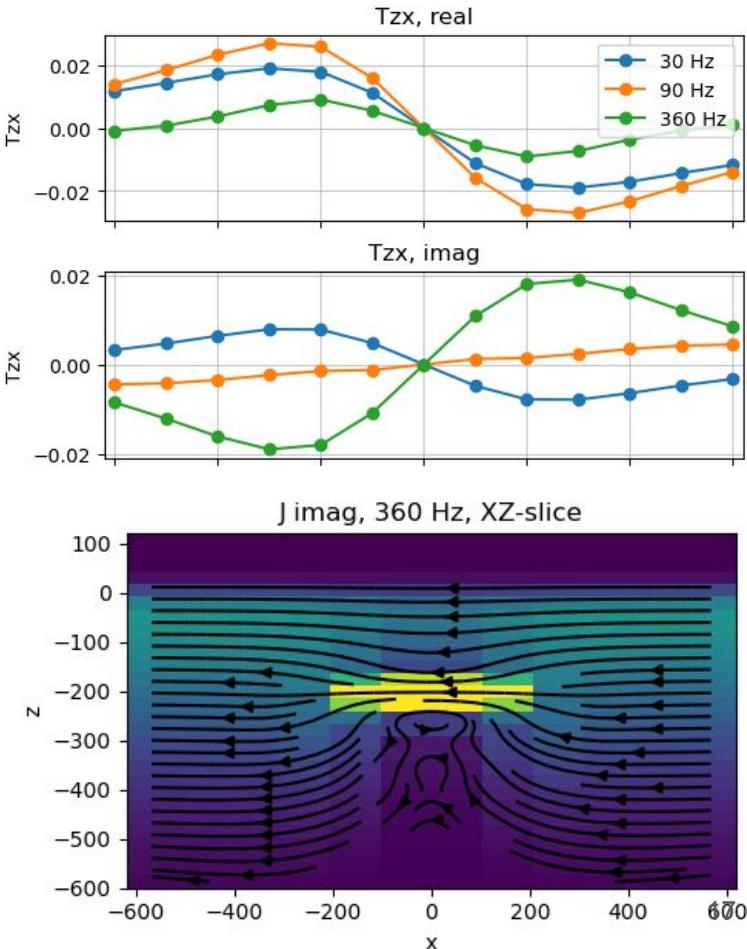
*time and frequency domain*

Controlled source + natural source methods

- Rely on the same methods for calculating fields, sensitivities
- 1D, 2D (in progress), 3D simulation on Tensor, OcTree meshes
  - Primary-secondary approach
  - Boundary-conditions in progress

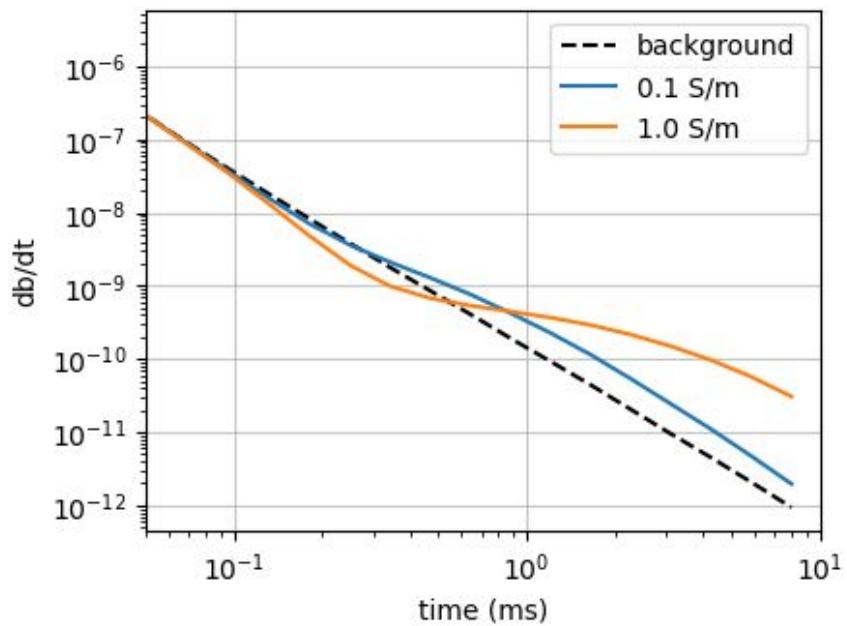
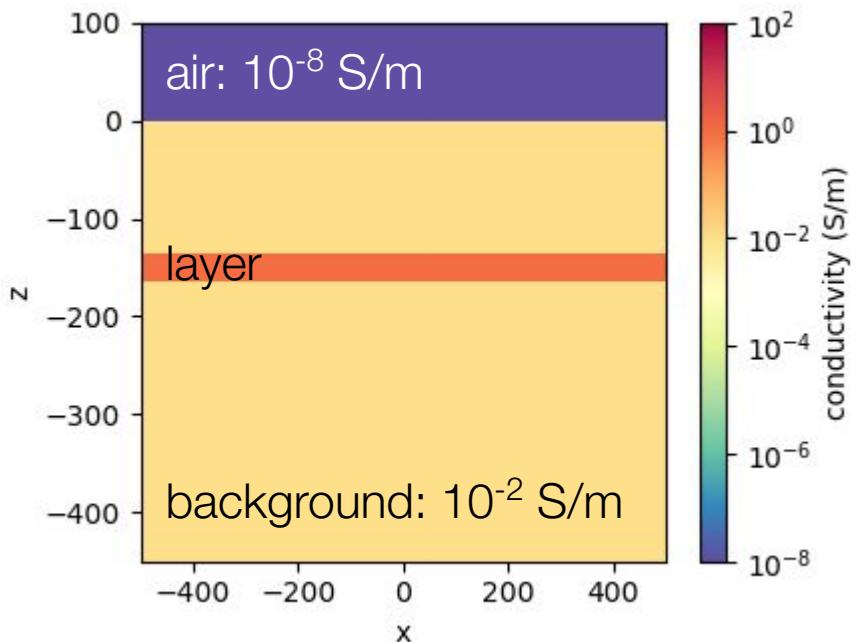
Conventions:

- right handed coordinate system
- z positive up
- $e^{i\omega t}$  Fourier transform



# Time domain simulation: VTEM

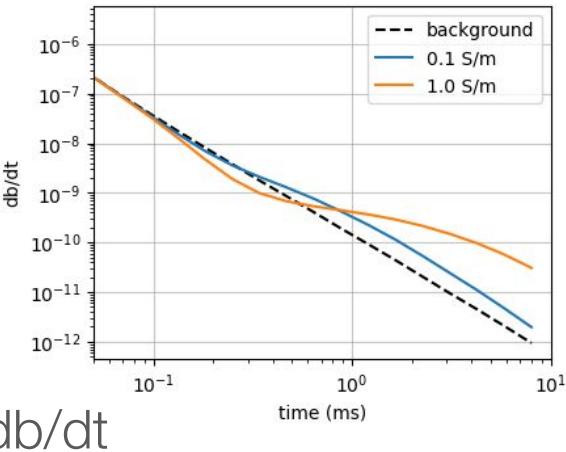
response of a layered Earth



# Time domain simulation:VTEM

halfspace

current density

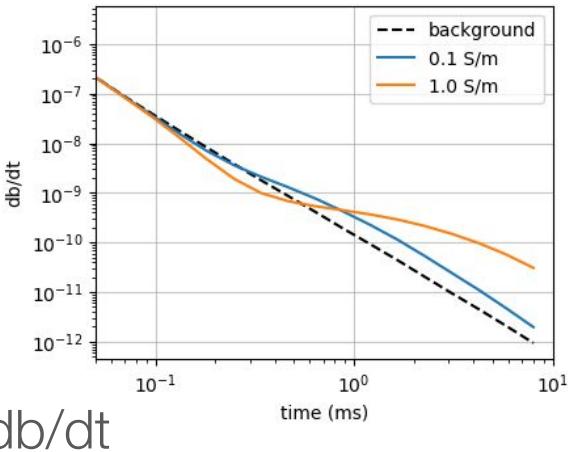
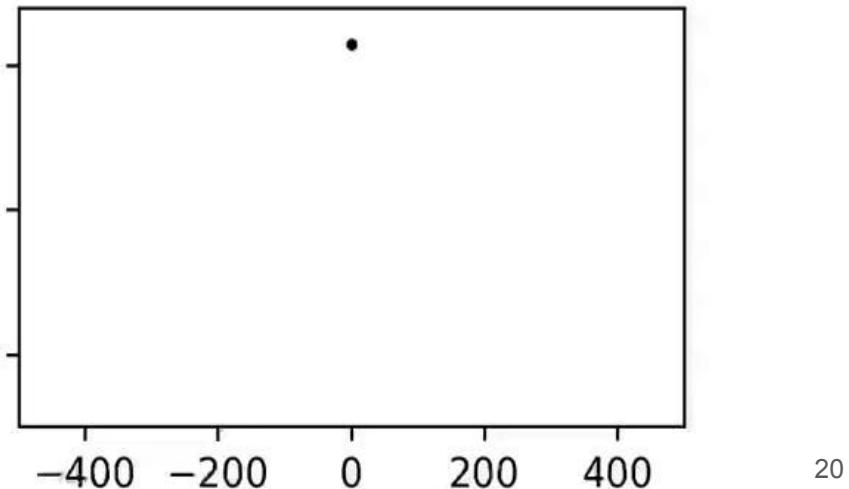
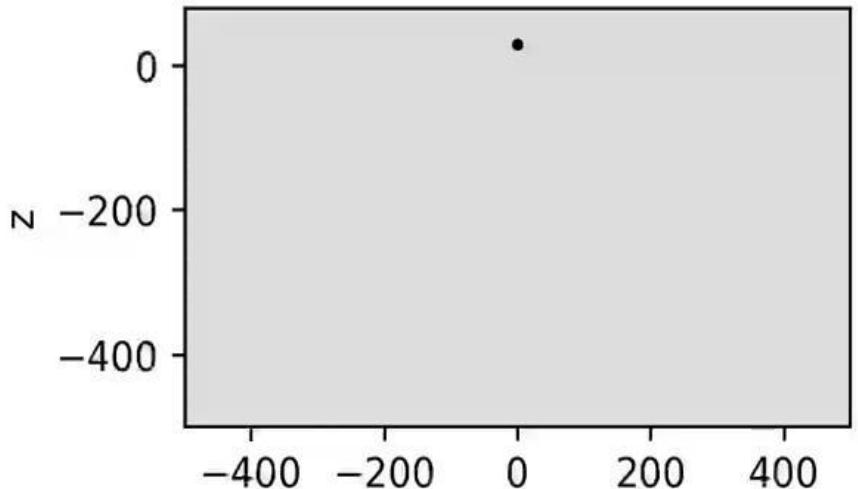


# Time domain simulation:VTEM

halfspace with a conductive layer

current density

$t=0.00$  ms



# Time domain inversion: VTEM

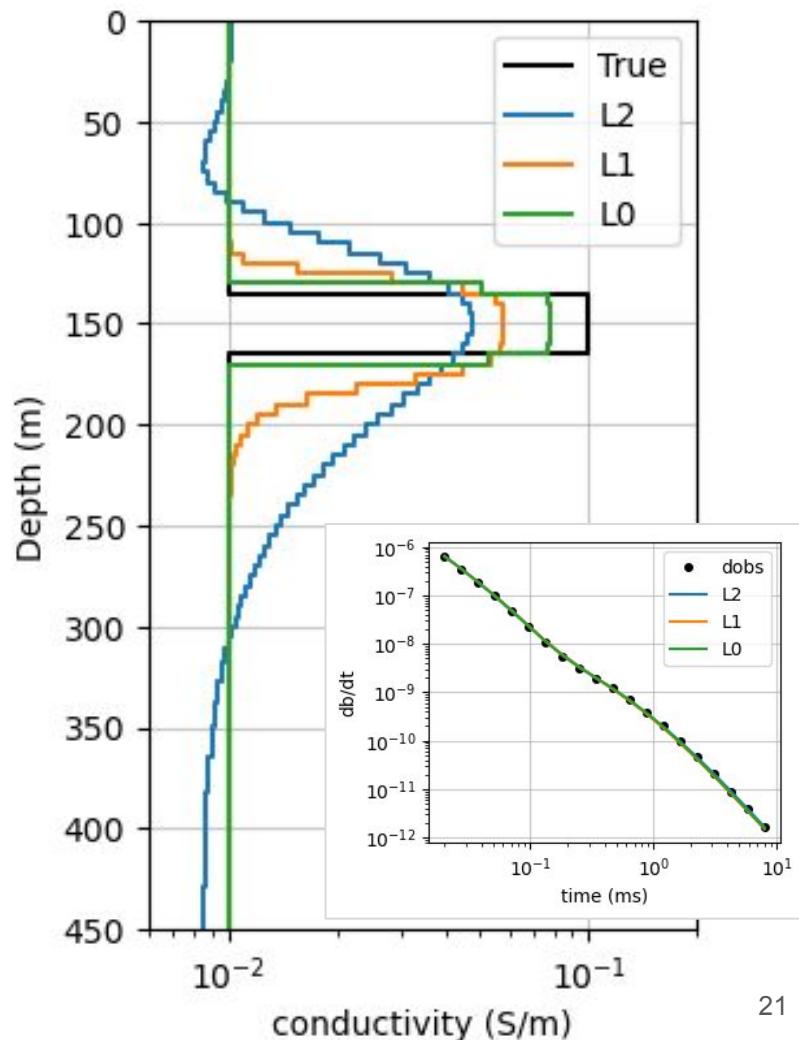
solving the inverse problem

$$\min_{\mathbf{m}} \phi(\mathbf{m}) = \phi_d(\mathbf{m}) + \beta \phi_m(\mathbf{m})$$

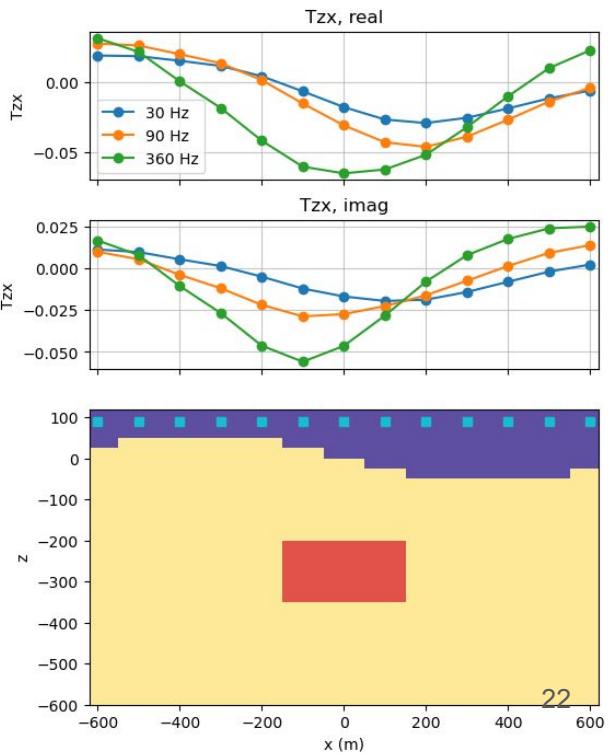
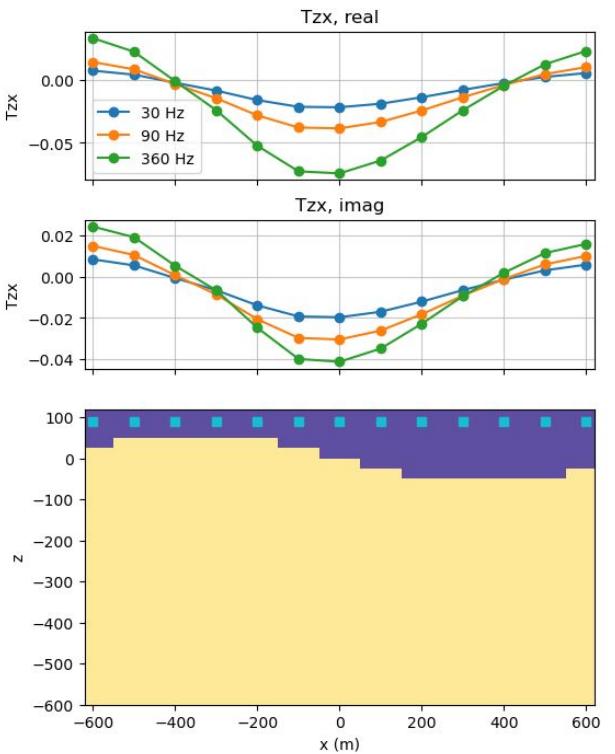
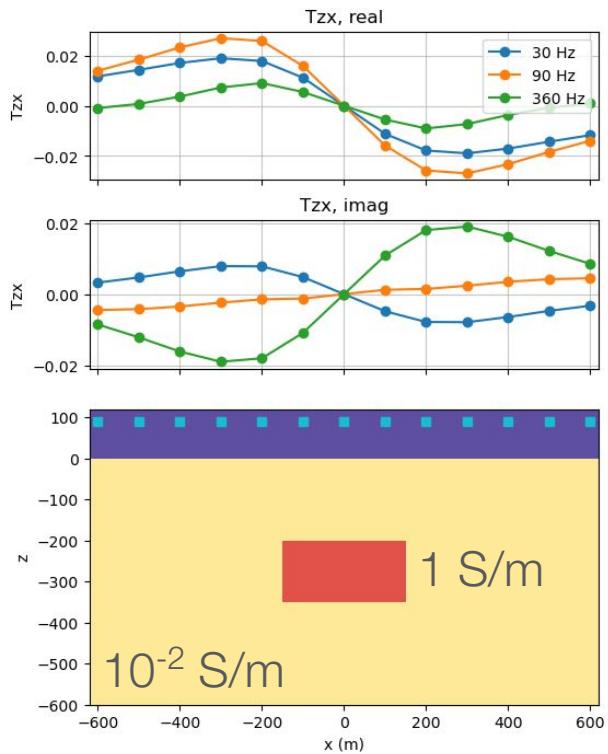
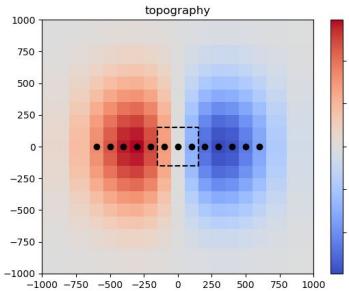
$$\text{s.t. } \phi_d \leq \phi_d^* \quad \mathbf{m}_L \leq \mathbf{m} \leq \mathbf{m}_U$$

choice of model norm

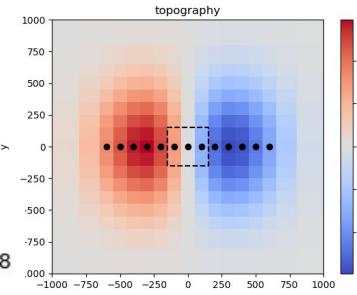
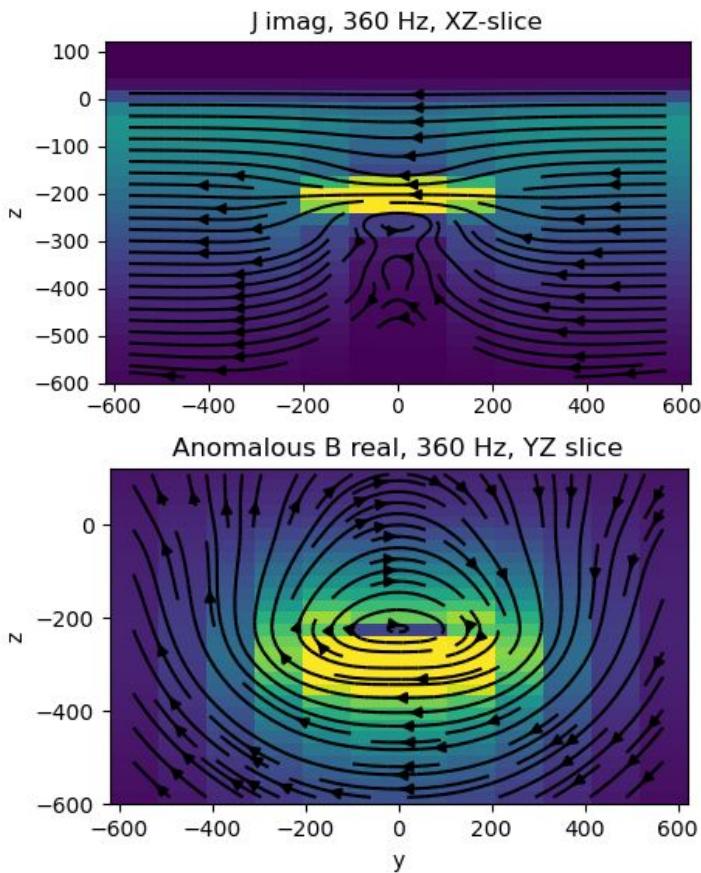
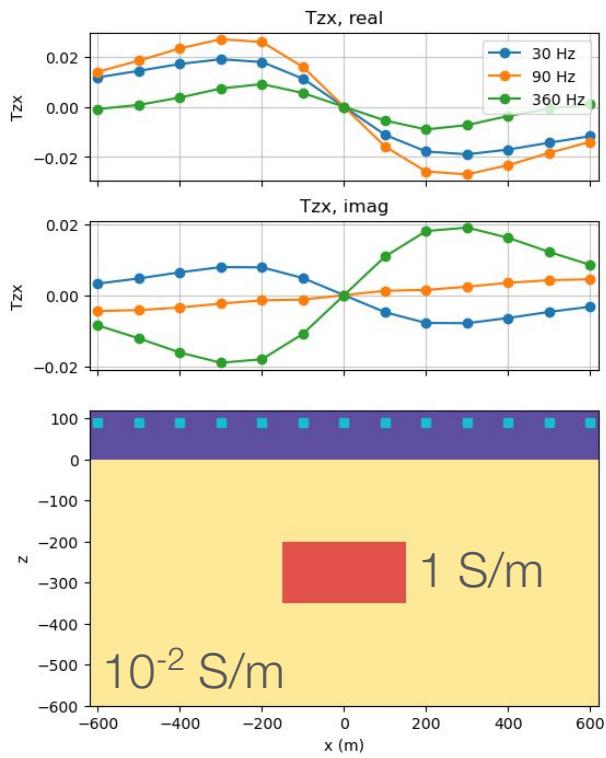
$$\phi_m = \alpha_s \int_v w_s |\mathbf{m} - \mathbf{m}_{ref}|^{p_s} dV + \alpha_x \int_V w_x \left| \frac{d\mathbf{m}}{dz} \right|^{q_z} dV$$



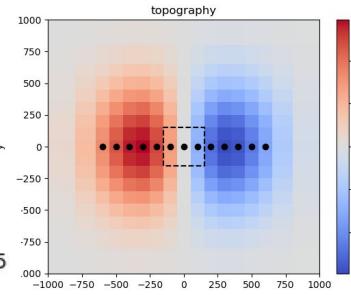
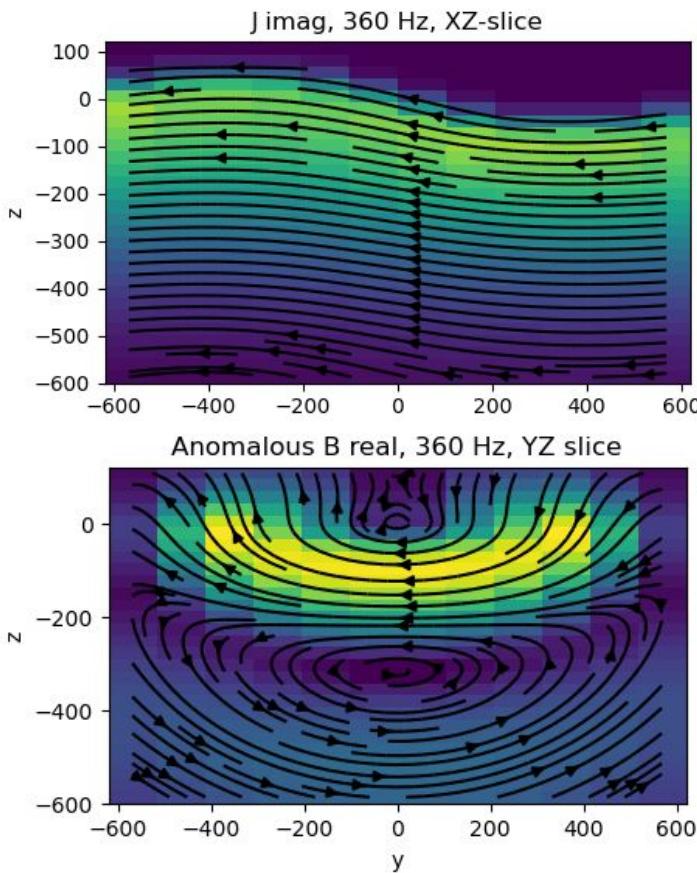
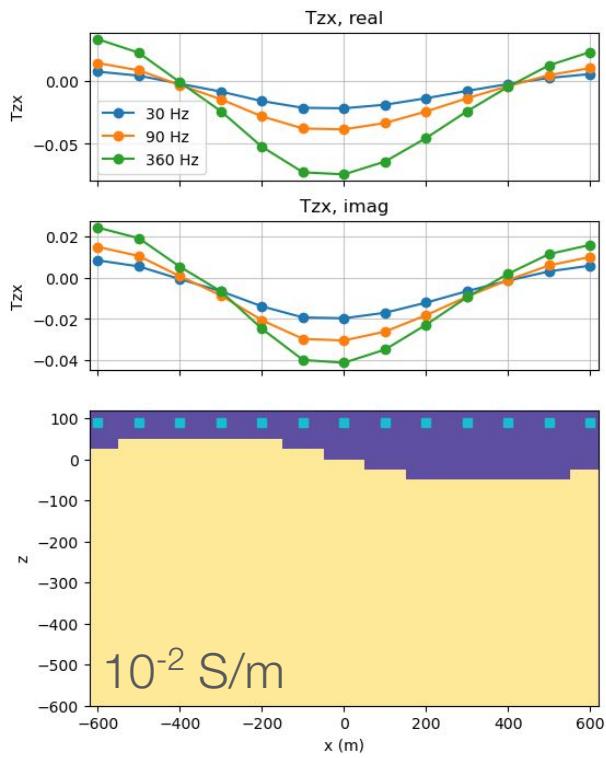
# Frequency domain: ZTEM



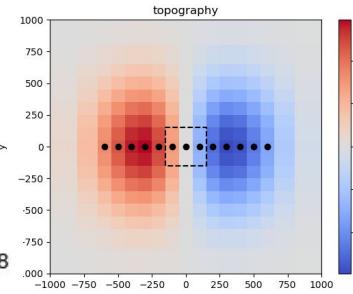
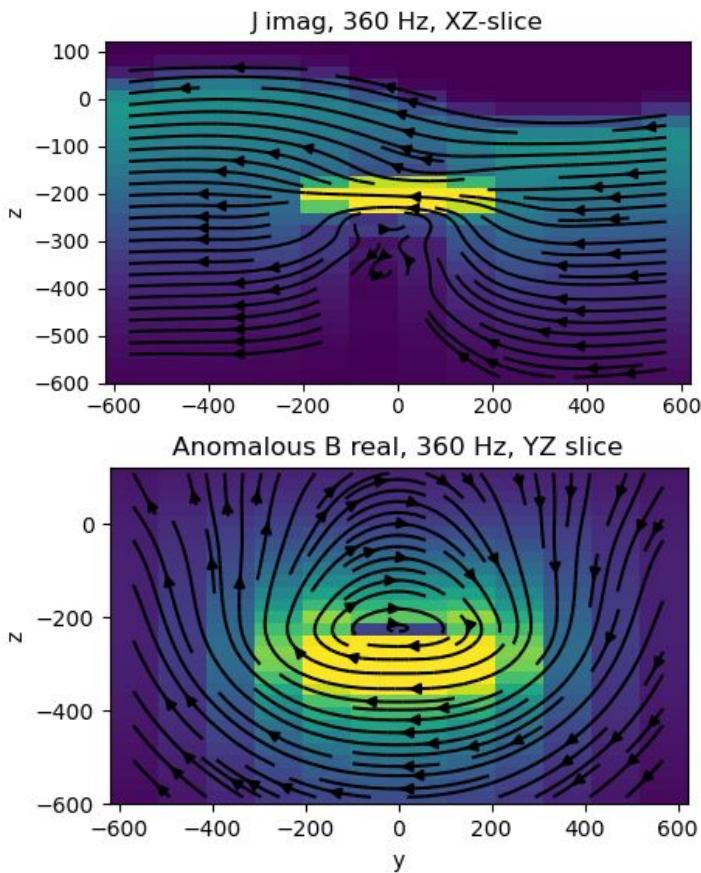
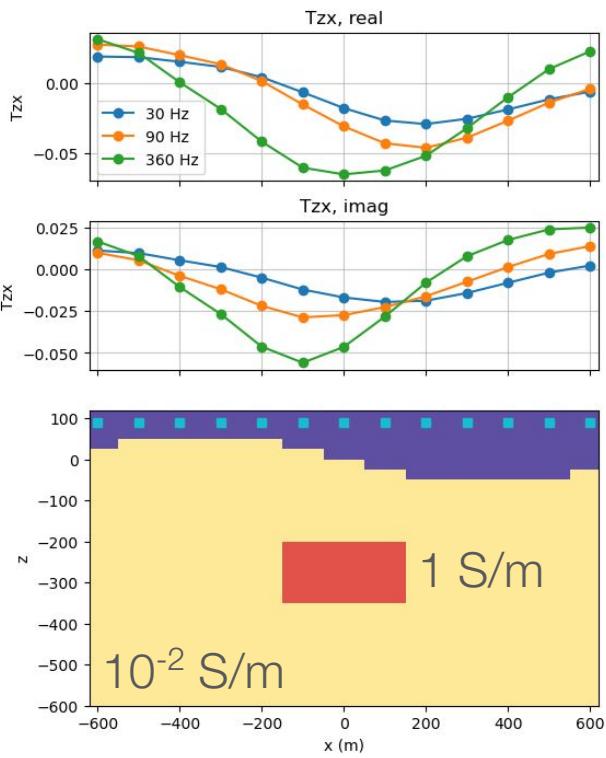
# Frequency domain: ZTEM



# Frequency domain: ZTEM



# Frequency domain: ZTEM



# functionality & status

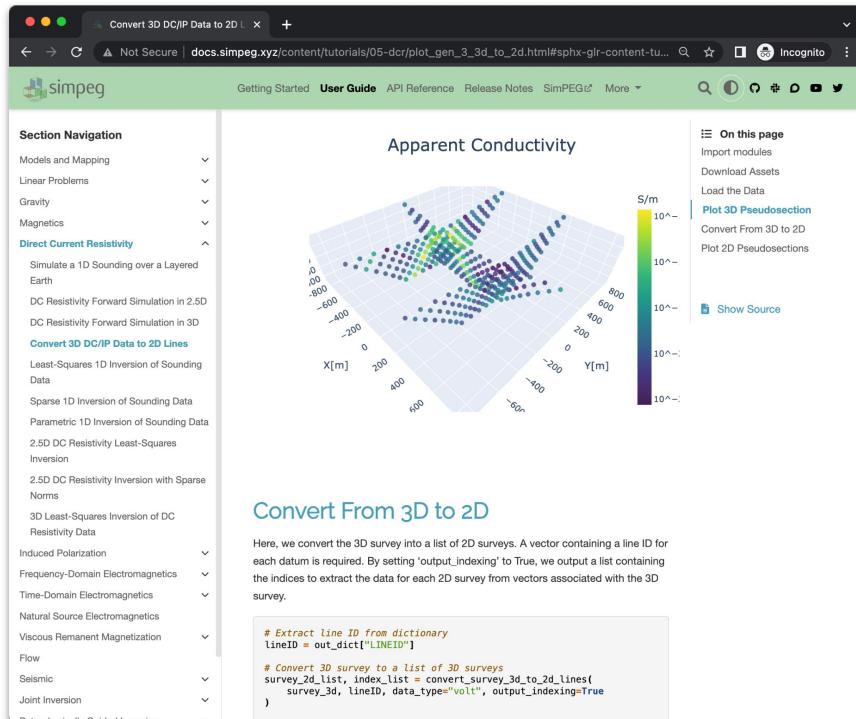


Many aspects production-ready

- OcTree meshes
- Gravity
- Magnetics
- DC Resistivity
- Induced Polarization
- EM1D

Other areas in research and development

- 3D FDEM
- 2D, 3D NSEM
- 3D TDEM
- Joint Inversions



# simpeg & GIF Fortran codes

GIF Fortran codes: widely used, applied, and trusted

- plan to maintain Fortran codes and make functional improvements
- code validations project: comparing, testing Fortran & simpeg codes

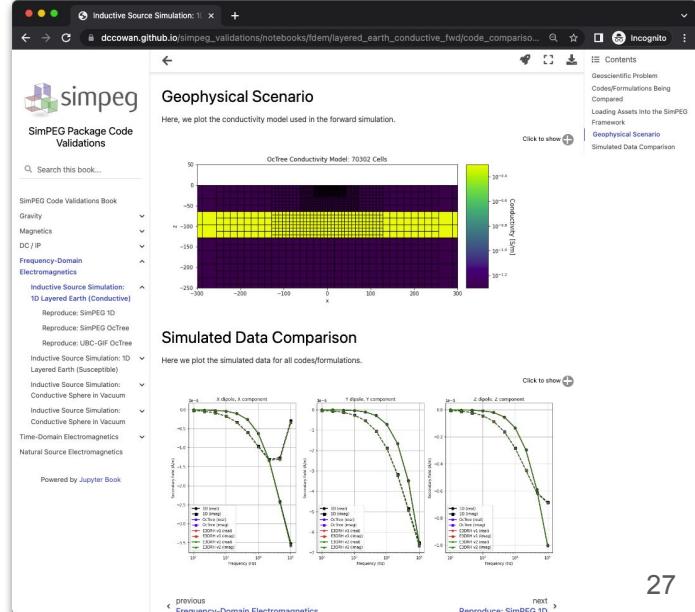
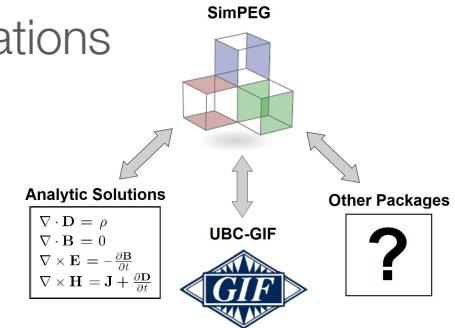
Moving forwards with open source

- focus of new research & development efforts
- streamlines technology transfer
- accelerates research...?

code validations



Devin Cowan

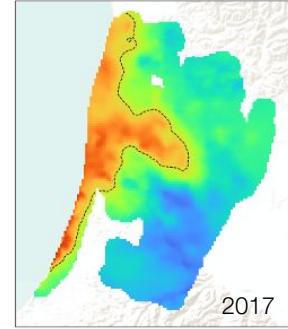
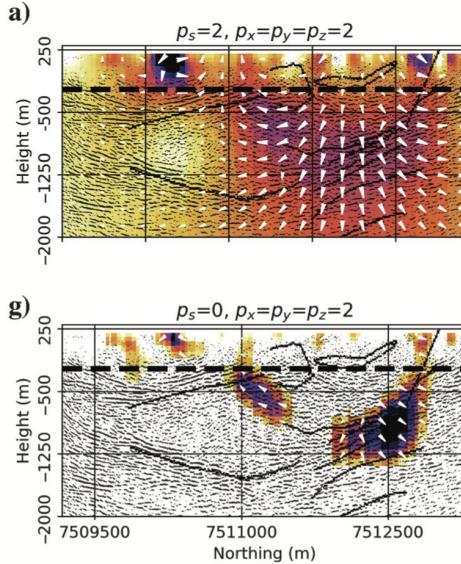


# sparse norms

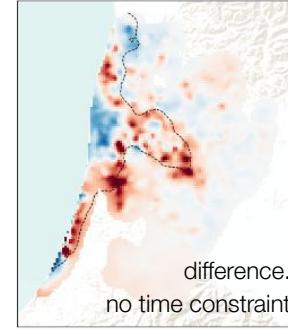
developed in potential fields



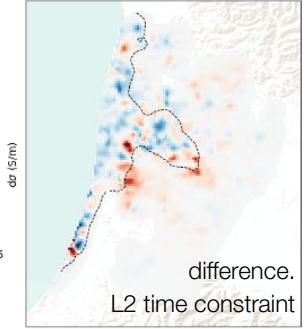
[\(Fournier et al., 2019\)](#)



adapted to time-lapse AEM



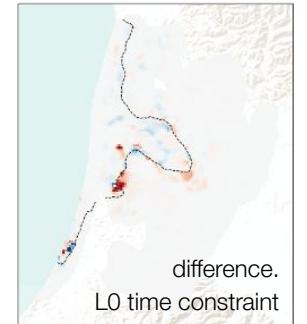
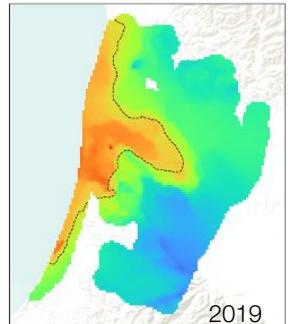
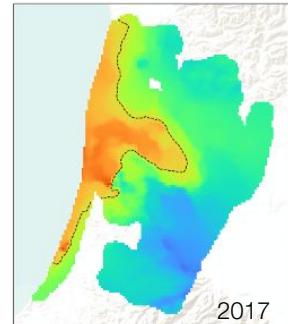
difference,  
no time constraint



difference.  
L2 time constraint



[\(Kang & Knight, 2022\)](#)



difference.  
L0 time constraint

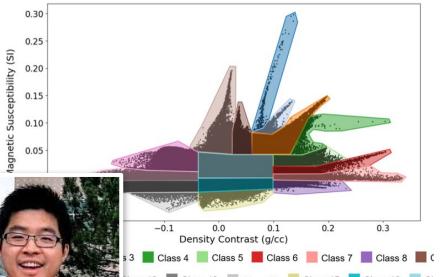
$$\phi_m = \alpha_s \int_v w_s |\mathbf{m} - \mathbf{m}_{ref}|^{p_s} dV + \alpha_x \int_V w_x \left| \frac{d\mathbf{m}}{dz} \right|^{q_z} dV$$

# joint inversions

$\phi_{\text{data}} = \phi_{\text{grav}} + \phi_{\text{mag}} \# \text{ one earth?}$

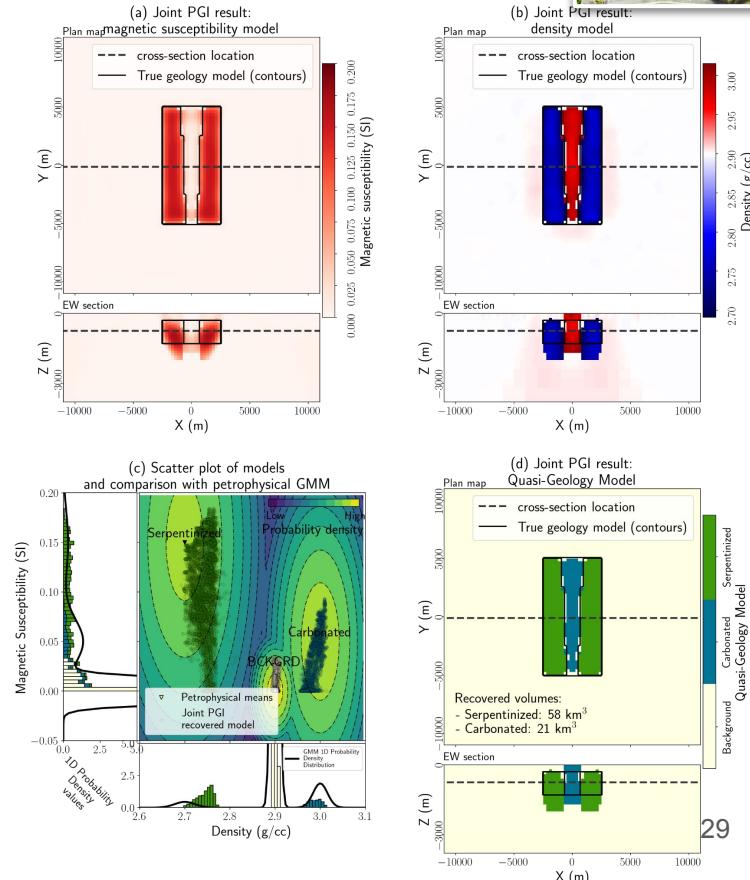
- modular framework enables extension to joint inversions
- PGI, cross-gradient, joint total variation implemented

Cross gradient & geologic differentiation at QUEST  
[Kim et al., 2020](#)



Jiajia Sun

Petrophysically and Geologically guided Inversion (PGI).  
[Astic et al., 2021](#)



# parallelization: an industry-academic collaboration

parallelization motivated initially by large scale 3D DCIP

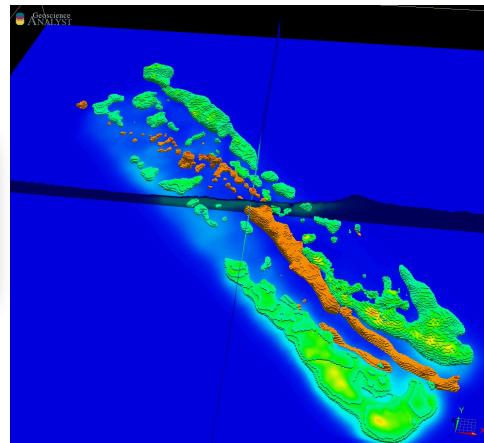


use of dask for parallelization (on any architecture)

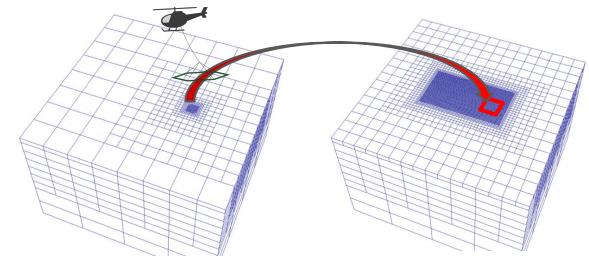
now adapted for Mag, Grav, and working on AEM, MT



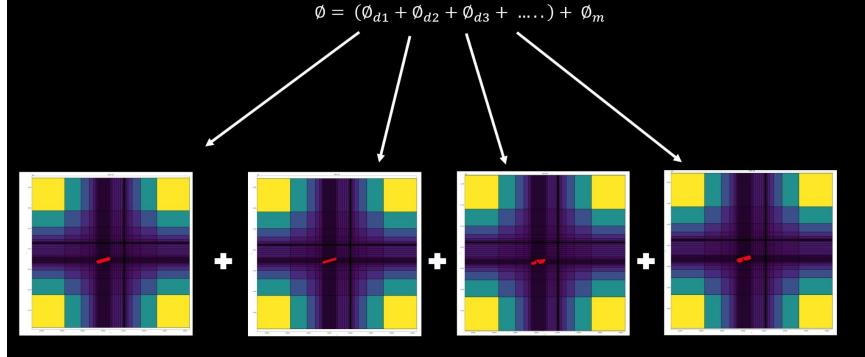
John Kuttai



17.1 million cells 783 injections



Parallelizing the large-scale inversion example:

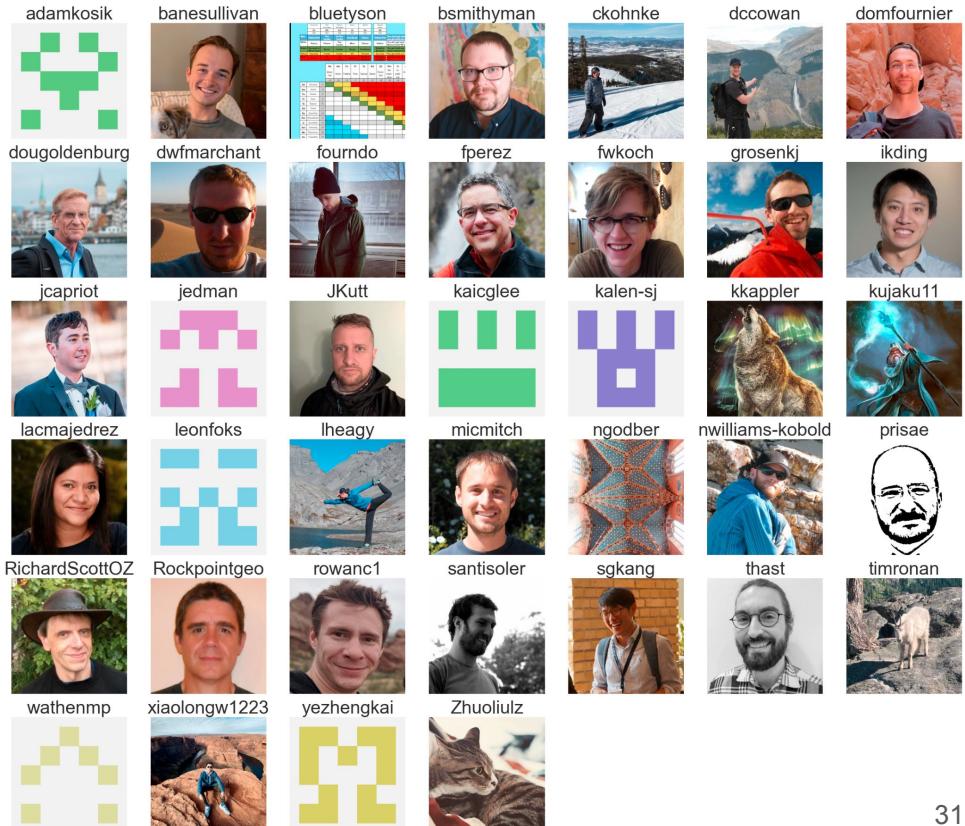


# what have we done in 10 years?

built a foundation for new research  
 → framework & code

importantly, growth of the community  
 → academic & industry

facilitated by open-source approach:  
 → enables collaboration, adaptation,  
 re-use and so much more!



# thank you and questions



Jean  
Legault



Paolo  
Berardelli



David  
Hitz

